



Seasonality in planktic foraminifera of the central California coastal upwelling region

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Abstract. The close association between planktic foraminiferal assemblages and local hydrography make foraminifera invaluable proxies for environmental conditions. Modern foraminiferal seasonality is important for interpreting fossil distributions and shell geochemistry as paleoclimate proxies. Understanding this seasonality in an active upwelling area is also critical for anticipating which species may be vulnerable to future changes in upwelling intensity and ocean acidification. Two years (2012–2014) of plankton tows, along with conductivity–temperature–depth profiles and carbonate chemistry measurements taken along the north-central California shelf, offer new insights into the seasonal dynamics of planktic foraminifera in a seasonal coastal upwelling regime. This study finds an upwelling affinity for *Neogloboquadrina pachyderma* as well as a seasonal and upwelling associated alternation between dominance of *N. pachyderma* and *Neogloboquadrina incompta*, consistent with previous observations. *Globigerina bulloides*, however, shows a strong affinity for non-upwelled waters, in contrast to findings in Southern California where the species is often associated with upwelling. We also find an apparent lunar periodicity in the abundances of all species and document the presence of foraminifera even at very low saturation states of calcite.

association between planktic species and local hydrography means that fossil foraminiferal assemblages have often been used to reconstruct the distribution of water masses through time (e.g., Berger, 1968; McIntyre et al., 1972; Oberhänsli et al., 1992; Ufkes et al., 1998). However, at sites where overlying water masses change seasonally, the foraminiferal fossil record will represent a combination of individuals that may have grown under vastly different conditions. This averaging of short-term variability has the potential to impact the interpretation of any proxy based on foraminifera. Seasonality in a variety of environments has been shown to have a pronounced effect on foraminiferal communities, with species assemblages changing throughout the year (Thunell et al., 1983; Reynolds and Thunell, 1985; Thunell and Honjo, 1987; Thunell and Sautter, 1992; Ortiz et al., 1995; Marchant et al., 1998; Eguchi et al., 2003; Pak et al., 2004). Previous studies have explored seasonal assemblage shifts in the North Pacific, including at Station Papa (50° N, 145° W; Thunell and Reynolds, 1984; Reynolds and Thunell, 1985), in the California Current off of Oregon (> 130 km offshore) (Ortiz et al., 1995), in the Santa Barbara Basin (Kincaid et al., 2000; Darling et al., 2003), off Southern California (Sautter and Thunell, 1991), and in the western Pacific (Eguchi et al., 2003). The majority of this prior work has focused on open-ocean assemblages, however, leaving a gap in understanding the seasonal dynamics in coastal upwelling regions, as well as a significant spatial gap within the California Current system between the Southern California Bight and Oregon.

An improved understanding of coastal upwelling fauna is also important for interpreting the paleoclimate record of these conditions (Reynolds and Thunell, 1986; Naidu and

1 Introduction

Planktic foraminifera have a long history as paleoceanographic proxies due to their environmental sensitivity, cosmopolitan distribution, and extensive fossil record. The close

Malmgren, 1995; Vénec-Peyré and Caulet, 2000; Ishikawa and Oda, 2007). Many modern surveys have characterized upwelling-associated foraminifera through plankton tow and sediment trap studies in the tropics and subtropics (e.g., Thiede, 1975; Naidu, 1990; Thunell and Sautter, 1992; Pak et al., 2004; Salgueiro et al., 2008). Temperate and subpolar upwelling communities such as those found along the central California shelf, however, remain poorly understood. On-shelf assemblages are particularly important for regions dominated by coastal upwelling processes where the alternation between upwelling and relaxation (periods of reduced wind strength in between upwelling periods) has large regional impacts on oceanography and planktic communities (Botsford et al., 2006; Dugdale et al., 2006; Largier et al., 2006; Garcí-Reyes et al., 2014). From a paleontological perspective, modern nearshore assemblages are of interest because sediments in these regions are among those most likely to contain a preserved carbonate fossil record, and thus intact fossil assemblages, due to the high sedimentation rates and the limitations of a narrow continental shelf above a shallow lysocline.

Understanding planktic foraminiferal assemblages in coastal upwelling regions is also relevant for predicting future climate and ecosystem perturbations. The California Current and other Eastern Boundary Current upwelling systems have been identified as especially susceptible to ocean acidification due to the incorporation of anthropogenic CO₂ into the surface ocean superimposed on naturally corrosive waters (Feely et al., 2008; Hofmann et al., 2010; Hauri et al., 2013). The pronounced influence of upwelling in this region is also likely to intensify due to anthropogenic impacts (Bakun, 1990; Garcí-Reyes and Largier, 2012; Sydesman et al., 2014), compounding the impacts of ocean acidification. Planktic calcifiers such as pteropods (Bednaršek et al., 2014; Busch et al., 2014), coccolithophorids (Beaufort et al., 2011; Iglesias-Rodríguez et al., 2008; Langer et al., 2006), and foraminifera (Barker and Elderfield, 2002; Manno et al., 2012; Moy et al., 2009) may be especially vulnerable to reductions in ocean calcite and aragonite saturation state. Upwelled waters are already becoming more acidic along the California margin, and the seasonal duration for which fauna are exposed to waters undersaturated with respect to aragonite and calcite is predicted to increase in the near future (Feely et al., 2008; Gruber et al., 2012; Harris et al., 2013; Hauri et al., 2013). The response of planktic foraminiferal assemblages to 20th century warming has been documented in Southern California (Field et al., 2006). An understanding of the modern seasonality of planktic foraminifera in this intense upwelling region can therefore serve as a baseline for future climate-driven change and may help to identify which upwelling species may already be living at low-saturation state and be potentially tolerant of low calcite saturation state waters that may resemble future conditions in the open ocean.

Here we focus on planktic foraminiferal assemblages sampled along a cross-shore transect over the central California shelf extending from 1 km offshore to the shelf break (30–60 km offshore). Plankton tows, supported by in situ water column data and discrete bottle samples, allow a documentation of species associations based on instantaneous (as opposed to time-averaged) water column conditions. Our goal was to understand (1) the spatial and temporal distribution of planktic foraminifera along the central California shelf and (2) the manner in which species assemblages respond to high frequency changes in water mass, especially those associated with upwelling. These efforts may offer a general framework for interpreting seasonality in foraminiferal records drawn from analogous oceanographic regions and could yield new insights into how this important group of marine calcifiers responds to ongoing climate change and acidification in coastal upwelling systems.

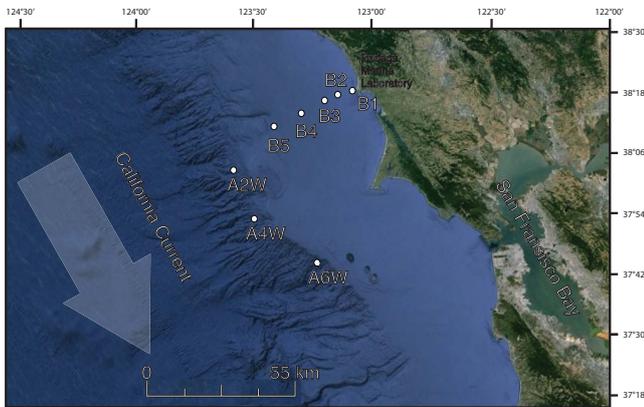
Regional setting

The California Current is the southward flowing arm of the North Pacific Subtropical Gyre and, along with the seasonal Davidson Countercurrent, flows adjacent to the central Californian coastline to the west of our study sites. At many locations along the coast, wind-driven coastal upwelling brings deeper, colder, nutrient-rich and low-O₂ water to the surface, with the strongest upwelling signal found in a 10 to 25 km band just offshore (Hickey, 1979, 1979; Huyer, 1983; Lynn and Simpson, 1987).

At the latitudes of our study sites (37–39° N), wind-driven coastal upwelling is generally strongest in April–June (García-Reyes and Largier, 2012). During the upwelling season, wind-driven upwelling events are interspersed with relaxation periods, the combination of which is responsible for large changes in productivity in the plankton (Botsford et al., 2006; Dugdale et al., 2006; Largier et al., 2006; Garcí-Reyes et al., 2014). During the upwelling season, further complexity is introduced through the advection of upwelled water masses both away from the continent and alongshore, with water parcels in the region which are dominantly sourced from the north (Kaplan and Largier, 2006). Outside of the upwelling season (~September–March), upwelling events are generally absent and there is occasional occurrence of downwelling, with net northward flow of water. Advection rates are variable but have been reported in the range of 10–30 km d⁻¹ (Kaplan and Largier, 2006). This stable post-upwelling season generally lasts into December when the stability can be punctuated by storm conditions (Kaplan and Largier, 2006; Garcí-Reyes and Largier, 2012). Together, these conditions create an environment of strong seasonality in terms of productivity, temperature, O₂, carbon chemistry, and water mass, all of which would be expected to influence the species of planktic foraminifera present in the region.

Table 1. Station locations, depths, and the number of times sampled over the course of this study.

Station	Latitude	Longitude	Depth sampled (m)	No. of times sampled
BL1	38°16'59"	−123°04'60"	25	15
BL2	38°23'38"	−123°13'00"	45	15
BL3	38°21'05"	−123°14'20"	90	15
BL4	38°26'20"	−123°27'01"	120	15
BL5	38°21'05"	−123°37'59"	200	14
A2W	38°02'45"	−123°33'47"	200	5
A4W	37°52'55"	−123°28'30"	200	4
A6W	37°43'20"	−123°13'59"	200	5

**Figure 1.** Map of tow stations BL1–5, A2W, A4W, and A6W.

2 Methods

2.1 Study area

Plankton collection took place at eight stations located at increasing distances from shore across the continental shelf (Fig. 1). Bodega Line (BL) (38°) sites start at nearshore station BL1, 1 km offshore, and extend across the shelf to station BL5, 32 km offshore. These stations were sampled monthly to bimonthly from September 2012 to September 2014. Three additional stations were sampled in 2013 and 2014 as part of the Applied California Current Ecosystem Studies (ACCESS) cruises (three times per year) and are located just over the shelf break at 40–60 km offshore, spanning a latitudinal range from 37–39° N (Table 1). All sampling stations are shoreward of the central core of the California Current (Lynn and Simpson, 1987) and are strongly influenced by both spring/summer upwelling as well as winter storms (Fig. 1).

2.2 Sample collection

Vertical net tows integrated foraminifera across the water column from the surface to a depth of 200 m or to 10 m above the sea floor at shallower sites. All foraminifera were sampled with a 150 µm mesh net. This approach po-

tentially excludes juveniles and small adults and therefore limited the samples to foraminifera of readily identifiable adult developmental stages and to species normally included in fossil analyses. Most samples were placed in ambient surface seawater and kept chilled without further preservation to be picked immediately upon return to shore. When this procedure was not feasible, samples were preserved shipboard in 95 % ethanol, buffered to a pH > 8.5 with Tris(hydroxymethyl)aminomethane. Foraminifera were picked wet from bulk tow material, rinsed in DI water and archived in slides. All archived foraminifera were identified to the lowest possible taxonomic level, with *N. incompta* and *N. pachyderma* defined primarily by shell coiling direction (following Darling et al., 2006). No distinction was made between living and dead individuals although almost all shells still contained some cytoplasm at the time of sorting. Taking into account the conservative end of the range of sinking rates for shells (e.g., 29–552 m d^{−1}; Takahashi and Be, 1984) and that foraminifera were sampled from the upper 200 m of the water column, we can assume that all foraminifera were likely alive within 6 days of collection. Transport data from the region allow us to further estimate a maximum horizontal transport of 50 km in 5 days, indicating that all shells still within the water column were locally sourced (Kaplan and Largier, 2006).

2.3 Environmental measurements

Water column profiles for temperature, salinity, dissolved O₂ (DO), and fluorescence were obtained across the plankton tow depths using a Seabird Electronics SBE 19 conductivity–temperature–depth (CTD) profiler. Plankton tow nets were equipped with a flow meter for each cast; however, due to frequent failures, flow rates were unreliable and are not reported here. At each station, discrete bottle samples of surface water and water from the bottom of each CTD cast were collected using a Niskin sampler. All water samples were analyzed spectrophotometrically for pH (total scale) using either a Sunburst SAMI (Submersible Autonomous Moored Instrument) modified for benchtop use (SD ± 0.009) or an Ocean Optics Jaz Spectrophotometer EL200 (SD ± 0.003) using *m*-cresol purple (Dickson et al., 2007). Total alkalinity was determined via automated Gran titration on a Metrohm 809 Titrando (SD ± 2.809 µmol kg^{−1}), with acid concentrations standardized to Dickson certified reference materials (A. Dickson, Scripps Institution of Oceanography, CA, USA). Measurements of pH and alkalinity were carried out at UC Davis Bodega Marine Laboratory and used to calculate other inorganic carbon system parameters, including calcite saturation state (Ω_{Ca}) and [CO₃[−]], using the software CO₂Calc (Robbins et al., 2010). Thermocline depths were defined as the depth (below 5 m) at which the greatest gradient in temperature occurred, exclusive of any temperature change with a slope of less than 0.1 °C m^{−1}, in which case the thermocline was assumed to be deeper than the profiled

water. Upwelling index is taken from the PFEL upwelling index modeled for 39° N (<http://www.pfel.noaa.gov/products/PFEL/modeled/indices/upwelling/upwelling.html>), which is in general agreement with temperature measurements from the Bodega Ocean Observing Node (BL1).

2.4 Data analysis

For the four most abundant species, *G. bulloides*, *Turborotalia quinqueloba*, *N. incompta*, and *N. pachyderma*, we performed a CCA (canonical correspondence analysis) on relative abundances using the “vegan” package in R (R Core Team, 2013; Oksanen et al., 2016). Potential explanatory variables included day of the lunar cycle relative to the new moon, upwelling index, duration of sustained upwelling as indicated by the PFEL upwelling index, surface and deep water carbonate system parameters, and CTD temperature, salinity, fluorescence, and DO. CTD data were binned into depths at 5 m intervals to a depth of 25 m and then at 10 m intervals. A subset of variables was selected for CCA by exclusion of all variables not related to any total abundances by pairwise correlation at a 95 % confidence level. Strongly interrelated or redundant parameters were additionally excluded (i.e., a parameter correlated at multiple consecutive depths would have been considered at only one of these depths).

3 Results

The assemblage was heavily dominated by the planktic species *N. pachyderma*, *N. incompta*, *T. quinqueloba*, and *G. bulloides*, representing 35.3, 23.1, 13.5, and 11.7 % of all recovered foraminifera, respectively. Less common forms included *Globigerinita glutinata*, *Globorotaloides hexagonus*, *Globigerinella calida*, *Globigerinita uvula*, and *Globorotalia* spp., as well as the occasional cosmopolitan species, *Orbulina universa* and subtropical *Neogloboquadrina dutertrei*, and, rarely, benthic species of foraminifera. The presence of these latter taxa was sporadic and in low abundance (all < 1 % of the overall recovered foraminifera, with the exception of *G. glutinata* at 2.1 %); therefore, further analysis will be confined to the four most abundant species.

At offshore stations BL3, BL4, and BL5 and off-shelf stations A2W, A4W, and A6W, foraminifera displayed a clear seasonality. The year can be divided between spring / summer and fall / winter faunas that coincide with the upwelling-dominated and non-upwelling season (Fig. 2). Beginning in May, shortly after the onset of upwelling, samples began to show a high abundance of *T. quinqueloba*. A bloom of *N. pachyderma*, seen here as an increase in total abundance to >200 individuals, occurs in July or August, after several months of sustained upwelling, followed by a decrease in both total and relative abundance to less than 50 % by the end of summer (Fig. 3). *N. pachyderma* was

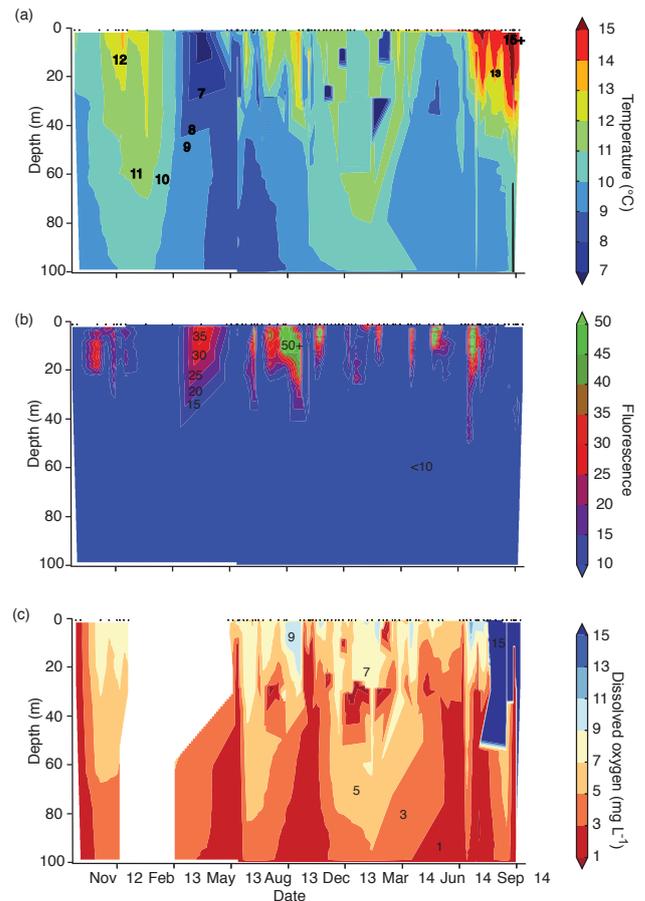


Figure 2. A time series of CTD cast profiles for (a) temperature, (b) productivity, and (c) DO taken between September 2012 and October 2014. Time series are compiled from CTD casts at BL5 in conjunction with plankton tows and supplemented with data from weekly CTD casts taken at BL1 as a part of the Bodega Ocean Observing Node. Black points at the top of each figure denote the location of each CTD cast.

also present through much of the winter in lower numbers in 2012–2013. By contrast, this species was virtually absent in the winter of 2013–2014, before reappearing after a period of sustained upwelling in July 2014 (Fig. 3). In both years, the earliest *N. pachyderma* blooms appeared to initiate farther offshore, although abundances within a given samples did not appear to be directly linked to specific upwelling events.

Following the end of the summer season, the fall–winter fauna shows a more even distribution of species and a distinct shift in the ratio of *N. pachyderma* to *N. incompta* (Fig. 3). *N. incompta* was equally or more abundant than *N. pachyderma* during the non-upwelling season although it was present year-round. *G. bulloides* also began to appear in the water column in the fall, strongly associated with non-upwelled waters, and is present throughout the winters. *G. bulloides* was present primarily in the winter and either ab-

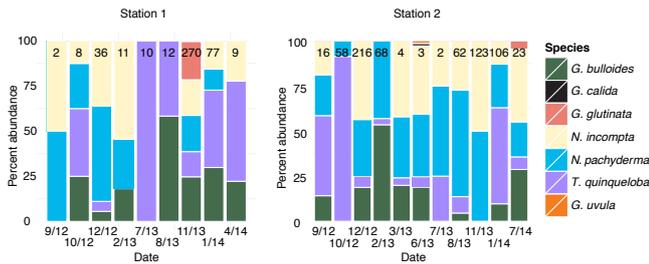


Figure 3. Relative abundance of all species at nearshore stations BL1 and BL2, with total number of foraminifera marked at the top of each bar.

sent or found only in very low numbers during the summer season.

The same suite of species was present at nearshore stations BL1 and BL2, but counts were lower year-round and most seasonal patterns seen offshore were not evident. *N. pachyderma* did appear to increase in relative abundance during the summer at these stations but remained in low abundance along with *N. incompta* year-round (Fig. 4). *T. quinqueloba* was also observed year-round at these nearshore stations. A greater proportional abundance of *G. bulloides* was seen during the fall and winter at nearshore sites, consistent with findings at the offshore stations (Fig. 4).

3.1 Environmental measurements

In spring and summer, surface hydrographic conditions were highly variable, reflecting alternating upwelling events and relaxation periods. Frequent changes in thermocline depth were observed, as well as intermittent blooms of near-surface productivity (Fig. 2). This resulted in a more surface-stratified and productive water column, with a shallow thermocline and high fluorescence in the upper water column. During upwelling season, near-surface temperatures cool to 8–9 °C, and subsurface waters approach calcite undersaturation ($\Omega_{Ca} < 1$) and display low DO ($< 4 \text{ mg L}^{-1}$ at $< 90 \text{ m}$) (Fig. 5). Despite consistently lower subsurface DO and pH, high near-surface productivity often increased DO and pH near-surface values, creating a noticeable down-profile gradient in these parameters.

Beginning in the late fall, and continuing into early spring, a consistently deep thermocline was observed at all stations. This trend often had the effect of confining the entire on-shelf water column (including all tow samples) to this deep mixed layer, which dominated the shelf in winter. Temperatures were generally warmer (11–14 °C) than during the upwelling season with relatively low fluorescence in the upper water column (< 4) (Fig. 5) and $\Omega_{Ca} > 1.5$ throughout the sampled water column. García-Reyes and Largier (2012) describe storm conditions, which are likely to have contributed to the deep mixed layer, observed outside of upwelling season (especially between January and March).

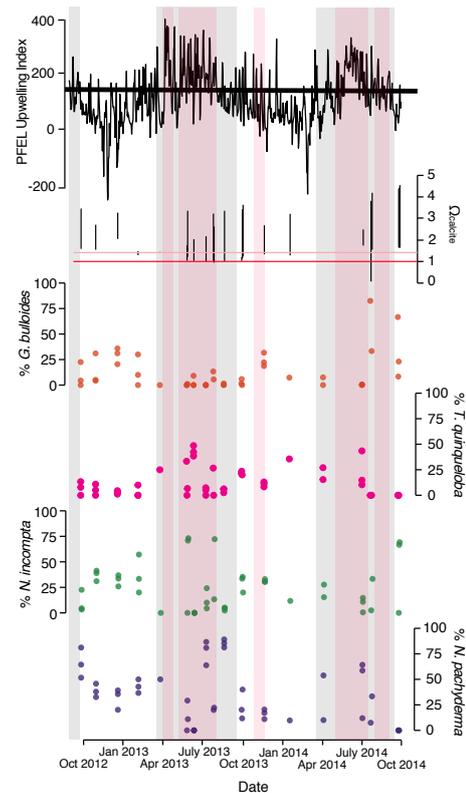


Figure 4. Upwelling index for 39° N (PFEL; <http://www.pfel.noaa.gov/products/PFEL/modeled/indices/upwelling/upwelling.html>), with “upwelling season” shaded in grey and periods of sustained upwelling conditions during the relevant years shaded in red. The range of Ω_{calcite} observed on each tow is marked in black, with $\Omega_{\text{calcite}} = 1$ marked in red and $\Omega_{\text{calcite}} = 1.5$ in pink. Percent abundances from vertical tows of *G. bulloides*, *T. quinqueloba*, *N. incompta*, and *N. pachyderma* from offshore stations BL3, BL4, BL5, and off-shelf stations AW2, AW4, and AW6 shown in orange, pink, green, and purple, respectively.

3.2 Canonical correspondence analysis

A CCA shows that the highest relative abundances of both *G. bulloides* and *N. incompta* fall along dimensions strongly influenced by environmental variables characteristic of non-upwelling season, such as higher temperatures and increased dissolved oxygen. The highest relative abundances of *N. pachyderma* are most closely associated with fluorescence, especially near the surface (5 m) and at 60–70 m water depth, more associated with upwelling waters. *T. quinqueloba* is also distinct, falling within the quadrant most associated with tows taken during upwelling (Fig. 6). Figure 5 additionally shows the range of environmental conditions captured by sampling around the full moon (lunar days 14–18).

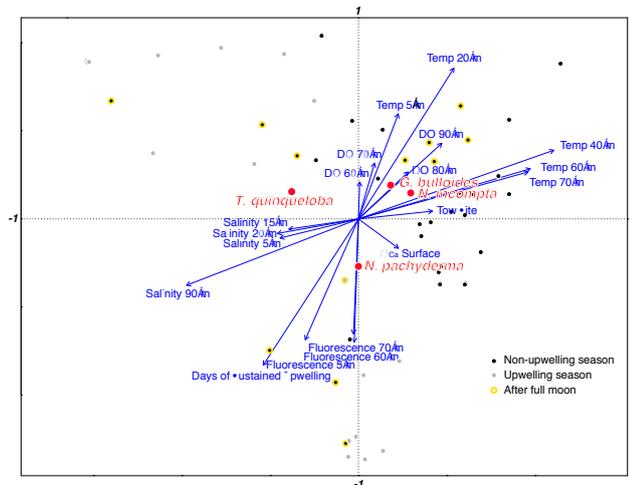


Figure 5. Biplot of canonical correspondence analysis of relative abundance of the four most abundant species compared with environmental variables at the time of collection for all tows in which foraminifera were found. Dominant species are represented by large red points, and small points represent individual tows. Grey points are tows that fall within upwelling season, as shown in Fig. 2, and black points outside of this season. Points with a golden halo represent tows during lunar days 14–18.

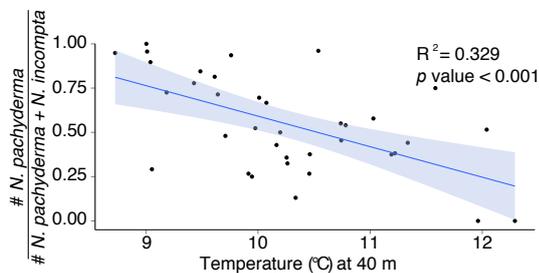


Figure 6. Ratio of *N. pachyderma* to *N. pachyderma* and *N. incompta* at 40 m depth with 95 % confidence envelopes.

3.3 Neogloboquadrina coiling direction

Coiling direction for Neogloboquadrinids is recognized as an empirical proxy for sea-surface temperature in the sedimentary record (Ericson, 1959; Bandy, 1960; Kennett, 1968; Bé and Tolderlund, 1971; Vella, 1974; Arikawa, 1983; Reynolds and Thunell, 1986). We tested whether the relationship is consistent on shorter timescales with mixed assemblages of *N. pachyderma* (primarily sinistral coiling) and *N. incompta* (primarily dextral coiling) (Darling et al., 2006). A very weak linear correlation with surface temperature is observed, between the ratio of *N. pachyderma* to all *N. pachyderma* and *N. incompta* ($r^2 = 0.09626$; p value = 0.02). Correlations improved deeper in the water column, with a weak but notable relationship at 40 m ($r^2 = 0.3285$; p value < 0.001) (Fig. 7).

4 Discussion

4.1 Foraminiferal seasonality

A key finding of this study is the clear seasonality of the four most abundant species of planktic foraminifera at offshore stations along the central California shelf. Our findings highlight the importance of seasonal-scale water column shifts in dictating foraminiferal species abundances, as well as suggest which species may be most vulnerable to ocean acidification in the region. It may also act as a guide to paleoceanographers in deciphering the specific species most likely to be recording seasonal signals along the shelf. *T. quinqueloba* appears to be associated mainly with the early summer months and the beginning of upwelling season as indicated by the PFEL Upwelling Index for the relevant study years. *N. pachyderma* increased in both total number and relative abundance in the late summer months following the onset of upwelling. *G. bulloides* is largely confined to the winter non-upwelling season while *N. incompta* is present in all seasons. The year-round presence of *N. incompta* combined with the high summer abundance of *N. pachyderma* creates the appearance of a seasonal switch in the relative abundances of the two Neogloboquadrinids (Fig. 3). These trends are described in more detail for each of the four species below.

4.1.1 Neogloboquadrinids

The seasonal trade-off observed at offshore stations between *N. pachyderma* and *N. incompta* is in agreement with previous studies interpreting seasonality from the geochemistry of the two species. Sediment trap data from the western North Pacific found that *N. incompta* and *G. bulloides* reflect winter sea-surface temperature while *N. pachyderma* reflects summer (Sagawa et al., 2013). Similarly, Mg/Ca ratios in recent fossils from the Norwegian Sea indicate that *N. pachyderma* is primarily a summer bloom species while *N. incompta* records winter conditions (Nyland et al., 2006). The close association between *G. bulloides* and *N. incompta* seen here has also been noted previously both in the water column and in coretop records (Reynolds and Thunell, 1986; Giraudeau, 1993; Ufkes et al., 1998).

The ratio of *N. pachyderma* to *N. incompta* (previously *N. pachyderma* var. sinistral and *N. pachyderma* var. dextral, respectively) has long been recognized to be paleoceanographically significant in marine sediments, with *N. pachyderma* associated with subpolar water masses, *N. incompta* associated with subtropical to temperate waters, and the ratio between the two acting as a proxy for sea-surface temperature (Ericson, 1959; Bandy, 1960; Kennett, 1968; Bé and Tolderlund, 1971; Vella, 1974; Arikawa, 1983; Reynolds and Thunell, 1986). The relationship observed here between coiling direction of Neogloboquadrinids and temperature is weak, at best, at the surface. The relationship is slightly stronger at 40 m depth (Fig. 7), with an equal ratio between

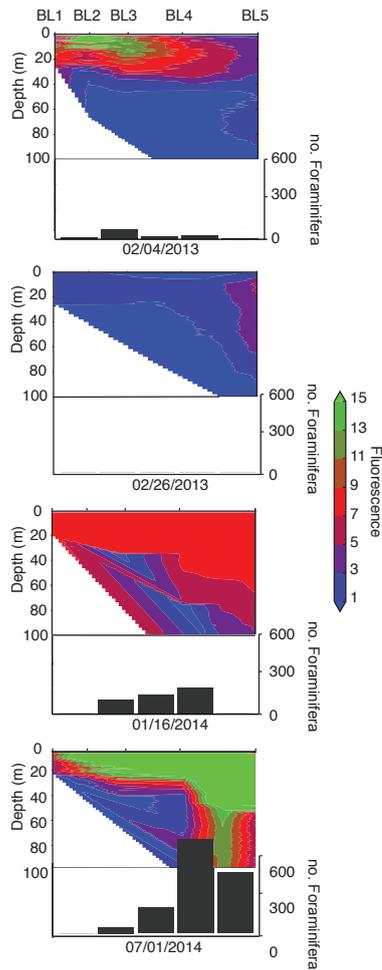


Figure 7. Water column fluorescence data and total number of foraminifera recovered at stations BL 1–5 on 4 days of extremely low productivity. CTD data from these five stations demonstrate small-scale variability from 1 to 32 km offshore along the continental shelf. This is compared with the total number of foraminifera retrieved at each of these stations.

N. incompta and *N. pachyderma* found around 10.5 °C. This ratio can largely be explained by the year-round presence of *N. incompta*, punctuated by an increase in *N. pachyderma* in the summer along with cooler temperatures, especially in the subsurface. These findings validate on short timescales what has been seen to be empirically true over longer timescales: *N. pachyderma* is found primarily in high-latitude waters and, when occurring in temperate regions, occurs mixed with *N. incompta*, in both the water column or sediment. This pattern is suggestive of an incursion of these cooler northern waters and not solely the impact of upwelled waters in this region (< 10 °C conditions).

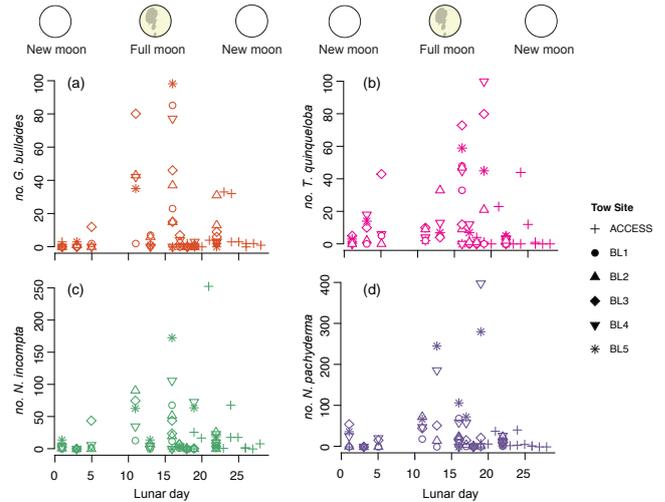


Figure 8. *G. bulloides*, *T. quinqueloba*, *N. incompta*, and *N. pachyderma* counts by lunar day from the new moon (day 0) to full moon (Day ~ 14). Abundances are taken from integrated tows, at 25, 45, 90, 120, and 200 m depending on the station (see Table 1). Symbols denote tow station.

4.1.2 Globigerina bulloides

Globigerina bulloides has previously been associated with active upwelling in Southern California (Sautter and Thunell, 1991; Field et al., 2006) and the Arabian Sea (Peeters et al., 2002). Observations along the central California shelf are in direct contrast to this, with *G. bulloides* observed to be far more abundant during the fall/winter relaxation and storm season (Fig. 3). It is notable that in at least one previous study, *G. bulloides* has shown a bimodal abundance in Southern California, with one population of *G. bulloides* associated with winter and another population with the spring / summer upwelling season (Sutter and Thunell, 1991). Furthermore, two distinct genotypes of *G. bulloides* have been identified in Southern California, one of which is present in winter samples and was previously recognized in “subpolar” regions (Darling et al., 2003). We interpret the *G. bulloides* observed along the central California coast as connected to this “subpolar”/winter population, accounting for the differences in seasonal abundance seen at our northern site compared to Southern California.

4.1.3 Spatial dynamics

Nearshore stations BL1 and BL2 are shoreward from the primary band of coastal upwelling (Huyer, 1983) and show less seasonality in species abundances with the exception of *G. bulloides*, which is more abundant in the fall and winter nearshore as well as offshore. Although non-spinose forms are also occasionally present at both nearshore sites, they do not show the seasonality that they do at offshore sites (Fig. 4). Some of the differences seen in the fauna at BL1 and BL2

compared to offshore stations may be due to shallower tow depths at these sites (25 and 45 m, respectively) and therefore a bias in favor of species living closer to the surface, which may include *G. bulloides*. However, shallow tows conducted at BL4 and BL5 confirm that all four species considered here are present in the upper water column (<30 m) at these sites, so depth alone cannot completely account for the nearshore/offshore difference in foraminiferal abundances. Nearshore stations may be sheltered from larger-scale transitions in source water that happen over most of the shelf and be more impacted by terrestrial processes.

Short-term spatial dynamics were also observed to impact foraminiferal abundance. On days when overall productivity was low, abundances of all foraminifera species were higher at sites with higher fluorescence (indicating higher biomass and suggesting higher primary productivity). Especially low fluorescence (near-surface fluorescence <2) was observed on collection days 4 February 2013, 16 January 2014, 1 July 2014, and 26 February 2013. On these days, foraminifera were recovered in much greater numbers at stations associated with peak fluorescence regardless of where along the transect the station was located (Fig. 8). No foraminifera were recovered at very low-productivity stations BL1 and BL5 on 16 January 2014 or at BL1 on 1 July 2014, while other sites yielded >100 individuals. On 4 February 2013, BL2 was associated with the only observation of surface fluorescence >10 and yielded more foraminifera than all other sites combined. Fluorescence was low at all sites on 26 February 2013 and no foraminifera were recovered from these tows (Fig. 8). These data indicate that phytoplankton productivity may ultimately be a limiting factor for all species. On days with higher measured fluorescence (productivity), the dominant spatial trend was towards higher abundances further offshore regardless of where peak productivity was observed.

4.1.4 Foraminifera in reduced pH waters

Upwelling associated waters with low Ω_{Ca} were observed on multiple occasions during plankton tows. Foraminifera are widely thought to be susceptible to ocean acidification (i.e., Barker and Elderfield, 2002; Manno et al., 2012; Moy et al., 2009; Orr et al., 2005), although this susceptibility may not always manifest through a reduction in shell weight in open-ocean conditions (i.e., Beer et al., 2010; Aldridge et al., 2012; Weinkauff et al., 2016). The association of multiple species of foraminifera already living at $\Omega_{Ca} < 1$ or very low Ω_{Ca} (<1.5) waters is notable. In particular, more than a quarter (26 %) of all observed *N. pachyderma*, with its strong upwelling association, was found to occur in a water column with $\Omega_{Ca} < 1$ in the upper 160 m. Culture studies with this species have indicated a decrease in shell weight associated with low Ω_{Ca} well within the range of those that *N. pachyderma* was found in during upwelling season, indicating the potential to impact carbonate flux in areas where this organism is an important

calcifier (Manno et al., 2012). If *N. pachyderma* is already living near its Ω_{Ca} tolerance, this species may be exceptionally vulnerable to a continued increase in ocean acidification in this region. However, it is possible that upwelling-adapted *N. pachyderma* may prove to be an example of a calcifying plankton able to tolerate undersaturated waters.

4.2 Causes of seasonality and fluctuations in abundance

Seasonality has previously been identified in temperate regions globally for all four of the species addressed here (Jonkers and Kucera, 2015). One important mechanism contributing to the seasonal progression of foraminifera species along the shelf in central California is the alternation between the direction of net water transport between upwelling and non-upwelling seasons. This phenomenon would account for the occurrence of *G. bulloides* in greater numbers outside of upwelling season when net poleward water transport is expected (Kaplan and Largier, 2006). Similarly, the influx of subpolar associated *N. pachyderma* could be due to this species being carried into the region during the southward transport of water that occurs during upwelling season (Kaplan and Largier, 2006). This is in slight contrast to the hypothesis of *N. pachyderma* dormancy outside of upwelling season suggested in the Arabian Sea (Ivanova et al., 1999), since in the California Current as well as on the Namibian margin there is some synchronicity between the preferred seasonality of *N. pachyderma* and greater transport into the region from high-latitude waters (Ukkes and Zachariasse, 1993). An alternation between the foraminiferal fauna of source waters additionally offers an explanation for the seasonal absence and reappearance of both *N. pachyderma* and *G. bulloides*. *N. incompta*, found year-round in the study region, may be present in both water masses.

In addition to the broad oscillation of source waters, higher counts of each species are associated with some specific water column characteristics. In most cases, species abundances could not be linked strongly to single environmental parameters, but rather a suite of hydrographic and temporal variables were required to account for faunal assemblages. For some species, particular variables can be identified through pairwise correlation as having a significant effect on abundance (Fig. S2 in the Supplement). In *G. bulloides*, higher abundance correlates with higher water temperatures throughout the water column. In *N. pachyderma*, higher abundances are associated with higher fluorescence and thus enhanced primary production. For *N. incompta*, counts seem to loosely correlate with higher O_2 and lower salinities, while *T. quinqueloba* is not clearly correlated with any single measured parameter (Fig. S2).

A CCA shows the relative abundances of *G. bulloides*, *N. incompta*, and *N. pachyderma* to be dependent upon environmental variables strongly associated with coastal upwelling conditions (Fig. 6). However, the directions of those

associations vary. *N. pachyderma* and, to a lesser extent, *T. quinqueloba* seem to be associated with upwelling-like water conditions. Sediment trap time series have previously linked *T. quinqueloba* to productivity in the North Atlantic (Chapman, 2010), which could explain the association of this species with the higher-productivity season, although it is not directly associated with shifts in productivity as is *N. pachyderma*. *G. bulloides*, however, is negatively associated with upwelling-like water conditions and more associated with warmer waters and higher DO levels seen outside of upwelling season. *N. incompta* relative abundances also increase outside of upwelling season, although in this case relative activity may be key, especially with regards to *N. pachyderma* (see Sect. 4.1.1). This outcome is supported by total counts, which indicate that this species is the only one clearly present at the tow sites year-round.

4.2.1 Lunar periodicity

Abundances of *G. bulloides*, *T. quinqueloba*, *N. incompta*, and *N. pachyderma* all display an abundance cycle with a 28-day period that appears to coincide with the lunar cycle (Fig. 9). Peak counts for each species occur within 7 days of the full moon, before dropping off before the new moon (Fig. 8). This trend offers further evidence that planktic foraminifera reproduce on a lunar cycle (Spindler et al., 1979; Bijma et al., 1990, 1994; Schiebel et al., 1997; Jonkers et al., 2015; Venancio et al., 2016; Erez et al., 1990). The peak abundance for *G. bulloides* occurs before that in the other species, starting 3 days before the full moon and remaining high until 4 days after the full moon. Abundances in *N. pachyderma* and *N. incompta* begin to increase around the same time, but high abundances in these species continue until 5 and 7 days after the full moon, respectively. Whether the observed offsets in peak abundance around the full moon represent inter-species differences in reproductive timing or are an artefact of sampling against a background of strong seasonality in a highly variable environment cannot be resolved from this dataset, even though sampling days surrounding the full moon occurred across seasons (Fig. 5).

4.3 Application to the fossil record

The presence of seasonally distinct faunas along the central California margin can be used in increasing the resolution of paleoceanographic and paleoecological records, as different species clearly represent different states of the seasonal upwelling regime. Single-species geochemical records are likely to show a strong bias towards either upwelled or non-upwelled water masses and, therefore, could potentially be harnessed as a record of changes in upwelling intensity and associated water chemistry. Our findings reaffirm a strong relationship between the dominance of *N. pachyderma* in conditions favorable to upwelling. This pattern has been noted along the African margin (Giraudeau, 1993), in the Ara-

bian Sea (Ivanova et al., 1999), and on the Namibian margin (Ufkes and Zachariasse, 1993). It remains possible that other genotypes of *N. pachyderma* have distinct and ecological preferences particularly those associated with the Arctic and Antarctic (Darling et al., 2006, 2007). As our record is based on discrete tows and not a continuous record, the percent composition of species cannot be directly translated into a sediment flux or to what would be preserved in aggregate in the fossil record. However, the upwelling season bloom of *N. pachyderma* seen here is strong enough that this signal would likely dominate the annual assemblages, although the vast majority of *N. pachyderma* (81 % of those seen in tows) occur between July and November.

Globigerina bulloides has been associated with upwelling at other sites globally (Sautter and Thunell, 1991; Peeters et al., 2002; Field et al., 2006) and even used as an upwelling indicator in the fossil record (e.g., Naidu, 1990; Kroon et al., 1991; Anderson and Prell, 1993; Naidu and Malmgren, 1996). However, within our study region, this species was present almost exclusively outside of upwelling season. The majority (88 %) of the *G. bulloides* seen in our samples was observed between November and February. CCA supports these observations in indicating that the relative abundance of this species is negatively associated with upwelling-like conditions in the region. This situation contrasts with findings in Southern California and the Oman Margin (Peeters et al., 2002; Field et al., 2006), highlighting the importance of using regionally specific associations where possible when interpreting planktic assemblages in the sediment record.

5 Conclusions

Surveys of planktic foraminifera retrieved from plankton tows both confirm and contrast with findings of studies in analogous regions. Along the north-central California shelf there is a clear association between upwelling and *N. pachyderma*, which experiences a summer bloom and increases in relative abundance, as seen at other sites. This summer population of *N. pachyderma* appears to routinely experience low Ω_{Ca} waters, conditions that are predicted to increase in the near future. *N. pachyderma* associations with upwelling and low temperatures are also reflected in the empirical relationship between *Neogloboquadrina* coiling direction, seen previously in the sediment record, with a switch in dominance between *N. incompta* and *N. pachyderma* at 10.5°. *G. bulloides*, however, is associated in our study with non-upwelled waters, in contrast to populations in Southern California and other upwelling regions. All species showed a lunar periodicity in their abundances, evidence of lunar timed reproduction. This study highlights the wealth of information on seasonal-scale processes that is contained within foraminiferal shells. To access this information, however, it is of great importance to ground interpretations of foraminiferal proxies in species and regional ecology to the greatest extent possible.

6 Data availability

All of the foraminiferal abundance data will be included as a supplement. Oceanographic data are available on request from BOON (Bodega Marine Laboratory) and Point Blue (<http://www.pointblue.org/datasharing>).

The Supplement related to this article is available online at doi:10.5194/bg-13-5139-2016-supplement.

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