



Wet season cyanobacterial N enrichment highly correlated with species richness and *Nostoc* in the northern Australian savannah

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Abstract. The Boodjamulla National Park research station is situated in the north-western Queensland dry savannah, where the climate is dominated by summer monsoons and virtually dry winters. Under shrub canopies and in between the tussock grasses cyanobacterial crusts almost entirely cover the flood plain soil surfaces. Seasonality drives N fixation, and in the savannah this has a large impact on both plant and soil function. Many cyanobacteria fix dinitrogen that is liberated into the soil in both inorganic and organic N forms. We examined cyanobacterial species richness and bioavailable N spanning 7 months of a typical wet season. Over the wet season cyanobacterial richness ranged from 6 to 19 species. N-fixing *Scytonema* accounted for seasonal averages between 51 and 93 % of the biocrust. Cyanobacterial richness was highly correlated with N fixation and bioavailable N in 0–1 cm. Key N-fixing species such as *Nostoc*, *Symploca* and *Gloeocapsa* significantly enriched soil N although *Nostoc* was the most influential. Total seasonal N fixation by cyanobacteria demonstrated the variability in productivity according to the number of wet days as well as the follow-on days where the soil retained adequate moisture. Based on total active days per month we estimated that N soil enrichment via cyanobacteria would be $\sim 5.2 \text{ kg ha}^{-1}$ annually which is comparable to global averages. This is a substantial contribution to the nutrient-deficient savannah soils that are almost entirely reliant on the wet season for microbial turnover of organic matter. Such well-defined seasonal trends and synchronisation in cyanobacterial species richness, N fixation, bioavailable N and C fixation (Büdel et al., 2018) provide important contributions to multifunctional microprocesses and soil fertility.

1 Introduction

The northern Australian savannah is one of the largest natural savannahs remaining on Earth with grasslands and shrublands that cover more than 1.5 million km² (Nix et al., 2013). Over the past century there have been several major degradation episodes, leaving about half of these important ecosystems in a degenerated state (Smith et al., 2007). It is a harsh environment where climate shapes the ecology, distribution and abundance of resources, affecting plant and animal species (Nix et al., 2013). There is a pronounced dry season, often lasting around 6 months, followed by violent storms and flooding rains. Across the savannah landscapes broadscale livestock grazing is the primary land use; however, managing these extensive perennial grasslands and woodlands demands an approach on several different scales. On a continental scale, empirical evidence clearly demonstrates the negative impact grazing has exerted on ecosystem structure including key aspects of soil function (Eldridge et al., 2016). Thus, to understand the scope and variability of the northern Australian savannah, soil function is important in the context of a holistic approach to land management (Vanderduys et al., 2012), and more importantly the preservation of soil microprocesses.

In the northern Australian savannah the interspaces between grass plants are inhabited by microorganisms that form an almost continuous cover dominated by cyanobacteria and liverworts, occasionally lichens and mosses, bacteria, algae and fungi (Williams et al., 2014). Cyanobacterial crust communities exist where there are only small fractions of organic nutrients, where diazotrophs (bacteria that fix dinitrogen into a more useable form) fuel soil food webs through photosynthesis and N fixation (Elbert et al., 2012). In these savannah

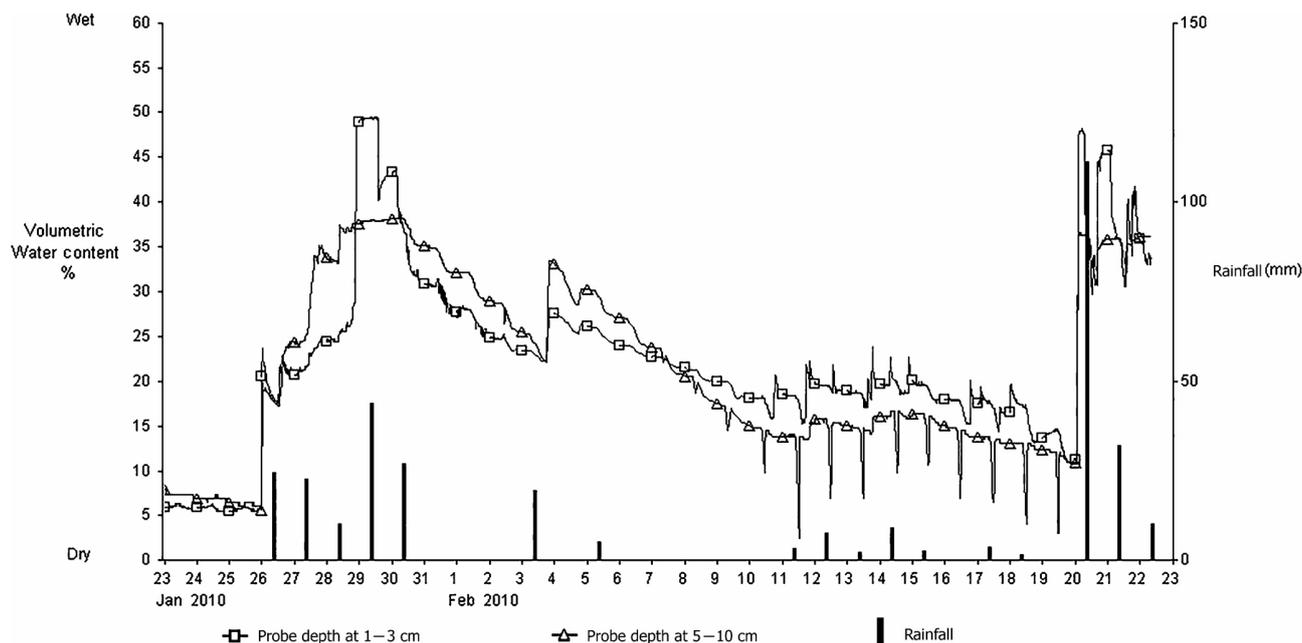


Figure 2. Example of soil moisture (volumetric water content %) patterns plotted against rainfall (mm) for 23 January to 23 February 2010 illustrating the retention of soil moisture after rain. A drying trend following significant rain was intermittently supplemented by smaller rain events, and then after 45 mm of rain on the 20 February the soil profile was recharged from 15 to 50 %. Note: downward spikes for 5–10 cm depth after 11 February are sensor artefacts and should be ignored. Moisture was measured with MEA ThetaProbe and GPI data logger probes (mea.com.au).

rains throughout January (279 mm) and February to April (555 mm) (Fig. 1).

2.2 Field sampling

To determine seasonal patterns of cyanobacterial mediated bioavailable N, multiple sample sets were taken from November 2009 (pre-wet season rains) to May 2010 (end of wet season). Sampling was timed before, during and after major rain events to provide a snapshot in time (also see Williams and Eldridge, 2011), and incorporated cyanobacterial surface crusts (0–1 cm) and immediately below the crust (1–3 cm) ($n = 5$ each depth). Cyanobacterial surface crusts and subsoils were removed using a 10 cm metal spatula to extract exactly 0–1 cm vertical depth with care to include the complete sample, followed by 1–3 cm depth directly below the crust. For analysis we allocated the samples into nine time frames: one sample set for November, December, March, April and May, and two sample sets from early and late January and February. Each time (monthly and twice-monthly) represented at least two separate sample periods, before and after rain. For pigment content (chlorophyll *a*), rates of N fixation (acetylene reduction assays, ARA) and identification (for species richness and abundance studies), an additional four petri dishes of 0–1 cm were collected at the same time (total $n = 100$).

2.3 Laboratory analysis

2.3.1 Seasonal trends in N

Cyanobacterial crusts (0–1 cm) and subsurface soils (1–3 cm) were analysed for bioavailable N (NH_4^+ and NO_3^-) according to Method 4, Gianello and Bremner (1986) and Williams and Eldridge (2011). The samples had been immediately dried in the field ($>40^\circ\text{C}$) and returned to the laboratory where they were stored dry and analysed after the end of the 2010 wet season. Duplicates for each of the nine time periods and each depth (minimum four reps) were sieved (1.86 mm) and weighed (20 mg) for both hot and cold analysis (total across all time periods for each depth, 0–1 cm $n = 125$, 1–3 cm $n = 99$). This method determines ammonium (NH_4^+) produced from organic soil N when the soil is heated with 2 M KCl in a stoppered tube at 95°C for 16 h. NH_4^+ is determined by the difference between the NH_4^+ liberation during the hot distillation of 20 mg soil and the NH_4^+ present prior to heating (Gianello and Bremner, 1986). This analysis provides an index for the fraction of N that would become available to plants in a growing season and at the time of sampling (also see Tongway and Ludwig, 1996; Williams and Eldridge, 2011).

Seasonal N fixation was determined through acetylene reduction based on the Hawkes (2003) method for acetylene reduction assays and Stewart et al. (1968). To complete the monthly estimates of rates of N fixation, $\delta^{15}\text{N}$ of the crust

(see Sect. 2.3.2) was calculated for each sample and used as a conversion factor for each month's results. Petri dish samples of cyanobacterial crusts (reserved for ARA) for each month (November–May) were reactivated in the glasshouse for approximately 2 weeks. Full resurrection during the wet season (when humidity increases) was critical due to the inability of these cyanobacteria to reactivate during the dry season (see Williams et al., 2014). This was carried out by daily wetting to field capacity but not saturated, then allowing them to dry naturally. An effort was made to ensure the surface crust was unbroken. Following reactivation, 18 mm diameter plugs (six reps per month, $n = 36$) representative of 10 % airspace were carefully removed and inserted into 40 mL glass vials with two-way septa lids. Dry weight was calculated prior to a light rewetting (~ 1 mL liquid) with care not to oversaturate. The vials were then placed in natural light conditions in the glasshouse for a further 2 days to acclimatise prior to ARA. The incubation was carried out from time zero (T0) and measured at 24 h (T2) and 48 h (T3). In between measurements the samples were maintained in the glasshouse at 28 °C (previously determined as an optimum temperature for these crusts) and in natural light conditions (Büdel et al., unpublished data).

Calculation of ethylene production ($\text{C}_2\text{H}_4 \mu\text{m mL}^{-1}$) was carried out using the following standard formula (see Hawkes, 2003; Weaver and Danso, 1994):

$$V_{\text{hs}} = V_{\text{t}} - V_{\text{w}} - V_{\text{s}},$$

where V is the volume, V_{hs} is the head space (volume of air in vial), V_{t} is the tube (volume of tube), V_{w} is the volume of water (mL added to sample) and V_{s} is the volume of solids (V_{s} is the weight of sample/soil bulk density for these soils of 1.6). Daily rates were calculated by T3–T2 (48–24 h) and then converted to grams per square metre. Monthly averages for $\delta^{15}\text{N}$ (derived from ARA samples, see Sect. 2.3.2) were then applied as conversion values and care was taken to ensure units were equivalent prior to final calculations. To estimate seasonal N fixation, the mean values of N fixation were calculated for each month and multiplied by photosynthetically active days. As we have clear data that indicate the periods of activity for these biocrusts (see Williams et al., 2014; Büdel et al., 2018), the number of photosynthetically productive days were calculated by the number of rain days plus the number of days that soil moisture was available following rain. This was based on moisture meter data from the site (example shown in Fig. 2) and in situ photosynthetic yield tests previously carried out by Williams et al. (2014).

2.3.2 Carbon, nitrogen and pigment content

Total C and total N, C:N ratio, and $\delta^{15}\text{N}$ and ^{13}C were determined with high temperature digestion using a vario MACRO elemental analyser (Elementar) and the Sercon Hydra 20–22 mass spectrometer (Griffith University laboratories). For each month for ARA analysis the samples were

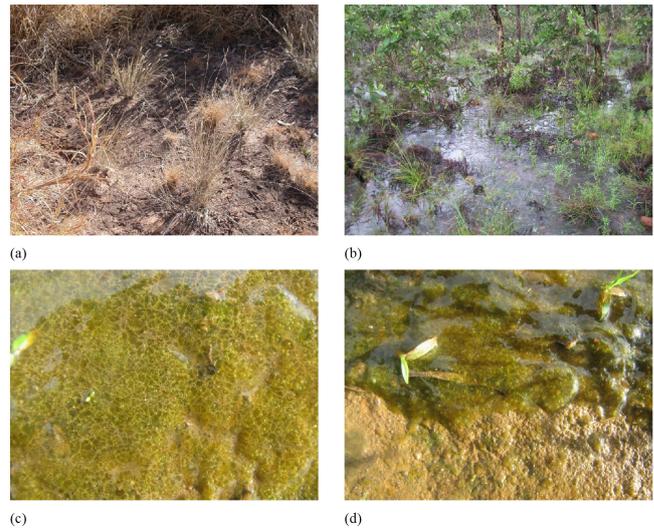


Figure 3. Seasonal cyanobacterial crust functions: (a) dry cyanobacterial crust, (b) flooded crust at the commencement of heavy rains in January, (c) rapid regrowth with extracellular polysaccharide (EPS) hyperproduction from *Nostoc* sp. and (d) gelatinous EPS during hyperproduction phase compared with bare area being recolonised.

amalgamated, dried and sieved to provide three samples for each time period. Data were averaged to provide the conversion factors used in rates of N fixation; however, there was only one replicate available for Nov–Dec as there was insufficient sample. For pigment content, the chlorophyll *a* extractions were carried out on the cyanobacterial soil crusts (Barnes et al., 1992) and calculated using Wellburn's (1994) equations.

2.3.3 Cyanobacterial richness and abundance

Morphological features and measurements were carried out from wet mounts prepared from each sample set for nine time periods (total $n = 625$). Abundances were determined from five subsamples of the five samples for each time ($n = 25$). The samples for the nine time periods were rehydrated for 24 h and examined using bright-field, phase contrast and differential interference contrast illumination systems with a Jena Zeiss and an Olympus BX51 compound microscope to a maximum magnification of 1000x. Photomicrographs were obtained using an Olympus DP12 digital microscope camera. Identification was performed to a species level (wherever possible) in the laboratory using the closest available keys (Anagnostidis and Komárek, 2005; Komárek, 2013; Komárek and Anagnostidis, 1999).

2.3.4 Statistics

We used linear regression models to examine potential relationships between bioavailable N (BioN) and cyanobacterial richness separately for N-fixing (Nrich) and non-N-

Table 1. Cyanobacterial species richness and abundance over 7 months (nine time periods) expressed as a percentage. N-fixing species (bold) were shown only if they produced heterocystes although we also determined (see main text) several non-fixing species were associated with N enrichment.

Cyanobacteria	Nov	Dec	Jan early	Jan late	Feb early	Feb late	Mar	Apr	May
<i>Scytonema</i> sp. 1	50.6	93	78.4	84.4	64	52.8	46.2	32.2	27.2
<i>Scytonema</i> sp. 2			5.8	8.4	20.6	1.8	14.8	21	9.6
<i>Scytonema</i> sp. 3						11.2	5.8	14.6	18
<i>Nostoc commune</i>	12.2	1.4	5.9	4	10.2	10.6	10.4	11.6	17
<i>Nostoc</i> sp. 2						0.8			
<i>Symploca</i> sp.	13	1.4	3.9	0.6	4.6	0.4			3.6
<i>Gloeocapsa</i> sp.							10	0.8	
<i>Symplocastrum</i> sp.	6.4	1.2	2			1.2	0.2	8	7
<i>Schizothrix</i> sp.	4	0.8	3.7	0.2		3.8	2.6	6	
<i>Porphyrosiphon</i> sp. 1	13.8	2.2		1.6	0.4	4.8		1.2	4
<i>Porphyrosiphon</i> sp. 2			0.1	0.2			0.6		0.2
<i>Porphyrosiphon</i> sp. 3								1.4	1.8
<i>Porphyrosiphon</i> sp. 4							1		
<i>Microcoleus vaginatus</i>			0.1	0.6	0.2			0.2	
<i>Microcoleus paludosus</i>						4.8	1.8		1.8
<i>Microcoleus lacustris</i>						7.8	4	0.6	4
<i>Oscillatoria</i> sp.									0.2
<i>Phormidium</i> sp.							2.6	2.4	5.2
<i>Chroococcus</i> sp.			0.1						0.4
N fix	75.8	95.8	94	97.4	99.4	77.6	87.2	80.2	75.4
Non-N fixers	24.2	4.2	6	2.6	0.6	22.4	12.8	19.8	24.6

fixing cyanobacteria (see Table 1). We examined differences in bioavailable N between the two depths across time with mixed-model ANOVA. Our model had two strata: one that accounted for the differences among the nine time periods and a second stratum accounting for depth and its interaction with time. All of these analyses were run in Minitab version 16.1.0 (2010). Least significant difference (LSD) testing was used to examine differences in means among the nine time periods. Tests for homogeneity of variance, independence and normality in the data, using Levene's test and other diagnostic tools in the Minitab (2010) statistical package, indicated that no transformations were necessary.

3 Results

Three species of the nostocalean N-fixing *Scytonema* accounted for 74 % of the biocrust in varying proportions (range 55–93 %) throughout the season (Table 1). Microscopic examination showed *Scytonema* was also the dominant structural component of the biocrust, and this cyanobacterium was found to be the major contributor to the breakdown of the crust and its re-establishment. This took place through the disintegration of EPS and sheath material (November–January); resurrection of a portion of desiccated filaments, followed by mass release of hormogonia (asexual reproductive cells) across January–February; and then vigorous growth of new material (Figs. 3–

Table 2. Pigment content (Chlorophyll *a* = g m^{-2}), total carbon (%C), nitrogen (%N), C:N ratio, and $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ isotopes for the time period from November to May (2009–2010).

	Ca	C	N	C:N	$\delta^{13}\text{C}$	$\delta^{15}\text{N}$
Nov	112.1	0.95	0.06	17	−19.1	2.1
Dec	146.7	0.9	0.06	14.9	−19.1	1.0
Jan	120.7	1.61	0.12	13.9	−17.6	0.9
Feb	171.9	1.8	0.1	17.2	−16.6	0.9
Mar	150.6	1.58	0.09	18.1	−15.5	0.7
Apr	159.5	1.05	0.07	14.1	−18.6	0.3
May	153.8	1.07	0.08	13	−22.2	0.7

4). Average cyanobacterial chlorophyll *a* (pigment content) increased from $112.1 \pm 21.3 \text{ SE } \mu\text{g C a g m}^{-2}$ soil (November) throughout the wet season, peaked in February ($171.9 \pm 2.4 \text{ SE } \mu\text{g C a g m}^{-2}$ soil) and declined towards the end of the wet season ($153.8 \pm 19.9 \text{ SE } \mu\text{g C a g m}^{-2}$) (Table 2).

3.1 Seasonal trends in bioavailable N

Bioavailable N was elevated in November 2009 ($\sim 6 \text{ mg NH}_4^+ \text{ kg}^{-1}$ soil), before falling by almost 50 % across January to February, followed by an exponential increase between March and May when it peaked at $> 13 \text{ mg NH}_4^+ \text{ kg}^{-1}$ soil (Fig. 5). There were significant differences in depth (with



Figure 4. Micrographs of cyanobacterial growth and reproduction, scale bars 20 μm : (a) *Scytonema* sp. with desiccated cells and filaments encased in outer sheath containing a high level of pigmentation (arrows), heterocysts (circled) and heavy cyanophycin granulation; (b) *Scytonema* sp. new growth filaments illustrating hormogonia release (arrow); (c, d) new colonies of *Nostoc* sp. illustrating extracellular polysaccharide (EPS) capsules and surrounding EPS that delivers a microenvironment for other cyanobacteria species cohabitation; (e) mature colonies of *Nostoc* sp. with heterocysts; (f) example of distinct EPS encapsulating *Nostoc* filaments within the overall colonial structure also bound together by EPS.

more in the 0–1 cm layer) and times (all except November); however, the effect of depth was consistent across time but not significant (depth \times time interaction: $P = 0.63$).

3.2 N productivity driven by cyanobacterial richness

Between November 2009 and May 2010 cyanobacterial richness ranged from 6 to 19 species, seven of which were known N fixers (Table 1). Bioavailable N where $\text{BioN} = 1.616 + 0.2072 \text{ Nrich4}$ ($P = 0.001$) was significantly correlated with the presence of four key N-fixing cyanobacteria (where $\text{Nrich4} = \textit{Nostoc commune}$, *Nostoc* sp. 2, *Symploca* and *Gloeocapsa*). Of these four cyanobacteria, *Nostoc* was the most influential where $\text{BioN} = 0.89 + 0.475 \text{ Nostoc}$ ($P = 0.004$). There was a strong positive relationship between total cyanobacterial richness and bioavailable N in the top cm ($F_{1,7} = 39.3$, $P < 0.001$, $R^2 = 0.83$) (Fig. 6).

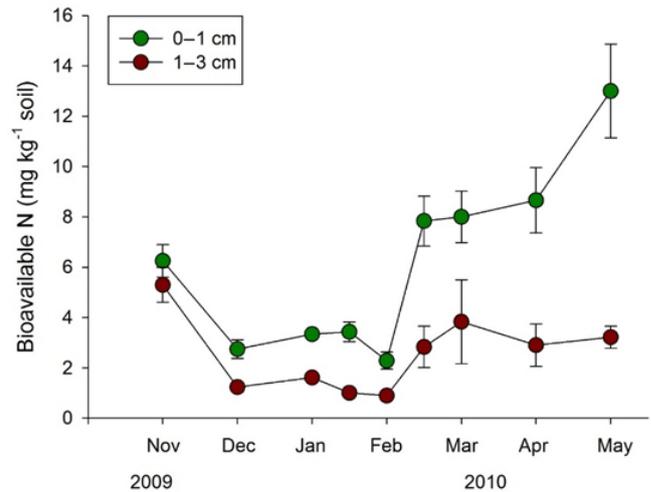


Figure 5. Seasonal trends to bioavailable N over 7 months (November–May 2009–2010) representing nine time periods (early and late January–February) for 0–1 cm and 1–3 cm depths; error bars represent standard error of the mean (SEM).

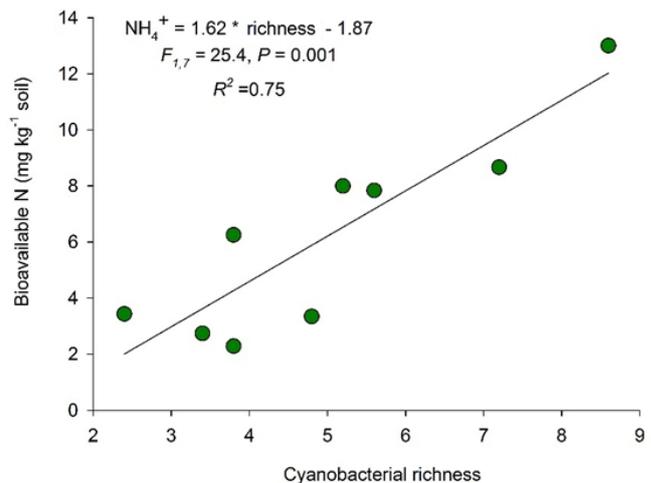


Figure 6. Relationship between bioavailable N and cyanobacterial richness over nine time periods.

There was no relationship between richness and bioavailable N deeper in the profile (1–3 cm; $P = 0.38$). For both N fixers and non-N fixers, increasing richness was associated with an increase in N fixation at a consistent rate (Fig. 7).

3.3 N fixation and chemistry

Rates of N fixation and total seasonal N fixation are reported in Figs. 8 and 9. Rates of N fixation over seven months were significantly correlated with cyanobacterial richness for all species ($P = 0.002$). Average $\delta^{15}\text{N}$ isotope across the season was 0.9 (range -0.2 – 2.1) and $\delta^{13}\text{C}$ isotope was -19.1 to -22.0 with C : N ratios reasonably stable (average 19.1).

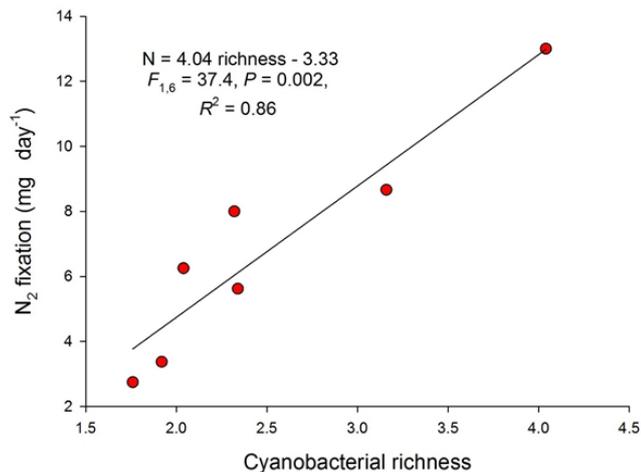


Figure 7. Relationship between cyanobacterial richness and N-fixation rates. As shown, N fixation increases significantly with increases in cyanobacterial richness.

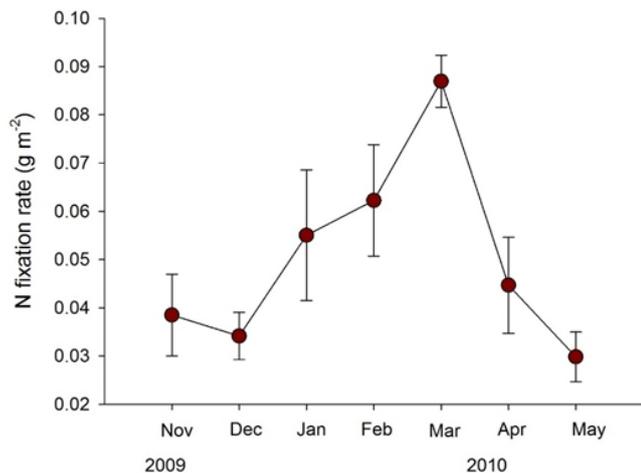


Figure 8. Daily rates of N fixation based on ARA over 7 months; error bars represent standard error of the mean (SEM).

Total C and N both doubled across the course of the wet season (Table 2).

4 Discussion

Atmospheric N₂ transformed to organic N through biological fixation by cyanobacteria has $\delta^{15}\text{N}$ values close to zero (Aranibar et al., 2004; Nadelhoffer and Fry, 1994). Across 7 months of active N fixation, isotopic N values averaged 0.9‰, indicating cyanobacteria were likely the primary source of bioavailable N (Evans and Ehleringer, 1994). Yet, given the rapid wetting following the dry season, potential cell lysis (Williams et al., 2014; Williams and Eldridge, 2011), the presence of other N fixers (Hawkes, 2003) and resource partitioning (e.g. Baran et al., 2015), it is difficult to

separate the exact source of bioavailable N (also see Dojani et al., 2011; Johnson et al., 2007). At the Boodjamulla research sites, bioavailable N was highly correlated with species richness (Fig. 6). This was further underpinned by the analysis of cyanobacterial richness that established key N-fixing species such as *Nostoc commune*, *Symploca* and *Gloeocapsa* significantly enriched soil N. We had hypothesised that N fixation and bioavailable N would pulse at times of high rainfall and gradually decline in the latter stages of the wet season as the rainfall events decreased. For seasonal N fixation the hypothesis was true as it peaked at the height of the wet season. On the other hand, bioavailable N pulsed after the first rains at the beginning of the wet season then declined but was followed up by incremental increases and an exponential rise at the end of the wet season (Fig. 5).

Dark cyanobacterial crusts dominated by species such as *Nostoc*, *Scytonema* and *Microcoleus* that were all influential in the northern savannah crusts are known for their association with high rates of N fixation and absorption (Barger et al., 2016). At the commencement of the wet season, *Scytonema*, due to its macroscopic size and colonial form, was dominant (Table 1). After the first rains in November the crust structure broke down (Williams et al., 2014). Subsequently, bioavailable N was elevated in November, most likely due to the disintegration of the EPS and some cell lysis (Williams et al., 2014) as EPS is known to store N (Otero and Vincenzini, 2003). This was followed by a reduction in bioavailable N after rain in December. We suggest this reflected the favoured investment in C fixation by cyanobacteria (Helm and Potts, 2012) to rebuild their colonies. This was demonstrated later, when in December 2010 to January 2011 there was a net loss in productivity coinciding with rainfall and growth (Büdel et al., 2018). On the other hand, the significant increase in bioavailable N in May appeared to be related to late season rains (in April) indicative of the investment in the storage of N in cyanophycin (granules) and EPS. Other records of seasonal influence on both C and N fixation have been previously demonstrated; however, the synchrony between these events on a monthly and bimonthly basis shows how well balanced the cyanobacterial biophysical and chemical functions were, dictated by rainfall and consequently soil moisture (Büdel et al., 2009; Castillo-Monroy et al., 2010).

We had anticipated that N fixation and bioavailable N enrichment of the surface soils would be correlated to the abundance of N-fixing cyanobacteria. This prediction was true with a significant relationship to both N fixation and bioavailable N at 0–1 cm depth but not for 1–3 cm depth (Fig. 7). This observation points to the importance of biocrusts in the maintenance of N in the soil, thus giving access to this nutrient for the microbial communities in the crust. It had been previously shown that the biocrust had high enzymatic activity relative to the underlying soil (Chen et al., 2014). In this study there was also a significant relationship between non-N-fixing species and N fixation which may be explained from other research: for example, the mutualistic beneficial rela-

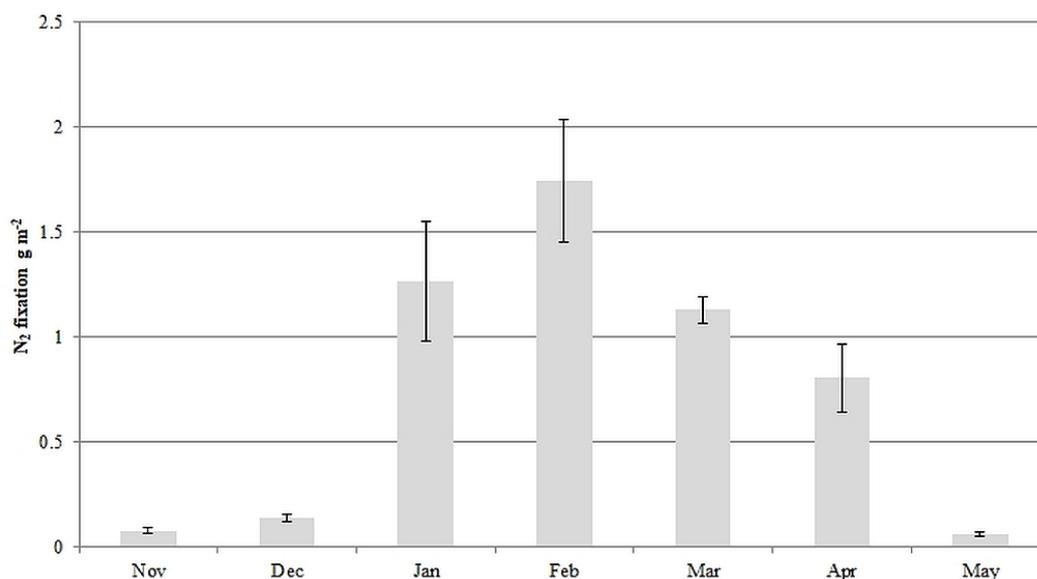


Figure 9. Total cyanobacterial N fixation estimated monthly for 2009–2010 wet season that ranged from 0.08 to 1.74 g m⁻², shown with standard error.

relationship with heterotrophic bacteria that fix N that is subsequently taken up by the non-N-fixing cyanobacterium *Microcoleus vaginatus* (Baran et al., 2015), and the role of bacteria and mycorrhizal fungi in rapid N transformation (Hawkes, 2003). In this study three species of *Microcoleus* were first identified in December but were more prominent between February and May (Table 2), an observation that could provide insight into the relationships with N enrichment and non-N-fixing cyanobacteria. It is now understood that there is a broad range of N-rich metabolites that are continually released and consumed by *Microcoleus* (Baran et al., 2015). It has also been demonstrated that N enrichment was associated with *Gloeocapsa* (Wyatt and Silvey, 1969), *Porphyrosiphon* (Prasanna et al., 2000) and *Schizothrix* (Berrendero et al., 2016). Indeed, many cyanobacteria obtain N by scavenging from mutually shared EPS (Rossi et al., 2017) or have multiple mechanisms for N fixation either in the dark (Lüttge, 1997), through O₂ inhibition (Stal, 1995), in anaerobic circumstances (Murukesan et al., 2016) or aquatic cyanobacterial mats when submerged under water (Berrendero et al., 2016; Stewart, 1980).

4.1 Seasonal trends in N fixation

Total seasonal N fixation by cyanobacteria was based on the number of wet days as well as the follow-on days whilst the soil retained adequate moisture (Fig. 2) that extended the period of photosynthetic activity (Williams et al., 2014). Based on the calculated number of photosynthetically active days each month (also supported by Büdel et al., 2018), we estimated that seasonal N soil enrichment via cyanobacteria would be ~5.2 kg ha⁻¹ (Fig. 9). This estimation must be

treated with caution as in the field there are multiple environmental variables that could result in this figure being higher or lower. Notwithstanding, this indicates an important contribution to the nutrient-deficient savannah soils that are almost entirely reliant on the wet season for microbial turnover of organic matter (Holt and Coventry, 1990). These estimations are comparable to global averages of biocrust N fixation of 6 kg N ha⁻¹ yr⁻¹ (Elbert et al., 2012). There are numerous examples with a broad range of values such as those of cyanobacterial crusts in grasslands from the Loess Plateau in China of 4 kg ha⁻¹ yr⁻¹ (Zhao et al., 2014) or in situ results from the Negev of 10–41 kg ha⁻¹ yr⁻¹ (Russow et al., 2005). Yet, many studies do not take into account a range of mitigating factors or fail to determine the δ¹⁵N conversion factor (Aranibar et al., 2004; Barger et al., 2016). The δ¹⁵N values in this study ranged between 0.3 and 2.1 (Table 2), pointing towards cyanobacterial fixed N₂, whereas theoretical δ¹⁵N conversion rates of 3–4 or higher may be overestimating N production (Barger et al., 2016). Isotopic measurements were taken from the cyanobacterial crusts used in ARA. The conversion rate often creates uncertainty although these values are comparable to other studies that have tested for δ¹⁵N (e.g. Aranibar et al., 2004; Russow et al., 2005). Other limitations of N-fixation estimates lie in the variability of cyanobacterial cover, species richness, and in this study seasonal conditions conducive to *Nostoc commune* productivity and growth.

At the height of the wet season following supersaturation of the soil profile there were two EPS hyperproduction events attributed to *Nostoc commune* (Williams et al., 2014). There is a tight link between rainfall, soil moisture, bioavailable N and EPS excretion that in turn trigger a range of metabolic processes (Chen et al., 2014; Rossi and De Philip-

pis, 2015). C and N fixation in cyanobacteria are closely interconnected as N fixation is energy demanding and dependent on carbohydrates provided by photosynthesis (Murukesan et al., 2016). It has been reported that diazotrophic growth by cyanobacteria occurs when the N to C balance is 1 to 1.5, and the EPS is used as a sink for excess C when the C:N ratio is unbalanced (Otero and Vincenzini, 2004) with C assimilation and diversion to EPS favoured over N fixation (Murukesan et al., 2016; De Philippis et al., 1996; Rossi and De Philippis, 2015). We were unable to make a direct comparison between C:N ratios (see Table 2) and EPS hyperproduction. Nevertheless, some in situ measurements of C fixation at this time (unpublished data) and the following year showed that during storm events optimal temperatures, humidity, moisture and light intensity resulted in extremely high CO₂ uptake (Büdel et al., 2018). With wet season storms this would potentially result in a high C concentration when N could prove a limiting factor. In other research, authors have reported laboratory and field conditions where optimum conditions lead to EPS hyperproduction (Helm and Potts, 2012; Otero and Vincenzini, 2003; Rossi and De Philippis, 2015). This balancing mechanism (Otero and Vincenzini, 2004) could explain the decline in bioavailable N in January–February 2010 at a time when it would be anticipated that a substantial increase in N would occur. In this study *Nostoc commune*, known for its secretion of large amounts of EPS in optimum conditions, was the key species influencing N enrichment, which suggests that *Nostoc* growth and EPS production is an important sequence in the seasonal trends in N bioavailability. EPS creates a microenvironment for the cyanobacterial community that has low oxygen concentrations to enable N fixation under anaerobic conditions (Rossi and De Philippis, 2015).

5 Conclusions

Considering the limited knowledge of N enrichment by both heterocyte-forming cyanobacteria and cyanobacteria that rely on other strategies under different environmental conditions, we need to better understand their function especially in terms of the importance of species richness. Importantly, we discovered that at this site N enrichment was highly correlated to cyanobacterial richness and especially the presence of *Nostoc*. The seasonal patterns in atmospheric N fixation and transformation to a bioavailable form were also present in C fixation results from parallel research for cyanobacterial crusts at the same research site (Büdel et al., 2018). Well-defined seasonal trends and synchronisation in cyanobacterial species richness, N fixation, bioavailable N and C fixation clearly provide significant contributions to multifunctional microprocesses and soil fertility. Further research in the northern Australian savannah is needed to support these findings. Land management practices and climate variability, especially rainfall timing and frequency, could re-

sult in changes to cyanobacterial species richness, which is vital in the sustainability of soil N enrichment, and directly affects ecosystem productivity.

Data availability. Currently data can only be accessed in the form of Excel spreadsheets via the corresponding author.

Sample availability. Limited access to sample material is only available via the corresponding author.

Competing interests. The authors declare that they have no conflict of interest.

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