

Biogenic isoprenoid emissions under drought stress: different responses for isoprene and terpenes

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Abstract. Emissions of volatile organic compounds (VOCs) by biogenic sources depend on different environmental conditions. Besides temperature and photosynthetic active radiation (PAR), the available soil water can be a major factor controlling the emission flux. This factor is expected to become more important under future climate conditions, including prolonged drying-wetting cycles. In this paper we use results of available studies on different tree types to set up a parameterization describing the influence of soil water availability (SWA) on different isoprenoid emission rates. Investigating SWA effects on isoprene (C_5H_8), monoterpene ($C_{10}H_{16}$) and sesquiterpene $(C_{15}H_{24})$ emissions separately, it is obvious that different plant processes seem to control the individual emission fluxes, providing a measure to which plants can react to stresses and interact. The SWA impact on isoprene emissions is well described by a biological growth type curve, while the sum of monoterpenes displays a hydraulic conductivity pattern reflecting the plant's stomata opening. However, emissions of individual monoterpene structures behave differently to the total sum, i.e., the emissions of some increase, whereas others decline at decreasing SWA. In addition to a rather similar behavior to that of monoterpene emissions, total sesquiterpene fluxes of species adapted to drought stress tend to reveal a rise close to the wilting point, protecting against oxidative damages. Considering further VOCs as well, the total sum of VOCs tends to increase at the start of severe drought conditions until resources decline. In contrast to declining soil water availability, OH and ozone reactivity are enhanced. Based on these observations, a set of plant protection mechanisms are displayed for fighting drought stress and imply notable feedbacks on atmospheric processes such as ozone, aerosol particles and cloud properties. With increasing lengths of drought periods, declining storage pools and plant structure effects yield different emission mixtures and strengths. This drought feedback effect is definitely worth consideration in climate feedback descriptions and for accurate climate predictions.

1 Introduction

Emissions of biogenic volatile organic compounds (VOCs, E) are known to represent the largest contribution to global carbon flux besides carbon dioxide (CO₂) (IPCC, 2014; Niinemets et al., 2014; Holopainen et al., 2017). The emissions and the corresponding deposition rates (exchange = emission + deposition) are driven by the temperature of vegetation (T_{veg}) and soil (T_{soil}) , photosynthetic active radiation (PAR), ambient CO₂ mixing ratio, defense against herbivores (Manninen et al., 1998) and reactive air pollutants (Bourtsoukidis et al., 2012), plant-plant communication and competition, fire, and drought (Lappalainen et al., 2009; Peñuelas et al., 2010; Guenther et al., 2012). The total amount of individual BVOC fluxes is linked to their production (de novo synthesis or online) and the storage (offline) capacities of individual plant types and species (Ghirardo et al., 2010) and is additionally affected by abiotic and biotic conditions. These include the temperatures of vegetation (T_{veg}) and soil (T_{soil}) , photosynthetic active radiation (PAR), ambient CO₂ mixing ratio, defense against herbivores (Manninen et al., 1998) and reactive air pollutants (Bourtsoukidis et al., 2012), plant-plant communication and competition, fire, and drought (Lappalainen et al., 2009; Peñuelas et al., 2010; Guenther et al., 2012). While biogenic VOCs (BVOCs) represent a large variability of different structures and species (Goldstein and Galbally, 2007; Laothawornkitkul et al., 2009), it is commonly accepted that isoprene (C_5H_8) and monoterpenes ($C_{10}H_{16}$) display the majority (38 %–50 % and 30 %, respectively) of total BVOC exchange (1–1.3 Pg yr⁻¹) (Goldstein and Galbally, 2007; Guenther et al., 1995, 2012) if methane is excluded. The exchange *E* is commonly described by the exchange at standard conditions E_0 ($T = 30 \,^{\circ}$ C) and scaled by factors for individual driving forces such as temperature (*T*), light (*L*) (Guenther et al., 1995, 2006, 2012) or soil water content (SWC), denoted here as SM in the index in accordance with the formulation of Guenther et al. (2006, 2012):

$$E = E_0 \cdot \gamma_T \cdot \gamma_L \cdot \gamma_{\rm CO_2} \cdot \gamma_{\rm SM}.$$
 (1)

An overview of all the abbreviated terms and parameters can be found in Table 1. Some of the scaling factors of driving forces (γ) for exchange are reasonably well described regarding temperature and PAR, while other parameters like soil water availability (SWA) are mostly ignored due to a lack of understanding of influencing plant processes, although a simplified parameterization exists (Guenther et al., 2012). This neglect is confronted with changing climate conditions, such as warming of 0.15 °C every 20 years (Schölzel and Hense, 2011), with predicted increasingly long drought periods in Central European forest ecosystems such as the Black Forest in southern Germany (Keuler et al., 2016; Kreienkamp et al., 2018). So far, a single simplified formulation using a linear increase from no emission at the permanent wilting point (PWP) to maximum emission shortly above this point is commonly applied for different conditions. This is based on an experiment using poplar (Pegoraro et al., 2004a), analyzing isoprene only. The changing soil water conditions and the resulting change in BVOC exchange fluxes are expected to influence plant responses and protection capacities (Peñuelas et al., 2010; Rennenberg et al., 2006), the exchange following up ozone formation strength, and further climate feedback processes (Bonn, 2014). In this study we aim to investigate the individual behavior of isoprene, monoterpene (MT) and sesquiterpene (SQT) exchange fluxes and their correlation with SWA for available tree species to allow identification of processes controlling their emissions. As European beech (Fagus sylvatica), which is SWA sensitive (Gessler et al., 2007; Dalsgaard et al., 2011), is one of the most common tree species in Central Europe and of economic importance (Kändler and Cullmann, 2016), we will focus on the impact of the SWA effect on emissions.

2 Materials and methods

2.1 Review of available studies and transfer between different water content parameters used

In order to develop an advanced description of the available water on BVOC exchange, available studies were collected, focusing on isoprenoid emissions in the context of different drought conditions. These studies include different tree types and several herbs, predominantly under controlled conditions and for selected compounds or compound groups (Table 2).

2.2 Adapting available studies to a comparative scheme

2.2.1 Different conditions of individual studies

First, available studies on BVOC exchange for different tree species at different soil water conditions have been collected. It is important to note that these studies have been described in various ways using

- a. different soil water parameters (SWA, SWC_v and water matrix potential Ψ_m),
- b. different plant ages,
- c. different soil types,
- d. different lengths of investigation and time resolutions,
- e. with and without rewatering,
- f. different environments (laboratory, greenhouse and ambient).

The named different parameters complicate a direct comparison and include the assumption of an applicable transfer of laboratory or greenhouse experimental results with predominantly seedlings to ambient conditions, a transfer including notable challenges and questions as discussed, e.g., by Niinemets (2010a).

2.2.2 Different parameters used for describing the effect of soil water

The plants access to soil water is best described by the accessible water SWA (index "%": fraction of accessible water, index "v": volumetric amount) or by the suction pressure Ψ_m . Both parameters denote the soil water status and the ability of a plant to extract water from it (Blume et al., 2010). However, the easiest parameter to quantify soil water conditions by measurement continuously is volumetric SWC (SWC_v). Therefore, SWC_v has so far been used to estimate soil water effects on plant processes, e.g., by Guenther et al. (2006, 2012). However, different soil types possess different PWP preventing water being extracted by the plant below. Without the knowledge of the PWP of the soil investigated the information of SWC_v to describe the plant water access remains

Abbreviation	Name	Abbreviation	Name
a, b, c	fitting parameters, indices "g" and "h" reflect growth and hydraulic- conductivity-based fits	NO _x	nitrogen oxides (= $NO + NO_2$)
β	temperature-dependent increase in emission rates	PAR	photosynthetically active radiation
CCN	cloud condensation nuclei	PWP	permanent wilting point
Ε	emission rate of compound (class) $[mgg(dw)^{-1}h^{-1}]$	ROS	reactive oxidation species
E_0	emission rate at standard conditions	RSE	residual standard error
FASW%	fraction of available soil water [%]	SAR	structure-activity relationships
FK	field capacity	SM	soil moisture
$\gamma_{\rm T}, \gamma_{\rm L}, \gamma_{\rm CO_2}, \gamma_{\rm SM}$	emission-regulating factor for the impact of T , light, CO ₂ and soil moisture	SQT	sesquiterpenes (C ₁₅ H ₂₄)

 $\gamma_{\rm SM}$ described by growth or hydraulic

reaction rate constants with OH and

conductivity curve

net field capacity

monoterpenes (C10H15)

ozone

SWA%, SWAv

SWC_v, SWC_m

VPD

 ψ, ψ_{PD}

Table 1. Overview of abbreviated terms and parameter names.

YSM,g,YSM,h

 k_{OH}, k_{O_3}

MT

nFK

incomplete and makes different studies difficult to compare. Therefore, we focus on SWA% as a primary parameter for describing the soil water effect and provide a set of equations to transfer between the different quantities. This makes different amounts of soil water comparable for different soil conditions and furthermore allows for easy usage in model studies that formerly used SWC_v (Guenther et al., 2006, 2012). To support this, all parameterizations are stated by both, i.e., (a) by SWA_% and (b) by SWC_v. The SWA describes the amount of soil water above PWP ($\Psi = -4.2$ MPa) up to field capacity (amount of water held by the soil against gravity, FK, $\Psi = -0.0063$ MPa); this is called available or net field capacity (nFK) for the corresponding soil type and depends on the soil capacity to fix water (van Genuchten et al., 1980). Over the short term, SWA can exceed nFK (e.g., after an intense rain fall), but this excess water infiltrates the soil and is not available to the plant thereafter. A list of the different parameters used in the available literature is given and the transfer in between is assumed to be as follows (Blume et al., 2010):

$$SWA_{\%} = \frac{SWA_{v}}{nEK} = \frac{SWC_{v} - PWP}{nEK},$$
 (2a)

$$SWA_{v} = SWC_{v} - PWP, \tag{2b}$$

$$SWC_v = SWA_v + PWP, \qquad (2c)$$

$$= SWA_{\%} \cdot nFK + PWP, \qquad (2d)$$

$$nFK = max(SWA_v) = FK - PWP,$$
 (2e)

$$FASW_{\%} = SWC_{\%} = \frac{SWC_{v}}{nFK + PWP}.$$
(2f)

 $FASW_{\%}$ and $SWC_{\%}$ abbreviate the volumetric soil water content relative to its maximum value (field capacity FK).

Converting water or matrix potential ψ values is a function of the actual soil mixture of clay, sand and silt, i.e., the different corn size distributions and classes and the inter-corn spaces available for water storage, which can be described by a water retention curve (van Genuchten diagram; van Genuchten et al., 1980; see e.g., Blume et al., 2010):

matrix or water potential [MPa]

soil water availability: percentile [%]

soil water content: volumetric [% vol]

and volumetric [v]

and by mass [mass %]

vapor pressure deficit

$$SWA_{\%} = \frac{1}{nFK}$$

$$\cdot \left(SWC_{v,r} + \frac{SWC_{v,max} - SWC_{v,r}}{1 + (a \mid \Psi \mid \cdot 10^4)^{n1 - \frac{1}{n}}} - PWP\right), \quad (3a)$$

$$SWC_{v,max} - SWC_{v,r} \qquad (3b)$$

SWC_v = SWC_{v,r} +
$$\frac{5 \times C_{v,\text{max}} - 5 \times C_{v,r}}{1 + (a \mid \psi \mid \cdot 10^4)^{n 1 - \frac{1}{n}}}$$
. (3b)

SWC_{v,r} is the residual SWC_v at completely air-dried conditions, SWC_{v,max} is the SWC_v at saturation, α the inverse of the air entry suction and *n* represents a measure of the pore size distribution. Representative data for different soil textures can be found, e.g., in Leij et al. (1996) and at http://soilphysics.okstate.edu/software/water/conductivity. html (last access: 17 January 2019).

2.2.3 Different soil types, corresponding PWPs and FKs

PWP and FK values for different soil types were calculated from the articles information based on data summarized by Chen et al. (2001) and their common soil type classification. Different soil types, predominantly expressed by different corn sizes (corn diameter D_p) and pore space volume, available for soil water, can be classified according to the contribution of sand (0.002 mm $< D_p < 2$ mm), silt (0.002 mm <

$(\phi m m a)$ and ϕ	ламн (рРД ш ин а							
Tree specie	86	Soil type	PWP ass.	VOC species classified	Parameter describ-	Rewa-	Environment	Reference
Latin name	English name		$[m^3 m^{-3}]$		ing water status	tering		
Castanea sativa (M)	sw. chestnut	70 % sand, 30 % organic	2.8	different MTs	SWC_{v}	no	plant chamber	Lüpke et al. (2017a)
Cistus albidus (M)	rockrose	calcareous soil	6.7	MTs, SQTs	ψ , SWC _{vm}	no	pots, greenhouse	Ormeño et al. (2007)
Cistus ladanifer (M)	common gum cistus	clay loam	10.2	Ox. MTs, MTs, SQT:	S_{v}, SWC_{v}, ψ	no	ambient	Haberstroh et al. (2018)
				diterpenes				
Cistus sinensis (M)	orange	clay loam	10.9	trans- β -ocimene (MT), β -caroyphyllen (SQTs)	e SWC _v , ψ	yes	pot, greenhouse	Hansen and Seufert (1999)
Fagus sylvatica (T)	European beech	CPS and fertilizer: 100 % org	6.0	total MTs	stem diameter	yes	growth room	Šimpraga et al. (2011)
Fagus sylvatica (T)	European beech	Bedrock	9.0 ± 1.0	isoprene, MTs, SQTs	+ Leaf water content,	no	ambient	Rombach (2018)
				70 VOCs	ψ_{PD} and ψ , SWC _v nearby			
Pinus sylvestris (T, A)	pine	70 % sand, 30 % humus	2.8	1,8-cineole, differer MTs	lt SWC _v	yes	plant chamber	Lüpke et al. (2017b)
Populus alba (T)	white poplar	CPS: 100 % org	6.0	isoprene	$SWA_{\%}$	yes	growth chamber	Brilli et al. (2007)
Populus deltoides (T)	cottonwood	60 % bare soil, 40 % organic	7.4	isoprene	SWC_v, ψ_{PD}	no	greenhouse	Pegoraro et al. (2004a)
Populus nigra (T)	black poplar	50 % org., 50 % sand	3.5	isoprene	$FASW_{\%}$	yes	growth chamber	Fortunati et al. (2008)
Prunus serotina (T)	cherry	70 %, 30 %	4.7	MTs, SQTs, selecte OVOCs	d SWA _%	no	greenhouse	Bourtsoukidis et al. (2014)
Rosmarinus officialis (M)	rosemary	calcareous soil	6.7	MTs, SQTs	$\psi, \mathrm{SWC}_{\%}$	no	pots, greenhouse	Ormeño et al. (2007)
Quercus ilex (M)	oak	Silty clay loam	13.6	different MTs	¢	yes	ambient	Lavoir et al. (2009)
Quercus robur (M, T)	oak	70~% , $30~%$	4.7	MTs, SQTs, selecte OVOCs	d SWA _%	no	greenhouse	Bourtsoukidis et al. (2014)
Quercus virginiana Mill. (M)	live oak	CPS: 100 % org.	5.0	isoprene	SWC_v, ψ_{PD}	no	phytotron	Pegoraro et al. (2004b)

available soil water (SWA% in %); soil water content by volume (SWCv), mass (SWCm), and ratio of water volume to soil mass (SWCvm); stem diameter; and water potential during Table 2. Overview of studies, tree species, and corresponding conditions, as well as soil water parameters used for deriving the _{YSM} parameterization for isoprene, MT, and SQT emissions. References to the individual studies are listed in the final column. Soil water status parameters include the fraction of available soil water (FASW in %); leaf water content;

CPS: commercial potting soil (e.g., by Agrofino, Miracle Grow) including fertilizer or with additional fertilizer (Basacol); (A): arctic; (M): Mediterranean; (T): temperate.

 $D_{\rm p} < 0.050$ mm) and clay ($D_{\rm p} < 0.002$ mm). Those mimic the potential of the soil texture to fix water, against which the suction pressure of plants needs to work for extracting soil water. This determines not only the PWP of the corresponding soil, i.e., the point at which a plant is unable to suck out any water of the soil pores, but also the maximum amount of water a soil can hold against gravity (FK or SWC_{v,max}). The individual studies published refer to different soil types, PWPs and SWC_{v,max} that need to be considered and included in a more general parameterization. Here, we use the contribution of individual soil types to the mixtures applied in the corresponding studies for calculating the PWP and the SWC_{v,max}. In order to make different soil types and conditions comparable, all soil water describing parameters were converted to SWA_% using PWP and nFK.

As this study is part of a beech and silver fir research project in southwestern Germany, most figures shown are displayed with two *x* axes, (i) the reference SWA[%] in percentage and (ii) an exemplary SWC_v describing the one at the project's field site in Freiamt, southwestern Germany (Magh et al., 2018). This is representative for natural European beech soil common in Central Europe (Gessler et al., 2007) with a PWP of 4.7 % and a nFK of 31.2 %. In order to make results applicable to a wider range, any fitted equations take into account the soil properties and are provided for general conditions as functions of (i) SWA[%] (general reference) and of (ii) SWC_v (requires adaptation to local conditions).

2.3 Different fitting approaches and corresponding driving forces

The effect of soil moisture on the emission of BVOCs is described by Guenther et al. (2006, 2012) using a three-step pattern: assuming (1) no emission below the PWP, (2) a linear increase in emissions between PWP ($\gamma_{SM} = 0$) and PWP + 4 % vol of SWC_v ($\gamma_{SM} = 1$), and (3) a soil-moisture-independent emission above the PWP. This empirical parameterization was based on isoprene emission measurements of Canadian black poplar (*Populus deltoides*) at their Biosphere 2 facility (Pegoraro et al., 2004a). A different suggestion, exponential dependence of emission on SWC_v, was published recently by Genard-Zielinski et al. (2018) for isoprene emissions as well but for downy oak (*Quercus pubescens*) instead of black poplar.

For the overall soil moisture dependency of BVOC emissions at standard conditions (T = 30 °C) several processes become important depending on the molecular size of the compound and its production, storage behavior, and water solubility.

i. *Stomata-controlled effect.* If the size of the BVOC molecule does not allow penetration of the leaf or needle surface layer, the last barrier between plant and atmosphere are the stomata. Thus, the emission process is controlled by stomatal opening behavior of the plant species (Simpson et al., 1985). Low SWA_% and SWC_v

causes low Ψ ; please note the negative scale, which triggers closure of stomata in order to increase the resistance for water molecules between leaf and atmosphere, helping to avoid loss of water by transpiration. Thus, γ_{SM} alters according to the hydraulic conductivity pattern (index "h") given as

$$\gamma_{\text{SM,h}} = a_{\text{h}} + \frac{b_{\text{h}} \cdot \text{SWA}_{\%}}{c_{\text{h}} + \text{SWA}_{\%}}$$
$$= a_{\text{h}} + b_{\text{h}} \cdot \frac{\text{SWC}_{\text{v}} - \text{PWP}}{\frac{c_{\text{h}}}{\text{nFK}} + \text{SWC}_{\text{v}} - \text{PWP}}.$$
(4a)

The curve starts at a_h , i.e., the residual emitted fraction at PWP, and increases to unity as SWA_% reaches 100 % at SWC_{v,max} (FK). Both coefficients b_h and c_h determine the exact shape and slope of increase and are linked via

$$c_{\rm h} = (b_{\rm h} + \text{SWC}_{\rm v}(\text{PWP}) - 1) \cdot (\text{SWC}_{\rm v,max} - \text{PWP}).$$
 (4b)

The shape is therefore different to the approach used in the MEGAN formulation (Guenther et al., 2006, 2012) as the effect sets in already below $SWC_{v,max}$.

ii. *Diffusion-controlled effect.* If the emission can take place at least partially through the cuticula, in smaller chemical species for example, loss of plant water will cause a bending of the tissue surface. This will increase the cell pressure and influence conditions and forces of contained smaller BVOCs, leading to diffusion towards the ambient, i.e., emission. This process acts in a similar manner to biological growth processes and can be described by

$$\gamma_{\text{SM},g} = \exp\left(-\exp\left(b_g \cdot \exp(1) \cdot \left(c_g - \frac{\text{SWC}_v - \text{PWP}}{\text{nFK}} + 1\right)\right)\right)$$
$$= \exp\left(-\exp\left(b_g \cdot \exp(1) \cdot \left(c_g - \text{SWA}_{\%}\right) + 1\right)\right).$$
(5)

Similar to the above examples, b_g as well as c_g represent curve shape parameters. The implicit γ_{SM} value at PWP, i.e., a_g , is given by $\exp(-\exp(b_g \cdot \exp(1) \cdot (c_g) + 1))$.

- iii. Water solubility and transport effect. Further important processes, such as water-dependent productivity or transport (e.g., via sap flow), will display either a linear behavior or a mixture of Eqs. (4a) and (5). This depends on the limitations in the entire process chain from production, storage, potential transport and emission, which may be controlled by stomata opening or diffusion through the cuticula.
- iv. Plant defense or interaction effect. Finally, a rise of BVOC emissions shortly above the PWP is apparent in some MT-related studies but are mainly seen in SQTrelated studies. This may be explained by plant defensive strategies such as detoxification and reduction of

radical oxidative species (ROS) species (Niinemets et al., 2014; Parveen et al., 2018; Piechowiak et al., 2019; Yalcinkaya et al., 2019), as most of these chemical species possess a high reactivity concerning ozone and radicals. The observations indicate an increase with decreasing SWA_% or SWC_v, until there is a rapid collapse close to PWP. We consider these observations as appearing to be like a gamma function with a maximum near the PWP, i.e., $a \cdot (SWA_{\%} + b)^c \cdot \exp(-d \cdot (SWA_{\%} + b))$, with a set of parameters a, b, c and d characteristic for individual plant species. They may potentially reflect a species' ability to respond to oxidative stress.

2.4 Exemplary field studies: Application to ambient conditions

2.4.1 Plant nursery, Freiburg

A total of 11 examples of 8-10 year old Fagus sylvatica seedlings were studied in the plant nursery of the Chair of Ecosystem Physiology (48.014695° N, 7.832494° E) of the Albert-Ludwigs-Universität Freiburg with different soil water potentials, which were determined at predawn and during daytime between 19 and 31 July 2018. Six trees served as control plants, which were watered regularly, while rainwater was excluded by self-constructed roofs above the soil for the other five trees. The distance between the edges of control and stressed group were approximately 10 m with 1 m distance between individual tree stems. Water was added from a watering pot by moderate flows to each of the trees in order to control the amount of water and to minimize the effects on neighboring trees. Note that summer 2018 was extremely dry, with less precipitation (below $35 \,\mathrm{mm}\,\mathrm{m}^{-2}$ during June and July), which is only one-third of the 30-year average $(90.5 \pm 40.5 \text{ mm m}^{-2})$, Freiburg airport, distance: ca. 900 m NNW, German Weather Service. https://opendata.dwd.de/climate environment/CDC/ observations_germany/climate/monthly/kl/, last access: 15 July 2019) for those 2 months. Mean SWC_v of control and stressed beeches were approximated by two methods: (i) the water potential of the experimental plants was determined with a Scholander pressure chamber (Scholander, 1966) at a cut branch before dawn and at around noon, and (ii) SWC_v was estimated using measurements at a nearby field (ca. 600 m NW) with a similar soil structure (VWC, 10HS, Decagon, Washington, USA). BVOC emission was measured with the method described by Haberstroh et al. (2018). For this purpose, BVOCs emitted from beech leaves were collected during daytime using air-sampling tubes filled with Tenax (Gerstel). Analysis occurred with a gas chromatograph-mass spectrometer (GC-MS) system (GC model: 7890 B GC System; MS model: 5975 C VL MSD, with a triple-axis detector, Agilent Technologies, Waldbronn) equipped with a multipurpose sampler (MPS 2, Gerstel, Mülheim, Germany) (Magh et al., 2019). More details of the analysis can be found here (Kleiber et al., 2017).

Based on the observed emission rates and corresponding forest air composition, relative changes in OH and ozone reactivity of the emission cocktail observed were derived as follows (Nölscher et al., 2014; Mogensen et al., 2015). The sum of the individual products of emission rates $[molec. m^{-2}h^{-1}]$ and their corresponding reaction rate constants $[cm^3 molec.^{-1} s^{-1}]$ was calculated and compared to the related sum of the undisturbed reference trees. This includes the knowledge of a large set of compound reaction rates of which some are not obtainable in the available literature. Those have been approximated using structureactivity relationship (SAR)-based algorithms developed by Neeb (2000) and McGillen et al. (2011). In order to do so, the molecule of interest was split in its functional groups and the respective coefficients used for the estimate. As this approach is relative (i.e., disturbed to undisturbed trees), emissions of any vertical mixing are unimportant here but will have to be considered for the detailed impact on atmospheric chemistry, cloud properties and radiation (Seinfeld and Pandis, 2016), and the range of the impact. In this case, it is worth mentioning that forests usually extend over tens of kilometers in the Black Forest area. This indicates the importance of BVOC emissions in local atmospheric chemistry.

2.4.2 Black Forest conditions at Freiamt

SWC_v values were measured at five different depths (5, 10, 25, 50 and 75 cm) with a time resolution of 2 h. Sensors were placed at different locations to test their heterogeneity, especially beneath different tree types, i.e., European beech (*Fagus sylvatica*), silver fir (*Abies alba*) and in between the two species. For further information about these measurements, see Magh et al. (2019). In order to approximate ambient SWC_v effects on isoprene and terpene emissions, Eqs. (1), (4a), (5) and (10b) have been applied to meteorological measurements nearby and SWC_v.

3 Results

As indicated in Sect. 2.3, we standardized and tested the influence of available soil water on emission rates of isoprene (E_{isop}) , MT (E_{MT}) and SQT (E_{SQT}) using different hypotheses: (a) a stepwise effect (Guenther et al., 2006), (b) a growthrate-like behavior (Eq. 5), (c) a hydraulic conductivity pattern (Eq. 4a and b), (d) a stress defense response of SWA_% (lower *x* axis) and, for comparison, the SWC_v for a selected condition (upper *x* axis).

3.1 Isoprene

Most studies of plant BVOC emissions affected by limited amounts of soil water have investigated isoprene. A summary plot of individual studies rescaled to comparable con-



Figure 1. Scaling parameter $\gamma_{SM}(isop)$ for the effect of soil moisture on isoprene emissions as a function of SWA_% (lower x axis) and, for comparison, the SWC_v for a selected condition (upper x axis). Data points represent observations, and different colors refer to different

studies. The dashed black line displays the approach by Guenther et al. (2006, 2012). The red line displays the present form.

ditions is provided in Fig. 1. In Fig. 1, the corresponding SWA_% conditions are plotted vs. γ_{SM} . The upper x axis of Fig. 1 displays typical SWC_v representative of Black Forest conditions at the Freiamt site (PWP = 4.7; nFK = 31.2%; soil: loamy sand; Magh et al., 2018). These conditions are similar to what has been found for European beech forest conditions elsewhere (Dalsgaard et al., 2011). Values of $\gamma_{\rm SM}$ displayed were obtained from measured emission rates (Pegoraro et al., 2004a, b; Brilli et al., 2007; Fortunati et al., 2008; Bourtsoukidis et al., 2014) and were divided by either (a) the corresponding measurements at well-watered standard conditions or (b) by literature values for the individual plant species. More information about the individual datasets and studies is provided in Table 2. Some values such as PWP and nFK are assumed according to the information provided by the authors of the corresponding studies. The figure clearly indicates that most of the studies focused on a specific smaller frame of the entire SWA% range, and only a single value has been measured below the PWP. This smaller range tended to yield a simplified description, e.g., the dataset of Pegoraro et al. (2004a), which was used by Guenther et al. (2006, 2012)(residual standard error (RSE) = 0.163; $r^2 = 0.65$). Testing different models for the

dominating processes, the biological growth curve was found to fit the available datasets best (RSE = 0.06).

$$\gamma_{\text{SM}}(\text{isop, fit, growth}) = \exp(-\exp((0.056 \pm 0.001) \cdot \exp(1)$$

 $\cdot (-(2.3 \pm 1.5) - \text{SWA}_{\%}) + 1))$

(6a)

$$= \exp(-\exp(\frac{(0.056 \pm 0.001)}{nFK}) + \exp(1) \cdot (-(5.4 \pm 3.5)) - (SWC_v - PWP) + 1))$$
(6b)

Apparently, limited soil water access seems to influence isoprene emissions predominantly through growth stress and, to a lesser extent, stomatal opening, although growth stress and stomatal opening share similar features in a plot like Fig. 1, and the entire SWA% range allows discrimination between the two. Isoprene may diffuse out of the plant because of its smaller molecular size. Thus, a nearly closed stomata may not protect the plant from the release of isoprene because of enhanced cellular concentrations at reduced stomatal opening (Simpson et al., 1985; Fall and Monson, 1992).

3.2 Monoterpenes

The situation for monoterpenes (MT) is more complex, as different MT isomers display a different behavior with decreasing SWA% (and SWCv, Fig. 2). For example, in the case of European beech (Fagus sylvatica) the dominant MT sabinene (Moukhtar et al., 2006) declines drastically with reduced water availability, but limonene reduces to a smaller extent, whereas *trans*- β -ocimene stays constant within the uncertainty limit (Rombach, 2018). Lüpke et al. (2017a) report declines in α - and β -pinene, limonene, and myrcene emissions in response to increasing drought stress on Scots pine (*Pinus sylvestris*), while Δ^3 -carene emissions remained unaffected. An isotopic labeling test displayed a tendency for negatively influenced de novo production and contribution with declining SWC_v and thus SWA_%, while emission of stored monoterpenes was less affected. This may explain the overall behavior of MT emission and provide the storage pools to determine the total amount of emissions. If so, any damage to cell walls, or other plant structure elements, and extensive drought length will cause (a) a significant decline in total emissions beyond the drought period in the long term and (b) a decline in the ability of the plant structure to recover, form, and store new MT as before. Individual MT structure behavior, however, will depend on their detailed production pathway and storage location within the plant. Ormeño et al. (2007) describe an increase in MT emissions by rosemary (Rosmarinus officialis) at reduced soil water supply under Mediterranean conditions. The individual contribution of different structures even display a further process to be noted: the relative contribution of α -pinene increased at reduction of SWC_v from 21 % to PWP, dropping thereafter in a similar manner to a stress response until cell damage was observed. A secondary contribution increase at very low SWC_v values observed for sweet chestnut (Castanea sativa; Lüpke et al., 2017b) may result from changes in very small emission amounts and appears as a large uncertainty range awaiting further investigation. The remaining emission flux below PWP may result from de novo production or from different storage locations that are still functioning properly. Emissions of sabinene behave in a very similar way to those of α -pinene, while the fluxes of myrcene transiently increase around PWP. The opposite is true for cineole contribution that drastically declines with reducing SWC_v and displays only at notable soil water presence. MT emissions of Aleppo pine (Pinus halepensis) display a pattern insignificantly influenced by γ_{SM} vs. SWC_v for α -pinene, Δ^3 -carene and linalool, while β -pinene and myrcene tend to increase. Total kermes oak (Quercus coccifera) emissions of MT do not change significantly as well, but the contribution of α -pinene to total MT emissions enhances continuously even beyond the PWP. Therefore, the effect of individual MT structures depends on production rate as well as on storage pools and locations, which differ for different plant species. By changing the MT mixture, plants may adapt to different stress conditions for improved defense (more details in Sect. 3.5). However, the total MT emission effect of limited soil moisture seems to be more general.

A representative plot indicating the influence of SWA_% on total MT emissions is shown in Fig. 2. The setup is identical to Fig. 1 for isoprene, i.e., the different studies included are marked with different symbols and colors. Lighter colors represent conditions after rewetting if performed in the same way as, for instance, for Scots pine (Pinus sylvestris) (Lüpke et al., 2017a). Lighter points indicate a severe plant structure damage and not entirely recovered plants reaching a lower emission intensity, as has been observed in several studies. Excluding those and assuming a soil water supply above PWP, different fits can be applied to the entire dataset. While all data scatter notably, the best performance is obtained with the hydraulic conductivity approach (residual standard error (RSE) = 0.02), which is stomata-opening-controlled. Eventually the biological growth fit (RSE = 0.14) works appropriately if the smallest values are neglected. However, the three-step approach by Guenther et al. (2006, 2012) used for isoprene fails (RSE = 0.69) to appropriately describe the observations. Due to different capabilities of storing MT, commonly denoted as offline and temperature-driven emissions, the lower end of the fit displays a notable variance.

$$\begin{split} \gamma_{\text{SM}}(\text{MT, fit, hyd.}) &= (0.22 \pm 0.05) + (0.78 \pm 0.05) \\ &\cdot (1.3 \pm 0.1) \cdot \frac{\text{SWA}_{\%}}{(33 \pm 3) \,\% + \text{SWA}_{\%}}, \quad (7a) \\ &= (0.22 \pm 0.05) + (0.78 \pm 0.05) \\ &\cdot (1.3 \pm 0.1) \\ &\cdot \frac{(\text{SWC}_{\text{V}} - \text{PWP}))}{\frac{(33 \pm 3) \,\%}{\text{nFK}} + (\text{SWC}_{\text{V}} - \text{PWP})} \quad , \quad (7b) \end{split}$$

 $\gamma_{\rm SM}({\rm MT, fit, growth}) = \exp(-\exp((0.019 \pm 0.004) \cdot \exp(1)))$

$$(-(5.0 \pm 1.2) - SWA_{\%}) + 1))$$
 (8a)

$$= \exp(-\exp(\frac{(0.019 \pm 0.004)}{nFK} \cdot \exp(1))$$

$$\cdot (-(5.0 \pm 1.2)) - (SWC_v - PWP) + 1))$$
 (8b)

Note that the SWA% effect is less punctual at a specific SWA% than in the case of isoprene and the specific slope is soil-dependent, i.e., on nFK and PWP. The decline in emission strength already becomes larger than 5 % by 70 ± 5 % of the accessible SWA_%, depending on the fitting model chosen, not at about 6 ± 2 in the case of isoprene, as also found by Guenther et al. (2012). However, the lower edge of the fitting is different but important, especially for arid regions and their corresponding emissions. On average about $22 \pm 5\%$ of the total MT emissions remain unaffected by soil water content (at PWP).



Soil water dependency of total monoterpene emission

Figure 2. Scaling parameter $\gamma_{SM}(MT)$ for the effect of soil moisture on isoprene emissions as a function of SWA_% (lower *x* axis) and, for comparison, the SWC_v for a selected condition (upper *x* axis). The upper plot (**a**) displays the full data range, while the lower plot (**b**) zooms vertical area between 0 and 1.2. Data points represent observations, and different colors refer to different studies. The dashed black line displays the approach by Guenther et al. (2006, 2012) for isoprene only. The red line displays the present form using hydraulic conductivity as driving force, and the dark red line displays the biological growth stress as driving force.

3.3 Sesquiterpenes

Based on molecular properties, such as saturation vapor pressure (Seinfeld and Pandis, 2016), water solubility (Sander, 2015) and capability of being stored (Kosina et al., 2012), the effect of SWA_% on the emissions of sesquiterpenes (SQT, $C_{15}H_{24}$) (Duhl et al., 2008) behaves the same way as seen for MT (Fig. 3), i.e., emissions are predominantly affected by plant hydraulic conductivity (RSE = 0.153), which is controlled by stomatal opening. Applying the same approach as that for monoterpenes above, yields the following results for the soil moisture effect on total sesquiterpene emission rates $\gamma_{SM}(SQT)$:

$$\gamma_{\rm SM}({\rm SQT, fit, hyd.}) = (0.22 \pm 0.04) + (0.78 \pm 0.04)$$
$$\cdot \frac{(1.75 \pm 0.5) \cdot {\rm SWA}_{\%}}{(81 \pm 57) \% + {\rm SWA}_{\%}} , \quad (9a)$$

$$= (0.22 \pm 0.05) + (0.78 \pm 0.05)$$

$$\cdot (1.75 \pm 0.5)$$

$$\cdot \frac{(SWC_v - PWP)}{\frac{(81\pm57)\%}{n^{FK}} + (SWC_v - PWP)} . (9b)$$

The onset of emission reduction with decreasing SWA_% sets in even earlier compared to MTs (Rombach, 2018), as indicated by the different slopes in Fig. 3 and in Eqs. (9a) and (9b). However, individual structures reveal a large scattering at lower SWA_% and thus SWC_v values (< 10%) that are supposed to result from plant defensive strategies and mechanisms and perhaps from damage to cell walls and membranes.



Soil water dependency of total sesquiterpene emission

Figure 3. Scaling parameter $\gamma_{SM}(SQT)$ for the effect of soil moisture on SQT emissions as a function of SWA_% (lower *x* axis) and, for comparison, the SWC_v for a selected condition (upper *x* axis). Data points represent observations; different colors refer to different studies. The dashed black line displays the approach by Guenther et al. (2006, 2012) for isoprene only. The blue line displays the present form. The pointed green line represents an exemplary fit to $\gamma_{SM}(SQT)$ based on *Cistus* and *Rosmarinus officialis* data.

3.4 Potential additional effects near the wilting point for terpenes

For some plant species, e.g., Cistus albidus and rosemary (Rosmarinus officialis), the effects of limited access to soil water on MT and SQT emission fluxes reveal a second maximum other than the one at $\gamma_{SM} = 1$, appearing near the PWP (Figs. 4, S3-S8 in the Supplement). Highly temperaturestress-adapted and drought-stress-adapted species with the ability of to produce large amounts of terpenoids (e.g., Cistus ladanifer, Haberstroh et al., 2018, and Rosmarinus officialis, Ormeño et al., 2007) display a second maximum close to PWP not only for some MTs (Cistus ladanifer: d-limonene, myrcene, and *trans*- β -ocimene Haberstroh et al., 2018; *Ros*marinus officialis: α -pinene and β -myrcene; Ormeño et al., 2007) but also for some SQTs and diterpenes (DT) (Cistus *ladanifer*: α -cubebene, α -farnesene, ent-16-kaurene, and β chamigrene Haberstroh et al., 2018; Rosmarinus officialis: Δ -cardinene and α -zingiberene Ormeño et al., 2007, Fig. 4) and even the sum of all species. This is apparent, for instance, in the studies of Ormeño et al. (2007) and Haberstroh et al. (2018) of a broader range of plant species. As can be seen in Figs. S3-S8 in the Supplement, there is a statistically significant difference between individual chemical structures for different plant species, indicating a specifically adopted plant species. If those plots are investigated in more detail it is apparent that, for example, the monoterpene α -pinene increases with decreasing SWA_% for *Rosmar*inus officalis (S6) and Quercus coccifera (S5) but persists for Pinus halepensis (S4) (Ormeño et al., 2007). The emission of the monoterpene β -pinene is declining for *Cistus al*bidus (S3), stays put for Rosmarinus officialis (S6), and enhances for Pinus halepensis and Quercus coccifera. Similar observations can be made for SQT emissions with SM values near the PWP: Δ -cadinene emissions increase with decreasing SM, while α -zingiberene emissions decline for Rosmarinus officialis. Haberstroh et al. (2018) monitored decreasing emission fluxes for all of the individual chemical SQT structures for Cistus ladanifer and Quercus robur in a very



Soil water dependency of monoterpene emission

Soil water dependency of sesquiterpene emission



Figure 4. Exemplary scaling parameter γ_{SM} for the effect of soil moisture with a maximum at PWP on individual and total terpene emissions as a function of SWA_{\(\%\)} (lower *x* axis) and, for comparison, the SWC_{\(\yeta\)} at Freiamt soil conditions (upper *x* axis) (*Cistus albidus*) (Ormeño et al., 2007). (a) SM effect on monoterpene emission fluxes of individual structured monoterpenes, as well as of the total sum of monoterpenes. (b) The same data as (a) but for individual structures and the total sum of sesquiterpene emissions (*Rosmarinus officialis*) (Ormeño et al., 2007).

narrow SM range (SWA_% = ca. 0 %–6 %). This is the opposite of the behavior of *Cistus albidus* (Ormeño et al., 2007). Δ -germacrene emissions tend to increase near the PWP for *Cistus albidus*, *Quercus coccifera* and *Rosmarinus officialis* (Figs. S3, S5 and S6).

The enhanced emitted chemical species are known to possess a substantially higher reactivity with respect to ozone, which is usually elevated at drier conditions on longer timescales because of enhanced production rates and accumulation (see Sect. 3.7). For example, α -pinene possesses an O₃ reaction rate constant (k_{O_3}) of 8.66×10^{-17} cm⁻³ molec.⁻¹ s⁻¹, compared to those of β -pinene and Δ^3 -carene of 1.5×10^{-17} and 3.7×10^{-17} cm³ molec.⁻¹ h⁻¹, respectively (Atkinson et al., 1990). With respect to the OH reaction, β -pinene reacts faster ($k_{\text{OH}} = 7.89 \times 10^{-11} \text{ cm}^3 \text{ molec.}^{-1} \text{ h}^{-1}$), followed by α -pinene and Δ^3 -carene, with corresponding rate constants k_{OH} of 5.37×10^{-11} and $0.88 \times 10^{-11} \text{ cm}^3 \text{ molec.}^{-1} \text{ h}^{-1}$, (Atkinson and Arey, 2003). This can be extended for other monoterpenes if more data is available, especially for dlimonene, terpinolene and other very reactive MTs. However, the amount of data for individual species is quite limited and depends on different plant ages, which does not allow for deriving a trustworthy equation at this stage. Sometimes the number of data points is equal to the number of parameters to constrain. The situation looks similar for SM effects on SQT emission rates. Consider, e.g., the bottom plot for SQT in Fig. 4, displaying allo-aromadendrene, Δ -germacrene, Δ -cadinene and α -zingiberene. Reaction rate constants k_{O_3} with respect to ozone are 6.5×10^{-16} , unknown, $3.2 \times$ 10^{-16} and unknown, respectively. The corresponding $k_{\rm OH}$ for OH reactions are 1.5×10^{-10} cm³ molec.⁻¹ h⁻¹ for aromadendrene, with further reactions not accessible based on experiments. However, structure-activity relationships such as that from US EPA (2018), indicate constants of ca. 10^{-10} cm³ molec.⁻¹ h⁻¹. Thus, we assume this secondary maximum to occur as detoxification and defensive strategies of the plants (Kaurinovic et al., 2010; Höferl et al., 2015). Because of notable variations in measurements, this effect is not significant for all the Mediterranean species but only for the Cistus ladanifer and Rosmarinus officialis measurements. To incorporate this feature near the PWP exemplarily, which tends to be linked to higher ozone concentrations, this secondary maximum can be taken into account using a gamma function term added to the hydraulic conductivity curve (RSE = 0.147) in Eq. (10a) and (10b):

$$\gamma_{\rm SM}({\rm SQT, fit, hyd.})^* = (0.22 \pm 0.04) + (0.78 \pm 0.04) \cdot (1.75 \pm 0.5) \cdot \frac{{\rm SWA}_{\%}}{(81 \pm 57)\,\% + {\rm SWA}_{\%}} + a \cdot ({\rm SWA}_{\%} + b)^c \cdot \exp(-d \cdot ({\rm SWA}_{\%} + b)) ,$$
(10a)

$$\gamma_{\rm SM}({\rm SQ1, int, nyd.}) = (0.22 \pm 0.04) + (0.78 \pm 0.04) \\ \cdot ((1.75 \pm 0.5)) \\ \frac{({\rm SWC_v} - {\rm PWP})}{\frac{(81 \pm 57)\%}{{\rm nFK}} + ({\rm SWC_v} - {\rm PWP})} \\ + a \cdot \left(\frac{{\rm SWC_v} - {\rm PWP}}{{\rm nFK}} + b\right)^c \\ \cdot \exp\left(-d \cdot \left(\frac{{\rm SWC_v} - {\rm PWP}}{{\rm nFK}} + b\right)\right).$$
(10b)

a, *b*, *c* and *d* state parameters for the individual response of plant species to severe drought stress (gamma function), and they need to be adapted to the exact plant or ecosystem response studied on the local scale. An example is shown by the green line in Fig. 3 (applied parameters in here for *Cistus ladanifer* and *Rosmarinus officialis*: a = 0.4; b = 1.57; c = 0.6; d = 1). It is of interest that different structures seem to behave differently according to their historical stress adaptation, as presented by Fortunati et al. (2008) and Lüpke et al. (2016). In contrast to the hydraulic-conductivity-related description, the stepwise approach (Guenther et al., 2006) applied for isoprene fails at reproducing the declining pattern accurately (RSE = 0.895).

3.5 Further BVOC emission rates

In addition to isoprene, MT and SQT, other VOCs are released by plants in notable amounts, but the impact of soil moisture on them is even less studied. Because of the huge number of different oxidized species and related publications, including those within the parameterization would extend this study greatly. A snapshot of individual correlation coefficients of different compound groups and species can be seen in Fig. S1. A non-negligible role of those larger VOCs seems plausible, either via formation pathway or oxidation processes within the plant. Those species supply further molecules that can buffer oxidative stresses onto the plant's internal processes. Additionally, they have substantial implications especially for OH reactivity (Atkinson et al., 1990; Atkinson and Arey, 2003; Neeb, 2000; Müenz, 2010; Seinfeld and Pandis, 2016; US EPA, 2018). Overall, there is a tendency (Rombach, 2018) for an increased total BVOC emission under drought-stressed conditions, but there is a large scatter of individual species reactions. However, the implications for OH reactivity and thus atmospheric chemistry are of relevance, as this parameter describes the plant's ability to still counteract oxidative damages by BVOC emissions.

3.6 Effects on reactivity

As chemical oxidation and ROS cause substantial damage to plants (Loreto and Velikova, 2001; Oikeawa and Lerdau, 2013), especially at stressed conditions such as warm and dry conditions, any strategy of plants to counteract or detoxify will be beneficial for survival and plant competition. An indication for this may be seen in the local emission increase in SQT near PWP for some plant species. Saunier et al. (2017) found a stable ratio of catabolic to anabolic BVOC emissions, i.e., emission fluxes related to oxidation stress and to non-oxidative stress production, respectively, during recurring drought periods for downy oak (Quercus pubescens) at Mediterranean conditions. This is different for different plant types and their adaptation to stress conditions (Niinemets, 2010a, b), i.e., the need to compensate for oxidative stress for survival. Looking at the pattern of absolute and relative MT and SQT emissions with declining SWA% values, it becomes obvious that more reactive structures tend to increase until plant cell damage occurs (Ormeño et al., 2007; Bourtsoukidis et al., 2012; Nölscher et al., 2016). SQT emissions displaying this transient peak around PWP especially cause a remarkable change. For example, reducing the SWC_v for European beech by approximately 3 % vol close to PWP increases the total emission of SQTs by 170 %. The individual SQT species, however, show different patterns (Table 3). While emission of junipene declines (-50%), those of longifolene stays constant within the uncertainty limits and the corresponding fluxes of isolongifolene, α -bergamotene and α -farnesene increase drastically

Species	SW	Ά‰	Total amount $[\mu g g(dw)^{-1} h^{-1}]$		Rel. amount [%]	
	А	В	A	В	A	В
Monoterpenes	6_4^8	3 ⁵ ₁	$11.7^{35.7}_{2.8}$	$0.8^{2.4}_{0.4}$		100
Sabinene	6_{4}^{8}	3^{5}_{1}	$11.2^{21.9}_{2.6}$	$0.2^{3.0}_{0.1}$	91 ⁹⁵	37^{75}_{23}
Limonene	6_4^{8}	3^{5}_{1}	$0.3^{1.5}_{0.1}$	$0.1^{0.3}_{0.1}$	5^{9}_{3}	$12_{8}^{\tilde{1}\tilde{9}}$
<i>trans</i> - β -ocimene	6_4^{8}	3 ⁵ ₁	$0.2_{0.0}^{1.2}$	$0.2_{0.1}^{0.9}$	1^{10}_{0}	39_{10}^{64}
Sesquiterpenes	6_{4}^{8}	3_1^5	$10.43_{2.88}^{17.23}$	$27.65_{12.12}^{42.27}$		100
Isolongifolene	6_{4}^{8}	3^{5}_{1}	$9.10^{15.57}_{2.22}$	$12.43_{6.60}^{30.17}$	8863	82_{35}^{97}
Longifolene	6^{8}_{4}	3^{5}_{1}	0.390.79	$0.30_{0.20}^{1.56}$	4_{2}^{12}	1^{6}_{1}
α -farnesene	$6_{4}^{\dot{8}}$	$3^{\hat{5}}_{1}$	$0.34_{0.05}^{0.57}$	$2.64_{0.21}^{6.56}$	$\bar{3}_{1}^{7}$	9_1^{49}
Junipene	6^{8}_{4}	3^{5}_{1}	$0.12_{0.04}^{0.15}$	$0.06_{0.04}^{0.50}$	1^{2}_{1}	16 10
α -bergamotene	6^{8}_{4}	3^{5}_{1}	$0.05_{0.02}^{0.07}$	$0.29_{0.05}^{0.77}$	$ 0_0^{1}$	1_{0}^{6}

Table 3. Change for individual monoterpene and sesquiterpene species emission fluxes of *Fagus sylvatica* for two SM conditions (Rombach, 2018). Because of abnormally distributed data, values are stated as the median, 25th percentile (lower) and 75th (upper) percentile.

(+50%, +500% and +700%). Other VOCs analyzed (total: 81 VOCs; Rombach, 2018) operate in various ways: sabinene emission nearly vanishes at PWP. In contrast, emission of α -terpineol increased. Overall, the total emissions of VOCs and C increased tentatively (Fig. 5). This caused total OH reactivity and especially ozone reactivity to tentatively enhance (OH: $+39^{+88}_{-46}$ %; O₃: $+131^{+265}_{-104}$ %) at a SWC_v change of approximately 3% vol. In this way the total amount, the changing composition and reactivity provide a tool for plants in this example for European beech to counteract oxidative stress and damage. The median change in OH reactivity corresponds nicely to the increase in total carbon emitted, as most organic species react with OH. However, the change in ozone reactivity is primarily related to the changes in individual SQT emissions. The large uncertainty in this context results from partially unknown reaction rate constants estimated either from similarity or from structure-activity relationships (SAR) (Neeb, 2000; McGillen et al., 2011). Since the storage pools, especially of larger terpenes, will be depleted with prolonged duration of drought, this effect will weaken over time.

3.7 Application to atmospheric conditions: ambient SWC_v and its effect on emission rates at the reference site

In order to test the impact of the described effects of SWA_% and thus SWC_v on isoprenoid emissions, we apply the following emission flux parameterizations for isoprenoids from European beech and silver fir, which were quantified in earlier studies in this region excluding the SWA_% effect (Moukhtar et al., 2005, 2006; Šimpraga et al., 2011; Bonn et al., 2017).

$$E_{\rm isop}(Fs) = 0\,\mu g\,g(dw)^{-1}\,h^{-1} \cdot C_{\rm T} \cdot C_{\rm L}$$
(11a)

$$E_{MT}(Fs) = 43.5 \,\mu g \,g(dw)^{-1} \,h^{-1} \cdot C_T \cdot C_L$$
(11b)
$$E_{SQT}(Fs) = 0.04 \,\mu g \,g(dw)^{-1} \,h^{-1}$$

$$\cdot \exp(0.11 \,^{\circ}\mathrm{C}^{-1} \cdot (T_{\mathrm{L}} - 30 \,^{\circ}\mathrm{C}))$$
 (11c)

$$E_{\rm isop}(Aa) = 0.038\,\mu g\,{\rm g}(dw)^{-1}\,{\rm h}^{-1} \cdot C_{\rm T} \cdot C_{\rm L} \tag{11d}$$

$$E_{\rm MT}(Aa) = 28.8\,\mu g\,g(dw)^{-1}\,h^{-1} \cdot (0.71 \cdot \exp(0.14\,^{\circ}{\rm C}^{-1} \\ \cdot (T_{\rm L} - 30\,^{\circ}{\rm C}))) + 0.29 \cdot C_{\rm T} \cdot C_{\rm L})$$
(11e)

$$E_{\text{SQT}}(\text{Aa}) = 0.13 \,\mu\text{g}\,\text{g}(\text{dw})^{-1}\,\text{h}^{-1} \cdot \exp(0.04\,^{\circ}\text{C}^{-1}$$
$$\cdot (T_{\text{L}} - 30\,^{\circ}\text{C}))) \tag{11f}$$

SM effects on specific emission rates have been calculated by including and excluding the individual γ_{SM} factors as multiplicands on the results. In Eq. (11a–f) C_{T} and C_{L} abbreviate the temperature- and light-scaling factors of Guenther et al. (1995), as well as T_{L} the temperature of the needle or leaf surface in °C.

The base effect of soil water on isoprenoid emission rates can be exemplarily shown assuming a temperature of 30 °C, i.e., standard conditions (see Fig. S2 in the Supplement). The available soil water for a plant has different impacts at distinct SWA_% values. At a SWA_% (i.e., SWC_v (Freiamt) = 20%) of 35%, isoprene emission rates display no decline, while the corresponding rates of MT and SQT reduce by 15 % and 25 %, respectively. Assuming an SWA $_{\%}$ of 10 % $(SWC_v = 10\%$ with a PWP of 4.8% at Freiamt), isoprene, MT and SQT emission rates represent only 95%, 50% and 40% of the standard rates, respectively. Therefore, the amount of soil water has the strongest effects on the larger terpenes, shrinking with molecular size. Isoprene emissions will stay unaffected longer than MT and will keep ozone production high if isoprene is the dominant class of isoprenoids released by the ecosystems in the nearby region. Depending on the tree-specific emission, i.e., online- and offline-



Figure 5. Change for the total sum of VOCs (left) and total carbon (right) emission fluxes of *Fagus sylvatica* for two SM conditions (Rombach, 2018). Because of the abnormally distributed data, values are stated as the median, 25th percentile (lower) and 75th (upper) percentile.



Figure 6. Calculated mean isoprene and monoterpene emissions under different temperature and drought stress conditions using the parameterizations of Guenther et al. (2012) and the SWC dependency derived in this study. "Black Forest region" represents the forest inventory mean tree species mixture of the Black Forest area: (a) total isoprene emission, (b) total monoterpene emission, and baseline *Fagus sylvatica* and *Abies alba* MT emission rates (c and d).



Figure 7. Calculated γ_{SM} effects during 2016 at Freiamt on isoprene, total MT and SQT emission rates. Note that there were two induced drought periods during the summer and quite saturated conditions otherwise. The applied drought period is indicated by a reddish background.

including temperature effects, different locations of emission optima or maxima are found (Fig. 6). The predominantly de novo-emitting European beech will emit the most close to 30 °C, while the maxima of silver fir emission rates are shifted to larger temperatures. Similarly, the effects of changing SM will occur at the very same temperatures.

The overall impact of variable SM pattern on isoprenoid emissions was investigated using exemplary SWC_v data at five different depths. Those measurements were located within the reference stand of a European beech (*Fagus sylvatica*, Fs) and silver fir (*Abies alba*, Aa) of the BuTaKli project in Freiamt within the Black Forest (southwestern Germany) region (Magh et al., 2018). In general, the uppermost layer dries out fastest, especially during summertime, while lower layers, i.e., below 30 cm, retain a higher SWC_v for a longer time period. This implies different effects on the tree species with different rooting depths and thus access to soil water. Regarding the effect of SM on isoprenoid emissions, i.e., the γ factor, the γ_{SM} values for one vegetation period (during 2016) are shown in Fig. 7. Individual plots represent the SWC_v emission effects of different compounds or compound classes and different depths. Apparently during summertime, when the evapotranspiration flux and the emission rates

are highest due to elevated temperatures and radiation intensity, the effect is most intense. During the rather wellwatered summer 2016, $\gamma_{SM}(\text{isop})$ caused reduction by almost 40 % for more shallow-rooting species and by 20 % to 30 % for trees with access to deeper soil layers (60 cm in depth). $\gamma_{SM}(\text{MT})$ declined to nearly 55 % on average with only limited access to more water for deeper roots and $\gamma_{SM}(\text{SQT})$ decreased to nearly 60 % of emission rates.

During 2017, with an improved roof cover, the differences between different soil depths became more obvious, and they changed by 10 % or more between 40 and 60 cm in depth (not shown, Magh et al., 2019). Those effects are estimated to intensify remarkably during extensive drought periods such as those that occurred in summer 2018, of which no dataset is available. This is going to reduce most of the BVOC emission rates over time, progressing through the drought period, depending on the tree species rooting depth and soil water access. Therefore, the total sum of isoprenoid emissions of a specific forest ecosystem will be reflected by the individual contribution of tree species and their specific emission patterns, and the sum of SM affects the drought sensitivity of the ecosystem.

4 Conclusions

In this study we derived three sets of parameterizations for the influence of soil moisture on isoprenoid emission rates. These depend on the details of soil characteristics, plant properties of online and offline production of isoprenoids and potential defensive plant feedback mechanisms against oxidative stress. The effect on isoprene emissions acts similarly to a small molecule able to penetrate the plant surface, acting like a biological growth stress. While the cuticula provides some minor resistance it cannot hinder isoprene from penetration at nearly closed stomata. Thus, the present description allows minor isoprene emissions below the PWP. The influence of soil water access on larger isoprenoid emission occurs on a notably larger range of SWA% as a function of stomata opening (hydraulic conductivity pattern). Depending on the specific tree species, some molecular structures of terpenes, e.g., the more reactive ones like α -pinene, β -myrcene (Haberstroh et al., 2018; Ormeño et al., 2007) and α -farmesene (Duhl et al., 2008; Haberstroh et al., 2018) display an elevated relative terpene flux contribution at lower soil moisture (Figs. S3-S8), indicating a defensive strategy of plants against oxidative stress at lower soil water concentrations, and thus warmer conditions, that is critical for the plant's survival. This is plant species dependent and can be incorporated by a secondary maximum close to the permanent wilting point. The named SWA% impact on isoprenoid emission rates is expected to induce several feedback processes in the atmosphere and climate system, including local tropospheric ozone production and sink (Seinfeld and Pandis, 2016), new particle formation (Bonn et al., 2014; Bonn, 2014) and growth, and cloud condensation nuclei (Bonn, 2014; Seinfeld and Pandis, 2016). With these findings we conclude that more studies on available soil water dependence of emissions of higher terpenes, oxygenates and follow-up processes in different environments are required to realistically describe the natural feedback processes of forests and ecosystems in changing climate conditions and the anthropogenic impact on these ecosystems. Because of this, the SWA% concepts described should definitely be included in corresponding global atmospheric chemistry transport models in order to cover not only plant response to the expected increase in drought periods but also future development of biosphere–atmosphere processes, resulting in adaptation or extinction in a changing climate.

Data availability. The datasets have been provided as a zipped set of Excel files and pdf within the Supplement. Of these, one Excel file corresponds to the manuscripts datasets and one to the datasets displayed in the supplementary pdf-file.

Supplement. The supplement related to this article is available online at: https://doi.org/10.5194/bg-16-4627-2019-supplement.

Author contributions. BB was responsible for the methodology, analysis and writing. RKM conducted the SWC_v measurements and soil texture analysis, contributed to the data analysis and parameterization approach, and was responsible for the soil water concepts. The formal analysis of *Fagus sylvatica* BVOC emissions and water potential measurements at the campus field site was done by JR. JK contributed to the conceptualization of the *Fagus sylvatica* campus field site experiment and discussion about accurate plant physiology process description, as well as by reviewing and editing the paper.

Competing interests. The authors declare that they have no conflict of interest.

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