



Distribution and behaviour of dissolved selenium in tropical peatland-draining rivers and estuaries of Malaysia

Yan Chang¹, Moritz Müller², Ying Wu¹, Shan Jiang¹, Wan Wan Cao¹, Jian Guo Qu¹, Jing Ling Ren³, Xiao Na Wang¹, En Ming Rao³, Xiao Lu Wang¹, Aazani Mujahid⁴, Mohd Fakharuddin Muhamad², Edwin Sien Aun Sia², Faddrine Holt Ajon Jang², and Jing Zhang¹

¹State Key Laboratory of Estuarine and Coastal Research, East China Normal University, Shanghai 200062, China

²Faculty of Engineering, Computing and Science Swinburne, University of Technology, Kuching 93350, Malaysia

³Key Laboratory of Marine Chemistry Theory and Technology, Ministry of Education, Ocean University of China, Qingdao 266100, China

⁴Faculty of Resource Science and Technology, Universiti Malaysia Sarawak, Kuching 93350, Malaysia

Correspondence: Yan Chang (ychang@sklec.ecnu.edu.cn)

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Abstract. Selenium (Se) is an essential micronutrient for aquatic organisms. Despite its importance, our current knowledge of the biogeochemical cycling of dissolved Se in tropical estuaries is limited, especially in Southeast Asia. To gain insights into Se cycling in tropical peat-draining rivers and estuaries, samples were collected from the Rajang, Maludam, Sebuyau, Simunjan, Sematan, Samunsam and Lunda rivers and estuaries in western Sarawak, Malaysia, in March and September 2017 and analysed for various forms of Se (dissolved inorganic and organic). Mean total dissolved Se (TDSe), dissolved inorganic Se (DISe) and dissolved organic Se concentrations (DOSe) were 2.2 nmol L^{-1} (range: 0.7 to 5.7 nmol L^{-1}), 0.18 nmol L^{-1} (range: less than the detection limit to 0.47 nmol L^{-1}) and 2.0 nmol L^{-1} (range: 0.42 to 5.7 nmol L^{-1}), respectively. In acidic, low-oxygen, organic-rich blackwater (peatland-draining) rivers, the concentrations of DISe were extremely low (near or below the detection limit, i.e. $0.0063 \text{ nmol L}^{-1}$), whereas those of DOSe were high. In rivers and estuaries that drained peatland, DOSe / TDSe ratios ranged from 0.67 to 0.99, showing that DOSe dominated. The positive relationship between DISe and salinity and the negative relationship between DOSe and salinity indicate marine and terrestrial origins of DISe and DOSe, respectively. The positive correlations of DOSe with the humification index and humic-like chromophoric dissolved organic matter components in freshwater river reaches suggest that peat soils are probably the main source of DOSe.

The DOSe fractions may be associated with high molecular weight peatland-derived aromatic and black carbon compounds and may photodegrade to more bioavailable forms once transported to coastal waters. The TDSe flux delivered by the peat-draining rivers exceeded those reported for other small rivers and is quantitatively more significant than previously thought.

1 Introduction

Selenium (Se) is an essential trace element for aquatic organisms (Bodnar et al., 2014). Low levels of Se in the food chain lead to disease or death, whereas high levels are toxic (Lobanov et al., 2009; Winkel et al., 2015). Selenium depletion in the Phanerozoic oceans may have contributed to three major mass extinction scenarios (Long et al., 2016). Thus, there has been great interest in Se biogeochemical cycling in aquatic systems for many decades (e.g. Cutter and Bruland, 1984; Cutter and Cutter, 1995, 2001; Mason et al., 2018).

The bioavailability of Se is determined by its concentrations and species (Fernández and Charlet, 2009). The behaviour of Se in natural waters is complicated, as it exists in several oxidation states (–II, IV, VI) (Conde and Sanz Alaejos, 1997). A number of field and laboratory studies have found that selenite (Se(IV)) and selenate (Se(VI)) can be assimilated by marine phytoplankton with Se(IV) being

the preferred species (Wrench and Measures, 1982; Apte et al., 1986; Vandermeulen and Foda, 1988; Baines and Fisher, 2001). Direct uptake of selenomethionine and selenocystine has been demonstrated in diatoms (Doblin et al., 1999). Laboratory study showed that organic selenide released by algal cell lysis of the diatom or viral of the chrysophyte were also bioavailable to the marine phytoplankton (Gobler et al., 1997; Baines et al., 2001). An understanding of Se speciation may therefore be important for determining the bioavailability of Se that is transported from land to oceans.

The chemical behaviour of Se in estuarine mixing plays an important role in overall geochemical cycling. From their investigation into dissolved Se species in various estuaries, Chang et al. (2016) found that Se speciation was controlled by biological, physical and redox processes in the estuaries; non-conservative processes resulting from phytoplankton uptake; absorption by suspended particles; and generation of particulate organic selenide in the water. Thus far, the behaviour of Se in estuaries has been studied mainly in the temperate zone of the Northern Hemisphere (between 20 and 60° N) (Measures and Burton, 1978; Takayanagi and Wong, 1984; Van der Sloot et al., 1985; Cutter, 1989a; Guan and Martin, 1991; Hung and Shy, 1995; Abdel-Moati, 1998; Yao et al., 2006; Chang et al., 2016). The behaviour of Se in tropical estuaries, however, is still poorly understood.

In the high-latitude peatland-draining rivers, dissolved Se concentrations are spatially variable, with concentrations of up to 13 nmol L⁻¹ being observed in northern Minnesota, USA (Clausen and Brooks, 1983), from 0.38 to 5 nmol L⁻¹ in the Krycklan catchment, Sweden (Lidman et al., 2011), and from 0.25 to 1.25 nmol L⁻¹ in the lakes and rivers of western Siberia, Russia (Pokrovsky et al., 2018). Although these various studies did not report different species of Se (Clausen and Brooks, 1983; Lidman et al., 2011; Pokrovsky et al., 2018), the dissolved organic Se concentrations (DOSe) is probably the dominant species in peatland-draining river. In the open ocean, DOSe was assumed to mainly associate with soluble peptides with low molecular weight in surface waters and to be relatively refractory (Cutter and Cutter, 1995, 2004). Substantial amounts of dissolved Se also are known to be associated humic substances, Gustafsson and Johnsson (1994) assumed that Se was preferentially incorporated into low molecular weight humic substances fractions by means of microbial reductive incorporation, while Kamei-Ishikawa et al. (2008) found that Se associated with high molecular weights humic acid fractions. The current paucity of information about DOSe characteristics and its export by tropical peat-draining rivers remains a major gap in our understanding of Se biogeochemical cycling. The highest concentrations of dissolved organic carbon (DOC) globally were reported in tropical peat-draining rivers in Borneo (Moore et al., 2013; Wit et al., 2015). More works of Se in the fluvial systems of this region are therefore needed to provide an improved understanding of the biogeochemical processing of Se and its associations with organic matter.

Table 1. The physico-geographical parameters of sampled rivers.

River names	Total basin ^a	Runoff (km ³ yr ⁻¹)	Peatland cover (%) ^a	Degree affected by palm plantations (%) ^a
Rajang	50 000	114 ^b	7.7	9.1
Maludam	197	0.14 ^c	87	8.1
Sebuyau	538	NA	54	4.5
Simunjan	788	NA	44	30
Samunsam	163	NA	10	0
Sematan	287	NA	0	0

^a Modified from Bange et al. (2019). ^b Cited from Staub et al. (2000).

^c Cited from Müller et al. (2016). NA – not available.

To the best of our knowledge, the present study is the first analysis of the distribution and behaviour of dissolved species of Se in seven rivers and estuaries in western Borneo (Sarawak, Malaysia, Southeast Asia). We hypothesize that DOSe is the major species in those peatland-draining rivers that come mainly from peat soils and that sizable Se from peatland is delivered to the coastal areas. The main objectives of the study were to (1) evaluate the fate of dissolved Se species in peatland-draining estuaries, (2) characterize the DOSe fractions and (3) estimate the magnitude of Se fluxes delivered from to coastal ocean. The results of this study should contribute to an improved understanding of how Se behaves in tropical peat-draining rivers and estuaries.

2 Materials and methods

2.1 Study areas and sample collection

Sarawak, Malaysia's largest state, is in the northwest of the island of Borneo, Malaysia (Müller et al., 2016). Sarawak has a tropical climate, with a mean annual air temperature at the capital Kuching of 26.1 °C (Müller et al., 2016). Rainfall is abundant throughout the year but is pronounced during the northeastern monsoon, which occurs between November and February (wet season). The period from May to September, before the southwestern monsoon, constitutes the dry season (Sa'adi et al., 2017). About 12 % of the coastal area of western Sarawak is covered by peatlands, of which approximately 41 % has been converted to palm plantations (Müller et al., 2016).

Two sampling campaigns were conducted in peat-draining rivers and estuaries in Sarawak in 2017. The first was at the end of the northeastern monsoon (from 4 to 12 March 2017, just after the wet season), and the second was shortly before the beginning of the southwestern monsoon (from 4 to 17 September 2017, in the dry season) (Fig. 1). Six rivers, namely the Rajang, Maludam, Simunjan, Sebuyau, Sematan and Samunsam, were sampled in March and September, and the Lundu River was sampled only in September (Fig. 1). The physico-geographical parameters of sampled river basins are

summarized in Table 1. Four of the rivers (the Maludam, Simunjan, Sebuyau and Samunsam) drain catchments with high peatland coverages and are known as blackwater rivers, whereas the Sematan and Lundu drain catchments with high proportions of mineral soils (Table 1). The Rajang River drains mineral soils in its upper reaches (Staub et al., 2000) but branches into multiple distributary channels at Sibul (the Igan, Paloh and Rajang) that flow from north to south through land covered with thick peat (Staub et al., 2000) (Fig. 1). Water samples were collected from a boat. As the boat moved forward, surface water was collected upstream and to the side of the boat into an acid-cleaned polyethylene bottle attached to the end of a plastic pole sampler (3–4 m long). Water temperature, salinity, pH and dissolved oxygen (DO) concentrations were measured in situ using a portable multifunction water quality meter (AP-2000, Aquaread Company, UK) at the time of sample collection. Water samples were filtered within 12 h of collection through pre-cleaned 0.4 μm filters (Nuclepore) at a laminar air flow clean bench (Class 100). The filtrates were placed in acid-cleaned polyethylene bottles and were frozen and stored until analysis.

2.2 Analytical methods

The Se(IV), dissolved inorganic Se (DISe) and total dissolved Se (TDSe) concentrations were determined in carbon-containing plasma using a hydride generation (HG) system (Hydride FAST, ESI) combined with a sector field inductively coupled plasma–mass spectrometry (ICP–MS) instrument, as outlined in the operationally defined hydride generation-based speciation analysis methods described by Chang et al. (2014, 2017). Selenium was measured at $m/z^{-1} = 82$ with low resolution. By adding methane (2 mL min^{-1}) to the carbon-containing plasma, Se sensitivity was increased and spectral interference was suppressed, which improved the detection limits. Briefly, Se(IV) at an acidity of 2 mol L^{-1} HCl was reacted with NaBH_4 to produce hydrogen selenide and then quantified using HG–ICP–MS. Se(VI) was quantitatively reduced to Se(IV) by heating a sample acidified with 3 mol L^{-1} HCl to 97°C for 75 min and then quickly cooling to room temperature using an ice-water bath. The steps used to determine Se(IV) were then followed to obtain the concentration of DISe. The reduction recoveries ranged from 95 % to 103 %. The Se(VI) concentration was calculated as the difference between DISe and Se(IV). The total dissolved selenium (TDSe) concentrations were determined using the same method as that used for DISe, following ultraviolet digestion (Li et al., 2014). The concentration of DOSe was calculated as the difference between the TDSe and DISe concentrations ($\text{DOSe} = \text{TDSe} - \text{DISe}$). Detection limits for Se(IV), DISe and TDSe were 0.0025, 0.0063 and 0.0097 nmol L^{-1} , respectively. The accuracy of the methods was tested with standard solutions, and Se(IV) GSBZ 50031-94, Se(VI) GBW10032, selenocysteine GBW10087 and selenomethionine GBW10034 showed

differences within 3.0 %, 0.7 %, 1.6 % and 1.4 %, respectively.

2.3 The calculation of TDSe flux and yield

TDSe fluxes (Q) transported to the ocean were estimated according to the following equation:

$$Q = C \times V, \quad (1)$$

where C is the mean concentration of TDSe at the fresh end-member of the river (salinity < 1) and V is the river water discharge. Riverine TDSe yield is the ratio of TDSe flux to the drainage area.

2.4 Data statistics and analysis

The Statistical Package for Social Sciences (SPSS) version 23.0 was used to perform a Student's t test, a Mann–Whitney U test and linear regression analyses. The significance level for all the analyses was $p < 0.05$.

3 Results

3.1 Water chemistry

The water chemistry in the freshwater reach of the Maludam, Simunjan, Sebuyau and Samunsam rivers is typical of blackwater rivers draining peatlands with acidic pH and low DO concentrations, and the mixing with coastal water increased the pH and DO (Table S1, Fig. S1 in the Supplement). Values of pH and DO concentrations in the Sematan and Lundu, which drain mostly mineral soils, were higher than those in the blackwater rivers (Fig. S1 in the Supplement). In the Rajang estuary, values of pH and DO were lower in the riverine side, especially in the distributaries covered by the peat (Fig. S2 in the Supplement).

3.2 Se species distributions

TDSe concentrations in the studied rivers and estuaries ranged from 1.0 to 5.7 nmol L^{-1} in March and from 0.70 to 3.9 nmol L^{-1} in September (Fig. 2, Table S1 in the Supplement). DOSe / TDSe ratios ranged from 0.56 to 0.99, indicating that DOSe was the major species of Se in the peat-draining rivers and estuaries in both the dry and wet seasons (Fig. 2, Table S1 in the Supplement). Considerable variation was observed in Se speciation between the studied rivers. The TDSe and DOSe concentrations and DOSe / TDSe ratios were higher in the Maludam, Simunjan, Sebuyau and Samunsam that drain catchments with higher peatland coverage than those in the Rajang, Sematan and Lundu estuaries, whereas the opposite trend is seen for DISe (Table 1, Fig. 2). Student's t test results showed that the concentrations of TDSe, DISe and DOSe did not differ between the wet and dry seasons in the sampled rivers and estuaries ($p > 0.05$),

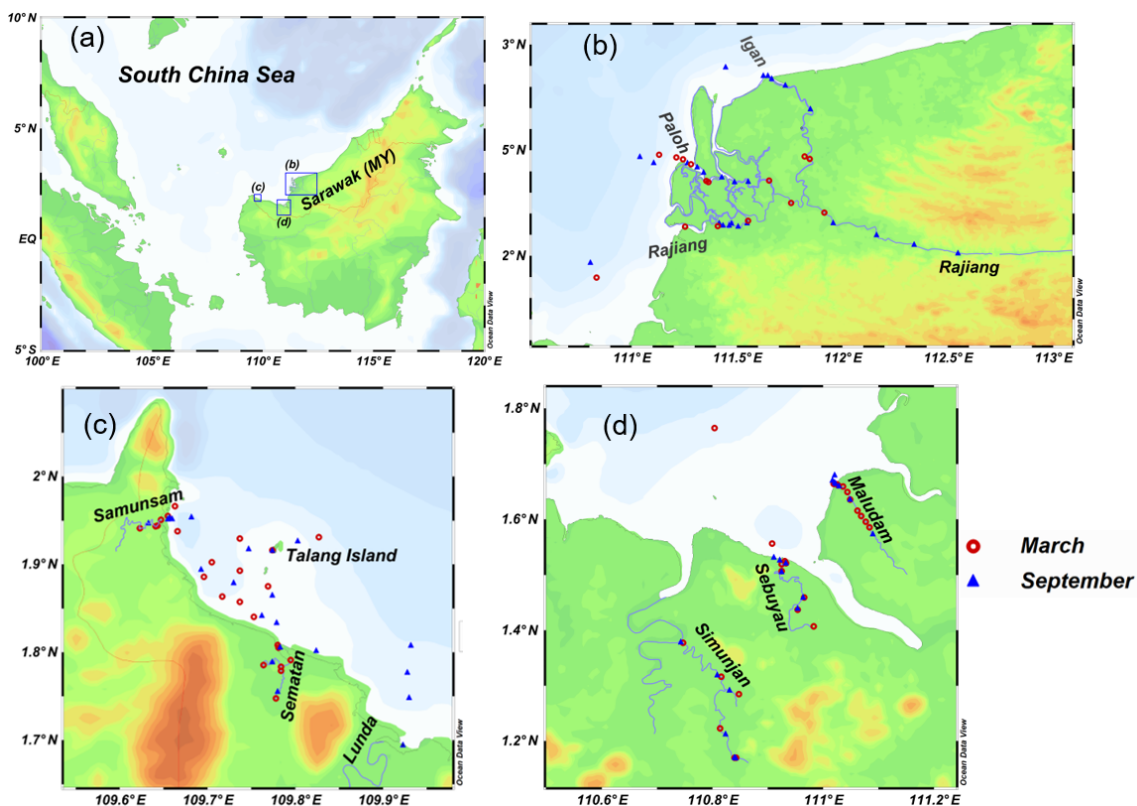


Figure 1. (a) Map of the study area showing the location of Sarawak on the island of Borneo. Blue boxes with letters indicate the areas shown in (b)–(d). (b–d) Station locations for the Rajang River (b); the Samunsam, Sematan, and Lunda rivers (c); and the Maludam, Sebuyau, and Simunjan rivers (d) in March and September 2017. The maps were made with Ocean Data View (2019).

which may reflect the La Niña conditions that caused high precipitation and high discharge rates in Malaysia in 2017 (Jiang et al., 2019).

In the Rajang estuary, Se(IV), Se(VI) and DISe concentrations were high in the coastal areas, whereas DOSe concentrations were higher in the distributaries than in the upper reach in both seasons (Fig. S3 in the Supplement). DOSe/TDSe ratios were high in the distributaries and decreased in a seaward direction to around 0.7 (Fig. S3 in the Supplement). In the Maludam estuary, DISe concentrations were extremely low (near or below the detection limit) in the freshwater reach and increased towards the sea, whereas the DOSe concentrations decreased in a seaward direction in both seasons (Fig. S4 in the Supplement). DISe and DOSe concentrations followed similar patterns in the Sebuyau, Simunjan, Samunsam, Sematan and Lunda estuaries (Fig. S4 in the Supplement). Se(IV) and Se(VI) concentration are not presented but were even lower than those of DISe and were commonly below the detection limit, especially in the freshwater reaches. As TDSe is the sum of the DISe and DOSe concentrations and DOSe generally dominated in the sampled rivers and estuaries, the distributions of TDSe and DOSe were similar (Fig. S4 in the Supplement). The DOSe/TDSe ratios were between 0.8 and almost 1 in

the Maludam, Sebuyau, Simunjan and Samunsam estuaries as salinity < 1, indicating that DOSe was the only (or dominant) species in the freshwater of the blackwater rivers. The DOSe/TDSe ratios were between 0.6 and 0.9 in the Sematan and Lunda, indicating that more than half of the Se was still present in the form of DOSe in those rivers and estuaries with high proportions of mineral soils (Fig. S4 in the Supplement).

3.3 Se species relationship with salinity

Variation in Se species concentrations along a salinity gradient in the three tributaries (the Igan, Lassa and Rajang) of the Rajang Estuary in March and September are shown in Fig. 3. Theoretical mixing lines (TMLs) were developed using two endmembers, namely, a freshwater endmember in the freshwater reach of the Rajang River and a marine endmember with a salinity of > 30. DISe concentrations increased with salinity and, compared with the TML, DISe removals were observed in the Rajang branches in March, and DISe additions were observed in the upper reaches of the Rajang and Paloh branches (Fig. 3a and b). DOSe concentrations decreased with salinity and were much higher than the TML in the Rajang and upper Paloh branches in both March and September and in the Igan Branch in September (Fig. 3c

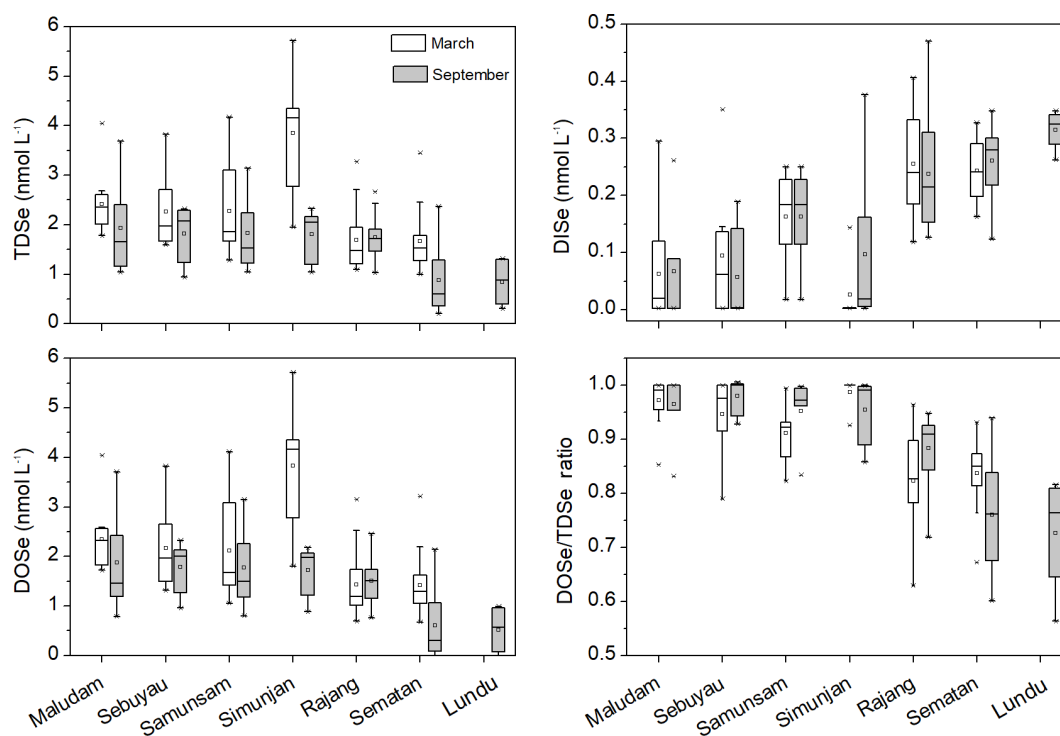


Figure 2. The box plot of TDSe, DISe and DOSe concentration and DOSe / TDSe ratio in the sampled rivers and estuaries in Malaysia in March and September 2017, respectively. The ends of the boxes, the ends of the whiskers, and the line across each box represent the 25th and 75th percentiles, the fifth and 99th percentiles, and the median, respectively; the open square indicates the mean value.

and f). TDSe concentrations in the mixing zone of the Rajang and Igan branches were also higher than the TML (Fig. 3i and j).

Plots of DISe concentration against salinity show a positive linear regression between DISe and salinity in the Maludam, Sebuyau and Samunsam estuaries ($p < 0.05$) in both seasons but not in the Sematan estuary ($p > 0.05$), where DISe concentrations in the freshwater and marine water endmembers were similar in both seasons (Figs. 3, S5 in the Supplement). The salinities varied little, either between the two seasons in the Simunjan and Lundu estuaries or in the Sebuyau estuary in September, and therefore Se concentration–salinity relationships were not examined. As shown in Fig. 3j and h, DOSe concentrations in the freshwater parts of the Maludam and Sebuyau rivers varied widely and increased downstream, so the geographical location nearest to the river mouth with a salinity of < 1 was selected as the freshwater endmember in the linear mixing models. A negative linear correlation was observed between DOSe concentration and salinity ($p < 0.05$) in the Maludam, Sebuyau and Samunsam estuaries for both seasons, but DOSe concentrations did not vary significantly with salinity ($p > 0.05$) in the Sematan estuary (Figs. 3 and S5 in the Supplement). TDSe concentrations were also negatively correlated with salinity ($p < 0.05$) in the Maludam, Sebuyau and Samunsam estuaries but not in the Sematan Estuary (Figs. 3 and S5 in the Supplement).

Generally, relationships between the Se species and salinity fell into three groups. In the blackwater estuaries (the Maludam, Sebuyau and Samunsam), DISe concentrations were positively correlated with salinity; DOSe and TDSe concentrations were negatively correlated with salinity (Fig. 3). In the Rajang estuary, DISe increased with salinity but behaved non-conservatively and was removed in the brackish water region; whereas DOSe and TDSe decreased with salinity, behaved non-conservatively and were added during estuarine mixing (Fig. 3). In the Sematan estuary, TDSe, DOSe and DISe behaved non-conservatively and showed little change during estuarine mixing (Fig. 3).

3.4 Correlation between Se species with DO, pH and dissolved organic matter (DOM)

For the freshwaters ($S < 1$) of the studied rivers, DISe concentrations were positively correlated with the DO concentrations and pH values and the DISe / DOSe ratio was negatively related to DOC concentration (data from Martin et al., 2018; Fig. 4a and b). DOSe concentrations correlated positively with the humification index (HIX) and the sum of the humic-like chromophoric dissolved organic matter (CDOM components, C1, C2, C3 and C4) ($p < 0.05$) (data from Zhou et al., 2019) (Fig. 4c and d).

In the Maludam Estuary, DOSe concentrations were negatively correlated with the CDOM spectral slope from 275

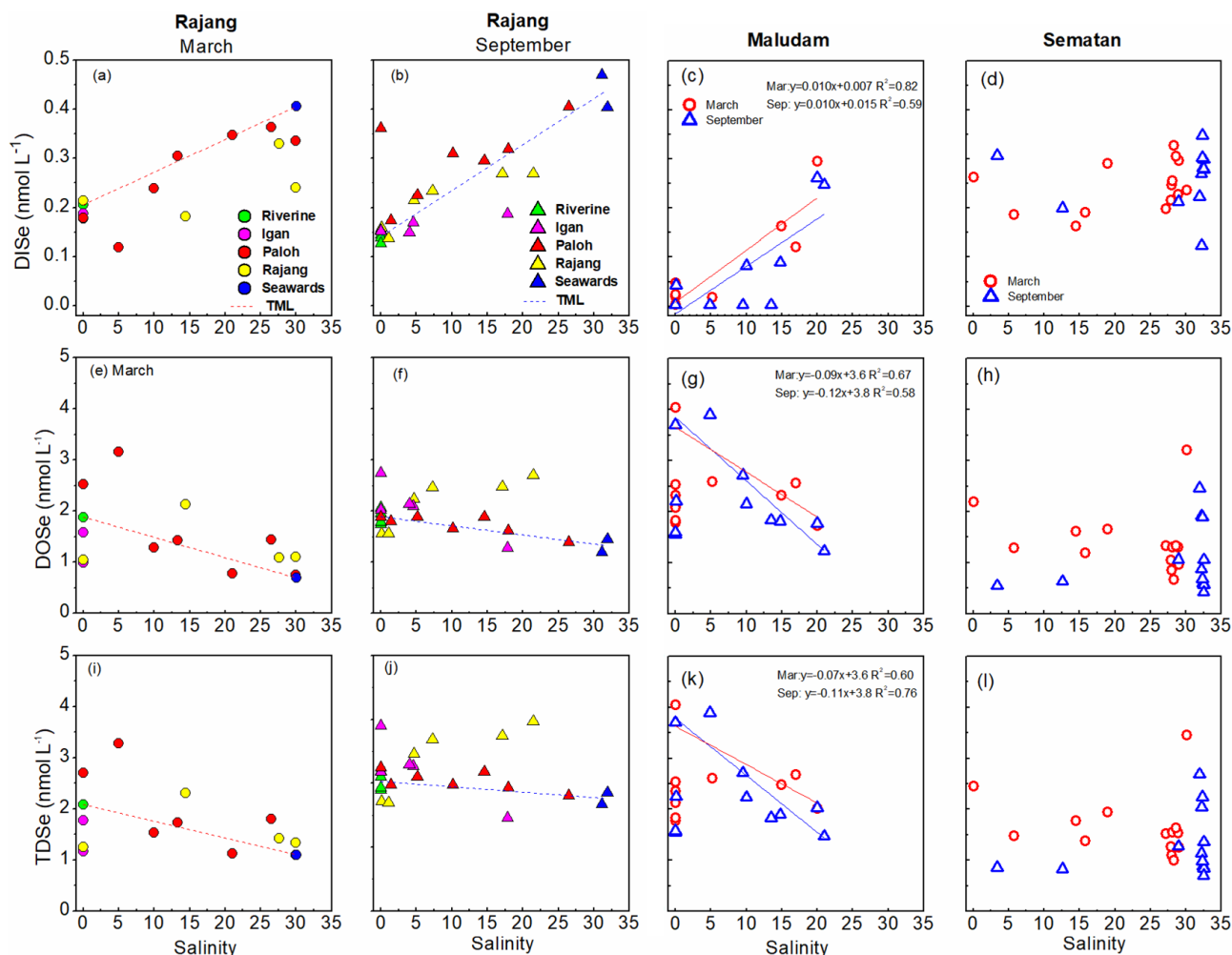


Figure 3. Relationships between DISe (a–d), DOSe (e–h) and TDSe (i–l) concentrations with salinity in the Rajang and three Rajang tributaries (Igan, Lassa and Rajang) and in the Maludam and Sematan estuaries in March and September 2017. TML refers to the theoretical mixing line, which was defined using two endmembers: freshwater in the riverine system and seawater.

to 295 nm ($S_{275-295}$) and were positively correlated with the humic-like C3 component and specific UV absorbance at 254 nm ($SUVA_{254}$) during estuarine mixing in both seasons (data from Martin et al., 2018; Zhou et al., 2019, Fig. 5a–c). In addition, DOSe / DOC and DOSe / DISe ratios were negatively correlated with C2 / C1 component ratios (Fig. 5d, e).

4 Discussion

4.1 Fate of Se species during estuarine mixing

On a global perspective, TDSe concentrations in the sampled rivers were comparable with those in other reported rivers (between 0.2 and 30 nmol L⁻¹); however, in contrast to our findings, DISe generally dominates in other rivers (Table 2, Cutter, 1989b; Conde and Sanz Alaejos, 1997; Pilarczyk et al., 2019). DOSe concentrations in rivers worldwide

range from < 0.02 to 0.82 nmol L⁻¹ (Takayanagi and Wong, 1984; Huang and Shy, 1995; Cutter and Cutter, 2001, 2004). In the blackwater systems of the Orinoco in South America, TDSe concentrations were found to range from 0.07 to 0.25 nmol L⁻¹ (Yee et al., 1987). Although they did not analyse DOSe fractions directly, Yee et al. (1987) assumed that DOSe was likely to constitute about 10%–15% of the total Se, a much lower value than the DOSe proportions observed in peat-draining rivers in Sarawak.

Species of Se are very sensitive to redox conditions and pH values (Sharma et al., 2015). Se(IV) and the Se(VI) are soluble in water which exists under mild and strong oxidizing conditions (Torres et al., 2010), thus DISe concentrations are expected to increase with DO values (Fig. 4a). Sorption to solid surfaces is a pH-dependent process, with substantial sorption of Se(IV) and Se(VI) occurring at pH values of 4 to 6 and negligible sorption under more alkaline conditions (pH > 8) (Bar-Yosef and Meek, 1987; Balistrieri and Chao,

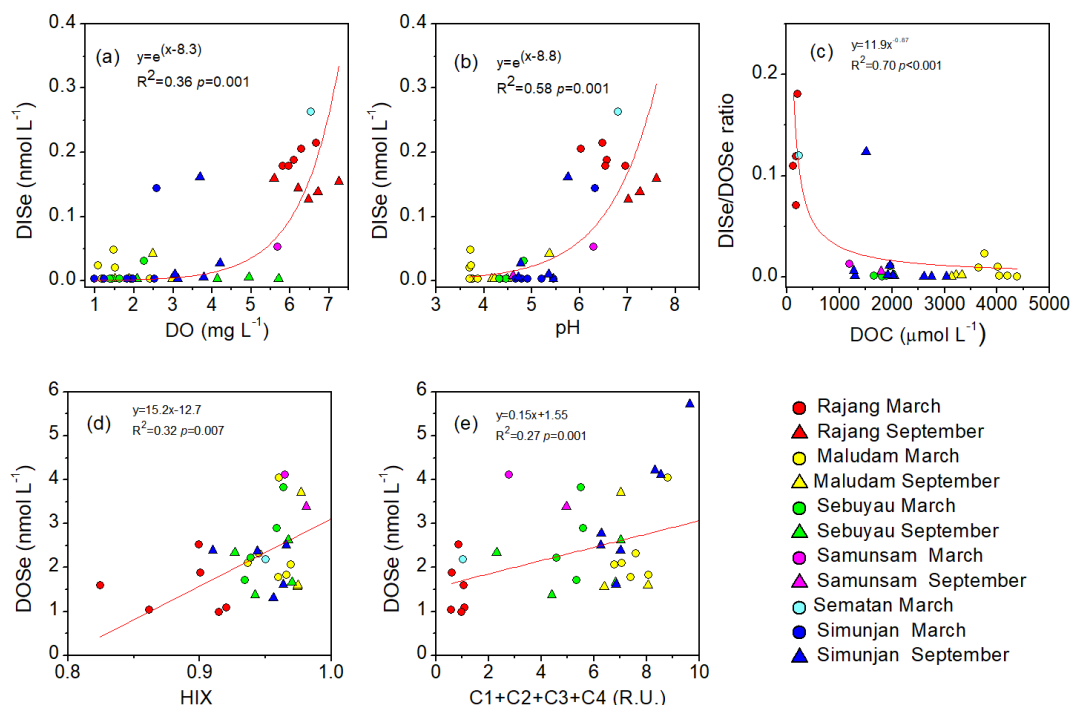


Figure 4. Relationships between (a, b) DISe concentrations and DO and pH values, (c) DISe / DOSe ratios and DOC concentration values, and (d–e) DOSe concentrations with the humification index (HIX) and the sum of humic-like CDOM components (C1, C2, C3 and C4) in freshwater (salinity < 1) for the Rajang, Sematan, Maludam, Sebuyau, Samunsam and Simunjan rivers in March and September. The HIX and C1, C2, C3, and C4 components are from Zhou et al. (2019) from the same cruises. DO concentrations and pH values were not available for the Sematan River for September, and the HIX and CDOM components were not available for the Rajang River for September.

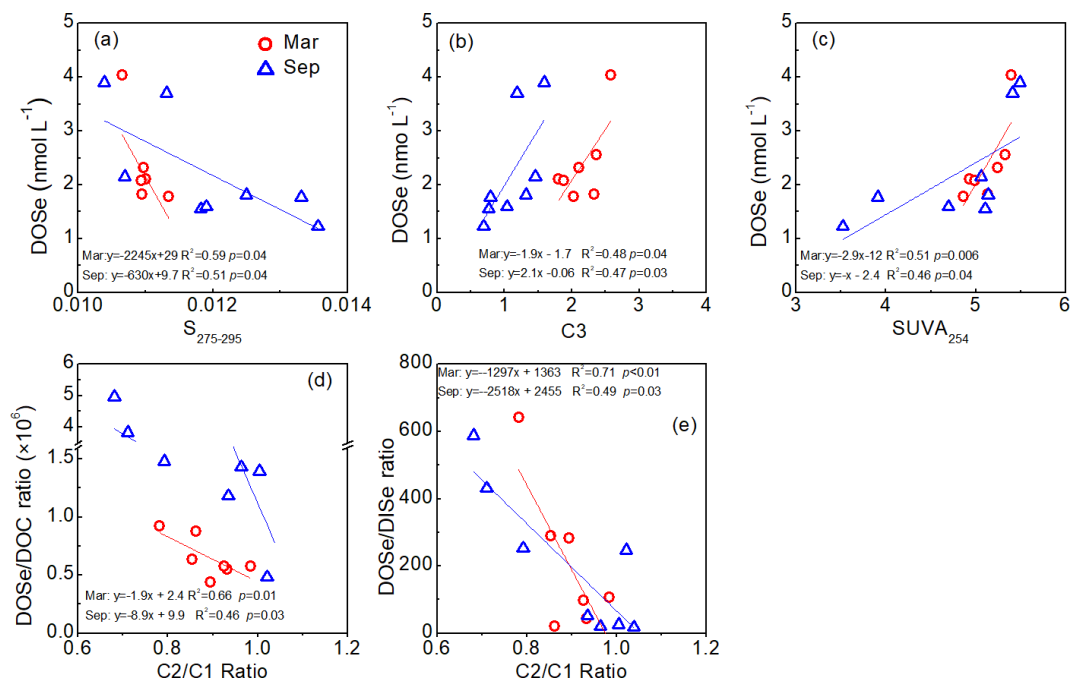


Figure 5. Relationships between DOSe concentrations and $S_{275-295}$, C3 components and $SUVA_{254}$, DOSe / DOC ratio and C2 / C1 component ratios, and DOSe / DISe ratios and C2 / C1 component ratios in the Rajang and Maludam estuaries. The $S_{275-295}$, $SUVA_{254}$, C1, C2 and C3 components are from Martin et al. (2018) and Zhou et al. (2019) taken during the same cruises.

Table 2. Overview of the TDS_e concentrations and DOSe / TDS_e ratios in the river and the magnitude of riverine TDS_e flux and TDS_e and DOSe yield to the ocean.

River name	TDS _e (nmolL ⁻¹)	DOSe / TDS _e ratio	TDS _e flux ^a (10 ³ kg yr ⁻¹)	TDS _e yield ^a (kg km ⁻² yr ⁻¹)	DOSe yield ^a (kg km ⁻² yr ⁻¹)	Reference
Rajang	1.76	0.90	16	0.32	0.28	This study
Maludam	4.04	0.99	0.044	0.22	0.22	This study
Amazon	0.48	0.85	250	0.041	0.035	Cutter and Cutter (2001)
Changjiang	4.59 ^b	NA ^c	652	0.72	NA ^c	Chang et al. (2016)
Zhujiang	4.87 ^b	NA ^c	100	0.20	NA ^c	Yao et al. (2006)
Orinoco	0.45	NA ^c	39	0.036	NA ^c	Yee et al. (1987)
St. Lawrence	2.12	0.11	57	0.047	0.0051	Takayanagi and Wong (1985)
Rhone	2.18	0.14	9.3	0.10	0.013	Guan and Martin (1991)
James	2.08	0.40	1.4	0.020	0.008	Takayanagi and Wong (1983, 1984)
Sacramento	0.91	0.38	1.2	0.023	0.009	Cutter and Cutter (2004)
San Joaquin	15.8	0.23	5.0	0.060	0.014	Cutter and Cutter (2004)
Jiulongjiang	2.44	0.21	1.6	0.11	0.022	Hu et al. (1995)
Kaoping	1.19	0.47	0.26	0.081	0.038	Hung and Shy (1995)
Erhjen	1.11	0.47	0.044	0.13	0.059	Hung and Shy (1995)
Shinano	0.50	< 0.1	0.55	0.046	0.006	Suzuki et al. (1981)
Scheldt	29.2 ^b	NA ^c	13.83	0.63	NA ^c	Van der Sloot et al. (1985)

^a The calculation used river basin areas and discharge rate were cited from Milliman and Farnsworth, 2013. ^b The data were DISe species. ^c The DOSe were not measured.

1987; Papelis et al., 1995; Sharma et al., 2015). Adsorption of Se(IV) and Se(VI) by solid surfaces when pH is between 4 and 6 may help to explain the low DISe concentrations in the sampled freshwater, and DISe concentrations are expected to increase as pH increases (Fig. 4b). In addition almost 15 % of Se(IV) is removed by adsorption to peat (Kharkar et al., 1968). Se(IV) and Se(VI) associated with humic and fulvic substances appear to be responsible for the immobilization of inorganic Se (Kang et al., 1991; Zhang and Moore, 1996; Wang and Gao, 2001). The DISe / DOSe ratios negatively related with DOC concentrations (Fig. 4c). DO, pH and DOC concentrations of the water probably contributed to the observed variations in Se species, and the acidic, low-oxygen and organic-rich blackwater rivers were not a suitable environment for DISe.

During estuarine mixing, reversed DISe concentration–salinity relationships were observed in the Rajang, Maludam, Sebuyau and Samunsam estuaries (Figs. 3, S5 in the Supplement), which contrasts with those reported for other estuaries (Measures and Burton, 1978; Takayanagi and Wong, 1984; Van der Sloot et al., 1985; Cutter, 1989a; Guan and Martin, 1991; Hung and Shy, 1995; Abdel-Moati, 1998; Yao et al., 2006; Chang et al., 2016). The marine endmember of the DISe concentrations in the sampled estuaries (salinity > 31) was 0.30 nmolL⁻¹ (range: 0.12 to 0.47 nmolL⁻¹), encompassing or close to the values reported for surface water in the South China Sea (around 0.38 nmolL⁻¹, Nakaguchi et al., 2004) and the Pacific (mean of 0.24 nmolL⁻¹, range: 0.02 to 0.69 nmolL⁻¹) (Cutter and Bruland, 1984; Sherrard et al., 2004; Mason et al., 2018). The salinity-related increases in DISe in a seaward direction indicate that the patterns

of distribution of DISe in those peat-draining estuaries are controlled mainly by conservative mixing of ocean-derived DISe. In addition, DISe was removed in March but was added in September in the Rajang estuary. Laboratory studies have shown that Se(IV) can be adsorbed by peat and that 60 % of the adsorbed Se(IV) can be desorbed upon exposure of the solid phase to seawater (Kharkar et al., 1968). DISe may have been added to the Rajang estuary in September via release of Se(IV) from peat in brackish waters. Other studies have reported removal of the humic fractions of DOM, colloidal iron and phosphorus by flocculation in the river–sea mixing zones (Eckert and Sholkovitz, 1976; Forsgren et al., 1996; Asmala et al., 2014). Some of the DISe may exist in colloidal form in natural water (Takayanagi and Wong, 1984), and DISe may be removed by flocculation. In peat-draining estuaries, ocean-derived DISe may be adsorbed to peat and may be associated with DOM, which is then converted to DOSe and/or flocculated to particulate Se.

In contrast to DISe, DOSe concentrations were high in the rivers and decreased in a seaward direction as salinity increased (Figs. 3, S5 in the Supplement). DOSe has been shown to behave non-conservatively in other estuaries, with concentrations decreasing along salinity gradients or with mid-estuarine input (Cutter, 1989a; Guan and Martin, 1991; Hung and Shy, 1995; Abdel-Moati, 1998). DOSe concentrations in the estuaries studied in Sarawak were higher than those reported in other estuaries (0.1 to 2.5 nmolL⁻¹) (Cutter, 1989a; Guan and Martin, 1991; Hung and Shy, 1995; Abdel-Moati, 1998). The marine endmember of the DOSe concentrations in the sampled estuaries (salinity > 31) ranged from 0.42 to 2.91 nmolL⁻¹ (mean: 1.32 nmolL⁻¹)

and exceeded those in surface water of the South China Sea (mean: 0.20 nmol L^{-1} , range: 0.33 to 0.14 nmol L^{-1} , Nakaguchi et al., 2004) and the Pacific (mean: 0.36 nmol L^{-1} , range: 0.01 to 0.67 nmol L^{-1} (Cutter and Bruland, 1984; Sherrard et al., 2004; Mason et al., 2018). The high DOSe concentrations in coastal waters in Sarawak ($S > 30$) suggest a significant contribution from terrigenous DOSe. In the distributary channels of Rajang, there are large inputs of organic matter from peat, thus higher DOSe concentrations than the TML values be expected in most of the brackish waters (Fig. 3).

4.2 Character of the DOSe fractions

Coal deposits in Kanawha County in the USA have been interpreted as a dome-shaped peat swamp, analogous to those in Malaysia. Coal Se contents reached 10.7 mg kg^{-1} , and sequential extraction results showed that the concentrations of the organically bound fraction were the highest (Vesper et al., 2008). It is therefore expected that organic matter that is solubilized and leached from peat would cause Se concentrations to increase, and therefore leaching from Se-rich peat soils is inferred to be the major source of DOSe in our sampled rivers. Moreover the peat-draining rivers demonstrated a linear relationship between DOSe concentrations and HIX and humic-like CDOM components (Fig. 4d and e) indicating that DOSe may be associated with dissolved humic substances. In addition, DOSe correlated with $S_{275-295}$ and $SUVA_{254}$ (Fig. 5a and c) suggesting that DOSe was associated closely with high molecular weight and highly aromatic DOM. In addition, the positive correlations between DOSe and the humic-like C3 component (Fig. 5b), i.e. aromatic and black carbon compounds, suggest a strong association of DOSe to these high molecular weight DOM (Fig. 6). Pokrovsky et al. (2018) also found that Se was transported in the form of high molecular weights organic aromatic-rich complexes from peat to the rivers and lakes in the Arctic. Bruggeman et al. (2007) and Kamei-Ishikawa et al. (2008) both found that 50 % to 70 % of Se(IV)–humic substance associates had high molecular weights ($> 10 \text{ kDa}$), consistent with our findings.

During the estuarine mixing, the negative correlation between DOSe/DOC and DOSe/DISE ratios with C2/C1 ratios is enhanced by photodegradation (Wang et al., 2019; Fig. 5d and e), indicating that, compared to bulk DOM, the DOSe fractions were more susceptible to photodegradation, and that DOSe was probably photodegraded to DISE. As suggested by Martin et al. (2018), most photochemical transformations of DOM in Sarawak likely take place after DOM reaches the sea. Thus, photodegradation plays an important role in DOSe processing once transported offshore, and DOSe might contain a significant photoreactive fraction that facilitates photodegradation of DOSe into lower mean molecular weights or gaseous Se or photomineralization to DISE (Fig. 6). Considerable amounts of Se may be

volatilized when methylselenide compounds form (Lidman et al., 2011). A field study found that volatile species of Se were naturally emitted from peatland at concentrations of around 33 nmol L^{-1} (Vriens et al., 2015). As a result of the method used in the present study, volatile methylselenide compounds in the DOSe fractions may not have been detected, so DOSe may have been underestimated. In future work, particular attention should be given to methylselenide. Studies have shown that photodegradation of DOM results in a range of bioavailable products (Miller and Moran, 1997). Peatland-derived DOSe might be degraded to a lower molecular weight or DISE in the coastal areas, both of which are bioavailable for phytoplankton and may stimulate their growth and thereby impact the marine animals via the food chain. The photoreactive DOSe fractions are probably transported across the marginal sea and circulated globally. Given that the bioavailability and biogeochemical cycling of the peatland-derived DOSe fractions may differ from those of peptides produced in situ by phytoplankton in the ocean, the impact on coastal and open ocean ecosystems should be evaluated in the future.

4.3 TDSe flux

Information about the biogeochemistry of peat-draining rivers is scarce, thus their possible quantitative significance for the oceanic TDSe budget is unexplored as of yet. The TDSe flux was estimated to be 16×10^3 and $0.044 \times 10^3 \text{ kg yr}^{-1}$ for Rajang and Maludam, respectively (Table 2). On a global scale, the TDSe delivered from Rajang were less than those large rivers, including the Changjiang, Amazon, Zhujiang, Orinoco and St. Lawrence rivers, but exceeded other small rivers reported so far (Table 2). The TDSe delivered by Rajang and Maludam contributed nearly 1 % of the total riverine TDSe input to the ocean with only 0.3 % of freshwater discharge (Nriagu, 1989; Milliman and Farnsworth, 2013). The TDSe yields for Rajang and Maludam were just below the second-largest river, the Changjiang, and the polluted Scheldt River but exceed the other rivers (Table 2). The magnitude of DOSe yields obtained from Rajang and Maludam was 1 to 2 degrees of magnitude higher than those in other reported rivers (Table 2). This indicates that the numerous small blackwater rivers draining from peatland are very efficient TDSe and DOSe sources for the coastal waters. The roughly estimated TDSe flux from tropical peatland ($439\,238 \text{ km}^2$, Page et al., 2011) could be around $120 \times 10^3 \text{ kg yr}^{-1}$, which was nearly 35 % of the current total riverine TDSe flux, based on average TDSe yield from Rajang and the Maludam ($0.27 \text{ kg km}^{-2} \text{ yr}^{-1}$). From a global perspective, TDSe export from peat-draining rivers is quantitatively more significant than previously thought. It can be expected that increasing anthropogenic disturbing of peat can release a great amount of Se to rivers and then transport it to the coastal areas. The impact of peatland-derived Se to the ecosystem should receive more attention in future studies.

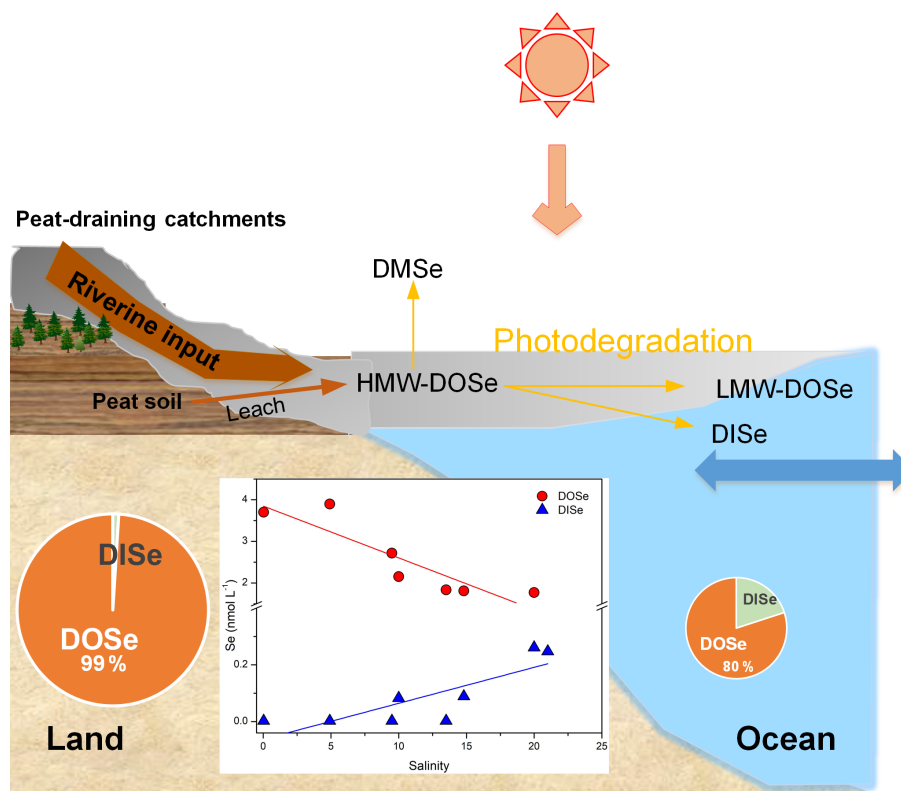


Figure 6. Conceptual diagram of the behaviour of Se species in the Maludam estuary. HMW, LMW and DMSe represent high molecular weight, low molecular weight and dimethylselenide, respectively.

5 Conclusions

To the best of our knowledge, this is the first study of seasonal variations in Se speciation in peat-draining rivers and estuaries in Southeast Asia. Contrary to the results from studies elsewhere, DOSe, not DISe, was the major species in the peat-draining rivers and estuaries of Sarawak, Malaysia. Contrary to our expectations, reversed DISe concentration–salinity relationships were observed in those estuaries, indicating a marine origin, while DOSe concentrations decreased with salinity, indicating terrestrial sources. The DOSe fractions may be associated with high molecular weight peatland-derived aromatic and black carbon compounds. The DOSe yields in the peatland-draining rivers were 1 or even 2 orders of magnitude higher than other reported rivers. The TDOSe flux delivered by the peat-draining rivers exceeded other small rivers, and it is quantitatively more significant than previously thought. The impact of the sizable Se from increasing anthropogenic disturbing of peat to the ecosystem should be evaluated in the future.

Data availability. The datasets used and/or analysed in this study are available from the corresponding author on reasonable request.

Supplement. The supplement related to this article is available online at: <https://doi.org/10.5194/bg-17-1133-2020-supplement>.

Author contributions. JZ, MM, YW, SJ and YC conceptualized the research project and planned the field expeditions. SJ, AM, ESAS, FHAJ and MM performed sample collection and in-situ measurement for the cruises. YC, WWC, JGQ, JLR, EMR and XLW completed laboratory analyses. YC, XLW, YW, SJ, XNW, JZ and MM processed and analysed the data. All co-authors participated in the interpretation and discussion of the results. YC prepared the manuscript with suggestions from all co-authors.

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