

Growth rate rather than temperature affects the B/Ca ratio in the calcareous red alga *Lithothamnion corallioides*

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Abstract. The B/Ca ratio in calcareous marine species is informative of past seawater CO_3^{2-} concentrations, but scarce data exist on B/Ca in coralline algae. Recent studies suggest influences of temperature and growth rates on B/Ca, the effect of which could be critical for the reconstructions of surface ocean pH and atmospheric pCO_2 . In this paper, we present the first laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) analyses of Mg, Sr, Li, and B in the coralline alga Lithothamnion corallioides collected from different geographic settings and depths across the Mediterranean Sea and in the Atlantic Ocean. We produced the first data on putative temperature proxies (Mg/Ca, Li/Ca, Sr/Ca, Mg/Li) and B/Ca in a coralline algal species grown in different basins from across the photic zone (12, 40, 45, and 66 m depth). We tested the B/Ca correlation with temperature proxies and growth rates in order to evaluate their possible effect on B incorporation. Our results suggested a growth rate influence on B/Ca, which was evident in the sample with the lowest growth rate of 0.10 mm yr^{-1} (Pontian Isl., Italy; 66 m depth) and in Elba (Italy; 45 m depth), where the algal growth rate was the highest (0.14 mm yr⁻¹). At these two sites, the measured B/Ca was the lowest at $462.8 \pm 49.2 \,\mu\text{mol}\,\text{mol}^{-1}$ and the highest at $757.7 \pm 75.5 \,\mu\text{mol mol}^{-1}$, respectively. A positive correlation between B/Ca and temperature proxies was found only in the shallowest sample from Morlaix (Atlantic coast of France; 12 m depth), where the amplitude of temperature variation (ΔT) was the highest (8.9 °C). Still, fluctuations in B/Ca did not mirror yearly seasonal temperature oscilla-

tions as for Mg/Ca, Li/Ca, and Sr/Ca. We concluded that growth rates, triggered by the different ΔT and light availability across depth, affect the B incorporation in *L. corallioides*.

1 Introduction

Warming and acidification are major anthropogenic perturbations of present-day oceans (Callendar, 1938; Fairhall, 1973; Brewer, 1997; Gattuso, 1999; Caldeira, 2005; Hönisch et al., 2012; Masson-Delmotte et al., 2021). Ocean acidification reduces the saturation state of calcite and aragonite, lowering the dissolution threshold of biominerals and threatening habitat-forming species of critical ecological importance such as coralline red algae and corals (Morse et al., 2006; Hoegh-Guldberg et al., 2007; Andersson et al., 2008, 2011; Basso, 2012; Ragazzola et al., 2012; Ries et al., 2016). Coralline algae, which precipitate high-Mg calcite (> $8 \mod \% - 12 \mod \% MgCO_3$) (Morse et al., 2006), are particularly suitable as proxy archives for paleoclimate reconstruction because of their worldwide distribution and longevity. Importantly, they show indeterminate growth with no ontogenetic trend (Halfar et al., 2008), which means the growth trend of coralline algae does not slow down asymptotically with age, as in bivalves, thus preserving the resolution of the geochemical signals in all stages of growth (Adey, 1965; Frantz et al., 2005; Halfar et al., 2008). Moreover, coralline algae thin sections under optical microscopy reveal bands that reflect the growth pattern (Cabioch, 1966; Basso, 1995a, b; Foster, 2001), similar to tree rings (Ragazzola et al., 2016) that can be targeted for high-resolution geochemical analyses. Seasonal growth bands, indeed, consist of the perithallial alternation of dark and light bands that together constitute the annual growth patterns (Freiwald and Henrich, 1994; Basso, 1995a, b; Kamenos et al., 2009). Dark bands correspond to slow-growing cells produced in the cold season, which are shorter, thick-walled, and with lower Mg contents, while light bands are fast-growing cells produced in the warm season, which are longer, less calcified, and with higher Mg concentrations (Kamenos et al., 2009; Ragazzola et al., 2016). The high-Mg calcite of calcareous red algae records ambient seawater temperature (Halfar et al., 2000; Kamenos et al., 2008; Nash et al., 2016; Hetzinger er al., 2018), primary productivity (Chan et al., 2017; Hou et al., 2019), and salinity (Kamenos, 2012), proving to be a suitable paleoclimate archive. Most of the data were collected from high-latitude (Kamenos et al., 2008; Anagnostou et al., 2019) and tropical species (Caragnano et al., 2014; Darrenougue et al., 2014), whereas less attention has been given to coralline algae from mid-latitudes.

Trace element variations in marine calcareous species inform the reconstruction of changes in the environmental parameters which characterized the seawater during their growth (Hetzinger et al., 2011; Montagna and Douville, 2017). Boron is incorporated into the mineral lattice of calcareous marine species during calcite precipitation. In the ocean, B occurs in two molecular species: boric acid B(OH)₃ and borate ion $B(OH)_4^-$ (Dickson, 1990), which are related by the following acid–base equilibrium reaction:

$$B(OH)_3 + H_2O \leftrightarrow B(OH)_4^- + H^+, \tag{R1}$$

which shows the dependence of the two species concentrations on pH. The first analyses of the isotopic signal of marine carbonates evidenced a strong similarity with the isotopic composition of $B(OH)_4^-$ in solution, suggesting that borate would preferentially be incorporated into marine carbonates (Vengosh et al., 1991; Hemming and Hanson, 1992; Zeebe and Wolf-Gladrow, 2001; DeCarlo et al., 2018). The B content and its isotopic signature (δ^{11} B) in calcareous marine species record information about the seawater carbonate system. The δ^{11} B is used to reconstruct past seawater pH (Hönisch and Hemming, 2005; Foster, 2008; Douville et al., 2010; Paris et al., 2010; Rae et al., 2011). The boronto-calcium ratio (B/Ca) proved to be informative about past seawater CO_3^{2-} concentrations in different empirical studies on benthic foraminifera (Yu and Elderfield, 2007; Yu et al., 2007; Rae et al., 2011) and in synthetic aragonite (Holcomb et al., 2016). Most of the literature on boron studies is focused on its isotopic composition (Hemming and Hönisch, 2007; Klochko et al., 2009; Henehan et al., 2013; Fietzke et al., 2015; Cornwall et al., 2017; Ragazzola et al., 2020), whereas less attention has been given to B/Ca records, especially in coralline algae. Recent studies suggest that B/Ca is a function of seawater pH, as well as of other environmental variables such as temperature, the effect of which should be considered in the attempt to reconstruct surface ocean pH and atmospheric pCO_2 (Wara et al., 2003; Allen et al., 2012; Kaczmarek et al., 2016).

To achieve the best reliability of geochemical proxies for climate reconstructions, it is important to recognize the influence of multiple factors on a single proxy (Kaczmarek et al., 2016; Donald et al., 2017). For instance, more recently the effects of temperature and growth rate on B incorporation have been investigated through experiments on both synthetic and biogenic carbonates (Wara et al., 2003; Yu et al., 2007; Gabitov et al., 2014; Mavromatis et al., 2015; Uchikawa et al., 2015; Kaczmarek et al., 2016; Donald et al., 2017). In particular, a culture experiment on the coralline alga Neogoniolithon sp. showed a positive correlation of B/Ca with growth rate and a negative correlation with Sr/Ca, which was proposed as a proxy for dissolved inorganic carbon (DIC) (Donald et al., 2017). Moreover, a culture experiment on the high-latitude species Clathromorphum compactum (Kjellman) Foslie 1898 revealed non-significant temperature influences on B/Ca and a significant inverse relationship with growth rate (Anagnostou et al., 2019). The factors which influence the B incorporation in calcareous red algae are therefore still debated. Recent experiments also suggest that coralline algae can control the calcifying fluid pH (pH_{cf}) (Cornwall et al., 2017), as already observed in corals (Comeau et al., 2017). Both organisms have a speciesspecific capability to elevate pH at calcification sites in response to variations of ambient pH, also influencing precipitation rates (Cornwall et al., 2017). Differences between carbonate polymorphs were also highlighted (McCulloch et al., 2012; Cornwall et al., 2018), showing more elevated pH_{cf} in aragonitic corals than calcites, pointing to the relevance of the mineralogical control on biological up-regulation. So far, no investigations on pH_{cf} modifications in natural systems have been performed on calcareous red algae.

No studies have been conducted so far on the correlation between temperature proxies (Mg, Sr, Li/Ca) and B/Ca. The Mg/Ca ratio is extensively used as a temperature proxy in coralline algae (Halfar et al., 2008; Kamenos et al., 2008; Fietzke et al., 2015; Ragazzola et al., 2020), since the substitution of Mg^{2+} with Ca^{2+} ions in the calcite lattice is an endothermic reaction. Accordingly, Mg incorporation increases with temperature (Moberly, 1968; Berner, 1975; Ries, 2006; Caragnano et al., 2014, 2017). Sr/Ca and Li/Ca ratios in calcareous red algae have also been investigated as climate proxies, showing significant positive correlations with temperature in different species (Kamenos et al., 2008; Hetzinger et al., 2011; Caragnano et al., 2014; Darrenougue et al., 2014). The Mg/Li ratio showed a strong correlation with seawater temperature in cultured C. compactum (Anagnostou et al., 2019) and in empirical studies on high-Mg calcites, including coralline algae (Stewart et al., 2020). Conversely, the Mg/Li calibration did not reveal improvements

in the Mg/Ca or Li/Ca proxies in *Lithophyllum* spp. (Caragnano et al., 2014, 2017).

Here, we present laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) conducted on a wildgrown coralline alga with a wide geographic scope. This technique, which allows high-resolution analysis of a broad range of trace elements in solid-state samples, has been widely used in biogenic carbonates to extract records of seawater temperature, salinity, and water chemistry (Schöne et al., 2005; Corrège, 2006; Hetzinger et al., 2009, 2011; Fietzke et al., 2015; Ragazzola et al., 2020). Measurements were made on the non-geniculate coralline alga Lithothamnion corallioides (P. Crouan and H. Crouan) P. Crouan and H. Crouan 1867, which is widely distributed in the Mediterranean Sea and in the north-eastern Atlantic Ocean from Scotland to the Canary Islands (Irvine and Chamberlain, 1999; Wilson et al., 2004; Carro et al., 2014), usually constituting maerl beds (Potin et al., 1990; Foster, 2001; Martin et al., 2006; Savini et al., 2012; Basso et al., 2017). It forms rhodoliths as unattached branches (Basso et al., 2016) with obvious banding in longitudinal sections (Basso, 1995b). These characteristics combine to make this species a suitable model for the measurement of geochemical proxies by comparing different environmental settings.

In this paper, we provide the first LA-ICP-MS data on putative temperature proxies (Mg/Ca, Sr/Ca, Li/Ca, Mg/Li) and B/Ca measured on *L. corallioides* collected from different geographic settings and depths across the Mediterranean Sea and in the Atlantic Ocean. We test the influence of temperature and growth rate on the B/Ca ratio, which could be crucial in assessing the reliability of B/Ca as a proxy for the seawater carbonate system.

2 Materials and methods

2.1 Sampling sites and collection of *Lithothamnion corallioides*

Samples of the coralline alga *L. corallioides* were collected in the western Mediterranean Sea and in the Atlantic Ocean (Fig. 1). In the Mediterranean Sea, the samples collected in the Pontian Islands (Italy) at 66 m depth were gathered by grab during the cruises of the R/V *Minerva Uno* in the framework of the Marine Strategy Campaigns 2016 (Table 1). The last two Mediterranean samples were collected by one of the authors (DB) by scuba diving during local surveys at 45 m off the coasts of Pomonte (Elba Island, Italy) (Basso and Brusoni, 2004) and at 40 m depth in the Aegadian Islands (Marettimo, Italy). The Atlantic sample was collected by grab at 12 m depth in Morlaix Bay (Brittany, France) (Table 1).

The identification of the algal samples was achieved by morphological analyses of epithallial and perithallial cells using a field emission gun scanning electron microscope (SEM-FEG) Gemini 500 Zeiss. Samples were prepared for



Figure 1. Map showing the distribution of sampling sites where *Lithothamnion corallioides* samples were collected: Morlaix Bay (48°34′42″ N, 3°49′36″ W), Aegadian Islands (37°97′36″ N, 12°14′12″ E), Elba (42°44′56.4″ N, 10°07′08.4″ E), and Pontian Islands (40°54′ N, 12°45′ E). Service layer credits: source Esri, GEBCO, NOAA, National Geographic, Garmin, https://www.geonames.org/ (last access: 21 December 2021), and other contributors.

SEM according to Basso (1995a). Morphological identification was based on Adey and McKibbin (1970) and Irvine and Chamberlain (1994). Other information about maerl species distribution in Morlaix was provided by Carro et al. (2014) and Melbourne et al. (2017). *L. corallioides* was selected as the target species because of its presence in both Mediterranean and Atlantic waters. The Atlantic sample (Morlaix) was used as voucher specimen for the subsequent identification of the Mediterranean samples, since *Phymatolithon* spp. and *L. corallioides* are the only components of the Morlaix maerl (Carro et al., 2014; Melbourne et al., 2017). Hence, once its inclusion under the genus *Phymatolithon* was excluded, the Morlaix sample identified as *L. corallioides* was used as a reference for the most reliable identification of the other Mediterranean samples.

2.2 Sample preparation

The selected algal branches were embedded in Epo-Fix resin, which was stirred for 2 min with a hardener (13 % w/w); they were then left to dry at room temperature for 24 h. Afterwards, the treated branches were cut by an IsoMet diamond wafering blade 15HC along the direction of growth. In the laboratory of the Institute of Geosciences and Earth Resources of the National Research Council (IGG-CNR) in Pavia (Italy), the sections were polished with a MetaServ grinder–polisher (400 RPM) using a diamond paste solution, finally cleaned ultrasonically in distilled water for 10 min, and dried at 30 °C for 24 h.



Figure 2. (a) Thalli of *Lithothamnion corallioides* collected in Morlaix (scale bar: 5 mm). (b) Longitudinal section through the *L. corallioides* branch sampled in Morlaix showing the LA-ICP-MS transects targeting each growth band (scale bar: 200 µm).

2.3 Trace elements analyses and environmental data

LA-ICP-MS analyses were carried out at the IGG-CNR laboratory on one algal branch per sampling site. 43 Ca, 7 Li, 25 Mg, 88 Sr, and 11 B contents were measured using an Agilent ICP-QQQ 8900 quadrupole ICP-MS coupled to an Excimer laser ablation system (193 nm wavelength, MicroLas with GeoLas optics). Element / Ca ratios were calculated for these isotopes, as was the Mg/Li ratio. Measurements were performed with laser energy densities of 4 J cm⁻² and helium as a carrier gas.

The laser transects crossed the algal growth direction with a spot size of 50 µm, attempting to target each growth band change which marked the transition between the cells usually produced in the warm season and those usually produced in the cold season, hereinafter referred to as long and short cells (Figs. 2, S1, S2, S3). NIST 612 was used as an external standard (e.g. Fietzke et al., 2010; Jochum et al., 2012), whereas Ca was adopted as an internal standard. Accuracy and precision were better than 4 % for NIST 612 and 8 % for Ca standard. Minimum detection limits (99 % confidence) of measured elements were Ca = 16.91, Li = 0.07, Mg = 0.11, Sr = 0.004, and B = 2.64 ppm. Each analysis was carried out in MS/MS mode for 3 min by acquiring 60 s of background before and after the sampling period by the laser on the polished surface. The first part of the signal was not used for the integration to avoid surface contamination. The Glitter software (v. 4.4.4) was used for data reduction.

In the absence of in situ environmental data, the seawater temperature has been extracted by 11 years of monthly reanalysis spanning 1979–2016 from from the Ocean ReAnalysis System 5 (ORAS5) at 0.25° horizontal resolution (Zuo et al., 2019). The time interval of extraction for each site corresponded to 11 years before sample collection (Bracchi et al., 2021). Minimum, maximum, and mean values, as reported in Table 1, refer to the temperature at sampling depth and have been measured over the entire time interval of extraction.

For the purpose of this work, we considered temperature data in terms of (a) the amplitude of temperature variation (ΔT) and (b) the temperature maxima and minima. ΔT represents the temperature fluctuations during the algal growth and has been measured as the difference between the maximum and minimum temperature registered at the site over 11 years. We used ΔT when comparing the sampling sites, given their differences in depth and geographical regions. The temperature peaks (maximum and minimum) have rather been used when considering data corresponding to long and short cells, since they are related to warm and cold periods of growth, respectively. We used the temperature peaks over the entire time interval of extraction (11 years) when comparing the mean elemental ratios of long and short cells per sampling site. The maximum and minimum temperature within each year have been used instead for the reconstruction of the algal age model. In the sample from Morlaix Bay, indeed, the good visibility of the growth bands allowed us to relate the temperature record with the algal growth at annual resolution. We therefore plotted all the element ratios against the average seawater temperature values of the coldest and warmest months of the year to reconstruct the temperature variations during the algal growth (Moberly, 1970; Corrège, 2006; Williams et al., 2014; Ragazzola et al., 2020; Caragnano et al., 2014). Missing element ratios, possibly due to non-targeted consecutive bands, were calculated as the means of known values.

Carbon system parameters for each site have also been estimated. Even if they were not available in the same time interval of temperature data, the seasonal variations occurring in the extracted period allowed the characterization of the sampling sites. Monthly mean seawater pH has been derived by the EU Copernicus Marine Service Information (CMEMS) global biogeochemical hindcast spanning 1999–2017 at 0.25° horizontal resolution (Perruche, 2018). Monthly means of DIC in the time interval 1999–2017 have been extracted by CMEMS biogeochemical reanalysis for the Mediterranean Sea at 0.042° horizontal resolution



Figure 3. Correlation plots of Mg/Ca with Li/Ca and Sr/Ca. For each analysis the Spearman's coefficient r, the p value, and the line equation are given.



Figure 4. Box plot of the statistical tests performed to evaluate the differences of Mg/Ca in *L. corallioides* collected at different sampling sites. The horizontal black lines indicate the median values. The black dots and the numbers inside the plot indicate the mean values.

(Teruzzi et al., 2021). At the Atlantic site, monthly means of DIC in 1999–2017 were derived from the CMEMS IBI biogeochemical model at 0.083° horizontal resolution (Copernicus Marine Environmental Monitoring Service, 2020). The pH and DIC data showed consistent variations among sites, despite being derived from different biogeochemical models. The mean values of DIC and pH, as reported in Table 1, refer to sampling depth and have been measured over the entire time interval of extraction. The complete timeline of temperature and carbon data used in this paper is shown in Supplement Figs. S5, S6, S7, and S8.

2.4 Growth rate estimation

Growth rates were estimated under light microscope by measuring the length of the LA-ICP-MS transect and dividing it by the number of annual growth bands crossed by the transect (Bracchi et al., 2021). The obtained values are expressed in linear extension per year (mm yr⁻¹). In the samples wherein the growth bands were not easily detectable under microscope, i.e. the Elba sample, we also used the Mg/Ca results in order to check for the correspondence of Mg peaks with growth bands.

2.5 Statistical analysis

Statistics were calculated for both the dataset with all the spot analyses and the dataset with the records from long and short cells separately. Short cells refer to slow-growing cells in dark bands, usually produced in the cold season; long cells correspond to fast-growing cells in light bands, usually produced in the warm season (Kamenos et al., 2009; Ragazzola et al., 2016). For each spot, a distinction between the cells was thus made by image analyses, except for the Elba sample, given the poor resolution of the growth bands. The Spearman's correlation was tested to provide the statistical comparisons between Mg/Ca, Li/Ca, Sr/Ca, and B/Ca records from the LA-ICP-MS analyses in L. corallioides. The Kruskal–Wallis test, followed by the Dunn's test for comparisons, and the one-way analysis of variance (ANOVA), followed by the Tukey's test for post hoc analysis, were used to compare the geochemical signals among sampling sites and to evidence the differences between group medians and means. All statistical analyses were run in R 3.6.3 software.



Figure 5. Box plots of the statistical tests performed to evaluate the differences of Mg/Ca in the long and short cells of *L. corallioides* collected at different sampling sites. The horizontal black lines indicate the median values. The black dots and the numbers inside the plot indicate the mean values.



Figure 6. Correlation plot between Mg/Li and seawater temperature. Data are shown for cultured *C. compactum* (Anagnostou et al., 2019) and *L. corallioides* (this paper). *L. corallioides* results are shown separately in long and short cells per sampling site.

3 Results

3.1 Environmental data

The temperature data obtained by ORAS5 reanalysis revealed a lower amplitude of the seasonal temperature change in the Mediterranean samples with respect to the Atlantic site, as shown by the standard deviation and ΔT values in Table 1. This difference is explained in terms of the different sampling depths, with the seasonal variations decreasing with increasing depth.

Temperature variations in Morlaix Bay (Atlantic Ocean) were higher, registering an overall mean seawater temperature of 12.4 °C (Table 1). Among Mediterranean samples, mean seawater temperatures were highest in the Aegadian Isl., followed by Elba and the Pontian Isl. (Table 1). Aegadian Isl. also registered the highest temperature variations among the Mediterranean sites (Table 1). Moderate temperature variations characterized the site in Elba, which registered the lowest monthly mean temperature among Mediterranean sites (Table 1). At the Pontian Isl., consistent with the fact that it is the deepest sampling site at 66 m depth, the lowest seawater temperature variations were found (Table 1).

The pH estimates at the Mediterranean sites were all similarly high at ~ 8.13 and less variable than the Atlantic site (8.05). The mean pH had slightly higher values in Pontian Isl. and Elba than Aegadian Isl. (Table 1). Similarly, DIC was higher in the Mediterranean sites and decreased in Morlaix, as this largely dictates the pH (Table 1).

3.2 Mg/Ca, Li/Ca, Sr/Ca, and Mg/Li

Both Li/Ca and Sr/Ca records had positive correlations with Mg/Ca in our samples of *L. corallioides* (Figs. 3 and S4). The overall mean Mg/Ca was 225.3 ± 30.4 mmol mol⁻¹, registering the minimum value in the sample from Aegadian Isl. (171.7 mmol mol⁻¹) and the maximum value in Morlaix (311.2 mmol mol⁻¹) (Fig. 4; Table 2). The Kruskal–Wallis test did not show significant differences in Mg/Ca among samples (Table A1 in the Appendix; Fig. 4). Among Mediterranean sites, the algal sample from Aegadian Isl. had the highest Mg/Ca mean value, followed by Elba and Pontian Isl., which had the lowest Mg/Ca mean value of all sampling sites (Fig. 4). The highest mean Mg/Ca was registered in the sample from Morlaix Bay, which also showed a large dispersion of data above the median Mg/Ca value (Fig. 4).



Figure 7. Mg/Li ratio of *L. corallioides* collected in Morlaix Bay. Note the lack of cyclic variations in Mg/Li results. In the timeline, the coldest and the warmest months have been reported. Mg/Li ratios in the missing bands (asterisks) have been calculated as the means of the values measured in warm and cold periods. Monthly means of seawater temperature have been extracted by ORAS5 reanalysis.



Figure 8. Correlation plots of B/Ca with Mg/Ca, Li/Ca, and Sr/Ca in *L. corallioides* collected in Morlaix Bay. For each analysis the Spearman's coefficient r, the p value, and the line equation are given.

Long cells had high Mg/Ca values; conversely, short cells corresponded to areas with a low Mg/Ca ratio.

The ANOVA test followed by the Tukey's test for multiple comparisons evidenced a significant variability of algal Mg/Ca among three sites in long cells (Table A2; Fig. 5). In the long cells of *L. corallioides* collected from Aegadian Isl. and Pontian Isl., the Mg/Ca data showed a quite similar distribution (Table A2; Fig. 5). The Mg/Ca of the alga from Pontian Isl. was the lowest (Fig. 5). In Morlaix a higher Mg/Ca mean value was registered, which is significantly different compared to Aegadian and Pontian Isl. (Table A2; Fig. 5). In short cells, the differences in Mg/Ca among samples were not statistically significant (Table A1). The magnesium incorporation was slightly higher in Morlaix and very similar between Aegadian and Pontian Isl. samples (Fig. 5).

Mg/Li values in long and short cells fell in the range found by Anagnostou et al. (2019) for cultured *Clathromorphum compactum* (Fig. 6). When plotted against the extracted seawater temperature in Morlaix (Fig. 7), Mg/Li results did not reflect the seasonal oscillations in temperature.

3.3 B/Ca

The B/Ca ratio in the sample collected from Morlaix showed a moderate positive correlation with all the examined temperature proxies (Mg/Ca, Li/Ca, Sr/Ca), with a more defined trend when plotted against Li/Ca (r = 0.68) and slightly less defined against Mg/Ca (r = 0.58) and Sr/Ca (r = 0.57) (Fig. 8). On the contrary, the Spearman's analyses did not evidence significant correlations between B/Ca and the temperature signals in the algae collected elsewhere (p > 0.05).

Overall, the B/Ca ratio in *L. corallioides* was $661.9 \pm 138.9 \,\mu\text{mol}\,\text{mol}^{-1}$, registering the minimum value in the long cells of the sample from Pontian Isl. $(356.0 \,\mu\text{mol}\,\text{mol}^{-1})$ and the maximum value in Elba $(954.1 \,\mu\text{mol}\,\text{mol}^{-1})$ (Fig. 9; Table 2).

The Kruskal–Wallis coefficient showed a highly significant difference in the B/Ca value among sites, particularly in the *L. corallioides* from the Pontian Isl., which had the lowest boron incorporation (Table A3; Fig. 9). The algae collected in Aegadian Isl. still had significantly lower B/Ca compared to those collected in Elba and Morlaix (Table A3; Fig. 9).

Table 1. Environmental data from each sampling site. The minimum and maximum monthly means of temperature are indicated, as are the highest temperature variation (ΔT), the mean, and the standard deviation of the time series. Data from monthly means were extracted by 11 years of ORAS5 reanalysis. The pH and DIC for each sampling site are also indicated. The minimum, maximum, mean, and standard deviation values have been measured over the time interval 1999–2017. Carbonate system parameters have been extracted from monthly mean biogeochemical data provided by CMEMS.

Sampling site	Depth (m)	Collection date	Temperature (°C)			рН		DIC (µmol kg ⁻¹)		Growth rates $(mm yr^{-1})$		
			min	max	ΔT	mean	SD	mean	SD	mean	SD	
Aegadian Isl.	40	25/08/1993	13.7	18.8	5.1	15.9	1.3	8.11	0.02	2303.72	11.58	0.13
Elba	45	01/12/1990	12.9	17.5	4.6	15.0	1.2	8.13	0.02	2322.24	7.08	0.14
Pontian Isl.	66	27/07/2016	13.9	16.7	2.8	15.0	0.6	8.13	0.01	2321.29	6.71	0.10
Morlaix	12	02/05/1991	8.3	17.2	8.9	12.4	2.5	8.05	0.04	2155.60	15.27	0.13



Figure 9. Box plot of the statistical tests performed to evaluate the differences of B/Ca in *L. corallioides* collected at different sampling sites. The horizontal black lines indicate the median values. The black dots and the numbers inside the plot indicate the mean values.

The highest B/Ca mean value was registered in Elba, with medians comparable to Morlaix (Table A3; Fig. 9).

The ANOVA test followed by the Tukey's test for multiple comparisons by site, for long (Table A4) and short cells (Table A5) separately, showed lower values at the Mediterranean sites and higher values at the Atlantic site (Fig. 10).

The sample from Pontian Isl. had the lowest mean B/Ca in both seasons, being significantly different from the samples from both Morlaix and Aegadian Isl. (Tables A4, A5; Fig. 10). Morlaix had the highest mean B/Ca in both long and short cells (Tables A4, A5; Fig. 10). *L. corallioides* from Aegadian Isl. had an intermediate B/Ca mean value in long cells, differing significantly from both the Morlaix and Pontian Isl. samples (Table A4; Fig. 10). In short cells, the sample from Aegadian Isl. slightly differed from the one in Morlaix (Table A5; Fig. 10). Interestingly, the long cells of all samples had higher median B/Ca values compared to short cells (Fig. 10), although only in Morlaix, the differences between B/Ca measured in long and short cells were statistically significant ($\chi^2 =$ 8.4899, p < 0.01).

3.4 Growth rates

In the sample from Aegadian Isl., the LA-ICP-MS transect was 1.31 mm long, and 10 years of growth was detected by coupling microscopical imaging and Mg/Ca peaks, resulting in a growth rate of 0.13 mm yr^{-1} . In the Elba sample the laser transect was 1.15 mm long, crossing 8 years of growth, with a resulting growth rate of 0.14 mm yr^{-1} . The Pontian Isl. sample had 1.08 mm of transect including 11 years of growth and hence a growth rate of 0.10 mm yr^{-1} . Finally, the transect from the Morlaix sample was 1.38 mm long, counting 11 years and resulting in a growth rate of 0.13 mm yr^{-1} .

Growth rates did not show any linear relationship with Mg, Li, and Sr/Ca, but they were positively correlated with the sample mean B/Ca values (Fig. 11).

4 Discussion

Temperature variations affect many physiological processes involved in biomineralization, and the rate of calcification, along with the preservation state of mineral structures, influences the content of trace elements in carbonates (Lorens, 1981; Rimstidt et al., 1998; Gussone et al., 2005; Noireaux et al., 2015; Kaczmarek et al., 2016). Trace element concentrations recorded from the four *L. corallioides* branches analysed in this study were consistent with previously published values for other calcareous red algae (Chave, 1954; Hemming and Hanson, 1992; Hetzinger et al., 2011; Darrenougue et al., 2014). Particularly, the Mg/Ca ratios recorded in this study ranged from 172 to 311 mmol mol⁻¹, which is comparable to previous studies on rhodoliths of *Lithothamnion glaciale* Kjellman 1883 grown at 6–15 °C (148–326 mmol mol⁻¹) (Kamenos et al., 2008). The B/Ca



Figure 10. Box plots of the statistical tests performed to evaluate the differences of B/Ca in the long and short cells of *L. corallioides* collected at different sampling sites. The horizontal black lines indicate the median values. The black dots and the numbers inside the plot indicate the mean values.



Figure 11. Correlation plots of growth rates and seawater temperature with B/Ca in *L. corallioides* samples analysed in this study. Spearman's coefficient *r*, the *p* value, and the line equation are given. Temperature variations (ΔT) correspond to the differences between the maximum and minimum temperature registered over 11 years of monthly reanalysis (ORAS5). The B/Ca means measured in long and short cells respectively correspond to the maximum and minimum temperature.

ratios in *L. corallioides* from our results range from 356 to $954 \,\mu\text{mol}\,\text{mol}^{-1}$, which is wider than the range measured by solution ICP-MS on bulk samples of *Neogoniolithon* sp. $(352-670 \,\mu\text{mol}\,\text{mol}^{-1})$ (Donald et al., 2017) and *C. compactum* (320–430 $\mu\text{mol}\,\text{mol}^{-1})$ (Anagnostou et al., 2019), both cultured with controlled *p*CO₂ and a pH ranging from 7.2 to 8.2. The high resolution given by laser ablation should be more effective in measuring the heterogeneity of B/Ca across the thallus, thus explaining the wider range of our data.

The Mg/Ca results evidenced a strong relationship with the seawater temperatures extracted from ORAS5 (Table 1; Fig. 12), as expected. *L. corallioides* from Aegadian Isl. had slightly higher Mg/Ca values, followed by Elba and Pontian Isl. (Fig. 4). This was consistent with local temperature values in the Mediterranean (Table 1), since Pontian Isl. registered the lowest mean value and the lowest ΔT , while Aegadian Isl. showed the highest mean temperature and ΔT . On the contrary, the sample from Morlaix, collected at 12 m depth, showed high Mg/Ca values in both long and short cells (Table A2; Fig. 5). The monthly mean temperatures had the highest variations during the year (ΔT in Table 1) due to the shallower depth (12 m) and the geographical location. Temperature correlates with seasons, influencing primary production, respiration, and calcification in *L. corallioides* (Payri, 2000; Martin et al., 2006) as well as other calcareous red algae (Roberts et al., 2002). The high seasonality that characterized the sample from Morlaix, represented by the high ΔT (Table 1), was probably responsible for the highest variation of Mg/Ca values and undoubtedly



Figure 12. Elemental ratios in *L. corallioides* collected in Morlaix Bay (scale bar: $200 \,\mu$ m). Mg, Li, and Sr/Ca show cyclic variations mirroring the local seawater temperature. In the timeline, the coldest and the warmest months have been reported, which correspond to dark and light bands of growth. Element / Ca ratios in the missing bands (asterisks) have been calculated as the means of the values measured in warm or cold periods. Monthly means of temperature have been extracted by ORAS5 reanalysis.

accounted for most of the differences with Mediterranean samples. For the first time, we confirmed the reliability of the temperature proxies Li/Ca and Sr/Ca on a wild-grown coralline alga collected at different depths and locations. Li/Ca and Sr/Ca records were positively correlated with Mg/Ca in *L. corallioides* (Fig. 3), which, in turn, showed a strong relationship with seawater temperature. Moreover, both Li and Sr/Ca showed periodical oscillations in correspondence to long and short cells, related to seasonal temperature variations (Fig. 12). Therefore, Li/Ca and Sr/Ca could be regarded as temperature proxies in *L. corallioides*, as could Mg/Ca. The coupling of the Mg/Ca ratio with Li/Ca and Sr/Ca can be considered a useful technique to gather information about past temperature for paleoclimate reconstructions (Halfar et al., 2011; Caragnano et al., 2014; Williams et al., 2014; Fowell et al., 2016; Cuny-Guirriec et al., 2019).

The B/Ca ratio in coralline algae has rarely been measured, and it is not clear how the environmental factors control its incorporation. The carbonate system primarily drives the changes in B incorporation (Hemming and Hanson, 1992; Yu and Elderfield, 2007). In benthic foraminifera, B/Ca increases with $[CO_3^{2-}]$ (Yu and Elderfield, 2007), whereas there is no consensus on the effect of $[CO_3^{2-}]$ on Mg/Ca

Sampling site	B/Ca (µmol mol ⁻¹)		$ \begin{array}{ c c c c } Mg/Ca & \\ (mmol mol^{-1}) & (\mu r) \end{array} $		Li/ (µmol n	Li/Ca (µmol mol ⁻¹)		Sr/Ca (mmol mol ⁻¹)		Mg/Li (mol mmol ⁻¹)	
	mean	SD	mean	SD	mean	SD	mean	SD	mean	SD	
Aegadian Isl.	610.8	63.9	224.9	30.3	81.9	18.4	3.7	0.4	2.8	0.5	
Elba	757.7	75.5	223.4	26.4	85.2	14.3	3.6	0.4	2.7	0.3	
Pontian Isl.	462.8	49.2	216.1	21.9	79.6	14.6	3.3	0.3	2.8	0.4	
Morlaix	726.9	102.8	239.5	41.2	113.1	32.9	4.1	0.5	2.2	0.4	
Total	661.9	138.9	225.3	30.4	89.0	23.3	3.7	0.5	2.6	0.4	

Table 2. Element / Ca ratio measurements in L. corallioides.

and Sr/Ca (Rosenthal et al., 2006; Dueñas-Bohórquez et al., 2011). Experiments with inorganic calcite showed a positive correlation between B/Ca and [DIC] (Uchikawa et al., 2015). Nevertheless, in culture experiments of the coralline algae *Neogoniolithon* sp. (Donald et al., 2017) and corals (Gagnon et al., 2021), [DIC] had a negative effect on B/Ca. DIC and B/Ca values showed a negative relationship in the samples from Morlaix, Aegadian Isl., and Pontian Isl., which was not observed in Elba (Fig. 9; Table 1). Significant differences among B/Ca values in the Mediterranean samples were not expected, since DIC concentrations were similar (Table 1). This evidence suggests influences other than DIC on the B signal.

The mean estimated growth rate of L. corallioides was 0.13 ± 0.02 mm yr⁻¹, and it was supposed to decrease with increasing depth as a direct consequence of lower light availability (Halfar et al., 2011); indeed, the growth rate of the sample from Pontian Isl. was the lowest $(0.10 \text{ mm yr}^{-1})$. As already suggested by previous studies on both synthetic and biogenic calcite, B incorporation is likely affected by growth rate (Gabitov et al., 2014; Mavromatis et al., 2015; Noireaux et al., 2015; Uchikawa et al., 2015; Kaczmarek et al., 2016). Indeed, in the cultured calcareous red alga Neogoniolithon sp. B/Ca increases with increasing growth rate (Donald et al., 2017). The slowest growth rate found in Pontian Isl. possibly contributed to the lowest B/Ca value; similarly, the highest growth rate $(0.14 \text{ mm yr}^{-1})$ in the sample from Elba could be responsible for the highest B/Ca (Figs. 9, 11). The mean annual growth rate of the shallowest sample (Morlaix) was intermediate $(0.13 \text{ mm yr}^{-1})$ and likely not constant during the year. In Morlaix, the alga probably significantly slowed down the growth in cold months, when the monthly mean seawater temperature was the lowest of all sampling sites (Table 1). Nevertheless, its growth rate likely sped up in the warm season due to the abundant light availability at shallow depth and the warming of seawater (Table 1), contributing to the significantly higher B/Ca values in long cells (Fig. 10). According to this interpretation, the effect of the growth rate on B/Ca might be significant across depth and geographical regions (Fig. 11).

In Morlaix, B/Ca showed a weak positive correlation with temperature proxies (Mg/Ca, Li/Ca, and Sr/Ca; Fig. 8). A positive correlation between B/Ca and Mg/Ca was already observed in planktonic foraminifera (Wara et al., 2003; Yu et al., 2007). Therefore, we reconstructed the elemental variations during algal growth in the Morlaix sample at annual resolution in order to highlight the influence of temperature (Fig. 12). Contrary to Mg, Li, and Sr/Ca, the B/Ca did not mirror the seasonal temperature variations as accurately as the other proxies.

In the sample from the Pontian Isl., the seasonal ΔT , Mg/Ca, and B/Ca values were the lowest among sites. In particular, B/Ca was significantly low (Fig. 9), differing more from the other samples than the results for Mg/Ca (Fig. 4). This suggests that in this sample the B incorporation should be influenced by factors other than those affecting Mg. In general, the poor correlation with seawater temperature (Fig. 11), and most of all the lack of distinct seasonal oscillations in B/Ca across the algal thallus (Fig. 12), excludes the suitability of B/Ca as a temperature proxy and suggests a closer relationship with growth rate than temperature.

Knowing the biogeochemistry and the variation of the environmental variables of seawater is crucial for a more comprehensive picture of the reliability of geochemical proxies like the ones investigated here (Mg, Li, Sr/Ca, and B/Ca). Boron incorporation in marine carbonates is still debated, raising questions about the boron isotopic fractionation, the seawater isotopic composition, and the so-called "vital effects" (i.e. the metabolic activities that can bias the proxy record). Vital effects include species-specific biologically controlled up-regulations of pH in the calcifying fluid in both corals and coralline algae in response to pH manipulations mimicking the ongoing ocean acidification (Cornwall et al., 2017, 2018). Natural marine pH undergoes pH fluctuations with characteristic ranges of about 0.3 pH units in very shallow coastal water, in areas with restricted circulation, or in shallow nearshore communities supporting high rates of benthic metabolism, such as seagrass beds (Cornwall et al., 2013; Duarte et al., 2013; Hofmann et al., 2011). However, this study analysed the trace element record in a single high-Mg-calcite species grown in a natural wide bathymetric interval (from 10 to 66 m depth) characterized by normal marine pH (range 8.05–8.13, Table 1). Therefore, we considered upregulation, if present, to be constant among our conspecific specimens and thus irrelevant for the measured B/Ca variations. Moreover, no significant yearly pH fluctuations are expected at each site. Thus, within a single specimen, the observed differences in B/Ca between short and long cells (i.e. cold and warm periods of growth) (Fig. 10) are unlikely to be related to changes in pH_{cf}.

The paucity of B/Ca measurements from coralline algae and, most of all, the complete absence of these data on specimens grown in nature make it difficult to compare our B/Ca data with the literature. This observation takes stock of the significance of our results and emphasizes the importance of collecting more representative B/Ca data in coralline algae. Further studies on *L. corallioides* and other calcareous red algae should be carried out to clarify the environmental factors influencing the B/Ca in these organisms and to ensure the reliability of this proxy for paleoclimate reconstructions.

5 Conclusions

This paper presents the first measures of trace elements (Mg, Sr, Li, and B) from the coralline alga L. corallioides collected across the Mediterranean Sea and in the Atlantic Ocean at different oceanographic settings and depths (12, 40, 45, and 66 m depth), providing the first geochemical data on a wildgrown coralline algal species sampled at different depths and geographical locations. LA-ICP-MS records of Mg/Ca, Sr/Ca, and Li/Ca have shown a similar trend, primarily controlled by seawater temperatures in the algal habitat. On the contrary, Mg/Li did not provide a valuable temperature proxy in this species. In order to evaluate the control exerted by temperature over B incorporation, we also tested the correlation of B/Ca with Mg/Ca, Li/Ca. and Sr/Ca. This led us to provide the first B/Ca data on wild-grown coralline algae from across the photic zone in different basins. The correlation between B/Ca and Mg/Ca in L. corallioides was statistically significant only in the shallow Atlantic waters of Morlaix, where seasonality, and hence the seasonal temperature variations, during the algal growth was the strongest among sites. Accordingly, B incorporation differences between long and short cells of L. corallioides strongly depend on seasonality, being statistically significant just in Morlaix. Nevertheless, in contrast to Mg, Li, and Sr/Ca, B/Ca oscillations across the algal growth showed a poor relationship with seasonal variations in seawater temperature. We found high B/Ca values in the Atlantic sample, wherein pH and DIC were the lowest. Carbon data did not explain the low B concentration in the Pontian Isl. sample (66 m depth), though, wherein pH and DIC were similar to the other Mediterranean sites. The estimation of growth rate, which was low in the Pontian Isl. sample $(0.10 \text{ mm yr}^{-1})$ and got higher in the other Mediterranean samples and in Morlaix

 $(\sim 0.13 \text{ mm yr}^{-1})$, led us to conclude that B/Ca relates to growth rate rather than seawater temperature. B incorporation is therefore subject to the specific algal growth patterns and rates, knowledge of which is essential in order to assess the reliability of B/Ca in tracing seawater carbon variations.

Appendix A

Table A1. (a) Statistically non-significant results of tests performed to evaluate (a) the differences of Mg/Ca in *L. corallioides* and (b) the differences of Mg/Ca in the short cells of *L. corallioides* collected at different sampling sites. Test significance at $\alpha = 0.05$.

	(a) Kruskal–Wallis test (Mg/Ca)								
		Df	χ ²	Р					
	Site	3	3.799	0.284					
(b) One-way ANOVA test (Mg/Ca)									
-	Short cells								
	Df	Sum sq.	Mean sq.	F value	$\Pr\left(>F\right)$				
Site	2	788.1	394.0	1.4647	0.2496				
Residuals	26	6994.5	269.0						
Shapiro–Wilk normality test Bartlett's <i>K</i> squared				P = 0.6442 P = 0.5856					

Table A2. Results of statistical tests performed to evaluate the differences of Mg/Ca in the long cells of *L. corallioides* collected at different sampling sites. Statistically significant *p* values are given in bold. ANOVA test significance at $\alpha = 0.05$; Tukey's test significant at $p \le \alpha$.

One-way ANOVA test (Mg/Ca)								
	Long cells							
	Df	Sum sq.	Mean sq.	F value	$\Pr\left(>F\right)$			
Site	2	10897.7	5448.9	16.413	0.0001			
Residuals	20	6639.8	332.0					
Shapiro–Wilk normality test $P = 0.1440$								
Bartlett's K squared $P = 0.5826$								
Tukey's test								
		Multiple com	parisons of mean	ns				
Site	Mean difference	95 % confid	dence interval <i>P</i> . adjust		ljusted			
	Site	lower bound	upper bound					
Morlaix–Aegadian Isl.	38.32918	15.09816	61.56019	0.0	0130			
Pontian Isl.–Aegadian Isl.	-10.84361	-35.48382	13.79661	0.5	1716			
Pontian IslMorlaix	-49.17278	-72.40380	-25.94177	0.0	0009			

Table A3. Results of statistical tests performed to evaluate the differences of B/Ca in *L. corallioides* collected at different sampling sites. Statistically significant *p* values are given in bold. Kruskal–Wallis test significance at $\alpha = 0.05$; Dunn's test significant at $p \le \alpha/2$.

Kruskal–Wallis test (B/Ca)									
	Df	χ ²	Р						
Site	3	79.816	$< 2.2 \times 10^{-16}$						
Dunn's test Comparisons	Dunn's test Comparisons by site (Bonferroni)								
Z P. adjusted	Aegadian Isl.	Elba	Morlaix						
Elba	-4.64580 0.00000								
Morlaix	-3.07755 0.00630	1.17249 0.72300							
Pontian Isl.	2.80564 0.01510	8.38673 0.00000	6.15663 0.00000						

Table A4. Results of statistical tests performed to evaluate the differences of B/Ca in the long cells of *L. corallioides* collected at different sampling sites. Statistically significant *p* values are given in bold. ANOVA test significance at $\alpha = 0.05$; Tukey's test significant at $p \le \alpha$.

One-way ANOVA test (B/Ca)								
	Long cells							
	Df	Sum sq.	Mean sq.	F value	$\Pr\left(>F\right)$			
Site	2	428 364	214 182	33.066	0.0000			
Residuals	20	129 546	6477					
Shapiro–Wilk normality test $P = 0.5527$								
Bartlett's K squared $P = 0.5470$								
Tukey's test								
		Multiple com	parisons of mea	18				
Site	Mean difference	95% confide	ence interval	P. adjusted				
	Site	lower bound	upper bound					
Morlaix–Aegadian Isl.	190.11730	87.50374	292.73094	0.0	0.00040			
Pontian Isl.–Aegadian Isl.	-135.42490	-244.26303	-26.58672 0.0134		1342			
Pontian IslMorlaix	-325.54220	-428.15581	-222.92862	0.0	0000			

Table A5. Results of statistical tests performed to evaluate the differences of B/Ca in the short cells of *L. corallioides* collected at different sampling sites. Statistically significant *p* values are given in bold. ANOVA test significance at $\alpha = 0.05$; Tukey's test significant at $p \le \alpha$.

One-way ANOVA test (B/Ca)								
	Short cells							
	Df	Sum sq.	Mean sq.	F value	$\Pr\left(>F\right)$			
Site Residuals	2 26	216 232 79497	108 116 3058	35.360	0.0000			
Shapiro–Wilk normality test $P = 0.1699$								
Bartlett's K squared $P = 0.0576$								
Tukey's test								
		Multiple comp	arisons of mear	18				
Site	Mean difference	95 % confider	nce interval	<i>P</i> . ad	ljusted			
	Site	lower bound	upper bound					
Morlaix–Aegadian Isl. Pontian Isl.–Aegadian Isl. Pontian Isl.–Morlaix	43.09640 -156.90170 -199.99810	-19.61932 -223.66771 -260.58727	105.81212 -90.13574 -139.40898	0.22 0.00 0.00	2146 0001 0000			

Data availability. Data resulting from this study are available at https://doi.org/10.1594/PANGAEA.932201 (last access: 28 January 2022, Piazza et al., 2021). Environmental data were provided by EU Copernicus Marine Service Information. DIC data in the Mediterranean: https://doi.org/10.25423/CMCC/MEDSEA_ MULTIYEAR_BGC_006_008_MEDBFM3 (last access: 28 January 2022, Teruzzi et al., 2021). DIC data in the Atlantic https://resources.marine.copernicus.eu/?option=com_ Ocean: cswview=detailsproduct_id=IBI_MULTIYEAR_BGC_005_003 (last access: 28 January 2022, Copernicus Marine Environmental Monitoring Service, 2020.) The pH data: https://resources.marine.copernicus.eu/?option=com_cswview= detailsproduct_id=GLOBAL_REANALYSIS_BIO_001_029 (last access: 13 October 2021, Perruche, 2018) Temperature https://icdc.cen.uni-hamburg.de/daten/reanalysis-ocean/ data: easy-init-ocean/ecmwf-oras5.html (last access: 28 January 2022, Zuo et al., 2019).

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