



Supplement of

Molecular-level characterization of supraglacial dissolved and water-extractable organic matter along a hydrological flow path in a Greenland Ice Sheet micro-catchment

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S1. Supplementary methods

S1.1 Hydrology

Hydraulic conductivity and water table height were calculated from recharge curves in five auger holes with an initial radius of 70 mm and depth of 400 – 450 mm. Throughout the duration of the experiment, surface lowering caused an effective
5 reduction in auger hole depth. Recharge rate was recorded using ultrasonic rangefinders at a 1Hz resolution, and recharge curves were processed as described in Stevens et al. (2018). The height of the water table was defined following the completion of recharge, i.e. when water in the auger hole stopped rising. Full recharge holes were emptied using a siphon prior to the subsequent recharge measurement.

10 Orthoimages and DEMs of the site were created using aerial imagery from a UAV (DJI Mini 2) at 9:00, 14:00, and 21:00 on the sampling day. Flight elevation was 20 m, whilst the camera was set with a fixed aperture ($f/2.8$) and ISO (100), with shutter speed and UAV airspeed adjusted to suit the prevailing light conditions. One orthophoto and DEM were constructed per image-set, at a resolution of 0.024 m/pixel using the commercially available Agisoft MetashapePro (Agisoft, Russia) and GNSS-derived position from the UAV (i.e. this was not corrected using ground control points).

15 Hydrological modelling of the weathering crust was undertaken to establish water flow direction and magnitude, which was used to drive a particle tracking tool to establish hydrological connectivity between auger hole D and the supraglacial stream. The model used was the ‘Darcy Velocity’ groundwater flow direction and magnitude model included in the Spatial Analyst package of ArcMap 10.8 (esri, USA). This tool requires the following inputs: hydraulic head, effective porosity, saturated thickness and transmissivity (itself a function of the saturated thickness and hydraulic conductivity). As the weathering crust
20 is an unconfined aquifer (e.g. (Müller and Keeler, 1969)), the hydraulic head is equivalent to the absolute elevation of the water table (Cohen and Cherry, 2020). The water table (and therefore hydraulic head) was interpolated across the study area using the ‘Topo-to-Raster’ tool, tied with manually digitized streams and point measurements of absolute water table from each borehole (i.e., DEM elevation minus water depth from the surface). Effective porosity was parameterized as 0.35, assuming a saturated ice density of 550 kg m^{-3} (after (Cooper et al., 2018)). Saturated thickness was calculated using the
25 interpolated water table, and a parameterized weathering crust depth of 50 cm from the surface (Irvine-Fynn et al., 2021; Stevens et al., 2018), and is multiplied by median hydraulic conductivity for each time window, derived from point measurements, to establish transmissivity. Note that there is no hydraulic conductivity measurement at site A at 14:00 and 21:00, and these values are parameterized using available data from this hole and consideration of temporal hydraulic conductivity trends in holes B-E. The modelled flow vector field was used to drive the ‘Particle Track’ tool for auger hole D,
30 producing modelled flow path of water from this hole.

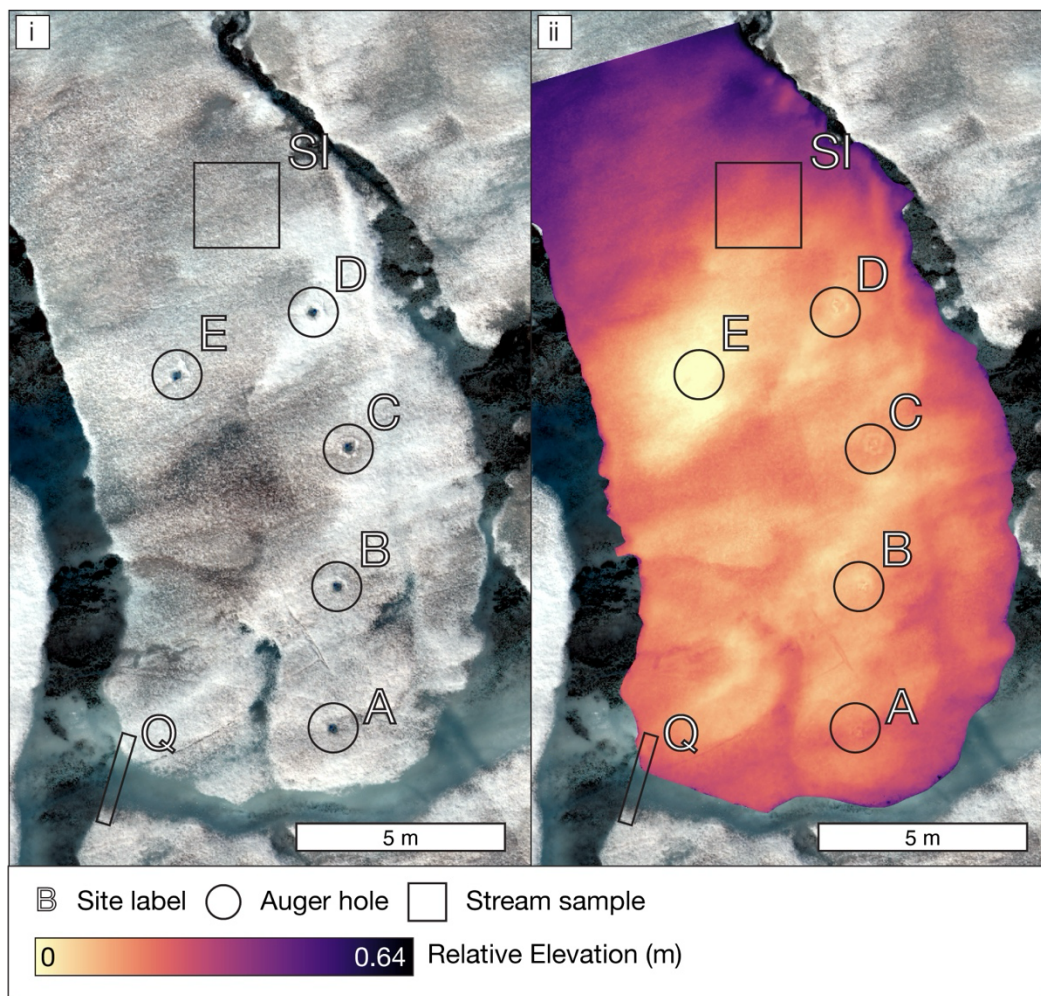


Figure S1 | A DSM (panel ii) of the study site, with sampling locations marked. Note that elevations are relative to a reference datum, not mean sea level. The orthophoto provided in Fig 1 is included for reference (panel i).

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S1.2 FT-ICR MS instrumentation and acquisition settings

Sample solution was infused via a micro-electrospray source (Emmett et al., 1998) (50 μm i.d. fused silica emitter) at 500 nL/min by a syringe pump. Typical conditions for negative ion formation were: emitter voltage, -2.4-2.9 kV; tube lens, -250 V; and heated metal capillary current, 7 A. DOM extracts were analysed with a custom-built hybrid linear ion trap FT-ICR mass spectrometer equipped with a 21 T superconducting solenoid magnet (Hendrickson et al., 2015; Smith et al., 2018). Ions were initially accumulated in an external multipole ion guide (1-5 ms) and released m/z -dependently by decrease of an auxiliary radio frequency potential between the multipole rods and the end-cap electrode (Kaiser et al., 2013). Ions were excited to m/z -dependent radius to maximize the dynamic range and number of observed mass spectral peaks (32-64%), and excitation and

detection were performed on the same pair of electrodes (Chen et al., 2014). The dynamically harmonized ICR cell in the 21
 45 T FT-ICR is operated with 6 V trapping potential (Boldin and Nikolaev, 2011; Kaiser et al., 2013). Time-domain transients of
 3.1 seconds were acquired with the Predator data station that handled excitation and detection only, initiated by a TTL trigger
 from the commercial Thermo data station, with 100 time-domain acquisitions averaged for all experiments (Blakney et al.,
 2011). Mass spectra were phase-corrected (Xian et al., 2010) and internally calibrated with 10-15 highly abundant homologous
 series that span the entire molecular weight distribution based on the “walking” calibration method (Savory et al., 2011).
 50 Experimentally measured masses were converted from the International Union of pure and Applied Chemistry (IUPAC) mass
 scale to the Kendrick mass scale (Kendrick, 1963) for rapid identification of homologous series for each heteroatom class (i.e.
 species with the same $C_cH_hN_nO_oS_s$ content, differing only by the degree of alkylation) (Hughey et al., 2001).

55 **Table S1 The number of formulae assigned per FT-ICR MS file, prior to blank correction, and the RMS error per assigned file. All
 calibrated peak lists and assigned files, and a list of all elemental compositions assigned across the dataset, are publicly available via
 the Open Science Framework (<https://osf.io/>) via DOI 10.17605/OSF.IO/JRBTH.**

File name in OSF	Sample group	Number assigned formulae	RMS error
X2021December02_weathering_crust_R1_sum100_assigned.csv.RelAbun	Weathering crust	11,537	0.054
X2021December02_weathering_crust_R2_sum100_assigned.csv.RelAbun	Weathering crust	13,761	0.054
X2021December02_weathering_crust_R3_sum100_assigned.csv.RelAbun	Weathering crust	12,756	0.055
X2021December02_weathering_crust_R4_sum100_assigned.csv.RelAbun	Weathering crust	11,930	0.046
X2021December02_field_blank_sum100_assigned.csv.RelAbun	Procedural blank	4,365	0.100
X2021December02_supraglacial_stream_R1_sum100_assigned.csv.RelAbun	Supraglacial stream	13,347	0.050
X2021December02_supraglacial_stream_R2_sum100_assigned.csv.RelAbun	Supraglacial stream	12,307	0.068
X2021December02_supraglacial_stream_R3_sum100_assigned.csv.RelAbun	Supraglacial stream	12,634	0.050
X2021December02_supraglacial_stream_R4_sum100_assigned.csv.RelAbun	Supraglacial stream	9,774	0.054
X2021December02_supraglacial_stream_R5_sum100_assigned.csv.RelAbun	Supraglacial stream	10,096	0.056
X2021December02_dark_ice_R1_sum100_assigned.csv.RelAbun	Surface ice	11,155	0.050
X2021December02_dark_ice_R2_sum100_assigned.csv.RelAbun	Surface ice	8,684	0.067
X2021December02_dark_ice_R3_sum100_assigned.csv.RelAbun	Surface ice	9,763	0.052
X2021December02_dark_ice_R4_sum100_assigned.csv.RelAbun	Surface ice	11,710	0.056
X2022January26_laboratory_leachate_R1_sum100_assigned.csv.RelAbun	SIP-WEOM	6,617	0.078
X2022January26_laboratory_leachate_R2_sum100_assigned.csv.RelAbun	SIP-WEOM	8,340	0.099
X2022January26_laboratory_leachate_R3_sum100_assigned.csv.RelAbun	SIP-WEOM	6,514	0.056
X2022January26_laboratory_leachate_R4_sum100_assigned.csv.RelAbun	SIP-WEOM	7,543	0.064

Table S2 Principal Component Analysis Structure Matrix of DOM parameters

Parameter	PC1	PC2
Formulae (#)	-0.462	0.059
Mass ^{wa} (Da)	0.553	0.796
NOSC ^{wa}	1.026	0.102
AI _{mod} ^{wa}	1.021	-0.081
Aliphatic High O/C (%RA)	-0.624	0.754
Aliphatic Low O/C (%RA)	-0.903	-0.499
HUP High O/C (%RA)	1.008	0.012
HUP Low O/C (%RA)	-0.525	0.863
Peptide-like (%RA)	-0.896	0.170
Condensed aromatic (%RA)	0.993	-0.144
Polyphenolic (%RA)	0.990	-0.174
CHO (%RA)	0.797	0.640
CHON (%RA)	-0.686	0.755
CHOS (%RA)	-0.609	-0.821

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