

Implications of CO₂ pooling on $\delta^{13}\text{C}$ of ecosystem respiration and leaves in Amazonian forest

A. C. de Araújo¹, J. P. H. B. Ometto^{2,5}, A. J. Dolman¹, B. Kruijt³, M. J. Waterloo¹, and J. R. Ehleringer⁴

¹Department of Hydrology and Geo-Environmental Sciences, Faculty of Earth and Life Sciences, Vrije Universiteit, De Boelelaan 1085, 1081 HV, Amsterdam, The Netherlands

²Laboratório de Ecologia Isotópica, Centro de Energia Nuclear na Agricultura, USP, Av. Centenário, 303, Piracicaba, SP, CEP 13416-190, Brazil

³Earth System Science-Climate Change, Centre for Water and Climate, Wageningen-UR, PO Box 47, 6700 AA, Wageningen, The Netherlands

⁴Department of Biology, University of Utah, 257S 1400E, Salt Lake City, UT, 84112-0840, USA

⁵IGBP Regional Office/INPE, Av. dos Astronautas, 1.758 - Jd. Granja, São José dos Campos, SP, CEP 12227-010, Brazil

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Abstract. The carbon isotope of a leaf ($\delta^{13}\text{C}_{\text{leaf}}$) is generally more negative in riparian zones than in areas with low soil moisture content or rainfall input. In Central Amazonia, the small-scale topography is composed of plateaus and valleys, with plateaus generally having a lower soil moisture status than the valley edges in the dry season. Yet in the dry season, the nocturnal accumulation of CO₂ is higher in the valleys than on the plateaus. Samples of sunlit leaves and atmospheric air were collected along a topographical gradient in the dry season to test whether the $\delta^{13}\text{C}_{\text{leaf}}$ of sunlit leaves and the carbon isotope ratio of ecosystem respired CO₂ ($\delta^{13}\text{C}_{\text{Reco}}$) may be more negative in the valley than those on the plateau.

The $\delta^{13}\text{C}_{\text{leaf}}$ was significantly more negative in the valley than on the plateau. Factors considered to be driving the observed variability in $\delta^{13}\text{C}_{\text{leaf}}$ were: leaf nitrogen concentration, leaf mass per unit area (LMA), soil moisture availability, more negative carbon isotope ratio of atmospheric CO₂ ($\delta^{13}\text{C}_a$) in the valleys during daytime hours, and leaf discrimination (Δ_{leaf}). The observed pattern of $\delta^{13}\text{C}_{\text{leaf}}$ might suggest that water-use efficiency (WUE) is higher on the plateaus than in the valleys. However, there was no full supporting evidence for this because it remains unclear how much of the difference in $\delta^{13}\text{C}_{\text{leaf}}$ was driven by physiology or $\delta^{13}\text{C}_a$. The $\delta^{13}\text{C}_{\text{Reco}}$ was more negative in the valleys than on the plateaus on some nights, whereas in others it was not.

It is likely that lateral drainage of CO₂ enriched in ¹³C from upslope areas might have happened when the nights were less stable. Biotic factors such as soil CO₂ efflux (R_{soil}) and the responses of plants to environmental variables such as vapor pressure deficit (D) may also play a role. The preferential pooling of CO₂ in the low-lying areas of this landscape may confound the interpretation of $\delta^{13}\text{C}_{\text{leaf}}$ and $\delta^{13}\text{C}_{\text{Reco}}$.

1 Introduction

The use of isotopic tracers in organic matter, water, and atmospheric gases has emerged as a powerful tool that integrates biotic and physical processes over space and time, improving our understanding of plant physiology, biogeochemistry, and ecosystem function (Pataki et al., 2003b; Pataki et al., 2007). The mean carbon isotope ratio of atmospheric CO₂ ($\delta^{13}\text{C}_a$) is currently –8‰ (Keeling et al., 2005a; Keeling et al., 2005b). Plants utilizing the C₃ photosynthetic pathway (the majority of terrestrial plants) typically have values of $\delta^{13}\text{C}$ that range from –21 to –35‰ (Pataki et al., 2007). An expression for discrimination in leaves (Δ_{leaf}) of C₃ plants can be stated as follows

$$\Delta_{\text{leaf}} = a + (b - a) \frac{c_i}{c_a} \quad (1)$$

where a is the fractionation associated with diffusion (4.4‰), b is the net kinetic/enzymatic fractionation associated with carboxylation (27‰), and c_i and c_a are the intercellular and atmospheric concentrations of CO₂, respectively



Correspondence to: A. C. de Araújo
(alessandro.araujo@falw.vu.nl)

(Farquhar et al., 1982, 1989a, b). The Δ_{leaf} therefore is a function of c_i/c_a ratio, which is sensitive to a variety of factors that influence the balance of stomatal conductance and assimilation rate, for example light and water availability (Pataki et al., 2003b).

The carbon isotope ratio of a leaf ($\delta^{13}\text{C}_{\text{leaf}}$) is a measure that integrates the photosynthetic activity over the period of weeks to months during which the leaf tissue was synthesized (Ometto et al., 2002; Dawson et al., 2002; Ometto et al., 2006). In tropical forests, $\delta^{13}\text{C}_{\text{leaf}}$ is strongly correlated with the height of the leaf within the canopy. Low $\delta^{13}\text{C}_{\text{leaf}}$ (more negative) is observed in the understory vegetation, and high $\delta^{13}\text{C}_{\text{leaf}}$ (less negative) is observed in the upper canopy (Medina and Minchin, 1980; Medina, 1986; Medina et al., 1991; Vandermerwe and Medina, 1989; Sternberg, 1989; Zimmerman and Ehleringer, 1990; Broadmeadow, 1992; Buchmann et al., 1997; Guehl, 1998; Martinelli, 1998; Bonal, 2000; Ometto et al., 2002; Ometto et al., 2006). This trend in $\delta^{13}\text{C}_{\text{leaf}}$ through the canopy is related to the reassimilation of respired CO₂ and differences in conditions such as light and vapor pressure deficit through the canopy, resulting in changes in c_i/c_a ratios (Sternberg et al., 1989; Lloyd et al., 1996; Sternberg et al., 1997; Buchmann et al., 1997; Ometto et al., 2002; Ometto et al., 2006).

The temporal and spatial variability of $\delta^{13}\text{C}_{\text{leaf}}$ in forested landscapes and along environmental gradients has shown that $\delta^{13}\text{C}_{\text{leaf}}$ is more negative in riparian zones or in areas with high soil moisture content or rainfall input than in areas with low soil moisture content or rainfall input (Medina and Minchin, 1980; Ehleringer et al., 1986; Ehleringer et al., 1987; Korner et al., 1988; Korner et al., 1991; Garten and Taylor, 1992; Marshall and Zhang, 1993; Stewart et al., 1995; Sparks and Ehleringer, 1997; Hanba et al., 2000; Bowling et al., 2002). Yet, plants growing in dry environments have shown lower Δ_{leaf} and higher water-use efficiency (WUE) than those grown at low altitudes or in wet environments.

In Central Amazonia, the small-scale topography is composed of plateaus and valleys. These give rise to a high variability of soil moisture contents (θ) in the unsaturated zone, with plateaus generally having a lower θ than the valley edges in the dry season (J. S. de Souza, data not published). De Araújo et al. (in press) made nocturnal measurements of c_a along a topographical gradient at a site in Central Amazonia. They showed that in the dry season, depending on the atmospheric stability, larger amounts of CO₂ were stored on the slopes and in the valleys than on the plateaus of this undulating landscape. Lateral drainage of respired CO₂ downslope and high soil CO₂ efflux (R_{soil}) in the valleys were considered as possible causes for the observed variability in c_a . In addition, these authors observed that the CO₂ stored in the air in the valley took longer to be released than that on the plateau, and that c_a in the valley did not decrease to the same level as on the plateau at any time during the day. This leads to two hypotheses for dry season conditions. The first is that the $\delta^{13}\text{C}_{\text{leaf}}$ in the valleys may be more negative than

that on the plateaus due to both higher soil water availability and longer time of exposure to high c_a with low $\delta^{13}\text{C}_a$. The second is that the carbon isotope ratio of ecosystem respired CO₂ ($\delta^{13}\text{C}_{\text{Reco}}$) may be less negative on the plateaus than in the valleys.

This study aims to investigate how $\delta^{13}\text{C}_{\text{leaf}}$ and $\delta^{13}\text{C}_a$ vary in time and space along a topographical gradient at a site in Central Amazonia and analyses the biotic and physical factors controlling the stable carbon isotope discrimination.

2 Site description

Measurements were made at the Manaus LBA site (2°36'32"S, 60°12'33"W, 45–110 m a.s.l. – above sea level), located in the Asu catchment in the Reserva Biológica do Cuieiras. The forest belongs to the Instituto Nacional de Pesquisas da Amazônia (INPA). The exchange of CO₂, sensible and latent heat, momentum transfer, and meteorological variables have been measured almost continuously on two micrometeorological towers, installed in July 1999 and in May 2006, respectively. The first tower, known as K34 (Araújo et al., 2002), is on a medium sized plateau, whereas the second one, known as B34 (de Araújo et al. in preparation¹), is at the bottom of a U-shaped valley.

The mean air temperature was 26°C between July 1999 and June 2000 (Araújo et al., 2002). Average annual rainfall is about 2400 mm, with a distinct dry season during July, August and September when there is less than 100 mm rainfall per month (Araújo et al., 2002).

There is very little large-scale variation in topography in the region, but at a smaller scale, the dense drainage network has formed a pattern of plateaus and valleys. The mean elevation is about 100 m a.s.l. with about 40–60 m difference between plateaus and valleys bottoms. The soils along a typical toposequence consist of well-drained Oxisols and Ultisols on plateaus and slopes, respectively, and poorly drained Spodosols in the valleys (Chauvel et al., 1987). From the plateau down to the valley, the soil (top 5 cm) clay fraction decreases (from about 75% to 5%) and the sand fraction increases (from about 10% to 85%) (Ferraz et al., 1998; Chambers et al., 2004; Souza, 2004; Luizao et al., 2004).

The vegetation is old-growth closed-canopy *terra firme* (non-flooded) forest. Variation in soil type, topography and drainage status has created distinct patterns in forest vegetation composition. On the plateaus, well drained clay soils favor high biomass forests 35–40 m in height with emergent trees over 45 m tall: typical *terra firme* forest. Along the slopes, where a layer of sandy soil is deepening towards the valley bottom, forest biomass is lower and height is around 20–35 m with few emerging trees. In the valleys, the sandy soils are poorly drained and usually they remain waterlogged

¹ de Araújo, A. C., Dolman, A. J., Waterloo, M. J., Kruijt, B., et al.: Comparative measurements of CO₂ fluxes from a plateau and valley in Amazonian forest, in preparation, 2007.

during the rainy season, supporting low biomass and low tree height (20–35 m), with very few emerging trees. A distinct forest type, classified as *Campinarana* (as it resembles the *Campina* forest that develops on white sand areas), also occurs between the lower slope and valley bottom areas. This vegetation has lower biomass, tree diversity and tree height (15–25 m) (Guillaumet, 1987; Luizão et al., 2004). The forest canopy is stratified in four layers. The first layer is that formed by emergent trees, reaching heights of 35–45 m above ground level (a.g.l.). Below this layer, there are trees with their canopies between 20 and 35 m. The third layer is formed by understory regeneration, whereas shrubs and seedlings form a fourth layer close to the ground. More elaborate descriptions of the site can be found in Araujo et al. (2002), Chambers et al. (2004), Luizao et al. (2004) and Waterloo et al. (2006).

3 Material and methods

3.1 Air sampling collection and data conditioning

All sampling was carried out in representative plots along a transect that was divided into 3 topographical sections: plateau, slope and valley (see Fig. 1 in de Araujo et al., in press). In each plot, air samples were collected at different levels above and within the canopy for δ¹³C_a and c_a analysis. Each profile sampling system consisted of high-density polyethylene (HDPE) tubes (Dekorun 1300, 6.25 mm OD, non-buffering ethylene copolymer coating, USA) with intakes at different heights. Nylon funnels with stainless steel filters were installed on the air intakes to avoid sample contamination by particles. A battery-operated air pump (Capex V2X, UK) was used to draw air through the tubing, a desiccant tube containing magnesium perchlorate and a glass sample flask. The flow rate was 10 L min⁻¹. The longest air sampling tube had an internal volume of about 0.65 L that corresponds to a maximum residence time of 4 s. All air samples were collected in pre-evacuated 100 mL glass flasks that were closed with two high-vacuum Teflon stopcocks (34-5671, Kontes Glass Co., USA) after air had been pumped through the flask for about 3 minutes. The c_a was measured at the same time with an infrared gas analyzer (IRGA) (LI-800, LI-COR, Inc., USA). For this a “T” piece was connected at the air pump output, which allowed a low subsampling flow of about 800 mL min⁻¹ to be passed through the IRGA.

Plateau air samples were collected at K34 tower (118 m a.s.l.) with a tube system attached to it. The slope profile system was suspended from the highest branch of a tall tree located about midway down the slope (89.2 m a.s.l.) at 550 m from the K34 tower, whereas the valley profile system, which was suspended in the same way as that on the slope, was installed in the valley (77.3 m a.s.l.) at about 850 m from the K34 tower (de Araujo et al., in press). This latter system had its highest intake at 11 m a.g.l. in 2002 (de Araujo et

al., in press) and was extended to reach up to 30 m a.g.l. in August 2004. In October 2006, the valley profile system was relocated 500 m to the west and attached to the newly erected B34 tower.

Nighttime sampling began about one hour after sunset (about 19:30 local time) and ended about one hour before sunrise (about 05:30 local time) to avoid any effects of photosynthesis on the estimates of δ¹³C_{Reco}. In order to increase statistical confidence in δ¹³C_{Reco} values, we aimed to collect samples with a minimum c_a difference of about 75 ppm between samples, which was set a priori as the minimum difference that shall be observed among the flasks sampled at each nighttime sampling (Pataki et al., 2003a; Lai et al., 2004; Lai et al., 2005). Daytime values of δ¹³C_a and c_a within and above the canopy were obtained between 07:00 and 18:00 h. Due to both strong rainfall on 10 October and technical problems, the last sampling of atmospheric air during daytime hours occurred on 16 October at both plateau and valley.

The flasks were shipped to the Centro de Energia Nuclear na Agricultura (CENA) in São Paulo, Brazil, for stable isotope ratio and c_a analyses. Details about the analytical procedures at CENA are given by Ometto et al. (2002).

3.2 Sampling of foliage and litter

In August 2004, leaf samples were collected once from trees at each topographical section by a tree climber, sampling a vertical profile through the forest canopy. The sampling heights were not uniform among the topographical sections, as follows: plateau (3, 10, 17, 21, 24, 26, and 30 m a.g.l.), slope (3, 8, 10, 12, 20, 26, 28, 30 m a.g.l.), and valley (3, 7, 20, 25 m a.g.l.). There was no botanical classification for the trees sampled in August 2004. In October 2006, sun leaves at the top of the canopy were collected once by a tree climber at plateau and valley sections. Trees with botanical classification to species level were now systematically selected according to either their importance value index (IVI) or occurrence at both plateau and valley areas (Oliveira and Amaral, 2004; Oliveira and Amaral, 2005; I. L. do Amaral, personal communication). Each sample from a single tree consisted of at least five healthy leaves that were combined according to their status (either mature or young). In August 2004, litter samples were randomly collected at each topographical section. These were bulked by topographic section to form single samples. The samples were pre-dried at ambient air temperature for 3 days in a home-made greenhouse located in an open-sky area and shipped to CENA for stable isotope ratio and elementary analyses.

3.3 Soil-respired CO₂ sampling

In August 2004, CO₂ released from the soil was sampled at each topographical section using the protocol described by Flanagan et al. (1999) and Ometto et al. (2002). The

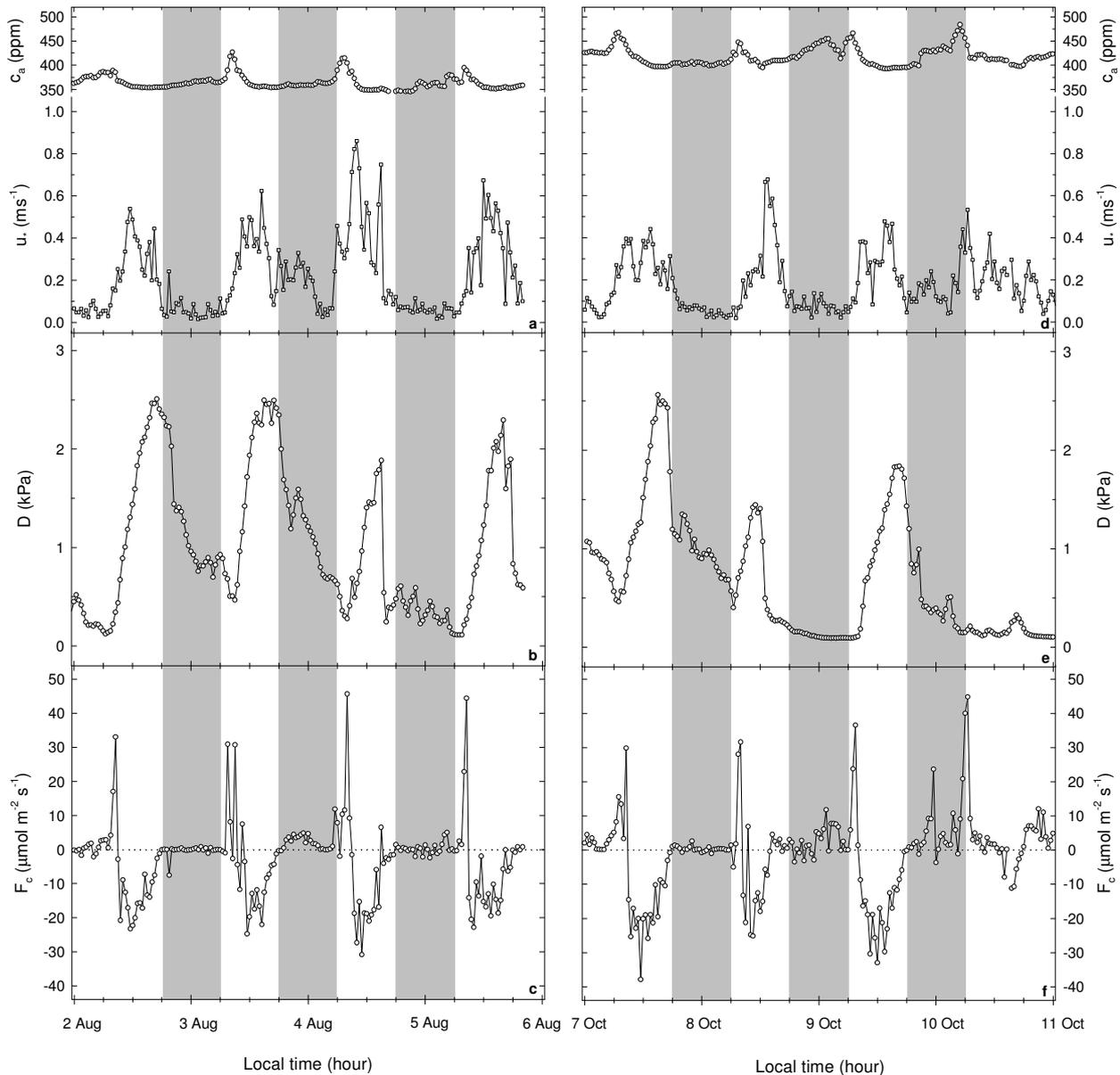


Fig. 1. Diurnal variation of some meteorological and turbulent variables measured at the top of K34 tower (53 m a.g.l. on the plateau): from 2–5 August 2004 (a–c) and from 7–10 October 2006 (d–f). Points correspond to half-an-hour averages. The shaded boxes indicate the nighttime periods. Points above the horizontal dotted line (c and f) denote CO₂ release, and below the line CO₂ uptake.

sampling was repeated at plateau and valley sites in October 2006 and now included the *Campinarana* site. Samples were collected using a stainless steel chamber with an internal volume of about 40 L and a small electric fan to enhance mixing within the chamber. Samples were collected at two sites in each topographical section. At each site, five sample flasks were filled using five minutes time intervals between sampling for determining the carbon isotope ratio of soil respired CO₂ (δ¹³C_{Rsoil}) (explained further). All samples were shipped to CENA for stable isotope ratio and *c_a* analyses.

3.4 Laboratory analyses

The δ¹³C_{*a*} in sample flasks were measured using a continuous-flow isotope-ratio mass spectrometer (IRMS) (Delta Plus, Finnigan MAT, Germany) as described by Ehleringer and Cook (1998) and Ometto et al. (2002). Measurement precision of this method was 0.13‰ for ¹³C (Ometto et al., 2002). The air remaining in the flask after stable isotope ratio analysis was used to measure *c_a* using a system similar to that described by Bowling et al. (2001a). Measurement precision and accuracy of this method were 0.2

and 0.3 ppm, respectively (Ometto et al., 2002).

Leaf and litter samples were dried at 65°C to constant weight, then ground with mortar and pestle to a fine powder. A 1–2 mg subsample of ground organic material was sealed in a tin capsule and placed into an elemental analyzer (Carlo Erba Instruments, Model EA 1110 CHNS-O, Milan, Italy) for combustion and subsequent elemental C and N analysis. The CO₂ generated by combustion was purified in a gas chromatograph column and passed directly to the inlet of the IRMS (Delta Plus, Finnigan MAT, USA) operating in continuous-flow mode.

These provided stable isotope ratios of carbon, oxygen and nitrogen (¹³C/¹²C; ¹⁸O/¹⁶O; ¹⁵N/¹⁴N) with a measurement precision of 0.2‰ (Ometto et al., 2006). The carbon isotope ratio was expressed in the delta notation (δ), which relates the measured ¹³C/¹²C molar ratio of the sample and the international Pee Dee Belemnite (PDB) limestone standard (Ehleringer and Rundel, 1989). The δ¹³C values are presented in parts per thousand (‰).

3.5 Correlation between δ¹³C_{Reco} and water vapor saturation deficit

The possibility of a correlation with fine time lag between environmental variables and δ¹³C_{Reco} may confound our analysis, so we investigated the relation between ambient vapor pressure deficit and δ¹³C_{Reco} (Ekblad and Hogberg, 2001; Bowling et al., 2002; McDowell et al., 2004; Knohl et al., 2005; Werner et al., 2006). We selected the water vapor saturation deficit in the air (*D*) as a suitable variable because it may influence δ¹³C_{Reco} at a short time scale, possibly through changes in photosynthetic discrimination (Bowling et al., 2002). Daytime mean values were used unless specified otherwise. Because there may be a significant delay between the time that a given carbon atom is assimilated by photosynthesis and the time that it is respired by various ecosystem components (Ekblad and Hogberg, 2001), correlations between *D* and δ¹³C_{Reco} were examined over a range of time lags (e.g., relationships between *D* on day *X* and δ¹³C_{Reco} on day *X*+*n*). We calculated averages of daytime *D* (from 10:00 to 17:00) from 1–5 days, and then shifted these averages back in time by 0–15 days. A 1-day average and a 0-day time lag correspond to the average daytime *D* on the day prior to the night of sampling. For a more detailed description of lag analysis see Bowling et al. (2002), Ekblad and Hogberg (2001) and McDowell et al. (2004).

3.6 Statistical analyses

3.6.1 Organic samples (δ¹³C_{leaf} and δ¹³C_{litter})

Statistical comparisons were made using Model I ANOVA, and comparisons between means were evaluated with Bonferroni *t* tests (Glover and Mitchell, 2002; Sokal and Rohlf,

1995). Unless otherwise indicated, a significance level of 99% was used in all hypotheses testing (Table 1).

3.6.2 Keeling plots for δ¹³C_{Reco} and δ¹³C_{Rsoil}

A two-source mixing model proposed by Keeling (1958) was used to obtain δ¹³C_{Reco} and δ¹³C_{Rsoil} (Flanagan et al., 1999; Ometto et al., 2002; Pataki et al., 2003a). The Model II regression or geometric mean regression (GMR) has been recommended to determine the Y-intercept (Pataki et al., 2003a). However, Zobitz et al. (2006) argued that the use of Model II regression to obtain δ¹³C_{Reco} is inappropriate because it is a biased estimator of δ¹³C_{Reco} and the relative error in the δ¹³C_a measurements is significantly greater than the relative error in *c_a* measurements. They suggested therefore the use of Model I regression or ordinary least squares (OLS) to determine the Y-intercept (*a*) and slope (*b_{y.x}*). We have decided to follow their recommendation, though we also present the slope (*v_{y.x}*) and Y-intercept (*a_v*) of Model II regression for δ¹³C_{Reco} (Table 2). Uncertainty for the Y-intercept is reported as standard error estimate from a Model I regression or standard linear regression intercept (SE_{*a*}) (Sokal and Rohlf, 1995; Pataki et al., 2003a; Zobitz et al., 2006). Because we followed the suggested guidelines made by Pataki et al. (2003a) to reduce errors when using the two-source mixing approach for estimating δ¹³C_{Reco}, the majority of the standard errors of the Y-intercept reported here are smaller than 1‰ (Table 2).

3.6.3 Correlation between δ¹³C_{Reco} and *D*

First-order linear regression was used except in cases where scatter plots suggested nonlinear or second-order equations were appropriate. The Pearson product-moment correlation coefficient, usually known by correlation coefficient (*r*), was used as the index of association of two variables (Glover and Mitchell, 2002).

4 Results

4.1 Meteorological and turbulent variables

4.1.1 Dry season campaign on 2–5 August 2004

Figure 1 shows the diurnal variation of selected meteorological and turbulent variables measured at the top of K34 tower (53 m a.g.l.) during the sampling period. Rainfall (12 mm) occurred in the late afternoon of 4 August (data not shown), with a corresponding increase in the friction velocity (*u_{*}*) from about 0.2 m s⁻¹ to 0.8 m s⁻¹ and *D* decreased from about 2 kPa to 0.5 kPa (Fig. 1a, b). Nighttime values of *u_{*}* were higher at 0.2 m s⁻¹ from 3–4 August than during the other nighttime periods (Fig. 1a), implying enhanced vertical mixing on 3–4 August. The nighttime periods from 3–4 and

Table 1. Stable carbon isotope and C:N ratios of leaves and litter along a topographical gradient in central Amazonia. Leaves were sampled in the canopy layer in August 2004 and at the top of the canopy in October 2006. The δ¹³C_{leaf}, δ¹³C_{litter}, C:N_{leaf} and C:N_{litter} are presented as average (±standard error). Averages in the same column followed by different letters are significantly different at α=0.01 (Bonferroni *t*-tests).

Date	Place	δ ¹³ C _{leaf} (‰)	C:N _{leaf}	δ ¹³ C _{litter} (‰)	C:N _{litter}	<i>n</i> ₁	<i>n</i> ₂
August 2004	Plateau	-28.86 ^a ±0.84	28.01 ^a ±5.5	-30.75 ^a	27.84 ^a	5	1
	Slope	-31.54 ^{ab} ±1.00	35.15 ^a ±5.1	-30.22 ^a	26.67 ^a	4	1
	Valley	-32.34 ^b ±0.79	35.34 ^a ±5.5	-30.12 ^a	30.84 ^a	5	1
October 2006	Plateau	-29.71 ^a ±0.32	30.92 ^a ±2.3			13	
	Campinarana	-30.31 ^b ±0.36	35.24 ^a ±2.8			11	
	Valley	-30.55 ^b ±0.25	37.88 ^a ±2.7			11	

The number of samples collected at each topographical section is presented as *n*₁ and *n*₂, and refer to leaf and litter samples, respectively.

Table 2. Statistics of Keeling plots used to obtain the δ¹³C_{Reco} along a topographical gradient in central Amazonia. The values of *a* and *a*_v are expressed in ‰ (per mil), *c*_a in ppm, and *b*_{*y*.*x*} and *v*_{*y*.*x*} in ‰ ppm.

Place	Period	<i>b</i> _{<i>y</i>.<i>x</i>}	<i>a</i>	<i>v</i> _{<i>y</i>.<i>x</i>}	<i>a</i> _v	<i>r</i>	<i>r</i> ²	SE _{<i>a</i>}	<i>c</i> _a min	<i>c</i> _a max	<i>c</i> _a range	<i>n</i>
Plateau	2–3 Aug 2004	8066.06	-29.76	8221.10	-30.14	0.98	0.96	0.91	379.5	483.7	104.2	20
	3–4 Aug 2004	8487.91	-30.79	8562.05	-30.97	0.99	0.98	0.71	388.1	476.3	88.2	16
	4–5 Aug 2004	8069.02	-29.62	8107.45	-29.71	0.99	0.99	0.43	374.3	551.9	177.6	20
	7–8 Oct 2006	7377.16	-28.02	7403.43	-28.09	0.99	0.99	0.37	374.2	513.2	139	17
	8–9 Oct 2006	7698.18	-28.70	7780.70	-28.89	0.99	0.98	0.68	408.7	456.5	47.8	17
	9–10 Oct 2006	5609.03	-23.64	5797.74	-24.04	0.96	0.94	0.81	419.4	553.4	134	17
Slope	2–3 Aug 2004	9441.98	-32.46	9605.08	-32.81	0.98	0.97	0.89	394.2	545.3	151.1	18
	3–4 Aug 2004	8278.19	-29.39	8447.39	-29.74	0.98	0.96	0.82	428.5	545.7	117.2	20
	4–5 Aug 2004	7496.50	-28.30	7517.47	-28.35	0.99	0.99	0.28	410.8	578.7	167.9	18
Valley	2–3 Aug 2004	9658.42	-32.44	9779.97	-32.69	0.99	0.98	0.73	447.7	560.6	112.9	20
	3–4 Aug 2004	6489.37	-26.08	6679.55	-26.46	0.97	0.94	0.77	455.7	558.6	102.9	19
	4–5 Aug 2004	8299.85	-29.69	8340.31	-29.78	0.99	0.99	0.40	422.2	559.6	137.4	20
	7–8 Oct 2006	7523.61	-28.13	7615.30	-28.30	0.99	0.98	0.59	476.2	595.0	118.8	16
	8–9 Oct 2006	6671.87	-26.64	7052.79	-27.48	0.93	0.89	1.29	405.1	488.4	83.3	17
	9–10 Oct 2006	5809.47	-25.10	5989.76	-25.48	0.94	0.94	0.79	418.7	523.8	105.1	14

The slope of Model I regression (*b*_{*y*.*x*}), the Y-intercept of Model I regression (*a*), the slope of Model II regression (*v*_{*y*.*x*}), the Y-intercept of Model II regression (*a*_v), the correlation coefficient (*r*), the coefficient of determination (*r*²), the standard error of the Y-intercept of Model I regression (SE_{*a*}), and the *c*_a minimum, maximum and range are presented. The SE_{*a*} higher than 1‰ is in bold.

4–5 August were ranked as the least and most stable, respectively. This was observed in the nighttime values of above-canopy turbulent exchanges of CO₂ (*F*_c) that were higher from 3–4 August (about 2.7±0.6 μmol m⁻² s⁻¹) than from 4–5 August (about 0.2±0.4 μmol m⁻² s⁻¹) (Fig. 1c).

4.1.2 Dry season campaign on 7–10 October 2006

Due to a lightning strike on the B34 valley tower neither meteorological nor turbulent data were measured at this site during the sampling period (between 27 September and 12 October). Figure 1d–f therefore shows the diurnal variability of meteorological and turbulent variables measured at the top of K34 tower on the plateau. Rainfall occurred in the morning of 8 October (0.2 mm) and 10 October (27 mm), respectively

(data not shown). Nighttime values of *u*_{*} were higher for 9–10 October than at the other nighttime periods, implying enhanced vertical mixing on 9–10 October (Fig. 1d). The nighttime periods from 9–10 and 7–8 October were ranked as the least and most stable, respectively. This was observed in the nighttime values of *F*_c that were higher from 9–10 October than from 7–8 October (Fig. 1d, f). On 8 and 9 October, between 06:00 and 08:00 local time (LT), the values of *F*_c were very high showing considerable amounts of CO₂ from ecosystem respiration (*R*_{eco}) released in a short time interval (Fig. 1f). Daytime values of *F*_c were less negative on 10 October than on 8 October due to a high morning rainfall on 10 October (Fig. 1f).

4.2 Spatial variability of δ¹³C_{leaf}, δ¹³C_{litter} and canopy and litter C:N ratio

In August 2004, the vertical profile of δ¹³C_{leaf} through the canopy showed a similar pattern for every topographical section, decreasing with depth into the canopy (data not shown). However, δ¹³C_{leaf} of canopy layer was significantly more negative in the valley than on the plateau, with values of -32.34‰ and -28.86‰, respectively (Table 1). On other hand, δ¹³C_{litter} showed no significant difference among the topographical sections (Table 1). Although the litter samples had been collected randomly, they comprise a mix of litter-fall from different canopy heights and decomposition stages on the soil surface. It is very likely that the reduced number of litter samples per topographical section may have not been representative of the site variability. The C:N ratios of leaves from the canopy layer and litter showed no significant difference between the means (Table 1).

In October 2006, δ¹³C_{leaf} at the top of the canopy was significantly more negative in the valley than on the plateau, about -30.55‰ and -29.71‰, respectively (Table 1). Yet, the C:N ratios of leaves from the top of the canopy were higher in the valley and *Campinarana* than on the plateau, though there was no significant difference between the means at the 99% level (Table 1). No significant difference was observed between the δ¹³C_{leaf} of old and new leaves sampled in October 2006 (data not shown).

4.3 Temporal and spatial variability of δ¹³C_a, c_a, δ¹³C_{Reco} and δ¹³C_{Rsoil}

4.3.1 Campaigns on 8–9 October and 17–18 November 2002

These trial campaigns, on 8–9 October and 17–18 November 2002, provided the first insights into the variation of δ¹³C_a and c_a with time and topography. They suggested that the atmospheric air below the canopy was more ¹³C depleted in the valley than on the plateau (data not shown). Yet, δ¹³C_a was uniform with height a.g.l. in the valley, whereas at both slope and plateau it was quite variable. In addition, c_a was higher in the valley and slope than on the plateau, and it was uniform with height a.g.l. in the valley (data not shown).

4.3.2 Dry season campaign on 2–5 August 2004

As in the trial campaigns, nighttime values of δ¹³C_a were significantly different among the topographical sections. The δ¹³C_a was more negative in the valley and slope plots than on the plateau (Fig. 2a–c). The δ¹³C_a difference between the canopy layer (35–20 m a.g.l.) and shrub layer (from 5 m a.g.l. downwards) was larger on the plateau than on the slope and in the valley. Post-sunset or pre-dawn values of the δ¹³C_a in the canopy layer were always more negative at both slope and valley plots than on the plateau by at least 1.5‰ or 2.5‰ respectively (Fig. 2a–c). Before dawn, on 4 August, the δ¹³C_a

measured at 30 m a.g.l. increased sharply at both slope and valley, this was not observed on 3 and 5 August (Fig. 2b, c), thus suggesting that the erosion of the nighttime buildup in the valley had already started.

Opposing the isotopic signatures along the topographical sections, the nighttime values of c_a were higher at both slope and valley than on the plateau (Fig. 2d–f). In addition, the c_a difference between the canopy layer and shrub layer was larger on the plateau than on the slope and in the valley. Before dawn, on 4 August, the c_a measured at 30 m a.g.l. decreased sharply at both slope and valley, this was different on 3 and 5 August (Fig. 2e, f). As mentioned above, this suggests that the erosion of the nighttime buildup in the valley had already started.

Figure 3 shows the relationship between c_a and δ¹³C_a for each topographical section during the three consecutive nighttime periods. Although the second-order regressions for the plateau were quite similar, their curvatures showed that for the same values of c_a the values of δ¹³C_a from 4–5 August were slightly less negative than from 2–3 August and 3–4 August, respectively (Fig. 3a). At both slope and valley, the regressions were quite variable and hard to interpret (Fig. 3b, c). The nighttime variability of c_a and δ¹³C_a was also observed in the values of δ¹³C_{Reco} (Table 2). From 2–3 August, δ¹³C_{Reco} was less negative on the plateau than at both slope and valley (Fig. 4a). On the following day, 3–4 August, δ¹³C_{Reco} became progressively less negative moving from the valley to the slope and onto the plateau. Finally, from 4–5 August, δ¹³C_{Reco} was less negative on the slope than on both plateau and valley (Fig. 4a). On the plateau, the values of δ¹³C_{Reco} agreed very well with the predictions based on the regressions in Fig. 3a. The δ¹³C_{Rsoil} was also variable among the topographical sections. It was less negative on the plateau than in the valley, but the minimum was found on the slope (Fig. 4b).

4.3.3 Dry season campaign on 7–10 October 2006

Nighttime values of δ¹³C_a were more negative in the valley than on the plateau. High δ¹³C_a was observed early in the night, whereas low δ¹³C_a occurred before dawn. The δ¹³C_a was highest on 7 October on the plateau and on 8 October in the valley, and it was lowest on October 8 at both plateau and valley (Fig. 5a, b). Before dawn on 8 October, the δ¹³C_a measured at 42 m a.g.l. on the plateau and at 30 m a.g.l. in the valley were about 4‰ and 1‰ more enriched in ¹³C, respectively, than the levels below them (Fig. 5a, b). This suggests that the erosion of the nighttime buildup on the plateau and in the valley had already started. From 8–9 October, the values of δ¹³C_a from the canopy layer downwards on the plateau and from middle layer downwards in the valley were almost uniform with height a.g.l. (Fig. 5a, b).

Nighttime values of c_a were higher in the valley than on the plateau. Low c_a was observed early at night and high c_a before dawn (Fig. 5c, d). Before dawn on 8 October, the

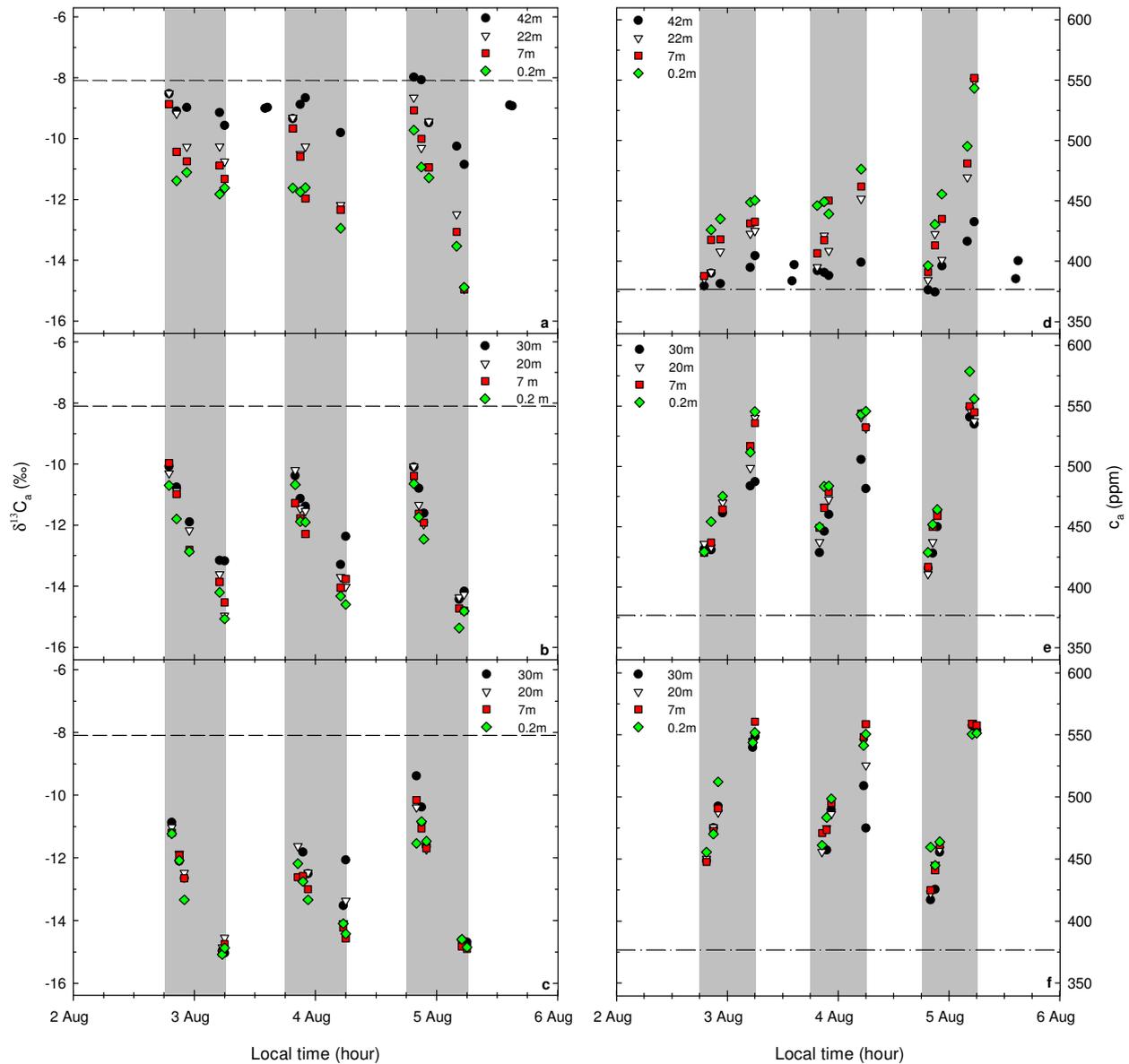


Fig. 2. Evolution of $\delta^{13}\text{C}_a$ and c_a along a topographical gradient in Central Amazonia from sunset until dawn on 2–5 August 2004: on the plateau (**a, d**), slope (**b, e**) and valley (**c, f**); and measurements of $\delta^{13}\text{C}_a$ and c_a at 42 m a.g.l. on the plateau late in the afternoon on 5 August 2004 (**a, d**). Points correspond to the single measurement made at each sampling level. The dashed line in (**a–c**) and the dash-dotted line in (**d–f**) represent the carbon isotope ratio of tropospheric background CO₂ ($\delta^{13}\text{C}_b$) and the tropospheric background [CO₂] (c_b) measured in the marine boundary layer at Ascension Island, UK (7.92°S 14.42°W; 54 m a.s.l.), respectively, on August 2004 (about -8.094% and 376.71 ppm) (White and Vaughn, 2007; Conway et al., 2007). The shaded boxes indicate the nighttime periods.

c_a measured at 42 m a.g.l. on the plateau and at 30 m a.g.l. in the valley were about 120 ppm and 30 ppm lower, respectively, than at lower levels (Fig. 5c, d). Again, as mentioned before, this suggests that the erosion of the nighttime buildup on the plateau and in the valley had already started. On 8–9 October, the values of c_a from the canopy layer downwards on the plateau and from middle layer downwards in the valley were almost uniform with height a.g.l. (Fig. 5c, d).

The relation between c_a and $\delta^{13}\text{C}_a$ for each topographical section during the three consecutive nighttime periods was also investigated. The second-order regressions for the plateau had the same shape as those shown in Fig. 3a. In addition, the curvatures of the regressions showed that for the same values of c_a the values of $\delta^{13}\text{C}_a$ from 9–10 October were less negative than those of 7–8 October and 8–9 October, respectively (data not shown). The regressions for the

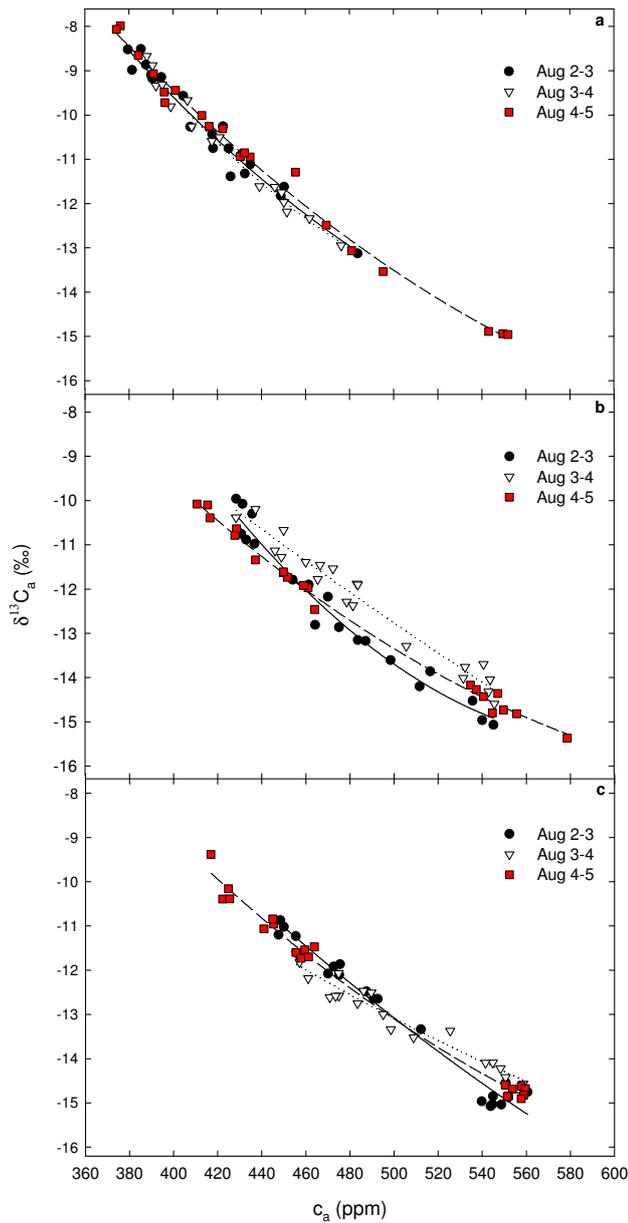


Fig. 3. Relationship between c_a and $\delta^{13}\text{C}_a$ along a topographical gradient in Central Amazonia from sunset until dawn on 2–5 August 2004: (a) on the plateau, (b) slope and (c) valley. Points correspond to the single measurement made at each sampling level. The solid, dotted and dashed lines denote the second-order regressions for the periods of 2–3 August, 3–4 August and 4–5 August, respectively.

valley also had a similar shape as those shown in Fig. 3c (data not shown).

Nighttime variability of c_a and $\delta^{13}\text{C}_a$ was also observed in the values of $\delta^{13}\text{C}_{\text{Reco}}$ (Table 2). The values of $\delta^{13}\text{C}_{\text{Reco}}$ were more negative in the valley than on the plateau on 7–8 October and 9–10 October and more positive on 8–9 October (Fig. 4c). On the plateau, the values of $\delta^{13}\text{C}_{\text{Reco}}$ agreed very well with the predictions based on the regressions. In

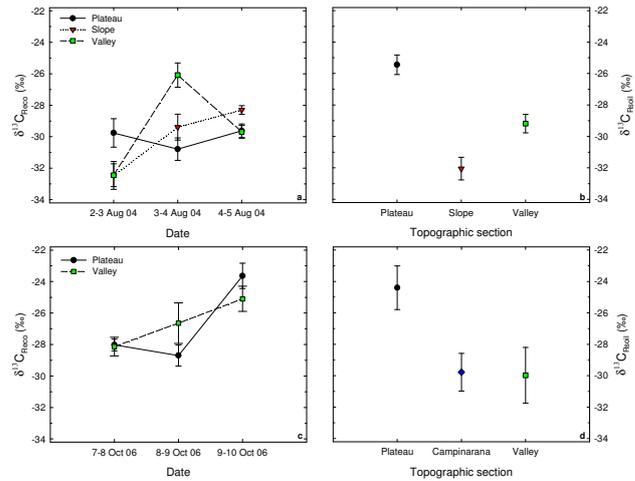


Fig. 4. $\delta^{13}\text{C}_{\text{Reco}}$ and $\delta^{13}\text{C}_{\text{Rsoil}}$ measured along a topographical gradient in Central Amazonia in August 2004 (a, b) and October 2006 (c, d). Points in a and c correspond to the Y-intercept of the Keeling plot (\pm standard error) for every nighttime period at each topographical position. Each point in b and d corresponds to the Y-intercept of the Keeling plot (\pm standard error) at each topographical position on 6 August 2004 and 9 October 2006, respectively.

addition, the values of $\delta^{13}\text{C}_{\text{Reco}}$ were higher than those measured on August 2004. The $\delta^{13}\text{C}_{\text{Rsoil}}$ was greater on the plateau than at the *Campinarana* and in the valley, respectively (Fig. 4d).

Daytime values of $\delta^{13}\text{C}_a$ were typically less negative on the plateau than in the valley by about 1‰ (Table 3).

4.4 Correlation between $\delta^{13}\text{C}_{\text{Reco}}$ and D

There were strong correlations between $\delta^{13}\text{C}_{\text{Reco}}$ and D at all topographical sections. In August 2004, the highest correlations were observed with 1 and 3-day average and 2 and 4-day lag times for plateau, slope and valley (Table 4). Figure 6a shows that on the plateau, according to Table 4, the $\delta^{13}\text{C}_{\text{Reco}}$ on 2–3, 3–4 and 4–5 August had a maximum correlation with the averaged D of 31 July, 1 and 2 August, respectively. In October 2006, the highest correlations were observed with 1 and 2-day average and 7 and 6-day lag times for plateau and valley, respectively (Table 4). Figure 6b shows that in the valley, according to Table 4, the $\delta^{13}\text{C}_{\text{Reco}}$ on 7–8, 8–9 and 9–10 October had a maximum correlation with the averaged D for the period from 30 September to 1 October, from 1 to 2 October, and from 2 to 3 October, respectively. Figure 6c shows the relationship between $\delta^{13}\text{C}_{\text{Reco}}$ and D for each topographical section, according to the results of Table 4. In August 2004, the $\delta^{13}\text{C}_{\text{Reco}}$ was more responsive to changes on the slope and in the valley than on the plateau (Fig. 6c). However, on the slope and in the valley, $\delta^{13}\text{C}_{\text{Reco}}$ ranged from about -26 to -33 ‰ with almost no variation in D . In October 2006, on both plateau and valley, $\delta^{13}\text{C}_{\text{Reco}}$

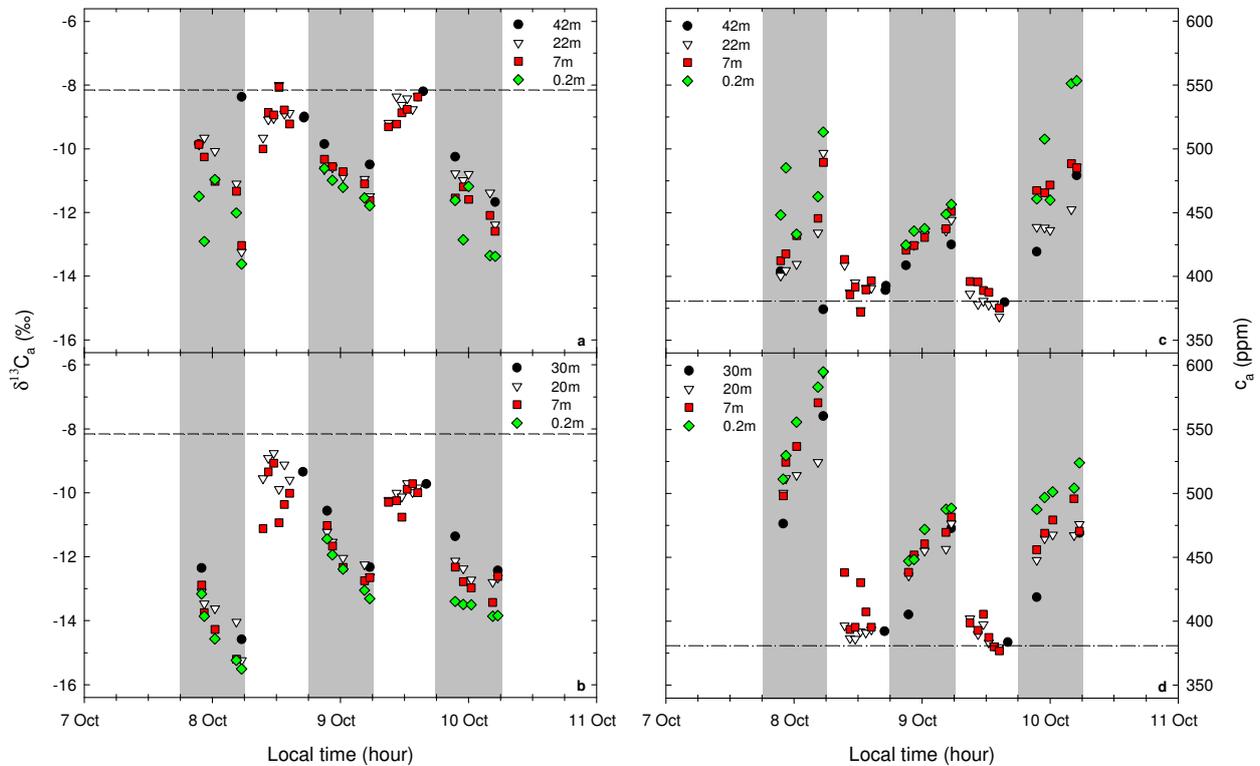


Fig. 5. Diurnal curves of $\delta^{13}C_a$ and c_a along a topographical gradient in Central Amazonia on 7–10 October 2006: on the plateau (**a**, **c**) and in the valley (**b**, **d**). Points correspond to the single measurements made at each sampling level. The dashed line (**a**, **b**) and the dash-dotted line (**c**, **d**) represent the carbon isotope ratio of tropospheric background CO₂ ($\delta^{13}C_b$) and the tropospheric background [CO₂] (c_b) measured in the marine boundary layer at Ascension Island, UK (7.92° S 14.42° W; 54 m a.s.l.), respectively, on October 2006 (about -8.156% and 380.77 ppm). The $\delta^{13}C_b$ was determined by linear interpolation using the data from October 2004 and 2005, as these were the last data available (about -8.105 and -8.152% respectively). The c_b was determined by adding the annual [CO₂] growth rate for the year 2006 (about 1.73 ppm year⁻¹) to the [CO₂] on October 2005 (about 379.04 ppm) (White and Vaughn, 2007; Conway et al., 2007). The shaded boxes indicate the nighttime periods.

Table 3. Statistics of daytime values of $\delta^{13}C_a$ measured along a topographical gradient in central Amazonia. The averaged $\delta^{13}C_a$ (\pm standard error), $\delta^{13}C_a$ min and $\delta^{13}C_a$ max are expressed in ‰ (per mil).

		Date											
		08 Oct 06				09 Oct 06				16 Oct 06			
Place	<i>z</i> (m, a.g.l.)	$\delta^{13}C_a$	$\delta^{13}C_a$ min	$\delta^{13}C_a$ max	<i>n</i>	$\delta^{13}C_a$	$\delta^{13}C_a$ min	$\delta^{13}C_a$ max	<i>n</i>	$\delta^{13}C_a$	$\delta^{13}C_a$ min	$\delta^{13}C_a$ max	<i>n</i>
Plateau	42*	-9.01	-9.02	-8.99	2	-8.20			1	-9.72			1
	22	-8.76 ± 0.2	-9.12	-7.99	6	-8.61 ± 0.1	-9.19	-8.34	6	-9.08 ± 0.2	-9.56	-8.68	6
	7	-8.90 ± 0.2	-9.33	-8.07	6	-8.91 ± 0.2	-9.31	-8.37	5	-9.61 ± 0.3	-10.87	-8.48	6
Valley	30*	-9.35			1	-9.72			1	-10.14			1
	20	-9.36 ± 0.2	-9.89	-8.91	6	-10.01 ± 0.2	-10.23	-9.84	6	-10.11 ± 0.3	-10.56	-9.74	6
	7	-10.09 ± 0.4	-11.12	-8.76	6	-10.12 ± 0.2	-10.76	-9.70	6	-10.97 ± 0.8	-12.02	-9.76	6

* Air samples collected late in the afternoon (between 15:00 and 17:00 Local Time) at these heights.

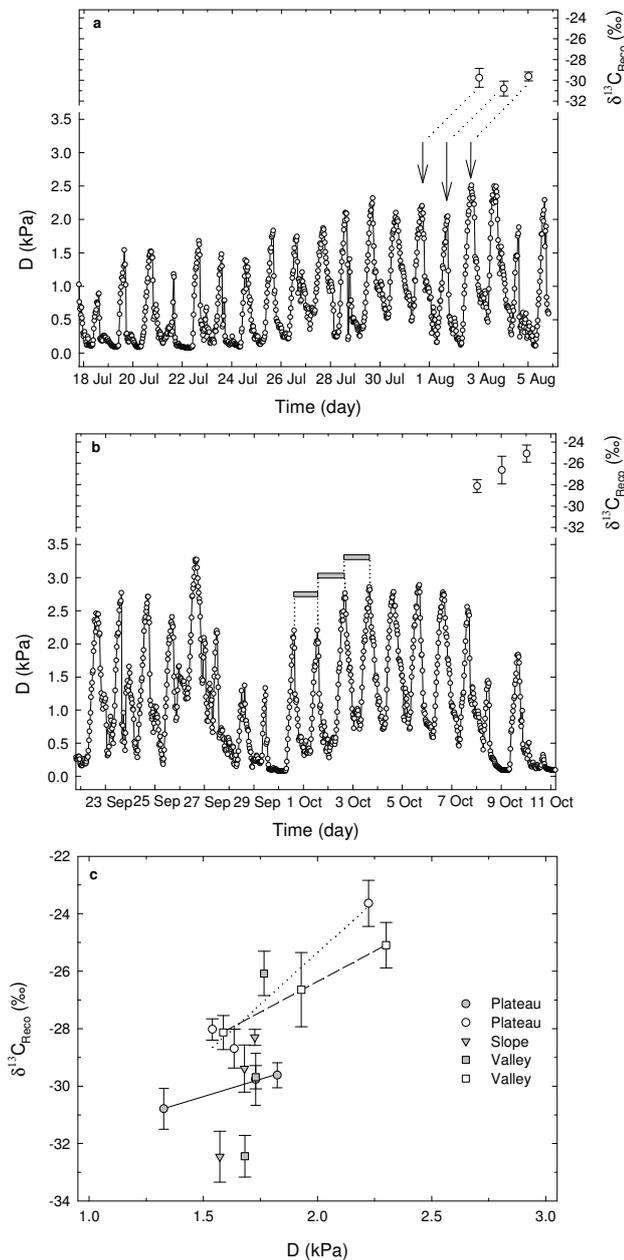


Fig. 6. Diurnal curves of half-an-hour averages of water vapor saturation deficit in the air (D) measured at the top of K34 tower (53 m a.g.l. on the plateau): from 18 July to 5 August 2004 (a) and from 22 September to 10 October 2006 (b). Carbon isotope ratio of ecosystem respired CO₂ ($\delta^{13}\text{C}_{\text{Reco}}$) (\pm standard error) on the plateau from 2–5 August 2004 (a) and in the valley from 7–10 October 2006 (b). The n-day average and n-day time lag that presented the maximum correlation between $\delta^{13}\text{C}_{\text{Reco}}$ and D on the plateau and in the valley (as in Table 4) are indicated by arrows (a) and horizontal bars (b), respectively. The length of the horizontal bar denotes the 2-day averaged (as in Table 4), and each horizontal bar has a corresponding $\delta^{13}\text{C}_{\text{Reco}}$, from left to right. (c) Relationship between $\delta^{13}\text{C}_{\text{Reco}}$ and D (averaged and time-lagged) along a topographical gradient in Central Amazonia in August 2004 (shaded symbols) and October 2006 (open symbols) according to the results of Table 4. The linear regressions for slope and valley in August 2004 were omitted for clarity. The error bars denote the standard error.

was positively correlated with D (Fig. 6c). The $\delta^{13}\text{C}_{\text{Reco}}$ was here more responsive to changes in D on the plateau than in the valley.

5 Discussion

5.1 Spatial variability of $\delta^{13}\text{C}_{\text{leaf}}$

The $\delta^{13}\text{C}_{\text{leaf}}$ decreased from plateau towards the valley (Table 1). This result is consistent with the work of Medina and Minchin (1980) in Amazonian rainforests in the southern part of Venezuela. These authors reported averaged $\delta^{13}\text{C}_{\text{leaf}}$ of -28.7 and -30.5 ‰ for the upper canopy levels of forests on lateritic outcrops and sandy spodosols soils, respectively. Increased leaf-level photosynthetic capacity of plants has been linked to higher leaf nitrogen content and leaf mass per unit area (LMA), and increased leaf thickness (Sparks and Ehleringer, 1997; Hanba et al., 2000; Vitousek et al., 1990; Korner and Woodward, 1987; Friend et al., 1989). Increased leaf-level photosynthetic capacity would decrease c_i at the carboxylation site, thus reducing Δ_{leaf} and consequently increasing $\delta^{13}\text{C}_{\text{leaf}}$ (Sparks and Ehleringer, 1997). At our study site, Luizão et al. (2004) observed that leaf nitrogen concentration was significantly higher on the plateau than in the valley. Furthermore, Nardoto (2005) showed that LMA was higher on the plateau than in the valley. These results support our findings.

The $\delta^{13}\text{C}_{\text{leaf}}$ may also be affected by c_a , $\delta^{13}\text{C}_a$ and soil moisture. Even though daytime values of c_a were about 20 ppm lower on the plateau than in the valley (Fig. 5c, d) it is unlikely that this difference would have contributed much to the observed pattern in $\delta^{13}\text{C}_{\text{leaf}}$. Daytime values of $\delta^{13}\text{C}_a$ were about 1‰ lower in the valley than on the plateau (Table 3, Fig. 5a, b). Lower $\delta^{13}\text{C}_a$ may have a significant contribution to lowering the values of $\delta^{13}\text{C}_{\text{leaf}}$ in the valley. Schulze (1986) demonstrated that leaf conductance might be more sensitive to soil moisture than photosynthesis. In this manner, it is very likely that decreased soil moisture content on the plateau would cause a decrease in leaf conductance, which implies in less diffusion of CO₂ to the interior of the stomatal chamber therefore lowering the c_i . As c_i decreases in the carboxylation site, the c_i/c_a ratio of a leaf decreases as well Δ_{leaf} and consequently $\delta^{13}\text{C}_{\text{leaf}}$ increases on the plateau. Thus, at this site, it seems that it is not only Δ_{leaf} that explains the pattern in $\delta^{13}\text{C}_{\text{leaf}}$, but rather the combination of factors such as the $\delta^{13}\text{C}_a$ in air surrounding the leaves, soil moisture availability, leaf nitrogen concentration, and LMA.

5.2 Temporal and spatial variability of $\delta^{13}\text{C}_a$ and c_a

In general, $\delta^{13}\text{C}_a$ was more negative in the valley than on the plateau at night, whereas c_a showed an opposite pattern, i.e. it was higher in the valley than on the plateau. This is consistent with the findings of de Araújo et al. (in press),

Table 4. Summary of correlation analyses between δ¹³C_{Reco} and water vapor saturation deficit in the air (*D*) at a forest site in central Amazonia in the dry season. The values of *a* are expressed in ‰ (per mil) and *b*_{*y*.*x*} in ‰ kPa⁻¹. Unless otherwise indicated, *n*=3 for every topographical section.

Date	Place	<i>n</i> -day average	<i>n</i> -day time lag	<i>r</i>	<i>r</i> ²	<i>P</i>	<i>a</i>	<i>b</i> _{<i>y</i>.<i>x</i>}
August 2004	Plateau	1	2	0.99	0.99	0.046	-33.98	2.41
	Slope	3	4	0.99	0.99	0.003	-75.36	27.31
	Valley	1	4	0.99	0.98	0.094	-158.65	74.90
October 2006	Plateau	1	7	0.97	0.94	0.160	-39.71	7.18
	Valley	2	6	0.99	0.99	0.008	-34.88	4.26

The number of days averaged (*n*-day average), the number of days lagged (*n*-day time lag), the correlation coefficient (*r*), the coefficient of determination (*r*²), the *P*-value of the regression, the *Y*-intercept of Model I regression (*a*), and the slope of Model I regression (*b*_{*y*.*x*}) are presented.

who observed that in the dry season, depending on the atmospheric stability, there was a preferential pooling of *c*_{*a*} in the lower topographical areas of this landscape.

Larger differences in δ¹³C_{*a*} and *c*_{*a*} between canopy layer and shrub layer on the plateau than in the valley may result from horizontal stratification of the nocturnal CO₂ buildup. According to de Araújo et al. (in press), *c*_{*a*} was stratified horizontally in layers of increasing concentration (from top to bottom) along the topographical gradient. They argued that horizontal stratification was caused by inversion layers that develop above and underneath the canopy. Figure 7 shows the evolution of vertical profiles of δ¹³C_{*a*} along the topographical gradient during 3–4 and 4–5 August, which is consistent with the pattern described by de Araújo et al. (in press). The δ¹³C_{*a*} measured at about 160 m a.s.l. (42 m a.g.l. on the plateau) was consistently higher than that measured at the levels below, most likely as consequence of an inversion layer that separated the canopy air from the free atmospheric air (Fig. 7). The vertical stratification of δ¹³C_{*a*} was clearer on the plateau than on the slope or in the valley, particularly before dawn, when the δ¹³C_{*a*} profiles on the slope and in the valley were fairly uniform with altitude (Fig. 7). The uniformity of δ¹³C_{*a*} and *c*_{*a*} with height in the valley suggests that the air is well mixed (vertical mixing), most likely as a consequence of the nocturnal thermal belts that might have occurred (Goulden et al., 2006; de Araújo et al., in press). Vertical mixing might also have happened on the plateau and in the valley during the nighttime period from 8–9 October, when the values of δ¹³C_{*a*} and *c*_{*a*} from the canopy layer downwards and from middle layer downwards were almost uniform with height a.g.l., respectively (Fig. 5).

The δ¹³C_{*a*} was more negative in the valley than on the plateau during daytime periods, whereas *c*_{*a*} was higher. In addition, the decrease of *c*_{*a*} with time of the day was faster on the plateau than in the valley (Fig. 5). This is consistent with the findings of de Araújo et al. (in press), who observed that in the dry season the CO₂ stored in the valley took longer to be released than that on the plateau, and that *c*_{*a*} in the valley

did not decrease to the same level as on the plateau at any time during the day. Weak vertical mixing and high *R*_{soil} (discussed below) in the valleys were considered driving the observed variability in *c*_{*a*}.

5.3 Temporal and spatial variability of δ¹³C_{Reco} and δ¹³C_{Rsoil}

The δ¹³C_{Reco} is more closely associated with that of sun foliage than with the shade foliage across a variety of ecosystems (Pataki et al., 2003a). In addition, δ¹³C_{Reco} of an entire ecosystem can be either more enriched or more depleted in ¹³C than sun foliage. This association also holds at our site, even considering the high variability in δ¹³C_{Reco} in both space and time (Table 1, Fig. 4). Yet, the averaged δ¹³C_{leaf} of the most exposed sun foliage of the dominant tree species of some tropical forests was similar to the δ¹³C_{Reco} value, thus suggesting that the major portion of recently respired CO₂ in these forests was metabolized carbohydrate fixed by the sun leaves at the top of the forest canopy (Buchmann et al., 1997; Ometto et al., 2002). The δ¹³C_{Reco} is a dynamic indicator of plant physiological response to short-term changes in environmental conditions. Tu and Dawson (2005) showed that δ¹³C_{Rsoil} (root plus microbial) is often enriched in ¹³C relative to δ¹³C_{Reco} whereas aboveground respiration (leaf plus stem) is often depleted across a variety of ecosystems. In our site, δ¹³C_{Rsoil} was higher than δ¹³C_{Reco} only on the plateau, whereas on the slope and in the valley it was the opposite (explained below) (Fig. 4). The pattern observed on the plateau agrees with Buchmann et al. (1997) and Flanagan et al. (1999), who also observed higher δ¹³C_{Rsoil} than δ¹³C_{Reco} in both tropical and boreal forests. In contrast, Ometto et al. (2002) observed that δ¹³C_{Rsoil} and δ¹³C_{Reco} were of similar magnitude in a forest about 700 km east of our site.

During the present study, the associations among δ¹³C_{leaf}, δ¹³C_{Reco} and δ¹³C_{Rsoil} held reasonably well for the plateau, whereas for the valley they did not. For example, there were periods such as on 3–4 August and on 8–9 October in which

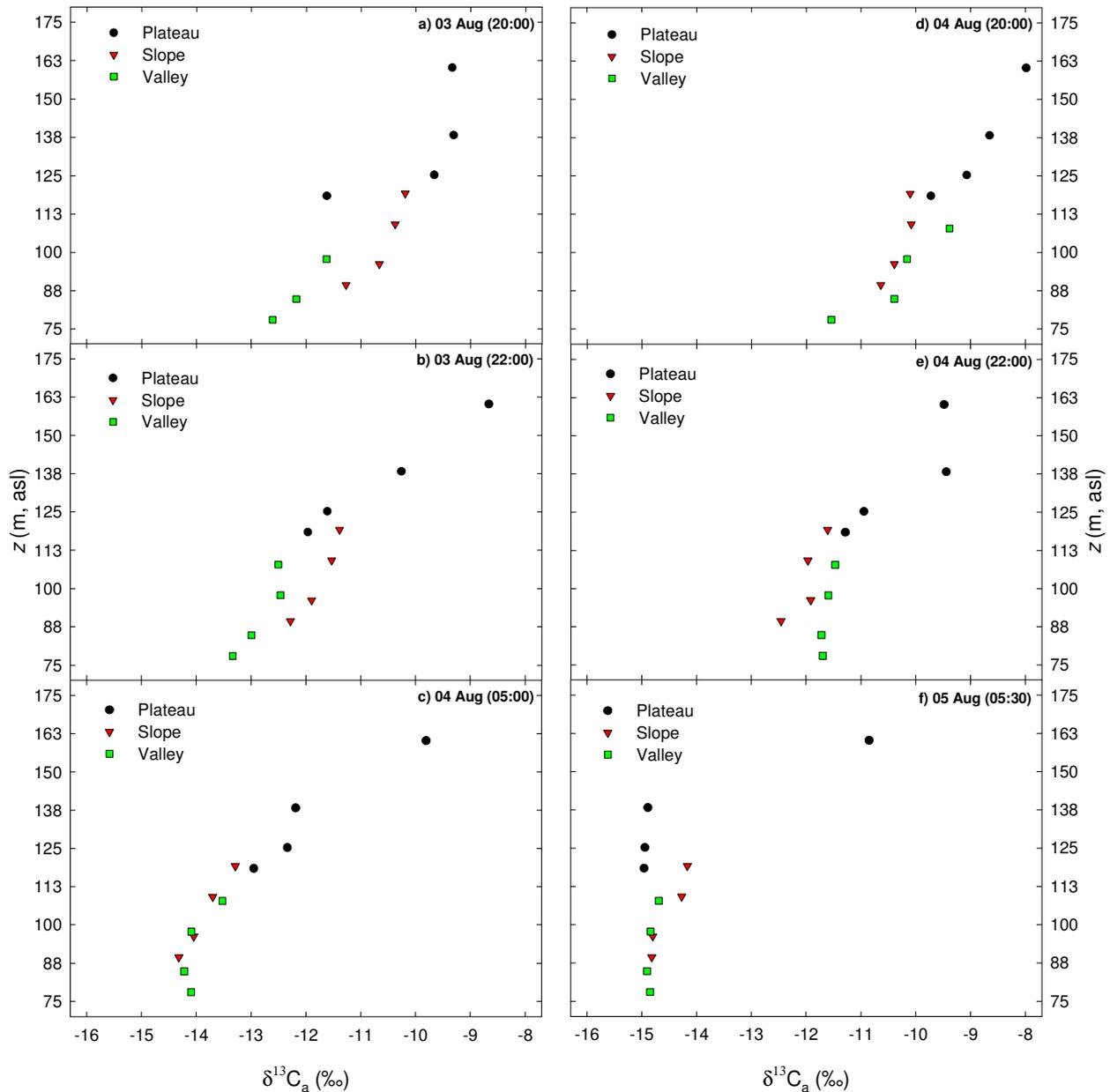


Fig. 7. Composite of vertical profiles of $\delta^{13}\text{C}_a$ from three different positions along a topographical gradient in Central Amazonia during the nighttime periods from 3–4 August (a–c) and 4–5 August (d–f) 2004. The reference altitude corresponds to the soil surface level at 850 m in the valley (about 77.3 m a.s.l.). Points correspond to the single measurement made at each sampling level. Time is presented as Local Time.

$\delta^{13}\text{C}_{\text{Reco}}$ was higher in the valley than on the plateau (Fig. 4a, c). This was somewhat unexpected because in the valley the $\delta^{13}\text{C}_{\text{leaf}}$ and $\delta^{13}\text{C}_{\text{Rsoil}}$ were well depleted in ^{13}C relative to the plateau. Galvão (2005) also observed that $\delta^{13}\text{C}_{\text{Rsoil}}$ was lower in the valley than on the plateau in the dry season of 2003. Leaf respiration and R_{soil} correspond to about 80% of R_{eco} at our site (Chambers et al., 2004). Yet, R_{soil} measurements made during the dry season of 2003 showed that R_{soil} was lower on the plateau than in the valley (Souza, 2004). Particularly during these two nighttime periods, there

was some vertical mixing as shown by the measurements of u_* and F_c (Fig. 1). It is unlikely that R_{soil} from the valley might have contributed to $\delta^{13}\text{C}_{\text{Reco}}$ being enriched in ^{13}C rather than being depleted in ^{13}C . Higher $\delta^{13}\text{C}_{\text{Reco}}$ in the valley than on the plateau therefore points to a combination of physical (mixing and transport) and biotic (respiration) processes as in de Araújo et al. (in press). These authors argued that respired CO₂ drains downslope and high R_{soil} in the valleys was driving the observed variability in c_a along this topographical gradient. Because there were no advection

measurements during the periods sampled at this study, we cannot corroborate the lateral drainage with empirical data.

5.4 Correlation between δ¹³C_{Reco} and *D*

According to the highest correlations between δ¹³C_{Reco} and *D*, the time elapsed for a given carbon atom to be assimilated by photosynthesis and to be respired by various ecosystem components varied between 2 and 7 days at this site (Table 4). Similar investigations in boreal, temperate and Mediterranean forest ecosystems have shown time lags ranging between 0 and 10 days (Bowling et al., 2002; Knohl et al., 2005; Werner et al., 2006; Ekblad and Hogberg, 2001; McDowell et al., 2004; Mortazavi et al., 2005). It is important to note that at these forest ecosystems the diversity of species per unit area is very low, whereas the forests of Central Amazonia have more than 200 species ha⁻¹ (Oliveira and Mori, 1999; Oliveira and Amaral, 2004; Oliveira and Amaral, 2005).

Prior to nocturnal sampling periods, there were several rainfall events in August 2004 rather than in October 2006 (data not shown), as it can be seen in the values of *D* from August 2004 and October 2006 (Fig. 6a, b). This might explain the time lags being shorter in August 2004 than in October 2006. We do not have a clear explanation for the time lag of plateau being shorter than for the slope and valley in August 2004 (Table 4) as it is somewhat counterintuitive, because trees are taller on the plateau than on the slope and in the valley. It is likely that the high variability of δ¹³C_{Reco} on the slope and in the valley may have contributed to this. As mentioned before, it is very likely that lateral drainage of air enriched in ¹³C from upslope areas occurs at our site, which leads to unexpected values of δ¹³C_{Reco} on the slope and in the valley. Nevertheless, in October 2006, the time lags for plateau and valley were 7 and 6-days, respectively (Table 4). This shows that the time elapsed for a carbon atom to move from foliage to the site of respiration is not constant, but rather, it probably shifts with changes in carbon allocation, tissue metabolism, dark discrimination, assimilation rates, environmental conditions, etc. (McDowell et al., 2004; Bowling et al., 2002; Knohl et al., 2005; Werner et al., 2006).

In October 2006, δ¹³C_{Reco} was more responsive to changes in *D* on the plateau than in the valley (Fig. 6c). As mentioned before, prior to nocturnal sampling periods, there was almost no rainfall in October 2006. Thus, it is likely that high *D* in combination with low soil water availability on the plateau have driven the observed pattern. Unfortunately, there were no measurements of *θ* available for the valley to corroborate our hypothesis with empirical data. However, for the plateau, there were. There was a strong negative correlation between δ¹³C_{Reco} and *θ*, and the maximum correlation was observed with 1-day average and 1-day time lag (*r*²=0.91). This is consistent with Lai et al. (2005), McDowell et al. (2004), Werner et al. (2006), Ponton et al. (2006)

and Mortazavi et al. (2005). The shorter time lag for *θ* in comparison with that for *D* suggests that soil conditions have a faster and likely more direct effect on δ¹³C_{Reco}. For example, it may indicate that the proportion of δ¹³C_{Reco} released from heterotrophic *R*_{soil} responds faster to changes in edaphic conditions (Werner et al., 2006). This effect was shown by Goulden et al. (2004) at a site in Central Amazonia.

6 Conclusions

We formulated two hypotheses to be tested in this study. The first one proposed that δ¹³C_{leaf} is more negative in the valley than on the plateau as a consequence of both higher soil water availability in the valley and longer time of exposure to high *c*_a with low δ¹³C_a in the valley than on the plateau during daytime hours. The second one proposed that the δ¹³C_{Reco} is more negative in the valley than on the plateau.

There is substantial evidence that δ¹³C_{leaf} is more negative in the valley than on the plateau (Sect. 5.1). The processes and factors that may be playing a role at our site are leaf nitrogen concentration, LMA, soil moisture availability, Δ_{leaf} and lower δ¹³C_a in the valley during daytime hours.

According to the literature, there is a strong positive relationship between δ¹³C_{leaf} and WUE. Thus, at this site, the observed pattern of δ¹³C_{leaf} might suggest that WUE is higher on the plateaus than in the valleys. However, there was no full supporting evidence for this because it remains unclear how much of the difference in δ¹³C_{leaf} was driven by physiology or δ¹³C_a.

The δ¹³C_{leaf}, δ¹³C_a and δ¹³C_{Rsoil} were more negative in the valley than on the plateau. Thus, δ¹³C_{Reco} is expected to be also more negative in the valley than on the plateau. This was observed on some nights, whereas on others it was not. The most likely explanation for this was sought in lateral drainage of CO₂ enriched in ¹³C from upslope areas, when the nights are less stable. This argument is purely based on physical factors only, such as stability parameters, lateral drainage, nocturnal thermal stratification, thermal belts, etc.

However, biotic factors, such as *R*_{soil} and the responses of plants to environmental variables such as *D* may also play a role. For example, *R*_{soil} varies spatial and seasonally along this topography and the response of heterotrophic *R*_{soil} to hydration is faster than that of autotrophic *R*_{soil}. The soluble sugars produced at the top of the trees are used at the sites of respiration (e.g. stem, leaves, and roots) and their signature should reflect the environmental conditions that prevailed when they were biosynthesized. The relationship between δ¹³C_{Reco} and *D* sheds light on this issue.

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