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# Upper Arctic Ocean water masses harbor distinct communities of heterotrophic flagellates

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#### Abstract

The ubiquity of heterotrophic flagellates (HFL) in marine waters has been recognized for several decades, but the phylogenetic diversity of these small (ca. 0.8–20 µm cell diameter), mostly phagotrophic protists in the pelagic zone of the ocean is underap-<sup>5</sup> preciated. Community composition of microbes, including HFL, is the result of past and current environmental selection, and different taxa may be indicative of food webs that cycle carbon and energy very differently. While all oceanic water columns can be density stratified due to the temperature and salinity characteristics of different water masses, the Arctic Ocean is particularly well stratified, with nutrients often limiting in <sup>10</sup> surface waters and most photosynthetic biomass confined to a subsurface chlorophyll maximum (SCM) layer. This physically well-characterized system provided an opportunity to explore the community diversity of HFL across a wide region, and down the water column. We used high-throughput DNA sequencing techniques as a rapid means of surveying the diversity of HFL communities in the southern Beaufort Sea (Canada),

- targeting the surface, the SCM and just below the SCM. In addition to identifying major clades and their distribution, we explored the micro-diversity within the globally significant but uncultivated clade of marine stramenopiles (MAST-1) to examine the possibility of niche differentiation within the stratified water column. Our results strongly implied that HFL community composition was determined by water mass rather than geograph-
- <sup>20</sup> ical location across the Beaufort Sea. Future work should focus on the biogeochemical and ecological repercussions of different HFL communities in the face of climate driven changes to the physical structure of the Arctic Ocean.

#### 1 Introduction

Small phagotrophic protists, often referred to as heterotrophic flagellates (HFL), are ubiquitous and can account for a significant proportion of the microbial eukaryotic biomass in the marine pelagic zone, including in the Arctic Ocean (Sherr and Sherr,



2009). HFL are the defining feature of microbial food webs and impact the global carbon cycle by grazing bacteria that take up organic carbon, thus recycling carbon in the water column and restraining carbon burial to the deep sea (Azam and Malfatti, 2007). As well, HFL are grazed by zooplankton, channelling carbon and energy to higher food webs (Jürgens and Massana, 2008). However, unlike large phytoplank-5 ton, most HFL lack easily preserved and readily identifiable morphological characters, and potential taxonomic and functional diversity of HFL is rarely considered. Indeed, most ecological studies and biogeochemical models place them into a single functional guild (e.g., Forest et al., 2013). However, the single functional grouping is guestionable because molecular techniques have revealed that representatives of HFL are found 10 across nearly the entire spectrum of eukaryotic diversity (Massana, 2011). In addition, HFL communities are rarely dominated by a single species or species complex (Lovejoy et al., 2006; Massana et al., 2004), suggesting that taxa are sensitive indicators of external environmental forces (Jones and Lennon, 2010) and different taxa may be indicative of food webs that cycle carbon and energy very differently. Alternatively, if 15 taxa occur randomly, grouping them as a single guild may be sufficient for conceptual and numeric modeling applications.

The Arctic Ocean is considered more quiescent than other oceans because of extensive ice cover and strong salinity-stratification (Rainville et al., 2011). This persistent stratification means that for much of the Arctic Ocean and surrounding seas the euphotic zone is nutrient-limited, and that much of the productivity takes place within a subsurface chlorophyll maximum (SCM) layer where inorganic nutrients are available and irradiance levels are sufficient for photosynthesis by eukaryotic phytoplankton (Lovejoy et al., 2007; Martin et al., 2010). Given the recent changes in the hydrological

regime of the Arctic Ocean (Jackson et al., 2012; Tsubouchi et al., 2012), identifying characteristic taxa that occur in different water masses and depths will be crucial for predicting the consequences of such changes on carbon and energy cycling. Sequence-based 18S rRNA gene surveys are rapidly becoming a standard method to identify and compare distributions of small eukaryotes in marine environments,



providing a tool for investigating diversity and distribution of taxa in pelagic systems (Massana et al., 2006). To date, environmental 18S rRNA gene surveys in the Arctic Ocean have focused on surface or SCM (Lovejoy et al., 2010; Comeau et al., 2011) water masses, but rarely both or over synoptic scales; no studies have yet system-<sup>5</sup> atically addressed the vertical distribution of microbial eukaryotes inhabiting waters

immediately below the SCM, that is, outside the zone of active photosynthesis.

To test whether the HFL communities can indeed be considered a single guild down the water column or if they reflect a potential functional partitioning, we targeted three distinct water masses in the upper Beaufort Sea water column and hypothesized that

- HFL would be sensitive indicators of their vertically structured environment. Microbial eukaryotes were surveyed by way of high-throughput "tag pyrosequencing" (Sogin et al., 2006) targeting ~ 400 nucleotides within the V4 hyper-variable region of the 18S rRNA gene (Comeau et al., 2011). Using this molecular taxonomic approach, our goal was twofold. First, by identifying the major HFL that live in surface, SCM and be-
- <sup>15</sup> low the SCM water masses, we tested whether or not these communities were distinct from each other. Our second objective was to determine whether pyroreads covering this region of the 18S rRNA gene contain sufficient phylogenetic signal to explore diversity at high resolution. To examine the possibility of niche differentiation, we analysed the micro-diversity of a widespread, but as yet uncultivated, clade of heterotrophic ma-
- rine stramenopiles known as MAST-1 (Massana et al., 2006). We chose MAST-1 because they are likely very specious, and based on comparable similarity at the level of 18S rRNA genes in the clade, they have undergone recent rapid radiation with the potential for ecological specialization (e.g., Hawlitschek et al., 2012).

#### 2 Material and methods

To study Arctic HFL assemblages, we collected 50–0.2 µm plankton in the Beaufort Sea in August 2009 aboard the CCGS Amundsen as part of the French-Canadian International Polar Year Program Malina. Specifically, four stations were sampled over



one week (Fig. 1a, Tables 1 and S1); three of these stations were east of the Mackenzie Canyon (stations 430, 460 and 540) and one was west of the canyon (station 760). Conductivity, temperature and depth (CTD) profiles were taken using a Sea-Bird SBE-911 mounted on a rosette system which was also fitted with fluorometer (SeaPoint) and

- in situ optical nitrate probe (Satlantic MBARI-ISUS). Samples for community analysis were selected on the downward cast from the fluorescence profiles to determine the depth of the SCM. To investigate the vertical structure of HFL communities, we sampled four depths at each station; surface arctic mixed layer (AML), just above the SCM, the SCM and waters immediately below the SCM where the fluorescence signal had
   disappeared (these are referred to here as the SML, SCM*a*, SCM and SCM*b*, respectively.
- tively). Water samples for DNA were collected from 12-L Niskin type bottles closed on the upward cast.

DNA was extracted from cells collected on filters representing both "small" (0.2–3 μm) and "large" (3–50 μm) size fractions. Extracted DNA from the >3 μm size fractions was amplified using 18S rRNA gene V4 region specific 454 primers as described in Comeau et al. (2011). Amplicons were mixed in equal quantities and run on 2/8th of a plate, which was sequenced on a Roche 454 GS-FLX Titanium platform at the IBIS/Université Laval Plateforme d'Analyses Génomiques. The raw pyroreads have been deposited in the NCBI Sequence Read Archive (SRA) with accession number X (NCBI BioProject submission identifier: SUB153435; *final identifier pending*).

Raw pyroreads were quality controlled and chimeras were detected using UCHIME (Edgar et al., 2011); in addition, pyroreads smaller than 300 nucleotides were discarded from the study. Retained pyroreads were aligned using the Silva eukaryotic alignment (Pruesse et al., 2007) as a template and clustered into OTUs at the  $\geq$  98 % similarity

<sup>25</sup> level using Mothur v1.21.1 (Schloss et al., 2009) as described in Comeau et al. (2011). We discarded singleton OTUs from our subsequent analyses (note that we used pyroreads from two additional Beaufort Sea stations to cluster OTUs and to identify singleton OTUs in order to maximize the number of sequences to keep; these additional station sequences were removed and will be treated in a separate study). Resulting





OTUs were taxonomically identified using Mothur and a user-designed taxonomy outline and reference sequence database (Comeau et al., 2011), trimmed only to the V4 region. Following initial analysis of both temperature-salinity (TS) data and community similarity, the pyroreads from SCM and SCM*a* for each station, were merged and are

5 referred to as SCM samples from here on. Sequences classified as known HFL were selected and all other sequences were discarded from this study and will be dealt with in a subsequent article.

Distribution and statistical analyses were computed in the R environment v2.12.1 (ww.r-project.org) and plots were generated using the R package ggplot2 (Wickham, 2009). All statistical analyses were performed on subsampled pyroread and OTU datasets (sizes: 2900 and 800, respectively).

We used the Fast-UniFrac webserver (Hamady et al., 2010) to compute all UniFrac distances between sampled HFL communities and generate corresponding UPGMA clusters and principal coordinate analysis (PCoA) plots. UniFrac distances were cal-

<sup>15</sup> culated using a large scale phylogenetic tree reconstructed based on the Mothur/Silva alignment of all heterotroph-like 454-reads and using FastTree v2.1 (Price et al., 2010) in "accurate mode" (-mlacc 2 -slownni) with the general time reversible (GTR) model and pseudocounts. Corrected p-values from UniFrac significance tests were computed using normalized UniFrac weighted distances with 500 permutations.

To generate the MAST-1 reference phylogenetic tree, we first retrieved "long" reference 18S rRNA gene sequences (i.e., originating from Sanger sequencing) from Genbank based on published MAST phylogenies (e.g., Lin et al., 2012). We also added putative MAST-1 18S rRNA gene sequences recently submitted to Genbank (as of June 2012) identified through MAST-1 specific BLASTn searches (Altschul et

al., 1997) as well as a group of most immediate outgroup sequences. These reference and outgroup 18S rRNA gene sequences and the pyroreads classified as MAST-1 (total 970 sequences) were then aligned with Mothur using the Silva eukaryotic alignment as a template. Resulting alignment was inspected and 5'/3' trimmed using Seaview v4 (Gouy et al., 2010). Aligned reference sequences were extracted and used to



reconstruct the reference MAST-1 phylogenetic trees with RAxML v7.2.8 (Stamatakis, 2006). Specifically, we used GTR model and gamma rate variation among sites, and the best phylogenetic tree was identified from 100 maximum-likelihood reconstruction runs; node statistical supports of this MAST-1 reference tree were computed from 1000 rapid bootstrap trees. Pyroreads classified as MAST-1 were placed onto the reference phylogenetic tree using RAxML Evolutionary Placement Algorithm (Berger and Stamatakis, 2011).

#### 3 Results

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#### 3.1 Environmental characteristics and general results

<sup>10</sup> Although all sampled stations were close to the main discharge of the Mackenzie River, we found no freshwater or river signal in CTD measurements (Fig. S1, Table 1). TS properties of the four depths indicated that the samples fell within three water masses previously identified in the Beaufort Sea (Carmack and Macdonald, 2002; McLaughlin et al., 2005) with the surface samples taken from the AML, the SCM and the SCM*a* 15 samples taken from Bering Sea Summer Water (BSSW) and the SCM*b* from Bering Sea Winter Water (BSWW).

After sequence quality filtering and chimera detection (Table S2), we clustered all pyroreads into operational taxonomic units (OTUs or phylotypes) using a ≥ 98 % sequence-identity cutoff; all OTUs comprised of a single pyroread (singletons) were discarded at this point. We then classified the OTUs taxonomically using an in-house curated database based on NCBI taxonomy (Comeau et al., 2011). For this study, we targeted and kept only pyroreads assigned to known HFL taxonomic groups. After preliminary taxonomic analyses, we merged the pyroreads from both SCM*a* and SCM samples because the communities were globally similar in terms of composition for

those two depths at all stations; in addition, a detrended correspondence analysis using untransformed values of temperature, salinity, nitrate concentration, fluorescence



and photosynthetically active radiation (PAR) also grouped the SCM*a* and SCM together. These two merged depths are hereafter referred as SCM. The final dataset of 12 samples consisted of four stations spanning three water masses each.

Based on taxonomic prediction outputs, we kept 8697 pyroreads unambiguously
classified as HFL out of a total of 59 409 pyroreads representing 14.6% of the total (Table S3). HFL sequences (Fig. 1b) ranged from 12.4% of all OTUs (station 460 SCM; 7.5% of total station pyroreads) to 23% (station 760 SCMb; 28.7% of total station pyroreads). The HFL pyroreads were grouped as 784 distinct OTUs out of a total of 4052 OTUs, representing 19.35% of the total. At the higher taxonomic group level (Adl et al. 2012), the HFL OTUs were classified (Fig. 1c, Table S3) as Marine Stramenopiles (MASTs; 156 OTUs from 2390 pyroreads), which are paraphyletic, mostly uncultivated protists assumed to be phagotrophic (Massana et al., 2004), biliphytes (115/1198), choanoflagellates (37/219), Rhizaria (296/2758) and Telonemia (177/2128). We found a negligible number of sequences classified as Centroheliozoa and Katablepharido-phyta (2 and 3 sequences, respectively) and none classified as Apusozoa or Amoebozoa.

#### 3.2 Vertical HFL taxonomic distributions

The proportion of HFL pyroreads to total pyroreads (a proxy for relative abundance) as well as the number of HFL OTUs (a measure of taxonomic diversity) did not vary significantly by station (Fig. 1b, Table S3). In contrast, there were significant differences with relative depth. The SCM samples were consistently poorer in both HFL pyroreads and HFL OTUs compared to samples from surface (ANOVA and Tukey's HSD test; pyroreads: p = 0.09; OTUs: p < 0.05) and SCMb ( $p \le 0.01$ ; p < 0.05). We caution here that this is likely due to an increase in phytoplankton species in SCM samples occupyion "accuration energy".

ing "sequencing space". SCMb samples were populated with more HFL pyroreads and OTUs compared to other samples, consistent with HFL being favored in the deeper waters. The only exception was station 430 for which both HFL OTUs and pyroreads were more abundant surface waters.



We found no significant differences in distributions for the HFL taxonomic groups among stations, but differences across depth category were marked. Specific major groups accounted for many of these differences. For example, overall diversity of biliphyte OTUs was greater in the surface and SCM samples than in deeper wa- $_5$  ters (p < 0.05; Fig. 1c). In addition, the proportion of biliphyte pyroreads tended to be greater in SCM than SCM*b* waters, but this was not significant (p = 0.08). Surface waters were significantly richer in OTUs and pyroreads classified as MASTs (p < 0.001) and Telonemia (p < 0.005). Although not significant, choanoflagellate sequences were consistently more abundant in surface samples (Fig. S2). For all SCM*b* samples, HFL communities were overwhelmingly dominated by Rhizaria (p < 0.001), which represented more than half of all HFL OTUs and pyroreads.

#### 3.3 Phylogenetic structure of HFL communities

To investigate the composition of the different HFL communities and their (dis)similarities, we conducted phylogenetic beta diversity analyses using UniFrac distances (Lozupone and Knight, 2005). Similar to HFL sequence distributions, phylogenetic structures of HFL communities were significantly different between water masses (UniFrac weighted significance test;  $p \le 0.012$ ) but not between stations. Likewise, when applying weighted UniFrac metrics to jackknifed cluster analysis, we observed a significant relationship between HFL community compositions and water mass cate-

- <sup>20</sup> gory, indicating that HFL communities collected from surface, SCM and SCMb waters were distinct from each other (Fig. 2a). The same relative depth-dependent similarities in HFL community composition were recovered using principal coordinate analysis (PCoA; Fig. 2b) of UniFrac weighted distances. Furthermore, HFL communities clustered by water mass in both unweighted UniFrac and Bray-Curtis distance met-
- rics (Fig. S3a and S3b, respectively), which indicates that the rarer OTUs followed the same trends as the more abundant OTUs. Altogether, these results indicate a clear water mass segregation in the composition of HFL communities in the Beaufort Sea, with distinct assemblages occurring in the AML, SCM and SCM*b* samples.



To determine if the water mass signal could be found at finer scales within main HFL lineages, we next deconstructed the weighted UniFrac cluster analysis by HFL taxonomic groups (Fig. S4). The clustering patterns were conserved overall for most major HFL clades, with the exception of the choanoflagellate and Telonemia lineages. Biliphyte communities from surface samples were globally more similar to each other than 5 biliphytes from deeper samples (i.e., SCM and SCMb). Rhizaria and MASTs were the HFL lineages with the most clearly depth-influenced composition. Within the Rhizaria, Cercozoa OTUs occurred at the surface, whereas deeper communities were mainly Polycystinea OTUs (Fig. 3). In addition, the various communities of MASTs perfectly clustered based on their water mass of origin (Fig. S4), indicating MASTs were specific 10 to the depth categories we sampled, and were good water mass biomarkers. Within the two exceptional lineages that did not follow water mass distributions, Telonemia communities followed station location, whereas choanoflagellates did not trend with either water mass or station.

#### **3.4** MAST subclade distribution using a focused phylogenetic approach

Because of the strong water mass signal that structured their community, we next focused on the distributions of the different MAST clades across depths, as well as among stations. Our taxonomic classifications assigned OTUs to distinct previously delineated MAST clades: MAST 1, 2, 3, 7, 8 and unclassified MASTs. MAST-1 sub-

- <sup>20</sup> clades (1a, 1b and 1c), MAST-2 and MAST-7 were the most abundant in terms of OTU numbers, indicating they were the most diverse (Fig. 4). Overall, we identified clear trends between water masses and MAST diversity from OTU distributions. OTUs within MAST-1a, which was globally the most diverse MAST-1 subclade, were more diverse in surface waters and less diverse at the SCM. In contrast, OTUs classified as MAST-
- 1c were more diverse at the SCM compared to surface or deeper waters. Generally, MAST-2 OTUs were more abundant in surface waters while MAST-7 OTUs were more abundant in SCM and SCMb waters. In sum, we detected water mass sorting of MAST phylogenetic diversity at the clade and subclade level.



To further investigate taxonomic distribution in the water masses by way of in depth taxonomic resolution, we used a phylogenetic placement approach (Berger and Stamatakis, 2011; Matsen et al., 2010; Monier et al., 2008). To this end we used the MAST-1 clade as a case study because of their clear water mass influenced OTU distri-

- <sup>5</sup> bution, and the availability reference sequences that are needed to produce a reference phylogeny with a strong phylogenetic signal. We first built a MAST-1 reference phylogenetic tree using an alignment of nearly full-length published 18S rRNA gene sequences originating from Sanger clone sequencing. The reference tree (Fig. 5, left phylogram) shows clear distinctions between four MAST-1 subclades, all with high bootstrap sup-
- port. Next, to gain semi-quantitative information, we mapped the pyroreads classified as MAST-1 (not the OTU numbers) onto the MAST-1 reference tree. This phylogenetic mapping revealed two patterns within the MAST-1a and MAST-1c sub-clades. First, although MAST-1a had greater diversity in surface waters, numbers were dominated by a single group; furthermore, a high number of reads was also detected at the SCMb.
- <sup>15</sup> Second, concordantly with MAST-1c diversity distribution, total and relative pyroreads were the highest at the SCM, confirming a water mass selectivity of MAST-1 (Fig. 5, right cladogram).

#### 4 Discussion

Here we provide the first comprehensive survey of HFL communities using highthroughput sequencing in a stratified upper water column. Our study highlighted the partitioning of HFL assemblages by taxa at all levels, showing evidence of environmental selection over relatively small vertical spatial scales. When using phylogenetic beta-diversity distance metrics, with and without relative abundance information, the water mass taxon distribution remained robust.

<sup>25</sup> The separation of taxa was evident at all taxonomic levels within the major recognized eukaryotic groupings (AdI et al., 2012). The separation within the Rhizaria, which includes three morphologically and functionally diverse lineages, Foraminifera,



Cercozoa and Radiolaria (Burki et al., 2010; Cavalier-Smith, 2003), was particularly striking. Surface Rhizaria (Fig. S4) pyroreads and OTU numbers were mainly in the Cercozoa (Fig. 3) with matches to taxa previously reported from Arctic surface waters (Lovejoy and Potvin, 2011; Lovejoy et al., 2006) including *Cryothecomonas*. Indeed,

- <sup>5</sup> Thaler and Lovejoy (Thaler and Lovejoy, 2012) used *Cryothecomonas* fluorescence in situ hybridization (FISH) to map the distribution of *Cryothecomonas* across the Canadian Arctic and found it exclusively in surface waters; their analysis pointed to *Cryothecomonas* being closely associated with sea-ice. Another recent study of sea-ice from the Beaufort Sea, using the same high throughput tag sequencing approach reported
- <sup>10</sup> here, found that *Cryothecomonas* was a major contributor to the sea ice communities (Comeau et al., 2013). The offshore surface waters of the Malina study region were particularly impacted by ice (Matsuoka et al., 2012), consistent with the occurrence of *Cryothecomonas*. On the other hand, Radiolarians, especially Acantharea, and Polycystinea sequences dominated the HFL assemblages of the SCM and SCM*b* (Fig. 3),
- <sup>15</sup> as was earlier reported from cloning and sequencing studies in the Canada Basin and Amundsen Gulf (Lovejoy and Potvin, 2011; Lovejoy et al., 2006; Terrado et al., 2009). Nearly identical sequences have been reported from deep Pacific low oxygen waters, and their occurrence in Pacific origin BSWW and BSSW has been rather constant across years and regions of the Canadian Arctic (Lovejoy and Potvin, 2011). It has long
- <sup>20</sup> been stated that Radiolaria follow species-specific depth distributions (Russell, 1927), and we found greater diversity and abundance of Radiolaria in the SCM and SCMb. They were likely the dominant HFL within and below the SCM, preying both on bacteria and other flagellates (Matsuoka et al., 2007). Additional temporal and geographic data may uncover more species-specific ecological information, and eventually Radiolaria
- <sup>25</sup> might be a good marker to detect the origin of upwelled surface waters. Along these same lines, a deeper understanding on the molecular diversity of this group would be useful for monitoring pulses of warm Atlantic waters entering the Arctic Ocean (Bjorklund et al., 2012).



While Cercozoa were common in surface waters, MAST and Telonemia pyroreads dominated HFL assemblages, both at the diversity and relative abundance levels (Fig. 1c). All of the MAST clades reported here (Fig. 4) have been previously recovered in clone libraries from the Beaufort and other Arctic Seas (Lovejoy and Potvin, 2011;

- Lovejoy et al., 2006; Terrado et al., 2009; 2011), although it is intriguing that none was found under perennially ice-covered waters near the North Pole (Bachy et al., 2011). Previous reports have shown that different MAST clades may be segregated across spatial scales, which might be due to different prey affinities and availability (Lin et al., 2012; Piwosz and Pernthaler, 2010). Using pyrosequencing, Logares and colleagues
- (Logares et al., 2012) reported differences in MAST assemblages, at the phylogenetic level, between deep chlorophyll maximum and surface samples from European coastal waters, from a variety of sampling sites. We found significant partitioning of MAST communities down the water column (Figs. 5 and S4), but not among sampling sites. The lack of geographical patterns could be due to the more limited area we sampled (i.e.,
- the coastal Beaufort Sea) compared to the study from Logares and colleagues, which included the coastal Mediterranean Sea and English Channel. Extending our survey of HFL communities beyond the Beaufort Sea may reveal geographical differences in Arctic MAST communities.

MASTs belonging to clades 1, 2, 3 and 7 are thought to be ubiquitous in Arctic waters (Lovejoy et al., 2006), and previous reports suggested that their distributions are shaped by prey affinities and availability (e.g., Lin et al., 2012). Comeau et al. (2013) recently noted that MASTs tended to co-occur with diatoms in sea ice. Similarly, in a late winter-spring study, MAST-1a and 1c were absent from clones libraries constructed based on rRNA template (Terrado et al., 2011), and these authors suggested that this

may have been linked to low concentrations of bacteria available for grazing. Comeau et al. (2011) reported a decrease in the proportion of MAST sequences in the Amundsen Gulf SCM following the accelerated seasonal ice-cover loss since 2007.

Overall, MASTS are not only taxonomically diverse but also likely occupy a number of ecological roles. Our analysis of MAST communities showed a strong vertical



specificity at all phylogenetic levels, likely indicating that prey preference and availability shape these communities. Surface MAST communities differed phylogenetically from those in deeper water masses. Specifically, MAST-1a and -2 were overall more diverse in all surface samples; this higher diversity in MAST-1a and -2 phylotypes appears as a characteristic of surface waters from the Beaufort Sea.

The use of a pyrosequencing approach enabled us to recover many more sequences than reported from clone libraries. Despite this, in our taxonomic survey we did not recover pyroreads classified as MAST-4 in any of the sampled communities. Although the MAST-4 clade is one of the most widely distributed and abundant MAST clades in most oceans, it is generally absent from polar waters (Massana et al., 2006; Rodríguez-Martínez et al., 2009) and the few records of its presence have been related to Pacific

Martínez et al., 2009) and the few records of its presence have been related to Pacific water inflows (Comeau et al., 2011; Lovejoy and Potvin, 2011). MAST-4 cells were first reported to feed on heterotrophic bacteria (Massana et al., 2006) but more recently, they have been shown to prey on the picocyanobacteria *Synechoccocus* (Lin et al.,

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<sup>15</sup> 2012). The Arctic Ocean is noteworthy for lacking picocyanobacteria (Tremblay et al., 2009; Waleron et al., 2006), a feature shared by the cold waters around Antarctica (Ghiglione and Murray, 2011; Wilkins et al., 2012). We hypothesize that the absence of MAST-4 from polar waters could be linked to the absence of its cyanobacterial prey.

MAST -3 is reported to be globally the most abundant of the MAST clades (Logares et al., 2011) and MAST-3 sequences have been previously reported from this same

- region of the Arctic (Lovejoy et al., 2006; Terrado et al., 2011). However, we found relatively few MAST-3 sequences in this study and their occurrence was sporadic, being found only in the surface and SCM depths of two stations. Recently Gómez and colleagues (Gómez et al., 2011) suggested that MAST-3 cells may be parasites of
- <sup>25</sup> diatoms, because of 18S rRNA gene phylogenetic affinity with Solenicola setigera, which is a well known diatom epibiont. However, the ecology of the clade is still largely speculative. For example, Terrado and colleagues (Terrado et al., 2012) cloned and sequenced 18S rRNA from surface and BSSW samples collected from mid-March to mid-May in Amundsen Gulf, using both DNA and RNA as templates. They found that



MAST-3 was most common in the RNA template libraries in mid-March, prior to diatoms becoming abundant.

Telonemia, was only recently described and, although accepted as a phylum level taxon (Shalchian-Tabrizi et al., 2007), its taxonomic affinities to other eukaryotes remains unresolved. These phagotrophic predators appear to be widely distributed in marine environments (Bråte et al., 2010). Microscopy-based and molecular-based surveys in the Arctic have reported Telonemia from Arctic surface waters (Bråte et al., 2010; Lovejoy et al., 2006; 2002; Terrado et al., 2011; Vørs, 1993), as well as sea-ice (Bachy et al., 2011; Ikävalko and Gradinger, 1997; Majaneva et al., 2011; Różańska
et al., 2008; Sazhin et al., 2004; Werner et al., 2007; Comeau et al., 2013). Telonemia could well be specialists since they are most often found in surface waters, and, for example, they were absent from deeper Western North Atlantic waters (Countway et al., 2007) as well as winter mesopelagic waters (150–200 m) in the Beaufort Sea (Terrado et al., 2009). We found that both Telonemia pyroreads and OTUs were signifi-

- <sup>15</sup> cantly greater in surface waters (Table S3); however, we did not detect any depth effect on Telonemia diversity. In contrast to the other groups, local scale geographical influence was noted (Fig. S4). An earlier study suggested that there may be at least one Arctic restricted clade of Telonemia (Shalchian-Tabrizi et al., 2007), but this was later reported to be an artefact caused by undersampling (Bråte et al., 2010). The uniformity
- <sup>20</sup> of very closely related phylotypes occurring sporadically over different geographic regions suggest that Telonemia may be able to survive long distance transport and then 'bloom' under precise conditions.

"Biliphyte" or "picobiliphyte" sequences are commonly retrieved in the Arctic Ocean (Comeau et al., 2011; Hamilton et al., 2008; Not et al., 2007; Terrado et al., 2011).
The term picobiliphyte was first used to describe this phylum-level taxon of uncertain phylogenetic affinities. The first sequences belonging to 'biliphytes' were retrieved from samples that had been pre-filtered through a 3 µm filter (picoplanktonic size) and an 18S rRNA specific FISH probe revealed phycobilin-like fluorescence in many of the FISH positive cells (Not et al., 2007). A subsequent review of related sequences



from additional sites suggested that they were not uniformly pico-sized and were referred to then as biliphytes (Cuvelier et al., 2008). A recent study based on single-cell genome fragments failed to detect any plastid-related genes implying they were strict phagotrophs (Yoon et al., 2011). The current phylogenetic position based on multiple

- genes from the Yoon and colleagues' study (Yoon et al., 2011) puts them outside of any known photosynthetic groups (Adl et al., 2012). Biliphytes have been shown to increase in number during dark incubations (Weber et al., 2012) and it is now thought that the phycobilin-like fluorescence reported earlier may be due to the presence of ingested cyanobacterial prey (Kim et al., 2011). However, the paucity of cyanobacteria
   in the Arctic Ocean (Waleron et al., 2006) suggests that biliphyes in the Arctic have
- alternative prey, for example heterotrophic bacteria or picophytoplancton.

Within the HFL community, the choanoflagellates did not segregate by depth or station (Fig. S4). Choanoflagellates are spherical or ovoid cells with a funnel-shaped collar of microvilli surrounding a unique flagellum (King et al., 2009). The use of molecular

- <sup>15</sup> markers has shown that there is an emerging diversity based on environmental surveys (Del Campo and Massana, 2011) as well as a cryptic diversity within described species (Stoupin et al., 2012). These organisms feed on small prey (< 1 µm; Marchant and Scott, 1993), creating a current with their flagellum to drive algal and bacterial prey to the collar of microvilli that acts as a filter (Kiørboe, 2011). Choanoflagellates are morphologically very diverse, and include both single cells and larger colonies, and</p>
- they could well specialize on particular prey and also be prey to larger heterotrophs. Although they are often found in Arctic Ocean and other waters, they are rarely abundant.

#### 4.1 HFL in a dynamic Arctic

<sup>25</sup> Until recently, the cold, perennially ice covered Arctic Ocean was considered a stable and predictable environment where food webs were short with nutrients-triggered diatom-dominated phytoplankton blooms following ice melt. These diatom blooms supported large zooplankton populations at the base of the food web. This view gives rise



to a basic nutrient-phytoplankton-zooplankton-detritus (NPZD) type food web model that is widely used in ecosystem studies and often forms the base of the biological component of global climate models (Poulin and Franks, 2010). The simple NPZD model in the Arctic ignores the contribution of microbial cells such as bacteria, archaea, small
 phytoflagellates and heterotrophic or parasitic protists that persist throughout the year (Galand et al., 2009a,b; Terrado et al., 2008). A corollary of the phylogenetic selection and diversity of HFL communities observed here draws a picture of further complexity in the transfer of carbon in microbial food webs. Figure 6 illustrates, at relatively short

- vertical scales (Surface, SCM and SCMb), the flow of carbon through HFL commu nities. The upper mixed layer in the Arctic remains depleted in nitrate through much of the year, driving the formation of a SCM at the pycnocline created by BSSW that becomes closely associated with the nitracline (Martin et al., 2010). The input of nitrate and silicate from the BSSW drives much of the primary production in the Beaufort Sea region, and the lack of nutrients in the surface waters limits productivity, while low
- light levels impede photosynthesis below the SCM. Polar regions are currently warming more rapidly than other regions of the planet, resulting in a freshening of the Arctic from terrestrial (permafrost) melt, increased discharge of large rivers and melting multiyear ice, impacting the hydrography of the region (McLaughlin et al., 2005). The loss of summer ice cover has also been linked to changes in the size structure of the phy toplankton communities (Li et al., 2009) and changes in major microbial species and species groups (Comeau et al., 2011).

#### 4.2 Future directions

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Recent studies using molecular markers have shown that that the species composition of small phytoplankton varies both seasonally and by depth (Giovannoni and Vergin, 2012; Treusch et al., 2011), implying strong environmental selection of particular phylotypes. In contrast to photosynthetic microbes, including phytoplankton, which are

vertically stratified due to a strong selection by easily identifiable abiotic factors such as irradiance levels or nutrient availability (Demir-Hilton et al., 2011; Johnson et al.,



2006), little is known about the drivers that structure of HFL communities within the water column.

Prey type, size and availability within the water column likely drives selection for particular HFL taxa. To reveal predator-prey interactions, detection of significant co-occurrences using a network analysis approach linking eukaryotic, bacterial and archaeal diversities would be desirable. Ecological studies integrating network analyses have so far focused on a single domain of life, specifically bacterial OTU co-occurrences (Barberán et al., 2011; Chaffron et al., 2010). The use of such approaches demands a large number of samples in order to identify significant statistical relation-ships (Barberán et al., 2011), and given the relatively low number of samples in our study, this kind of analysis is out of the scope of this paper. Earlier clone library results (Lovejoy and Potvin, 2011) showed that the physically stable Beaufort Sea water column (Carmack, 2007) harbours distinct eukaryotic communities within different water

masses. Our results from high throughput sequencing enabled us to explore taxonomic
 trends of HFL with much greater detail; a comparison of the nutrient-depleted surface
 mixed layer, the SCM and below the SCM, highlighted very different grazer communi ties, suggesting differences in the prey characteristics within the three water masses.

## Supplementary material related to this article is available online at: http://www.biogeosciences-discuss.net/10/3397/2013/

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**Table 1.** Metadata for the Beaufort Sea stations and depths from which eukaryotic microbial communities were sampled (Table S1 lists additional metadata).

430Surface $18/08/09$ $71.13^{\circ}$ N, $136.42^{\circ}$ W $3$ $-0.8$ $25.93$ $0.35$ $0.04$ $64.42$ 430SCMa $18/08/09$ $71.13^{\circ}$ N, $136.42^{\circ}$ W $55$ $-0.94$ $31.11$ $0.73$ $0.33$ $2.44$ 430SCM $18/08/09$ $71.13^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 430SCMb $18/08/09$ $71.13^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 430SCMb $18/08/09$ $70.40^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.25$ $32.07$ $11.11$ $0.27$ $0.41$ 460Surface $19/08/09$ $70.40^{\circ}$ N, $136.03^{\circ}$ W $45$ $-1.21$ $30.61$ $1.36$ $0.1$ $1.98$ 460SCM $19/08/09$ $70.40^{\circ}$ N, $136.03^{\circ}$ W $66$ $-1.23$ $32.07$ $10.43$ $0.29$ $0.21$ 460SCMb $19/08/09$ $70.45^{\circ}$ N, $136.03^{\circ}$ W $-1.23$ $32.07$ $10.43$ $0.29$ $0.21$ 540SCMa $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $-1.23$ $30.8$ $0.75$ $0.13$ $5.36$ 540SCMb $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $-1.12$ $31.7$ $4.33$ $0.55$ $1.79$ 540SCMb $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $-1.12$ $31.7$ $4.33$ $0.55$ $1.79$ 540SCMb $17/08/09$ $7$	Station	Depth Category	Date (d/m/yr)	Latitude, Longitude	Depth (m)	<i>Т</i> (°С)	Salinity (PSU)	Nitrate (ISUS)	Fluo (RU)	PAR (μE m <sup>-2</sup> s <sup>-1</sup> )
430SCMa18/08/09 $71.13^{\circ}$ N, $136.42^{\circ}$ W55 $-0.94$ $31.11$ $0.73$ $0.33$ $2.44$ 430SCM18/08/09 $71.13^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 430SCMb18/08/09 $71.13^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 430SCMb18/08/09 $71.13^{\circ}$ N, $136.42^{\circ}$ W $65$ $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 460Surface19/08/09 $70.40^{\circ}$ N, $136.03^{\circ}$ W $4$ $0.07$ $25.38$ $0.79$ $0.09$ $20.75$ 460SCMa19/08/09 $70.40^{\circ}$ N, $136.03^{\circ}$ W $4$ $0.07$ $25.38$ $0.79$ $0.09$ $20.75$ 460SCMa19/08/09 $70.40^{\circ}$ N, $136.03^{\circ}$ W $56$ $-1.21$ $30.61$ $1.36$ $0.1$ $1.98$ 460SCMb19/08/09 $70.40^{\circ}$ N, $136.03^{\circ}$ W $-1.23$ $32.07$ $10.43$ $0.29$ $0.21$ 460SCMb19/08/09 $70.45^{\circ}$ N, $137.53^{\circ}$ W $-0.61$ $26.24$ $0.66$ $0.04$ $95.91$ 540SCMa17/08/09 $70.45^{\circ}$ N, $137.53^{\circ}$ W $-1.12$ $31.7$ $4.33$ $0.55$ $1.79$ 540SCMb17/08/09 $70.45^{\circ}$ N, $137.53^{\circ}$ W $-1.13$ $32.15$ $10.18$ $0.2$ $0.67$ 760Surface12/08/09 $70.33^{\circ}$ N, $140.47^{$	430	Surface	18/08/09	71.13° N, 136.42° W	3	-0.8	25.93	0.35	0.04	64.42
430SCM $18/08/09$ $71.13^{\circ}N, \\ 136.42^{\circ}W$ 65 $-1.04$ $31.63$ $4.31$ $1.1$ $1.24$ 430SCMb $18/08/09$ $71.13^{\circ}N, \\ 136.42^{\circ}W$ 80 $-1.25$ $32.07$ $11.11$ $0.27$ $0.41$ 460Surface $19/08/09$ $70.40^{\circ}N, \\ 136.03^{\circ}W$ $4$ $0.07$ $25.38$ $0.79$ $0.09$ $20.75$ 460SCMa $19/08/09$ $70.40^{\circ}N, \\ 136.03^{\circ}W$ $45$ $-1.21$ $30.61$ $1.36$ $0.1$ $1.98$ 460SCM $19/08/09$ $70.40^{\circ}N, \\ 136.03^{\circ}W$ $56$ $-1.03$ $31.14$ $1.78$ $1.28$ $1.14$ 460SCMb $19/08/09$ $70.40^{\circ}N, \\ 136.03^{\circ}W$ $56$ $-1.03$ $31.14$ $1.78$ $1.28$ $1.14$ 460SCMb $19/08/09$ $70.40^{\circ}N, \\ 136.03^{\circ}W$ $60$ $-1.23$ $32.07$ $10.43$ $0.29$ $0.21$ 540Surface $17/08/09$ $70.45^{\circ}N, \\ 137.53^{\circ}W$ $-0.61$ $26.24$ $0.66$ $0.04$ $95.91$ 540SCM $17/08/09$ $70.45^{\circ}N, \\ 137.53^{\circ}W$ $-1.13$ $32.15$ $10.18$ $0.2$ $0.67$ 540SCMb $17/08/09$ $70.45^{\circ}N, \\ 137.53^{\circ}W$ $3$ $0.51$ $22.45$ $1.44$ $0.09$ $69.71$ 540SCMa $17/08/09$ $70.33^{\circ}N, \\ 140.47^{\circ}W$ $30.31$ $0.99$ $0.11$ $3.45$ 760SCMa $12/08/09$ $70.33^{\circ}N, \\ 140.47^{\circ}W$ $-1.26$ $32.06$ $9.81$	430	SCM <i>a</i>	18/08/09	71.13° N, 136.42° W	55	-0.94	31.11	0.73	0.33	2.44
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460Surface19/08/0970.40° N, 136.03° W40.0725.380.790.0920.75460SCMa19/08/0970.40° N, 136.03° W45 $-1.21$ 30.611.360.11.98460SCM19/08/0970.40° N, 136.03° W56 $-1.03$ 31.141.781.281.14460SCMb19/08/0970.40° N, 	430	SCM <i>b</i>	18/08/09	71.13° N, 136.42° W	80	-1.25	32.07	11.11	0.27	0.41
460SCMa19/08/0970.40° N, 136.03° W45 45 $-1.21$ 30.611.360.11.98460SCM19/08/0970.40° N, 136.03° W56 	460	Surface	19/08/09	70.40° N, 136.03° W	4	0.07	25.38	0.79	0.09	20.75
460SCM19/08/09 $70.40^{\circ}$ N, 136.03° W56 -1.03-1.0331.141.781.281.14460SCMb19/08/09 $70.40^{\circ}$ N, 136.03° W80 	460	SCM <i>a</i>	19/08/09	70.40° N, 136.03° W	45	-1.21	30.61	1.36	0.1	1.98
460SCMb19/08/0970.40° N, 136.03° W80 $-1.23$ $-1.23$ $32.07$ 10.43 $0.29$ $0.21$ 540Surface17/08/0970.45° N, 137.53° W $-0.61$ $26.24$ $0.66$ $0.04$ $95.91$ 540SCMa17/08/0970.45° N, 	460	SCM	19/08/09	70.40° N, 136.03° W	56	-1.03	31.14	1.78	1.28	1.14
540Surface $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $3$ $-0.61$ $26.24$ $0.66$ $0.04$ $95.91$ 540SCMa $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $50$ $-0.93$ $30.8$ $0.75$ $0.13$ $5.36$ 540SCM $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $70$ $-1.12$ $31.7$ $4.33$ $0.55$ $1.79$ 540SCM $17/08/09$ $70.45^{\circ}$ N, 	460	SCM <i>b</i>	19/08/09	70.40° N, 136.03° W	80	-1.23	32.07	10.43	0.29	0.21
540SCMa17/08/0970.45° N, 137.53° W50 -0.93-0.9330.80.750.135.36540SCM17/08/0970.45° N, 137.53° W70 	540	Surface	17/08/09	70.45° N, 137.53° W	3	-0.61	26.24	0.66	0.04	95.91
540SCM $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $70$ $-1.12$ $31.7$ $4.33$ $0.55$ $1.79$ 540SCMb $17/08/09$ $70.45^{\circ}$ N, $137.53^{\circ}$ W $85$ $-1.3$ $32.15$ $10.18$ $0.2$ $0.67$ 760Surface $12/08/09$ $70.33^{\circ}$ N, $140.47^{\circ}$ W $0.51$ $22.45$ $1.44$ $0.09$ $69.71$ 760SCMa $12/08/09$ $70.33^{\circ}$ N, 	540	SCM <i>a</i>	17/08/09	70.45° N, 137.53° W	50	-0.93	30.8	0.75	0.13	5.36
540SCMb17/08/0970.45° N, 137.53° W85 $-1.3$ $-1.3$ 32.1510.180.20.67760Surface12/08/0970.33° N, 140.47° W0.5122.451.440.0969.71760SCMa12/08/0970.33° N, 	540	SCM	17/08/09	70.45° N, 137.53° W	70	-1.12	31.7	4.33	0.55	1.79
760Surface $12/08/09$ $70.33^{\circ}$ N, $3$ $0.51$ $22.45$ $1.44$ $0.09$ $69.71$ 760SCMa $12/08/09$ $70.33^{\circ}$ N, $50$ $-1.1$ $30.31$ $0.99$ $0.11$ $3.45$ 760SCM $12/08/09$ $70.33^{\circ}$ N, $70$ $-1.09$ $31.4$ $2.52$ $0.48$ $1.19$ 760SCMb $12/08/09$ $70.33^{\circ}$ N, $90$ $-1.26$ $32.06$ $9.81$ $0.17$ $0.34$	540	SCM <i>b</i>	17/08/09	70.45° N, 137.53° W	85	-1.3	32.15	10.18	0.2	0.67
760SCMa $12/08/09$ $70.33^{\circ}$ N, $50$ $-1.1$ $30.31$ $0.99$ $0.11$ $3.45$ 760SCM $12/08/09$ $70.33^{\circ}$ N, $70$ $-1.09$ $31.4$ $2.52$ $0.48$ $1.19$ 760SCMb $12/08/09$ $70.33^{\circ}$ N, $90$ $-1.26$ $32.06$ $9.81$ $0.17$ $0.34$	760	Surface	12/08/09	70.33° N, 140.47° W	3	0.51	22.45	1.44	0.09	69.71
760         SCM         12/08/09         70.33° N, 140.47° W         70         -1.09         31.4         2.52         0.48         1.19           760         SCMb         12/08/09         70.33° N, 140.47° W         90         -1.26         32.06         9.81         0.17         0.34	760	SCM <i>a</i>	12/08/09	70.33° N, 140.47° W	50	-1.1	30.31	0.99	0.11	3.45
760 SCM <i>b</i> 12/08/09 70.33°N, 90 –1.26 32.06 9.81 0.17 0.34 140.47°W	760	SCM	12/08/09	70.33° N, 140.47° W	70	-1.09	31.4	2.52	0.48	1.19
	760	SCM <i>b</i>	12/08/09	70.33° N, 140.47° W	90	-1.26	32.06	9.81	0.17	0.34

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Fig. 1. (A) Beaufort Sea stations sampled during the leg 2B of the Malina cruise (August 2009). (B) Overall distributions of HFL pyroreads and OTUs (clustered at  $\geq$  98% sequence identity) of sequenced Beaufort Sea microbial eukaryotic communities. Purple, green and blue bars represent surface, SCM and SCM*b* depths, respectively. Solid and transparent color bars represent OTU and pyroread distributions, respectively. (C) OTU and pyroread distributions of main HFL taxonomic groups across Beaufort Sea samples.



**Fig. 2. (A)** Weighted Unifrac cluster analysis of HFL OTUs (clustered at  $\geq$  98 % sequence identity). OTU relative abundances were normalized across samples. Clustering statistical supports were computed using 100 jackknife replicates. **(B)** Corresponding Principal Coordinate Analysis (PCoA) using UniFrac weighted distance metric.





**Fig. 3.** Overall Rhizaria OTU distributions across Beaufort Sea samples. OTUs (clustered at  $\geq$  98 % sequence identity) were taxonomically classified using an in-house curated sequence database based on NCBI taxonomy. The distributions of each sample in distinct Rhizaria clades are displayed according to color-legend.

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**Fig. 4.** Overall MAST OTU distributions across Beaufort Sea samples. OTUs (clustered at  $\geq$  98 % sequence identity) were taxonomically classified using an in-house curated sequence database based on NCBI taxonomy. The distributions of each sample in distinct MAST sub-clades are displayed according to color-legend.



**Fig. 5.** Phylogenetic mapping of MAST-1 pyroreads from the Beaufort Sea. MAST-1 reference phylogenetic tree (right phylogram) was reconstructed using maximum-likelihood and node statistical supports were computed using 1000 rapid bootstrap replicates using RAxML. Pyroreads classified as MAST-1 were mapped onto the MAST-1 reference tree using RAxML evolutionary placement algorithm (left cladogram). Phylogenetic placements are indicated by red branches. Distributions of phylogenetic placements are indicated by color bars (Purple, green and blue bars represent surface, SCM and SCM*b* depths, respectively). First number corresponds to percent of total depth and second number is total number of a given depth mapped.







**Fig. 6.** Schematic of the trophic networks in the upper Arctic Ocean in light of HFL taxonomic information within the upper water column: Surface, SCM and SCM*b*. Bacteria counts and phytoplankton are scaled to actual values available at the Malina database (http://malina.obs-vlfr.fr). For phytoplankton smaller circles represent the small fraction and outer circles the total from extracted chl *a* values.