

1 **Modeling the impact of agricultural land use and management on U.S. carbon budgets**

2 B. A. Drewniak¹, U. Mishra¹, J. Song², J. Prell¹, V. R. Kotamarthi¹

3 ¹Environmental Science Division

4 Argonne National Laboratory

5 9700 S. Cass Ave

6 Argonne, IL 60439

7

8 ²Northern Illinois University

9 Department of Geography

10 Davis Hall, Room 118

11 DeKalb, IL 60115

12

13 Correspondence: Beth Drewniak

14 E-mail: bbye@anl.gov

15 Phone: 630-252-3732

16 Fax: 630-252-5880

17

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19

20 **Abstract**

21 Cultivation of the terrestrial land surface can create either a source or sink of atmospheric CO₂,
22 depending on land management practices. The Community Land Model (CLM) provides a useful
23 tool to explore how land use and management impact the soil carbon pool at regional to global
24 scales. CLM was recently updated to include representation of managed lands growing maize,
25 soybean, and spring wheat. In this study, CLM-Crop is used to investigate the impacts of various
26 management practices, including fertilizer use and differential rates of crop residue removal, on
27 the soil organic carbon (SOC) storage of croplands in the continental United States over
28 approximately a 170-year period. Results indicate that total U.S. SOC stocks have already lost
29 over 8 Pg C (10%) due to land cultivation practices (e.g., fertilizer application, cultivar choice,
30 and residue removal), compared to a land surface composed of native vegetation (i.e.,
31 grasslands). After long periods of cultivation, individual subgrids (the equivalent of a field plot)
32 growing maize and soybean lost up to 65% of the carbon stored, compared to a grassland site.
33 Crop residue management showed the greatest effect on soil carbon storage, with low and
34 medium residue returns resulting in additional losses of 5% and 3.5%, respectively, in U.S.
35 carbon storage, while plots with high residue returns stored 2% more carbon. Nitrogenous
36 fertilizer can alter the amount of soil carbon stocks significantly. Under current levels of crop
37 residue return, not applying fertilizer resulted in a 5% loss of soil carbon. Our simulations
38 indicate that disturbance through cultivation will always result in a loss of soil carbon, and
39 management practices will have a large influence on the magnitude of SOC loss.

40

41 **1. Introduction**

42 Bioenergy crops are promoted as a renewable energy source capable of improving energy
43 security and mitigating greenhouse gas (GHG) emissions from fossil fuels. These crops are
44 considered environmentally friendly and economically competitive, because CO₂ emitted by
45 biofuel combustion is partially balanced by atmospheric uptake through photosynthesis (Hughes
46 et al., 2010). The Renewable Fuel Standard of the U.S. Energy Independence and Security Act
47 (EISA 2007) sets a national target of producing 136 billion liters of renewable fuels by 2022. Of
48 this, at least 61 billion liters is expected to come from cellulosic ethanol (U.S. Environmental
49 Protection Agency, 2010). Though maize grain and sugarcane are currently the major global
50 sources for bioethanol production, maize production in the United States is not sufficient to meet
51 the renewable fuel targets. Furthermore, recent studies suggest that production of ethanol from
52 corn grain might in fact increase GHG emissions because of changes in land use (Searchinger et
53 al., 2008; Kim et al., 2009; Melillo et al., 2009). For these reasons, cellulosic biofuels produced
54 from cellulose and hemicellulose plant biomass are considered a viable alternative to
55 conventional crop-based biofuels.

56 Cellulosic biofuels can be made from perennial feedstocks or from residues of annual
57 cropping and forestry activities, thereby reducing or eliminating the need for additional
58 agricultural land. The use of crop residues for bioethanol production shows promise for fulfilling
59 U.S. renewable fuel goals, but more research is needed on the effects on soil organic carbon
60 (SOC) of crop residue removal from croplands (Mishra et al., 2013) and net GHG balance
61 (McKone et al., 2011). Furthermore, crop residues play a crucial role in sustainability and
62 resilience of agroecosystems (Karlen et al., 2009). Therefore, to understand the environmental
63 consequences of using crop residues for bioenergy production on large spatial scales, it is

64 essential to know the impacts on the SOC pool of differential rates of crop residue removal and
65 nitrogenous fertilizer applications.

66 Crop residue is responsible for maintaining soil moisture, returning carbon and other
67 nutrients to soil, and erosion mitigation; in general, it provides a sustainable environment for
68 cultivation activities (Lal, 2009). Without residue cover, wind and water erosion will increase
69 (Van Pelt et al., 2013). Long-term residue harvest results in loss of yields and productivity by
70 decreasing the nutrient content of soils (Blanco-Canqui and Lal, 2009a). These arguments
71 demonstrate that using crop residues as a bioenergy fuel resource could have detrimental impacts
72 on agroecosystems (Blanco-Canqui and Lal, 2009a).

73 Globally, soils store more carbon than the atmosphere and biosphere combined, acting
74 both as a source and sink of atmospheric CO₂ (IPCC, 2013). However, cultivation loss of SOC
75 ranges from 50% to 70% (Lal and Bruce, 1999). Over the U.S. Midwest, land conversion led to a
76 25-50% reduction of soil carbon (Houghton et al., 1999; Lal, 2002). The result is large carbon
77 payback times, ranging from a few years to several centuries (Fargione et al., 2008; Gibbs et al.,
78 2008; Searchinger et al., 2008). On the other hand, conversion from cultivation to native
79 grasslands, such as through enrollment in the Conservation Reserve Program, resulted in
80 increased soil carbon (Anderson-Teixeira et al., 2009; Pineiro et al., 2009). Therefore, it is
81 critical to evaluate the impact of agricultural land use and management on regional carbon
82 budgets.

83 The influence of agriculture on the carbon cycle is complex; carbon capture and storage
84 in croplands are dependent on management practices, including tillage, fertilizer applications,
85 residue management, and crop sequence (West and Post, 2002; Hooker et al., 2005; Dou and

86 Hons, 2006; Huggins et al., 2007; Khan et al., 2007; Kim et al., 2009). SOC stocks and fluxes at
87 a particular location are soil and site specific and reflect the long-term balance between organic
88 matter inputs from vegetation and losses due to decomposition, erosion, and leaching. Some
89 studies have attempted to quantify carbon sequestration from mitigation strategies such as no-till
90 or conservation tillage practices, residue management, use of cover crops, and restoration and
91 reserve actions (Conant et al., 2001; West and Post, 2002). These studies showed that as farming
92 techniques are improved to maximize yield and minimize disturbance, SOC can be maintained
93 and perhaps even increased over time.

94 However, the effect of altered management on agricultural soil's ability to store or emit
95 carbon is unresolved, largely as a result of conflicting evidence. For example, some studies on
96 the effects of nitrogen fertilizer indicated a decrease in SOC caused by increased decomposition
97 (Khan et al., 2007; Russell et al., 2009), while others reported an increase in SOC from increased
98 biomass returned to the soil after harvest (Jung and Lal, 2011; Halvorson et al., 1999; Wilts et
99 al., 2004). SOC increases when crop residue is returned to the land (Buyanovsky and Wagner,
100 1998; Wilhelm et al, 2004; van Groenigen et al., 2011), but residue can also increase
101 decomposition in warm, moist areas (Johnson et al., 2005). Perhaps the disagreement is the result
102 of the large variability and uncertainty of field measurements, which make developing
103 conclusions difficult (Karlen et al., 2011). For example, Smith et al. (2012) found no differences
104 between the residue-returned and residue-harvested treatments, and in some cases the residue-
105 harvested sites had increased SOC. Thorburn et al. (2012) also found no consensus regarding
106 residue harvest and SOC response. Nonetheless, most studies found a loss of SOC with residue
107 harvesting. Although the variability of SOC measurements can be attributed to any number of
108 effects — including topography (Senthilkumar et al., 2009b), SOC baseline (Senthilkumar et al.,

109 2009a), aggregate protection (Ananyeva et al., 2013), and even depth (Kravchenko and
110 Robertson, 2011; Syswerda et al., 2011) — it is generally agreed that if crop residue is used as
111 feedstock for biofuels, additional carbon losses can occur (Karlen et al., 2011).

112 SOC losses can be mitigated through recommended management practices, but studies
113 disagree on the limits of harvestable crop residue to maintain SOC levels in soils. Estimates of
114 harvestable non-grain biomass range from 13% (Tan et al., 2012) to 50% (Blanco-Canqui and
115 Lal, 2009a), with an average of about 25%, although that might require stabilization of SOC
116 (Tan et al., 2012). These estimates consider erosion, soil productivity, maintaining SOC, surface
117 crusting, porosity, aggregate breakdown, compaction, and soil temperature, but the wide range in
118 estimated biomass available for harvest leaves questions regarding the sustainability of cellulosic
119 ethanol. However, because the rate of SOC loss tends to increase with increased biomass harvest
120 (Lemke et al., 2010), harvesting small amounts of residue for biofuel might be feasible.

121 Modeling studies can supplement observational data and explore possible differences in
122 SOC by investigating idealized cases. A benefit is that the wide study area can be extended to
123 regional or global scales without resorting to geospatial methods of interpolating sparse data. In
124 this study, we evaluated the influence of cultivation on SOC by using the agriculture version of
125 the Community Land Model (CLM), CLM-Crop (Drewniak et al., 2013). Our analysis includes
126 impacts of changes in land use and also in management practices, such as crop residue harvesting
127 and fertilizer application. A description of the model and the simulations performed is presented
128 in Sect. 2, followed by results and a discussion in Sect. 3 and Sect. 4, respectively.

129 **2 Methods**

130 **2.1 CLM-Crop model description**

131 CLM-Crop, the agriculture version of CLM, includes representations of maize, spring
132 wheat, and soybean crop types with fully coupled carbon-nitrogen cycling (Drewniak et al.,
133 2013). The variation of carbon and nitrogen allocation to plant components with the growth
134 phase of crop development is based on the dynamic vegetation model Agro-IBIS (Kucharik and
135 Brye, 2003). The growth phases are defined as planting, emergence, grain fill, and harvest. Plant
136 date and growth period are determined from the Crop Calendar Dataset (Sacks et al., 2010), and
137 each phase is reached according to a phenological heat unit (PHU) method (see Drewniak et al.,
138 2013).

139 Several processes governing nitrogen cycling are included in CLM-Crop to represent
140 nitrogen retranslocation, fertilization, and nitrogen fixation in soybean. Nitrogen retranslocation
141 occurs during the grain fill growth phase, when nitrogen in the leaves and stem are mobilized to
142 meet organ demands. Fertilizer is applied during the emergence phase for 20 days at constant
143 rates of 150 kg/ha for maize, 80 kg/ha for spring wheat, and 25 kg/ha for soybean. The 20-day
144 fertilization period is designed to optimize nitrogen usage and reduce loss of excess nitrogen
145 through denitrification. Soybean nitrogen fixation allows soybean crops to behave as legumes
146 fixing additional nitrogen through roots — a treatment similar to that of the SWAT model
147 (Neitsch et al., 2005).

148 Harvest occurs as soon as maturity is reached. Grain is removed from the system to
149 represent the consumption of that plant component. The remaining stems and leaves are
150 considered residue and are split into litter and product pools. Litter is returned to the soil through

151 the decomposition process, while product is removed with the grain for uses such as biofuels,
152 animal bedding, etc. The amount of residue returned as litter can be varied for different
153 scenarios. High returns represent sustainable agriculture practices to maintain soil fertility, and
154 low returns are indicative of high cellulosic biofuel usage.

155 **2.2 Input data**

156 CLM-Crop requires two types of input: climate data and surface data. The climate data
157 from the National Center for Environmental Protection reanalysis for 1948-2004 (Kalnay et al.,
158 1996) include temperature, wind speed, humidity, precipitation, solar radiation, and surface
159 pressure at 3-hr intervals. Because the spin-up of the model requires over 600 yr of simulation,
160 we cycled through the reanalysis data to reach a steady state (Thornton and Rosenbloom, 2005).

161 Surface data sets assign the proportion of each land type and plant functional type in a
162 grid cell; crops are grown separately from natural vegetation to eliminate competition for
163 resources. Natural vegetation prescribed from Bonan et al. (2002) includes a generic crop area.
164 Crop distribution for 1992 from Leff et al. (2004) is used to construct maize, wheat, and soybean
165 coverage from the total generic crop area. Because the wheat coverage includes both spring and
166 winter wheat, we model winter wheat as spring wheat in CLM-Crop. Some crop areas
167 overestimated as double cropping in the data set might result in a crop area being counted twice.

168 In addition to land use, the surface data include the planting dates and growth period of
169 each crop type from the Crop Calendar Dataset (Sacks et al., 2010). Planting date is the average
170 day of year when planting occurs, aggregated from 0.5° resolution to 2.8° for CLM-Crop. In
171 regions where data are not available, Sacks et al. (2010) used nearest-neighbor extrapolation to
172 infer planting date. Growth period is calculated in Sacks et al. (2010) as the average number of

173 PHUs between the average planting date and the average harvest date for the 30-yr Climatic
174 Research Unit data set (New et al., 1999).

175 **2.3 Simulations**

176 CLM-Crop was run at a resolution of $2.8^\circ \times 2.8^\circ$ by using the spin-up procedure in
177 Thornton and Rosenbloom (2005). During spin-up, only natural vegetation was active, and
178 croplands were simulated as grass until a steady SOC state was reached. At the end of the spin-
179 up, the land use was converted to include agriculture, representative of the early 1990's land use
180 maps from Leff et al. (2004). CLM does not have a dynamic vegetation capability when crops
181 are active, so land use/land cover is held constant for the remaining simulations. Several case
182 studies were designed and run to evaluate the influence of management practices on SOC (Table
183 1). Each case study was run for a total of 171 years (three complete cycles of the 1948-2004
184 data) at an hourly time step to represent the most intense cultivation period in North America
185 (Ramankutty and Foley, 1999). However, we consider only the last 57 yr of simulation for
186 analysis with averaged data. The control simulation, representing current fertilizer and
187 management practices over North America, is compared to an extension of the spin-up, with
188 crops represented as grass. Additional experiments compared the impact on soil carbon from four
189 agricultural practices (high, medium, and low residue levels and zero fertilizer) with our control
190 simulation.

191 To investigate the effects of land use changes on SOC, different residue management
192 practices, and varied fertilizer application, the results from six scenarios were analyzed (Table 1).
193 First, conventional crop management (control run, 70% residue) is compared with crops
194 simulated as grass (grass run). Second, effects of high (90%), medium (30-40%), and low (10%)

195 residue are compared with values for the control run. Third, the effect of no fertilizer application
196 (with 70% residue) is evaluated by comparison with the control run.

197 **3 Results**

198 **3.1 Soil organic carbon**

199 Simulated SOC values from the control run range from $< 2 \text{ kg C m}^{-2}$ in the Southwest to
200 $> 20 \text{ kg C m}^{-2}$ in the northern United States (Fig. 1). Average SOC values are lower in crop
201 ecosystems than in natural vegetation systems because of biomass removal and other land
202 management. The total stored SOC over all land surface types in the United States, as calculated
203 by CLM-Crop, is 84 Pg C, which falls within the range of previous estimates of 78-85 Pg C
204 (Kern, 1994). CLM-Crop-simulated SOC for agriculture sites over the contiguous United States
205 (CONUS) has a pattern similar to that of total SOC, with higher SOC in the northern part of the
206 country and lower SOC in the southern regions.

207 The general spatial pattern of the model-calculated SOC over CONUS is evaluated by
208 using available spatially gridded data sets of SOC. The data developed by the global soil carbon
209 International Geosphere-Biosphere Program (IGBP; Global Soil Data Task Group, 2000) for
210 CONUS are summarized in Fig. 1b. The SOC pattern and magnitude are similar to the model-
211 calculated values (Fig. 1a). The differences between the model-calculated SOC and the IGBP
212 data set are shown in Fig. 1c. In most regions, the percent difference between the data set and the
213 model simulation is $< 5\%$. Areas with higher percent differences are in boreal regions, where
214 CLM tends to underestimate soil carbon (Koven et al., 2013).

215 Figure 1 includes both managed and natural lands. To evaluate the model-simulated SOC
216 over agricultural lands, we selected self-identified measurements of SOC from agricultural lands

217 available from the International Soil Carbon Network (ISCN; 2014). This data set has over 4,000
218 unique SOC measurements to 1-m depth from croplands over CONUS. Although CLM soil
219 depth (3.8 m) is deeper than the observations (1 m), since nearly two-thirds of SOC is found
220 within the top 1 m (Jobbagy and Jackson, 2000), the bulk of the soil carbon is still captured in
221 the observations. Because the ISCN data were collected over a wide variety of soils, at different
222 points in the crop cycle and different times since the change in land used, variability is large, and
223 the number of outliers from the median of the sample is significant. The plot in Fig. 2 shows the
224 range of values with significant occurrences in the upper quartile and above the 90th percentile
225 of the distribution. We filtered out outliers with SOC measurements $> 50 \text{ kg C m}^{-2}$ in this figure
226 only to improve readability of the graph, since only a small portion (2.5%) of the measured
227 values were higher than 50 kg C m^{-2} and SOC in agriculture lands is typically less than 50 kg C
228 (Kern et al., 1994; Mishra et al., 2010). The model results for the grid cells identified as cropland
229 are included in Fig. 2. The model results have a smaller range than the ISCN data, as would be
230 expected for SOC values extracted at the end of the simulation period and post-harvest. In
231 addition, the SOC in the model is less variable because of the larger grid cells with uniform soil
232 type. Nevertheless, the median SOC values simulated by CLM-Crop fall within range of the
233 middle 50% of the ISCN measurements (Fig. 2), and thus the simulated values are comparable,
234 on average, with the observations. In order to compensate for the mismatch of soil depth, we
235 added an additional 36% of SOC to the observed stocks (to account for the $\sim 1/3$ carbon between
236 2-3 m soil depth; Jobbagy and Jackson, 2000). The resulting increase in observed SOC (not
237 shown) caused median CLM-Crop SOC stocks to fall outside the 50 percentile of the
238 observations, but the top 75 percentile of CLM SOC still fall within observed range.

239 In a further evaluation of the model's performance over agricultural lands, we completed
240 a site-by-site comparison of modeled SOC to observed SOC. We applied a filter to separate soil
241 over the modeling domain into three types (clay, sand, and silt), to examine the model behavior
242 against the different textures. Figure 3 plots simulation results versus observations of SOC for
243 values selected as described above. Each point indicates the mean observational SOC stock at the
244 model grid scale with the standard deviation. The plot indicates that although the model does
245 tend to underestimate soil carbon over croplands, CLM does reasonably well at catching a wide
246 range of SOC values at agricultural sites for all soil textures. The model does not capture the
247 individual site observations well ($RSME = 13.1 \text{ kg C m}^{-2}$; $R^2 = 0.016$), due to the high spatial
248 variability. CLM tends to simulate high SOC in sandy soils, low SOC for silt soils, and clay SOC
249 in between, however the soil texture is determined from the model data and therefore may not
250 accurately represent the soil texture of the observations. This result is encouraging, in view of
251 difficulties in comparing CLM-Crop-simulated SOC with observations at agriculture sites. First,
252 the large grid size used in the model simulation cannot resolve the small-scale variability
253 between farm-scale measurements, which are apparent from the large standard deviation in
254 observations. Second, the model is run with static management for long time periods and cannot
255 capture changes in management or land use over long temporal and large spatial resolutions
256 while observations are taken over various time frames with vastly different land use history.
257 Finally, measurements are 1 m depth, and CLM-Crop estimates SOC for the total soil column (>
258 300 cm). When we attempt to adjust the observed SOC to include carbon at deeper soil layers
259 (by adding $\sim 1/3$ more carbon as in Fig. 2), RSME increases to 18.8 kg C m^{-2} , although R^2 did not
260 change. Despite these challenges, CLM can capture the range of SOC present at many

261 agriculture sites and in many cases CLM SOC estimates fall within the standard deviation of the
262 observations.

263 In order to explore the model performance further, we examined the effect of climate
264 variability on SOC stocks. CLM SOC stocks decrease with increasing mean annual temperature
265 and total annual precipitation (Fig. 4), which is also supported by observations. Higher
266 temperatures and soil moisture generally result in higher below ground activity and therefore
267 faster turnover of soil carbon (Wei et al., 2014). Natural vegetation follows the same temperature
268 trends, but regions with higher annual precipitation indicate higher SOC stock. This is possibly
269 the result of increased productivity when precipitation is high, however the variability in natural
270 vegetation is quite high making conclusions difficult.

271 Finally, we also consider the ability of the model to capture temporal changes in SOC
272 from land use conversion. Percent SOC loss since conversion from forest to agriculture, as
273 summarized in Wei et al. (2014), is plotted in Fig. 5 over temporal periods ranging from 1-207
274 years with a subset (500 points) of CLM SOC percent loss taken from random grids and time
275 periods. Although CLM does not simulate the rapid loss of SOC that occurs in some field
276 observations, by the end of the simulation, CLM does capture the range of SOC loss as seen in
277 observations. Initial lower SOC stocks likely cause the initial modest decline in SOC simulated
278 by the model, since SOC loss increases with increasing initial SOC concentration (Wei et al.,
279 2014). This result highlights CLMs ability to capture changes in SOC over long time periods.

280 **3.2 CLM-Crop-simulated changes in soil carbon**

281 Most grid cells lost between 3% and 45% of total SOC, averaged across the grid cell. The
282 amount of SOC lost was correlated with the size of the agriculture land base; higher agriculture

283 land use resulted in larger SOC loss. Individual crop soil columns indicate high losses of SOC,
284 up to a maximum of 75% of total SOC, although average soil loss is 33-51%. Total loss also
285 varied with crop type; maize and wheat lost about 10% less SOC than soybean. This is
286 understandable, given the low residue of soybean crops, although this result varied with location.
287 For example, total simulated SOC loss over maize and soybean soil columns at the Bondville site
288 in Illinois was 48%. At the Mead, Nebraska, site, losses of SOC for maize and soybean columns
289 were approximately 44% and 52%, respectively.

290 While these site-level SOC losses are comparable with observations (Lal, 2004),
291 comparison with the SOC values in the control simulation might be exaggerated as a result of the
292 subgrid hierarchy, because the accumulated SOC estimated by the grass simulation was
293 influenced by all vegetation types in the soil column, while the soil column in the control
294 simulation only included one crop type. In addition, Ramankutty and Foley (1999) showed that
295 most early croplands from the late 1800s were formed through deforestation and later prairie
296 removal. This implies that our estimation might be exaggerated, because grassland ecosystems
297 can hold more carbon than forests (Schlesinger, 1997). Overall, a 10% loss in total SOC over the
298 United States between the control run and the grass run accounts for a nationwide carbon loss of
299 more than 8 Pg (Fig. 6).

300 Residue management can have the largest impact on soil carbon. Increasing the residue
301 left on the field to 90% results in a 2.6% increase of SOC, but allowing a 10% residue amount
302 (as a potential result of increased cellulosic biofuel demand) leaves an SOC decrease of over
303 5.7%. The difference between these two scenarios is over 7 Pg C, almost the same amount as the
304 total carbon loss due to agricultural land use. Interestingly, we found no notable differences
305 between crop responses. Even a more modest decrease in the residue returned to the field (30-

306 40%) results in a 3.5% loss of SOC compared to the control simulation. Increasing the residue
307 harvest will increase the amount of SOC loss (Anderson-Teixeira et al., 2009; Blanco-Canqui
308 and Lal, 2009b). Harvesting residue results in the loss of not only soil carbon, but also soil
309 fertility, indicated by declining yields (data not shown). This implies that increased residue
310 harvest for cellulose might result in expansion of croplands to counter yield declines.

311 Eliminating fertilizer use showed the biggest impact on yields and SOC, simulating over
312 6% loss (Fig. 6). Globally, decreases in yields of roughly 60-70% occurred for maize and wheat,
313 but soybeans, relying less on fertilizer inputs, suffered a 22% decrease in yields. The different
314 response between plant types was large: individual maize and wheat soil columns lost an average
315 of 63% SOC, whereas soybean only lost 11%. Despite low yields, leaving 70% residue allowed
316 carbon inputs to maintain nearly the same SOC level as in the run with low residue return. This
317 indicates a critical role for fertilization in soil carbon storage, without which an additional 5 Pg C
318 might be lost due to cultivation. The observed result is not surprising, as fertilizer contributes to
319 the total biomass accumulated during crop development, and increased biomass returned as
320 residue will allow the soil to retain some of the nutrients taken up during crop growth, improving
321 the soil fertility.

322 **4. Discussion**

323 CLM-Crop has proven to be a valuable tool for evaluating changes in soil carbon under
324 various management practices. Our results indicate that the SOC for agricultural sites will be
325 reduced through any management practice while disturbance continues, with the total amount
326 lost depending on the management practice. Model-estimated U.S. losses of SOC due to current

327 cultivation practices are around 10%, with a potential for greater loss as the amount of harvested
328 residue increases.

329 The amount of biomass residue left on the field after grain harvest has the most
330 significant effect on SOC. Cellulosic biofuels rely on harvesting the stems and leaves of crops,
331 resulting in an additional 5% loss of carbon within the soil system. Currently, model subgrids
332 growing a single crop type on an independent soil column typically lose 33-51% of SOC, and
333 that loss increases to nearly 90% when residue is harvested. Over long time scales, this effect can
334 degrade the sustainability of the soil for crop growth and can negatively affect yield. For
335 example, plant nitrogen uptake (Fig. 7) decreased linearly with increasing residue harvest. The
336 high residue returns uptake 7.4% more N than the current residue runs, whereas medium and low
337 residue returns have 6.6% and 15.6% lower N uptake, respectively. When fertilizer is not
338 included, the resulting N uptake is 57% lower. This impact is transferred to yields (Fig. 8)
339 resulting in 9% and 17% lower yields for the medium and low residue returns, respectively.
340 Thus, the effects of residue management on SOC are very important, and increasing the amount
341 of residue used for cellulosic ethanol production could have a significant impact on soil carbon
342 storage and ultimately plant productivity. Leaving plant residue from crop production in the soil
343 decreases the amount of carbon lost to the atmosphere. However, meeting cellulosic biofuel
344 demand through cultivation of managed grasses such as switchgrass and Miscanthus has been
345 shown to increase soil carbon storage over time (Anderson-Teixeira et al., 2009), most likely
346 because nutrient demands and management practices are different for these types of biofuel
347 crops.

348 Disagreement between studies about the possible effect of fertilizer on SOC leaves this
349 management practice open for further research. Our findings suggest that fertilizer use might

350 improve yield and increase the amount of carbon returned to the soil in crop residue; however,
351 increased residue removal for biofuels could reduce this effect. As fertilizers improve and are
352 applied to maximize plant uptake while minimizing loss to leaching and denitrification, fertilizer
353 might provide an important tool for farmers to mitigate the soil carbon loss due to increasing
354 residue harvest for biofuel use. However, care must be taken to ensure that fertilizer inputs do
355 not exceed plant uptake, which could result in increased nitrogen leached into the groundwater
356 and increased greenhouse gas emission of N_2O via nitrification and denitrification pathways. The
357 effect of increased decomposition when fertilizer is used also needs to be explored.

358 Expanding the model to incorporate other management practices (rotation, tillage,
359 irrigation, etc.) is important activity for future model development. Erosion, for example, is
360 expected to increase as a result of crop residue harvest (Lal and Pimentel, 2007). This secondary
361 effect of residue harvest can have multiple consequences. First, soil fertility will decline with the
362 loss or transport of soil organic matter. Second, erosion processes result in the breakdown of soil
363 aggregates promoting oxidation of SOC. Both effects will reduce nutrient and water holding
364 capacities of the soil (Lal and Pimentel 2008). Finally, the loss of nutrients will result in a
365 decline of crop productivity, further enhancing SOC loss. As such, our result should be
366 considered a lower bound estimate of SOC loss from residue harvest. Including these effects and
367 expanding agricultural models to a global scale should be a priority for future model
368 development. Given the challenges comparing with observations, focusing on model
369 developments that capture cropland SOC dynamics is equally important as developing datasets
370 that can be used for climate model validation, especially considering the increasing complexity
371 of ESMs that include cropland representation. Although the crop representation in CLM-Crop is

372 flexible enough for expansion to a global scale, rigorous testing is needed to ensure that crop
373 behavior is consistent with regional observations.

374 There are some limitations to our modeling approach that lead to uncertainties in the
375 model prediction of SOC. For example, changes in land use and land cover are not included in
376 CLM. Historical changes in land use indicate a steady increase in cultivated land which peaked
377 in the 1940's and declined thereafter (Waisanen and Bliss, 2002). Using a modern land use cover
378 over the historical period may result in an over prediction of SOC loss, because the model will
379 overestimate the agricultural land base in some (early) years and the model won't capture
380 increases in SOC when agriculture land is abandoned. This also limits the influence of beneficial
381 agriculture practices such as crop rotation and fallowing. Historical changes in land management
382 are also not represented in the model, such as changes in residue harvest over time or organic
383 matter additions. For example, Lal et al. (1999) suggest early cultivation removed residue
384 following harvest until after 1940 when residue was returned to the field. The high spatial
385 variability and difficulty finding these types of historical data is a major challenge for trying to
386 add these features to CLM.

387 Finally, further research is needed for full evaluation of the importance of agro-
388 ecosystem impacts on soil carbon. We have shown here that SOC loss can vary greatly,
389 depending on management practices. Practices such as residue management can have significant
390 impact on SOC retained in agricultural soils, with higher residue removal from soil leading to
391 higher SOC losses. Use of fertilizer can compensate for some of the loss, but the benefit is
392 limited. Further modeling studies are important for simulating these competing effects on carbon
393 storage. Our study suggests that considerable care is needed in designing appropriate

394 management practices to realize the full carbon mitigation benefits of using biofuels from
395 cellulosic ethanol.

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583 Figure 1. (a) Total SOC (kg C m^{-2}) simulated by CLM-Crop over the contiguous United States.
584 (b) Total SOC from the IGBP over the same domain as in (a). (c) Percent difference between (a)
585 and (b).

586 Figure 2. Box plot of the weighted average total SOC over croplands, as simulated in CLM-Crop
587 and in observations from the ISCN. Observations reporting $> 50 \text{ kg C m}^{-2}$ were removed from
588 the analysis.

589 Figure 3. CLM-modeled SOC (kg C m^{-2}) versus ISCN observations for model derived soil
590 texture types clay, sand, and silt. Each point represents the mean observed SOC value in the grid
591 cell; error bars show the standard deviation. The black line represents the 1:1 ratio.

592 Figure 4. Top: The effects of temperature on SOC stock from CLM crops (blue) and natural
593 vegetation (green) and ISCN observations (red). Bottom: The effects of precipitation on SOC
594 stock from CLM crops (blue) and natural vegetation (green) and ISCN observations (red).

595 Figure 5. Percent decrease of SOC after conversion from natural vegetation to cropland. Percent
596 decrease data from Wei et al. (2014) are in red (US points are orange) and CLM percent loss is
597 blue.

598 Figure 6: Simulated change in total U.S. SOC (Pg C) due to agricultural land management for all
599 scenarios.

600 Figure 7. The effect of agricultural land management change on crop annual average nitrogen
601 uptake.

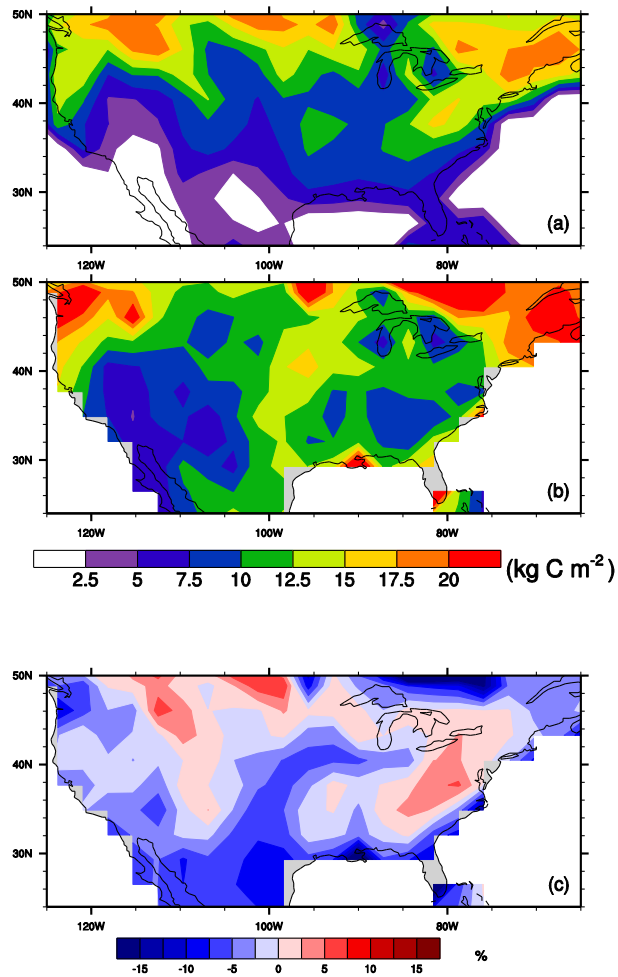
602 Figure 8. The effect of agricultural land management change on annual crop yield.

603 Table 1. CLM-Crop simulations performed.

Run name	Land use	Fertilizer	Residue
Control	Leff et al., 2004	Yes	70% — all crops
High residue	Leff et al., 2004	Yes	90% — all crops
Medium residue	Leff et al., 2004	Yes	30% — maize 30% — wheat 40% — soybean
Low residue	Leff et al., 2004	Yes	10% — all crops
No fertilizer	Leff et al., 2004	No	70% — all crops
Grass	Bonan et al., 2002	Not applicable	Not applicable

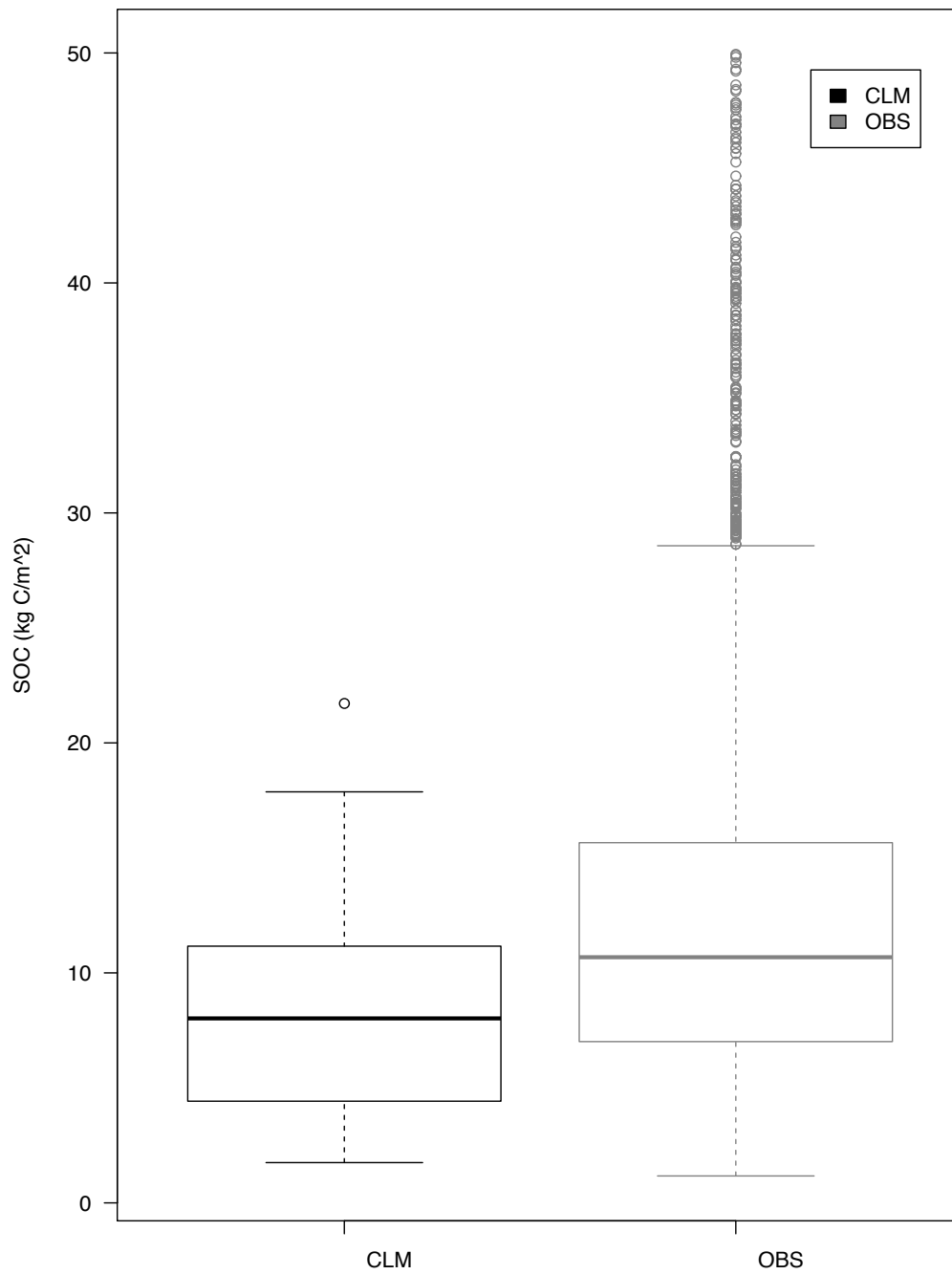
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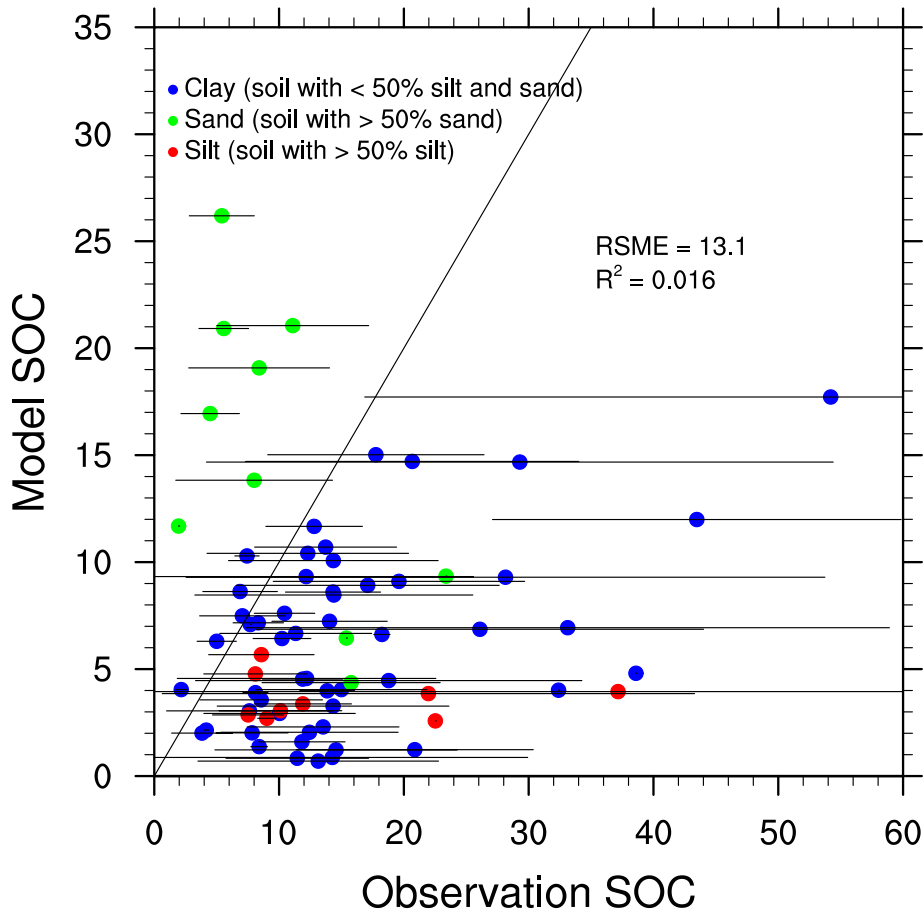
607 Figure 1. (a) Total SOC (kg C m⁻²) simulated by CLM-Crop over the contiguous United States. (b) Total
 608 SOC from the IGBP over the same domain as in (a). (c) Percent difference between (a) and (b).



609

610 Figure 2. Box plot of the weighted average total SOC over croplands, as simulated in CLM-Crop and in
 611 observations from the ISCN. Observations reporting > 50 kg C m⁻² were removed from the analysis.

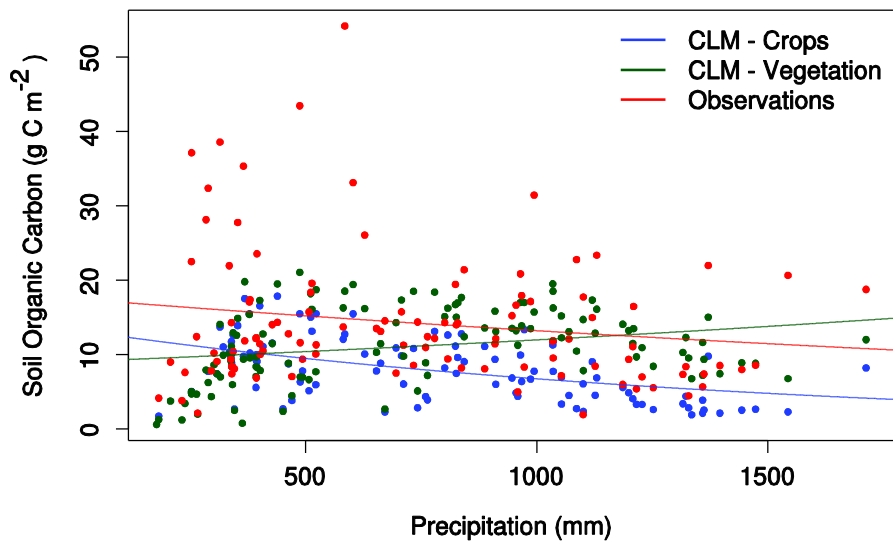
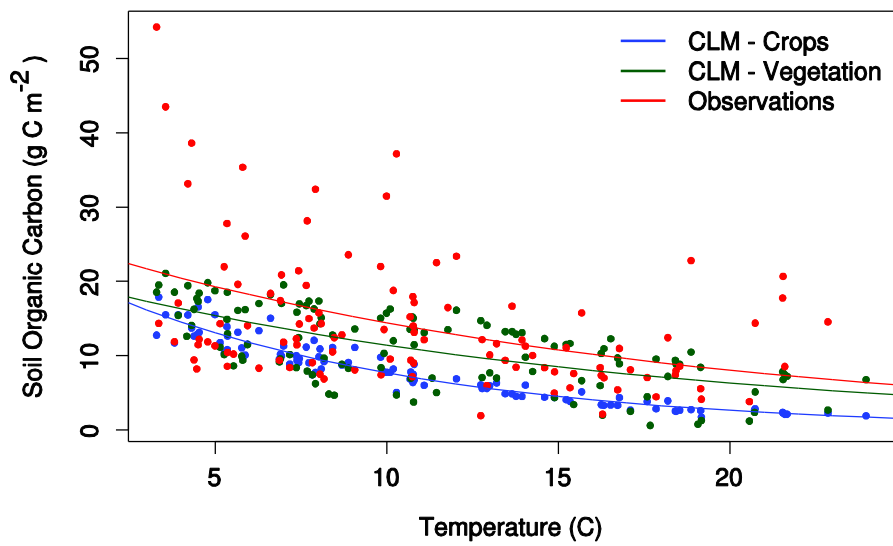
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613

614 Figure 3. CLM-modeled SOC (kg C m^{-2}) versus ISCN observations for model derived soil
 615 texture types clay, sand, and silt. Each point represents the mean observed SOC value in the grid
 616 cell; error bars show the standard deviation. The black line represents the 1:1 ratio.

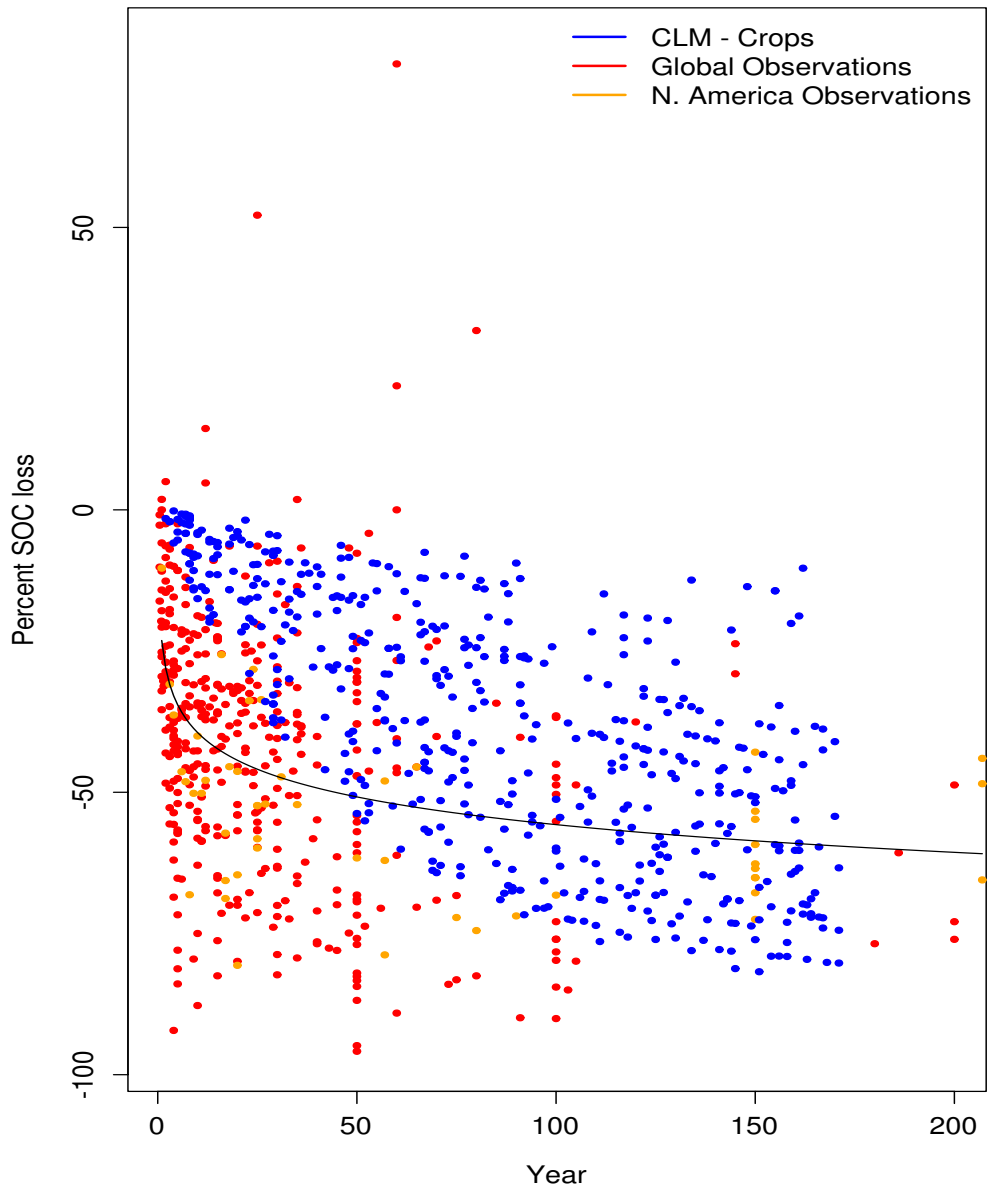
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619 Figure 4. The effects of temperature (top) and precipitation (bottom) on SOC stock from CLM crops
 620 (blue) and natural vegetation (green) and ISCN observations (red).

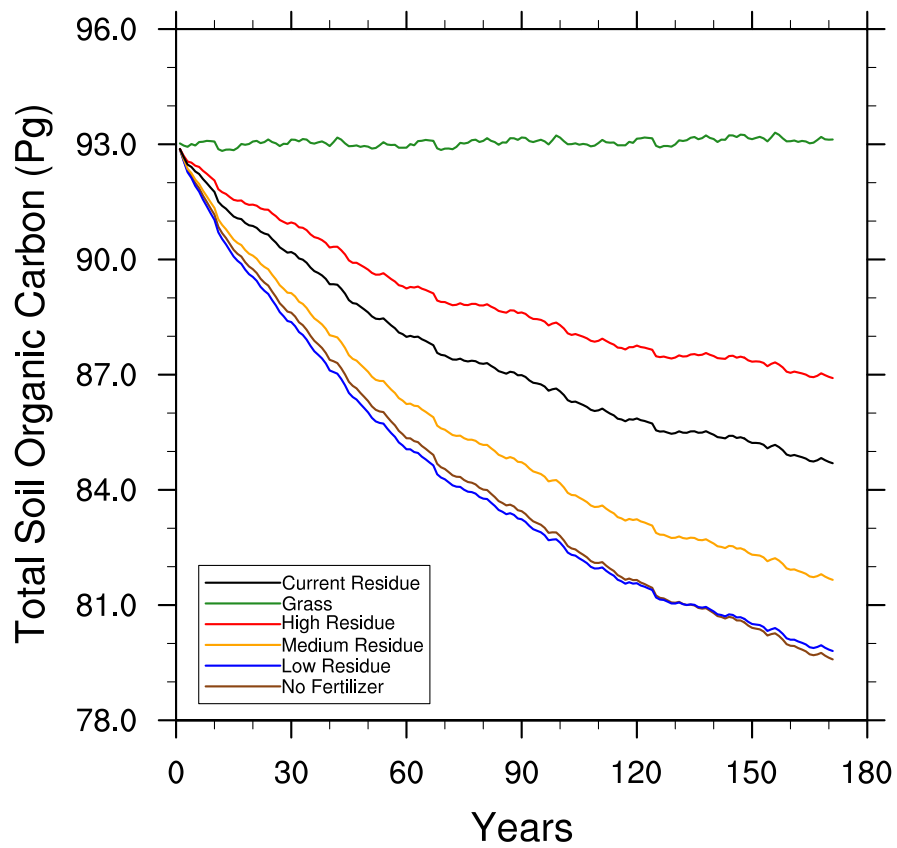
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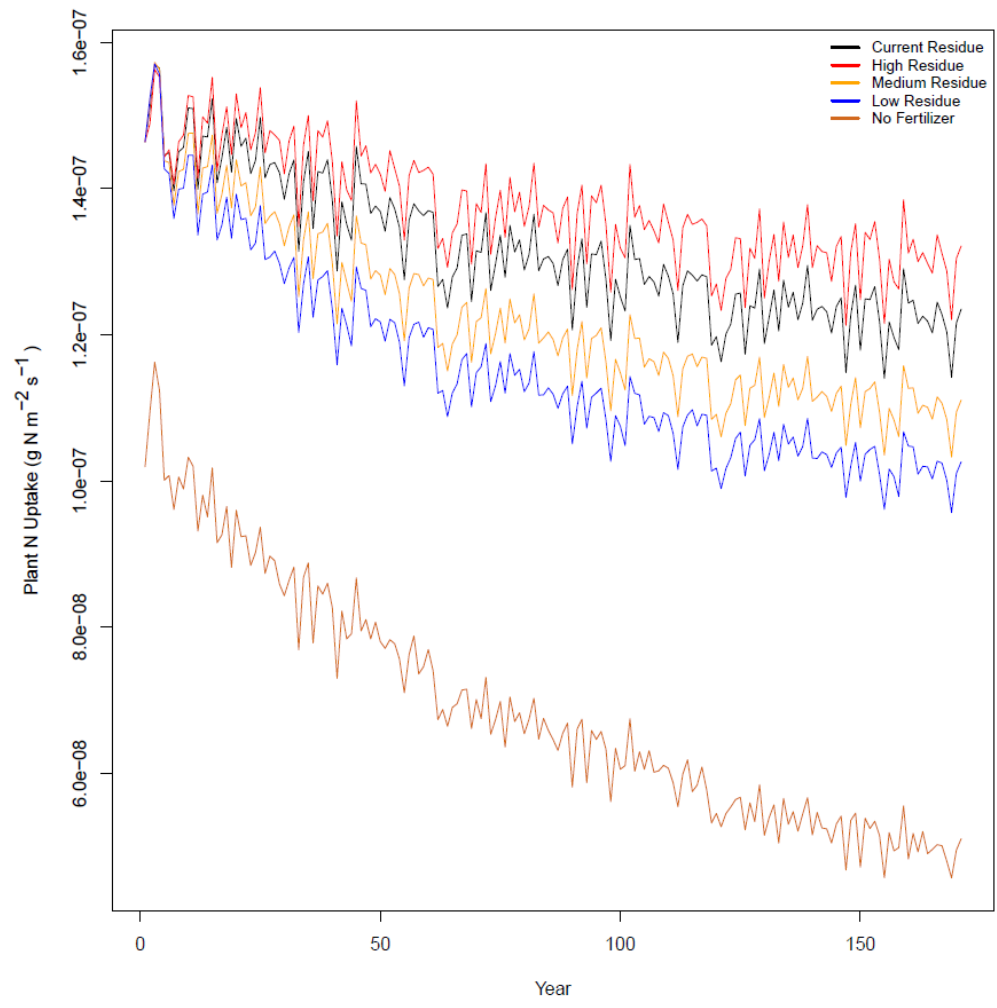
623 Figure 5. Percent decrease of SOC after conversion from natural vegetation to cropland. Percent decrease
 624 data from Wei et al. (2014) are in red (US points are orange) and CLM percent loss is blue.

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627 Figure 6: Simulated change in total U.S. SOC (Pg C) due to agricultural land management for all
 628 scenarios.

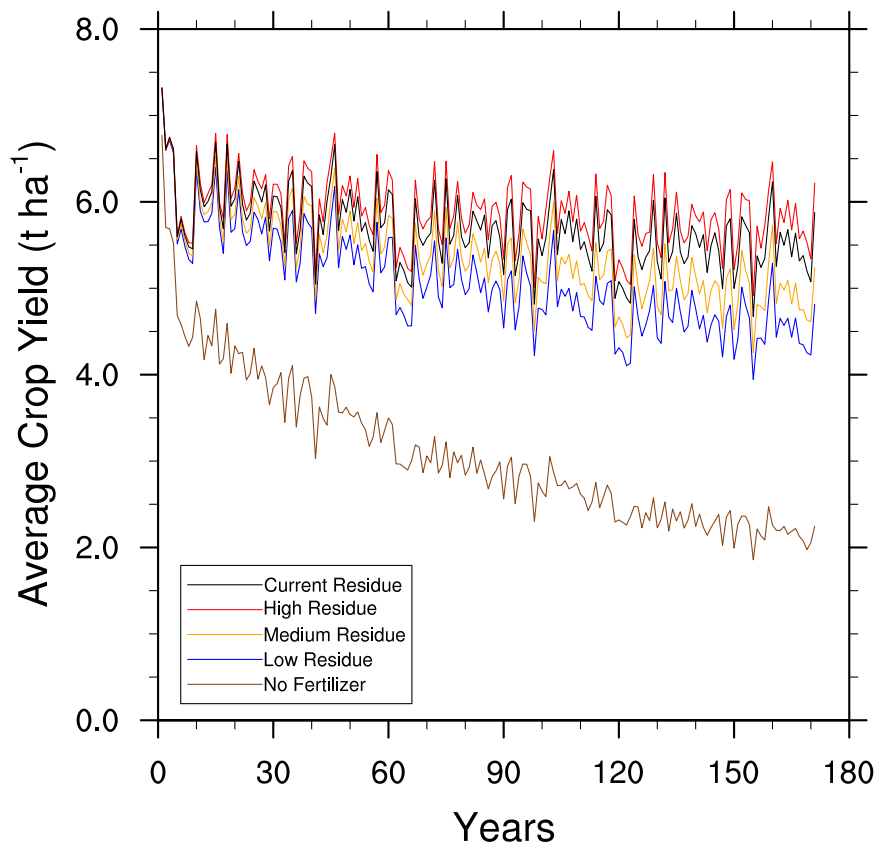


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630

631 Figure 7. The effect of agricultural land management change on crop annual average nitrogen uptake.

632



633

634 Figure 8. The effect of agricultural land management change on annual crop yield.

635