

#### Abstract

1

2 Concerning their sensitivity to ocean acidification, coccolithophores, a group of calcifying single-celled phytoplankton, are one of the best-studied groups of marine organisms. 3 4 However, in spite of the large number of studies investigating coccolithophore physiological responses to ocean acidification, uncertainties still remain due to variable and partly 5 6 contradictory results. In the present study we have used all existing data in a meta-analysis to 7 estimate the effect size of future pCO<sub>2</sub> changes on the rates of calcification and 8 photosynthesis and the ratio of particulate inorganic to organic carbon (PIC/POC) in different 9 coccolithophore species. Our results indicate that ocean acidification has a negative effect on 10 calcification and the cellular PIC/POC ratio in the most abundant coccolithophore species Emiliania huxleyi and Gephyrocapsa oceanica. In contrast the more heavily calcified species 11 Coccolithus braarudii did not show a distinct response when exposed to elevated 12 13 pCO<sub>2</sub>/reduced pH. Photosynthesis in Gephyrocapsa oceanica was positively affected by high CO<sub>2</sub>, while no effect was observed for the other coccolithophore species. There was no 14 15 indication that the method of carbonate chemistry manipulation was responsible for the 16 inconsistent results regarding observed responses in calcification and the PIC/POC ratio. The perturbation method, however, appears to affect photosynthesis, as responses varied 17 18 significantly between total alkalinity (TA) and dissolved inorganic carbon (DIC) 19 manipulations. These results emphasize that coccolithophore species respond differently to 20 ocean acidification, both in terms of calcification and photosynthesis. Where negative effects 21 occur, they become evident at CO<sub>2</sub> levels in the range projected for this century in case of 22 unabated CO<sub>2</sub> emissions. As the data sets used in this meta-analysis do not account for 23 adaptive responses, ecological fitness and ecosystem interactions, the questions remains how 24 these physiological responses play out in the natural environment.

## 1. Introduction

Coccolithophores, a globally distributed group of marine haptophytes, are major primary producers in the ocean and the most prolific calcifying organisms on our planet (Brownlee & Taylor, 2004; Shutler et al., 2010). By performing photosynthesis and calcification, they contribute to both biological carbon pumps, the soft tissue pump and the carbonate counter pump. While the former supports carbon sequestration in the ocean through production and sinking of organic matter to depth, the latter decreases the ocean's capacity to take up CO<sub>2</sub> due to the reduction of surface layer alkalinity. Moreover, by providing ballast material, which accelerates sinking velocities of organic particles to depth, coccolithophore-derived calcite contributes to enhancing carbon sequestration to depth (Klaas & Archer, 2002; Armstrong et al., 2002; Ziveri et al., 2007). Thus, changes in the contribution of coccolithophores to ocean primary production could potentially impact global carbon cycling (Riebesell et al., 2009).

In the face of global change phytoplankton are subjected to rapid alterations in their environmental conditions. Due to the sensitivity of calcification to ocean acidification, coccolithophores are considered to be among those, which may be adversely affected in a high  $CO_2$  future ocean. While impacts of ocean acidification on coccolithophores have been studied extensively (for review see e.g. Riebesell & Tortell, 2011), variable and partly conflicting responses were observed in different perturbation studies (for a summary see Tables 1 and 2). Differences in experimental conditions, such as in light intensity, temperature, salinity, nutrient concentration and  $pCO_2$  levels have been attributed as possible causes for those variations. But even studies with comparable experimental conditions provided deviating responses of coccolithophores. Some of this divergence was shown to be related to species- and strain-specific differences (Langer et al., 2006, 2009). But also the

method of carbonate chemistry manipulation, whether through changes in total alkalinity (TA) or dissolved inorganic carbon (DIC), was discussed as possible cause for some of the observed discrepancies (Iglesias-Rodriguez et al., 2008; Shi et al., 2009).

54

55

56

57

58

59

60

61

62

63

64

65

66

67

68

69

70

71

72

51

52

53

Building on the extensive literature on coccolithophore responses to ocean acidification, the present study aims to provide statistically and methodologically robust estimates for those responses. In particular, we intend to answer the question whether increasing seawater acidity alters calcification, photosynthesis and the PIC to POC ratio in acclimated cultures of coccolithophores. We further assess whether the observed responses are affected by the carbonate chemistry manipulation method and if they differ between coccolithophore species, thus trying to address some of the inconsistencies in the existing studies. Recent metaanalyses conducted by Kroeker et al. (2010, 2013) and Hendriks et al. (2010) did not specifically focus on coccolithophores but analyzed responses of many different taxa to ocean acidification. Although coccolithophores were included in those meta-analyses, only a few experiments (Kroeker et al., 2010: 13 experiments, Hendriks et al., 2010: 2 experiments for calcification responses, 12 experiments for photosynthetic responses, Kroeker et al., 2013: 19 experiments) were considered and no distinction was made between different coccolithophore species. The meta-analysis by Findlay et al. (2011) focused on *Emiliania*, huxleyi, but only investigated the species' PIC/POC response to ocean acidification (15 experiments were included in the analysis). In our approach a larger set of experiments and response variables was analyzed, allowing for a more robust prediction of the impact of ocean acidification and related changes in seawater chemistry on coccolithophore physiological performance.

73

74

#### 2. Materials and Methods

#### 2.1 Literature search

78 A literature search was conducted to assemble all published data sets on CO<sub>2</sub>/pH sensitivities 79 of coccolithophore calcification and photosynthesis. As a first step the ISI database Web of Science (www.webofknowledge.com) was scanned for appropriate studies. Additional 80 81 literature was obtained from the EPOCA (European Project on OCean Acidification) database 82 (www.epoca-project.eu) and from the associated blog 83 (www.oceanacidification.wordpress.com). Subsequently, the reference lists of all studies identified by this approach were scanned for other relevant literature. 84 85 Experimental data were extracted directly from the published papers or, if not reported therein, from the PANGEA® archive (www.pangaea.de). If the information could not be 86

retrieved from either source, the first author of the study was contacted directly.

88

89

90

91

92

93

94

95

96

97

98

99

87

76

77

#### 2.2 Data selection

All studies in which the carbonate system was altered and the effect on coccolithophores reported, comprising both laboratory and field experiments, were selected for this meta-analysis. Studies that varied other environmental factors in addition to seawater carbonate chemistry, such as light intensity, day length, temperature or nutrient availability, were also incorporated. Data of particulate inorganic (PIC) and organic carbon (POC) production rates, pH values, carbonate system parameters and experimental conditions (light level, day length, temperature, nutrients) were obtained for the control (ambient or pre-industrial  $pCO_2$  level) and the experimental treatments (elevated  $pCO_2$  level). If PIC and POC were provided as quota values on a per-cell basis, production rates were calculated by multiplying the growth rates ( $\mu$ ) with the cell quota of organic or inorganic carbon.

- The following  $pCO_2$  levels were chosen to compare the responses of *Emiliania huxleyi* to pre-
- industrial carbon dioxide concentrations of ~ 280 parts per million (ppm):
- 102 (1)  $\sim$  380 ppm reflecting the present day  $pCO_2$  level,
- 103 (2)  $\sim$  780 ppm the  $pCO_2$  level projected for the end of this century under the SRES A1B
- scenario, IPCC Report 2000 (Nakicenovic et al., 2000), and
- 105 (3)  $\sim 1000$  ppm the pCO<sub>2</sub> level projected for the end of the century under the 'worst case'
- emission scenario A1FI, IPCC Report 2000 (Nakicenovic et al., 2000).
- Since there was not a sufficient number of studies investigating the responses of *Coccolithus*
- braarudii and Gephyrocapsa oceanica at pCO<sub>2</sub> levels around 780 ppm, only concentrations
- 109 of  $\sim 380$  ppm and  $\sim 1000$  ppm were used to compare the responses of these species. All
- experiments where the  $pCO_2$  levels deviated no more than  $\pm$  50 ppm from the targeted 380
- ppm and no more than  $\pm 100$  ppm from the targeted 780 ppm and 1000 ppm were included in
- the analysis. Since the studies by Lefebvre et al. (2012) and Jones et al. (2013) did not meet
- these specifications, they were excluded from the meta-analysis.

114

116

117

118

119

120

121

122

Manipulation of the seawater carbonate chemistry can be achieved in various ways. First, the

carbonate system can be adjusted by bubbling with CO<sub>2</sub>. This approach increases [CO<sub>2</sub>],

[HCO<sub>3</sub>] and DIC, decreases pH and [CO<sub>3</sub><sup>2</sup>] and does not change the alkalinity. Second, acid

can be added, which increases  $[CO_2]$  and  $[HCO_3]$ , decreases the alkalinity and  $[CO_3^2]$  and

does not change DIC. In both manipulations the saturation state  $(\Omega)$  decreases as well.

Although there are other ways to adjust the carbonate system, the above-mentioned methods

are the ones most commonly used. It was noted which manipulation method was applied to

decrease the pH in each study. Subsequently, a separate meta-analysis was conducted in order

to analyse whether responses of coccolithophores varied between the methods. Here, only

responses to a pCO<sub>2</sub> elevation from pre-industrial levels to 780 ppm and 1000 ppm were included in the analysis. On this basis 22 experiments were excluded.

When studies reported results from multiple carbonate system perturbation experiments, all

individual experiments were included in the analysis. The same applied when there were

different experiments with various species or strains.

If not only the carbonate system, but also other factors such as light intensity or day length were changed in a study, the approach of Kroeker et al. (2010) was adopted and the ambient level of the factor, defined by the authors of the primary study, was used to ensure the comparability between the experiments. If the observed responses of a study did not differ significantly for the ambient and non-ambient levels of a given environmental factor (always regarding the same  $pCO_2$  value), both experiments were included.

The data on PIC and POC production obtained by Iglesias-Rodriguez et al. (2008) were normalized to POC biomass, following the approach suggested by Riebesell et al. (2008).

Data shown in Table 1 represent the original measurements reported by Iglesias-Rodriguez et

al. (2008) prior to normalization. Müller et al. (2010) did not report PIC and POC production

rates in their study, since the sampling time for those data varied and created a bias in the

data. By averaging the PIC and POC production rates over time, the bias was minimized and

the data were suitable to be included in this meta-analysis.

## 2.3 Data analysis

Determining differences between the control and treatment groups in response to changes in carbonate chemistry was the first step in our analysis. For this purpose the logarithmically transformed response ratio (L) was calculated for each experiment and response variable (PIC, POC and PIC/POC) as:

$$L = \ln(RR) = \ln(\bar{X}_E) - \ln(\bar{X}_C) \tag{1}$$

150

151

152

153

154

155

156

157

158

159

160

161

162

163

164

165

166

167

168

169

170

171

172

173

where  $\bar{X}$  is the mean of a treatment (E) and a control (C) group. The response ratio is logarithmically transformed and unit-less, thus allowing a comparison of data between experiments, which report responses in different units. The effect size is an easy measure of relative change between the control and the treatment group. When L < 0, the effect of acidification in the treatment group is negative and when L > 0, the effect is positive. A response ratio of zero indicates that there is no effect and that the responses in the control and treatment group are the same. Since not all studies are equally precise, meaning that they are based on different numbers of replicates and variable standard deviations, the simple computation of the mean effect sizes is not to be recommended. Instead, a weighted mean is computed where more precise studies are given more weight. This meta-analysis of the response ratios follows the approach of Hedges et al. (1999) with a few variations when weighting the effect sizes. A random effects model was used where the assumption is made that the effect of ocean acidification varies between studies (Borenstein et al., 2010). For example, the effect size might differ between strains or it might turn out significant if the response was measured more reliably or if the incubation time was longer. The random effects model accounts for this variation and includes the within-study variance  $(v_i)$  as well as the between study variance  $(\sigma^2)$  when calculating the mean effect for the response variables. Statistical significance for all effect sizes is displayed by the 95% confidence interval. The effect size is considered to be significant ( $\alpha = 0.05$ ), when the confidence intervals do not overlap zero. Traditionally, when studies report means, standard deviation, and sample size for both the control and treatment groups, a weighted meta-analysis is possible and the variance  $(v_i)$  within the experiment i can be calculated. Consequently, studies with a higher number of replicates

and lower variance are weighted more heavily, which results in a more robust meta-analysis where the estimate of the effect size is more precise than in unweighted meta-analyses (Hedges & Olkin, 1985). Some of the data required for a weighted meta-analysis, however, were not available for some studies. In those cases  $v_i$  was estimated as the average of the computed variances from those experiments where  $v_i$  was calculable. In this way it was possible to include all studies in the meta-analysis. Using the variance  $v_i$  and the mean of the response ratio  $L_i$  for each experiment i, Cochran's Q (Cochran, 1954) was computed. With the help of Q an estimate of the between experiment variance ( $\sigma^2_{\lambda}$ ) was obtained (Hedges et al., 1999). The weighted mean of the log response ratio  $\overline{L^*}$  is given by:

183 
$$\overline{L}^* = \frac{\sum_{i=1}^k w_i^* L_i}{\sum_{i=1}^k w_i^*}$$
 (2)

where k is the number of studies and  $w_i = 1/(v_i + \sigma^2_{\lambda})$ .

Subsequently, the standard error of the weighted mean was estimated (see Eq. 7 in Hedges et

al., 1999) and the confidence intervals were calculated. For all calculations Microsoft Excel ®

187 2008 was used.

A normal distribution of the mean response ratio was assumed. As described in Hedges

et al. (1999), this assumption can be made, because the single response ratios are normally

distributed as well.

## 2.3.1 Identifying heterogeneity

A test for heterogeneity in effect sizes was performed based on the Q-statistic. Q approximately follows the chi-squared distribution with k degrees of freedom. The Null hypothesis of homogeneity among the effects of different experiments is rejected if Q exceeds the 95 % quantile of the distribution. Heterogeneity results in a positive estimate for the between experiments variance  $\sigma^2_{\lambda}$ , which leads to a larger total variation, that is the sum of

the within and between experiment variance. Consequently, larger standard errors as well as wider confidence intervals for the effect size are computed from the weighted variances.

200

201

## 3. Results

23 studies were obtained from the literature, summarized in Tables 1 and 2. A total of 48 202 203 single experiments, which met the above-mentioned criteria, were extracted from these 204 studies to be included in this meta-analysis. 205 The carbonate chemistry perturbation experiments examining responses of *Emiliania huxleyi* 206 are depicted in Table 1. A total of 19 studies dealt with the responses of 14 different strains to 207 ocean acidification. In most experiments, strains of Emiliania huxleyi showed reduced 208 calcification rates with increased pCO<sub>2</sub> concentrations (Barcelos e Ramos et al., 2010; De 209 Bodt et al., 2010; Delille et al., 2005; Engel et al., 2005; Feng et al., 2008; Gao et al., 2009; 210 Hoppe et al., 2011; Langer et al., 2009; Müller et al., 2010; Riebesell et al., 2000; Rokitta & 211 Rost, 2012; Sciandra et al., 2003; Shi et al., 2009; Wuori, 2012; Zondervan et al., 2002). In 212 other experiments some strains showed an optimum curve in response to increasing pCO<sub>2</sub> 213 (Bach et al., 2011; Langer et al., 2009), no significant response (Langer et al., 2009; Richier et 214 al., 2011) or increased calcification rates (Fiorini et al., 2011; Iglesias-Rodriguez et al., 2008; 215 Shi et al., 2009). 216 Photosynthetic responses were more diverse. In six experiments no response was observed 217 (De Bodt et al., 2010; Delille et al., 2005; Engel et al., 2005; Feng et al., 2008; Fiorini et 218 al., 2011; Hoppe et al., 2011; Müller et al., 2010; Richier et al., 2011), while in another six 219 experiments the POC production increased in response to elevated pCO<sub>2</sub> (Barcelos e Ramos 220 et al., 2010; Hoppe et al., 2011; Iglesias-Rodriguez et al., 2008; Riebesell et al., 2000; Rokitta 221 & Rost, 2012; Shi et al., 2009; Wuori, 2012; Zondervan et al., 2002). Five experiments 222 showed decreasing photosynthesis rates (Bach et al., 2011; Langer et al., 2009; Sciandra et al., 223 2003; Shi et al., 2009), whereas in three experiments an optimum curve was obtained (Gao et 224 al., 2009, Langer et al., 2009). 225 The observed PIC/POC ratios are more homogeneous across experiments with most of them 226 decreasing with increased pCO<sub>2</sub> (Bach et al., 2011; Barcelos e Ramos et al., 2010; De Bodt et al., 2010; Delille et al., 2005; Engel et al., 2005; Feng et al., 2008; Gao et al., 2009; Hoppe et 227 228 al., 2011; Langer et al., 2009; Müller et al., 2010; Riebesell et al., 2000; Rokitta & Rost et al., 229 2012; Shi et al., 2009; Wuori, 2012; Zondervan et al., 2002). Only in four experiments the 230 PIC/POC ratio did not change with increasing pCO<sub>2</sub> (Iglesias-Rodriguez et al., 2008; Langer 231 et al., 2009; Richier et al., 2011; Sciandra et al., 2003) and in one an increase was observed 232 (Fiorini et al., 2011) 233 Experiments with other coccolithophore species also revealed varying responses (Table 2). Of 234 the four experiments with Coccolithus braarudii, two observed a decrease in PIC production 235 with increased CO<sub>2</sub> levels (Krug et al., 2011; Müller et al., 2010), whereas one observed no 236 response (Langer et al., 2006) and the other a slight increase in the calcification rate (Rickaby 237 et al., 2010). The POC production rates varied just as much and increased in two experiments 238 (Rickaby et al., 2010; Müller et al., 2010), while they did not change significantly in another 239 experiment (Langer et al., 2006). In a fourth experiment a non-linear response was observed 240 (Krug et al., 2011). 241 In two experiments conducted with Gephyrocapsa oceanica, the calcification rates decreased 242 (Riebesell et al., 2000) or did not change significantly (Rickaby et al., 2010) with increasing pCO<sub>2</sub>, whereas photosynthetic carbon fixation increased in one experiment (Riebesell 243 244 et al., 2000) and showed an optimum curve in the other one (Rickaby et al., 2010). The

245

PIC/POC ratio declined in both experiments.

In a fourth coccolithophore species, *Calcidiscus leptoporus*, the calcification response was non-linear, while the photosynthesis rate remained constant over the tested CO<sub>2</sub> range (Langer et al., 2006, Langer & Bode, 2011).

## 3.1 Effect of ocean acidification on calcification responses

The meta-analysis of calcification responses to elevated  $CO_2$  concentrations revealed different results between the examined species (Figure 1). Increasing  $CO_2$  concentrations from preindustrial to present day levels had no significant effect on calcification in *Emiliania huxleyi* (lnRR = -0.004). In contrast, the effect of near future  $CO_2$  concentrations under both the 'business as usual' and the 'worst case' scenario had significant negative effects on calcification in this species. This negative effect was more pronounced at 1000 ppm compared to 780 ppm (780 ppm: lnRR = -0.19, confidence interval = -0.07 to -0.30; 1000 ppm: lnRR = -0.38, confidence interval = -0,08 to -0,67).

In *Gephyrocapsa oceanica* an increase from preindustrial to present day  $CO_2$  concentrations had a slightly negative but non-significant effect on calcification. Projected future ocean acidification had a negative mean effect on calcification greater than in *Emiliania huxleyi*, but it was not significant (lnRR = -0.79, confidence interval = 0.61 to -2.19). In contrast, no significant effect of ocean acidification was detected in *Coccolithus braarudii*, where the mean effect sizes were slightly positive at both  $pCO_2$  concentrations. Significant

# 3.2 Effect of ocean acidification on photosynthetic responses

heterogeneity was detected for all calcification responses.

A significant effect of ocean acidification on photosynthesis was observed in *Gephyrocapsa* oceanica for the present-day as well as the high  $CO_2$  concentration, with the mean response at 1000 ppm being more than twice as high (lnRR = 0.57) as the mean response at 380 ppm

(lnRR = 0.24; Figure 2). For *Coccolithus braarudii*, a significant positive effect was observed at 380 ppm and a similar but non-significant positive effect at 1000 ppm. No effect of ocean acidification on photosynthesis was observed for *Emiliania huxleyi* at 380 ppm and 1000 ppm. Only at 780 ppm was the mean effect size slightly positive (lnRR = 0.044), but this effect was non-significant. A significant Q-statistic was calculated for all effect sizes.

responses.

## 3.3 Effect of ocean acidification on PIC/POC responses

The observed PIC/POC responses to an increased  $CO_2$  concentration are similar to those observed for the calcification responses (Figure 3). For *Emiliania huxleyi*, there was a larger negative effect on PIC/POC at 1000 ppm (lnRR = -0.39) than at 780 ppm (lnRR = -0.22), but both responses were significantly negative. No effect was observed at present day  $CO_2$  concentrations.

At both  $CO_2$  concentrations a small, non-significant negative effect of a similar magnitude (380 ppm: lnRR = 0.05, 1000 ppm: lnRR = 0.07) was observed for *Coccolithus braarudii*. The strongest effect of ocean acidification on the PIC/POC ratio was observed for *Gephyrocapsa oceanica*. The mean effect size was significantly negative at both  $pCO_2$  levels, with the negative mean effect size at 1000 ppm (lnRR = 1.37) being more than three times lower than at 380 ppm (lnRR = 0.36). There was significant heterogeneity in all PIC/POC

# 3.4 Relationship between effect sizes and methodological factors

For the three response variables (PIC, POC and PIC/POC) a further meta-analysis was conducted in order to test whether they varied between the two different carbonate chemistry manipulation methods (constant TA vs. constant DIC) used in the experiments.

This meta-analysis revealed that the mean effects of ocean acidification were not consistent between the two methods (Figure 4). Keeping TA constant and changing DIC resulted in a more negative mean effect size for calcification and photosynthesis as compared to constant DIC and variable TA. However, the observed difference between the mean effect sizes for calcification was not significant (p = 0.07) and the overall effect of ocean acidification on calcification was negative, regardless of the manipulation method. In contrast, the mean effect sizes for photosynthesis differed substantially. While no significant effect was observed at constant TA, the effect size at constant DIC was significantly positive. There was significant difference between the mean effect sizes (p = 0.0001). The difference between the effect sizes for PIC/POC was only small. Here, ocean acidification had a slightly more negative effect when keeping DIC constant and changing TA. Both effect sizes were, however, significantly negative. Interestingly, all experiments using Coccolithus braarudii and Gephyrocapsa oceanica manipulated the  $pCO_2$  in the culture medium by adding acid, i.e. changing TA while keeping DIC constant. Thus, all these experiments were included in the constant DIC treatments, while only experiments with *Emiliania huxleyi* were included in the constant TA treatments. In order to eliminate a possible bias due to the unequal distribution of coccolithophore species across carbonate chemistry manipulation methods, a separate meta-analysis was conducted. This analysis only included experiments of *Emiliania huxleyi* and determined the variation of effect sizes between carbonate chemistry manipulations (Figure S1, supplement). The results of this analysis were not significantly different from those obtained from the analysis performed on the full data set. A bias due to the unequal distribution of species between treatments can therefore be ruled out.

318

295

296

297

298

299

300

301

302

303

304

305

306

307

308

309

310

311

312

313

314

315

316

317

#### 4. Discussion

320

321

322

323

324

325

326

327

328

329

330

331

332

333

334

335

336

337

338

339

340

341

342

343

344

The difference in variance between single studies is statistically described as heterogeneity. The term indicates that there is more variability in results than would be expected from the sampling distribution. Differences in the experimental setup, deviations in the measuring method and biological differences between the examined organisms can generally explain the existence of heterogeneity. Heterogeneity in effect size was detected in all analyses in the present study. In retrospect, this finding justifies the use of a random-effect model in this meta-analysis. In contrast to the fixed effect model that only includes variance within the studies, the random effects model accounts for the variance between and within single studies. Our study revealed that heterogeneity in mean effect sizes is not due to different carbonate chemistry perturbation methods. The differences between TA and DIC manipulations in the carbonate chemistry were shown not to cause strong variations in biological responses in coccolithophores - with a possible exception in photosynthetic responses. Another proposed explanation for the high difference in variance between studies could be the morphological and genetic differences of single coccolithophore strains. A high physiological variability was already shown to exist in the coccolithophore Emiliania huxleyi (Iglesias-Rodriguez et al., 2006; Cubillos et al., 2007), with different strains and ecotypes exhibiting diverse responses to ocean acidification (Langer et al., 2009; Hoppe et al., 2011). Moreover, adaption processes of clones that are kept in culture over years could further result in variable responses in CO<sub>2</sub> perturbation experiments (Ridgwell et al., 2009). Thus, a large part of the variance between the analyzed studies is most likely due to intra-species variability of coccolithophore species, especially in Emiliania huxleyi. A further reason for heterogeneity in mean effect size could be discrepancies in calculating the carbonate system from measured parameters. As mentioned earlier in this study, all components of the carbonate system can be calculated if two variables, e.g. pH and DIC, are known. A recently published study suggests, that the pCO<sub>2</sub> concentration measured in CO<sub>2</sub> perturbation experiments differs strongly between calculations (up to 30%), when the input parameters for these calculations were different (Hoppe et al., 2012). The authors state that some publications may not be comparable with each other, as  $pCO_2$  values might have been underestimated when they were calculated from TA and DIC, influencing the interpretation of coccolithophore responses. This finding also has implications for the present study, as some heterogeneity in mean effect size might be due to inconsistencies in calculating  $pCO_2$ . The aim of this study was to synthesize the available data of coccolithophores biological responses to ocean acidification in order to more robustly estimate the actual effect of a lowered seawater pH on those calcifying organism. Despite known intra-specific variability, a negative effect of ocean acidification on calcification as well as on the cellular PIC/POC ratio was observed for the dominant and cosmopolitan species *Emiliania huxleyi*. Our results are in accordance with findings from a meta-analysis conducted by Findlay et al. (2011), who also identified a negative correlation between the cellular PIC/POC ratio in Emiliania huxleyi and the pCO<sub>2</sub> concentration in the culture medium. Although some strains of E. huxleyi appear to be less sensitive to ocean acidification (Langer et al., 2009), the species shows a negative response towards reduced pCO<sub>2</sub> levels in our meta-analysis, suggesting that strain-specific variations are small compared to the generally negative effect of ocean acidification on this species. Calcification and PIC/POC in the coccolithophore Gephyrocapsa oceanica was even more negatively affected by future ocean acidification than in *Emiliania huxleyi*, indicating that G. oceanica is even more sensitive to changes in pCO<sub>2</sub> and pH. Although the meta-analysis with this species was based on only two studies and a significant effect on the calcification response was not observed, the mean effect sizes were even more negative than those

345

346

347

348

349

350

351

352

353

354

355

356

357

358

359

360

361

362

363

364

365

366

367

368

observed for *Emiliania huxleyi* at 1000 ppm. We assume that the inclusion of more studies to the meta-analysis would likely decrease the confidence interval of the mean effect size, resulting in a significantly negative effect of ocean acidification on calcification in Gephyrocapsa oceanica. The strong negative effect of ocean acidification on the PIC/POC ratio in this species was not only due to the strong decrease in calcification, but also a consequence of an increase in the photosynthesis rate with increasing  $pCO_2$ . Apparently, this species profits more from high  $pCO_2$  levels during photosynthesis than the others. This might - at least for Gephyrocapsa oceanica - confirm the hypothesis that some coccolithophores might benefit from higher CO<sub>2</sub> concentrations, since their rate of carbon fixation is below CO<sub>2</sub> saturation at pre-industrial CO<sub>2</sub> levels (Riebesell et al., 2000, 2004; Rost et al., 2003; Nimer & Merrett, 1996). Higher CO<sub>2</sub> concentrations in the water would thus allow them to more efficiently assimilate and fix carbon during photosynthesis and thus increase their photosynthesis rate (Rost et al., 2008). It is further suggested that an increase in the photosynthesis rate might buffer a possible negative effect of ocean acidification on calcification (Ries et al., 2009). When photosynthesis becomes more efficient and additional energy is provided due to enhanced photosynthetic activity, the building and maintenance of coccoliths could be facilitated. This hypothesis, however, was not confirmed by the present analysis, since the species that showed the most positive effect on photosynthesis, Gephyrocapsa oceanica, was also the one where the effect of ocean acidification on calcification was most negative. For *Coccolithus braarudii* the results from the present study confirm the hypothesis that this species is insensitive to elevated pCO<sub>2</sub> levels within the tested range (Langer et al., 2006). To some extent, it might even benefit from higher CO<sub>2</sub> concentrations, as it exhibits a slightly positive photosynthesis response.

370

371

372

373

374

375

376

377

378

379

380

381

382

383

384

385

386

387

388

389

390

391

392

The results for the effect of ocean acidification on calcification gained by the present study are consistent with the observations by Kroeker et al. (2010, 2013) (Figure 5). These authors included responses of all coccolithophore species in one meta-analysis without distinguishing between species, and found a negative but non-significant effect of ocean acidification on calcification. They state that the absence of a significantly negative result might be due to the species-specific responses of coccolithophores, which can be confirmed by our study.

With some coccolithophore species being generally more sensitive with regard to ocean acidification than others, a replacement of sensitive strains by more tolerant strains of the same species or a shift in species composition is probable. It cannot be assessed if a general

same species or a shift in species composition is probable. It cannot be assessed if a general decline in the abundance of coccolithophores with a replacement by other photoautotrophic organism is possible, as long as the role of calcification in coccolithophores is not completely understood. What implications a reduced calcium carbonate production has on the physiological performance and ecological fitness of coccolithophores therefore needs to be further evaluated. Considering that the more prevalent coccolithophore species appear to be vulnerable to ocean acidification, a local or global shift in the species composition or a

replacement by other photoautotrophic organisms may occur and could affect higher trophic

levels and ocean biogeochemical cycling.

Differences between TA and DIC manipulations were not the cause of variable calcification and PIC/POC responses between experiments, confirming earlier results by Kroeker et al. (2009), Findlay et al. (2011) and Hoppe et al. (2011) and following the reviews of Schulz et al. (2009) and Ridgwell et al. (2009). In contrast, mean effect sizes on photosynthetic rates were significantly different between the two manipulation methods. Whereas no effect of ocean acidification on photosynthesis was observed for the constant TA manipulations, the effect in the constant DIC manipulations was significantly positive. This finding is surprising,

as the modifications of the carbonate system induced by the different manipulation methods are very similar, particularly in the range of carbonate chemistry changes projected to occur until the end of this century (Schulz et al., 2009). Although bubbling with CO<sub>2</sub> more closely resembles predicted changes in the oceans carbonate chemistry, because dissolved inorganic carbon increases while total alkalinity remains unchanged, the modification of each carbonate system parameter (pH,  $[CO_2]$ ,  $[CO_3^2]$  and  $\Omega$ Ca) is rather similar. An exception is the concentration of  $HCO_3$ , which increases slightly more in experiments where the  $pCO_2$ concentration is altered by CO<sub>2</sub> bubbling (constant TA manipulation). As not only CO<sub>2</sub>, but also HCO<sub>3</sub> is known to be a carbon source for photosynthesis in most phytoplankton species (Riebesell, 2004), one could assume that the higher HCO<sub>3</sub> concentration in the constant TA manipulations was responsible for the observed difference in photosynthetic responses between manipulation methods. However, a higher rather than a lower photosynthesis rate would be expected in the constant TA manipulations compared to the constant DIC manipulations, as more inorganic carbon in the form of HCO<sub>3</sub> would be available for photosynthesis. Thus, it does not seem likely that the slight deviation in the HCO<sub>3</sub> concentration is responsible for the difference in mean effect sizes between manipulation methods. Nevertheless, discrepancies between the two methods of CO<sub>2</sub> manipulation observed in the present study are consistent with findings of Kroeker et al. (2010). In their meta-analysis a comparison of photosynthetic responses between manipulation methods also showed that keeping TA constant while increasing DIC caused a more negative effect. The deviation between the mean effect sizes was also significant in their study. Although variable photosynthetic responses have been observed in different carbonate chemistry perturbation experiments, it remains to be clarified what causes these differences. To date, studies and reviews have mainly focused on revealing the reason for diverse calcification responses in coccolithophores (Ridgwell et al., 2009; Schulz et al., 2009). This is

419

420

421

422

423

424

425

426

427

428

429

430

431

432

433

434

435

436

437

438

439

440

441

442

probably because ocean acidification is regarded to have a greater impact on calcification in those species than on photosynthesis. While the present study shows that this assumption holds true, a clear understanding of all physiological processes and their relevance for coccolithophore ecological fitness is necessary to realistically assess the influence of future ocean acidification on these organisms. A limitation of the carbonate chemistry manipulation experiments included in this metaanalysis is the short duration of the experiments. As a result, they do not account for possible adaptation processes of coccolithophores that might occur over a longer time-period, and only test for non-adaptive responses. A recent study investigated evolutionary adaptation in E. huxleyi in a long-term experiment (Lohbeck et al., 2012). In this study a population adapted to higher  $pCO_2$  levels showed significantly higher calcification rates than the control population. Although adaptation did not restore calcification rates under elevated pCO<sub>2</sub> to those measured under ambient pCO<sub>2</sub> levels, this observation highlights the possibility of adaptive evolution in coccolithophores. If species like Emiliania huxleyi and Gephyrocapsa oceanica can adapt to decreased pH levels, consequences for the whole ecosystem might be averted. It remains speculative, however, whether results from monocultural experiments can be extrapolated to the natural environment. This also has to be acknowledged when interpreting results of the present study. Generalizations from laboratory observations must be drawn with great care and it has to be kept in mind that ocean acidification is not the only consequence of anthropogenic carbon emissions. Global warming and increased surface ocean stratification as well as changes in nutrient availability will further affect the physiological responses of marine organisms, including coccolithophores. Therefore, the effects of ocean acidification might differ when other potential stressors are included. Some studies have already examined the interactive effects of multiple stress factors on coccolithophore responses (e.g. Zondervan et al., 2002; Feng et al., 2008; De Bodt et al.,

444

445

446

447

448

449

450

451

452

453

454

455

456

457

458

459

460

461

462

463

464

465

466

467

2010; Sett et al., 2014). However, more studies are required that analyze responses of coccolithophores to multiple stressor within the marine ecosystem in order to better quantify community and ecosystem responses to ocean acidification and global warming.

# 472 References

- 473 Armstrong RA, Lee C, Hedges JI, Honjo S, Wakeham SG (2002) A new, mechanistic model
- 474 for organic carbon fluxes in the ocean based on the quantitative association of POC with
- ballast minerals. Deep Sea Research Part II, 49, 219–236.

476

- 477 Bach LT, Riebesell U, Schulz K (2011) Distinguishing between the effects of ocean
- acidification and ocean carbonation in the coccolithophore *Emiliania huxleyi*. Limnology and
- 479 Oceanography, **56**, 2040–2050.

480

- 481 Barcelos e Ramos J, Müller MN, Riebesell U (2010) Short-term response of the
- 482 coccolithophore Emiliania huxleyi to an abrupt change in seawater carbon dioxide
- 483 concentrations. Biogeosciences, 7, 177–186.

484

- Borenstein M, Hedges LV, Higgins J, Rothstein HR (2010) A basic introduction to fixed-
- effect and random-effects models for meta-analysis. Research Synthesis Methods, 1, 97–111.

487

- Brownlee C, Taylor A (2004) Calcification in coccolithophores: A cellular perspective. In:
- 489 Coccolithophores From Molecular Processes to Global Impact (eds Thierstein HR, Young
- 490 JR), pp- 99 125, Springer, Berlin, Germany.

491

492 Caldeira K, Wickett ME (2003) Anthropogenic carbon and ocean pH. Nature, 425, 365.

- Klaas C, Archer DE (2002) Association of sinking organic matter with various types of
- 495 mineral ballast in the deep sea: Implications for the rain ratio. Global Biogeochemical Cycles.
- **16**, 63-1–63-14.

497 Cochran W (1954) The contribution of estimates from different experiments. Biometrics, 10, 498 101–129. 499 500 Conway T, Tans P, NOAA/ESRL (www.esrl.noaa.gov/gmd/ccgg/trends/). 501 502 Cubillos JC, Wright SW, Nash G, de Salas MF, Griffiths B, Tilbrook B, Poisson A, 503 Hallegraeff GM (2007) Calcification morphotypes of the coccolithophorid Emiliania huxleyi 504 in the Southern Ocean: changes in 2001 to 2006 compared to historical data. Marine Ecology 505 Progress Series, 348, 47–54. 506 507 De Bodt C, Van Oostende N, Harlay J, Sabbe K, Chou L (2010) Individual and interacting 508 effects of pCO<sub>2</sub> and temperature on Emiliania huxlevi calcification: study of the calcite 509 production, the coccolith morphology and the coccosphere size. Biogeosciences, 7, 1401– 510 1412. 511 Delille B, Harlay J, Zondervan I, Jacquet S, Chou L, Wollast R, Bellerby RGJ, Frankignoulle 512 513 M, Borges AV, Riebesell U, Gattuso JP (2005) Response of primary production and 514 calcification to changes of pCO<sub>2</sub> during experimental blooms of the coccolithophorid 515 *Emiliania huxleyi*. Global Biogeochemical Cycles, **19**, 1–14. 516 517 Engel A, Zondervan I, Aerts K et al. (2005) Testing the direct effect of CO<sub>2</sub> concentration on

a bloom of the coccolithophorid *Emiliania huxleyi* in mesocosm experiments. Limnology and

518

519

520

Oceanography, **50**, 493–507.

- 521 Feng Y, Warner ME, Zhang Y, Sun J, Fu FX, Rose JM, Hutchins DA (2008) Interactive
- effects of increased  $pCO_2$ , temperature and irradiance on the marine coccolithophore
- 523 Emiliania huxleyi (Prymnesiophyceae). European Journal of Phycology, 43, 87–98.

524

- 525 Findlay HS, Calosi P, Crawfurd, K (2011) Determinants of the PIC:POC response in the
- 526 coccolithophore *Emiliania huxleyi* under future ocean acidification scenarios. Limnology and
- 527 Oceanography, **56**, 1168–1178.

528

- 529 Fiorini S, Middelburg JJ, Gattuso JP (2011) Testing the effects of elevated pCO<sub>2</sub> on
- coccolithophores (Prymnesiophyceae): comparison between haploid and diploid life stages.
- 531 Journal of Phycology, **47**, 1281–1291.

532

- 533 Gao K, Ruan Z, Villafañe VE, Gattuso JP, Helbling EW (2009) Ocean acidification
- exacerbates the effect of UV radiation on the calcifying phytoplankter *Emiliania huxleyi*.
- Limnology and Oceanography, **54**, 1855–1862.

536

- Hedges LV, Olkin I (1985) Statistical Methods for Meta-Analysis. Academic Press, London,
- 538 New York

539

- 540 Hedges LV, Gurevitch J, Curtis PS (1999) The meta-analysis of response ratios in
- experimental ecology. Ecology, **80**, 1150–1156.

542

- Hendriks IE, Duarte CM, Álvarez M (2010) Vulnerability of marine biodiversity to ocean
- acidification: A meta-analysis. Estuarine, Coastal and Shelf Science, **86**, 157–164.

546 Hoppe CJM, Langer G, Rost B (2011) Emiliania huxleyi shows identical responses to 547 elevated pCO<sub>2</sub> in TA and DIC manipulations. Journal of Experimental Marine Biology and 548 Ecology, 406, 54-62. 549 550 Hoppe CJM, Langer G, Rokitta SD, Wolf-Gladrow DA, Rost B (2012) Implications of 551 observed inconsistencies in carbonate chemistry measurements for ocean acidification studies. 552 Biogeosciences, **9**, 2401–2405. 553 554 Iglesias-Rodriguez MD, Schofield OM, Batley J, Medlin LK, Hayes PK (2006) Intraspecific 555 genetic diversity in the marine coccolithophore Emiliania huxlevi (Prymnesiophyceae): the 556 use of microsatellite analysis in marine phytoplankton population studies. Journal of 557 Phycology, 42, 526–536. 558 559 Iglesias-Rodriguez MD, Halloran PR, Rickaby RE et al. (2008) Phytoplankton Calcification 560 in a High-CO<sub>2</sub> World. Science, **320**, 336–340. 561 562 Jones BM, Iglesias-Rodriguez MD, Skipp PJ et al. (2013) Responses of the *Emiliania huxleyi* 563 Proteome to Ocean Acidification. PLoS ONE, **8**, e61868, 1–13. 564 565 Kroeker KJ, Kordas RL, Crim RN, Singh GG (2010) Meta-analysis reveals negative yet

variable effects of ocean acidification on marine organisms. Ecology Letters, 13, 1419–1434.

Kroeker KJ, Kordas RL, Crim R et al. (2013) Impacts of ocean acidification on marine

organisms: quantifying sensitivities and interaction with warming. Global Change Biology,

566

567

568

569

570

**19**, 1884–1896.

- Krug S, Schulz K, Riebesell U (2011) Effects of changes in carbonate chemistry speciation on
- 572 Coccolithus braarudii: a discussion of coccolithophorid sensitivities. Biogeosciences, 8, 771–
- 573 777.

574

- Langer G, Geisen M, Baumann KH, Kläs J, Riebesell U, Thoms S, Young JR (2006) Species-
- 576 specific responses of calcifying algae to changing seawater carbonate chemistry.
- 577 Geochemistry, Geophysics, Geosystems, 7, 1–12.

578

- Langer G, Nehrke G, Probert I, Ly J, Ziveri P (2009) Strain-specific responses of Emiliania
- *huxleyi* to changing seawater carbonate chemistry. Biogeosciences, **6**, 2637–2646.

581

- Langer G, Bode M (2011) CO<sub>2</sub> mediation of adverse effects of seawater acidification in
- 583 Calcidiscus leptoporus. Geochemistry, Geophysics, Geosystems, 12, 1–8.

584

- Lefebvre SC, Benner I, Stillman JH et al. (2012) Nitrogen source and pCO<sub>2</sub> synergistically
- affect carbon allocation, growth and morphology of the coccolithophore *Emiliania huxleyi*:
- potential implications of ocean acidification for the carbon cycle. Global Change Biology, **18**,
- 588 493–503.

589

- 590 Lohbeck KT, Riebesell U, Reusch TB (2012) Adaptive evolution of a key phytoplankton
- species to ocean acidification. Nature Geosciences, 5, 346–351.

592

- Müller MN, Schulz KG, Riebesell U (2010) Effects of long-term high CO<sub>2</sub> exposure on two
- species of coccolithophores. Biogeosciences, 7, 1109–1116.

596 Nakicenovic N, Alcamo J, Davis G et al. (2000) IPCC 2000: Special Report on Emissions 597 Scenarios: A Special Report of Working Group III of the Intergovernmental Panel on Climate 598 Change (eds Nakicenovic N, Swart R) Cambridge University Press, Cambridge, UK and New 599 York, NY, USA. 600 601 Nimer NA, Merrett MJ (1996) The development of a CO<sub>2</sub>-concentrating mechanism in 602 Emiliania huxleyi. New Phytologist, 133, 383–389. 603 604 Richier S, Fiorini S, Kerros ME, Von Dassow P, Gattuso JP (2011) Response of the 605 calcifying coccolithophore Emiliania huxleyi to low pH/high pCO<sub>2</sub>: from physiology to 606 molecular level. Marine Biology, 158, 551–560. 607 608 Rickaby RE, Henderiks J, Young JN (2010) Perturbing phytoplankton: response and isotopic 609 fractionation with changing carbonate chemistry in two coccolithophore species. Climate of 610 the Past, 6, 771–785. 611 612 Ridgwell A, Schmidt DN, Turley C, Brownlee C, Maldonado MT, Tortell P, Young JR 613 (2009) From laboratory manipulations to Earth system models: scaling calcification impacts 614 of ocean acidification. Biogeosciences, 6, 2611–2623. 615 616 Riebesell U, Zondervan I, Rost B, Tortell PD, Zeebe RE, Morel FM (2000) Reduced

calcification of marine plankton in response to increased atmospheric CO<sub>2</sub>. Nature, 407, 364–

617

618

619

367.

620 Riebesell U (2004) Effects of CO<sub>2</sub> Enrichment on Marine Phytoplankton. Journal of 621 Oceanography, **60**, 719–729. 622 623 Riebesell U, Bellerby RG, Engel A et al. (2008) Comment on "Phytoplankton Calcification in 624 a High-CO<sub>2</sub> World". Science, **322**, 1466b. 625 626 Riebesell U, Körtzinger A, Oschlies A (2009) Sensitivities of marine carbon fluxes to ocean 627 change. Proceedings of the National Academy of Sciences USA, 106, 20602–20609. 628 629 Riebesell U, Tortell PD (2011) Effects of ocean acidification on pelagic organisms and 630 ecosystems. In: Ocean Acidification. (eds Gattuso JP, Hansson L), pp. 99–121, Oxford 631 University Press, Oxford, UK. 632 633 Ries JB, Cohen AL, McCorkle DC (2009) Marine calcifiers exhibit mixed responses to CO<sub>2</sub>-634 induced ocean acidification. Geology, 37, 1131–1134. 635 Rokitta SD, Rost B (2012) Effects of CO<sub>2</sub> and their modulation by light in the life-cycle 636 637 stages of the coccolithophore *Emiliania huxleyi*. Limnology and Oceanography, **57**, 607–618. 638 639 Rost B, Riebesell U, Burkhardt S, Sültemeyer D (2003) Carbon acquisition of bloom-forming 640 marine phytoplankton. Limnology and Oceanography, 48, 55–67. 641 Rost B, Riebesell U (2004) Coccolithophores and the biological pump: responses to 642 643 environmental changes. In: Coccolithophores — From Molecular Processes to Global Impact

(eds Thierstein HR, Young, JR), pp. 99-125, Springer, Berlin, Germany.

- Rost B, Zondervan I, Wolf-Gladrow D (2008) Sensitivity of phytoplankton to future changes
- in ocean carbonate chemistry: current knowledge, contradictions and research directions.
- Marine Ecology Progress Series, **373**, 227–237.

648

- Sabine CL, Feely RA, Gruber N et al. (2004) The oceanic sink for anthropogenic CO<sub>2</sub>.
- 650 Science, **305**, 367–371.

651

- 652 Sciandra A, Harlay J, Lefèvre D, Lemée R, Rimmelin P, Denis M, Gattuso JP (2003)
- Response of coccolithophorid *Emiliania huxleyi* to elevated partial pressure of CO<sub>2</sub> under
- nitrogen limitation. Marine Ecology Progress Series, **261**, 111–122.

655

- 656 Sett S, Bach LT, Schulz KG, Koch-Klavsen S, Lebrato M, Riebesell U (2014) Temperature
- modulates coccolithophorid sensitivity of growth, photosynthesis and calcification to
- increasing seawater pCO<sub>2</sub>. PLoS ONE, **9**, e88308.

659

- Shi D, Xu Y, Morel FMM (2009) Effects of the pH/pCO<sub>2</sub> control method on medium
- chemistry and phytoplankton growth. Biogeosciences, **6**, 1199–1207.

662

- 663 Shutler JD, Grant MG, Miller PI, Rushton E, Anderson K (2010) Coccolithophore bloom
- detection in the northeast Atlantic using SeaWiFS: Algorithm description, application and
- sensitivity analysis. Remote Sensing of Environment, **114**, 1008–1016.

666

- Wolf-Gladrow D, Riebesell U, Burkhardt S, Bijma J (1999) Direct effects of CO<sub>2</sub>
- concentration on growth and isotopic composition of marine plankton. Tellus B, **51**, 461–476.

Wuori T (2012) Effects of elevated pCO<sub>2</sub> on the physiology of Emiliania huxleyi. M.Sc. 670 671 Thesis, Western Washington University, USA. 672 673 Ziveri P, de Bernardi B, Baumann KH, Stoll HM, Mortyn PG (2007) Sinking of coccolith 674 carbonate and potential contribution to organic carbon ballasting in the deep ocean. Deep Sea 675 Research Part II, **54**, 659–675. 676 677 Zondervan I, Rost B, Riebesell U (2002) Effect of CO<sub>2</sub> concentration on the PIC/POC ratio in 678 the coccolithophore Emiliania huxleyi grown under light-limiting conditions and different 679 daylengths. Journal of Experimental Marine Biology and Ecology, 272, 55–70.

Table 1. Summary of the available carbonate chemistry manipulation experiments and the responses of *Emiliania huxleyi* as reported by the authors of those studies.

Symbols indicate: — no response, / increased response, ^ non-linear response, \ decreased response

Reference	<i>E.huxlexi</i> strain	Experiment type	CO₂ mani- pulation	PIC production	POC production	PIC/POC	Specifics
Bach at al. (2011)	PML B92/11A	laboratory	constant DIC	$\overline{}$	\	\	large <i>p</i> CO₂ range
Barcelos e Ramos et al. (2010)	Raune Fjord, Norway 2005	laboratory	constant DIC	\	/	\	short-term incubation
De Bodt et al. (2010)	AC481	laboratory	constant TA	\		\	variable temperatures
Delille et al. (2005)	Raune Fjord, Norway 2001	mesocosm	constant TA	\		\	
Engel et al. (2005)	Raune Fjord, Norway 2001	mesocosm	constant TA	\		\	
Feng et al. (2008)	CCMP 371	laboratory	constant TA	\		\	variable light & temperature
Fiorini et al. (2011)	AC472	laboratory	constant TA	/		/	
Gao et al. (2009)	CS369	laboratory	constant TA	\	$\overline{}$	\	PAR & UVR
Hoppe et al. (2011)	RCC1256	laboratory	constant DIC and constant TA	\		\	
Hoppe et al. (2011)	NZEH	laboratory	constant DIC and constant TA	\	/	\	
Iglesias- Rodriguez et al. (2008)	NZEH	laboratory	constant TA	/	/	—	
Langer et al. (2009)	RCC1212	laboratory	constant TA	\	\	\	
Langer et al. (2009)	RCC1216	laboratory	constant TA	\	\	\	
Langer et al. (2009)	RCC1238	laboratory	constant TA	—	$\overline{}$	—	
Langer et al. (2009)	RCC1256	laboratory	constant TA	$\overline{}$	$\overline{}$	—	
Müller et al. (2010)	Raune Fjord, Norway 2005	laboratory	constant DIC	\		\	long-term incubation
Riebesell et al. (2000)	PML B92/11A	laboratory	constant DIC	\	/	\	variable day- lengths & light intensity
Richier et al. (2011)	RCC1216	laboratory	constant TA	—		—	,
Rokitta and Rost et al. (2012)	RCC1216	laboratory	constant TA	\	/	\	low and high light conditions
Sciandra et al. (2003)	TW1	laboratory	constant TA	\	\	—	chemostat
Shi et al. (2009)	NZEH	laboratory	constant TA	\	\	\	
Shi et al. (2009)	NZEH	laboratory	constant DIC	/	/	\	
Zondervan et al. (2002)	PML B92/11A	laboratory	constant DIC	\	/	\	variable day- lengths & light intensity
Wuori et al. (2012)	CCMP 2668	laboratory	constant TA	\	/	\	писнопу

Table 2. Summary of the available carbonate chemistry manipulation experiments and the responses of *Coccolithus braarudii*, *Gephyrocapsa oceanica* and *Calcidiscus leptoporus* found in those studies.

Reference	Species	Strain	Experiment type	CO₂ mani- pulation	PIC production	POC production	PIC/POC
Krug et al. (2011)	Coccolithus braarudii	RCC 1200	laboratory	constant DIC	\	$\overline{}$	\
Langer et al. (2006)		AC400	laboratory	constant DIC	—		
Müller et al. (2010)		RCC 1200	laboratory	constant DIC	\	/	\
Rickaby et al. (2010)		4762	laboratory	constant DIC	/	/	
Riebesell et al. (2000)	Gephyro- capsa oceanica	PC7/1	laboratory	constant DIC	\	/	/
Rickaby et al. (2010)		PZ 3.1	laboratory	constant DIC	—	$\overline{}$	\
Langer et al. (2006)	Calcidiscus Ieptoporus	AC365	laboratory	constant DIC	$\overline{}$		
Langer and Bode (2011)		AC365	laboratory	constant DIC	$\langle$	—	$\frown$

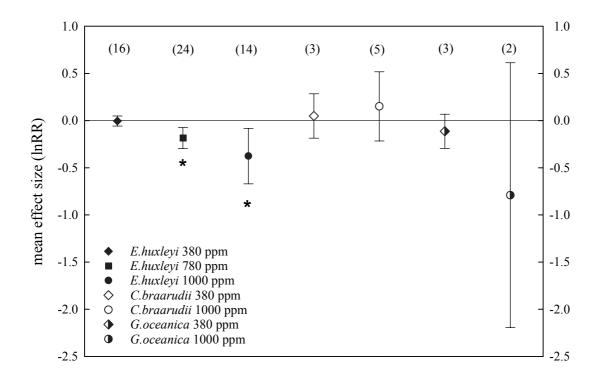
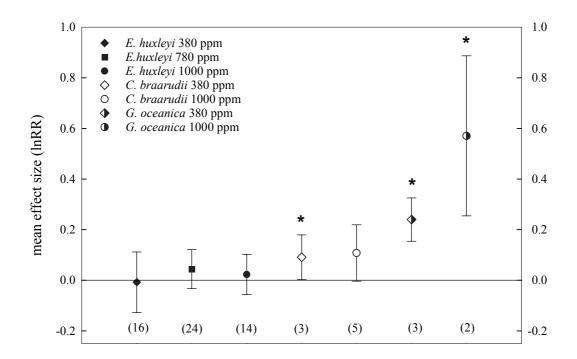


Figure 1. The effect of elevated CO<sub>2</sub> concentrations on the calcification rates of the three coccolithophore species *Emiliania huxleyi*, *Coccolithus braarudii* and *Gephyrocapsa oceanica* [mean effect size and 95% confidence interval]. Responses are relative to 280 ppm.

\* indicates a significant response, which is given when the confidence interval does not overlap zero. The number of experiments used to calculate mean effect sizes are shown in parentheses. The zero line indicates no effect.



**Figure 2.** Mean effect of elevated CO<sub>2</sub> concentrations (relative to 280 ppm) on the photosynthesis rates of three coccolithophore species, *Emiliania huxleyi*, *Coccolithus braarudii* and *Gephyrocapsa oceanica*. Error bars denote the 95% confidence intervals. \* indicates a significant response, which is given when the confidence interval does not overlap zero. The number of experiments included in the meta-analysis is shown in parentheses. The zero line indicates no effect.

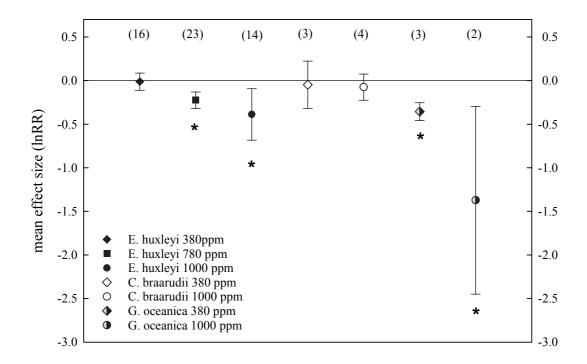
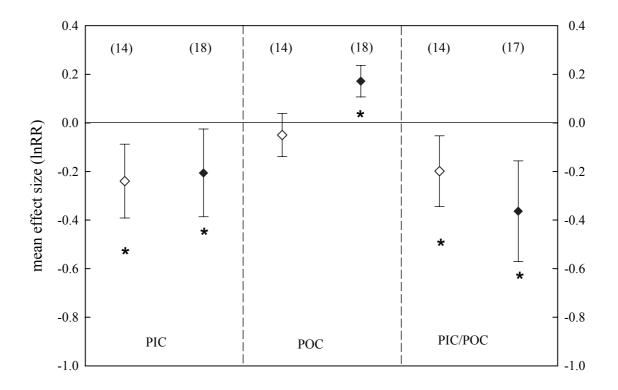


Figure 3. The effect of elevated CO<sub>2</sub> concentrations on the inorganic to organic carbon ratio of three coccolithophore species: *Emiliania huxleyi*, *Coccolithus braarudii* and *Gephyrocapsa oceanica* [mean effect size and 95% confidence interval]. Responses are relative to 280 ppm.

\* indicates a significant response, which is given when the confidence interval does not overlap zero. The number of experiments included in the meta-analysis is shown in parentheses. The zero line indicates no effect.



**Figure 4.** Comparison of effect sizes between the methods of carbonate chemistry manipulation. White diamonds symbolize treatments where total alkalinity [TA] was kept constant while dissolved inorganic carbon [DIC] changed. Black diamonds symbolize treatments where DIC was kept constant and TA varied. The number of experiments included in the meta-analysis are shown in parentheses. The mean effect size is significant when the 95% confidence interval does not overlap zero [\*].

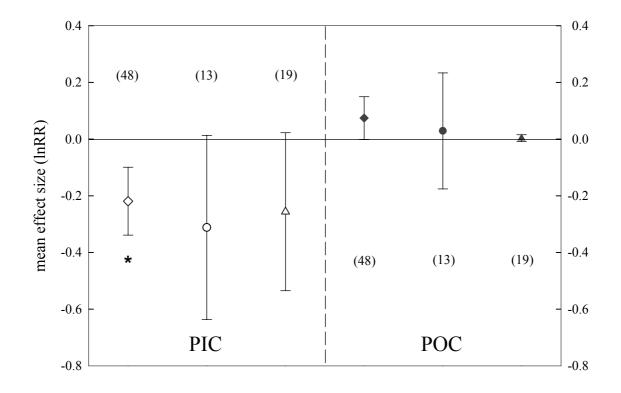


Figure 5. Comparison of effect sizes from PIC and POC analyses derived from the study by Kroeker et al. (2010) [circles], Kroeker et al. (2013) [triangles] and the present study [diamonds]. Data from Kroeker et al. (2010 and 2013) were extracted directly out of the study with the help of the Web Plot Digitizer Software [www.arohatgi.info/WebPlotDigitizer/]. The meta-analysis from the present study contains experiments of all coccolithophore species, including those of *Calcidiscus leptoporus* [see Table 2]. Error bars denote the 95% confidence intervals. \* indicates a significant response, which is given when the confidence interval does not overlap zero. The number of experiments included in the meta-analysis is shown in parentheses. The zero line indicates no effect.