

Response to comments by D. Wiedemeier:

We have a few technical comments: (1) the usage of a B5CA/B6CA ratio and interpretations about condensation could be better introduced and explained. For example, in the abstract, B5CA/B6CA is described as “less condensed to more condensed BC”, which is not fully correct because it actually is just a ratio of less carboxylated to more carboxylated building blocks of a part of BC (B3CA, B4CA exist, too) after BC digestion.

Dr. Wiedemeier has very recently published some interesting work where he examining the degree of BC condensation as a ratio of B6CA to total BPCAs. We have revised the abstract and discussion to use the B6CA:total BC to determine the effect of burn intensity and layer on the degree of condensation of BC in light of Dr. Wiedemeier’s findings. We also describe and maintain the use of B5CA:B6CA as an indication of heat of BC formation as described in Wolf et al. 2013.

The paper that you cited (Rodionov et al. 2006) does not explicitly make this link nor mention BPCA ratios. Moreover, the postulated link in the manuscript between BPCA ratios and degree of processing is still a topic of research and the link between the ratios and age seems rather hypothetical (p 16803).

The reviewer is correct, we provided the wrong citation. It should be Rodionov et al. 2010, Black carbon in grassland ecosystems of the world. Global Biogeochemical Cycles vol. 24 GB3013, doi:10.1029/2009GB003669. We agree the link between age and BPCA ratios has not been well established and have revised the discussion to focus on the potential to relate BPCA ratios with the heat of BC formation.

The fully factorial design of the study allows for a lot of statistics. However, as it is currently presented, it is hard to check if the used statistical methods (ANOVA and post-hoc (F-test?)) really aid data interpretation. Mentioning the number of replicates (also in figures/captions), conducting and showing the residual analysis and model plausibility of ANOVA (in supplement), including a table showing all data including BC (maybe in supplement or extending table 1), and a more transparent handling of statistics could enhance the credibility of the findings. For example, F-values do not add much information when p-values are shown and it is pretty risky to state that burn intensity and layer affect %C when their interaction is significant (p. 16807& table S1).

The statistical analyses were repeated using SAS instead of sigmaplot for increased flexibility and better visualization of residual analysis. The Data Analysis section (2.5) was updated to reflect this change. The number of replicates was added to all figure captions and stated more explicitly in the data analysis section. The results of residual analyses are not typically presented in supplemental information for Biogeosciences Discussion papers, so, similar to the full dataset, we have elected to make them available upon request. F-values were removed from data presentation and text in the results section was updated to note the independent effects of burn severity and layer could not be assessed due to the significant interaction term.

Generally, we think that the manuscript could considerably profit from focusing on the main story (post-fire BC distribution) while additional information can be given in the supplement and/or omitted. Shortening would be particularly useful in the Materials and Method section as it is very long, very descriptive and mostly published elsewhere in detail (e.g. density measurement, elemental, BPCA and statistical analysis). Moreover, the “BPCA method validation” has been published before (in our method paper and supplement), yielding the same CV’s and linearities on a variety of materials and can therefore be omitted or put into the supplement as a “BPCA setup check”.

The manuscript has been refocused to highlight the major findings on post-fire BC distribution, and the methods section was shortened for the BPCA method validation and bulk density determination, the BPCA method validation was moved to the Supplemental Information.

In contrast, some sections could potentially be slightly expanded, for example: Based on the entire BPCA distribution patterns (figure 3), could you distinguish litter BC from soil BC and what are the implications? Or why did you prefer slope as a factor instead of landform (0° can be on a peak or in a valley but with very different erosion characteristics)?

The BPCA distribution patterns discussion was expanded, although, as noted in the results, we could distinguish the soil layers (litter vs. 5-15 cm soils, and 0-5 cm soils vs. 5-15 cm soils) based on the proportion of B6CA and B5CA:B6CA. This sentence was added to explain our choice of soil slope versus landform, "We opted to constrain the study by slope rather than landscape position (e.g. hilltop versus valley location of flat surface) in order to constrain study site criteria to public lands within the patchy distribution of fire-impacted sites of Ponderosa pine vegetation on difficult to access terrain." (inserted in section 2.2 Litter and soil collection). We also added this statement in the discussion, "The position of our sites on the landscape may have also contributed to the lack of effect of slope on BC distribution. Because our aim was to address slope, rather than position, the sites were not oriented in a consistent up- or downslope manner, thus some 0-5° sites are located on hilltops and others at valley bottoms. In addition, the landscape position influences the location of Ponderosa pine through elevational temperature and moisture gradients (Peet, 1981). We focused on the Ponderosa pine because it is the dominant vegetation in the drainage located on a variety of slopes, whereas consideration of hillslope processes would require accounting for the differences in fire properties and BC inputs that would likely result from grass or shrub dominated areas (DeBano, 2000)."

Response to comments by Anonymous Referee #1

Some detail on already published methods section could be reduced, and I would recommend to refocus and polish the Discussion section: - "... BC incorporation processes at depth via water flow and biotic infiltration processes". I was not sure which " biotic infiltration processes" you consider.

The methods section was reduced and the discussion section was expanded and polished. The clause, "stimulated by soil fauna" was added to clarify "biotic infiltration processes" on pg. 16813.

The attempt to relate the ratio B5AC to B6CA to "age" is not substantiated, but if you think so you should explain why you think so

This has been removed and reframed in light of the relationship of B5CA:B6CA on heat of formation.

Is the last paragraph of the Discussion section necessary? "The distribution of BC on a landscapetransport into the soils bias dissolution and translocation". It did not become clear to me how these conclusions relate to your actual results. Or are these (admittedly reasonable) speculations, but still speculations.

The last paragraph of the discussion should be included under the conclusions heading. It provides context for the implications of the study. We studied the distribution of BC on the landscape post-fire, and describe how improving our understanding of distribution is meaningful in relationship to how BC functions in soil. Text was added to justify erosion as a dominant loss mechanism (beginning page 16813 following the discussion of rates of incorporation).

98 Then some more minor remarks: P 16800 L1: communicate clearly and avoid ambigu- ous wording, such
99 as “heterogeneous product of burned biomass”, “critical component of black carbon cycle” L 10: replace
100 “developed and implemented the BPCA method” simply with “used the BPCA method”. I could not see
101 “development” of the BPCA method. It basically reproduces an existing and published method
102 (Wiedemeier et al, 2013). building on previous work (Rodionov et al, 2006, Brodowski et al., Glaser et
103 al.). Adapt wording throughout the text. L15: Explain why you think that none of the post-fire BC has
104 been incorporated to the soil L16: Please explain why you choose to use the ratio B5CA:B6CA, and why
105 you think it can be linked to condensation, age and processing. L25: “is arguably the least understood
106 component of terrestrial C cy- cle” is spurious, there are many other components (microbial etc.), which
107 we do not understand too well yet. Rephrase.
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109 | Recommended sentences have [all](#) been rephrased to improve clarity.
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111 P16801 L14: “BC particles are composed of a refractory aromatic core and a reactive, oxidized patina” –
112 this may not always be true, especially when considering fresh char. Please rephrase. L11-20: include
113 some more recent publications, e.g. Keiluweit et al., 2010, Wiedemeier et al., 2014 L28: surface
114 topography and geomorphology. Specify why you think that there is a difference between the two.
115
116 The statement was rephrased to, “Persistent BC particles in soils are composed of a refractory, aromatic
117 core and a reactive, oxidized patina (Keiluweit et al., 2010; Lehmann et al., 2005) characterized by
118 carbonyl and carboxyl functionalities (Cheng et al., 2008; Cheng et al., 2006).”, and “geomorphology”
119 was removed.
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121 P16802 L11: the list of methods is not extensive, therefore use “or” instead of “and”.
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123 The suggested change was made.
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125 P16803 L4-6: Probably you mean BC stocks and not BPCA stocks L17: Could you add the WRB names
126 for the soils?
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128 The suggested change was made to BC stocks, and the WRB names for the soils were added in the
129 statement, “Soils in the montane forests are Alfisols from the great group cryoboralfs and Mollisols from
130 the suborder ustolls (Peet, 1981).”
131
132 P16804-P16807: Shorten significantly 16808: “BPCA method validation” seems to be a simple quality
133 check.
134
135 The section was shortened as suggested.
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137 P16810 L6: BPCA are molecular markers, not biomarkers. L7: . . . higher proportion of more condensed
138 BC
139
140 The suggested changes were made.
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142 P16814 L20: The conclusion’s section is extremely short and probably not complete.
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144 Additional text was added to the conclusions section rephrasing the major findings of the manuscript.
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146 Figure 1 Since you investigate the effect of topography, an elevation map would be more appropriate than
147 airborne/satellite pictures.
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149 The satellite pictures aid in showing the distribution of the vegetation so we have left the figure as is. If
 150 the editor requires, we can produce an elevation map as well.
 151
 152 Figure 3 Maybe separate litter visually a bit better from the two soil bars in order to guide the reader.
 153
 154 We feel the green color and indication of surface on the graph are adequate for making the litter distinct.
 155 None of the other reviewers commented on this, which suggests that the figure is sufficiently clear.
 156
 157 Table S1-S6 Remove the F-values and make sure you show what is mentioned in the caption, e.g. Table
 158 S5, B4CA:B6CA is not shown in the table.
 159
 160 F-values have been removed, and captions were updated.
 161
 162 Response to comments by Anonymous Referee #3

163 (i) It is clear from the data presented that unburnt litter has a lot less pyrogenic carbon than litter from the
 164 burnt sites, which means that the pyrogenic carbon from previous fires has gone 'somewhere'. The
 165 authors seem to favour erosional loss as the dominant mechanism for the difference, but its not clear to
 166 me why this mechanism is favoured over biotic/abiotic degradation/mineralization (possibly followed by
 167 leaching), given the 40-100 year fire return interval, which leaves plenty of time for both degradation and
 168 erosion. I don't think its possible to pick between these possibilities based on the data presented, so I
 169 would recommend expanding the text to elaborate on the possibility that degradation also a plausible
 170 mechanism for loss. Over a few years, with resampling of the sites, it may well be possible to demonstrate
 171 that erosion is dominant in this environment, but this can't be uniquely concluded to be the dominant
 172 mechanism from a single post-fire sampling. This probably also means softening/modifying some of the
 173 conclusions in the abstract as well as parts of the discussion.

174 Text on the potential for biotic and abiotic degradation mechanisms to contribute to BC losses in this
 175 system was added (pg 16813, new paragraph for loss mechanisms following discussion of potential rate
 176 of incorporation), and reference to preliminary data discussing the contribution of erosion on BC contents
 177 in the CLP was added as well as a reference that describes particulate black carbon export from the CLP
 178 following the HPF (Wagner et al. 2015). We have also softened the concluded loss mechanism in the
 179 abstract and discussion.

180 (ii) As I understand it, the BPCA technique likely under-estimates to pyrogenic carbon because the
 181 technique cannot completely break down large polyaromatic compounds. This is implied on pp16806 line
 182 27 where conversion factors are mentioned but it should be explicitly stated that to stocks are likely
 183 under-estimated by a significant amount - its not a problem for the comparisons in this paper, because like
 184 is being compared with like, but it is important in terms of comparing this data with other techniques.

185 Text added to address this comment (page 16811 at the end of the first paragraph of the discussion) and
 186 reinforce the idea that BPCAs are markers for BC comparative analyses, but their measurement
 187 underestimates 'true' BC stocks.

188 (iii) The BPCA analyses were calibrated through an analysis of a 'laboratory biochar'. Please provide
 189 further information on the characteristics of this char, and the degree to which it might be representative
 190 of natural bushfire chars

191 The laboratory char was combusted pine wood and details on its production have been added to the
 192 supplemental information.

193 (iv) can you add any information on the likely temperatures achieved in the fire at different intensities?
194 This bears particularly on the recalcitrance of the pyrogenic carbon produced in the fire at different
195 intensities, and probably on its observability by BPCA analysis. For example, it is possible that there was
196 'more' total PyC made during the high intensity fire, but less of it measured by BPCA analysis because a
197 higher proportion was not liberated for analysis (ie a correction factor closer to 5 than 2.27. This might
198 explain in part the apparent similarity of stocks between the medium and high intensity burn sites in the
199 litter layer.

200 As for most of wildfires, we unfortunately do not have any data on the temperature of the fire, although
201 the BPCA pattern is related to the highest temperature achieved, suggesting that, given the similar B6CA
202 content and B5CA:B6CA ratio in moderate and highly burned litter, the temperatures of the fires may
203 have been similar. Text to address this comment has been added (p 16814).

204 (v) it seems important to include information about whether the sites chosen are all sites of erosion rather
205 than colluvial accumulation, the latter potentially being sites enhanced burial of pyrogenic carbon from
206 previous fires?

207 It is true that the sites we sampled could experience erosion or accumulation, but given that our study was
208 designed to test the effect of slope rather than landscape position we are unable to evaluate the effect of
209 erosion versus accumulation in a balanced, replicated way. We've added text to the discussion of
210 experimental design (p 16804) and discussion on the effects of slope (p. 16812) that addresses this
211 comment.

212 Smaller things: Page 16801 line 2 – change in which to during which 16803 – 4 – highly is probably not
213 the right word here 16803 – 7 – I think you means BPCA-C? 16803 – 13 – delete first occurrence of the
214 16811 – 25 – how many rain events of what magnitude? 16813 – 10 – this is a place where the erosion
215 argument is made a bit strongly and could be reworded – there are others.

216 The suggested changes were made.

217 Comments by Anonymous Referee #4

218 1) Assumption of this study: The major constraint to compare the BC stocks in the litter/soil of the
219 unburned sites with a fire history and recently burnt site is that we do not have any knowledge of the
220 production rate of BC in these two sites after wildfire. Therefore, there is high uncertainty with respect to
221 loss of BC or incorporation into the soils. The assumptions made in this study and various constraint
222 should be highlighted.

223 This study focused on processes contributing to the distribution of BC in soils after it was deposited on
224 the forest floor, and does not address the amount of BC produced in the fire. Different amounts could be
225 deposited from different fires depending on the type of fire, and climate post-fire, however in our
226 discussion of incorporation of BC into soils we explicitly state, "This calculation contains a high degree
227 of uncertainty..." (p. 16813), and have added text that the incorporation rates will also depend on
228 characteristics of the fire (p. 16813).

229 2) Loss of BC: The presence of BPCA-C in the litter of unburned site indicates fire history. The data
230 indicated that recently burnt sites litter have higher BPCA-C as compared to the unburned sites leading
231 the authors to conclude that erosion is the dominant loss mechanism. As mentioned by the authors, the
232 return interval of fire in this site is 40-100 years, it is possible the BC was lost not only via erosion but
233 also through leaching, degradation leading to mineralization/decomposition both biotically as well as

234 abiotically. The authors should discuss in detail why erosion is the most dominant loss mechanism for this
235 ecosystem and if it is specific to this site. Are there any other study that highlight erosion as one of the
236 prevalent process for this specific site?

237 We have added text that addresses this comment, expanding the discussion of loss mechanisms (pg
238 16813, new paragraph for loss mechanisms following discussion of potential rate of incorporation), and
239 referenced preliminary data discussing the contribution of erosion on BC contents in the CLP, as well as a
240 reference that describes particulate black carbon export from the CLP following the HPF (Wagner et al.
241 2015).

242 3) Burn intensity and BPCAs: Could authors provide any detail regarding the temper- ature range for the
243 wildfire. Does the moderate and high intensity also had differences in the temperature of the wildfire?
244 This is important because the gradient of change in structure between say 300 and 600 degrees is very
245 steep, so a small change in temperature can induce a large change in BC structure and hence differences
246 in the relative proportions of BPCAs in different burn intensity sites.

247 As we said above, we had no data on the temperature range of this wildfire. There was no difference in
248 the proportion of B6CA or the B5CA:B6CA ratio between the moderate and highly burned sites (see table
249 S5 where layer is the only significant effect). The discussion of temperature and the BPCA patterns has
250 been expanded to include the effects that were observed by layer (p. 16814).

251 4) BPCA and age of BC: A word of caution when drawing conclusions from the ra- tios of B5CA and
252 B6CA to the age of BC. There are studies that indicate no change (Schneider et al., 2011) or an increase
253 in B6CA (Hammes et al., 2008) after 100 y or decrease in total aryl group indicating decrease in B6CA
254 (Hilscher and Knicker, 2011). BC with very high proportion of B6CA could also come from high
255 temperature freshly burnt charcoal. It is not clear yet if the changes in the relative contribution of various
256 BPCAs could be linked directly to age of BC. Please clarify what does this ratio implies.

257 The discussion has been modified to remove conclusions about B5CA:B6CA and BC age, and has been
258 refocused on a discussion of temperature of formation.

259 5) BPCA pattern of fresh BC produced after wildfire: The type of BC produced during the wildfire in this
260 particular study has not been made very clear. Did litter layer in moderately burned sites majorly
261 constituted of charred needles while high burn inten- sity sites also had some wood charcoal or the type of
262 BC produced were similar? This is important as BC produced would be qualitatively different and hence
263 would have a difference in BPCA pattern. Did authors also measured BPCA of the BC produced (only
264 charcoal pieces and not litter in general)? This is important to compare the BC produced and the
265 processing it underwent while getting incorporated into soil layers.

266 We described the sites, “Areas were classified as high burn when the fire had burned the entire tree and
267 no needles or small branches remained. Moderate burn areas had ground fire and some crown scorch, but
268 crowns did not burn and at least some needles remained on the trees. Unburned areas had no evidence of
269 ground fire and no evidence of burned material on the ground surface.” The types of BC produced were
270 similar in the highly and moderately burned sites. There were standing, and a few charred logs in both the
271 moderate and highly burned sites, but they were not sampled as part of the soil collection effort, the
272 samples that were collected were as described, “The litter layer was sampled first and then the soil
273 excavated with the help of a hand shovel separately for the 0-5 cm and 5-15 cm depth. Due to the high
274 surface variability, 4 additional litter samples and 3 surface (0-5cm) soil samples were collected at each
275 site, positioning the frame orthogonally to a distance of 2.5 m from the original position. All litter and
276 surface soil samples were pooled by plot.”

277 6) The data in this study indicated that the stocks of BC in unburned sites and highly burned sites are
 278 similar, however, they are distributed differently. It is obvious to have higher BC stocks in the litter layer
 279 of the high burn intensity sites compared to unburned sites after wildfire. However, what could be the
 280 reasons for lower BC stocks in soils of high burn intensity as compared to unburned sites? Discuss.

281 BC stocks were not significantly lower in the high burn intensity than unburned sites (see Table S4). The
 282 paragraph that discusses this (p. 16812) opens with the idea that burned soils can have lower BC stocks
 283 due to combustion of relict BC during recent fires, however we did not we did not find this trend in our
 284 system.

285 7) Validation of BPCA method should be addressed in method section rather than in results as the method
 286 was not developed in this study but adopted from previous studies based on BPCA analysis. Moreover,
 287 Methods and material section could be shortened with proper referencing of the method used.

288 The methods section was shortened and the BPCA method validation was moved to the supplemental
 289 information.

290 Minor corrections

291 BPCA is molecular marker and not biomarker. Please correct it throughout the manuscript.

292 It has been corrected.

293 P16800 L10-remove developed and, "We developed and implemented the benzene polycarboxylic acid
 294 (BPCA) method"

295 It has been removed.

296 P16800 L12: abbreviate black carbon
 297 P16800 L25: add " one of the least"
 298 P16801 L3: "BC" instead of Black C
 299 P16803 L7: it should be either BC or BPCA-C and not BPC-A

300 Suggested changes have been made.

301 P16803 L9-10: rephrase the sentence "We also expected that the BC age and de- gree of processing would
 302 increase with depth" . The manuscript does not highlight how BC age was calculated. Instead of degree of
 303 processing it should be degree of condensation.

304 We have made this change.

305 P16803 L24: HPF was abbreviated here. It should be stated when first used, for eg. in the abstract
 306 (P16800 L 4). Be consistent while using abbreviated terms once defined.

307 We have made this change.

308 P16813 L9-11: The authors calculated based on this study that 17% of the HPF fire- derived BC in litter
 309 would be transferred to the 0–15 cm soil and concluded that the bulk of the BC in this system likely
 310 moves off the landscape through surface runoff. How was this conclusion drawn is unclear. Again, the

311 loss of BC could occur via several other mechanisms including runoffs. This conclusion should be stated
312 as one of the possibilities rather than a clear mechanism occurring in this specific site.

313 We have added text that addresses this comment, expanding the discussion of loss mechanisms (pg
314 16813, new paragraph for loss mechanisms following discussion of potential rate of incorporation), and
315 referenced preliminary data discussing the contribution of erosion on BC contents in the CLP, as well as a
316 reference that describes particulate black carbon export from the CLP following the HPF (Wagner et al.
317 2015).

318 P16813 L13: what does biotic infiltration processes refers here? Please clarify.

319 We have added the phrase, 'stimulated by soil fauna' to clarify the meaning of biotic infiltration.

320 Conclusion: This section here is rather a highlight and appears incomplete. It should include major
321 findings and its implications. The authors should stress on the distribution pattern of BC in different soil
322 layers after wildfire as the major conclusion rather than the loss mechanisms which remain unclear.

323 The 'conclusions heading' should have been at the beginning of the previous paragraph, and the
324 paragraph has been re-written as suggested.

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Distribution of black carbon in Ponderosa pine forest floor and soils following the High
Park wildfire

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375 Abstract

376 Biomass burning produces black carbon (BC), effectively transferring a fraction of the
 377 biomass C from an actively cycling to a passive C pool, which may get stored in the soil. Yet
 378 the timescales and mechanisms for incorporation of BC into the soil profile are not well
 379 understood. The High Park Fire (HPF), which took place in northwestern Colorado in the
 380 summer of 2012, provided an opportunity to study the effects of both fire severity and
 381 geomorphology on properties of carbon (C), nitrogen (N), and BC in the Cache La Poudre
 382 River drainage. We sampled montane Ponderosa pine forest floor, 0-5 cm soils, and 5-15
 383 cm soils four months post-fire in order to examine the effects of slope and burn severity on
 384 %C, C stocks, %N and BC. We used the benzene polycarboxylic acid (BPCA) method for
 385 quantifying BC. With regard to slope, we found that steeper slopes had higher C:N than
 386 shallow slopes, but that there was no difference in BPCA-C content or stocks. BC content
 387 was greatest in the forest floor in burned sites (19 g BPCA-C kg⁻¹ C), while BC stocks were
 388 greatest in the 5-15 cm subsurface soils (23 g BPCA-C m⁻²). At the time of sampling,
 389 unburned and burned soils had equivalent BC content, indicating none of the BC deposited
 390 on the land surface post-fire had been incorporated into either the 0-5 cm or 5-15 cm soil
 391 layers. The ratio of B6CA:total BPCAs, and index of the degree of aromatic C condensation,
 392 suggested that BC in the 5-15 cm soil layer may have been formed at higher temperatures
 393 than the forest floor, and 0-5 cm soils. Total BC soil stocks were relatively low compared to
 394 other fire-prone grassland and boreal forest systems, indicating most of the BC produced in
 395 this system is likely lost from the system, either through erosion events, degradation or
 396 translocation to deeper soils. Future work examining mechanisms for BC losses from the
 397 forest soil will be required for understanding the role BC plays in the global carbon cycle.

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453 1. Introduction

454 While pyrogenic or black carbon (BC) is now recognized as a ubiquitous soil carbon (C)
455 fraction it is one of the least understood components of the terrestrial C cycle. Every year,
456 fire burns approximately 10-15 $\times 10^6$ ha of boreal and temperate forest and more than 500
457 $\times 10^6$ ha of tropical and subtropical forests and savannas (Goldammer and Crutzen, 1993;
458 Knicker, 2011), during which 0.12 to 9.5% of the burned biomass is converted to BC
459 (Forbes et al., 2006). Black C is utilized by soil microbes, but at a slow rate (Santos et al.,
460 2012). Thus it generally resides in the soil for a long time (from centuries to millennia)
461 (Singh et al., 2012), acting as a long-term C sink, with a potential negative feedback on
462 climate warming. However, BC stocks in soils are not only related to BC production rate
463 and decomposition, but may also be lost through runoff, leaching or burning (Czimczik and
464 Masiello, 2007; Foereid et al., 2011), and thus BC stocks are strongly dependent on surface
465 topography and soil physical-chemical environment (Bird et al., 2015; Knicker, 2011).

466 BC persistence and dynamics in soil seem to be controlled by mechanisms similar to
467 those that control soil organic matter dynamics including inherent chemical recalcitrance
468 and organo-mineral interactions (Knicker, 2011). Persistent BC particles in soils are
469 composed of a refractory, aromatic core and a reactive, oxidized patina (Keiluweit et al.,
470 2010; Lehmann et al., 2005) characterized by carbonyl and carboxyl functionalities (Cheng
471 et al., 2008; Cheng et al., 2006). The degree of condensation of the aromatic core has been
472 shown to be quite variable (McBeath and Smernik, 2009; Wiedemeier et al., 2015d), but
473 can be broadly characterized as dominated by C in condensed aromatic rings resistant to
474 decomposition (Baldock and Smernik, 2002). Besides its inherent chemical recalcitrance,
475 BC stabilization in soils likely occurs through bonding to minerals, which is thought to be
476 the most persistent mechanism of SOM stabilization (von Lutzow et al., 2006). The
477 presence of carboxyl functionalities on BC surfaces provides 'teeth' available to chelate soil
478 aluminum and iron, creating BC-mineral complexes that are highly refractory to microbial
479 decay and have longer mean residence times than non mineral-associated BC (Christensen,
480 1996; von Lutzow et al., 2006).

481 In order to become stabilized in soils, BC must first be transferred from burned surface
482 material to the subsurface, and the process of incorporation will be strongly related to
483 surface topography. The shape of a landscape and propensity for erosion versus deposition

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489 is dependent upon several variables including bedrock composition, slope, elevational
490 gradients in temperature and precipitation, and disturbance history such as the frequency
491 of wildfires. While the strong relationship between geomorphology and soil
492 erosion/sediment transport is fairly well understood (Ritchie and McCarty, 2003; Slater
493 and Carleton, 1938), the relationship between soil erosion and fate of different components
494 of SOM that are eroded, including BC, are relatively unknown (Bird et al., 2015; Rumpel et
495 al., 2006).

496 The difficulty in measuring BC contributes to our limited understanding of its transport
497 processes and function in the global C cycle. Because BC exists along a continuum of
498 combustion products, from charred biomass to soot, with differing physical and chemical
499 features, no single method can accurately quantify total BC content (Hammes et al., 2007;
500 Masiello, 2004). Visual counts of charcoal, resistance to oxidation methods, nuclear
501 magnetic resonance spectroscopy (NMR) ~~or~~ the quantification of BC-specific molecular
502 markers (e.g., benzene polycarboxylic acids (BPCAs)) have each been employed for
503 quantification of BC. While each approach has advantages and disadvantages, the BPCA
504 method has been shown to yield conservative estimates of BC with charred inputs and
505 more consistent results than many other quantification methods (Hammes *et al.* 2007).
506 Moreover, the BPCA method yields additional information about BC quality related to its
507 degree of aromatic condensation and aromaticity, (Schneider et al., 2010; Schneider et al.,
508 2013; Wiedemeier et al., 2015a; Ziolkowski et al., 2011).

509 A few estimates exist of BC production after fire (Santín et al., 2012), as well as of BC
510 stocks in soils for different ecosystems (Bird et al., 1999; Cusack et al., 2012; Schmidt et al.,
511 2002). Yet, estimates of BC production and losses are not balanced (Czimczik and Masiello,
512 2007; Rivas et al., 2012), clearly identifying our lack of understanding and the need for a
513 full accurate accounting of BC dynamics after fire at the watershed level.

514 Between June 9th-24th, 2012, the High Park fire (HPF) burned more than 35,000 ha in
515 northern Colorado along the Cache la Poudre (CLP) River in an area dominated by
516 ponderosa pine (*Pinus ponderosa*) (Figure 1). The aims of this work were to: 1. determine
517 the C, and BC stocks, and the proportion of C that was BC, in Ponderosa pine forest floor,
518 and soils following the HPF; 2. examine the effects of burn severity and landscape slope on
519 soil C, N and proportion of BC; and 3. use the distribution of individual BPCAs to

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understand the degree of condensation of BC through the soil profile. We expected that BC stocks would be the greatest in high burn severity sites, followed by moderate, then unburned sites, and that the hillslope would have the opposite effect, with lowest BC stocks on the steepest slopes, and greatest BC stocks on shallow slopes. We anticipated that BC and C stocks would be greater in the forest floor than in soils, and that soil C stocks would be diminished in high burn severity surface soils due to combustion during fire. We also expected that the molecular characteristics of BC would change with depth related to their degree of condensation.

2. Materials and Methods

2.1 Experimental Design and Site Identification

The sites were located within the montane forest (elevation 1750 to 2850 m) of the CLP drainage which is dominated by Ponderosa pine (*Pinus ponderosa*) and Douglas fir (*Pseudotsuga menziesii*) and also includes aspen (*Populus tremuloides*), Rocky Mountain juniper (*Juniperus scopulorum*), lodgepole pine (*Pinus contorta*) and other species (Veblen and Donnegan, 2005). Soils in the montane forests are Alfisols from the great group cryoboralfs and Mollisols from the suborder ustolls (Peet, 1981).

The montane ponderosa pine forest has a variable severity fire regime meaning there is a mixture of both high severity, full or partial stand replacing fires, and low severity, non-lethal, surface fires. The mean return interval is approximately 40 to 100 years and most fire events have both high and low severity components, and are caused by a combination of human and lightning strike ignition (Veblen and Donnegan, 2005). A lightning strike started the HPF on June 9, 2012. It burned over 35,000 ha in the mountainous region of the CLP River drainage through early July 2012.

Our study was a fully factorial, randomized block design with four replicate blocks for all treatments plots, including three levels of burn severity (unburned, moderate burn, high burn) and three slopes (0-5, 5-15 and 15-30 degrees), for a total of 36 plots. We opted to constrain the study by slope rather than landscape position (e.g. hilltop versus valley location of flat surface) in order to constrain study site criteria to public lands within the patchy distribution of fire-impacted sites of Ponderosa pine vegetation on difficult to access terrain. Geographic Information System (GIS) layers of land ownership, slope, fire intensity,

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582 and burn severity were obtained prior to site location. Potential sampling areas were
 583 chosen in state or federal land in areas of homogenous vegetation stands where all slope
 584 classes and fire classes were present within a close distance (Figure 1). Ground truthing
 585 was subsequently done to locate each specific slope and burn severity sampling treatment
 586 plot. Slopes were determined using a clinometer. Areas were classified as high burn when
 587 the fire had burned the entire tree and no needles or small branches remained, the litter
 588 layer was consumed in the fire and there were some small pieces of charcoal throughout
 589 the surface layer. Moderate burn areas had ground fire and some crown scorch, but
 590 crowns did not burn, at least some needles remained on the trees and the litter layer
 591 remained on the forest floor with some small pieces of charcoal. Unburned areas had no
 592 evidence of ground fire and no evidence of burned material on the ground surface.

593 insert Figure 1

594 2.2 Forest Floor and Soil Collection

595 Soil and forest floor samples were collected between October and November of 2012.
 596 At each of the 36 experimental plots, forest floor and soils were collected from within a 20
 597 by 20 cm wooden frame, and frame GPS coordinates were recorded. The forest floor layer
 598 was sampled first including any litter plus organic soils down to the mineral layer, and then
 599 the soil was excavated with a hand shovel separately for the 0-5 cm and 5-15 cm depth.
 600 Due to the high surface variability, 4 additional forest floor samples and 3 surface (0-5cm)
 601 soil samples were collected at each site, positioning the frame orthogonally to a distance of
 602 2.5 m from the original position. All forest floor and surface soil samples were pooled by
 603 plot.

604 Due to the extreme rockiness at all of the sampling locations soil bulk density was
 605 determined using pit excavation separately for each depth layer (Page-Dumroese et al.,
 606 1999). The volume of the pit was determined using volume displacement with millet seed
 607 (detailed description in Supplemental Information). Soil and forest floor samples were
 608 transported to the lab and stored at 4°C until processing.

610 2.3 Forest Floor Soil Pretreatment and Elemental Analyses

611 In the laboratory, forest floor samples were weighed field-moist and a subsample of
 612 each was dried at 105°C for 48 hours for dry weight correction. Forest floor samples were

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646 then air-dried and another subsample taken and heated in a muffle furnace at 600°C for 12
647 hours to correct forest floor dry weight for ash content. All remaining air-dried forest floor
648 samples were passed through an 8 mm sieve, and any large pieces of plant material were
649 broken up with clippers prior to the samples being ground with a 0.75 mm mesh screen
650 equipped Wiley mill, and dried overnight at 60°C.

651 Soil samples were weighed field-moist and a subsample of each was dried at 105°C for
652 48 hours for dry weight correction. Bulk density of each soil depth was calculated as the
653 weight of oven dry soil with rock removed (Throop et al., 2012) divided by the volume for
654 the depth determined by millet with rock volume removed. Soils were sieved air dry to 2
655 mm and a subsample was oven dried (105°C) and finely ground. All the ground, dry, forest
656 floor and soil samples were analyzed for total C and N by an elemental analyzer (LECO
657 CHN-1000; LECO Corporation, St. Joseph, MI, USA), and for BC by the BPCA method as
658 described below.

660 2.4 BPCA Analyses

661 The BPCA method converts condensed aromatic structures to single aromatic rings
662 with variable numbers of carboxylic acid moieties, and a greater degree of condensation
663 (i.e. number of fused rings) correlates with a greater number of carboxylic acid moieties on
664 the individual BPCAs such that more condensed structures result in greater relative
665 abundance of B6CAs and the least condensed BC would result in a greater proportion of
666 B3CAs (Glaser et al., 1998; Wiedemeier et al., 2015a; Ziolkowski et al., 2011). Black C was
667 determined on all forest floor and soil samples using high performance liquid
668 chromatography (HPLC) equipped with a photo diode array detector to quantify benzene
669 polycarboxylic acids (BPCA) as described by (Wiedemeier et al., 2013). The BPCA method
670 was validated with biochar amended soils from the field site (see Supplemental
671 Information). Briefly, 50-150 mg of finely ground, oven dried sample was digested with
672 70% nitric acid for 8 hours at 170°C. The solution was filtered with ashless cellulose filters,
673 an internal reference standard of phthalic acid was added to the solution, and the filtrate
674 was cleaned by cation exchange resin and freeze-dried. The freeze-dried sample was re-
675 dissolved in HPLC grade water. The re-dissolved solution containing the BPCAs was
676 separated with a reversed stationary phase column (Waters X-Bridge C18, 3.5 um particle

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size, 2.1 x 150 mm) using **standard gradient conditions**. Individual BPCAs were quantified with using a **five** point calibration from standard solutions of benzenetricarboxylic acids (1,2,3-B3CA/hemimellitic acid, 1,2,4-B3CA/trimellitic acid, 1,3,5-B3CA/trimesic acid), benzenetetracarboxylic acid (1,2,4,5-B4CA/pyromellitic acid), benzenepentacarboxylic acid (B5CA), and benzenhexacarboxylic acid (B6CA/mellitic acid). The B4CA standards that are not commercially available (1,2,3,4-B4CA/prehenitic acid, and 1,2,3,5-B4CA/mellophanic acid) were identified by their ultraviolet adsorption spectra and quantified using the calibration for 1,2,4,5-B4CA (Yarnes et al. 2011). Previous attempts to calculate a BPCA-C to BC conversion factor have resulted in values that range from 2.27 to 5, and have been difficult to reproduce (Brodowski et al., 2005; Glaser et al., 1998; Ziolkowski et al., 2011). Thus, to simplify empirical comparisons we report values as BPCA-C, either as a proportion of total C or as a stock.

2.5 Data Analyses

The effects of layer (**forest floor**, 0-5 cm soil, 5-15 cm soil, **n=4 per layer**), slope (0-5°, 5-15°, 15-30°, **n=4 per slope**) burn **severity** (unburned, moderate burn, high burn, **n=4 per severity**) and all interaction terms on each response variable (soil C, soil N, BPCA C stock, BPCA C as a proportion of total C, and relative abundances of B4CA, B5CA, B6CA and B5CA:B6CA ratio) were compared using the SAS mixed procedure (proc mixed); fixed variables were layer, slope and severity, and block and core were designated as random effects. Post-hoc analysis for significant terms was conducted using Tukey's test. When necessary, dependent variable data were log-transformed (%C, %N, C stock, BPCA-C g m⁻²) to meet assumptions of equal variance and normality, which were assessed with Studentized residual diagnostic plots. The null hypothesis, that the independent factor had no effect, or that no linear correlation existed between variables, for all tests was evaluated at $\alpha < 0.05$. Analyses were run using SAS 9.4.

3. Results

3.1 Percent and Stocks of C and N in Forest floor and Soil

Values for %C ranged from 29% in the **forest floor** to 0.9% C in the 5-15 cm soil layer, for %N from 0.8% in **forest floor** to 0.08% in the 5-15 cm soil layer, and for C:N from 40 in

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the **forest floor** to 13 in 5-15 cm soil. We tested for effects of layer, burn **severity**, slope and **their interactions and** found that the main effects were distinct for each response variable (%C, %N and C:N). **Effects of burn severity** ($p=0.002$) and layer (<0.001) **on %C, could not be independently assessed because** the burn \times layer interaction **was also significant** ($p<0.001$, **Table S1**). Only layer had an effect on %N ($p<0.001$), while the C:N ratio was affected by slope ($p=0.009$), burn intensity ($p<0.001$), layer ($p<0.001$) **and in** interaction ($p<0.001$).

insert Table 1

Post hoc comparisons (**Table S2**) confirmed expected decreases in %C and %N from **forest floor** to 5-15 cm soil ($p<0.001$ for each successive layer), along with a decreasing C:N from **forest floor** to 0-5 cm soil ($p<0.001$) and with no change between 0-5 cm and 5-15 cm soil ($p=0.703$). The burn **severity** \times layer interaction term illustrated that the effects of burn were confined to the **forest floor** layer for %C and C:N. Within the **forest floor** layer, unburned sites had greater %C than moderately burned ($p=0.009$) or highly burned sites ($p<0.001$), and moderately burned sites also had greater %C than highly burned sites ($p<0.001$). For the C:N ratio the pattern was the same: C:N was widest in unburned sites which decreased significantly at moderately burned sites ($p<0.001$) and further still at highly burned sites ($p<0.001$). Interestingly, slope also had an effect on the C:N. Post hoc comparisons indicated that C:N on 0-5 degree slopes was **lower** than 5-15 degree slopes ($p=0.028$) and the C:N on 0-5 degree slopes was significantly lower than 15-30 degree slopes ($p=0.012$), while the 5-15 and 15-30 degree slopes were not different ($p=0.916$).

Total C stocks varied considerably between the layers from 3.8 in **forest floor** to 25.3 tons C per hectare in the 5-15 cm soil layer. **The** only significant effect on total C stocks was depth ($p<0.001$) with the **forest floor** having a smaller C stock than 0-5 cm and 5-15 cm soil layers ($p<0.001$, for each). Soil bulk density values were not significantly different among any of the study sites (**Table S1**).

3.2. Benzene Polycarboxylic Acid-C in **forest floor** and soil

We determined BPCA-C both in reference to the amount of carbon and the stock by volume of soil or **forest floor** and found highly variable amounts of BPCA-C for both metrics. For **forest floor**, concentration values ranged from 0.09 g kg⁻¹ OC in unburned

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forest floor, to 40.0 g kg⁻¹ OC in highly burned forest floor, and for stocks from 0.1 g m⁻² in unburned forest floor to 19.52 g m⁻² in moderately burned forest floor. In soils, concentration ranged from 2.86 g kg⁻¹ OC in moderately burned 0-5 cm soil to 33.83 g kg⁻¹ OC in 5-15 cm highly burned soils, and stocks ranged from 2.92 g m⁻² in highly burned 0-5 cm soils to 96.66 g m⁻² in unburned 5-15 cm soil.

Burn severity and layer were the main effects on the concentration and stock of BPCA-C (Figure 2, Table S3). Results of a mixed model (slope, burn severity, layer and interactions) indicated that there was no significant effect of slope either independently (p = 0.446) or in interaction (slope x burn p=0.191, slope x layer p=0.740) on BC concentration. Mean values for BPCA-C stock did decrease with increasing slope in moderately burned forest floor (0-5 degrees: 18.2 +/- 7.1 g m⁻²; 5-15 degrees: 14.8 +/- 4.7 g m⁻², 15-30 degrees: 11.8 +/- 4.3 g m⁻²), however the trend was not significant due to high variability. The independent effects of burn severity (concentration: p=0.007, stock: p=0.012), and layer (concentration: p=0.610, stock p<0.001) could not be interpreted independently as the interaction of burn severity and layer was also significant (concentration and stock: burn x layer p<0.001).

Post hoc comparisons indicated that within the forest floor layer, high and moderately burned material contained significantly more BPCA-C, both by concentration and stock, than unburned material (Table S4, p<0.001 for both). Within the 0-5 and 5-15 cm layers, there was no statistically significant difference in BPCA-C concentration or stock regardless of burn severity. Within unburned layers, 0-5 and 5-15 cm soils had significantly greater amounts of BPCA-C than forest floor, both by concentration (p<0.001, p=0.004, respectively) and stock (p<0.001 for both). Within high burn severity, forest floor and soil BPCA-C stocks and concentrations yielded distinct results: the amount of C that was BPCA-C was greater in the forest floor than 0-5 and 5-15 cm soils (p=0.023, p=0.027, respectively), whereas the stock of BPCA-C was not significantly different in the high burn among forest floor and soil layers.

insert Figure 2

We expected that the layer (forest floor, 0-5 cm soil, 5-15 cm soil) and burn severity may contribute to the distribution of BPCAs with BC formed at different temperatures (B5CA:B6CA), or by a higher proportion of more condensed C (B6CA:total BPCAs) with

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995 increasing soil depth. Overall, the bulk of BPCAs were B5CA and B6CA varieties, together
 996 making up approximately 80% of the total BPCA-C. The B4CAs were the next most
 997 abundant (10-20%) and the B3CAs were less than 3%. Results from statistical analyses
 998 indicated that 'layer' was the main effect on the distribution of BPCAs ($p < 0.001$, Table S5).
 999 Layer also had a significant effect on the ratio of B5CA to B6CA, ($p = 0.002$, Table S5, Figure
 1000 4). Post hoc comparisons were used to evaluate the relative abundance of each BPCA by
 1001 layer: the proportion of B6CA was greater in the 5-15 cm soil than both 0-5 cm soil and
 1002 forest floor layers ($p < 0.001$); B5CA was greater in the forest floor than 0-5 cm soil and 5-15
 1003 cm soils ($p < 0.001$ for both), and greater in 0-5 cm soils than 5-15 cm soils ($p = 0.037$); B4CA
 1004 was greater in 0-5 cm soil than in forest floor ($p < 0.001$), and 5-15 cm soil ($p = 0.002$), with
 1005 no difference in forest floor and 5-15 cm soil ($p = 0.148$). The ratio of B5CA:B6CA decreased
 1006 with depth due to both decreasing amounts of B5CA and increasing amounts of B6CA. The
 1007 B5CA:B6CA ratio was significantly greater in the forest floor than 5-15 cm soils ($p = 0.001$)
 1008 (Figure 3, Table S6)).

1009 insert Figure 3

1011 4. Discussion

1012 Our primary objective was to determine the C_{socks}, BPCA-C stocks, and the proportion
 1013 of C that was BPCA-C in Ponderosa pine forest floor and soils following the HPF. BC can
 1014 account for 1% to 45% of the soil organic C depending upon fire return interval (Czimczik
 1015 et al., 2005; Saiz et al., 2014), ecosystem type, soil mineralogical properties (Preston and
 1016 Schmidt, 2006) other factors that influence OC stabilization (Knicker, 2011), as well as the
 1017 method used for quantification. Estimates of BC content based on BPCA-measurements are
 1018 generally lower than those made with chemical- thermal- or photo-oxidation based
 1019 measurements or with NMR (Preston and Schmidt, 2006). Only a few studies have
 1020 estimated the amount of BC in forest soils using the BPCA method with values that range
 1021 10 to 60 g/kg organic C and 0-80 g/m² (Czimczik et al., 2003; Czimczik et al., 2005;
 1022 Rodionov et al., 2006). Excluding unburned forest floor samples, we found values within
 1023 this range, averaging 14 (± 7) g BPCA-C kg⁻¹ C, and 19 (± 5) g BPCA-C per m². It is important
 1024 to note that BPCAs are markers for BC, and their total amount is two to five times lower

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1064 [than the amount of BC. This should be taken into consideration when comparing BPCA](#)
1065 [estimates with BC distribution values in systems that have been assessed with different](#)
1066 [methods](#) (Brodowski et al., 2005; Glaser et al., 1998; Ziolkowski et al., 2011).

1067 We also aimed to determine how the slope of the landscape and burn [severity](#) would
1068 influence the amount of BC in [forest floor](#) and soil layers following a major wildfire. We
1069 found that neither slope nor burn [severity](#) had an effect on BC concentration in soils.
1070 Interestingly, even the soils from unburned sites had an average BC content of 14 g BPCA-C
1071 kg⁻¹ of C, suggesting [a persistent BC pool](#) from past fires. Within the [forest floor](#) layer,
1072 however, unburned sites contained very low BPCA-C and moderate and highly burned sites
1073 contained significantly more, averaging 18 g BPCA-C per kg of OC suggesting that the
1074 majority of the BC remaining on the landscape [after the HPF](#) persisted in the [forest floor](#)
1075 rather than moving into the surface soil four months post-fire.

1076 We expected that during the interval between the HPF (June 2013), and sample
1077 collection (October 2013), HPF-derived BC would have begun to move off of steeper slopes
1078 during post-fire erosion events, resulting in lower BC deposits on steeper slopes. However,
1079 we observed consistent BC content across slopes with the HPF-derived BC isolated to the
1080 [forest floor](#) layer in both highly and moderately burned areas on a per unit C and per m²
1081 basis. Although slope did not contribute to the landscape pattern of BC distribution over
1082 the time period of our study, the summer of 2013 was particularly dry with very few high
1083 intensity rain events (Wohl, 2013). Thus, slope may only become a contributing variable to
1084 landscape-level post-fire BC distribution when there are precipitation events sufficient to
1085 produce [significant](#) sediment movement. In addition, steeper slopes generally have
1086 increased surface roughness in montane systems constraining overland sediment
1087 movement (Wohl, 2013). We qualitatively examined photos of each of the collection sites
1088 and noted increased surface roughness in some of the steeper replicates, thus increased
1089 surface roughness is [a plausible explanation for similar BPCA-C values on shallow vs.](#)
1090 [steeper slopes.](#)

1091 [The position of our sites on the landscape may have also contributed to the lack of effect](#)
1092 [of slope on BC distribution. Because our aim was to address slope, rather than position, the](#)
1093 [sites were not oriented in a consistent up- or downslope manner, thus some 0-5° sites are](#)
1094 [located on hilltops and others at valley bottoms. In addition, the landscape position](#)

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1106 [influences the location of Ponderosa pine through elevational temperature and moisture](#)
 1107 [gradients](#) (Peet, 1981). [We focused on the Ponderosa pine because it is the dominant](#)
 1108 [vegetation in the drainage located on a variety of slopes, whereas consideration of hillslope](#)
 1109 [processes would require accounting for the differences in fire properties and BC inputs that](#)
 1110 [would likely result from grass or shrub dominated areas](#) (DeBano, 2000).

1111 The only variable that we found responsive to slope was the C:N ratio which increased
 1112 with increasing slopes. The constituent %C and %N values were not significantly different
 1113 by slope, so the pattern [was](#) driven by both slight increases in %C and decreases in %N
 1114 (Table 1). The trend of higher C:N at steeper sites has been noted on the Colorado Plateau
 1115 (Norton et al., 2003), and was attributed to the accumulation of fresh, plant derived high
 1116 C:N [forest floor](#) on steeper slopes in a N immobilizing environment, and the movement of
 1117 lower C:N, partially decomposed material, downslope with rain events. Thus over time,
 1118 steeper slopes do preferentially move material downslope but this export mechanism [did](#)
 1119 [not](#) apply to the BC that was stabilized in soils over time.

1120 Concentrations of post-fire BC have been shown to be highest in the surface of
 1121 moderately burned soils [due to consumption of relict BC content in highly burned areas](#)
 1122 [\(Czimczik et al., 2003\)](#). [However, in](#) our study, on a per unit C basis the amount of BC in
 1123 surface 0-5 cm soils was [not distinguishable](#) across burn intensities (~ 14 g BPCA-C kg⁻¹),
 1124 while on a per m² basis, moderately burned material had greater BC content (20 g BPCA-C
 1125 m^{-2}) than unburned material (17 g BPCA-C m^{-2}). The cumulative difference between
 1126 unburned and moderately burned material was driven by low BC content in the [forest floor](#)
 1127 layer at unburned sites. While the highly burned material did not contain significantly less
 1128 BC than the moderately burned material, it was also not significantly different from
 1129 unburned material, largely driven by cumulative losses from both the [forest floor](#) and 0-5
 1130 cm soil BC stocks. Essentially, the stocks of BC in unburned and highly burned sites are the
 1131 same, they are just distributed differently: the highly burned sites have greater BC stocks in
 1132 [forest floor](#) than soil, and the unburned sites have greater BC stocks in soil than [forest floor](#)
 1133 (Figure 2).

1134 We were initially surprised to find the same amount of BPCA-C in soils from unburned
 1135 and burned sites. The BC in unburned sites must be from prior fires, making up a relatively
 1136 small stock twice the size of BC found in the [forest floor](#) from the HPF. These data suggest

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that eventually a proportion of the BC produced during the HPF will be introduced into the soils and retained in the ecosystem. Given a fire return interval of ~70 years in Ponderosa forests, and a mean residence time for BC stock in soils of approximately 300 years (Hammes et al., 2008; Schmidt et al., 2011), using first order decay, we calculated that 2.4 g m⁻², or 17% of the HPF fire-derived BC in **forest floor** (~14 g m⁻²) would be transferred to the 0-15 cm soil to maintain a steady state stock (~40 g m⁻²). This calculation contains a high degree of uncertainty; a greater residence time **for BC** would result in decreased incorporation, and the reverse for a shorter residence time, **and fires with different properties will deposit different amounts of BC on the soil**.

The estimate for BC incorporation described above is not meant to be used as characteristic value for this ecosystem, but instead **is** meant to illustrate that the bulk of the BC in this system likely moves off **the surface, either through incorporation into deeper soils, biotic or abiotic degradation, or export through erosion. BC incorporation at depth via water flow and biotic infiltration processes stimulated by soil fauna, has been suggested to be the prime mechanism by which BC is sequestered in the soil (Czimczik and Masiello, 2007), although we would have expected to see some increase in the BC content of surface soils in burned sites if incorporation to deep soil was the dominant mechanism. An additional alternative is loss of BC through biotic and abiotic degradation, as a proportion of BC is known to be labile (Zimmerman, 2010), however that proportion is small (Stewart et al., 2013; Zimmerman and Gao, 2013) and other mechanisms are most likely to contribute to major loss pathways. Erosion rates in montane ecosystems post-fire can increase up to three orders of magnitude depending on the severity of the fire and the intensity of precipitation (Wagenbrenner and Robichaud, 2014). Erosion has been shown to be important for BC distribution, as previous work has demonstrated approximately 50% of BC may be lost through erosion processes (Major et al., 2010; Rumpel et al., 2009). While each of these loss mechanisms, degradation, downward translocation and erosion, may be important for BC distribution in the Cache la Poudre (CLP) drainage, preliminary BC data from sediment fences and river banks (Boot et al., 2014), along with a report on dissolved and particulate BC export (Wagner et al., 2015), suggests that erosion may be a dominant source of BC loss in this system.**

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Our third objective was to describe the distribution of BPCAs within forest floor and soil layers to determine if the molecular structure of BC was characteristic by layer, or influenced by burn severity. Recently, Wiedemeir and others validated that the proportion of B6CAs relative to the total BPCAs measured directly correlated with both the degree of condensation and aromaticity of chars, thus we used the relative abundance of B6CA:total BPCA to describe the molecular features of BC (Figure 3). We found that B6CA relative abundance was greater in the 5-15 cm soils relative to forest floor, suggesting that greater more condensed BC, is present in deeper soils at these sites, and there was no effect of burn severity on BPCA abundances. The relative abundance of B6CA has also been associated with the highest heat treatment temperature (HTT) correlating increasing HTT with increasing condensation (Schneider et al., 2013).

Forest fire temperatures are difficult to determine and can range from approximately 1000 °C in the canopy, a maximum of 850 °C at the surface, averaging approximately 300 °C, and rarely exceed 150 °C at 5 cm in the mineral soil (DeBano, 2000; Wolf et al., 2013). While it is tempting to derive HTT of BC deposited on the soil surface following the fire, it must be noted that surface and soil BC is likely a pool integrated across sources pyrolyzed over the range of fire temperatures. The amount of B6CA has been shown to correlate directly with HTT for bark and wood materials, yet no clear relationship exists among B6CA concentrations and temperature of charring for pine needle or leaf derived chars (Schneider et al., 2010; Schneider et al., 2013). Information on HTT from B6CA alone, can be bolstered by also using the ratio of B5CA:B6CA which has a significant inverse linear relationship with combustion temperature. Natural chars range from B5CA:B6CA values 1.3 to 1.9 for cooler burning forest fires (~300 °C), 0.8 to 1.4 for hotter grass and shrub fires (~500 °C), and <0.8 for the hottest burning domestic fires (800 °C) (Wolf et al., 2013). In HPF impacted areas, the forest floor had a B5CA:B6CA ratio of 1.2 which would be at the border between grass/shrub and forest fires and yield an integrated predicted temperature of around 400 °C, whereas the B5CA:B6CA ratio for 5-15 cm soils was significantly lower averaging 0.8, corresponding to higher combustion temperature of approximately 600 °C, which matches well with the temperatures that would be predicted from the B6CA content alone. Other studies have suggested that the pattern of BPCAs may be informative for determining the amount of processing by microorganisms (Rodionov et al. 2010), although

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1408 these correlations have not been empirically validated, and abiotic degradation, such as
 1409 preferential leaching of less condensed forms of BC, would also shift the relative abundance
 1410 of the BPCA pattern (Abiven et al., 2011). Thus the greater B6CA content and decreasing
 1411 B5CA:B6CA ratio in deeper soils from our study may represent either BC derived from
 1412 greater average HTT in past events, selective removal of less condensed forms of BC
 1413 through preferential solubilization (Abiven et al., 2011) or other biotic or abiotic
 1414 degradation of less condensed forms of BC.

1415 5. Conclusion

1416 The distribution of BC on a landscape will influence how an ecosystem recovers
 1417 following a wildfire. Although BC is generally considered nearly biologically inert, its
 1418 impact on soil physical properties may alter biogeochemical cycling. For example, BC
 1419 amendments in agricultural systems (as biochar) have been shown to change water
 1420 holding capacity and nutrient retention (Lehmann, 2007), thus its persistence in post-fire
 1421 soils may be beneficial to, or otherwise alter vegetation recovery dynamics. BC has also
 1422 been shown to enhance growth of microorganisms potentially increasing the accumulation
 1423 of new SOM (Bird et al. 1999). In addition to altering post-fire recovery dynamics, the
 1424 movement of BC following wildfire also has implications for water quality including
 1425 municipal water treatment techniques as well as reductions in primary productivity in
 1426 streams and sediments through increased sediment load (Wood and Armitage, 1997). Our
 1427 results suggest the vast majority of HPF-derived BC deposited on the landscape persisted in
 1428 the forest floor four-months post burn regardless of slope, and was formed at an average
 1429 temperature of 400 °C. Stocks of BC in this montane ecosystem were relatively small and
 1430 were not altered by the HPF thus, subsequent distribution will be governed by modes of BC
 1431 loss likely related to erosion of the forest floor layer, and may also include transport into
 1432 the soils via dissolution and translocation as well as biotic or abiotic degradation.

1433 Author Contribution

1434 MFC and KP designed the experiment, MH coordinated and executed field sampling and
 1435 site characteristic analyses, MH and CMB performed BPCA analyses, CMB prepared the
 1436 manuscript.

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1466 **Acknowledgements**

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1474 Ecology Laboratory at Colorado State University (<http://ecocore.nrel.colostate.edu>).

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1646

1647 **Table 1.** Site characteristics (%C, %N, C:N, C stock) of **forest floor**, 0-5 cm soils and 5-15
1648 cm soil classified by burn **severity** and slope. Mean values reported with standard errors in
1649 parentheses (**n=4**).
1650

layer	burn severity	slope (degrees)	%C	%N	C:N	C stock (tC * ha ⁻¹)
forest floor	unburned	0 to 5	27.31 (7.72)	0.72 (0.18)	37.8 (2.4)	6.6 (3.9)
		5 to 15	23.39 (5.75)	0.73 (0.10)	40.5 (4.2)	5.7 (1.5)
		15 to 30	29.13 (5.81)	0.71 (0.11)	40.2 (2.1)	6.0 (1.2)
	moderate	0 to 5	15.68 (1.44)	0.70 (0.13)	23.5 (2.5)	10.4 (2.4)
		5 to 15	22.35 (7.43)	0.80 (0.15)	26.4 (3.8)	10.9 (3.6)
		15 to 30	17.51 (2.97)	0.62 (0.09)	28.1 (1.8)	6.0 (2.2)
	high	0 to 5	7.07 (2.18)	0.43 (0.13)	16.6 (0.2)	3.8 (1.1)
		5 to 15	13.98 (3.09)	0.68 (0.12)	20.6 (2.1)	7.5 (4.6)
		15 to 30	9.23 (2.31)	0.44 (0.10)	21.4 (3.7)	5.4 (1.9)
0-5 cm soil	unburned	0 to 5	2.89 (1.38)	0.18 (0.06)	14.5 (2.7)	14.1 (6.3)
		5 to 15	2.75 (0.43)	0.17 (0.03)	16.1 (1.1)	15.1 (1.8)
		15 to 30	2.99 (0.56)	0.17 (0.03)	17.9 (1.5)	13.3 (3.2)
	moderate	0 to 5	2.42 (0.38)	0.16 (0.02)	15.0 (0.2)	11.4 (2.1)
		5 to 15	3.62 (0.68)	0.19 (0.04)	19.4 (2.3)	14.1 (2.8)
		15 to 30	3.23 (0.26)	0.17 (0.01)	18.7 (1.6)	17.7 (2.1)
	high	0 to 5	2.09 (0.25)	0.14 (0.02)	14.9 (0.7)	9.0 (1.5)
		5 to 15	2.25 (0.37)	0.14 (0.03)	16.5 (1.6)	11.7 (1.7)
		15 to 30	2.63 (0.24)	0.16 (0.02)	16.4 (1.0)	14.4 (2.9)
5-15 cm soil	unburned	0 to 5	1.82 (1.19)	0.12 (0.06)	13.2 (2.4)	17.8 (9.4)
		5 to 15	1.87 (0.53)	0.11 (0.04)	17.7 (1.6)	25.3 (9.1)
		15 to 30	1.47 (0.50)	0.08 (0.03)	18.5 (2.9)	13.5 (3.8)
	moderate	0 to 5	1.61 (0.35)	0.10 (0.02)	17.0 (2.0)	17.4 (0.3)
		5 to 15	0.97 (0.21)	0.06 (0.01)	16.5 (1.8)	16.3 (8.5)
		15 to 30	1.21 (0.19)	0.08 (0.02)	15.5 (1.2)	15.8 (5.6)
	high	0 to 5	1.29 (0.16)	0.09 (0.01)	15.4 (1.1)	14.3 (5.4)
		5 to 15	1.20 (0.15)	0.08 (0.01)	14.7 (1.2)	13.0 (2.0)
		15 to 30	1.48 (0.32)	0.10 (0.02)	14.9 (0.7)	13.9 (3.3)

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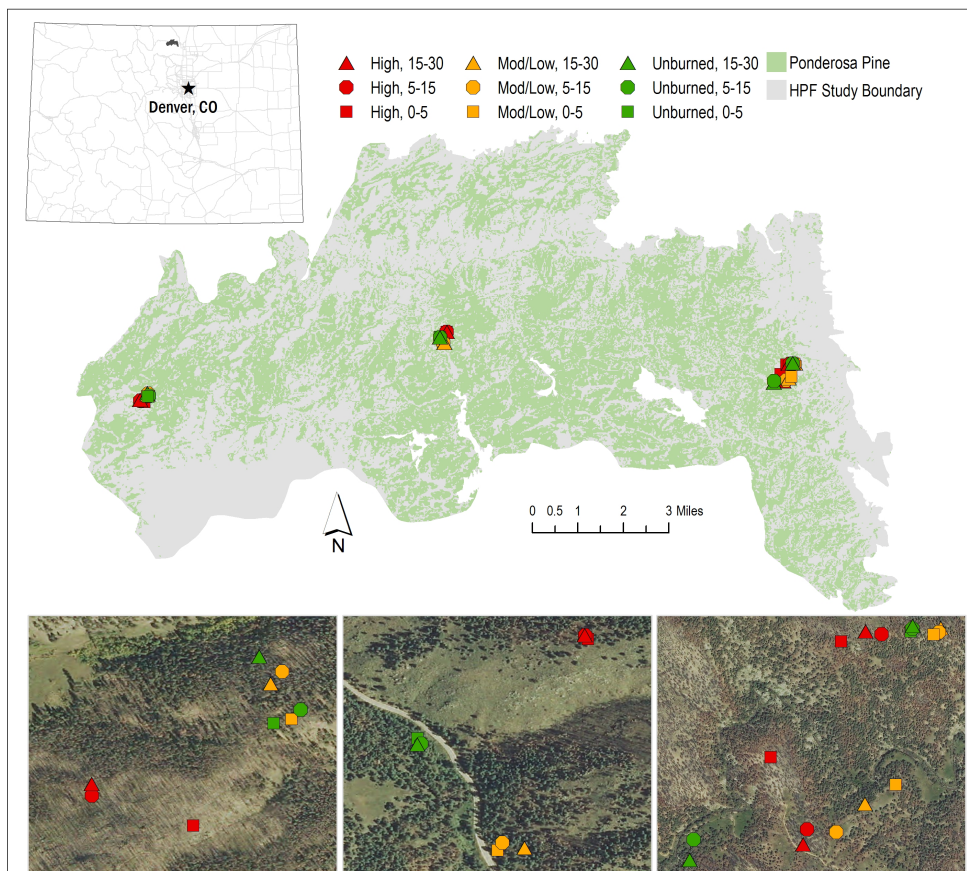


Figure 1. Location and classification (burn severity, and slope) of study sites ($n=36$) in the dominant Ponderosa pine vegetation highlighted in green within the High Park Fire burn area.

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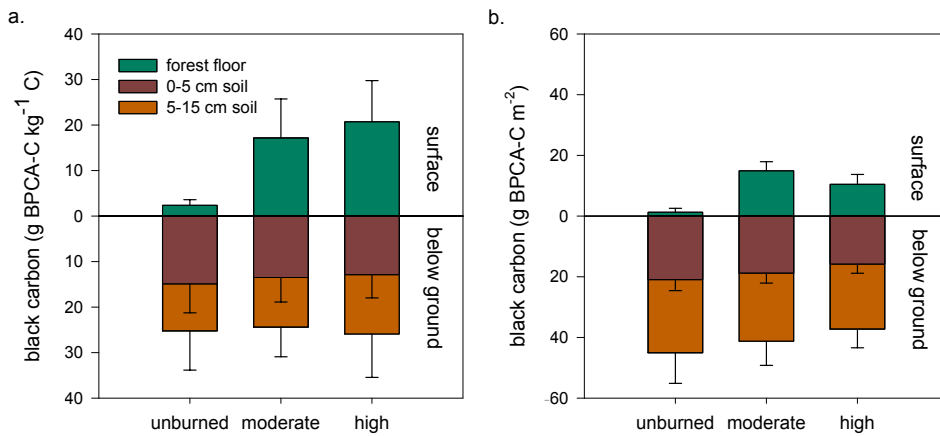


Figure 2. Distribution of black carbon in Ponderosa forest floor in (a.) g BPCA-C kg⁻¹ C, and (b.) g BPCA-C * m⁻² in forest floor (n=12), 0-5 cm (n=12), 5-15 cm (n=12) soils illustrating the significant differences in unburned versus moderately burned forest floor (p<0.001), highly burned forest floor (p<0.001), and unburned forest floor versus 0-5 cm soils (p<0.001) and 5-15 cm soils (p<0.001).

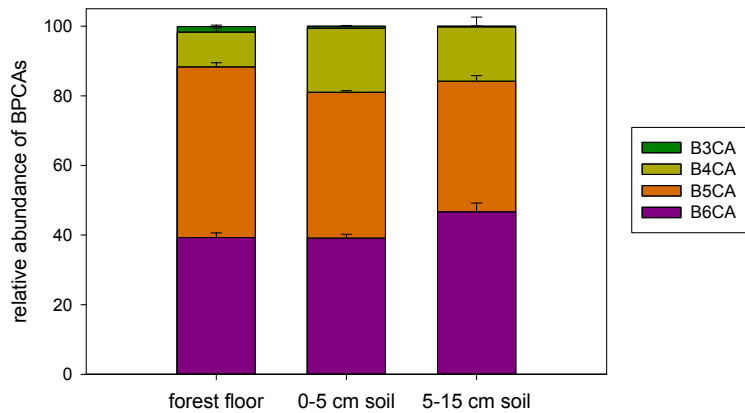
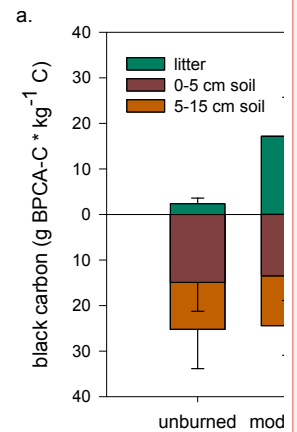


Figure 3. Distribution of BPCAs in each layer (n=36 per layer) illustrating greater relative abundance of B6CA in 5-15 cm soil versus 0-5 cm soil (p <0.001) and forest floor (p<0.001).

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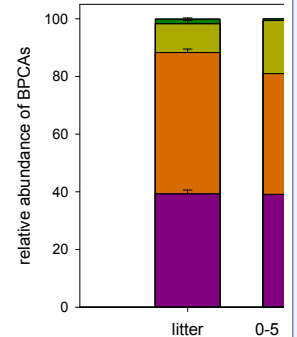


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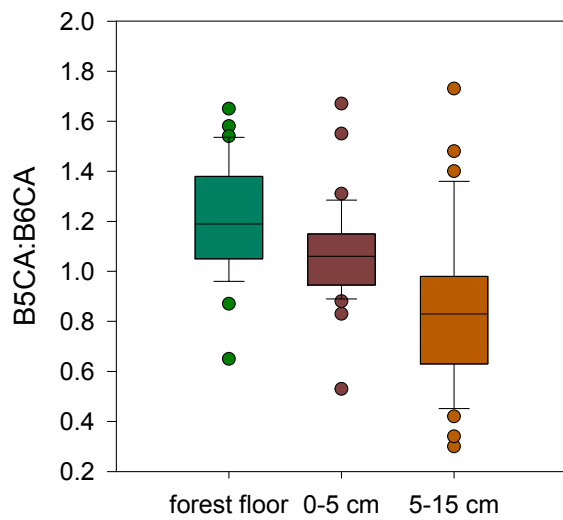


Figure 4. Ratio of B5CA to B6CA from 0.2 to 2.0, illustrating an increasing amount of B6CA and decreasing amount of B5CA with increasing soil depth (n=12 per layer), and significant difference between forest floor ratio and 5-15 cm ratio ($p < 0.001$).

1719 Supplemental Information

1720 Bulk Density Determination

1721 During excavation, if the majority of a rock resided inside the frame that rock was
1722 collected, if the majority of a rock was outside the frame the rock was left in place and the
1723 soil was excavated around the rock. Thin nylon fabric was used to line the pit and millet
1724 was added to the pit until level with the top of the excavation frame. The volume of the
1725 millet was determined with a graduated cylinder. The volume for the 0-5 cm depth was the
1726 volume of the millet that filled the entire 0-5 cm depth, with the volume of the frame
1727 thickness subtracted. The volume of the 5-15 cm depth was the volume of the millet the
1728 filled the entire 0-15 cm depth with the volume of the entire 0-5 cm depth subtracted.

1730 BPCA method validation

1731 The BPCA method was validated by evaluating an unburned field soil from our site that
1732 was mixed with a laboratory produced biochar at 5, 20 and 50% soil weight. The
1733 laboratory biochar was derived from beetle-killed pine sourced from the CLP drainage
1734 combusted at approximately 400 to 500 °C and ground to a fine texture. The BPCA method
1735 validation with biochar-amended soils generated a linear relationship among the samples
1736 ($BPCA-C = -1.506 + (0.89 * \% \text{ char})$, $r^2 = 0.94$, $p < 0.001$). The coefficients of variation at
1737 different amendment levels were 0.05 at 5% amendment, 0.07 at 20% amendment, and
1738 0.19 at 50% amendment. The majority of BPCA-C values in our field study were below the
1739 5% amendment, thus we concluded the method was robust for quantifying BC and
1740 progressed with field-based sample collections.

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1754 **Table S1.** Results of **mixed model** evaluating the effects of layer, burn **severity**, and slope
 1755 on C and N values **and bulk density presented with degrees of freedom (df) and p-values.**

effect	df	log %C	log %N	C:N	log C stock (g C m ⁻²)	bulk density
slope	2	0.306	0.872	0.009	0.774	0.927
burn	2	0.002	0.135	<0.001	0.080	0.796
layer	2	<0.001	<0.001	<0.001	<0.001	0.109
slope x burn	4	0.917	0.746	0.838	0.671	0.691
slope x layer	4	0.077	0.051	0.725	0.672	0.617
burn x layer	4	<0.001	0.005	<0.001	0.284	0.792
slope x burn x layer	8	0.093	0.169	0.685	0.534	0.875

1756
 1757
 1758 **Table S2.** Results of **Tukey's** post-hoc comparisons evaluating the source of significant
 1759 differences in %C, %N, C:N and C stocks by burn **severity**, layer, slope and the interaction
 1760 terms **presented as adjusted p-values.**

post hoc comparisons	log % C	log % N	C:N	log C stock (g C m ⁻²)
high vs. unburned	0.001	0.132	<0.001	0.320
moderate vs. unburned	0.371	0.835	<0.001	0.708
high vs. moderate	0.029	0.318	0.004	0.069
0-5 cm vs. forest floor	<0.001	<0.001	<0.001	<0.001
5-15 cm vs. forest floor	<0.001	<0.001	<0.001	<0.001
0-5 cm vs. 5-15 cm	<0.001	<0.001	0.703	0.696
0-5 vs. 5-15 degrees	0.998	0.941	0.028	0.755
0-5 vs. 15-30 degrees	0.932	0.984	0.012	0.915
5-15 cm vs. 15-30 degress	0.947	0.864	0.916	0.944
forest floor: unburned vs. moderate	0.009	0.968	<0.001	0.217
forest floor: unburned vs. high	<0.001	0.004	<0.001	0.400
forest floor: moderate vs. high	<0.001	0.006	<0.001	0.009
0-5 cm: unburned vs. moderate	0.631	0.978	0.471	0.938
0-5 cm: unburned vs. high	0.947	0.513	1.000	0.805
0-5 cm: moderate vs. high	0.166	0.376	0.448	0.582
5-15 cm: unburned vs. moderate	0.662	0.448	0.999	0.840
5-15 cm: unburned vs. high	0.947	0.999	0.674	0.706
5-15 cm: moderate vs. high	0.836	0.399	0.628	0.968

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response variable	df	f-v
slope	2	1.0
burn	2	7.1
layer	2	27
slope x burn	4	0.1
slope x layer	4	0.8
burn x layer	4	5.1
slope x burn x layer	8	0.0

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post hoc comparisons	p
high vs. unburned	
moderate vs. unburned	
high vs. moderate	
0-5 cm vs litter	
5-15 cm vs litter	
0-5 cm vs 5-15 cm	
0-5 vs 5-15 degrees	
0-5 vs 15-30 degrees	
5-15 cm vs 15-30 degress	
litter: unburned vs. moderate	
litter: unburned vs. high	
litter: moderate vs. high	
0-5 cm: unburned vs. moderate	
0-5 cm: unburned vs. high	
0-5 cm: moderate vs. high	
5-15 cm: unburned vs. moderate	
5-15 cm: unburned vs. high	
5-15 cm: moderate vs. high	

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1773 **Table S3.** Results of mixed model evaluating the effects of layer, burn severity, and slope on
1774 BPCA-C concentration and stocks presented as degrees of freedom (df) and p-values.

effect	df	BPCA C (g kg ⁻¹ C)	log BPCA C stock (g BPCA-C m ⁻²)
slope	2	0.446	0.801
burn	2	0.007	0.012
layer	2	0.610	<0.001
slope x burn	4	0.191	0.853
slope x layer	4	0.740	0.829
burn x layer	4	<0.001	<0.001
slope x burn x layer	8	0.545	0.679

1775
1776 **Table S4.** Results of Tukey's post-hoc comparisons evaluating the source of significant
1777 differences in BPCA-C concentration and stocks by burn severity, layer, and the burn by
1778 layer interaction presented as adjusted p-values.

post hoc comparisons	BPCA C (g BPCA-C kg ⁻¹ C)	log BPCA C stock (g BPCA-C m ⁻²)
high vs. unburned	0.005	0.135
moderate vs. unburned	0.072	0.009
high vs. moderate	0.576	0.408
0-5 cm vs. forest floor	0.980	<0.001
5-15 cm vs. forest floor	0.734	<0.001
0-5 cm vs. 5-15 cm	0.615	0.814
unburned: 0-5 cm vs. forest floor	<0.001	<0.001
unburned: 5-15 cm vs. forest floor	0.004	<0.001
unburned: 0-5 cm vs 5-15 cm	0.711	0.882
moderate: 0-5 cm vs. forest floor	0.418	0.484
moderate: 5-15 cm vs. forest floor	0.085	0.847
moderate: 0-5 cm vs 5-15 cm	0.646	0.817
high: 0-5 cm vs. forest floor	0.023	0.172
high: 5-15 cm vs. forest floor	0.027	0.162
high: 0-5 cm vs. 5-15 cm	0.998	0.999
forest floor: unburned vs. moderate	<0.001	<0.001
forest floor: unburned vs. high	<0.001	<0.001
forest floor: moderate vs. high	0.445	0.330
0-5 cm: unburned vs. moderate	0.874	0.999
0-5 cm: unburned vs. high	0.761	0.696
0-5 cm: moderate vs. high	0.974	0.709
5-15 cm: unburned vs. moderate	0.845	0.989
5-15 cm: unburned vs. high	0.986	0.948
5-15 cm: moderate vs. high	0.742	0.983

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		BPCA C (g C)
response variable	df	f-v
slope	2	
burn	2	
layer	2	
slope x burn	4	
slope x layer	4	
burn x layer	4	
slope x burn x layer	8	

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1788 **Table S5.** Results of mixed model evaluating the effects of burn, layer and their interaction
1789 on the relative abundances of B6CA, B5CA, B4CA and the ratio of B5CA:B6CA presented as
1790 degrees of freedom (df) and adjusted p-values.

effect	df	B6CA	B5CA	B4CA	B5CA:B6CA
burn	2	0.902	0.939	0.401	0.805
layer	2	<0.001	<0.001	<0.001	0.002
burn x layer	4	0.079	0.288	0.543	0.053

1791
1792 **Table S6.** Tukey's Post hoc comparisons among layers and burn by layer interaction for
1793 relative abundance of B6CA, B5CA, B4CA and the ratio of B5CA:B6CA presented as the
1794 adjusted p-values.

post hoc comparisons	B6CA	B5CA	B4CA	B5CA:B6CA
0-5 cm vs. forest floor	0.997	<0.001	<0.001	0.115
5-15 cm vs. forest floor	<0.001	<0.001	0.148	0.001
0-5 cm vs. 5-15 cm	<0.001	0.037	0.002	0.176

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post hoc comparisons
high vs. unburned
moderate vs. unburned
high vs. moderate
0-5 cm vs litter
5-15 cm vs litter
0-5 cm vs 5-15 cm
unburned: 0-5 cm vs litter
unburned: 5-15 cm vs litter
unburned: 0-5 cm vs 5-15 cm
moderate: 0-5 cm vs litter
moderate: 5-15 cm vs litter
moderate: 0-5 cm vs 5-15 cm
high: 0-5 cm vs litter
high: 5-15 cm vs litter
high: 0-5 cm vs 5-15 cm
litter: unburned vs. moderate
litter: unburned vs. high
litter: moderate vs. high
0-5 cm: unburned vs. moderate
0-5 cm: unburned vs. high
0-5 cm: moderate vs. high
5-15 cm: unburned vs. moderate
5-15 cm: unburned vs. high
5-15 cm: moderate vs. high

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response variable	df	f
burn	2	
layer	2	
burn x layer	4	

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post hoc comparisons	diff. r
0-5 cm vs litter	<
5-15 cm vs litter	
0-5 cm vs 5-15 cm	

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