

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

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Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Abstract

Iron (Fe) can limit phytoplankton productivity in approximately 40% of the global ocean, including high-nutrient, low-chlorophyll (HNLC) waters. However, there is little information available on the impact of CO₂-induced seawater acidification on natural phytoplankton assemblages in HNLC regions. We therefore conducted an on-deck experiment manipulating CO₂ and Fe using Fe-deficient Bering Sea waters during the summer of 2009. The concentrations of CO₂ in the incubation bottles were set at 380 and 600 ppm in the non-Fe-added (control) bottles and 180, 380, 600, and 1000 ppm in the Fe-added bottles. The phytoplankton assemblages were primarily composed of diatoms followed by haptophytes in all incubation bottles as estimated by pigment signatures throughout the 7 day incubation period. At the end of incubation, the relative contributions of diatoms to chlorophyll *a* biomass decreased significantly with increased CO₂ levels in the controls, whereas minimal changes were found in the Fe-added treatments. These results indicate that, under Fe-deficient conditions, the growth of diatoms was negatively affected by the increase in CO₂ availability. To confirm this, we estimated the expression and phylogeny of *rbcL* (which encodes the large subunit of RubisCO) mRNA in diatoms by quantitative reverse transcription PCR and clone library techniques, respectively. Interestingly, regardless of Fe availability, the expression and diversity of *rbcL* cDNA decreased in the high CO₂ treatments (600 and 1000 ppm). The present study suggests that the projected future increase in seawater *p*CO₂ could reduce the RubisCO activity of diatoms, resulting in a decrease in primary productivity and a shift in the food web structure of the Bering Sea.

1 Introduction

The atmospheric CO₂ concentration has risen from a pre-industrial level of approximately 280 ppm to the present level of approximately 400 ppm (WMO, 2013). Since the industrial revolution, the ocean has absorbed about one-third of CO₂ emitted

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

by human activity (Sabine et al., 2004). It is predicted that the atmospheric CO₂ concentration could reach more than 700 ppm by the end of the 21st century (Meehl et al., 2007), driving a surface seawater pH decrease of 0.3–0.4, the so-called “ocean acidification” (Caldeira and Wickett, 2003). Such a rapid decrease in seawater pH has most likely not occurred for at least millions of years in the earth’s history (Pearson and Palmer, 2000). Therefore, it has been suggested that these predicted changes in seawater carbonate chemistry would have enormous impacts on the health and function of marine organisms (Raven et al., 2005).

In the last decade, numerous studies have been performed to evaluate the impacts of ocean acidification on marine phytoplankton. In laboratory incubation experiments using individual species (a single strain), the response of phytoplankton to increased CO₂ levels differed among phytoplankton species, possibly depending on their ability to assimilate carbon (Riebesell and Tortell, 2011; Collins et al., 2014). In the natural environment, these taxon-specific differences in CO₂ response can cause a shift in the phytoplankton community composition (Engel et al., 2008; Meakin and Wyman, 2011; Endo et al., 2013) and subsequent changes in ocean trophic structures and biogeochemical cycles (Riebesell et al., 2007; Yoshimura et al., 2013). However, the current understanding of the effects of elevated CO₂ on marine phytoplankton is still insufficient at the community level.

In terms of physiology, CO₂ is fixed by the carboxylation enzyme ribulose biphosphate carboxylase/oxygenase (RubisCO) in the Calvin–Benson–Bascham (CBB) cycle. In general, the half-saturation constant of the enzyme ranges between 20 and 70 μmol kg⁻¹ CO₂ (Badger et al., 1998), whereas the ambient seawater CO₂ levels are between 10 and 25 μmol kg⁻¹. Therefore, the present CO₂ concentration could be insufficient to ensure effective RubisCO carboxylation. Consequently, algae have mechanisms to actively increase the CO₂ concentration at the site of carboxylation (i.e., carbon-concentrating mechanisms, CCMs) to overcome the low RubisCO CO₂ affinity (e.g., Raven, 2010; Reinfelder, 2011). Because the operation of CCMs requires nutrients and energy (Beardall and Giordano, 2002), additional materials

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

would be needed for photosynthetic carbon assimilation under CO₂-limited conditions. Therefore, the progression of ocean acidification could enhance photosynthetic carbon fixation in marine phytoplankton by reducing the investment in CCMs. Until now, most physiological work has been performed in laboratory incubation experiments using particular organisms, but there are few reports on natural phytoplankton assemblages.

Recent advances in molecular biology techniques have enabled us to examine the taxon-specific responses to environmental changes by quantifying functional gene expression in natural phytoplankton assemblages. For example, John et al. (2007a) developed a suite of quantitative reverse transcription PCR (qRT-PCR) assays to quantify *rbcL* (gene encoding the large subunit of RubisCO) mRNA in *Synechococcus*, haptophytes, and heterokonts including diatoms. John et al. (2007b) demonstrated a strong negative correlation between diatom-specific *rbcL* mRNA abundance and seawater *p*CO₂ in the Mississippi River plume, suggesting that diatoms were responsible for the greatest drawdown in seawater *p*CO₂. In addition, positive correlations between diatom-specific *rbcL* transcripts and light-saturated photosynthetic rates (*P*_{max}) in seawater were reported (Corredor et al., 2004; John et al., 2007b). These results suggest that *rbcL* expression in diatoms could be used to estimate the photosynthetic carbon-fixation capacity of natural phytoplankton assemblages. Therefore, quantification of clade-specific *rbcL* transcripts can be used to assess the physiological photosynthetic responses of individual phytoplankton taxa to environmental changes.

The oceanic Bering Sea investigated in this study is an HNLC region (Banse and English, 1999), where low iron (Fe) availability limits phytoplankton growth and nitrate utilization, so surface chlorophyll *a* (Chl *a*) concentrations usually remain low in the summer (Suzuki et al., 2002). Despite the low phytoplankton biomass, the oceanic domain has the greatest amount of total primary and secondary production in the Bering Sea (Springer et al., 1996). Suzuki et al. (2002) reported that diatoms were the dominant phytoplankton group in the oceanic regions of the Bering Sea in the summer. In addition, Takahashi et al. (2002) showed that diatoms had the greatest contribution in

the sinking particles in the area. However, less is known about the combined effects of ocean acidification and Fe enrichment on diatoms in such HNLC regions. Additionally, whether CO₂ level and/or Fe availability could affect *rbcL* transcription in diatoms is completely unknown.

The purpose of this study is to clarify the responses of phytoplankton, especially diatoms, to CO₂ enrichment under Fe-depleted and Fe-replete conditions in the Bering Sea basin using on-deck bottle incubation. Recently, Sugie et al. (2013) reported changes in phytoplankton biomass and nutrient stoichiometry in this experiment. They showed that Chl *a* biomass decreased with increased CO₂ levels only in Fe-depleted treatments, suggesting that Fe deficiency and increased CO₂ synergistically reduced the growth of phytoplankton in the study area. In addition, Yoshimura et al. (2014) demonstrated that the net production of particulate organic carbon (POC) and total organic carbon (TOC) decreased under high CO₂ levels only in the Fe-limited treatments, whereas those in the Fe-replete treatments were insignificantly different. These studies suggest that the increase in CO₂ could have negative impacts on phytoplankton growth and/or organic-matter production especially under Fe-depleted conditions. However, the molecular mechanisms of photosynthetic carbon assimilation in phytoplankton assemblages were not mentioned in the previous studies. Therefore, in the present paper, we primarily focused on changes in *rbcL* transcription in diatoms with different CO₂ and/or Fe availability.

2 Materials and methods

2.1 Experimental setup

The study was carried out aboard the R/V *Hakuho Maru* (JAMSTEC) during the KH-09-4 cruise in September 2009. The water samples for incubation were collected from 10 m depth at a station (53°05' N, 177°00' W) in the Bering Sea on 9 September with acid-cleaned Niskin-X bottles attached to a CTD-CMS system. A total of 300 L of

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



seawater was poured into six 50 L polypropylene carboys through tubing with a 197 μm mesh Teflon net to remove large particles. Subsamples were taken from each carboy and poured into triplicate acid-cleaned 12 L polycarbonate bottles (total 18 bottles) for incubation. Initial samples were collected from each carboy. All sampling was carried out using a trace-metal clean technique to avoid any trace metal contamination. Prior to incubation, FeCl_3 solutions (5 nmol L^{-1} in final concentration) were added to 12 bottles in order to reduce Fe limitation for the phytoplankton communities. The CO_2 levels in the incubation bottles were manipulated by injecting CO_2 controlled dry air purchased from a commercial gas supply company (Nissan-Tanaka Co., Japan). The CO_2 concentrations were set at 380 and 600 ppm for the non-Fe-added (control) bottles (hereafter referred to as “C-380” and “C-600”, respectively), and 180, 380, 600, and 1000 ppm for the Fe-added bottles (hereafter referred to as “Fe-180”, “Fe-380”, “Fe-600”, and “Fe-1000”, respectively). Incubation was performed on deck in temperature-controlled water-circulating tanks for 7 days at the in situ temperature (8.2°C) and 50 % surface irradiance adjusted by natural density screens.

2.2 Carbonate chemistry, nutrients, and Chl *a*

The detailed methodology and basic chemical and biological parameters were reported in Sugie et al. (2013). In brief, during the incubation experiment, samples were collected from the incubation bottles for dissolved inorganic carbon (DIC), total alkalinity (TA), nutrients, and Chl *a* determination. DIC and TA concentrations were measured with a total alkalinity analyzer using the potentiometric Gran plot method (Kimoto Electric) following Edmond (1970). The levels of $p\text{CO}_2$ and pH were calculated from the DIC and TA using the CO2SYS program (Lewis and Wallace, 1998). Concentrations of nitrate plus nitrite, nitrite, phosphate, and silicic acid were measured using a QuAATro-2 continuous-flow analyzer (Bran + Luebbe). The concentration of total dissolved Fe (TD-Fe) was determined by a flow-injection method with chemiluminescence detection (Obata et al., 1993). Chl *a* concentrations were

BGD

11, 18105–18143, 2014

Effects of CO_2 and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

determined with a Turner Design fluorometer (model 10-AU) with the non-acidification method (Welschmeyer, 1994).

2.3 HPLC and CHEMTAX analyses

Samples for high-performance liquid chromatography (HPLC) pigment analysis were collected on days 3 and 5 for the control treatments and on days 2, 4, and 6 for the Fe-added treatments. Water samples (400–1000 mL) were filtered onto GF/F filters under gentle vacuum (< 0.013 MPa) and stored in liquid nitrogen or a deep freezer (–80 °C) until analysis. HPLC pigment analysis was performed following the method of Endo et al. (2013).

To estimate the temporal changes in phytoplankton community structure during incubation, the CHEMTAX program (MacKey et al., 1996) was used following Endo et al. (2013). Briefly, optimal initial ratios were obtained following the method of Latasa (2007). Matrix **A** was obtained from Suzuki et al. (2002) (Table S1 in the Supplement), who examined phytoplankton community compositions in the Bering Sea. Matrices **B**, **C**, and **D** were also prepared to determine the optimal pigment/Chl *a* ratios (Table S1). The pigment ratios of matrices **B** and **C** were double and half the matrix **A** ratio, respectively. For matrix **D**, values of 0.75, 0.5, and 0.25 for dominant (rank in high pigment/Chl *a* ratio: 1–5), secondary (rank: 6–10), and minor (rank: 11–15) pigments, respectively, were multiplied by each pigment ratio of matrix **A**. We averaged the successive convergent ratios after the 10 runs among the 4 matrices to identify the most promising initial pigment ratios. The calculated final pigment/Chl *a* ratios in both the control and Fe-added treatments (Table S2) were within the range of values reported in Mackey et al. (1996), Wright and van den Enden (2000), and Suzuki et al. (2002).

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

2.4 qPCR and qRT-PCR

Water samples for DNA and RNA analyses were collected on days 3 and 5 for the control treatments and on days 2, 4, and 6 for the Fe-added treatments. DNA samples (400–500 mL) were collected onto 25 mm, 0.2 µm pore size polycarbonate Nuclepore filters (Whatman) with gentle vacuum (< 0.013 MPa) and stored in liquid nitrogen or a deep freezer at –80 °C until analysis. DNA extraction was performed following the method of Endo et al. (2013). Extracted DNA pellets were resuspended in 100 µL of 10 mM Tris-HCl buffer (pH 8.5).

For RNA analysis, seawater samples (400–500 mL) were filtered onto 25 mm, 0.2 µm pore size polycarbonate Nuclepore filters (Whatman) with gentle vacuum (< 0.013 MPa) and stored in 1.5 mL cryotubes previously filled with 0.2 g of muffled 0.1 mm glass beads and 600 µL RLT buffer (Qiagen) with 10 µL mL⁻¹ β-mercaptoethanol (Sigma, St Louis, USA). RNA samples were stored in liquid nitrogen or a deep freezer at –80 °C until analysis. Extraction and purification of RNA samples were performed using the RNeasy extraction kit (Qiagen) on a vacuum manifold with on-column DNA digestion using RNase-free DNase (Qiagen) according to the manufacturer's protocol. RNA was eluted using 50 µL of RNase-free H₂O. Total RNA was then reverse transcribed into complementary DNA (cDNA) using the PrimeScriptTM RT reagent Kit with gDNA Eraser (TaKaRa) following the manufacturer's specifications.

Standard curves for *rbcL* DNA were generated from plasmid DNA (pUC18, TaKaRa) containing an artificial gene fragment (113 bp in size) of *rbcL* from the diatom *Thalassiosira weissflogii* (CCMP1336). The plasmid DNA was linearized with *Hind*III (TaKaRa) and quantified using a Thermo NanoDrop spectrophotometer (ND-1000).

To produce a cDNA standard, a PCR-amplified *rbcL* gene fragment of *T. weissflogii* (CCMP1336) was inserted into a plasmid DNA (pCR2.1, Invitrogen). The plasmid DNA was purified using the Plasmid maxi kit (Qiagen) and linearized with *Bam*HI (TaKaRa), and in vitro transcription was performed using T7 RNA polymerase (Invitrogen) for 2 h at 37 °C with Recombinant RNase Inhibitor (TaKaRa). To eliminate DNA

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

contamination, RNA was digested for 2 min at 42 °C using gDNA Eraser (TaKaRa). RNA was purified using an RNeasy column (Qiagen) following the manufacturer's instructions and quantified with a Ribogreen RNA quantification kit (Molecular Probes) using the manufacturer's standard. RNA was reverse transcribed into cDNA using the PrimeScript™ RT reagent Kit with gDNA Eraser (TaKaRa).

Copy numbers of DNA and cDNA standards were calculated using the equation of Smith et al. (2006), where the molecular mass of each nucleotide (or nucleotide pair) in double- and single-stranded DNA is assumed to be 660 and 330 Da, respectively. Serial dilutions of DNA and cDNA standards were prepared using sterilized Milli-Q water.

To amplify the *rbcL* gene and cDNA fragments from diatoms, the following specific primer set designed by John et al. (2007a) was used: forward primer: 5'-GATGATGARAAYATTAAGTC-3', reverse primer: 5'-TAWGAACCTTTWACTTCWCC-3'. Real-time PCR amplification was performed using SYBR Premix Ex Taq II (Perfect Real Time, TaKaRa) with primer concentrations of 0.4 μM each and a Thermal Cycler Dice Real Time System (TP800, TaKaRa). Diluted nucleic acid standards were then added to the PCR mixture. The thermal cycling conditions were 95 °C for 60 s, then 40 cycles of 95 °C 5 s and 52 °C 60 s. The fluorescence intensity of the complex formed by SYBR green and the double-stranded PCR product was continuously monitored from cycle 1–40. Quantification was achieved by the second-derivative maximum method (Luu-The et al., 2005), and the copy number for each sample was determined by the standard curves generated by serial dilutions of the standards.

2.5 Clone libraries

Clone libraries of *rbcL* cDNA were constructed for the C-380 and C-600 samples on day 3, and Fe-380 and Fe-600 samples on day 2. The cDNA samples were PCR amplified with the diatom-specific primer set and thermal cycling condition described above using the TaKaRa Ex Taq Hot Start Version (TaKaRa). Triplicate PCR products were mixed and then purified with agarose gel electrophoresis and the

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



PureLink Quick Gel Extraction Kit (Invitrogen). Purified amplicons from cDNA samples were then cloned into the pCR2.1 vector using the TOPO TA cloning kit (Invitrogen) following the manufacturer's instructions. Thirty-five to 50 colonies were randomly picked from each clone library. Correct cDNA insertions were identified by PCR amplification using the M13 forward (5'-GTAAACGACGGCCAG-3') and reverse (5'-CAGGAAACAGCTATGA-3') primers flanking the cloning site. Plasmid DNA containing the inserts was cycle-sequenced using the Big Dye Terminator v3.1 Kit (Applied Biosystems) with the M13 forward primer. The cycle sequencing products were cleaned by isopropanol precipitation. Sequencing was performed with a 3130 Genetic Analyzer (Life Technologies). The obtained sequences were compared with *rbcL* sequences deposited in GenBank database (<http://www.ncbi.nlm.nih.gov>) using the BLAST query engine. Our *rbcL* cDNA sequences were deposited in the DDBJ database with the following accession numbers: AB985799–AB986033.

2.6 Phylogenetic and diversity analyses

The *rbcL* sequences obtained were assembled into operational taxonomic units (OTUs) with > 95% sequence identity, and rarefaction curves were plotted for each clone library with the software mothur v. 1.27 (Schloss et al., 2009). Genetic diversity was assessed based on the number of OTUs, the Shannon–Wiener index (H' , Shannon, 1948), and Simpson's index ($1-D$, Simpson, 1949). The statistical significance of differences in the compositions of pairs of *rbcL* sequences in the libraries was tested using LIBSHUFF (Singleton et al., 2001). The LIBSHUFF program determined the integral form of the Cramer–von Mises statistic for each pair of communities using 10 000 randomizations. Any two libraries were considered to be significantly different from each other if the lower of the significance values generated by the software was < 0.025 ($p < 0.05$).

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



2.7 Statistical analysis

Statistical analyses were performed with the program R (<http://www.r-project.org>). To assess the statistically significant differences between $p\text{CO}_2$ levels in the control treatments or between control and Fe treatments, Welch's t test was used. Differences among $p\text{CO}_2$ levels in the Fe-added treatments were evaluated with Kruskal–Wallis one-way analysis of variance (ANOVA). Holm's test for multiple comparisons was used to identify the source of the variance. For all of the analyses, the confidence level was set at 95 % ($p < 0.05$).

3 Results

3.1 Experimental conditions

Carbonate chemistry, macronutrients, and Fe parameters (value ± 1 SD, $n = 1$ or 3) during the incubation experiment are shown in Table 1, and the complete data set was reported in Sugie et al. (2013). The bubbling of CO_2 -controlled air succeeded in creating significant gradients in $p\text{CO}_2$, pH, and DIC in the different CO_2 treatments except on day 4 in the Fe-added treatments, when those values did not significantly differ between Fe-380 and Fe-600. The initial concentrations of nitrate, phosphate, and silicic acid were 18.06 ± 0.10 , 1.47 ± 0.01 , and $16.90 \pm 0.12 \mu\text{mol L}^{-1}$, respectively. In the control bottles, these macronutrients remained until the end of the incubation in both CO_2 treatments except for silicic acid, which was almost depleted on day 5 in the C-380 treatment. In the Fe-added bottles, macronutrients were depleted on days 4 or 6 in all CO_2 treatments. The TD-Fe concentration was 1.35 nmol L^{-1} in the initial seawater, and it remained low throughout the experiment in the control treatments. In the Fe-added treatments, the TD-Fe concentrations were $5.50 \pm 0.10 \text{ nmol L}^{-1}$ in the initial bottles and remained above 4 nmol L^{-1} until the end of incubation. The initial Chl a concentration was $1.96 \pm 0.14 \mu\text{g L}^{-1}$. In the control bottles, the Chl a concentration

BGD

11, 18105–18143, 2014

Effects of CO_2 and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

increased until the end of the incubation and reached $10.22 \pm 0.89 \mu\text{gL}^{-1}$ in the C-380 and $6.28 \pm 0.64 \mu\text{gL}^{-1}$ in the C-600 treatments. In the Fe-added bottles, the Chl *a* concentration increased rapidly and reached the maximum on day 4 in the Fe-180 and Fe-380 treatments (27.51 ± 0.71 and $28.45 \pm 3.40 \mu\text{gL}^{-1}$, respectively) and on day 5 in the Fe-600 and Fe-1000 treatments (27.68 ± 0.44 and $27.32 \pm 3.05 \mu\text{gL}^{-1}$, respectively), then declined toward the end of the incubation.

3.2 Phytoplankton pigments

Throughout the experiment, the concentrations of fucoxanthin (Fucox), mainly a biomarker for diatoms (Ondrusek et al., 1991; Suzuki et al., 2011), and 19'-hexanoyloxyfucoxanthin (19'-Hex), an indicator of haptophytes (Jeffrey and Wright, 1994), were relatively high among the phytoplankton pigments. Increased levels of Fucox concentration during the experiment were the highest among the phytoplankton pigments in all Fe and CO₂ treatments. The ratios of Fucox concentration on the final sampling days (days 5 or 6) relative to the initial decreased with an increase in CO₂ levels in both the control bottles (Welch's *t* test C-380 > C-600, $p < 0.05$) and Fe-added bottles (Holm's test Fe-180 > Fe-1000, $p < 0.05$) (Fig. 1a). The concentration of 19'-Hex also increased from initial to the final sampling day except for the Fe-1000 treatment (Fig. 1b). The ratios of 19'-Hex concentration on the final days relative to initial decreased in response to the increased CO₂ levels in both the control (Welch's *t* test C-380 > C-600, $p < 0.05$) and Fe-added bottles (Holm's test Fe-180 > Fe-600 and Fe-1000, $p < 0.05$) (Fig. 1b). Compared to the control treatments, the change in the ratio of Fucox levels was higher in the Fe-added bottles across all CO₂ treatments (Welch's *t* test, $p < 0.05$) (Fig. 1a). In contrast, the ratios of 19'-Hex concentration between the final and initial sampling days were significantly higher in the control bottles than those in the Fe-added bottles (Welch's *t* test, $p < 0.05$) (Fig. 1b).

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

3.3 CHEMTAX outputs

In the initial phytoplankton community, diatoms and haptophytes were the predominant numbers of the phytoplankton groups (i.e., they contributed 45 and 17% of the Chl *a* concentration, respectively). The initial contributions of chlorophytes, cryptophytes, peridinin-containing dinoflagellates, pelagophytes, prasinophytes, and cyanobacteria to the Chl *a* biomass were 10, 9, 8, 5, 4, and 2%, respectively. In the control bottles, the contributions of diatoms to the Chl *a* biomass increased with time, and their contributions reached the maximum (70% at the C-380 and 60% at the C-600 treatments) on day 5 (Fig. 2a). On day 5, the contribution of diatoms in the C-380 treatment was significantly higher than that in the C-600 treatment (Welch's *t* test, $p < 0.05$). However, the contribution of haptophytes to the Chl *a* biomass was higher in the C-600 treatment (21%) than in the C-380 treatment (14%) on day 5 (Welch's *t* test, $p < 0.05$). Increases in the contributions of diatoms were also observed in the Fe-added treatment, and the contributions reached the maximum (82–85%) on day 4 in all CO₂ treatments (Fig. 2b). In terms of diatom contribution, a significant difference among CO₂ treatments was not detected with Kruskal–Wallis ANOVA ($p > 0.05$) in the Fe-added bottles. The contributions of haptophytes to Chl *a* biomass did not differ significantly among CO₂ levels in the Fe-added bottles (Kruskal–Wallis ANOVA, $p > 0.05$).

3.4 Expression of diatom *rbcL* gene

A significant linear relationship between the Fucox concentration and the diatom-specific *rbcL* gene copy number was found (regression analysis: $r^2 = 0.677$, $p < 0.001$, $n = 28$) in our experiment (Fig. 3). In the control bottles, the transcript abundance normalized to gene abundance (i.e., cDNA/DNA) of the diatom-specific *rbcL* gene fragment for the C-380 treatment was significantly higher than that of the C-600 treatment on day 3 (Fig. 4; Welch's *t* test, $p < 0.05$). In the Fe-added bottles, the cDNA/DNA ratio of the diatom *rbcL* fragment in the lower CO₂ treatments (Fe-180

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



assemblages. Our HPLC and CHEMTAX results suggested that the increase in phytoplankton biomass was mainly due to an increase in diatoms (Figs. 1a and 2b).

We found that the ratios of Fucox concentration between initial and the final days decreased with increased CO₂ levels in both the control and Fe-added treatments (Fig. 1a), suggesting that the elevated CO₂ levels had a negative impact on the diatom biomass in the study area. Negative effects on diatoms induced by an increase in CO₂ availability were also reported in field incubation experiments conducted in the Bering Sea and the Okhotsk Sea (Hare et al., 2007 and Yoshimura et al., 2010, respectively). However, such trends have rarely been observed in other regions of the world's oceans (e.g., Tortell et al., 2002; Kim et al., 2006; Feng et al., 2009; Hoppe et al., 2013; Endo et al., 2013). Therefore, the responses of phytoplankton assemblages to ocean acidification can differ among geographic locations due to the differences in the biogeography of phytoplankton and/or environmental conditions. Our study indicates that the decrease in diatom biomass given elevated CO₂ levels was unique to the Bering Sea basin.

One possible cause of the geographic specificity in the open Bering Sea is the differences in the species composition of diatoms. Our microscope data showed that centric diatoms such as *Chaetocera* and *Rhizosolenia* were predominant at the beginning of the incubation in terms of carbon biomass, and the coastal diatom species *Chaetoceros* spp. became predominant in all incubation bottles after day 2 (Sugie et al., 2013). Therefore, the decrease in Fucox biomass with increased CO₂ levels might be partially explained by the decrease in *Chaetoceros* spp. A previous field incubation experiment conducted in the Bering Sea also showed that the carbon biomass of the *Chaetoceros* spp. decreased at higher CO₂ levels (600–960 μatm CO₂), although it increased at 1190 μatm CO₂ (Yoshimura et al., 2013). In the previous laboratory culture experiments, however, the effects of increased CO₂ on the growth and/or photosynthesis of *Chaetoceros* spp. were inconsistent. For example, Ihnken et al. (2011) demonstrated that the growth of diatom *Chaetoceros muelleri* decreased with elevated CO₂ and decreased pH levels although photosynthetic capacity

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



increased. In addition, no CO₂-related change in the growth and photosynthetic physiology of *Chaetoceros brevis* was found (Boelen et al., 2011). These results suggest that the responses to elevated CO₂ differ among *Chaetoceros* species.

The decrease in diatom biomass at elevated CO₂ levels may have been induced by decreased CCM activity of diatom cells. A recent study demonstrated that elevated CO₂ levels could decrease the growth of diatoms under high-light conditions, likely due to an increase in photoinhibition (Gao et al., 2012). The authors suggested that CCMs act as a diffusive process for photoinhibition in high-light conditions, and therefore, diatoms tend to be more vulnerable to light stress due to the downregulation of CCM activities at higher CO₂ levels.

The ratios of 19'-Hex concentration between initial and the final days decreased in response to increased CO₂ levels regardless of Fe availability (Fig. 1b), suggesting that ocean acidification can induce negative effects on not only the biomass of diatoms but also that of haptophytes in the study area. Similar results were obtained from the previous field studies in other regions (e.g., Feng et al., 2010; Endo et al., 2013). One possible factor underlying these decreases is that the reduced carbonate-saturation states under high CO₂ conditions. The energetic cost of calcification in coccolithophores will increase with a decrease in pH (Mackinder et al., 2010). Therefore, additional energy might be needed for cell growth in seawater with high CO₂ levels. In addition, non-calcifying haptophytes such as *Phaeocystis* spp. often dominate among haptophytes in the natural phytoplankton community (Schoemann et al., 2005), although the effects of ocean acidification on them are still not well understood. Therefore, additional study using a wide range of haptophyte species would be required for a detailed understanding of the responses of the haptophyte community to CO₂-induced ocean acidification.

Our CHEMTAX outputs showed that the relative contributions of diatoms decreased with increased CO₂ levels, whereas the contributions of haptophytes increased in both the control and Fe-added bottles (Fig. 2). This indicates that the negative impacts of increased CO₂ on diatoms were greater than those on haptophytes and other

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



phytoplankton groups. Larger diatoms can contribute to efficient transfer of energy and organic compounds to higher trophic levels because they would create a shorter food chain compared with nano- and pico-sized phytoplankton (Michaels and Silver, 1988). Because diatoms form a large part of phytoplankton biomass in the Bering Sea basin (Suzuki et al., 2002; Takahashi et al., 2002), the decrease in the relative contribution of diatoms with increasing CO₂ could reduce the energy transferred from the primary producers to the higher trophic levels.

The decreases in Fucox growth ratio and relative contribution of diatoms were larger in the control bottles than those in the Fe-added treatments (Figs. 1a and 2), suggesting that the negative effect of CO₂ enrichment was greater in the Fe-limited conditions. These results are consistent with Sugie et al. (2013) and Yoshimura et al. (2014), who observed significant decreases in diatom carbon biomass and particulate organic carbon (POC) production under high CO₂ levels in the control treatments, whereas those were insignificantly changed in the Fe-added treatments. Sugie et al. (2013) indicated that the Fe limitations for phytoplankton in the control bottles were enhanced at high CO₂ levels, likely due to the reduction of Fe bioavailability as reported in Shi et al. (2010). The combined effects of CO₂ and Fe availability were also tested in a diatom-dominated phytoplankton community in the Southern Ocean (Hoppe et al., 2013). In their study, net primary productivity in seawater decreased with increased pCO₂ levels in the Fe-depleted treatments but not in the Fe-enriched treatments. These studies indicate that an interactive effect of CO₂ enrichment and Fe limitation could occur in the diatom-dominated natural phytoplankton assemblages in the HNLC region.

4.2 *rbcL* expression in diatom

A significant correlation between *rbcL* copy number in diatoms and Fucox concentration was found in this study (Fig. 3), suggesting the usefulness of the *rbcL* gene fragment as a proxy for diatoms as well as Fucox. In addition, the cDNA sequences obtained from cloning were dominated by the diatom-derived *rbcL* gene (Fig. 6). These results indicate that the *rbcL* primers used successfully and selectively amplified the *rbcL*

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



gene of diatoms. Suzuki et al. (2011) showed that Fucox concentration significantly correlated with diatom carbon biomass in the subarctic Pacific. Furthermore, Matsuda et al. (2011) showed that the number of *rbcL* gene per cell varies among diatom species, and it was positively correlated with cell size. Therefore, we concluded that the *rbcL* gene could serve as a potential molecular marker for diatom biomass.

The transcription of the diatom-specific *rbcL* gene decreased with elevated CO₂ levels in both the control and Fe-added treatments (Fig. 4). Because RubisCO expression is primarily controlled at the transcriptional level in the natural phytoplankton community (Xu and Tabita, 1996; Wawrik et al., 2002), our results suggest that increased CO₂ levels could reduce the amount of RubisCO in diatoms. It should be noted that significant decreases in *rbcL* expression with increased CO₂ levels were observed on days 2 or 3, when macronutrients still remained (Sugie et al., 2013). This indicates that the downregulation of *rbcL* expression in diatoms was probably caused by the increase in CO₂ availability. It has been shown that some land plants can increase their nitrogen utilization efficiency under elevated CO₂ levels by reducing the investment of nitrogen in RubisCO (Curtis et al., 1989; Makino et al., 2003). Losh et al. (2012, 2013) also demonstrated a decreased RubisCO contribution to the total protein in the California Current phytoplankton community with an increase in CO₂ level. Because a decrease in the expression of RubisCO can result in a reduction of the potential capacity for carbon fixation in the natural environment (John et al., 2007b), our results indicate that an increase in CO₂ levels could have a negative impact on photosynthetic carbon fixation for diatoms in the study area.

The negative effects of increasing CO₂ on diatom biomass were not severe in the Fe-added bottles relative to Fe-limited control bottles (Fig. 1), whereas *rbcL* transcription decreased with increased CO₂ regardless of Fe availability (Fig. 4). This suggests that the diatoms could overcome the decrease in RubisCO activity in the Fe-added treatments potentially by induction of CCMs given that diatoms can upregulate CCM activity at elevated CO₂ levels (e.g., Trimborn et al., 2009). In the Fe-added treatments, photosynthetic carbon fixation in diatoms could not be limited by CO₂

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

availability as a consequence of the CCMs. This speculation is supported by the values of the maximum photochemical quantum yield (F_v/F_m) of photosystem II (PSII). In our incubation experiments, F_v/F_m values increased significantly with Fe enrichment (Sugie et al., 2013), suggesting that phytoplankton could produce sufficient energy for CCMs in the Fe-added bottles. However, we speculate that CCMs in the diatoms might not be active in the control treatments because Fe deficiency could reduce the functionality of algal CCMs due to a reduction in their light energy-harvesting ability (Giordano et al., 2005).

It is generally recognized that phytoplankton autonomously regulate the transcription of the *rbcL* gene in response to environmental conditions such as light and nutrient availability (Pichard et al., 1996; Granum et al., 2009; John et al., 2010). However, the mechanisms controlling the transcription of RubisCO operon in diatoms are largely unknown. Recently, Minoda et al. (2010) showed that the red alga *Cyanidioschyzon merolae* increased *rbcL* transcription at high levels of NADPH, 3-phosphoglyceric acid (3-PGA), or ribulose-1,5-bisphosphate (RuBP) under the influence of the transcription factor Ycf30. In addition, it has been reported that regeneration of RuBP could be a limiting factor for the CBB cycle in high CO₂ conditions (von Caemmerer and Farquhar, 1981; Stitt, 1991; Onoda et al., 2005). Thus, one possible mechanism underlying the reduction of diatom *rbcL* transcription observed in our study is related to a decrease in RuBP concentration in the chloroplasts due to the increase in CO₂ availability for diatoms. Because diatoms possess the same type of RubisCO (Form ID) and gene homologs encoding the Ycf30 protein (i.e., *ycf30*) (Kowallik et al., 1995), they could control *rbcL* gene expression using the same mechanisms as *C. merolae*. Further studies using marine diatom cultures are required to obtain a better understanding of the physiological mechanisms controlling the expression of RubisCO.

In our experiment, the rarefaction curves plateaued to some extent in all treatments (Fig. 5), indicating that the clone numbers screened from each library were statistically sufficient for further diversity analysis. Taxonomic compositions in the cDNA library were considerably different from those in the diatom carbon biomass revealed by

microscopic analysis by Sugie et al. (2013), which were composed primarily by *Chaetocerataceae*. This implies that the predominant diatoms did not necessarily become transcriptionally active *rbcL* phylotypes in our experiment. In addition, because 16–42 % of the sequences were classified as unidentified diatoms or other eukaryotes, the primer set used in this study might be insufficient to estimate diatom composition at the family level.

The *rbcL* cDNA libraries in the Fe-added treatments differed significantly from the initial library, whereas those in the control treatments were not significantly different (Table 3), suggesting that the diatom blooms induced by Fe infusion were associated with the change in the relative contribution of *rbcL* expression in diatoms. For example, compared to the initial seawater, the relative contributions of *Chaetocerataceae* and unidentified centrics to the *rbcL* cDNA library increased markedly in the Fe-added bottles whereas they remained minor components in the control bottles (Fig. 6). This indicates that the relative significance of the RubisCO activity of these phylotypes could be increased by Fe enrichment. In addition, cDNA libraries were significantly different from each other at different CO₂ levels in the Fe-added bottles (Table 3). This indicates that the transcriptionally active phylotypes in diatoms could shift in response to an increase in the CO₂ level.

The diversity of the diatom-specific *rbcL* cDNA sequences decreased with increased CO₂ levels in terms of number of OTUs and the Shannon–Wiener and Simpson’s indices (Table 2). These results suggest that the number of diatom phylotypes actively transcribing the *rbcL* gene decreased with increased CO₂ level. Because the different types of diatoms could have different functions in relation to sinking flux and trophic position in the ecosystem (e.g., Michaels and Silver, 1988), a decrease in diatom biodiversity could cause a decrease in the functional diversity of marine ecosystems. Furthermore, a decrease in biodiversity may weaken the ecological stability of phytoplankton ecosystems and make them more vulnerable to other environmental changes such as an increase in sea surface temperature (Overland and Stabeno, 2004).

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

5 Conclusions

The present study showed that an increase in CO₂ levels could have negative impacts on diatom biomass in the Bering Sea, especially under Fe-limited conditions. Because diatoms play pivotal roles in carbon sequestration and food webs in the Bering Sea (Springer et al., 1996; Takahashi et al., 2002), our results indicate that ocean acidification might alter the biogeochemical processes and ecological dynamics in the study area. Although the present results cannot be extrapolated to other HNLC ecosystems due to differences in other environmental conditions, our findings suggest that the combined effects of CO₂ and other environmental factors such as Fe availability need to be examined for a better understanding of the potential impacts of ocean acidification on marine ecosystems.

We examined, for the first time, the relationships between CO₂ levels or Fe availability and RubisCO expression of diatoms in the Bering Sea. Significant decreases in the *rbcL* expression of diatoms were observed at elevated CO₂ levels in both the Fe-limited and Fe-enriched treatments, suggesting that ocean acidification could reduce the primary productivity in the study area. Our results indicate that the amount of *rbcL* transcripts could be an important indicator to assess the physiological responses of RubisCO activity in diatoms to environmental drivers. However, because carbon fixation in diatoms is controlled not only by RubisCO activity but also by CCMs (Rost et al., 2003), more detailed studies of molecular mechanisms are required to clarify the physiological responses of the diatom community to CO₂ and Fe enrichments.

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)




[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)

[◀](#)

[▶](#)

[◀](#)

[▶](#)

[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)




[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

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Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Table 1. Carbonate chemistry, nutrients, and Fe parameters (value ± 1 SD, $n = 1$ or 3) during the incubation experiment. Carbonate parameters are the initial and mean values throughout the incubation. Macronutrients and Fe parameters are the values at the initial or final sampling days. See Sugie et al. (2013) for the complete data set.

	DIC ($\mu\text{mol kg}^{-1}$)	TA ($\mu\text{mol kg}^{-1}$)	$p\text{CO}_2$ (μatm)	CO ₂ ($\mu\text{mol kg}^{-1}$)	pH (total scale)
C-Initial	2086.4 \pm 2.8	2249.1 \pm 5.0	388.4 \pm 18.1	18.4 \pm 0.9	8.05 \pm 0.02
C-380	2075.5 \pm 8.1	2252.9 \pm 10.8	355.7 \pm 34.7	16.8 \pm 1.6	8.09 \pm 0.04
C-600	2151.6 \pm 7.8	2250.9 \pm 4.7	604.1 \pm 36.2	28.5 \pm 1.7	7.88 \pm 0.02
Fe-Initial	2085.3 \pm 0.8	2250.0 \pm 4.9	383.4 \pm 12.6	18.1 \pm 0.6	8.06 \pm 0.01
Fe-180	1959.9 \pm 62.0	2244.1 \pm 16.0	202.0 \pm 50.9	9.5 \pm 2.4	8.21 \pm 0.10
Fe-380	2068.5 \pm 27.7	2235.7 \pm 14.9	375.9 \pm 47.9	17.8 \pm 2.3	8.01 \pm 0.05
Fe-600	2120.6 \pm 33.5	2248.5 \pm 12.0	512.6 \pm 135.5	24.2 \pm 6.4	7.96 \pm 0.11
Fe-1000	2200.2 \pm 12.6	2248.4 \pm 9.8	913.8 \pm 159.8	43.2 \pm 7.6	7.72 \pm 0.07

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Table 1. Continued.

	Nitrate	Phosphate ($\mu\text{mol L}^{-1}$)	Silicic acid	TD-Fe nmolL ⁻¹
C-Initial	18.06 ± 0.10	1.47 ± 0.01	16.95 ± 0.12	1.35
C-380	7.09 ± 0.27	0.65 ± 0.02	0.28 ± 0.05	0.27 ± 0.03
C-600	12.01 ± 0.27	0.98 ± 0.02	3.04 ± 0.32	0.29 ± 0.04
Fe-Initial	18.09 ± 0.11	1.47 ± 0.01	16.90 ± 0.12	5.50 ± 0.10
Fe-180	0.13 ± 0.04	0.10 ± 0.01	0.66 ± 0.09	4.60 ± 0.19
Fe-380	0.09 ± 0.00	0.12 ± 0.04	0.50 ± 0.01	4.48 ± 0.12
Fe-600	0.08 ± 0.00	0.10 ± 0.00	0.50 ± 0.01	4.34 ± 0.08
Fe-1000	0.08 ± 0.00	0.08 ± 0.02	0.47 ± 0.02	4.18 ± 0.24

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[I ◀](#)[▶ I](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Table 2. Number of OTUs and diversity indices for *rbcL* cDNA libraries obtained from the initial seawater and the incubation bottles on day 2 (Fe-380 and Fe-600) and day 3 (C-380 and C-600).

Library	No. of sequences	No. of OTUs	H'	1- D
Initial	35	10	1.81	0.803
C-380	50	15	1.98	0.768
C-600	50	14	1.60	0.631
Fe-380	50	13	2.24	0.884
Fe-600	50	12	2.01	0.842

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

BGD

11, 18105–18143, 2014

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

Table 3. Significance levels for differences among *rbcL* cDNA libraries as calculated with LIBSHUFF. *p* values < 0.05 are bolded.

	Library (Y)				
	Initial	C-380	C-600	Fe-380	Fe-600
Library (X)					
Initial		0.434	0.573	0.383	0.587
C-380	0.153		0.086	0.101	0.898
C-600	0.523	0.500		0.004	0.033
Fe-380	< 0.001	< 0.001	< 0.001		0.002
Fe-600	0.009	0.004	< 0.001	0.030	

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

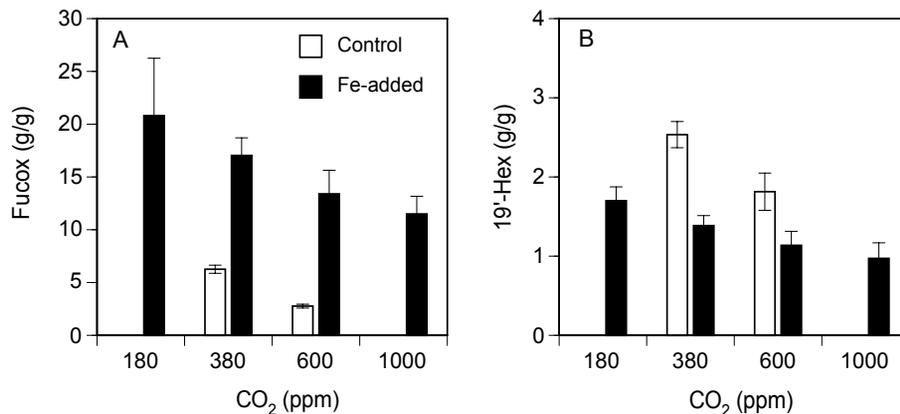


Figure 1. Ratios of (a) fucoxanthin (Fucox) or (b) 19'-hexanoyloxyfucoxanthin (19'-Hex) concentrations at the initial and final sampling day (final day/initial). Open bars and closed bars denote control and Fe-added treatments, respectively. Error bars denote ±1 SD ($n = 3$).

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[⏪](#)
[⏩](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

Effects of CO₂ and iron availability on *rbcl* gene expression in Bering Sea diatoms

H. Endo et al.

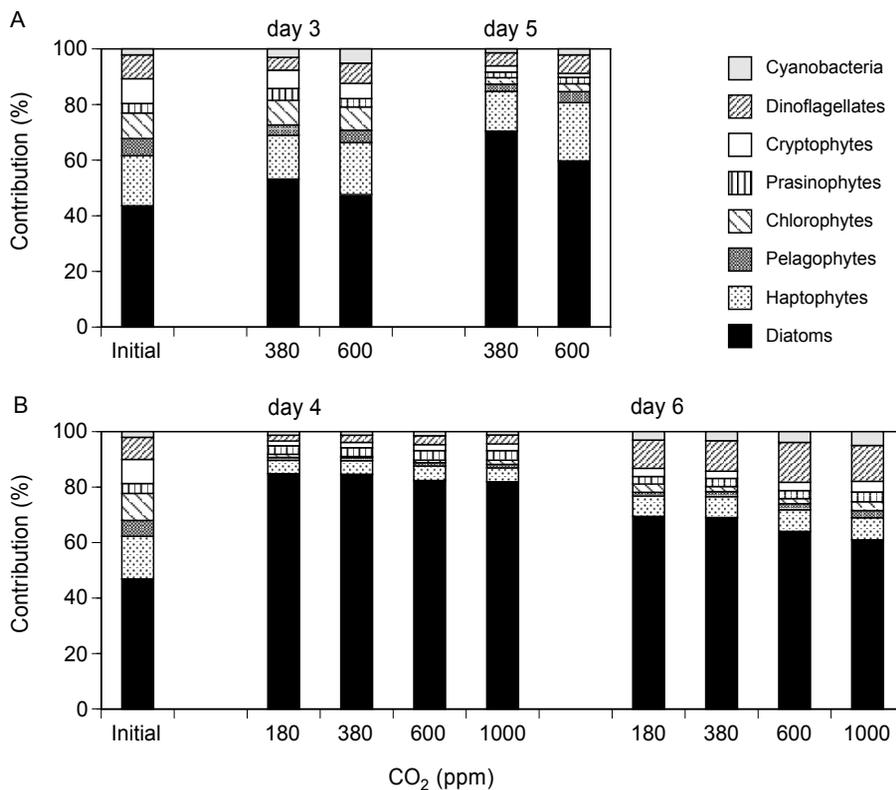


Figure 2. Mean contributions of each phytoplankton group to total Chl *a* biomass estimated by CHEMTAX in the **(a)** control bottles at 380 and 600 ppm CO₂, and **(b)** Fe-added bottles at 180, 380, 600 and 1000 ppm CO₂ ($n = 2$ or 3).

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

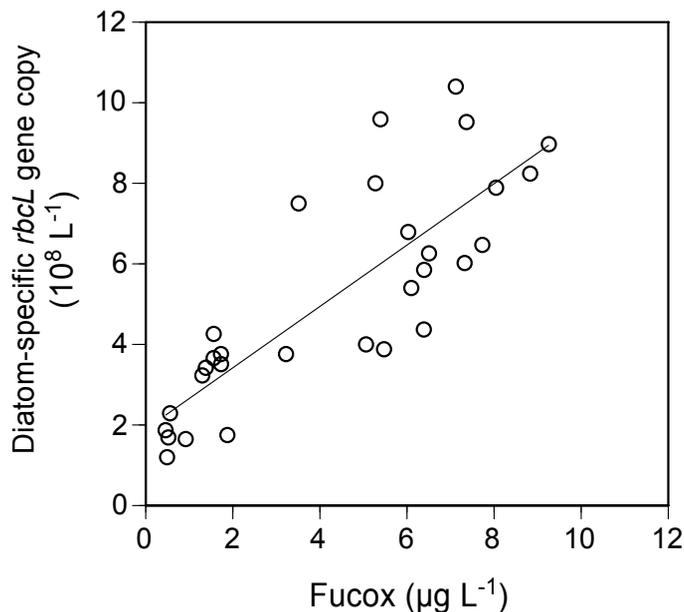


Figure 3. Relationship between fucoxanthin (Fucox) concentration and diatom-specific *rbcL* copy number ($y = 7.62 \times 10^8 x + 1.90 \times 10^8$, $r^2 = 0.677$, $p < 0.001$, $n = 28$).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

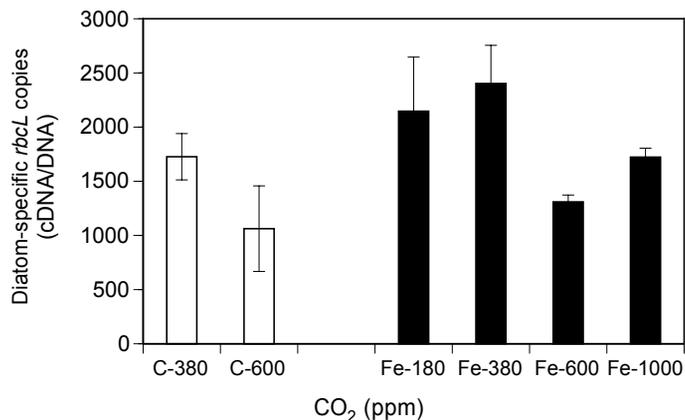


Figure 4. Abundances of *rbcL* mRNA (cDNA) normalized to *rbcL* gene copy number (*rbcL* cDNA/DNA) in the control bottles on day 3 and the Fe-added bottles on day 2. Open bars and closed bars denote control and Fe-added treatments, respectively. Error bars indicate ± 1 SD ($n = 3$).

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Effects of CO₂ and iron availability on *rbcL* gene expression in Bering Sea diatoms

H. Endo et al.

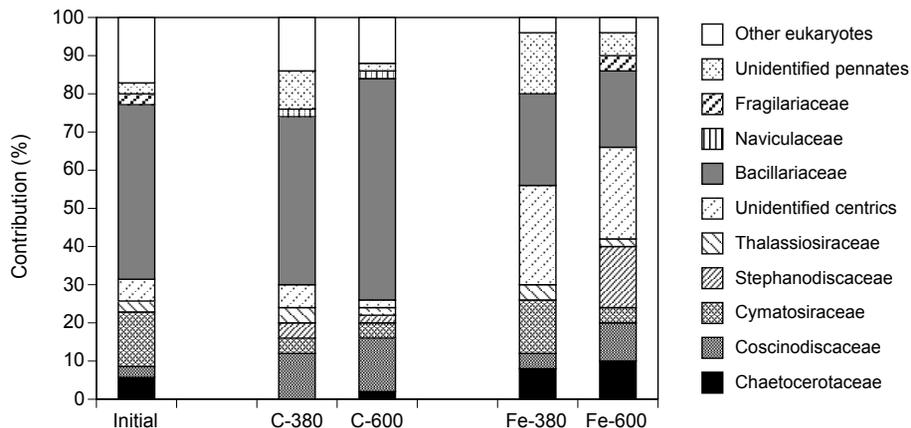


Figure 6. Relative phylotype contributions in the *rbcL* cDNA libraries obtained from the initial seawater and the incubation bottles at day 2 (Fe-380 and Fe-600) and day 3 (C-380 and C-600).

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion