1 **Title:** Local spatial structure of forest biomass and its consequences for remote sensing of

2 carbon stocks

3 **Running head**: Spatial sampling of forest biomass

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Abstract:

Advances in forest carbon mapping have the potential to greatly reduce uncertainties in the global carbon budget and to facilitate effective emissions mitigation strategies such as REDD+. Though broad scale mapping is based primarily on remote sensing data, the accuracy of resulting forest carbon stock estimates depends critically on the quality of field measurements and calibration procedures. The mismatch in spatial scales between field inventory plots and larger pixels of current and planned remote sensing products for forest biomass mapping is of particular concern, as it has the potential to introduce errors, especially if forest biomass shows strong local spatial variation. Here, we used 30 large (8-50 ha) globally distributed permanent forest plots to quantify the spatial variability in aboveground biomass density (AGBD in $Mg.ha^{-1}$) at spatial scales ranging from 5 to 250 m (0.025-6.25 ha), and to evaluate the implications of this variability for calibrating remote sensing products using simulated remote sensing footprints. We found that local spatial variability in AGBD is large for standard plot sizes, averaging 46.3% for replicate 0.1 ha subplots within a single large plot, and 16.6% for 1 ha subplots. AGBD showed weak spatial autocorrelation at distances of 20-400 m, with autocorrelation higher in sites with higher topographic variability and statistically significant in half of the sites. We further show that when field calibration plots are smaller than the remote sensing pixels, the high local spatial variability in AGBD leads to a substantial "dilution" bias in calibration parameters, a bias that cannot be removed with standard statistical methods. Our results suggest that topography should be explicitly accounted for in future sampling strategies and that much care must be taken in designing calibration schemes if remote sensing of forest carbon is to achieve its promise.

1   Introduction

Forests represent the largest aboveground carbon stock in the terrestrial biosphere, and forest degradation and regrowth are globally important carbon fluxes (Pan et al., 2011). Our ability to predict future atmospheric CO_2 concentrations or to implement effective carbon emission mitigation strategies (e.g. REDD+; Miles and Kapos, 2008) is limited by the accuracy of forest carbon stock estimates. The global monitoring of forest carbon stocks has thus come to the fore of the research agenda, with important implications in economics, policy and conservation (Gibbs et al., 2007).

Aboveground carbon stock estimates based on field inventories and on remote sensing approaches have led to substantial progress in mapping broad-scale carbon stocks (Asner et al., 2010; Baccini et al., 2012; Malhi et al., 2006; Saatchi et al., 2011). However, such carbon maps still have substantial uncertainties (Mitchard et al., 2014). The most common approach to quantifying forest carbon stocks at regional and national scales is to first to stratify the area of interest, and then to assign to each strata a mean carbon density value estimated from ground measurements. This approach inherently overlooks extensive spatial variation in carbon density within strata, and assumes that forest carbon fluxes from deforestation or regrowth may be deduced from changes in the extent of forest cover types and/or their mean carbon densities. Unfortunately, forest classifications are often insensitive to forest degradation and regrowth, both crucial components of forest carbon fluxes (Harris et al., 2012; Lewis et al., 2009). Thus, recent studies have moved toward more continuous measures of forest structure and carbon density, such as those provided by space-based and airborne sensing of vegetation (Asner et al., 2010, 2013; Goetz and Dubayah, 2011; Wulder et al., 2012).

Active remote sensing tools such as Light Detection and Ranging (LiDAR) and synthetic aperture radar (SAR) are currently the best candidates for forest carbon mapping at

The quality of a field-based calibration and resulting products depends fundamentally on how well forest biomass density in entire pixels is represented by the field data. In space-based remote sensing of forest biomass, sensor footprints are usually several to many times larger than field calibration plots (Baccini et al., 2007). If forest biomass is uniform within pixel-sized areas, this mismatch in sample area will have little impact on calibration; however, if there is substantial local spatial variability in biomass, then small calibration plots will have large sampling errors. In general, as the sampling area decreases, the variability associated with any field biomass estimate increases, as does associated sampling error. In addition, the remote sensing field of view often differs from the field-based view as a result of potential geolocalisation errors, the post-geoprocessing conversion of an ellipsoidal footprint into a square pixel, and fundamental differences between the forest components measured. Slant-range scale distortion is a typical example of such spatial mismatch in radar products (Henderson and Lewis, 1998) and remote sensing of canopy structure versus field-based tree stem measurements is a common source of spatial mismatch in high-resolution remote sensing products (Mascaro et al., 2011). Such spatial mismatches may considerably increase errors during the calibration step. There is thus a need to quantify these errors and test potential strategies to address them.

Here, we analyzed spatially explicit forest census data from a global network of 30 large permanent plots (8 to 50 ha) in natural forests (Condit, 1998; Losos and Leigh, 2004) to quantify local variation in aboveground biomass density (AGBD) and explore its consequences for calibrating large-footprint remote sensing products with field data for smaller plots (Fig. 1; Supplement, Table S1). Using these very large plots, we address three questions: (1) What is the local variability in aboveground biomass density (AGBD) for standard plot sizes, how does this variability scale with the area sampled, and how does it differ among sites, forest types, and continents? (2) What is the local spatial structure of AGBD, how does this spatial structure vary with the area sampled, and how does it differ among sites? (3) What are the implications of the observed AGBD variability for the accuracy of remote sensing calibration equations, when calibration plots are smaller than sensor footprints, and for different statistical procedures?

2   Material and methods

2.1   Field data

We used standardized measurements in 30 large forest plots across three continents (8–50 ha each, Fig. 1 and Table S1). In 28 of the plots, all free-standing trees $\geq$ 1 cm dbh (diameter measured at 130 cm above the ground or 50 cm above buttresses) were mapped, tagged, and identified taxonomically (Condit, 1998). In two additional plots, only trees $\geq$ 10 cm in dbh were included (Table S1). Trees < 10 cm dbh generally contribute less than 5% of the total aboveground biomass (AGB) in mature tropical forests (Chave et al., 2003). AGB of each individual stem was estimated using regression models based on the measured individual diameter and the wood specific gravity assigned to that species and site, or site-specific allometric equations (details in Table S1). We only used data for free-standing woody stems,

199 and excluded lianas from our analyses for the few sites where these were censused. Lianas

200 usually represent less than 5% of the total AGB (e.g. Schnitzer et al., 2012).

201   Elevation ranges were computed for each site based on 5 to 20 m elevation maps

202 generated from either field survey measurements (Condit 1998) or high-resolution airborne

203 LiDAR (in Paracou, Nouragues and Haliburton). Among 19 forest plots where elevation maps

204 were publicly available, the elevation range showed a strong and significant correlation with

205 the mean of the standard deviation of elevation within 1-ha subplots (Fig. S1). We therefore

206 used the elevation range, a metric available over all sites, as an indicator of topographic

207 variability.

2.2   Local spatial variability in AGBD

209 Each plot was gridded into subplots at spatial resolutions ranging from 5 to 250 m, to the

210 extent feasible given the plot dimensions. Within each subplot, AGBD (Mg.ha^{-1}) was

211 calculated by summing AGB estimates for all trees whose stems were located within the

212 subplot and expressing this on a per ha basis. We quantified the local spatial variability in

213 AGBD for subplots of area s (in ha) using the coefficient of variation of AGBD among

214 subplots within sites, calculated as

215
$$CV(s) = 100 \times \frac{\sigma(s)}{\mu} \quad (1)$$

216 where μ is the mean AGBD in the plot, $\sigma(s)$ is the standard deviation in AGBD computed

217 from subplots of area s, and $CV(s)$ is the coefficient of variation for plot area s in percent. A

218 higher CV value indicates a higher relative spatial variability of AGBD, and therefore greater

219 random sampling error when small subplots are used as samples to represent the full plot area.

220   We focused on the *CV* at the 1-ha scale, denoted CV(1) in our examination of

221 variation among sites. We evaluated whether CV(1) increased with AGBD among sites, and

222 whether it increased with topographic variability as represented by the elevation range, in

223 both cases using nonparametric Spearman rank correlations. We also tested whether CV(1)

224 varied significantly among continents or forest types using nonparametric Kruskal-Wallis

225 tests.

226   We examined the spatial scaling of variability with area both qualitatively with graphs

227 and quantitatively with fitted functions. Specifically, we graphed CV(s) vs. plot area (s) on

228 log scales, and fitted power functions to the relationship between the two. In the absence of

229 spatial autocorrelation (i.e. given independence of each grid cell), the logarithm of $CV(s)$

230 should decrease linearly with ln (s), with a slope of -½, just as the standard error of the mean

231 decreases with increasing sample size (that is, $CV(s) = \frac{CV(1)}{\sqrt{s}}$, thus $\log[CV(s)] =$

232 $\log[CV(1)] - 0.5 \; \log[s]$). Positive spatial autocorrelation will lead to a slower rate of decline

233 in the CV with increasing sample size over relevant spatial scales, and negative spatial

234 autocorrelation to a more rapid decline. We fitted power functions for the relationship of

235 CV(s) to *s* through linear regression on the log-transformed variables, and tested whether 95%

236 confidence intervals of the fitted exponents (slopes) included the value -0.5 expected in the

237 absence of autocorrelation. The confidence limits were calculated from the estimated standard

238 error of the slope and the Student's t distribution.

239 **2.3 Local spatial structure in AGBD**

240   We used empirical variograms to assess the spatial autocorrelation in AGBD for 20 ×

241 20 m (0.04 ha), 50 × 50 m (0.25 ha) and 100× 100 m (1 ha) subplots, with subplots created by

242 gridding each plot as above. We calculated variograms with the following formula:

243
$$\sigma^2(d) = \frac{1}{2N} \sum (AGBD_{xi+d} - AGBD_{xi})^2 \quad (2)$$

244   where $AGBD_{xi}$ is the AGBD observed at location *xi*, *d* is a class of spatial distance between

245   two locations and *N* is the number of pairs of observations, as implemented in the R package

246   geoR (Ribeiro Jr and Diggle, 2001). Distances between two subplots were based on the

247   coordinates of the center of each subplot. To make the variograms more usefully comparable

248   among plots, we transformed the variance $\sigma^2(d)$ to a coefficient of variation with $CV(d) =$

249     $100 \times \sqrt{\sigma^2(d)}/\mu$, where μ is the mean AGBD of the plot.

250   To further investigate the spatial structure of AGBD within field plots, we used

251   wavelet functions (Percival, 1995). The wavelet analysis decomposes the variance of a

252   process on a scale-by-scale basis, thus it is very useful for study of a variable influenced by

253   multiple processes operating simultaneously at different spatial scales (Detto and Muller-

254   Landau, 2013). A plot of wavelet variance versus scale indicates which scales are important

255   contributors to the total process variance. For example, global spatial variation in temperature

256   could be decomposed into the sum of large-scale variation due to latitude and smaller-scale

257   variation due to topography. In the absence of any spatial structure, the normalized wavelet

258   variance (the wavelet variance divided by the variance computed from the values of the

259   quadrats) is one at all scales. A value greater than one at scale *s* indicates that the variance of

260   the process at that specific scale is higher than expected under complete spatial randomness,

261   i.e., the scale-specific variation is spatially structured independent of the spatial variation

262   occurring at larger and smaller scales. In contrast, a normalized wavelet variance less than one

263   indicates that the scale-specific variation is lower than would be expected under complete

264   spatial randomness. Details of the methods for calculating the wavelet variances are given in

265   Appendix S1.

266   For each spatial scale, we then tested whether the scale-specific variation in AGBD

267   among sites is explained by topographic heterogeneity using repeated and independent

268   Spearman's rho correlation tests between the normalized wavelet variance and the elevation

269   range.

270   **2.4   Implications of local variability in AGBD for large-footprint remote sensing**

271   **calibration**

272   To assess the implications of local spatial variability in AGBD for remote sensing calibration,

273   we explored the joint influence of field plot size and of footprint size of a hypothetical remote

274   sensing observation on the sampling error associated with an AGBD estimate. We simulated

275   different plot sizes and footprint sizes under the best-case scenario in which the remote

276   sensing instrument was able to retrieve the exact value of AGBD as measured in field plots.

277   Because the remote sensing field of view often differs from the field-based one, we simulated

278   a spatial mismatch between the plot and footprint shape; for simplicity, we modeled the

279   remote sensing pixels as circles and the calibration plots as squares. We simulated field plots

280   of 0.04, 0.1, 0.25, 0.5, 1, 2 and 4 ha centered in remote-sensing circular footprints of 0.5, 1, 2

281   and 4 ha (Fig. 2). We then estimated the error associated with using the field plot to estimate

282   AGBD in the footprint, henceforth referred to as sampling error. Specifically, we calculated

283   *ErrCV* as the ratio between the root mean square error (*RMSE*) and the mean AGBD within

284   footprints (MAGBD) for each combination of areas in which the field plot area is less than or

285   equal to the footprint area:

286
$$RMSE = \sqrt{\frac{1}{N}\sum_{i=1}^{N}\left(AGBD_{footprint,i} - AGBD_{sublot,i}\right)^2} \quad (3)$$
287

288
$$MAGBD = \frac{1}{N}\sum_{i=1}^{n} AGBD_{footprint,i} \quad (4)$$
289

290
$$ErrCV = RMSE / MAGBD \quad (5)$$
291

292   where N is the number of simulations (1000 per combination), $AGBD_{footprint,i}$ is the AGBD

293   measured within the remote-sensing footprint (i.e. the circle) for the ith simulation, and

294      $AGBD_{subplot,i}$ is the AGBD measured within the field subplot for that simulation. Five of our

295   plots (the Haliburton plot and the four Ituri plots) were too small to accommodate a circular 4-

296   ha footprint and were thus not included in the calculation of $ErrCV$ at this scale.

297   To illustrate how this sampling error propagates into AGBD maps, we then fitted

298   calibration equations from the combination of simulated remote sensing pixels and field

299   calibration plots. For this exercise, we simulated square remote sensing pixels of 4 ha, thus

300   mimicking the expected resolution of the BIOMASS mission's future instrument (Le Toan et

301   al., 2011). Given the size of our field plots, we were able to simulate 60 such pixels (i.e. two

302   pixel per plot for 30 plots). Within each simulated pixel, we assumed that a single randomly

303   located field plot was available for calibration, of area 0.01, 0.04, 0.25, 0.5, 1 or 2 ha (i.e. 60

304   calibration plots, one per 4-ha pixel). For each field plot scale we calculated the regression

305   coefficients of an ordinary least squares (OLS) linear regression between the AGBD

306   estimated in the calibration subplots of a given area and the simulated pixels. We changed the

307   location of the subplots in each plot a thousand times and averaged the regression coefficients

308   for each subplot size.

309   It is well-established in the statistical literature that random error in the independent

310   variable, such as that which results from sampling error in field plots, leads to systematic

311   underestimation of the OLS regression slope, a bias referred to as attenuation or regression

312   dilution (Fuller, 1987). This phenomenon is easily understood as the OLS slope β is

313   calculated as $\beta = \sigma^2(X,Y)/\sigma^2(X)$, where $\sigma^2(X,Y)$ is the covariance of X and Y and $\sigma^2(X)$

314   is the variance of X. If W is a measure of X with measurement error (that is, $W = X + \varepsilon_X$),

315   then $\sigma^2(W) > \sigma^2(X)$ (Mcardle, 2003). Hence, the estimate of β tends to zero as the

316   measurement error in X increases to infinity, a phenomenon referred to as the dilution bias.

317   Several methods have been proposed to correct for this bias (Carroll and Ruppert,

318   1996; Frost and Thompson, 2000; Smith, 2009). The method of moments estimator (Carroll

319   and Ruppert, 1996; Fuller, 1987) assumes that a corrected slope, β_{MM}, could be calculated

320   from the observed slope, β, using a Reliability Ratio, R_r, with

321    $$\beta_{MM} = \frac{\beta}{R_r} \quad (6) \quad \text{where} \quad R_r = \frac{\sigma^2(W) - \sigma^2(\varepsilon_X)}{\sigma^2(W)} \quad (7)$$

322   To estimate $\sigma^2(\varepsilon_X)$, the variance of the sampling error in X, we generated new estimates of X

323   (here the AGBD of calibration plots) by bootstrapping over 0.01-ha (10 x 10 m²) subplots the

324   calibration plot (i.e. 100 bootstrapped values for each of the 60 calibration plots). The

325   reliability ratio R_r was estimated using the intra-class correlation coefficient (ICC), an

326   accurate proxy for R_r (Frost and Thompson, 2000), considering the bootstrapped values as

327   repeated measures grouped by calibration plot units. ICC was estimated through a one-way

328   analysis of variance of repeated measures considering the calibration plots as factor. This

329   approach was called "within subplot R_r". We also carried out a second reliability study based

330   on additional subplots (i.e. replicates) established randomly inside the 4-ha pixels (Appendix

331   S2).

332   We evaluated two alternatives to OLS that have the potential to produce less bias in

333   calibration equations. First, the Reduced Major Axis (RMA) regression minimizes the sum of

334   squared distances both horizontally (accounting for the error in X) and vertically (accounting

335   for the error in Y). Second, the nonparametric Theil-Sen estimator, also known as Sen's slope

336   estimator or the single median method, is the median of all the slopes determined by all pairs

337   of observations. Both methods have been proposed as preferred alternatives to OLS in remote

15

338   sensing studies (Cohen et al., 2003; Fernandes and Leblanc, 2005; Mitchard et al., 2013; Ryan

339   et al., 2012).

340   All analyses were performed using R version 3.0.2 (R Development Core Team,

341   2013). The R code for the analyses is available on request from the first author.

3   Results

3.1   Local spatial variability in AGBD

344   The coefficient of variation for AGBD at the 1-ha scale, $CV(1)$, varied among sites (n=30)

345   from 5.1% (Haliburton, Canada) to 29.9% (Palanan, Philippines), with a mean of 16.6%, and

346   a median of 15.2% (Table S2). The best predictor of variation in $CV(1)$ among plots was

347   within-plot elevation range, that is, the difference between the highest and lowest elevation

348   (Spearman's rho=0.70 and $p<10^{-4}$; Fig. 3a). Thus, topographic variability, represented in the

349   analyses by elevation range across the plot, explained considerable variation in AGBD

350   variability among sites at the 1-ha scale. In contrast, $CV(1)$ was not significantly correlated

351   with mean AGBD (Spearman's correlation test, p=0.15), and did not differ significantly

352   among forest types (tropical, subtropical and temperate; Kruskal-Wallis test, p=0.47) or

353   among continents (Kruskal-Wallis test: p=0.18). Asian tropical field plots tended to show

354   higher biomass variability than other tropical field plots (median CV(1) of 24.4 and 14.3 %

355   respectively), consistent with their higher average topographical variability (median elevation

356   range of 90 m for Asian tropical plots and 24 m for tropical non Asian).

357   Regressing the logarithm of CV(s) against ln (s), we found that in 15 of 30 sites the

358   slope was significantly greater (less negative) than -½, suggesting significantly positive

359   spatial autocorrelation in AGBD at the scales investigated. In contrast, in only two sites, the

360   Ituri Edoro1 plot in Democratic Republic of Congo and the Paracou plot in French Guiana

16

(Fig. 3b, Table S2-3), the slope was significantly lower than -½, suggestive of negative spatial

autocorrelation. Sites with greater elevation range showed gentler fitted slopes (Spearman's

rho= 0.47 and p=0.01). Such positive spatial autocorrelation means that extrapolation from 1

ha values under the assumption of no spatial autocorrelation will lead to a slight but

systematic overestimation of $CV(s)$ for areas (s) smaller than 1 ha, and underestimation for

areas larger than 1 ha (Fig. S3).

3.2 Local spatial structure in AGBD

Variograms revealed only weak spatial autocorrelation of AGBD at 20, 50 and 100-m

resolution over distances of 20-400 m (Fig. 4, Fig. S5). The average coefficient of variation

for AGBD was only slightly higher between distant subplots than between neighboring ones.

Though these increases with distance were generally very small, they were statistically

significant in half of the plots at 20 and 50-m resolution (Fig. S6-8), consistent with the

results of the analysis of the slope of spatial variability with plot scale (see above).

Wavelet analyses also showed a relative small departure from the complete spatial

randomness (Fig. 5, Fig. S9). The average normalized wavelet variances at scales above ~90

m were greater than one, indicating that a substantial part of the spatial structure of AGBD

occurs at these scales which might explain the observed weak spatial autocorrelation found in

the variograms. Interestingly, many sites showed a gap of variability at intermediate scales

(25-75 m). The plots with greater elevation range were characterized by larger wavelet

variances at scales >100 m (Fig. 5, Fig. S9), suggesting that the large scale variations are

driven by topographic effects.

3.3 Implications of local spatial variability in AGBD for large-footprint remote

sensing calibration

384   Field-based sampling error, depended on both field plot and remote sensing footprint areas.

385   For very small field subplots (0.1 ha and below), sampling error was due mostly to field

386   sampling and was relatively insensitive to the footprint size (Fig. 6). For subplots and

387   footprint size of 0.5 ha and larger, subplot area and footprint area had similar effects on the

388   sampling error. The error due to the spatial mismatch (circle versus square) was much higher

389   for small calibration plots: when the field calibration plot area was equal to the footprint area

390   (i.e. a ratio of one), the error ErrCV was higher for smaller sample areas (Fig. S10).

391   Field-based sampling error resulted in systematic underestimation of calibration

392   slopes, which could not be corrected through any currently available statistical approaches.

393   The OLS regression slope was underestimated by an average of 54% with 0.1-ha subplots and

394   by 37% with 0.25-ha subplots (Fig. 7a, see examples of fits on Fig. S11). The large sampling

395   errors associated with small field plots caused large dilution biases (i.e. slope

396   underestimation). Such dilution biases result in an underestimation of the variance in AGBD;

397   in particular, application of the resulting calibration equations would produce systematic

398   underestimation of AGBD in high AGB areas, and systematic overestimation in low AGBD

399   areas. Alternatives to OLS models, such as Reduced major axis (RMA) regression and the

400   Theil-Sen estimator, corrected for at best half of this bias (Fig. 7b). Our bias correction

401   approach, based on bootstrapping over spatial variability within subplots, outperformed the

402   RMA and the Theil-Sen estimator for plots ≥ 0.25 ha, but remained too conservative ("Within

403   subplot R_r" in Fig. 7b). The alternative reliability study approach involving replicate subplots

404   did somewhat better, but requires greatly increased ground sampling effort (Appendix S2,

405   Figure S2).

406   **4   Discussion**

Given the pressing need to monitor global forest carbon stocks, ecologists and remote sensing experts need to pay careful attention to quantifying the errors associated with forest carbon estimates. Our results quantify large spatial variability in mean AGBD for plot sizes smaller than 0.25 ha (the mean CV was of 26 % at the 0.25-ha resolution; table S2). This large local spatial variability in AGBD results in substantial sampling errors when small plots are used to estimate AGBD within larger areas. These sampling errors in turn bias calibration equations based on such estimates. Many forest inventory plots are much smaller than 0.25 ha and are regularly used for calibrating coarser resolution remote sensing products. Our findings suggest that using such small field plots to calibrate coarser resolution remote sensing products is likely to cause strong systematic biases in carbon maps.

4.1   Local spatial variability and spatial structure of AGBD

Spatial variability in mean AGBD decreased predictably with increasing plot area. We found that the coefficient of variation in AGBD averages ~16.6% at 1 ha, and scales roughly with $s^{-1/2}$ where s is the plot area. This present study confirms the findings of previous studies of individual sites or forest types (Baraloto et al., 2013; Chave et al., 2003; Holdaway et al., 2014; Keller et al., 2001; Wagner et al., 2010) and generalizes the results to many sites that encompass a wide range of forest types and topographical variation. We found that spatial variability of AGBD tended to be greater in hilly terrain, confirming that topography is a major driver of AGBD variability (e.g. de Castilho et al., 2006; Detto et al., 2013). This is an important finding given that 23% of the world's forests are on hilly terrain (Table S4). This result suggests that forest biomass maps in hilly areas have larger uncertainties, and that forest plot sampling designs should take topography into account (see below).

We found no other systematic differences in AGBD variability among continents, among forest types or with mean AGBD. Asian tropical forests displayed higher AGBD

19

431 variability than other tropical sites, but this could be explained by the larger topographic

432 variability in our tropical Asian study sites (Table S1). This finding is no accident of our study

433 locations; remaining old-growth tropical forests in Asia are disproportionately located in

434 topographically complex terrain, more so than on other continents (Table S4), probably

435 because these areas have disproportionately escaped human disturbance.

436   Our analyses revealed at most weak positive spatial autocorrelation of AGBD within

437 sites. The differences in AGBD among field plots separated by 100-400 m are not much

438 greater than the differences between adjacent plots (Fig. 4). Nonetheless, approximately half

439 of the sites individually exhibited significant spatial autocorrelation in AGBD. Decomposition

440 of the variance in AGBD at different spatial scales using wavelet analyses confirmed spatial

441 aggregation at scales >100 m, and the role of topography in explaining aggregation at these

442 scales (Fig. 5b). These results suggest that the weak spatial autocorrelation found in many

443 plots is due to broad-scale topographic differences. In a previous scale-wise analysis of a

444 single larger 5000 ha area of moist tropical forest, Detto et al. (2013) likewise found strong

445 wavelet coherence between spatial variation in topography and in canopy height (a proxy for

446 AGBD) at scales of 100-800 m. These scale-specific results are consistent with prior literature

447 (reviewed in Detto et al., 2013) documenting how forest structure and biomass vary with

448 topography (de Castilho et al., 2006; McEwan et al., 2011; Valencia et al., 2009).

449   In most plots, the wavelet analyses also uncovered more uniformity than expected by

450 chance at scales of 25-75 m. We hypothesize that this pattern may be associated with

451 neighborhood competition and gap-phase dynamics. That is, the forest can be thought of as a

452 mosaic of patches of different age, reflecting time since the last major treefall or other

453 disturbance opened a gap in the canopy, with patch age strongly influencing AGBD

454 (Moorcroft et al., 2001). Within such patches, biomass variation is reduced by the common

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455 time since disturbance, and also because local competition may cause large trees to be more

456 evenly spaced than would be expected by chance (Lutz et al., 2013). This local uniformity is

457 overlaid on the larger-scale topographic variation, and is evident only through scale-wise

458 wavelet analyses that separate the two.

4.2   Field sampling error and remote sensing of carbon stocks

460 As expected, sampling error associated with using a small field plot to estimate AGBD over a

461 larger remote sensing footprint depends on the sizes of both the field plot and the remote

462 sensing footprint. However, when field plots were very small (0.1 ha and below), the

463 sampling error was due mostly to the contribution from field sampling, and was relatively

464 insensitive to footprint area. Hence, with relatively high resolution pixels such as in the

465 Landsat (30 m) or ICESat/GLAS (~70 m) products, sampling errors are likely to be very high

466 if smaller plots are used or if spatial mismatches between the field and the sensor signal occur.

467 This is because most of the AGBD variability is at the local scale so that a small difference

468 between the areas sampled in the ground and by the sensor generates a large error. This is

469 well illustrated by our finding that error was much lower for large calibration plots even when

470 the same ratio of calibration plot area to footprint area was maintained (Fig. S10). This

471 reflects decreasing edge-to-area ratios for larger area, which also provide other advantages for

472 larger plots (see also Mascaro et al., 2011; Zolkos et al., 2013).

473   Remote sensing footprints and field plots often differ in their field of view, a mismatch

474 we simulated by assuming remote sensing footprints were circular and field plots were square.

475 In general, remote sensing footprints are ellipsoidal with different shapes depending on the

476 geometry of observation and imaging configurations. Ellipsoidal footprints can be processed

477 to represent rectangular areas, but such post-processing step may generate spatial mismatches

478 between what has been measured by the sensor and what is really occurring within the remote

21

479   sensing pixel (e.g. slant-range scale distortion in radar products; Henderson and Lewis, 1998).

480   Spatial mismatches between plot and footprint areas may also arise for other reasons,

481   including geolocalisation errors or difference between the forest components measured (e.g.

482   crown versus trunk measurements, Mascaro et al., 2011). Our approach here presents one

483   simplified example of the consequences of these kinds of mismatches; more precise

484   quantification of such errors could be obtained using sensor-specific and 3D simulation

485   approaches.

486   Our analyses show that field-sampling strategy may result in a serious bias in model

487   calibration of remote sensing products. When this bias is present, inversion models return

488   AGBD values that are regressed to the mean of the calibration plots (Fig. 7a), and thus

489   underestimate the true spatial AGBD variance. For instance, in a recent study that used 112

490   circular 0.13-ha plots to calibrate L-band radar products (Carreiras et al., 2012), the slope of

491   an OLS regression was found to be underestimated by 86% and the final AGBD map

492   displayed a much lower variance than the global map produced by Saatchi et al. (2011). The

493   dilution bias is independent of the number of calibration plots; it depends only on the

494   sampling error associated with these plots, which is determined largely by plot size. Though

495   the mean AGBD of the calibration plots is inherently correctly predicted (Fig. 7a), the

496   landscape mean AGBD and thus the landscape total AGBD will be correctly predicted only if

497   the landscape mean is identical to the mean of the calibration plots.

498   We tested alternative approaches to OLS regression, including Reduced Major Axis

499   (RMA) regression and the Theil-Sen estimator, and found that the best way to diminish the

500   dilution bias is to bootstrap over spatial variability using subplots within plots and to correct

501   the estimated slope using these simulated "replicates". Some remote sensing studies have

502   argued that alternative to OLS regression such as RMA or the Theil-Sen estimator are good

alternatives to OLS regression when errors occur in X (Cohen et al., 2003; Fernandes and Leblanc, 2005; Mitchard et al., 2013; Ryan et al., 2012). Here, we showed that these alternatives do not resolve the dilution bias and still provide strongly biased products. In theory, the dilution bias could be removed completely through Deming regression; however, this approach requires information on the ratio of the error variances in the two variables (Deming, 1944). The results we present here can assist in the estimation of error variances for field plots of different sizes. However, estimating error variances for remote sensing products – that is, their error in providing an estimate of the true value of AGBD – remains a challenge.

4.3 Implications for designing forest inventories and remote sensing calibration schemes

Our careful quantification of local spatial variability and spatial structure in AGBD should be useful for the design of national and regional forest inventories, as well as in remote sensing applications. Weak spatial autocorrelation at scales less than 100 m suggests that there is generally no gain in representativeness from locating multiple small plots within a small area or footprint (≤ 100 m) when compared to establishing one larger plot in the same area. That is, because neighboring small plots are on average almost as different as more distantly located small plots, thus expanding a single small plot provides similar information as adding another small plot nearby. A number of forest inventory designs use clusters of very small plots (≤ 0.04 ha); e.g., the US Forest Service Forest Inventory and Analysis program (Bechtold and Patterson, 2005). Based upon our results these cluster designs appear to have distinct disadvantages for calibrating remote sensing products as their small dimensions are below the resolution of most sensors, and their edge to area ratios are higher than single larger plots for the same total area. Although small plots may have practical advantages in time needed for field sampling and reduced equipment costs, these advantages should be carefully weighed

against the disadvantages for biomass measurements. Such small plots may induce strong biases when used individually for calibrating coarser resolution remote sensing products.

Our results reinforce the importance of topography as a factor that should be taken into account in designing forest inventories. AGBD variation at scales of >100 m was strongly associated with topographic variation in our analyses as was also found in previous studies (Detto et al., 2013). This suggests that sampling should generally be stratified by topography, especially if landscape AGBD is to be estimated purely from a field-based approach. In contrast, where the aim of field sampling is to calibrate coarse resolution remote sensing products, this might suggest that topographically complex areas should best be avoided to minimize sampling errors associated with local spatial variability. However, the gain from reducing such sampling errors would have to be weighed against the potential to bias the calibration sample if forests in topographically complex areas differ systematically in the relationship between remote sensing signals and AGBD.

The best way to avoid the dilution bias is to use calibration plots covering entire remote sensing pixels. For remote sensing tools with a resolution on the order of 4 ha, such as the planned BIOMASS mission, it is realistic to invest in a network of similarly sized field calibration plots. Though such field sampling is expensive, it would greatly improve the basis for mapping forest biomass, and its cost would remain small compared with the investment in the satellite itself. An alternative is to use a two-step approach in which a coarse-resolution remote sensing product is calibrated against a higher resolution remote sensing product itself calibrated with field plots. For instance, airborne LiDAR approach may retrieve forest carbon stocks with an error of ca. 10-15% at 1-ha resolution (Mascaro et al., 2011; Zolkos et al., 2013). This compares favorably with errors from purely field-based estimates for 1-ha and smaller plots (Fig. 3). Errors in LiDAR-based estimates are expected to be even lower for

larger areas, as random errors average out (Mascaro et al. 2011). Baccini and Asner (2013) found that using wall-to-wall airborne LiDAR AGBD estimates to calibrate a 500-m resolution MODIS product led to much less error than using nested AGBD estimates from Geoscience Laser Altimeter System (GLAS) footprints (60 to 75-m resolution). This shows that even if the operational cost associated with LiDAR coverage is high, the use of LiDAR technology has the potential to greatly reduce the errors during the calibration step. In this case, care must be taken that errors are carefully and appropriately propagated through the two-stage calibration to the final map (Asner et al., 2013).

Future research should integrate the results of this study with information on other sources of error in order to assess the relative importance of field sampling errors to forest carbon estimation and make appropriate recommendations. Other important sources of error in forest carbon estimates include field measurement errors (Flores and Coomes, 2011; Larjavaara and Muller-Landau, 2013), biomass allometries (Chave et al., in press, 2004; Molto et al., 2013), data cleaning procedures (Muller-Landau et al., 2014), and wood carbon content (Thomas and Martin, 2012). At the scale of forest inventories and calibration schemes, a major source of error is the uneven and non-random distribution of plots at broad spatial scales, an outstanding problem in the tropics where, for example, the central Amazon, the central Congo basin, and swamp forests all remain insufficiently sampled.

5   Conclusions

Accurate measurements of forest carbon stocks are critical to reduce uncertainties in the global carbon budget and for the REDD programme. However, uncertainty associated with forest carbon maps remains poorly quantified (but for notable exceptions see Asner et al., 2013; Gonzalez et al., 2010; Mermoz et al., in press). In this paper, we used a large-scale global dataset to illustrate that high local spatial variability in AGBD leads to large sampling

25

errors when plots of standard sizes (e.g., 0.1, 0.25, 1 ha) are used to estimate AGBD over larger areas (e.g., 4 ha, the expected resolution of BIOMASS products). We also show that remote sensing products that rely on field data for calibration may be highly biased if such field-sampling errors are large. Such biases have previously been ignored by the remote sensing community and, as we show, can only be partially corrected by available statistical tools. Overall, our results strongly suggest that calibration of coarse-resolution remote sensing products to estimate forest carbon would benefit greatly from more investment in large forest plots – plots large enough to encompass entire pixels. We hope that this contribution will stimulate further work on the propagation of field sampling errors to remote sensing products and that future studies will pay more careful attention to field sampling and calibration strategies.

Acknowledgments

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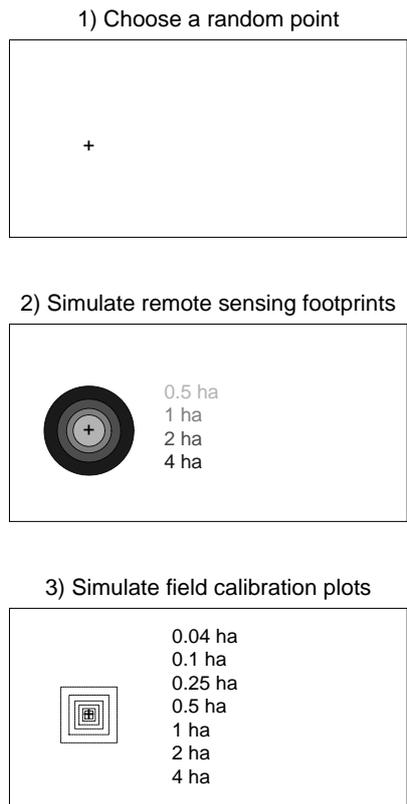804 review, Remote Sensing of Environment, 121, 196–209, doi:10.1016/j.rse.2012.02.001, 2012.

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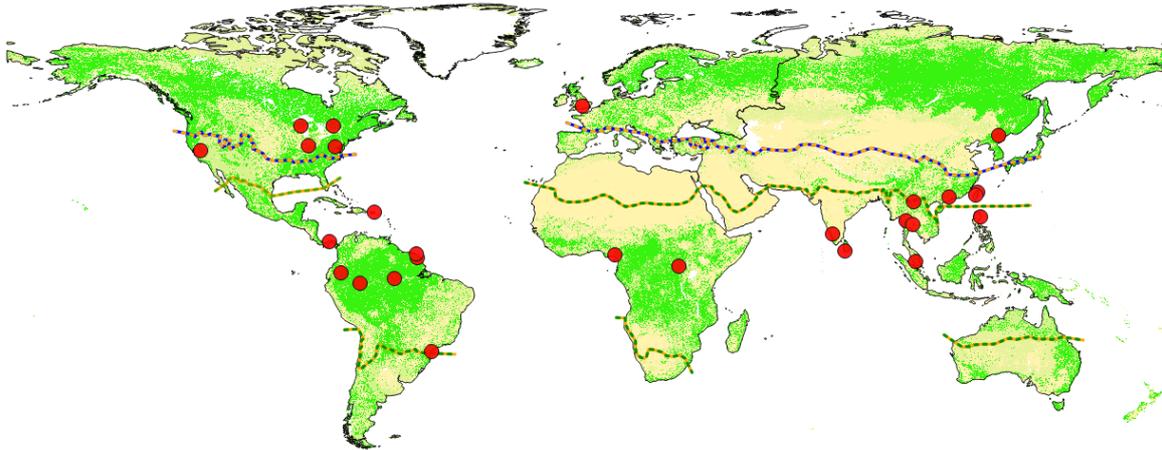
809
810   Figure 1. Geographical distribution of the 30 study sites (red points) included in the present

811   study, relative to the global distribution of forest (green) from GLOBCOVER2009 (Bontemps

812   et al., 2011), and the boundaries between temperate and subtropical areas (blue and orange

813   dashed lines) and between subtropical and tropical areas (orange and green dashed lines) from

814   Fischer et al. (2012). The four sites at Ituri (Democratic Republic of Congo) are represented

815   by a single dot due to their proximity. Note that Fischer et al. (2012) classify the Yosemite

816   site as subtropical, but we considered it as temperate due to its high elevation. Details on

817   study sites are provided in Table S1.

818

1) Choose a random point

2) Simulate remote sensing footprints

0.5 ha
1 ha
2 ha
4 ha

3) Simulate field calibration plots

0.04 ha
0.1 ha
0.25 ha
0.5 ha
1 ha
2 ha
4 ha

819

820   Figure 2. Schematic representation of the simulations used to assess expected errors when the

821   calibration/validation plots and the remote sensing footprint differ in shape and size. 1)

822   Within each large mapped plot, a point is chosen to be the center of both the simulated remote

823   sensing footprints and the simulated calibration subplots; it is chosen randomly from all

824   points for which the largest footprints and calibration plots are fully inside the mapped large

825   plot. 2) $AGBD_{footprint}$ is calculated within circular areas centered on this point, simulating the

826   remote sensing footprint, for the listed sizes. 3) $AGBD_{subplot}$ is calculated within square areas

827   centered on this point, simulating calibration/validation plots, for the listed sizes. We

828   replicated this procedure 1000 times and then calculated the root mean squared error of

829    $AGBD_{subplot}$ relative to $AGBD_{footprint}$ for each combination of areas in which the subplot area is

830   less than or equal to the footprint area, and normalized by the mean $AGBD_{footprint}$ to obtain a

831   measure of relative error specific to that combination of scales, $ErrCV$ (see equations 3-5).

a Spatial variability (1 ha) and topography   **b** Spatial scaling of AGBD variability
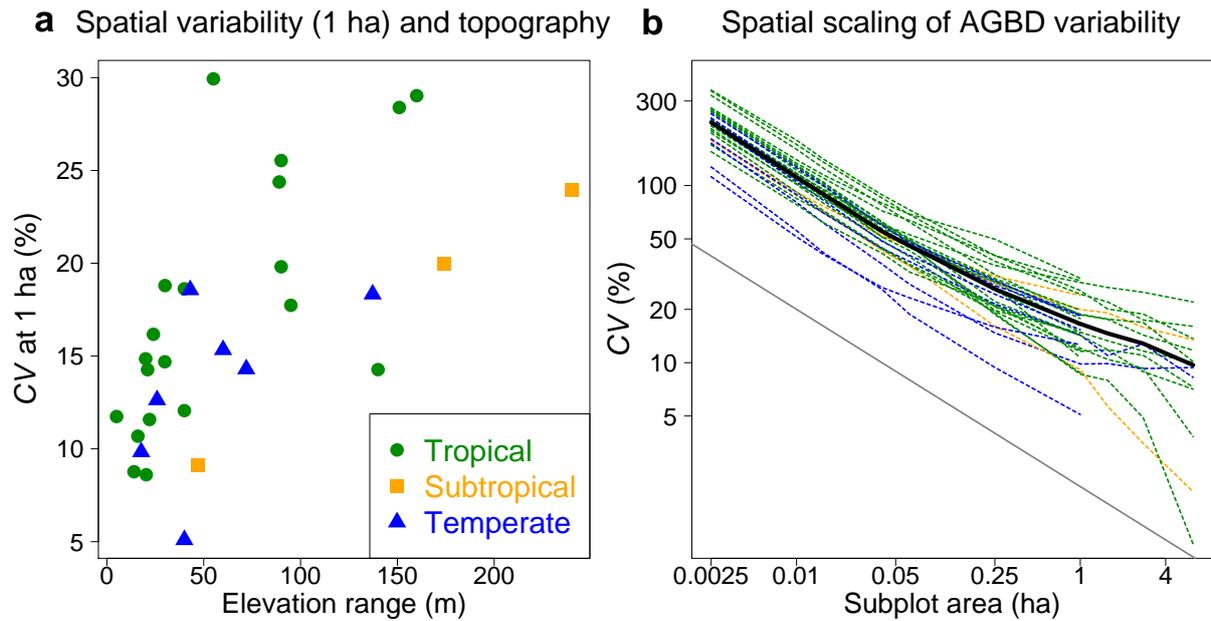
832
833   Figure 3. Local spatial variability in AGBD as a function of topographic variability and of

834   spatial scale. For each site, the spatial variability at a given spatial scale was quantified as the

835   coefficient of variation of AGBD for square subplots that gridded the entire large plot. (a) The

836   variability at the 1-hectare scale, $CV(1)$, was positively correlated with elevation range (the

837   difference between highest and lowest altitude) among plots (one point per site). (b) The

838   variability declined with increasing spatial scale within each site (one dashed line per site) and

839   in the cross-site mean (solid black line); the rate of decline deviated somewhat from the slope

840   of -0.5 (on log-log scales) expected in the absence of spatial autocorrelation in AGBD (solid

841   grey reference line), especially at larger spatial scales. Separate graphs for each individual site

842   are provided in Fig. S3 and standardised CV measures within 4-ha subplots are shown in Fig.

843   S4.

844

845
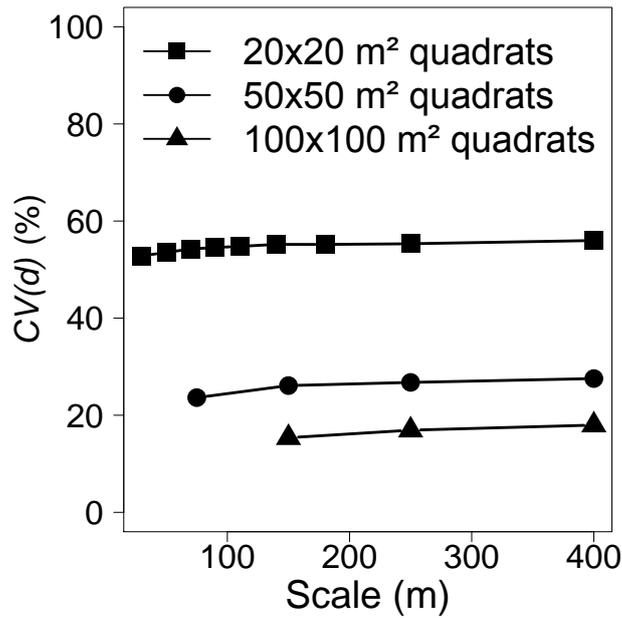
Figure 4. Spatial autocorrelation of AGBD for three different spatial resolution. Ensemble

average variograms for AGBD in square subplots of 20, 50 and 100-m resolution, with

variances transformed into distance-specific coefficients of variation (CV(d)) to facilitate

combining and comparing data across sites (see methods). Variograms for individual plots at

each spatial resolution are shown in Fig. S5. Separate graphs for each site, with confidence

intervals for the null hypothesis of no spatial correlation, are shown in Fig. S6-8.

852

a Scale-wise decomposition of AGBD variation

b Scale-specific variation and topography
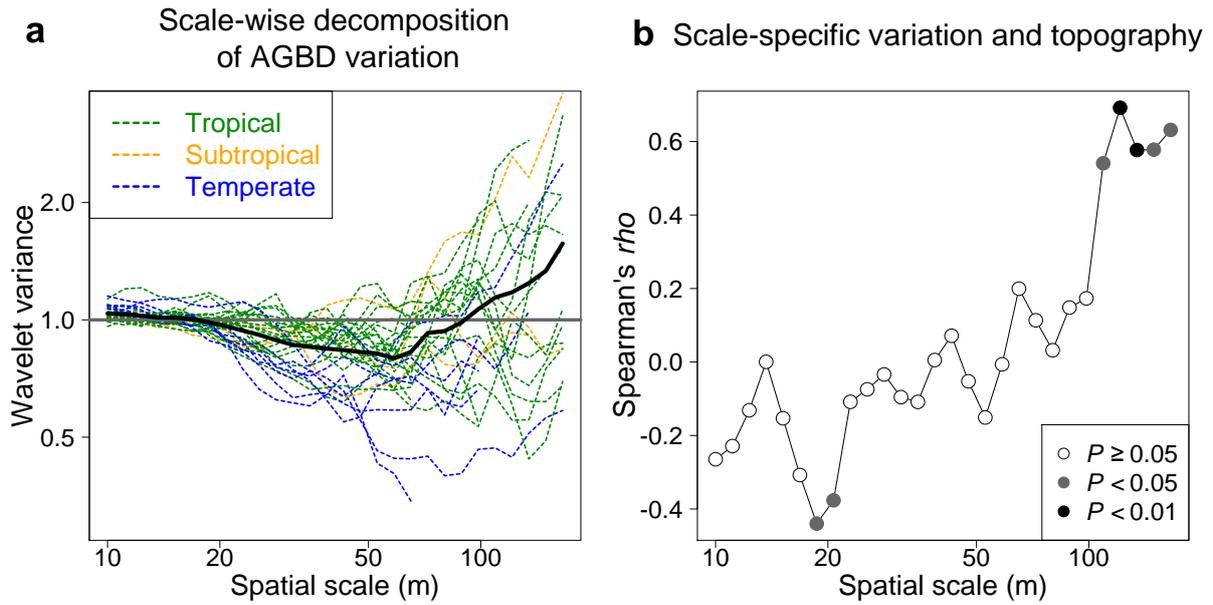
853
854 Figure 5. Scale-wise decomposition of spatial variation in AGBD and its relationship to

855 topographic heterogeneity. (a) The normalized wavelet variance of AGBD as a function of

856 spatial scale for individual plots (colored lines) and for the ensemble average across plots

857 (solid black line). Wavelet analysis decomposes the variance of AGBD on a scale-by-scale

858 basis; the normalized wavelet variance at a given scale reflects the spatial structure of AGBD

859 specific to that scale, with a value of one (solid grey line) indicating no spatial

860 autocorrelation, lower values indicating negative spatial autocorrelation, and higher values

861 positive spatial autocorrelation. Separate graphs for each site, with confidence intervals for

862 the null hypothesis of no spatial correlation, are shown in Fig. S9. (b) Among-site Spearman's
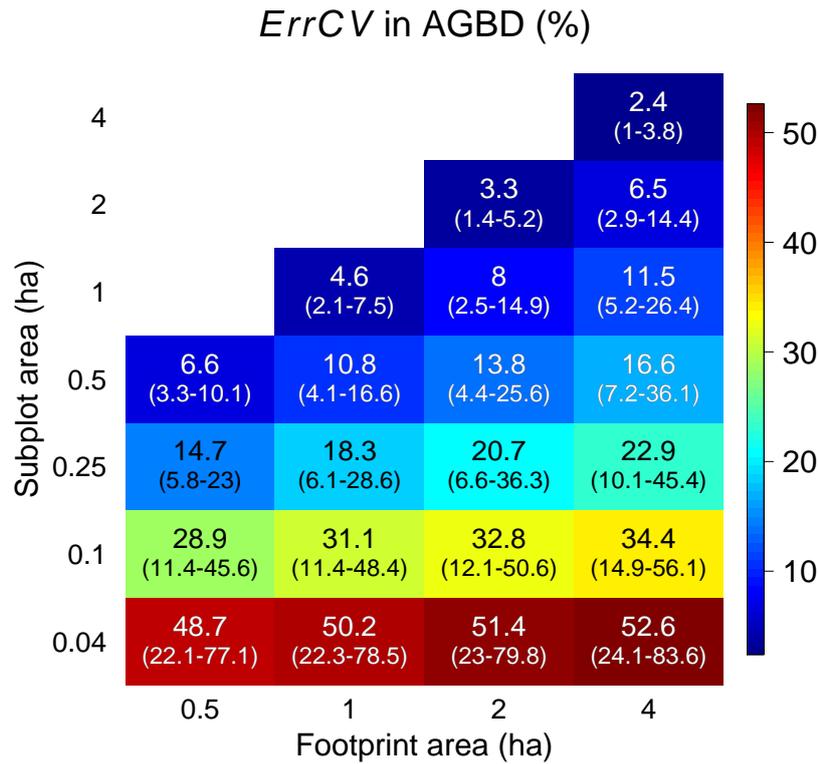
863 rho correlation of the elevation range with the wavelet variance for different spatial scales. P-

864 values of the Spearman's rho correlation tests are provided within the panel.

865

ErrCV in AGBD (%)

Figure 6. Expected sampling errors when the calibration/validation plots and the remote sensing footprint differ in shape and size. The remote sensing footprint is assumed circular, and subplots are assumed to be square to simulate the spatial mismatch between the remote sensing signal and the calibration plot (Fig. 2). The mean *ErrCV* in AGBD estimates across all sites (n=30) is given within the figure and the range of *ErrCV* across sites is given in parentheses below the mean.

a Dilution bias

b OLS corrections and alternatives

874
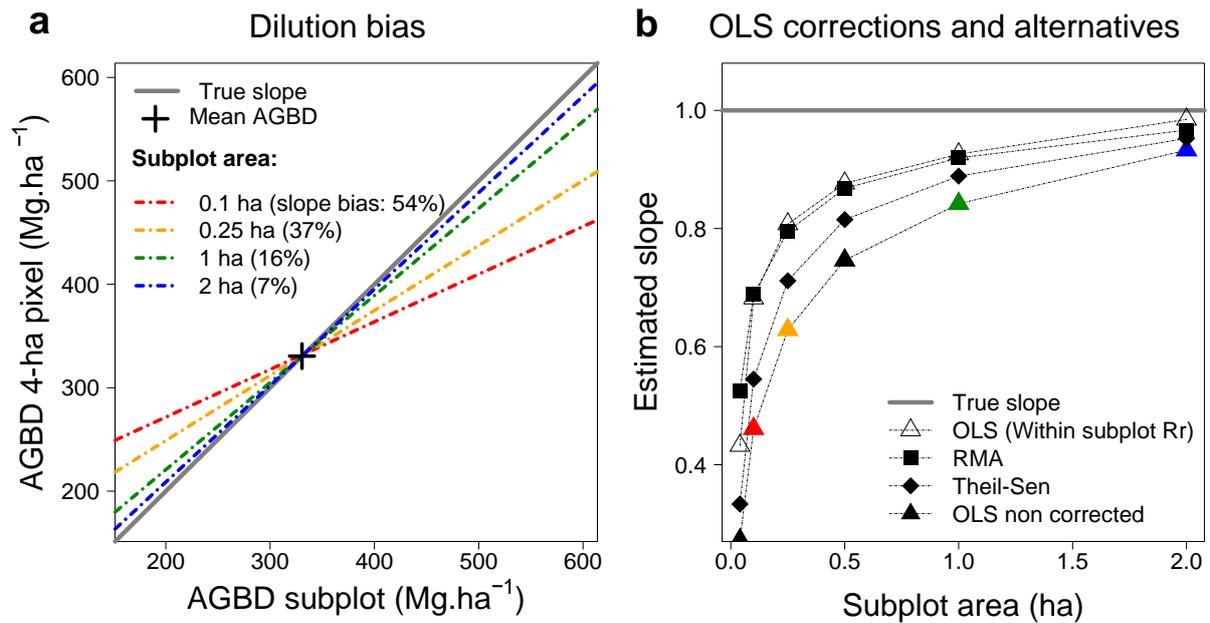
Figure 7. Propagation of field sampling error to remote sensing products: the dilution bias. (a) The mean regression lines obtained from an OLS linear regression between the AGBD estimated within 4-ha pixels randomly established in large plots (n=60, dependent variable) and variable-size subplots located within these pixels (independent variable) differ depending on subplot areas (see key), and are biased with respect to the true slope of one (slope dilution biases associated with each subplot area are provided in parentheses). All the lines cross at the mean AGBD over all sites, because regression always correctly predicts the mean. However, the smaller the subplot, the more regressed to the mean the predictions. (b) Different potential correction methods (see key) result in improved estimates of the slopes, but still retain considerable bias. The points corresponding to the lines in panel (a) are shown with matching colors. The true slope of one, i.e. the slope that would have been obtained without bias, is illustrated by the solid grey line.