We have made some edits throughout the revised manuscript, and marked modifications regarding to referees' comments using the track change function in Word or font color change.

Reviewer 1:

1. *This study employed a well-established process-based forest model (PnET-CN) to evaluate the effect of clearcut on carbon (C) flux trajectory in two widespread plant functional types (deciduous broadleaf forest and evergreen needleleaf forest) in the upper Midwest region of Wisconsin and Michigan. The trajectory analysis of C flux after clearcut makes this study quite interesting. Results suggested that harvest have a big influence on early stage of forest succession, but only had little effects on late stage. The method used in this study is solid, and results met with expected recovery trajectory in forest ecosystems. The manuscript is concise and well written, and the topic falls within the scope of BG.*

   Author response: Thanks for spending time to review the paper and for your constructive comments.

2. *However, it is surprised to me that only one scenario was used. I understand that this study was designed to quantify the C flux trajectory following clearcut. Since no ecological model can exactly reproduce the natural system, it is maybe more interesting to compare how forest recovery trajectories vary after different management alternatives. But I realized that this will completely change the objective of this study. And also, PnET-CN may have limited ability to simulate different harvest regimes and forest regeneration, and I will leave this comment to authors for their future exploration.*

   Author response: We agree that it is interesting and useful to quantify the effects of different management operations (e.g., thinning and selective cutting) on the C fluxes and stocks using ecosystem models. Currently, very few biogeochemical models can directly include these processes. Rather, we focus on harvest intensity to represent partial and clear cutting with a sensitivity analysis. Our analysis shows how forest responds to clearcuts (100%) and two partial cutting (80% and 60% removal of living trees in terms of biomass) (section 3.3), although we are not able test the results of the alternative management scenarios against observations in the region. More field measurements in carbon fluxes and stocks in managed forest sites (e.g., forest thinning: Saunders et al., 2012, Dore et al., 2012) are expected to be useful for model testing in the future.

3. *P8791 L23-26: See the latest debates on respiration, GPP, NPP/GPP trajectories during*
    succession (Tang et al., 2014 PNAS).

    Author response: Tang et al recently argue that the shaper decline in GPP than respiration
    might explain the age-related decline in NPP by using a global forest carbon database. We
    agree that the derived GPP and NPP changes with age, which is similar to traditional
    hypotheses. Therefore, we added the reference in the previous statement (see section 1).

4. *The in-situ measurement data reflect real world condition, which was affected by changes*
    in climate, atmospheric composition (e.g., CO2 rising, N deposition), and disturbance,
    while model simulation only included some of these factors (e.g., N deposition and
    disturbance). As I understand, climate data was used repeatedly from 1981 –2010. It is not
    clear how CO2 was parameterized in the PnET-CN. Does this influence your validation
    results?

    Author response: For the period 1959-2010, we used the CO_2 concentrations data from
    Mauna Loa. For the time period 1901–1958, we derived the time series of the historical
    atmospheric CO_2 mixing ratio using a spline fit to the ice-core record of Etheridge et al.
    (Etheridge et al. 1996), as described by McGuire et al. (McGuire et al. 2001) and used by
    Xiao et al. (2009). We have described this in the revised manuscript (Section 2.3).

    The simulations in this study ended in 2010, and the monthly CO_2 concentrations are
    shown in the following figure (Fig. R1). CO_2 fertilization has little effects on our testing
    sites during our simulation period. We did a sensitivity test for CO_2 fertilization (see
    Supplementary figure S3 and section 3.3 added in the revision). There was no noteworthy
    change in simulated trajectories in carbon fluxes and stocks.

5. *For the sensitivity analysis, were dead wood removal fraction and soil removal fraction*
    also changed, and how? Soil removal fraction may have a big influence (e.g., Peters et al.,
    2013, Ecosystems) on C flux and how these parameters was set deserved to be explained.

    Author response: The other disturbance parameters (i.e., wood removal fraction, soil
    removal fraction) were kept unchanged in the sensitivity analysis. We only focused on the
    harvesting parameter (stand mortality) in this paper (Fig. 8). Harvested wood removal
    fraction was parameterized according to expert's experience, ranging from 60%-90%.

We agree that soil carbon removal influence soil C fluxes and secondary forest growth. But several meta-analyses (Johnson and Curtis, 2001; Yanai et al., 2003; Nave et al., 2010) have showed that no significant change in soil carbon pools after harvest operations. Furthermore, the sensitivity analysis conducted by Peters et al, 2013 has shown the model sensitivity to soil removal. We have discussed this in the revision (Section 2.4, 3.3, 4.1 and 4.2). We have now added an analysis of sensitivity to soil removal fraction.

6. *It is not clear to me how CO2 concentration trend was parameterized in the model?*

Author response: We added sentences to describe the CO_2 concentration data used in this study (Section 2.3). We also added a figure (Fig. S3) to show the model sensitivity to CO_2 fertilization in the supplementary material.

7. *Figure 3: Do you have validation results for NEP, GPP, and ER?*

Author response: The figures 1 and 2 have shown the validation results for carbon fluxes. Since the model has been applied for major forest types in Midwestern USA (Peters et al., 2013; Ryu et al., 2012), most parameters were left unchanged in this study to facilitate the generalization (see Table S1).

References:

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Johnson, D.W. and Curtis, P.S., 2001. Effects of forest management on soil C and N storage: meta analysis. For. Ecol. Manage., 140(2-3): 227-238.

McGuire, A.D. et al., 2001. Carbon balance of the terrestrial biosphere in the Twentieth Century: Analyses of CO2, climate and land use effects with four process-based ecosystem models. Glob. Biogeochem. Cycle, 15(1): 183-206.

Nave, L.E., Vance, E.D., Swanston, C.W. and Curtis, P.S., 2010. Harvest impacts on soil carbon storage in temperate forests. For. Ecol. Manage., 259(5): 857-866.

Peters, E.B., Wythers, K.R., Bradford, J.B. and Reich, P.B., 2013. Influence of disturbance on temperate forest productivity. Ecosystems, 16(1): 95-110.

Ryu, S.-R., Chen, J., Noormets, A., Bresee, M.K. and Ollinger, S.V., 2008. Comparisons between PnET-Day and eddy covariance based gross ecosystem production in two Northern Wisconsin forests. Agric. For. Meteorol., 148(2): 247-256.

Saunders, M. et al., 2012. Thinning effects on the net ecosystem carbon exchange of a Sitka spruce forest are temperature-dependent. Agric. For. Meteorol., 157(0): 1-10.

Xiao, J. et al., 2009. Twentieth-century droughts and their impacts on terrestrial carbon cycling in China. Earth Interactions, 13(10): 1-31.

Yanai, R.D., Currie, W.S. and Goodale, C.L., 2003. Soil carbon dynamics after forest harvest: An ecosystem paradigm reconsidered. Ecosystems, 6(3): 197-212.

Reviewer 2:

We have addressed Reviewer 2's comments point by point and marked modifications with orange color.

8. *Although the authors have done a good job in data collecting, study design, modelling, data analyses and writing, I still have some major concerns about the MS in your discussion section. The one thing is that authors did not compare their results comprehensively with other studies all over the world. E.g. the trajectory analyses of other disturbances such as fire, etc. And other case studies for rainforests or boreal forests.*

> **Author response:** We agree that it is valuable to discuss the carbon trajectories for other forest types. In fact, we have compared our results with other forest types such as boreal forests (e.g., jack pine and black spruce), temperate forests (e.g., Douglas-fir forests in British Columbia, slash pine forests in Florida, white pine forests in Ontario) in the discussion (section 4.1), although we excluded a couple of chronosequence studies that comprise only two age groups.

> It is also interesting to discuss the effects of fire on successional trajectories, although our study focuses on harvesting. For example, the simulated LAI change with age is similar with previous fire chronosequence study (Goulden et al., 2011). We have discussed several case studies for fire disturbances in the revision (Section 4.1).

> Goulden, M.L. et al., 2011. Patterns of NPP, GPP, respiration, and NEP during boreal forest succession. Glob. Change Biol., 17(2): 855-871

9. *Moreover, for a MS to be submitted to a Special Issue "Impacts of extreme climate events and disturbances on carbon dynamics", maybe authors should be focused more on extreme ones not only regularly anthropogenic disturbances(such as clearing, etc.), they should discuss about extreme climate or other events' effect on carbon dynamics and their trajectory. You should add more references about extreme disturbance to make your MS fallen well in the scope of this special issue.*

> **Author response:** We have communicated with the guest editors of this special issue regarding the scope of the special issue. "Extreme climate events and disturbances" actually refers to "extreme climate events" and "disturbances", which means that this issue is interested in disturbances in general as well as extreme climate events. We have also

briefly discussed the importance of extreme climate events to make the manuscript at the end of the Discussion section (Section 4.3).

10. *You should reorganize your objectives to be constant with your results and discussion. One is to evaluate PnET-CN's performance for temperate forests and the other is to study the trajectory variation of carbon dynamics after clearing and their density.*

Author response: We first tested the model against field measurements along the chronosequence sites (section 3.1 and 4.3), analyzed the simulated and the hypothesized carbon trajectories (section 3.2 and 4.1), and compared the difference of successional trajectories in carbon fluxes and stocks between two forest types associated with harvest intensity (section 3.2, 3.3 and 4.2). We will emphasize the three objectives in the revision.

11. *I suggested that you added a subsection to compare your results with different disturbances as well as with different forest types.*

Author response: We have compared our results with different disturbances as well as with different forest types. Please see our response to your comment 1.

Reviewer 3:

We have addressed Reviewer 3's comments in the revised manuscript point by point and marked using track changes:

12. *This manuscript examined the ability of an ecosystem model PnET-CN to capture the trajectories of forest C dynamics after a stand-replacing disturbance and two hypotheses in two northern forest chronosequences. They showed that PnET-CN can reasonably simulate stand characteristics and capture the changes of C fluxes after clearcut. The work is good and interesting, and their conclusions are clear. It is well within the scope of BG.*

Author response: Thanks for reviewing the manuscript and for your constructive comments.

13. *It is better to add more information on the two plant functional types, especially main tree species composed.*

Author response: We have added tree species information about the two plant function types in section 2.1 and Table 1.

14. *Page 8798.the MS mentioned "The parameter values used in this study are given in Table 2", but I did not find it. Same as in page 8802 on maximum relative growth rate (Table 2). I guess it should be Table S1 instead.*

Author response: We have corrected the typos.

15. *Table 2: Is the statistics calculated from all data (DBF and NEF)? It is interesting to see it separately since you simulated two plant functional types. Why have they different sample size (n)?*

Author response: Our post-processing code removed a couple of data points that ER was estimated to be zero. We have corrected it in the revision.

16. *The discussion on the difference between DBF and NEF is more attractive. It seems to meet the objective on testing the role of forest composition on successional question on trajectories of forest C dynamics. However, I do not think "forest compostion" is the right word in this case, it is better to use plant functional types consistently.*

    Author response: We have changed the term of "forest composition" to "plant function types (PFTs)" in the revision.

17. *Type error in page 8792. Please delete comma in"Noormets et al., (2007) reported: ".*

    Author response: It has been deleted.

Quantifying legacies of clearcut harvesting on carbon fluxes and biomass carbon stock in northern temperate forests

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Abstract

Stand-replacing disturbances including harvests have substantial impacts on forest carbon (C) fluxes and stocks. The quantification of these effects is essential for better understanding of forest C dynamics and informing forest management in the context of global change. We evaluated a process-based forest ecosystem model, PnET-CN, for how and by what mechanisms clearcuts alter ecosystem C fluxes, aboveground C stocks (AGC), and leaf area index (LAI). We compared the effects of stand-replacing harvesting on C fluxes and stocks using two chronosequences of eddy covariance flux sites for deciduous broadleaf forests (DBF) and evergreen needleleaf forests (ENF) in the Upper Midwest region of northern Wisconsin and Michigan, U.S.A. The average normalized root mean square error (NRMSE) and the Willmott index of agreement (*d*) for carbon fluxes, LAI, and AGC in the two chronosequences were 20% and 0.90, respectively. Simulated gross primary productivity (GPP) increased with stand age, reaching a maximum (~1200–1500 g C m^{-2} yr^{-1}) at 11–30 years of age, and leveled off thereafter (~900–1000 g C m^{-2} yr^{-1}). Simulated ecosystem respiration (ER) for both plant function types

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(PFTs) was initially as high as ~700–1000 g C m^{-2} yr^{-1} in the first or second year after harvesting, decreased with age (~400–800 g C m^{-2} yr^{-1}) before canopy closure at 10–25 years of age, and increased to ~800–900 g C m^{-2} yr^{-1} with stand development after canopy recovery. Simulated net ecosystem productivity (NEP) for both PFTs was initially negative with net C losses of ~400–700 g C m^{-2} yr^{-1} for 6–17 years after clearcuts, reached the peak values of ~400–600 g C m^{-2} yr^{-1} at 14–29 years of age, and became stable and a weak C sink (~100–200 g C m^{-2} yr^{-1}) in mature forests (>60 years old). The decline of NEP with age was caused by the relative flattening of GPP and gradual increase of ER. ENF recovered slower from net C source to net sink and lost more C than DBF, suggesting that ENF is likely slower to recover to full C assimilation capacity after stand-replacing harvests arising from slower development of photosynthesis with stand age. Model results indicated that increased harvesting intensity would delay the recovery of NEP after clearcut, but it had little effect on C dynamics during late succession. Future modeling studies of disturbance effects will benefit from the incorporation of forest population dynamics (e.g., regeneration and mortality), relationships between age-related model parameters and state variables (e.g., LAI), and silvicultural practices into the model.

1   Introduction

Disturbance has been widely recognized as a key factor influencing ecosystem structure and function at decadal to century scales (Magnani et al., 2007; Williams et al., 2012; Kasischke et al., 2013). Harvest is an important anthropogenic disturbance shaping North American forest landscapes. Approximately 61,000 km^2 of forests were affected by harvests every year during the 2000s (Masek et al., 2011). Harvests affect forest age structure and alter the forest carbon (C) balance (Magnani et al., 2007; Pan et al., 2011; Williams et al., 2012). Quantifying the legacies of harvest disturbances under the context of climate change is essential for predicting forest C dynamics, informing climate policy-making, and improving forest management. Here, we focus on assessment of an ecosystem model C cycle response from one type of harvest, the clear-cut.

Harvests transfer living biomass C to harvested wood C and litter C, resulting in successional changes in C fluxes and stocks. Leaf biomass increases rapidly in secondary succession and then typically stabilizes at a certain level that is determined by light, water, nutrient availability, and forest type (Sprugel, 1985). Gross primary productivity (GPP) thus increases gradually over time, reaches maximum in middle age, and in response to nutrient

1   limitations and aging responses slightly declines thereafter (Odum, 1969; Chapin et al., 2002;

2   Tang et al., 2014). The successional change in plant respiration (autotrophic respiration) after

3   stand-replacing harvesting is similar to that of GPP, although C use efficiency (the ratio of net

4   primary productivity to GPP, NPP/GPP) could decline with forest age (DeLucia et al., 2007). As

5   a result, living tree biomass C gradually increases following a typical logistic growth curve

6   (Odum, 1969; Sprugel, 1985).

7   Heterotrophic respiration following stand-replacing harvesting could be stimulated at the

8   beginning of stand development because the removal of trees alters the environmental conditions

9   (e.g., soil temperature, moisture, and nutrients) and possibly leads to changes in litter quantity

10   depending on harvest types (e.g., stem-only harvesting) (Chapin et al., 2002). Heterotrophic

11   respiration is expected to gradually decrease thereafter because the regrowing forest reduces net

12   radiation, water, and nutrient availability to the soil (Chapin et al., 2002) and the amount of

13   decomposable soil organic matter from the prior forest and harvest residue (e.g., litter, coarse

14   woody debris, and soil organic C) also gradually decreases. Over time, however, heterotrophic

15   respiration could again be enhanced because of accumulation of woody debris and litter with

16   stand development. This theorized successional trajectory in ecosystem respiration (ER; the sum

17   of autotrophic and heterotrophic respiration) may also be strongly influenced by harvest types

18   and forest composition. Unlike GPP or NPP, the trajectory of heterotrophic respiration (and

19   consequently total ecosystem respiration) with age is not as well-understood (Amiro et al., 2010).

20   Observational studies to date have shown that forest ecosystems generally become C sources

21   (negative net ecosystem productivity, NEP) immediately following stand-replacing harvests,

22   approach the maximum NEP as they mature, and then experience a gradual decline in NEP

23   thereafter (e.g., Law et al., 2003; Gough et al., 2007; Goulden et al., 2011), following the

24   trajectories hypothesized by Odum (1969) and Chapin et al. (2002).

25   The changes of C fluxes and stocks after harvesting have been examined in many forest

26   ecosystems using ecological measurements (e.g., eddy covariance observations)from

27   chronosequences using a space-for-time substitution approach (e.g., Law et al., 2003; Gough et

28   al., 2007; Goulden et al., 2011). The trajectories and amplitude of C fluxes and stocks vary with

29   forest ecosystem types (Amiro et al., 2010). For example, Noormets et al. (2007) reported that a

30   young red pine (*Pinus resinosa*) stand at 8 years of age was a net C sink (313 ± 14 g C m^{-2} yr^{-1}),

31   but a young hardwood site at age of 3 was a net C source (-128 ± 17g C m^{-2} yr^{-1}) over the growing

1 season in northern Wisconsin, U.S.A. Young stands in northern Wisconsin may become net C

2 sinks within 10-15 years after harvesting (Noormets et al., 2007). More rapid recovery after

3 stand-replacing harvesting (< 6 years) was found for temperate forests in northern Michigan

4 (Gough et al., 2007). These studies have produced a wealth of information on ecosystem C

5 dynamics after stand-replacing disturbances, and this information can be translated to more

6 process-based and quantitative understanding of disturbance effects on the C cycle using

7 ecosystem models (Goulden et al., 2011). Process models require evaluation on how source/sink

8 transition and long-term carbon flux dynamics respond to differences in vegetation type, harvest

9 intensity, and age since clearing.

10   Although using the chronosequence approach to evaluate the changes of ecological

11 processes with age after disturbances is attractive, this approach is often limited by the lack of

12 biological and climatic data (Yanai et al., 2003; Bond-Lamberty et al., 2006) and full

13 representation of stand development stages. Process-based ecosystem models provide a means of

14 quantifying the effects of disturbances on C dynamics under changing climate over various

15 spatial and temporal scales. Ecosystem models have been used to assess the effects of clearcuts

16 and climate change on forest C dynamics at the stand/ecosystem (e.g., Bond-Lamberty et al.,

17 2006; Grant et al., 2009; Wang et al., 2012b) or regional scales (Desai et al., 2007; Dangal et al.,

18 2014). Moreover, ecosystem models can also be used to assess forest C dynamics under various

19 scenarios of climate change and harvesting regimes (e.g., Albani et al., 2006; Peckham et al.,

20 2012), since these models have been developed based on physiological, biogeochemical, and

21 ecological theories. However, few studies have used ecosystem models to examine the changes

22 of C fluxes and stocks with stand regrowth after stand-replacing disturbances for forest

23 chronosequences.

24   The objectives of this study were to evaluate the ability of an ecosystem model to capture

25 the trajectories of forest C dynamics after stand-replacing harvests for two northern temperate

26 plant function types (PFTs: deciduous broadleaf forests, DBF; evergreen needleleaf forests,

27 ENF), to examine which processes influence successional trajectories in these ecosystems, and to

28 test the role of plant function type on successional trajectory of C fluxes. We applied a process-

29 based forest ecosystem model, PnET-CN (Aber et al., 1997; Ollinger et al., 2002), to simulate

30 the effects of clearcut on forest C dynamics, and evaluated the simulated C fluxes and stocks for

31 both PFTs using in-situ measurements (e.g., eddy covariance observations and aboveground

4

1   biomass C, AGC). We hypothesized that (1) both DBF and ENF will have similar successional

2   patterns in C fluxes (GPP, ER, and NEP) and aboveground biomass C stocks after stand-

3   replacing harvests, but (2) DBF will recover faster than ENF from a net C source to a net C sink

4   and lose a smaller amount of C (negative NEP) following a stand-replacing harvest.

5   **2   Methods**

6   **2.1   Study sites and data**

7   Our study sites consist of 8 eddy covariance sites in the Upper Midwest region of

8   northern Wisconsin and Michigan (Chen et al., 2008; Table 1). The study area is characterized

9   by a humid-continental climate with hot summers and cold winters. The mean annual

10   temperature is 4.4 ˚C and the mean annual precipitation is 768.9 mm (as measured between 1981

11   and 2010 at Rest Lake weather station, 46.12˚ N 89.87˚ W, http://www.ncdc.noaa.gov). The

12   dominant soil type is glacial sandy loam and loamy tills (Noormets et al., 2008). The region has

13   been strongly influenced by forest industry. Most forest stands less than 100 years old in this

14   region regenerated following harvesting operations (Amiro et al., 2010).

15   Our sites consist of four DBF sites (YHW, IHW, WIC, and UMBS) and four ENF sites

16   (YRP, YJP, IRP, and MRP). The four DBF sites range from 3 to 86 years in age and constitute a

17   chronosequence. Dominant tree species are maple (*Acer* spp.), basswood (*Tilia american*), birch

18   (*Betula allghaniensis*), ash (*Fraxinus* spp.) and aspen (*Populus termuloides*). The four ENF sites

19   also represent a chronosequence with stand age ranging from 8 to 66 years. Red pine and jack

20   pine (*Pinus banksiana*) are dominant tree species in the four ENF sites. The two

21   chronosequences, most sites were initiated by stand-replacing harvests. We obtained monthly C

22   fluxes (observed NEP and its inferred data products GPP and ER) from AmeriFlux

23   (http://public.ornl.gov/ameriflux/) for the eight eddy covariance flux tower sites (Table 1).

24   Harmonized level 4 data were used in this study. These flux data have been described and used

25   in our previous studies (e.g., Noormets et al., 2007; Chen et al., 2008; Desai et al., 2008; Xiao et

26   al., 2011; Xiao et al., 2014). We also obtained LAI and AGC data from the literature for each site

27   (Table 1).

2.2   Model description

1

2   The PnET-CN model is a process-based forest ecosystem model that can simulate C,

3   nitrogen, and water dynamics at monthly time steps. PnET-CN is driven by temperature,

4   precipitation, photosynthetically active radiation (PAR), wet and dry nitrogen deposition, and

5   atmospheric CO_2 concentration (Aber and Driscoll, 1997; Aber et al., 1997; Ollinger et al.,

6   2002). The model has been applied and tested in the USA and Europe for simulating the effects

7   of climate variability, rising atmospheric CO_2, ozone pollution, and disturbance on ecosystem

8   processes and functions (e.g., Aber et al., 2002; Pan et al., 2009; Peters et al., 2013).

9   One of the unique features of PnET-CN is to simulate potential photosynthesis using

10   foliar nitrogen concentration and light use efficiency in a multilayered canopy (Aber and

11   Federer, 1992). Photosynthesis is then constrained by air temperature, vapor presser deficit, and

12   soil water availability for simulating actual GPP. The effects of elevated CO_2 concentration on

13   leaf photosynthetic rates are calculated using constant ratios of leaf internal to ambient CO_2

14   concentration (C_i/C_a) (Ollinger et al., 2002). PnET-CN incorporates a total of seven C pools, five

15   of which are structural C pools (foliage, woods, fine roots, woody debris, and soil organic

16   matter) and two of which are non-structural C pools stored in woods and fine roots.

17   Photosynthetic production is allocated to each living plant component (i.e., foliage, woods, and

18   fine roots) and to growth and maintenance respiration. Living biomass is transferred to dead

19   woody biomass and/or to soil organic C through leaf and root turnover, tree mortality, and

20   disturbance. The decomposition of coarse woody debris is a constant fraction of its C content.

21   The decomposition of soil organic C is calculated as a function of maximum decomposition rate

22   and effects of temperature and soil moisture.

23   PnET-CN includes a complete nitrogen cycle, and simulates nitrogen mineralization and

24   nitrification, plant nitrogen uptake, allocation, and leaching losses. Nitrogen depositions are

25   imposed into corresponding soil nitrogen pools (NH_4 and NO_3). As with C pools, nitrogen is

26   divided into five structural pools (foliage, woods, fine roots, woody debris, and soil organic

27   matter) and one non-structural nitrogen pool stored in the trees. C and nitrogen cycles interact

28   closely in the model. High leaf nitrogen concentration increases net photosynthesis rate in the

29   absence of water stress, thereby resulting in the high demand for non-structural nitrogen in plant

30   tissues (Aber et al., 1997). When plant non-structural nitrogen is low, plant nitrogen uptake

efficiency from available soil mineral nitrogen is increased in the model (Aber et al., 1997). In addition, high C: N ratios in biomass, litter, and soil organic matter reduce net mineralization rates. In general, the nitrogen cycle in the model is governed by a negative feedback loop.

The model also simulates key hydrological processes including rainfall interception, evaporation, transpiration, surface runoff, and drainage at each time step. Rainfall interception is treated as a constant fraction of precipitation. Transpiration is estimated based on water use efficiency. Surface runoff is calculated as a constant fraction of the difference between precipitation and evaporation. Drainage is estimated when potential soil water exceeds soil water holding capacity.

Prescribed disturbance events can be simulated in the model through four parameters: disturbance year, disturbance intensity, biomass removal fraction, and the loss rate of soil organic matter. In this study, when stand-replacing disturbance events occur, a uniform plant function type was assumed to be regenerated on-site. For the first year after clearcuts, minimum leaf area index (LAI) of 0.1 was assumed to regulate maximum potential foliage mass that controls leaf production. The photosynthetic production is transported to plant non-structural C pool where C could be allocated to leaves, stems, and roots. There is therefore no need for initialization (e.g., stand density) after disturbances in the model. More details about the model structure and processes have been described elsewhere (Aber et al., 1997; Ollinger et al., 2002).

2.3   Model inputs

The model inputs include temperature, precipitation, PAR, wet and dry nitrogen deposition, atmospheric CO_2 concentrations, and disturbance history. The climate data used in all simulations were derived from the Daymet database (Thornton et al., 2012). For each site, monthly maximum temperature, minimum air temperature, and precipitation were calculated from the daily Daymet data for the period 1981-2010. PAR (mol $m^{-2} s^{-1}$) was estimated from solar radiation (RAD, MJ m^{-2} day^{-1}) using the empirical relationship (PAR = 2.05 RAD) (Aber et al., 1996). The data from 1981 to 2010 were repeated as needed to create the time series from 1850 to 1980.

Annual rates of wet and dry nitrogen deposition were obtained from the United States Environmental Protection Agency (EPA; http://java.epa.gov/castnet/clearsession.do). The nitrogen deposition rates were measured at the Wellston station (44.22° N; 85.82° W) for the

7

1 period 1994-2011. We also obtained the nitrogen deposition rates in 1860 estimated by Galloway

2 et al. (2004). For each year prior to 1994, we used an exponential ramp function to estimate the

3 annual deposition rates by interpolating the historical (1860) and current nitrogen deposition

4 rates. Monthly wet deposition rates, needed for the model, were generated from annual wet

5 nitrogen deposition through the weighted coefficients (the ratio of monthly precipitation to total

6 precipitation from March to November). We assumed that there is no wet nitrogen deposition in

7 the winter. The soil water holding capacity in the rooting zone (100 cm) for each site was derived

8 from the gridded multi-layer soil characteristics dataset (STATSGO, Miller and White, 1998).

9 For the period 1959-2010, we used the CO_2 concentrations data from Mauna Loa. For the time

10 period 1901–1958, we derived the time series of the historical atmospheric CO_2 mixing ratio

11 using a spline fit to the ice-core record (Etheridge et al., 1996), as described by McGuire et al.

12 (2001) and used by Xiao et al. (2009). We used the CO_2 concentration in 1901 for the simulation

13 period prior to 1901 and spin up.

14   For each site, we prescribed the disturbance events using the site disturbance history

15 (Table 1). For each stand-replacing harvest, stand mortality was assumed to be 100%. The

16 merchantable wood removal (biomass removal out of the ecosystem) fraction was assumed to be

17 0.8 in this study. The soil removal fraction was assumed to be zero, given that the content of soil

18 organic C might not be considerably affected by harvesting (Johnson and Curtis, 2001; Yanai et

19 al., 2003). We also conducted a sensitivity analysis to these assumptions as described below in

20 section 2.4.

21 **2.4   Parameterization, initialization, validation, and sensitivity analysis**

22   PnET-CN has been parameterized and tested for temperate DBF (Aber et al., 1997;

23 Ollinger et al., 2002; Peters et al., 2013), temperate ENF(Aber et al., 1997; Peters et al., 2013),

24 and mixed forests (Aber et al., 1997) for forest productivity, net nitrogen mineralization, and

25 foliar nitrogen concentrations. The parameter values used in this study are given in supplement

26 Table S1. To apply the model to the transient simulation period (1860-2010), a 200-year spin up

27 run was conducted to ensure that the equilibrium (Δ NEP<10 g m^{-2} month $^{-1}$ and Δ soil organic C

28 <1%) was reached for each chronosequence site. The climate normals (1981-2010), pre-industry

29 nitrogen deposition rates, and historical CO_2 concentrations were used for the spin up runs.

1 To examine the stand-replacing harvest legacies, we conducted all simulations using the

2 site disturbance history (Table 1), vegetation parameters (Supplement Table S1), climate,

3 nitrogen deposition, and atmospheric CO_2 for each of the chronosequence sites. The model

4 simulations were evaluated against C fluxes (GPP, ER, and NEP), AGC, and LAI data collected

5 at the eddy covariance flux sites. We used two statistical measures to evaluate the overall model

6 performance: the normalized root mean square error (NRMSE) and the Willmott index of

7 agreement. The NRMSE (Eq. (1)) was used to assess the difference between predicted (P) and

8 observed (O) variables, and can be expressed as:

9 $$NRMSE = (O_{max} - O_{min})^{-1} \left[\frac{\sum_{i=1}^{n}(P_i - O_i)^2}{n} \right]^{0.5} \times 100\% \qquad (1)$$

10 where O_{max} and O_{min} are the maximum and minimum observed values, respectively; i is the ith

11 observation; and n is the total number of observations. A value close to 0 indicates perfect

12 agreement and a value of 100% suggests poor agreement. The Willmott index of agreement (d) is

13 an indicator of modeling efficiency and is expressed as:

14 $$d = 1 - \left[\frac{\sum_{i=1}^{n}(P_i - O_i)^2}{\sum_{i=1}^{n}(|P_i| + |O_i|)^2} \right] \qquad (2)$$

15 A value of 1 indicates perfect agreement and a value near 0 indicates weak agreement (Willmott,

16 1982).

17 The sensitivity of ecosystem C dynamics to changes in harvesting practices during the

18 secondary succession was assessed using sensitivity analysis. The model was run at WIC and

19 MRP for 100 years after scenario harvests in 1910 using the same climate data sequence.

20 Sensitivity scenarios involved applying the stand mortality (80% and 60%, compare to 100% in

21 the model test), soil organic matter loss (20% and 40%, compare to zero in the model test) to

22 reveal effects of different harvest intensity and soil organic matter loss scenarios on C dynamics.

23 We also tested the model sensitivity to CO2 fertilization for evaluating potential climate change

24 effects.

25

3 Results

3.1 Evaluation of modeled carbon fluxes and stocks

The simulated C fluxes were generally consistent with eddy covariance derived C fluxes for both DBF and ENF sites (Figs. 1 and 2). The NRMSE between simulated and tower fluxes (GPP, ER, and NEP) were between 10-21% (Table 2). The Willmott index of agreement between simulated and tower C fluxes for both plant function types ranged from 0.91 to 0.94 with the exception of NEP (d= 0.73, n= 235). The model underestimated GPP for the DBF sites and predicted ER fairly well for all DBF sites except for the intermediate-aged hardwood site, IHW. As a result, the model underestimated NEP for most DBF sites. For IHW, the model substantially underestimated both GPP and ER but predicted NEP fairly well. For the ENF sites, the model underestimated GPP. The model predicted ER fairly well for YRP (8 years old), YJP (15-16 years old) and IRP (23 years old), but overestimated ER for the older MRP sites. Thus, the model underestimated NEP for the ENF sites.

The simulated and observed stand characteristics (LAI and AGC) showed good agreement (Table 2 and Fig. 3). The model slightly underestimated LAI for the young forest sites, and overestimated LAI for the mature forest sites. Generally, the model overestimated AGC for the mature forest sites. The NRMSE was 28% for AGC and 31% for LAI. The Willmott index of agreement was 0.95 and 0.96 for AGC and LAI, respectively. Overall, the model evaluation metrics indicated that the model performed better in the DBF sites than in the ENF sites.

3.2 Legacy of clearcut on carbon fluxes and stocks

PnET-CN generally captured the changes of C fluxes following the clearcuts for each chronosequence site (Fig. 4). The predicted annual GPP generally increased with time since disturbance and approached the peak values (~1200–1500 g C m^{-2} yr^{-1}) between 11 and 26 years of age and between 29 and 30 years of age for the DBF (IHW, WIC, and UMBS) and the ENF (IRP and MRP) sites, respectively; thereafter, the forest stands reached maturity and GPP became relatively stable with mean values of 940–1000 g C m^{-2} yr^{-1}.

Predicted annual ER was initially as high as 860–1030 and 710–860 g C m^{-2} yr^{-1} within the first two years for the DBF and the ENF sites, respectively. During canopy recovery, predicted ER generally decreased to 620–780 g C m^{-2} yr^{-1} between 10 and 25 years of age for the

10

1   DBF sites and to 360–380 g C m^{-2} yr^{-1} between 14 and 17 years of age for the ENF sites (Fig. 4).

2   For forest age older than 60 years, the predicted annual ER for both PFTs showed a relatively flat

3   pattern, contrary to theoretical expectations, arising from the little change of both autotrophic

4   and heterotrophic respiration with age (Supplement Fig. S1). Average annual ER for mature

5   forests was 810–880 and 780 g C m^{-2} yr^{-1} for the DBF sites (WIC and UMBS) and the ENF

6   (MRP) site, respectively.

7   As expected, the ratio of annual GPP to annual ER (GPP: ER) simulated by PnET-CN

8   was low during the early years after clearcutting for both DBF and ENF (Fig. 5). Within ~6 years

9   for the DBF sites and ~17 years for the ENF sites, the GPP:ER ratio gradually increased and its

10   average value became larger than 1 (NEP>0). The simulated peak GPP:ER ratio for DBF (1.6)

11   occurred at 18 years after stand-replacing harvests, and the simulated peak ratio for ENF was 1.8

12   at 26 years. After those peaks, the ratio became relatively stable, with the mean values of 1.1 and

13   1.2 for mature DBF and mature ENF, respectively.

14   The model predicted negative NEP (C source) for the 6 and 17 years after stand-replacing

15   harvests for the DBF and the ENF, respectively (Fig. 4). The simulated peak annual net C loss

16   occurred in the first or second year after clearcutting. The average C loss was 530–710 g C m^{-2}

17   yr^{-1} for the DBF sites and 380–400 g C m^{-2} yr^{-1} for the ENF sites. The total C loss was 3.2–4.3

18   and 6.4–6.9 kg C m^{-2} for the DBF and the ENF sites, respectively. The maximum net C gain was

19   387–433 g C m^{-2} yr^{-1} at 14-26 years of age for the DBF sites (WIC and UMBS) and was 567–

20   602 g C m^{-2} yr^{-1} at 29 years of age for the ENF sites (IRP and MRP). Simulated annual NEP

21   decreased thereafter and became as low as 120–180 g C m^{-2} yr^{-1} after 17–31 years for the DBF

22   sites and 170 g C m^{-2} yr^{-1} after 44 years for the ENF sites.

23   Forest canopy as measured by LAI gradually recovered over time following clearcuts

24   (Fig. 6). LAI fully recovered within 10–15 years after disturbance for the DBF sites and within

25   40 years of age for the ENF sites. The gradual recovery of LAI led to the gradual increase in

26   GPP and the subsequent increase in AGC (Fig. 7). In general, AGC recovered much more slowly

27   than C fluxes and LAI. The changes of simulated AGC followed the logistic growth curve with

28   slow accumulation in the early years, fast accumulation in the middle age, and slow

29   accumulation afterwards. The predicted LAI and AGC generally fell within the range of

30   observed values across two chronosequences (Figs. 3, 6, 7). For mature forests (>60 years of

1 age) in 2010, the DBF sites generally stored more C in aboveground biomass than the ENF sites

2 (10–12 *vs.* 8.5 kg C m^{-2}; Fig. 7).

3 **3.3 Sensitivity analysis**

4   Harvest intensity had little effect on the long-term C dynamics for both PFTs, but it had

5 sizeable effects during early succession (Fig. 8). Greater harvest intensity led to earlier rising

6 GPP (Fig. 8 a and f) and LAI (Fig. 8 d and i) but delayed reduction in ER (Fig. 8 b and g), ~~finally~~

7 resulting in later rising NEP (Fig. 8 c and g). High harvest intensity (e.g., 100% removal of living

8 trees) also directly reduced living tree AGC (Fig. 8 e and i). By reducing harvest intensity

9 parameter to 80% and 60% from 100% used in the original model, average annual NEP over 100

10 years for DBF decreased by 104 and 88 g C m^{-2} yr^{-1}, respectively. The increased remaining tree

11 biomass resulted in an increase in AGC about 12% and 16%, respectively, after a 100-yr harvest

12 cycle. For ENF average annual NEP decreased about 1% and AGC decreased nearly 6% for both

13 reduced harvest intensity scenarios. Increasing soil removal fraction parameter resulted in lower

14 GPP and ER along succession and lower NEP in middle succession for both DBF and ENF (Fig.

15 S2 a-c and f-h). Greater soil removal fraction promoted the leaf production of DBF in middle and

16 late succession (Fig. S2 d), but restricted the leaf production of ENF in late succession (Fig.S2 i).

17 Increasing soil removal fraction parameter (20% and 40% removal of soil organic matter)

18 strongly reduced living AGC (16% and 39%, respectively) for DBF (Fig. S2 e) but slightly

19 decrease living AGC (up to 5%) for ENF (Fig. S2 j). There were little effects of CO_2 fertilization

20 on carbon dynamics for both DBF and ENF in our sensitivity analysis (Fig. S3).

21 **4   Discussion**

22 **4.1   Carbon fluxes and stocks following clearcuts**

23   PnET-CN generally simulated the expected post-harvest trajectories in C fluxes (GPP,

24 ER, and NEP) and stock (LAI and AGC). The model was unable to simulate high GPP rates

25 estimated by eddy covariance in mature forests regardless of vegetation type, suggesting that

26 there is room for improvement in model simulation of secondary succession.

27   Our simulations showed that LAI increased rapidly first and then stabilized during the

28 following development stages, given that the model estimates foliage growth through the

29 parameter of maximum relative growth rate (Table S1) with the limitation of current foliage

1   biomass and resource availability. This modeled response is consistent with the previous finding

2   that foliage biomass increased rapidly after disturbance and then stabilized (Sprugel 1985). Our

3   chronosequence-based results are generally consistent with previous results. For example,

4   Goulden et al. (2011) observed that LAI along a chronosequence of boreal forest stands

5   increased rapidly from 0.3 m^2 m^{-2} 1 year after fire, and then generally leveled off at 5.3–7.2 m^2

6   m^{-2} from 23 to 154 years after stand-replacing crown fire. A modeling study based on a modified

7   version of Biome-BGC (Bond-Lamberty et al., 2005) also showed a similar successional change

8   in LAI for boreal DBF and ENF.

9   The simulated successional change in annual GPP for both PFTs generally followed the

10   trajectory hypothesized by Odum (1969). However, despite a slight decrease in GPP

11   hypothesized in Odum's trajectories, our simulations show a relatively flat GPP in mature forests

12   (Figs. 4 and 10). In the model, GPP tracks LAI in the absence of significant changes in light,

13   water or nutrient stress. As LAI stabilizes in mature forests, GPP also stabilizes. Our results are

14   consistent with previous studies showing relatively flat pattern in GPP after 20 years following

15   harvests in temperate pine forests in Florida (Clark et al., 2004), northern temperate DBF in

16   Wisconsin (Desai et al., 2008), and boreal jack pine forests in Saskatchewan (Zha et al., 2009).

17   Furthermore, Humphreys et al. (2006) reported continuous increases of GPP with increasing

18   forest age for temperate rainforests using three different stands at different stages of development

19   (2, 14, and 53 years of age) following clearcuts in British Columbia, Canada. However, northern

20   temperate ENF showed a small difference in GPP between young and mature sites (Noormets et

21   al., 2007; Desai et al., 2008). Desai et al. (2005) found that a nearby old-growth mixed forest had

22   slightly lower GPP and significantly higher ER than nearby DBF sites. Site-to-site variations in

23   species and soil fertility could result in variations in the successional trajectory of GPP after

24   clearcuts so that the observed trajectories may deviate from hypothesized or modeled

25   trajectories. In addition, our chronosequences lack old-growth sites and do not encompass the

26   full range of forest development stages, which limits the representativeness of the C flux and

27   stock trajectories derived from chronosequence studies based on eddy covariance or other

28   ecological observations (e.g., Clark et al., 2004; Humphreys et al., 2006; Noormets et al., 2007).

29   Our simulations show that annual ER for secondary temperate forests slightly declined in

30   the first ten years because of low autotrophic respiration at first after the removal of trees. Amiro

31   et al. (2010) reported that ER reduced in the very first year following harvests for a number of

eddy covariance flux sites over North America. Previous field studies showed that ER following

clearcuts increased with forest age (e.g., Humphreys et al., 2006; Zha et al., 2009), partly

supporting our results that ER slightly increased after the short decline period (10-25 years of

age) in northern temperate forests until the stands reached maturity. Martin and Bolstad (2005)

showed that chamber-based soil respiration in DBF of northern Wisconsin ranged from 857–

1143 g C m^{-2} yr^{-1} in 1998 and 1013–1357 g C m^{-2} yr^{-1} in 1999, which is higher than tower ER

(825±133 g C m^{-2} yr^{-1}, WIC) from 1999-2006 in the same region. Soil respiration of 690 g C m^{-2}

over the growing season of 2005 in a mature DBF near WIC tower site was reported (Tang et al.,

2009). Our simulated respiration components (e.g., soil respiration) for DBF were lower than

those reported values (Supplement Fig. S1). The model underestimated GPP but estimated ER

well for mature DBF sites, indicating that the model likely overestimated root autotrophic

respiration. Eddy covariance derived ER were usually lower than chamber-based estimate at the

WIC site due to uncertainties induced by measurement methods, decoupling of surface and

canopy fluxes at night, and spatial scaling (Bolstad et al., 2004; Cook et al., 2004). For ENF, the

model overestimated ER for the mature site because of overestimated soil decomposition rate.

Our simulations also show that DBF had slightly higher soil respiration rate than ENF

(Supplement Fig. S1), which is consistent with the finding that chamber-based soil respiration

was slightly higher for DBF than for ENF (no significant difference) in Wisconsin (Euskirchen et

al., 2003). The changes of ER in secondary forests after clearcutting differ among sites because

of different site conditions (e.g., quantity and quality of soil organic C and litter C) and

harvesting types (e.g., Tang et al., 2009).

The trajectory of our simulated GPP: ER ratio is similar to the curve derived by Amiro et

al. (2010) using EC observations and forest age from fire and harvest chronosequences across

North America (GPP: ER = 1.23*[1−exp (−0.224*AGE)]). Our simulated ratios are within the

observed range of 0.9–1.6 for the DBF sites (Fig. 5a), although the model underestimated the

ratios for mature sites. The growing season GPP :ER ratios are typically higher than the annual

ratios because winter soil organic C decomposition is important to annual C balance (Aanderud

et al., 2013). However, the simulated ratios for the ENF sites are much lower than tower-derived

growth season ratios (1.9–4.7, Fig. 5b), and close to the annual range of 1.6–2.2 estimated by

Desai et al. (2008). The standard gap-filling methods of the eddy covariance flux data may lead

1   to the overestimation of net ecosystem exchange due to the lack of winter C flux observations for

2   the ENF sites and two of the DBF sites (YHW and IHW).

3   Our simulated successional dynamics of NEP following clearcuts generally supported the

4   trajectory hypothesized by Odum (1969) and Chapin et al. (2002). The similar trajectories,

5   however, were caused by different reasons. Odum's trajectories show declining GPP and

6   relatively flat community respiration with time. Our simulated decline in NEP resulted from

7   relatively flat GPP and growing ER with stand development (Figs. 4 and 8). This has been

8   observed for northern temperate hardwood chronosequence sites (Desai et al., 2008), northern

9   temperate pine forests (Peichl et al. 2010), and boreal DBF forests (e.g., Goulden et al., 2011). A

10 recent North American Carbon Program (NACP) synthesis study showed similar changes in NEP

11 after either stand-replacing fire or harvest based on eddy covariance chronosequence

12 measurements across North America (Amiro et al., 2010).

13   Chapin et al. (2002) hypothesized that heterotrophic respiration is initially high, declines

14 in middle succession, and rises thereafter, while NPP reaches a peak in middle age and declines

15 in old stands. The simulated successional trajectories in heterotrophic respiration support the

16 pattern hypothesized by Chapin et al. (2002), whereas our simulated NPP did not decline in

17 mature stands (Supplement Fig. S1). Previous studies also support our simulated trajectory in

18 heterotrophic respiration. For example, Pregitzer and Euskirchen (2004) reported that

19 heterotrophic respiration was high (mean value of 970 g C m^{-2} yr^{-1}) in young temperate forests,

20 declined with age in middle succession, and increased with time for mature forests, although for

21 old temperate forests (>120 years) the decline in NPP reduced heterotrophic respiration. The

22 decline of NPP with age was not predicted in this study. Validation of the simulated NPP was not

23 possible in this study due to the lack of NPP measurements across all sites. Our simulated

24 heterotrophic respiration for mature DBF is close to the observation of 502±86 g C m^{-2} yr^{-1} in a

25 mature DBF near UMBS tower site between 1999 and 2003 (Gough et al., 2008). However, ENF

26 chronosequence sites in this study show that NEP continually increased with age because of

27 relatively flat and low ER (340±96 g C m^{-2} yr^{-1}) and increasing GPP. Successional changes in C

28 fluxes after fire and harvest are similar over chronosequence sites of North America (Amiro et

29 al., 2010), although a specific chronosequence study in Saskatchewan, Canada, observed that C

30 fluxes are greater at the burned site than the harvested sites (Mkhabela et al., 2009).

15

1   Although our model underestimated NEP and GPP for both the DBF and ENF sites in the

2   Upper Midwest region (Figs.1 and 2), our predicted NEP was comparable to estimates from

3   other studies in similar regions. For example, our predicted maximum NEP for the ENF sites

4   (567–602 g C m^{-2} yr^{-1}, 29 years of age) was slightly lower than the estimates (690 g C m^{-2} yr^{-1},

5   15–20 years of age) for afforested white pine (*Pinus strobus*) forests in Ontario (Coursolle et al.,

6   2012). For a northern temperate forest chronosequence study in northern Michigan, NEP higher

7   than 200 g C m^{-2} yr^{-1} in young DBF forests could be derived from the reference forest (153± 115

8   g C m^{-2} yr^{-1}, 85 years of age) (Gough et al., 2007), suggesting that our predictions (390–430 g C

9   m^{-2} yr^{-1}, 14–26 years of age) for the DBF sites could be in the reasonable range. Furthermore, our

10   predicted mean annual NEP (123–177 g C m^{-2} yr^{-1}) for mature DBF sites (>60 years) was close

11   to estimates for other northern DBF, including a northern hardwood forest of central

12   Massachusetts (200±40 g C m^{-2} yr^{-1}, Barford et al., 2001) and four eastern North American DBF

13   (167–236 g C m-2 yr-1 Curtis et al., 2002).

14   We found that the simulated AGC during forest regrowth gradually increased following

15   the typical logistic growth curve (Sprugel, 1985). In the model, low NPP in the early stages

16   results in slow AGC accumulation. Once the amount of NPP approximately equals annual dead

17   biomass C that is largely controlled by the wood turnover rate, the trajectory of AGC reaches a

18   plateau. Previous chronosequence studies showed that AGC increased with increasing age (e.g.,

19   Peichl and Arain, 2006; Goulden et al., 2011; Powers et al., 2012). Powers et al. (2012) reported

20   that AGC increased rapidly with age in young red pine stands across a chronosequence in

21   northern Minnesota, USA. The representativeness and generalization of these findings were

22   limited by the small number of young stands (Powers et al., 2012).

23   Sensitivity analysis shows that more intensive harvests could have larger and longer

24   impacts on successional trajectories of C dynamics in early succession for both DBF and ENF.

25   Fewer flux tower based studies have investigated the effects of harvest intensity on forest C

26   fluxes (e.g., GPP, ER, and NEP) because of the high establishment cost of eddy covariance

27   system. Nevertheless, some modeling studies have provided insights into how forest C fluxes

28   and stocks are affected by harvest intensity. Our findings are supported by previous modeling

29   studies. For example, a recent modeling study of temperate forests reported that more intensive

30   harvests increased the recovery time of NPP for ENF and DBF in Minnesota and Wisconsin,

31   USA (Peters et al. 2013). In the boreal forest of central Canada less intensive harvest and longer

rotation length might increase total C sink (sum of biomass C and soil organic C) up to 40% (Peng et al. 2002), although recent studies indicate that longer rotation length could not necessarily increase C sequestration under changing climate conditions (Wang et al., 2012b; Wang et al., 2013). If harvesting operations largely reduce soil organic matter, C fluxes (e.g., GPP, NPP, ER, and NEP) and living AGC are reduced for both PFTs. Consistent with this, Peters et al. (2013) showed that simulated NPP could not recover to pre-harvesting levels due to greater removal of soil organic matter. Therefore, out model results suggest management practices should aim to decrease soil disturbance caused by harvest operations.

4.2 Differences between DBF and ENF

Our results showed that DBF may reach a peak in LAI and GPP faster than ENF after clearcutting, showing clear differences in pattern of ecosystem development between the DBF and ENF sites. More rapid recovery of LAI and GPP for DBF sites lead to sooner recovery of NEP and AGC regardless of harvest intensity, supporting our second hypothesis. The foliage related parameters such as FolRelGroMax and AmaxB mainly govern the differences in successional trajectories between the two PFTs (Supplement Table S1). DBF is assumed to have more productive foliage than ENF, and more photosynthetic production then can lead to more foliage production. With this positive feedback in the model, GPP, NEP, and AGC of the DBF sites recover more rapidly than those of the ENF sites. Our findings are consistent with the chronosequence studies showing that the temperate DBF in northern Michigan rapidly became a net C sink after six years following disturbances (Gough et al., 2007) and that ENF stands in northern Wisconsin became net C sinks within 10-15 years after harvesting (Noormets et al., 2007). Through the analysis of the Forest Inventory and Analysis (FIA) data, Williams et al. (2012) suggested that faster growth in AGC at high productivity sites caused higher C fluxes and stocks. Our findings are also consistent with a recent modeling study suggesting that temperate DBF switches to positive NEP faster than temperate ENF after clearcuts, and DBF has a higher peak in NEP compared to ENF (Peckham et al., 2012). A modeling study conducted in boreal forests also reported that low productive boreal ENF needed 1-3 more years to attain a positive NEP than boreal DBF after clearcuts in Saskatchewan, Canada (Wang et al., 2012b). These observed and modeled successional changes further indicate that DBF tend to have higher photosynthetic capacity than ENF in the early stage of stand development following stand-replacing harvests.

The sensitivity analysis suggests that more productive forests could be more strongly affected by greater soil removal fraction, as soil removal reduces soil organic matter thereby resulting in relatively low nitrogen mineralization in the model. Peters et al (2013) showed that NPP was more strongly reduced for Aspen than for jack pine in their simulations. However, productive lodgepole pine (*Pinus contorta* Dougl. ex Loud.) could maintain high productivity at 12 years after harvest disturbance regardless of soil organic carbon removal and soil compaction treatments because beneficial ectomycorrhizal fungi associated with lodgepole pine could help access nitrogen from organic matter; while hybrid white spruce (*Picea glauca × engelmannii* [Moench] Voss) was more sensitive to the treatments (Kranabetter et al., 2006). The discrepancy might be caused by the lack of representation of relationship between fungi and plants in the model.

4.3   Limitations and challenges

PnET-CN can explicitly simulate the effects of disturbance, pollution, and climate change on forest C dynamics (e.g., Ollinger et al., 2002; Pan et al., 2009; Peters et al., 2013). Despite the capability of the model, we do recognize that the model has some limitations in simulating harvesting effects, and that the accurate representation of the trajectories of C fluxes and stocks following harvests still remains a challenge.

First, the performance of the model to simulate forest regrowth after harvests is limited by the absence of regeneration and understory in the model. Most process-based models such as PnET-CN and TEM (Raich et al., 1991) have been mainly developed to simulate C balance for mature forests over the past decades (Landsberg, 2003), resulting in no provision for simulating regeneration such as shrub component and species succession in these models. Changes in forest composition (e.g., evergreen and deciduous tree species and understory shrubs) along the course of succession are not fully considered by most ecosystem models. PnET-CN does not simulate shrubs and herbs that likely dominate stands in the early successional stage after stand-replacing harvests. Therefore, the model is not able to simulate the particularly high GPP and ER in the young forests where forest canopy has not yet fully recovered.

Understory layer is also an important component for mature forest ecosystems in terms of C fluxes (Misson et al., 2007) and stocks. Misson et al. (2007) reported that understory can contribute 11% (range, 0–39%) of GPP at 10 sites across a wide range of forest type and climate.

1   PnET-CN slightly overestimated overstory LAI for the mature DBF sites and reasonably

2   predicted foliar nitrogen concentration compared to satellite-based estimates (data not shown).

3   The lack of understory layer in the model is possibly responsible for the underestimation of GPP

4   for mature DBF sites. Species competition and cohort methods that have been employed in other

5   models such as ED (Medvigy et al., 2009) and LPG-Guess (Smith et al., 2001) could be used to

6   improve the regeneration and understory components of PnET-CN in the future.

7   Second, parameter values used in the model were generally derived from specific

8   measurements for a given stand development stage particularly mature forests, although the

9   parameter values likely differ with stand development. For example, the canopy light attenuation

10 constant coefficient is typically measured in mature forests (e.g., Ryu et al., 2008), although the

11 coefficient is known to change with canopy cover (Brantley and Young, 2007). The use of the

12 canopy light attenuation coefficient measured in mature forest for whole forest life simulations

13 could slow down stand development due to the underestimation of photosynthesis in young

14 forests. Understanding the relationship between such parameters and state variables (e.g., LAI) is

15 thus one of the challenges to simulate the effects of stand-replacing harvests on forest C

16 dynamics.

17   Third, changing climate conditions can also affect the values of some parameters. For

18 example, wood turnover rate (%, tree mortality in terms of biomass losses), to which wood living

19 biomass C and soil organic C are sensitive, could be altered by extreme weather conditions

20 including droughts (Allen et al., 2010; Wang et al., 2012a). Most process-based models are not

21 able to simulate the mechanistic processes associated with tree mortality under changing climate

22 conditions (McDowell, 2011; Wang et al., 2012a), although there is growing interest in the

23 mechanistic modeling of forest mortality (e.g., McDowell et al., 2013; Powell et al., 2013).

24 Recent studies have revealed that climate and disturbance legacies govern forest C dynamics

25 (Magnani et al., 2007; Bond-Lamberty et al., 2013). Future modeling efforts can benefit from

26 improved understanding of the effects of climate change on parameter values that are assumed to

27 be constant in the model.

28   Finally, the whole silvicultural system (e.g., harvests) is not fully considered in the

29 model. Harvest methods depend on forest types, management needs, and species to be

30 regenerated. For example, selective harvesting or shelterwood system is typically used for

hardwoods in Wisconsin (Wisconsin Department of Natural Resources, 2011). Stand-replacing harvesting was assumed for both DBF and ENF chronosequence sites due to the lack of harvesting information and the types of clearing applied to the sites studied. The sensitivity analysis conducted in this study suggests that harvest intensity affects C dynamics in early succession after harvesting. Observations in residuals and post stands after each operation type (e.g., pre-commercial thinning and selective harvesting) are needed to parameterize process-based models for better mechanistic understanding of the harvest effects on forest C dynamics.

Besides disturbances, climate is also a key driver of ecosystem structure and function. Climate extremes such as drought induce forest die-off and reduction of carbon uptake globally (Ciais et al., 2005; Kurz et al., 2008; van Mantgem et al., 2009; Zhao and Running, 2010; Peng et al., 2011; Ma et al., 2012; Schwalm et al., 2012). In North America, droughts and disturbance are two main sources of interannual variability in carbon fluxes (Xiao et al. 2011, 2014). Future modeling studies should explicitly the effects of both disturbance and climate extremes.

5   Conclusions

The PnET-CN model is generally able to simulate the effects of stand-replacing harvests on forest C dynamics (C fluxes and AGC) for two northern temperate forest chronosequences. The predicted dynamics in NEP and AGC following clearcuts generally follow the hypothesized trajectories (Odum, 1969; Chapin et al., 2002), although our simulations show that the decline in NEP was due to relatively flattening GPP and gradually increasing ER. Our study also shows that DBF recovered faster (11 years) from net C sources to net sinks and accumulated more C in AGC than ENF. Northern temperate ENF is more vulnerable to stand-replacing harvests than northern temperate DBF. Future research is needed to better understand how respiration components (e.g., ecosystem and soil respiration) and production components (e.g., overstory and understory) change with forest age and their determinants. Modeling the combined effects of climate change and forest management will benefit from the incorporation of forest population dynamics (e.g., regeneration and mortality), relationships between age-related model parameters and state variables (e.g., LAI), and silvicultural system into the model. With these improvements, process-based ecosystem models can better simulate regional C balance associated with disturbance regime under changing climate.

Acknowledgements

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Tables

Table 1 Site characteristics for two chronosequences of deciduous broadleaf forests (DBF) and evergreen needleleaf forests (ENF) in Upper Midwest region of Wisconsin and Michigan, United States.

Site	ID	Location	Plant function type	Dominant species	Year of recent disturbance	AGC (Mg ha^{-1}, 2005)	LAI (m^2 m^{-2}, 2002)	Data period	Reference
Clearcut young hardwood	YHW	46.72°N 91.25°W	DBF	Aspen, red maple	1999	3.3 (1.3)	0.79 (0.6)	2002	Noormets et al. 2007
Intermediate hardwood	IHW	46.73°N 91.23°W	DBF	Aspen	1984	47.6 (15.6)	3.0	2003	Noormets et al. 2008
Willow creek	WIC	45.80°N 90.08°W	DBF	Sugar maple, basswood, green ash	1875[a], 1933	74.9[b]	5.36 (0.47)[c]	2000-2006	Cook et al., 2008, Curtis et al. 2002
University of Michigan Biological Station	UMBS	45.56°N 84.71°W	DBF	Aspen, white pine, red oak, sugar maple	1920	73.2 (3.1)[d]	3.54 (0.31)[e]	2000-2003	Gough et al. 2008
Young red pine	YRP	46.72°N 91.18°W	ENF	Red pine, jack pine	1993	7.7 (8.3)	0.52 (0.3)	2002	Noormets et al. 2007
Young jack pine	YJP	46.62°N 91.08°W	ENF	Jack pine	1987	4.9 (5.0)	0.93	2004-2005	Noormets et al. 2008
Intermediate red pine	IRP	46.69°N 91.15°W	ENF	Red pine	1980	47.7 (37.3)	3.0	2003	Desai et al. 2008
Mature red pine	MRP	46.74°N 91.17°W	ENF	Red pine, aspen	1939	56.9 (33.1)	2.7 (0.8)	2002-2005	Noormets et al. 2007

[a] estimated year of disturbance based on Ameriflux site description in AmericFlux.

[b] sum of wood and foliage biomass carbon from Curtis et al., 2002.

[c] estimated values based on measurements in 1998 to 2000 and 2002 from Cook et al., 2008.,

[d] value in 2003 from Gough et al. 2008.

[e] calculated based on multi-year (1999-2003) estimations with litter traps from Gough et al. 2008.

1   Table 2 PnET model performance in monthly carbon fluxes (GPP: gross primary productivity;

2   ER: ecosystem respiration; NEP: net ecosystem productivity), leaf area index (LAI), and

3   aboveground carbon stock (AGC) for the two chronosequences.

	NRMSE%[a]	d[b]	n
DBF			
GPP	10	0.95	147
ER	10	0.92	147
NEP	17	0.81	147
LAI	33	0.97	4
AGC	42	0.95	4
ENF			
GPP	28	0.91	64
ER	37	0.88	64
NEP	46	0.58	64
LAI	29	0.96	4
AGC	37	0.94	4
Overall			
GPP	11	0.94	211
ER	10	0.91	211
NEP	21	0.73	211
LAI	31	0.96	8
AGC	28	0.95	8
Total performance [c]	20	0.90	

4   [a] Normalized root mean square error.

5   [b] Willmott index.

6   [c] Mean value of evaluating statistics for all tested variables.
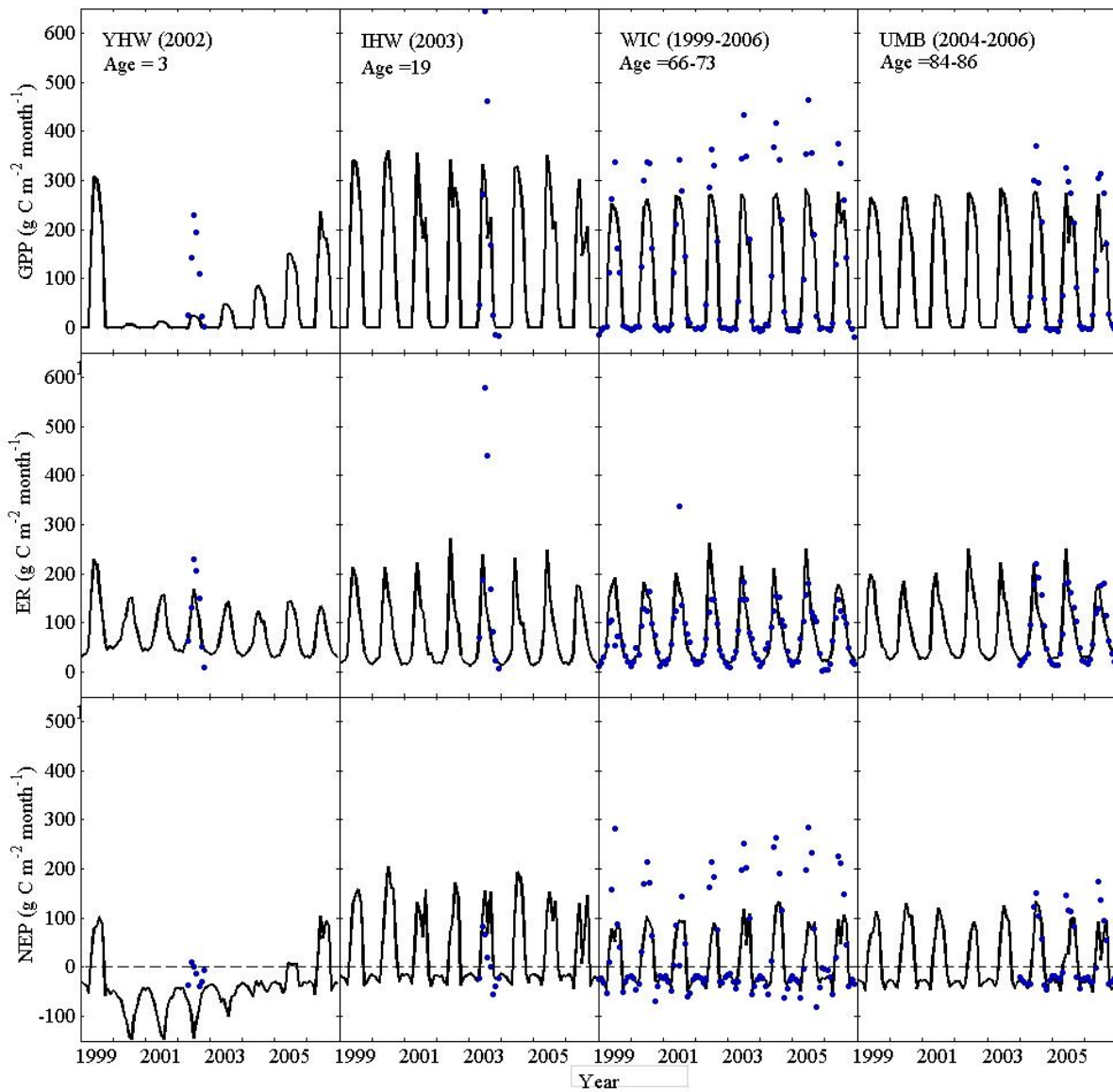
Figure 1. Simulated (lines) and observed (symbols) monthly carbon fluxes: GPP, ER, and NEP for the deciduous broadleaf chronosequence sites from 1999-2007.
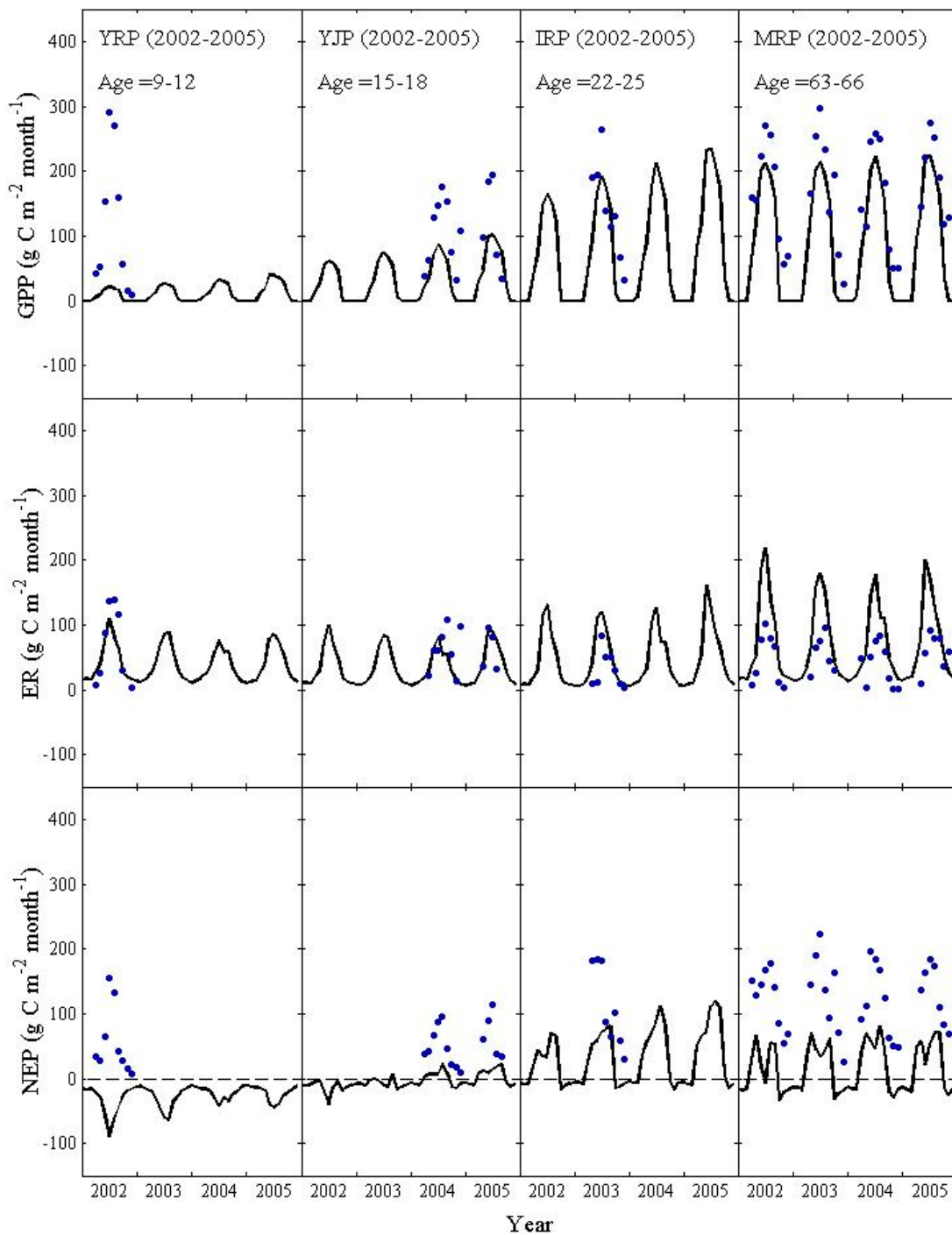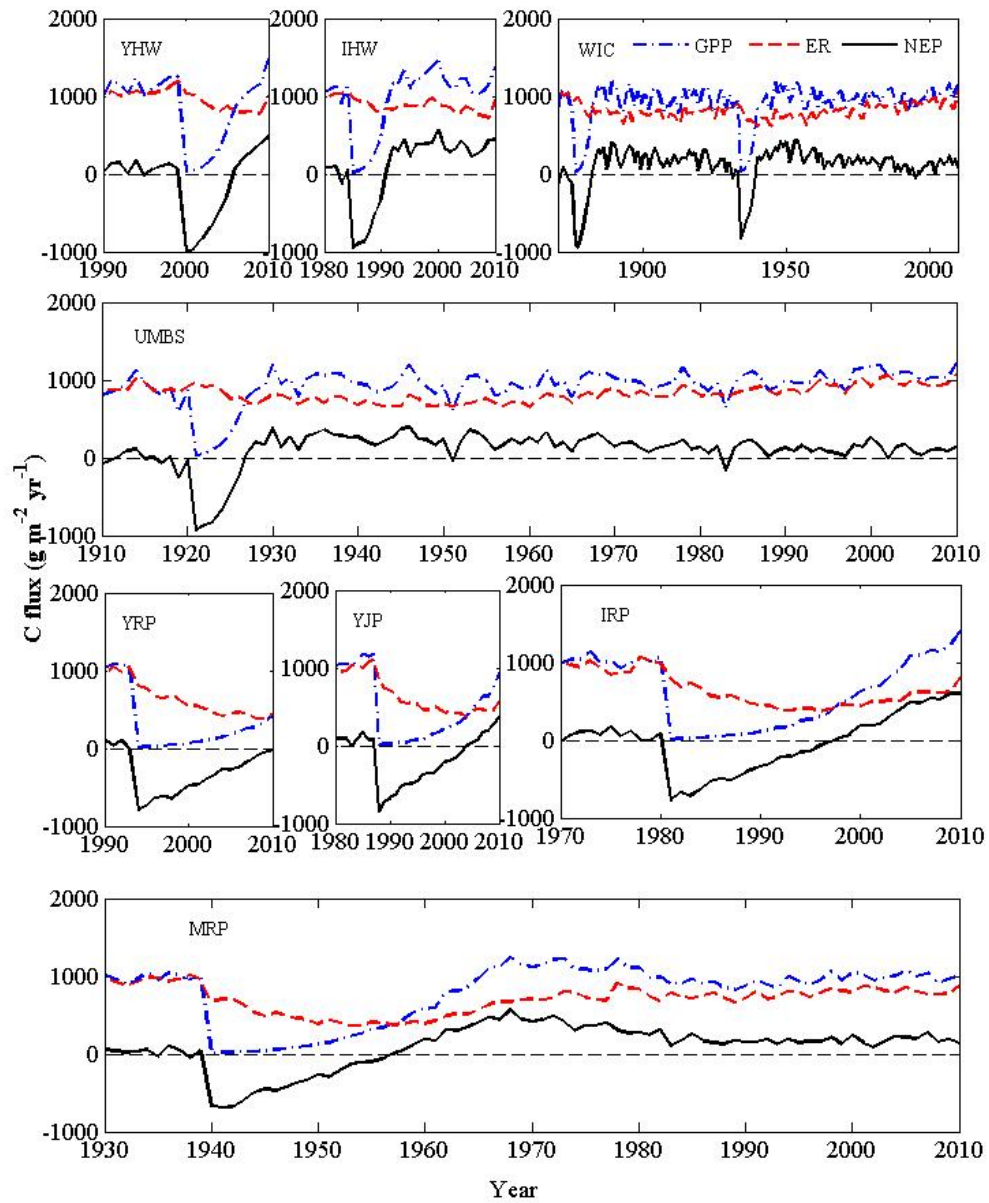
Figure 2. Simulated (lines) and observed (symbols) monthly carbon fluxes: GPP, ER, and NEP for the evergreen coniferous chronosequence study sites from 2002-2005.
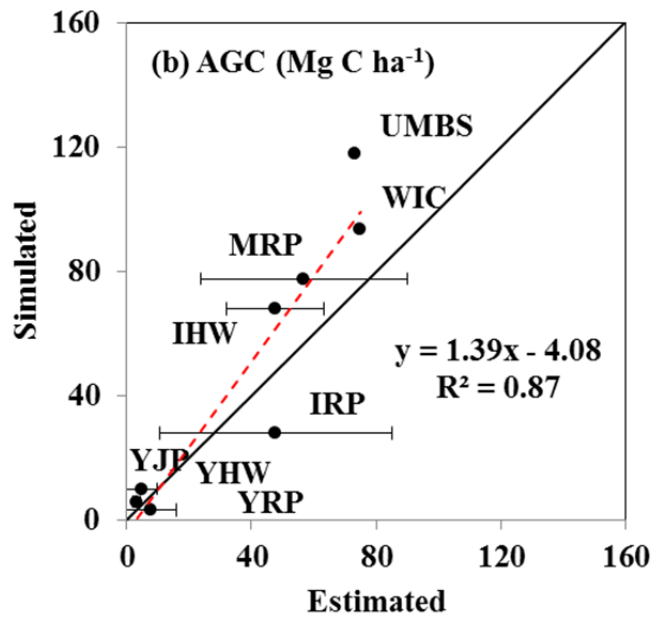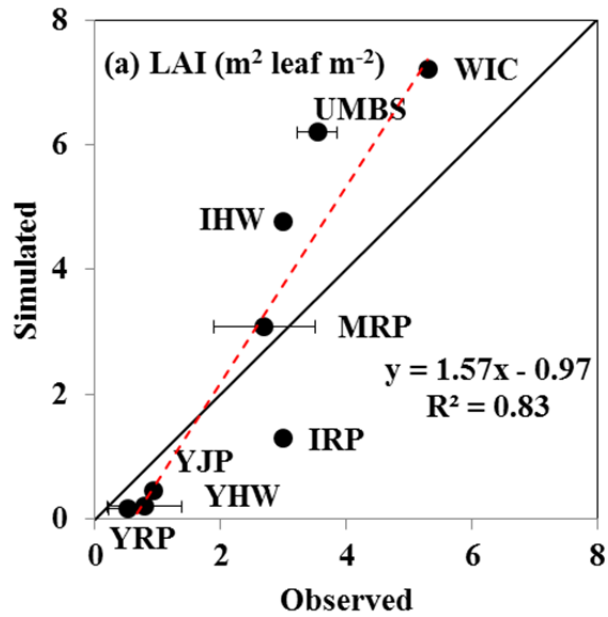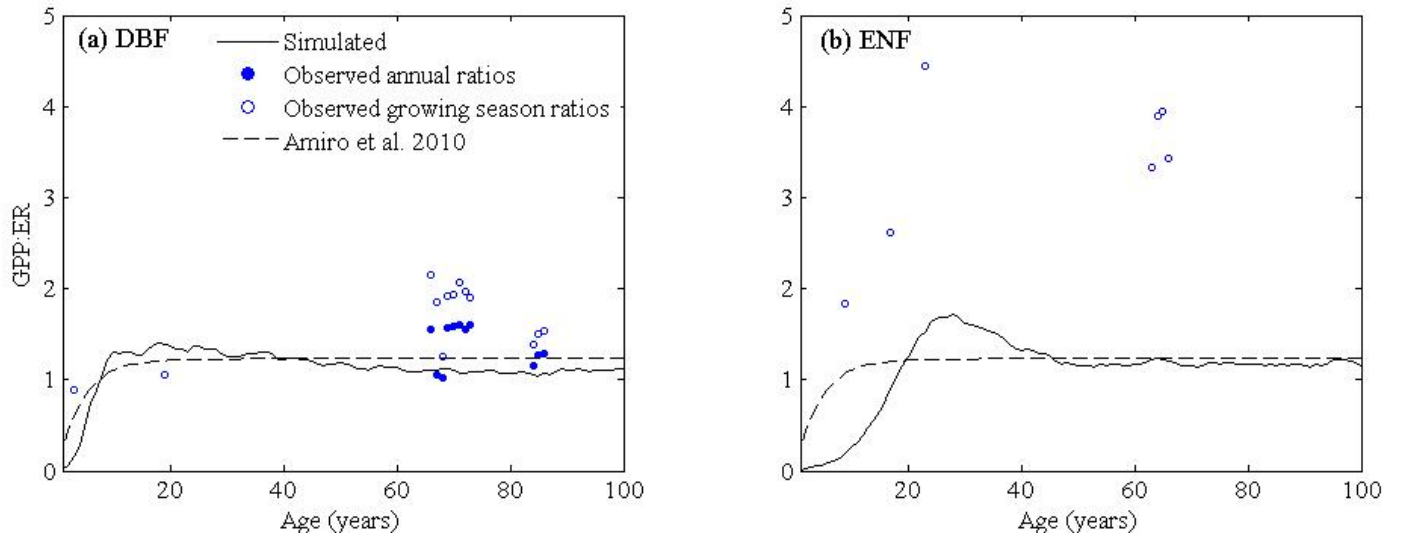
1

2   Figure 3. Comparisons of simulated and observed (a) leaf area index (LAI) and (b) aboveground

3   carbon stock (AGC) for all eight sites.

4

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2   Figure 4. Simulated trajectories of GPP, ER, and NEP for each site based on the site disturbance

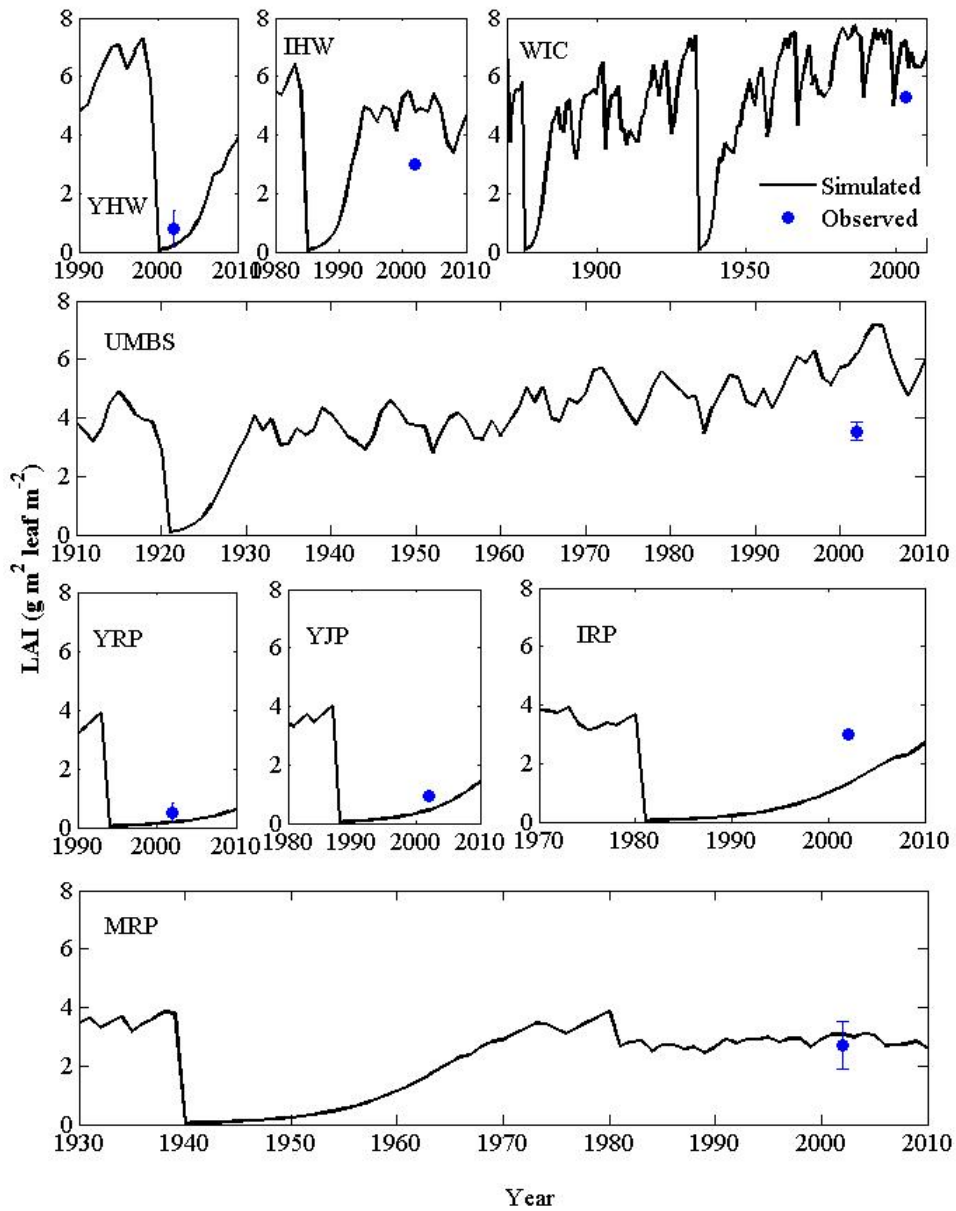3   history (Table 1).The time series started from the earliest major disturbance for each site.

4

1

2

3 Figure 5. Simulated trajectories of the annual GPP/ER ratio with stand age for (a) deciduous

4 broadleaf forests (DBF) and (b) evergreen coniferous forests (ENF). The dashed line is a fitted

5 curve derived by Amiro et al. (2010) using eddy covariance observations from chronosequence

6 forests in North America. Solid and hollow circles represent measured annual and growing

7 season (May to October) ratios, respectively. The simulated curves were smoothed using a

8 moving average filter with a span of 5.

9

Figure 6. Simulated trajectories of LAI for each site based on the site disturbance history (Table 1). The time series started from the earliest major disturbance for each site. Symbols represent measured LAI.
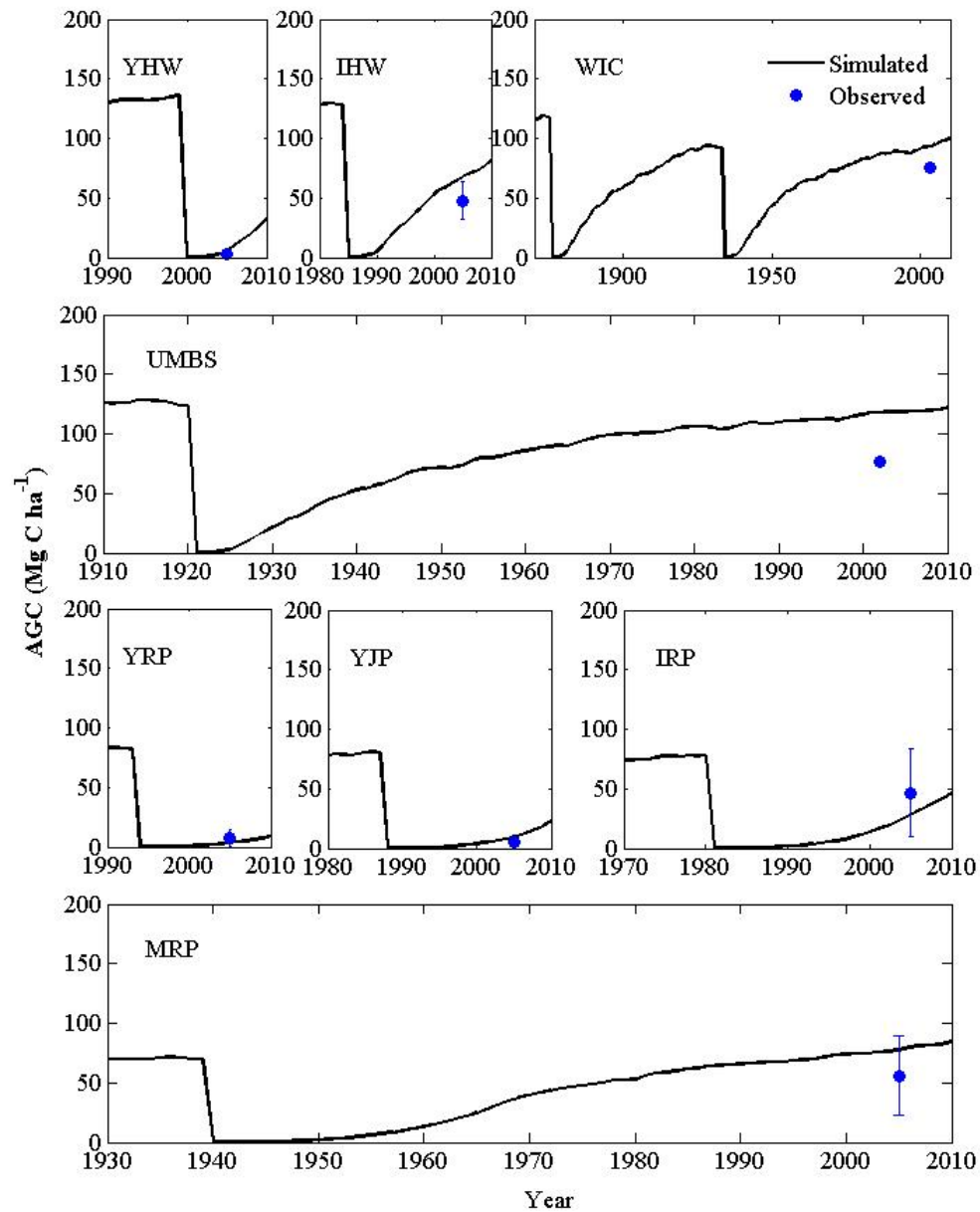
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2   Figure 7. Simulated trajectories of aboveground biomass carbon (AGC) for each site based on

3   the site disturbance history (Table 1). The time series started from the earliest major disturbance

4   for each site. Symbols represent estimated AGC.
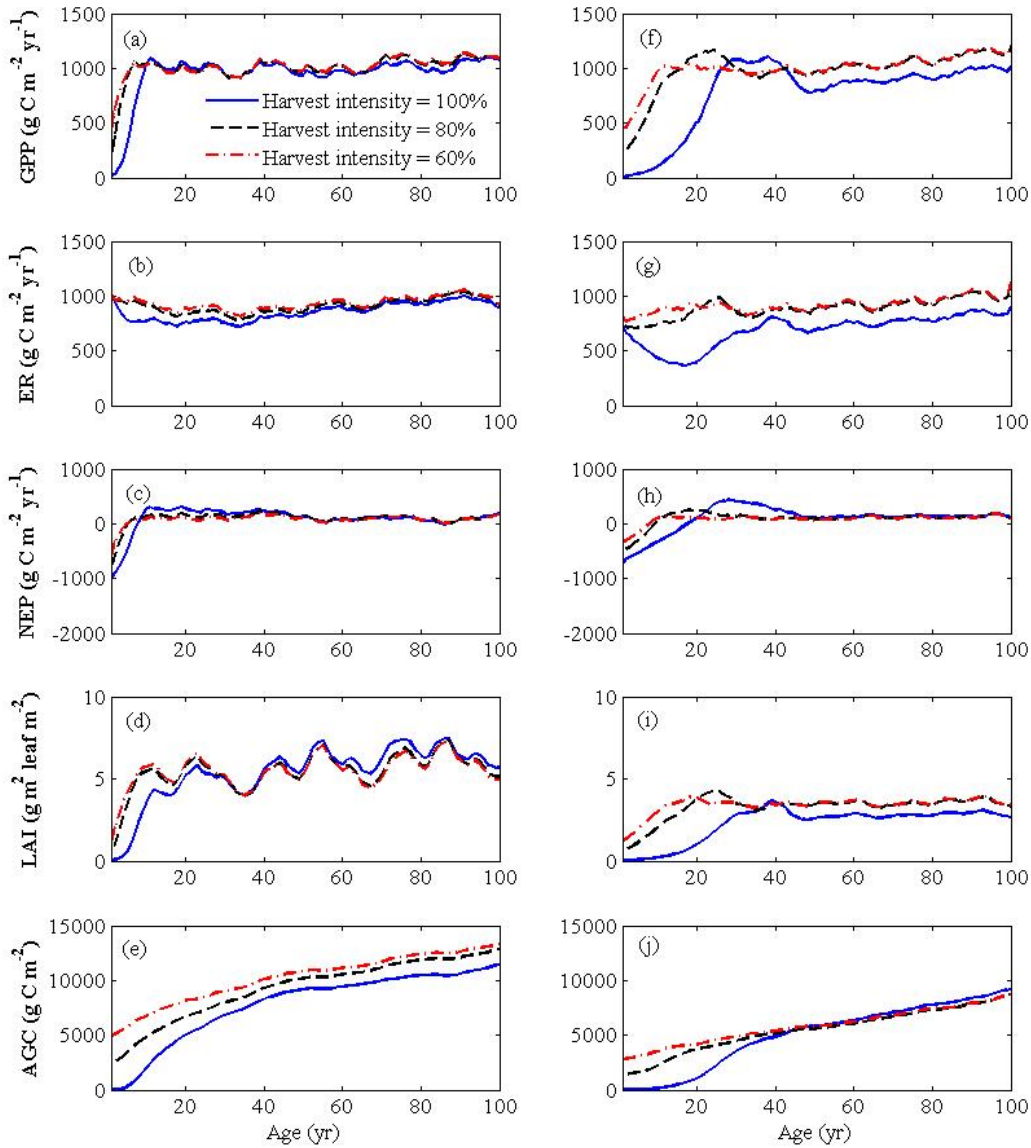
Figure 8. Sensitivity of carbon fluxes (GPP, gross primary production; ER, ecosystem respiration; NEP, net ecosystem production) and stand characteristics (LAI: leaf area index; AGC: aboveground carbon stock) to changes in harvest intensity (reduced by 0.2 and 0.4 compared to 1 for assumed clearcuts used in the model tests) for (a-e) deciduous broadleaf forests (DBF) at Willow creek and (f-j) evergreen coniferous forests (ENF) at Mature red pine site over a 100-yr harvest cycle. The simulated curves were smoothed using a moving average filter with a span of 5.