

1 **Substrate quality alters microbial mineralization of added substrate and soil organic**
2 **carbon**

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25 **Abstract.** The rate and extent of decomposition of soil organic carbon (SOC) is dependent on 
26 substrate chemistry and microbial dynamics. Our objectives were to understand the influence of
27 substrate chemistry on microbial processing of carbon (C), and to use model fitting to quantify
28 differences in pool sizes and mineralization rates. We conducted an incubation experiment for
29 270 days using four uniformly-labeled ^{14}C substrates (glucose, starch, cinnamic acid and stearic
30 acid) on four different soils (a temperate Mollisol, a tropical Ultisol, a sub-arctic Andisol, and an
31 arctic Gelisol). The ^{14}C labeling enabled us to separate CO_2 respired from added substrates and
32 from native SOC. Microbial gene copy numbers were quantified at days 4, 30 and 270 using
33 quantitative polymerase chain reaction (qPCR). Substrate C respiration was always higher for
34 glucose than other substrates. Soils with cinnamic and stearic acid lost more native SOC than
35 glucose- and starch-amended soils, despite an initial delay in respiration.  Cinnamic and stearic
36 acid amendments also exhibited higher fungal gene copy numbers at the end of incubation. We
37 found that 270 days was sufficient to model decomposition of simple substrates (glucose and
38 starch) with three pools, but was insufficient for more complex substrates (cinnamic and stearic
39 acid) and native SOC. This study reveals that substrate quality imparts considerable control on
40 microbial decomposition of newly added and native SOC, and demonstrates the need for multi-
41 year incubation experiments to constrain decomposition parameters for the most recalcitrant
42 fractions of SOC and added substrates.

43

44 **1 Introduction**

45

46 The chemistry of carbon (C) inputs into soils influences the rate and extent of microbial
47 decomposition of soil organic carbon (SOC) (Schmidt et al., 2011; Schnitzer and Monreal,

48 2011). Three hypotheses are used to explain the decomposition of fresh C according to chemistry
49 (Wickings et al., 2012): (i) chemical convergence, (ii) initial litter quality, and (iii) decomposer
50 control. The chemical convergence hypothesis suggests that regardless of the differences in
51 substrate quality and microbial diversity, all C substrates undergo decomposition through a
52 limited number of biochemical pathways and reactions resulting in SOC of homogeneous
53 chemistry (McGill, 2007; Fierer et al., 2009) and it supports the general understanding that
54 simple sugars and amino acids are preferentially decomposed over complex lignin and ligno-
55 cellulose. However, recent studies have also identified simple biopolymers of plant and
56 microbial origin in the stabilized SOC (Sutton and Sposito, 2005; Kelleher and Simpson, 2006).
57 According to the initial litter quality hypothesis, the chemical composition of substrates at the
58 start of the decomposition process (e.g. leaf litter) exhibits a strong influence on decomposition
59 rate and the chemistry of stabilized SOC (Angers and Mehuys, 1990; Berg and McClaugherty,
60 2008). The decomposer control hypothesis suggests that distinct decomposer communities
61 impose constraints on substrate decomposition regardless of the difference in quality of substrate
62 and stage of decomposition. Wickings et al. (2012) analyzed these three hypotheses through a
63 long-term litter decomposition experiment and found experimental evidence for an interactive
64 influence of both ‘initial litter quality hypothesis’ and ‘decomposer control hypothesis’ on the
65 chemistry of decomposing litter. While these two hypotheses appear to be complimentary in
66 nature, there have been few studies that expressly examine the combined influence of initial substrate
67 quality and the decomposer community on the decomposition of C inputs leading to SOC
68 formation and stabilization.

69 Most past studies addressed the initial C substrate quality effect by adding isotopically-
70 labeled and/or chemically distinct plant litters to soils in laboratory microcosms. Labeling with

71 ^{13}C or ^{14}C isotopes allows separate quantification of SOC-derived CO_2 and substrate-derived
72 CO_2 , and specifically resolves the effects of substrate additions on SOC turnover (Kuzyakov and
73 Cheng, 2001; Leake et al., 2006; Williams et al., 2006; Werth and Kuzyakov, 2008).
74 Isotopically-labeled natural plant litter, however, can't be used to identify the role of specific
75 litter constituents on SOC dynamics (Grayston et al. 1998; Loreau, 2001). One way to overcome
76 this issue is to apply isotopically-labeled C compounds representing different constituents of
77 plant residues, e.g. simple sugars, polysaccharides, proteins, lipids, and/or aromatic compounds
78 to observe their direct effect on SOC decomposition (e.g. Brant et al., 2006; Hoyle et al., 2008;
79 Schneckengerger et al., 2008; Strahm and Harrison, 2008; de Graaff et al., 2010). These studies
80 indicated increased, decreased or no change in SOC decomposition dynamics due to the addition
81 of substrates compared to unamended control treatments. Also, most of these studies used only
82 labile C compounds such as simple sugars and organic acids as C amendments, and did not
83 account for other, relatively more recalcitrant C compounds such as lignin, fatty acids, lipids etc.
84 Therefore, more studies with isotopically-labeled substrate additions are needed to determine the
85 role of initial litter quality on SOC decomposition.

86 In accordance with the decomposer community hypothesis, the magnitude of SOC change
87 largely depends on the abundance and diversity of soil microbial communities (Fontaine et al.,
88 2005). Bacteria and fungi are the major drivers of substrate and SOC decomposition comprising
89 more than 90% of the soil microbial biomass, and clear evidence exists that these groups
90 function differently in the decomposition process (de Graaff et al., 2010). There is a general
91 understanding that easily available simple C compounds are taken up by the fast growing r-
92 strategists in the early stages of decomposition, while in the later stages slow-growing k-
93 strategists break down more recalcitrant C, i.e., compounds having higher thermodynamic

94 activation energies (Wardle et al., 2002; Fontaine et al., 2003; Blagodatskaya and Kuzyakov,
95 2008). Among the r-strategists, bacteria are mostly considered responsible for utilizing labile C
96 sources immediately after their addition to soils (Paterson et al., 2007; Moore-Kucera and Dick,
97 2008). Fungi are commonly regarded as k-strategists utilizing C from more recalcitrant
98 substrates (Otten et al., 2001). However, this general paradigm has been challenged by other
99 studies. For example, Nottingham et al. (2009) reported that gram-negative bacteria also belong
100 to k-strategists and are responsible for the decomposition of complex C compounds, and Rinnan
101 and Bååth (2009) did not find evidence that bacteria were more efficient in utilizing simple
102 compounds than fungi. Evaluation of the interplay of these life-history strategies on SOC
103 turnover across a suite of substrates, soils and microbial communities is still lacking and is
104 essential to resolve the role of the decomposer community on SOC dynamics.

105 Lab-scale incubation studies have been instrumental to quantify the influence of initial litter
106 quality and decomposer community by modeling SOC pool sizes and mineralization rates.
107 Although laboratory incubations deviate from natural ecosystem environments in terms of
108 continuous C input, microbial community structure and environmental conditions, they help to
109 isolate specific mechanisms by systematically eliminating variations in certain environmental
110 variables. Since there is no continuous C input during the course of the experiment, incubation
111 studies can be used to quantify the mineralization kinetics of different fractions of C pools
112 according to different types of substrate addition (Schädel et al., 2013). Statistical models are
113 used to estimate the sizes and rates of SOC pools by curve fitting. Within these constraints, total
114 SOC is generally divided into three pools with fast, intermediate and slow mineralization rates
115 (Trumbore, 1997; Krull et al., 2003). The terminology, definitions and measurement techniques
116 of these pools, however, vary widely in the literature. The lack of experimental data using

117 multiple substrates in long-term incubation experiments, however, limits understanding of the
118 role of substrate complexity and decomposer community (von Lützow and Kögel-Knabner,
119 2009; Schädel et al., 2013).

120 In this paper we used long-term incubations to investigate how the chemistry of added C
121 substrates affected mineralization of the substrate C and of the SOC, and the composition of the
122 decomposer community in several different soils. We hypothesized that: (i) cumulative
123 respiration of substrate C and native C would be higher when soils are amended with easily
124 metabolized substrates compared to relatively more complex substrates, and that (ii) both
125 incubation time and the relative recalcitrance of the added substrate would favor soil fungi over
126 bacteria. To test these hypotheses, we conducted a long-term (270 day) laboratory incubation
127 experiment using four different uniformly-labeled ^{14}C substrates (monosaccharide,
128 polysaccharide, aromatic, fatty acid). The ^{14}C labeling enabled us to separate substrate-derived
129 CO_2 from native SOC-derived CO_2 . We tested the effect of different substrate additions on
130 substrate and native C respiration using a first order exponential decay model, and utilized
131 quantitative polymerase chain reaction (qPCR) to compare bacterial and fungal gene copy
132 numbers. Finally, we incubated four different soils that spanned a wide range in climate, soil
133 development, and type and quantity of organic C inputs.

134

135 **2 Materials and methods**

136

137 **2.1 Soil sampling and characterization**

138

139 Soils were collected from four contrasting climatic zones- temperate, tropical, sub-arctic and
140 arctic. The selected soils are from major soil orders of the respective climatic regions: the

141 Mollisol (temperate), the Ultisol (tropical), the Andisol (sub-arctic), and the Gelisol (arctic)
142 (Table 1). Multiple soil cores were collected randomly from each location to a depth of 15 cm,
143 pooled to form a composite sample per location and sieved to <2 mm. Subsamples (n=3) of the
144 soils were taken for the determination of organic C, total N, microbial biomass C (MBC), soil
145 pH, and soil texture (Table 1). Organic C and total N concentrations were determined by
146 combustion method using a Leco combustion analyzer (Leco Corp., St. Joseph, MI) (Nelson and
147 Sommers, 1996) after removing the inorganic C by treating with 3M HCl for 1 hr. Determination
148 of MBC was conducted by the chloroform fumigation extraction method (Vance et al., 1987).
149 Soil pH was determined by shaking 1 part soil in 2 parts Milli-Q (MQ) water and measuring the
150 pH of the supernatant (Thomas, 1996), and soil texture was determined by the bouyoucos
151 hydrometer method (Gee and Or, 2002).

152

153 **2.2 Carbon substrates**

154

155 Four uniformly-labeled ^{14}C substrates were used: glucose, starch, cinnamic acid and stearic acid,
156 representing several dominant C compounds present in plant litter and SOC, and spanning a
157 range of chemical lability. Glucose is a common simple sugar and starch is a common
158 polysaccharide in plant residues, cinnamic acid contains an aromatic ring and is a common
159 product of lignin depolymerization, and stearic acid represents a fatty acid (Orwin et al., 2006;
160 Rinnan and Bååth, 2009). Similar to Orwin et al. (2006), we selected compounds containing only
161 C, hydrogen, and oxygen and lacking nutrient elements such as nitrogen and phosphorus. These
162 nutrients are expected to cause confounding effects on microbial activities and C decomposition
163 (Orwin et al., 2006). Availability in uniformly-labeled ^{14}C form (U- ^{14}C) was also another
164 criterion for the compound selection.

165 2.3 Incubation experiments

166


167 We used five control (unamended) replicates of each of 4 soils for measuring native SOC
168 respiration. Two replicates were destructively harvested at days 4 and 30 and stored at -20°C for
169 microbial community analysis. The three remaining replicates were monitored for respiration
170 until they were destructively harvested for community analysis at 270 days. An identical scheme
171 was used for the 4 different substrates to measure $^{14}\text{CO}_2$ evolved from decomposition of
172 substrate and CO_2 evolved from native SOC. Our initial experiment thus had 4 soils each having
173 5 controls and five ^{14}C substrate additions, using 4 different substrates. Though we could include
174 only one replicate for the destructive sampling at day 4 and day 30 due to limitations of space,
175 soil, and ^{14}C substrate, we conducted three analytical replicates of the microbial community
176 measurements for these sampling times, and three experimental replicates for the 270 day
177 sampling time.

178 For the substrate addition experiments, 25 g (oven-dry basis) soils were amended with
179 0.4 mg C g⁻¹ soil substrates which were labeled with 296 Becquerel g⁻¹ soil U- ^{14}C substrate. The
180 substrates were added in dissolved form and mixed well with the soil using a spatula. The final
181 moisture content of substrate amended and unamended samples were maintained at 50% WHC
182 with MQ water. The solvents were MQ water for glucose and starch, ethanol for cinnamic acid
183 and toluene for stearic acid. Organic solvents were used for cinnamic acid and stearic acid
184 because these compounds are sparingly soluble in water. We introduced only a small amount of
185 organic solvents to the samples (4 μL ethanol g⁻¹ soil and 6 μL toluene g⁻¹ soil) and our
186 preliminary experiments revealed that the solvents did not influence the microbial activities (Fig.

187 S1 

188 2.4 Measurement of CO₂ respiration

189

190 Specimen cups containing the substrate amended and unamended control soils were placed in
191 1L, wide mouthed glass jars, along with a glass vial containing 17 ml of 0.5 N NaOH solution to
192 trap the evolved CO₂. The jars were tightly closed and incubated in the dark at 20 °C for up to
193 270 days in a temperature and humidity controlled room. The NaOH solution was exchanged 15
194 times during the experiment at daily to weekly intervals in the first two months and monthly
195 intervals thereafter.  The jars were sufficiently ventilated each time when they were opened for
196 NaOH solution exchange in order to avoid anaerobic conditions inside the jar.

197 The amount of total C respiration is defined as the sum of SOC-derived CO₂ and substrate-
198 derived ¹⁴CO₂, where the control (unamended) samples have no contribution from substrate.
199 Total mineralized CO₂ was determined by titrating an aliquot of NaOH solution collected at each
200 sampling time with 0.5 N HCl by an automatic titrator (Metrohm USA). Before the titration, the
201 CO₂ collected in NaOH solution was precipitated as barium carbonate (BaCO₃) by adding 2 ml
202 10% barium chloride (BaCl₂). The volume of acid needed to neutralize the remaining NaOH
203 (unreacted with CO₂) was determined by the titration, which was used to calculate the
204 concentration of CO₂ trapped in the NaOH solution (Zibilske, 1994). Evolution of substrate C
205 was determined by measuring the activity of ¹⁴CO₂ trapped in NaOH solution collected from the
206 substrate amended samples with a Packard Tri-Carb Liquid Scintillation Counter (LSC) after
207 mixing 5 ml of the NaOH solution with 10 ml of the scintillation cocktail Ultima Gold XR
208 (PerkinElmer). The CO₂ derived from SOC for the substrate-amended samples was calculated by
209 subtracting substrate-derived ¹⁴CO₂ from the total CO₂.

210

211 **2.5 Microbial gene copy numbers**

212
213 Microbial DNA extraction was conducted with 0.25 g of moist soil using the PowerSoil DNA
214 Isolation Kit (MOBIO Laboratories, Inc., CA, USA). The abundance of the ribosomal RNA
215 (rRNA) genes was determined by quantitative real time polymerase chain reaction (qPCR) on a
216 CFX96™ Real-Time PCR Detection System (Bio-Rad Laboratories, CA, USA) with group
217 specific ribosomal DNA gene primers using iQ SYBR Green Supermix (Bio-Rad, CA, USA). A
218 small segment of the sample DNA was amplified using primer pairs that targeted the conserved
219 region of the rRNA. Gene copy numbers for bacteria, fungi and archaea were determined in
220 analytical triplicates using standard curves constructed from group specific microorganisms. The
221 primers, PCR reaction conditions, composition of the reaction mixture and the pure cultures used
222 for preparing the standard curves are described in Table S1.

223

224 **2.6 Exponential decay modeling**


225

226 The respiration data (both the substrate C and SOC) were tested using a double and a triple pool
227 first order exponential decay model (Farrar et al., 2012):

$$228 \quad \text{Double pool model: } C_t = C_1(e^{-k_1t}) + C_2(e^{-k_2t}) \quad (1)$$

$$229 \quad \text{Triple pool model: } C_t = C_1(e^{-k_1t}) + C_2(e^{-k_2t}) + C_3(e^{-k_3t}) \quad (2)$$


230 where C_t is the total substrate C (in terms of % of added substrate C) or total SOC (in terms of %
231 of initial SOC) remaining in time t , C_1 , C_2 , and C_3 are pool sizes, and k_1 , k_2 and k_3 are associated
232 mineralization rates. For the double pool model, C_1 and C_2 are defined as fast and intermediate
233 pools, respectively, and for triple pool model, C_1 , C_2 and C_3 are defined as fast, intermediate and

234 slow pools, respectively. For each set of data, multiple pool models were fit using Sigma plot
235 v11 (Systat Software Inc., IL, USA) and dependency values and r^2 for fit parameters were
236 calculated. We followed two criteria to determine the best fits as outlined in Farrar et al. (2012):
237 (i) dependencies less than 0.98, and (ii) a statistically greater r^2 over a lower-order fit. 

238

239 **2.7 Statistical analysis**

240

241 Statistical analyses were conducted using SAS software (SAS Institute Inc., 2002). The effect of
242 substrate type on substrate-derived and SOC-derived respiration was determined by repeated
243 measures analysis using the PROC MIXED option of SAS with incubation length considered as
244 the repeated measure with autoregressive 1 covariance structure. The repeated measures analysis
245 with the PROC MIXED option of SAS is analogous to the generalized linear model analysis with
246 the PROC GLM option of SAS, except that the former allows modeling of the covariance
247 structure of the dataset to account for unevenly spaced sampling dates (Littel et al., 1996;
248 Schaeffer et al., 2007). *Post hoc* comparisons for determining the effect of substrate types on
249 respiration, and modeled mineralization parameters (pool sizes and rates) in each soil were
250 performed using PROC GLM of SAS. The treatment effects were separated using the Fisher's
251 protected least significant difference (LSD) test. T-test was performed to determine if fungal to
252 bacterial (F:B) gene copy ratio upon substrate addition was significantly different from F:B ratio
253 of unamended controls at each time point. In all statistical tests, the mean differences were
254 considered significant at $P \leq 0.05$. Error bars are represented as one standard error of the mean. 

255

256

257 **3 Results**

258

259 **3.1 Substrate-derived C respiration**

260

261 There was a significant effect of substrate chemistry on substrate mineralization ($P \leq 0.05$), with
262 respiration from glucose addition being the greatest (Fig. 1). Respiration rate was highly variable
263 among substrates in the first several days of incubation. After day 2 of incubation, the proportion
264 of added C respired as CO_2 for different soils was 18 to 28% from glucose, 12 to 16% from
265 starch, 0.2 to 5% from cinnamic acid and 0.1 to 0.4% from stearic acid. Thus, a considerable
266 initial delay was observed in the mineralization of C from cinnamic acid and stearic acid as
267 compared to glucose and starch. At the end of incubation, cumulative respiration for different
268 soils was 52 to 60% of added C for glucose, 39 to 49% for starch, 33 to 53 % for cinnamic acid
269 and 43 to 57 % for stearic acid. Respiration from substrates varied within a narrow range for the
270 Mollisol and the Andisol throughout the course of incubation compared to the Ultisol and the
271 Gelisol. At the end of incubation, the proportion of substrate C respired for all substrates
272 combined was 41 to 50% for the Mollisol, 43 to 54% for the Andisol, 33 to 57% for the Ultisol
273 and 39 to 60% for the Gelisol.

274

275 **3.2 SOC-derived C respiration**

276

277 The cumulative amount of native SOC mineralized at the end of experiments with unamended
278 soils varied from 2.4 to 4.1 mg C g^{-1} across the soils and substrate types (Table 2). Adding
279 substrates significantly affected the amount of native SOC mineralized from the Ultisol, the

280 Andisol and the Gelisol, but not from the Mollisol (Table 2, Fig. S2). Contrary to our hypothesis,
281 cinnamic acid and stearic acid additions resulted in mineralization of more native SOC than from
282 unamended control in all soils except the Mollisol. Compared to the unamended control,
283 cinnamic acid treatment caused 24% more mineralization of native SOC in the Ultisol, 36%
284 more in the Andisol, and 20% more in the Gelisol. Likewise, stearic acid addition caused 28%
285 more SOC mineralization in the Ultisol and the Andisol, and 30% more in the Gelisol.
286 Cumulative SOC mineralization from glucose and starch treated soils was statistically similar to
287 unamended soils (Table 2).

288

289 **3.3 Microbial community composition**

290

291 The fungal:bacterial (F:B) ratios were calculated from the fungal and bacterial gene copy
292 numbers measured by qPCR (Fig. S3, S4). $F:B \gg 1$ indicates fungal dominance and $F:B \ll 1$
293 indicates bacterial dominance. To compare the F:B ratios from the substrate amended and
294 unamended samples, we calculated the difference ($F:B_{\text{amended}} - F:B_{\text{unamended}}$) at each sampling
295 point (day 4, 30 and 270) (Fig. 2). Positive values indicate greater fungal (and lesser bacterial)
296 numbers in amended versus unamended soils, and negative values indicate smaller fungal (and
297 greater bacterial) numbers in amended versus unamended soils. Positive values were nearly
298 always observed for the Ultisol, the Andisol, and the Gelisol, and these values became more
299 positive over time, indicating increasing fungal dominance. At day 4, the difference between F:B
300 ratios between substrate amended and unamended soils was small, except for glucose addition to
301 the Mollisol and the Ultisol which showed fungal dominance. Cinnamic acid and stearic acid
302 addition exhibited enhanced fungal dominance by day 270 except for the Mollisol. Archaeal

303 gene copy numbers were also measured by qPCR and they were the lowest among the microbial
304 groups for all substrate types, soil orders and sampling times (data not shown). There was no
305 evident influence of substrate addition or length of incubation on archaeal numbers.

306

307 **3.4 Pools and rates associated with respiration**

308

309 Native SOC respiration was best modeled by the double pool exponential decay model.

310 Irrespective of the substrate treatments, the lowest proportion of the initial SOC was assigned to

311 labile pool (Pool 1) for the Andisol compared to other soils (Fig. 3A). The size of Pool 1 was

312 greater for stearic acid and cinnamic acid amended soils than for control soils and soils with

313 other substrates. For the Gelisol and the Ultisol, cinnamic acid and stearic acid addition yielded

314 lower mineralization rate k_1 associated with Pool 1, while no difference was observed for the

315 Mollisol or the Andisol (Fig. 3B). The mineralization rate k_2 corresponding to intermediate pool

316 (Pool 2) was statistically similar among the substrates for all soils, however, there was a notable

317 decrease in k_2 for the Andisol in comparison with other soils (Fig. 3C).

318 Modeling of substrate-derived respiration data was strongly dependent on substrate

319 chemistry: a triple pool exponential decay model was the best fit for the substrate-derived C

320 respiration following glucose and starch amendments, whereas a double pool model was the best

321 fit following cinnamic acid and stearic acid amendment (Fig. 4). When comparing modeled C

322 pools from cinnamic/stearic acid to glucose/starch amendments, Pool 1 of cinnamic/stearic acid

323 amended soils mostly equals or exceeds the combined size of Pool 1 and Pool 2 modeled from

324 glucose and starch respiration (Fig. 4). The mineralization rate k_1 associated with Pool 1

325 following glucose and starch amendments was one or two orders of magnitude greater than the

326 corresponding k_2 , which again was considerably greater than k_3 (Fig. 4B, 4C, 4D). Mineralization
327 rate k_1 of cinnamic acid and stearic acid respiration was closer to the k_2 of glucose and starch
328 respiration, and the k_2 following cinnamic acid and stearic acid respiration was equal to or lower
329 than k_3 following glucose and starch addition. Since two types of models were needed to best fit
330 the respiration data of two sets of substrates (~~3 pool model for glucose and starch, and 2 pool~~
331 ~~model for cinnamic acid and stearic acid~~), a statistical comparison of mineralization kinetics
332 across substrate types was not conducted. Pool sizes and rates showed variations across soil
333 types, but no overall consistent patterns were observed.

334

335 4 Discussion

336

337 4.1 Substrate-derived C respiration

338


339 In accordance with our hypothesis, substrate C mineralization rate and extent were influenced by
340 initial substrate quality (Fig. 1). Indeed, the greatest mineralization of substrate C occurred
341 following glucose addition (52-60% of added C); and in the initial days after substrate addition,
342 we observed more rapid mineralization of C from glucose and starch than from cinnamic acid
343 and stearic acid (Fig. 1). Our results with glucose and starch was quantitatively similar to
344 previous studies (Bremer and Kuikman, 1997; Jones and Murphy, 2007; Hoyle et al., 2008) and
345 in a similar experiment, Orwin et al. (2006) found that CO_2 respiration from sugars was greater
346 than respiration from fatty acids and tannin. Considerably higher CO_2 efflux in the first three
347 days of incubation was found when a synthetic root exudate cocktail containing 60% sugars,
348 35% organic acids and 2% amino acids was added to soils (de Graaff et al., 2010). The slower

349 degradation following starch addition in comparison to glucose addition in our study could be
350 due to the requirement of extracellular enzymes (α -glucosidase) for starch hydrolysis to occur
351 (Kelley et al., 2011; German et al., 2012), while glucose can be directly assimilated by microbes.

352 Contradictory to the general notion that the fast growing sugar feeders are composed mostly
353 of bacterial species (Paterson et al., 2007; Moore-Kucera and Dick, 2008), enhanced F:B values
354 at day 4 following glucose addition in our study indicates that some fungi responded quickly to
355 substrate addition (Broeckling et al., 2008; Chiginevaa et al., 2009; de Graaff et al., 2010).

356 Panikov (1995) and Rinnan and Bååth (2009) also observed fungal-controlled mineralization of
357 glucose in the initial phase of similar microcosm studies. Addition of a synthetic root exudate
358 mixture containing 60% simple sugars resulted in a higher proportion of fungal growth relative
359 to bacterial growth at day 3 (de Graaff et al., 2010).

360 Sugars and other easily assimilable substrates added to soil are used by microbes not only for
361 the production of energy and release of CO₂, but also for the biosynthesis of products including
362 extracellular enzymes, extracellular polysaccharides, cell wall polymers, storage compounds and
363 stress response compounds (Nguyen and Guckert, 2001; Dijkstra et al., 2011; Schimel and
364 Schaeffer, 2012). The proportion of C initially allocated for biosynthetic processes may take
365 more time to mineralize to CO₂. Consequently, we observed continued evolution of ¹⁴CO₂ even
366 after several months of incubation from all the added substrates (including the most labile
367 glucose), albeit at a slower rate. Therefore it is very likely that part of the added sugars may have
368 been used as biosynthetic precursors and those microbial byproducts contributed to the evolution
369 of ¹⁴CO₂ during the later stages of incubation.

370 **The metabolism of C substrates in soil is a function of microbial community structure,** 
371 relative access that different groups of microbes have to these substrates (Schimel and Schaeffer,

2012). Relative controls imposed by soil biology and physics, however, are not readily apparent. We observed that mineralization of C from cinnamic acid and stearic acid was delayed for several days (Fig. 1). However, this delay was not due to the decreased abundance of microbial activity because native SOC mineralization was similar to control. Specialized microorganisms might be responsible for the mineralization of these relatively complex compounds, and these organisms were either low in abundance in the beginning of the experiment, or the organisms simply took more time to consume and cycle these compounds. Degradation requires the production of specific extra-cellular enzymes before they can be utilized (German et al., 2011). Sorption to the soil mineral phase could be another reason for the delayed respiration, because our previous experiments showed considerably higher affinity of stearic acid to soil minerals in comparison with other compounds (Jagadamma et al., 2014). The eventual decomposition of stearic acid in this study, however, suggests that sorption did not protect stearic acid over longer time frames.

385

386 **4.2 SOC-derived C respiration**

387

We found that the chemical composition of substrates added to soils altered the stability of native SOC, but the results were different than what we originally hypothesized. Surprisingly, cumulative native SOC mineralization showed an increase due to cinnamic acid and stearic acid addition relative to glucose and starch additions and unamended soils (Table 2), and further, the increase in decomposition was only evident after several weeks of incubation (Fig. S2).

Literature on substrate-controlled difference in SOC mineralization is scanty and the limited studies available mostly used simple sugars and organic acids as substrates to understand the

395 SOC mineralization process. In our study, we consider cinnamic acid and stearic acid as more
396 complex C compounds than glucose and starch because of the higher hydrophobicity of both
397 compounds, aromatic structure of cinnamic acid, and strong mineral sorption capacity of stearic
398 acid (Orwin et al., 2006; Jagadamma et al., 2014). In a similar study, Brant et al. (2006)
399 measured SOC mineralization following the addition of glucose, glutamate, oxalate and phenol
400 from a forest soil in Oregon and found that more SOC was mineralized with oxalate and phenol
401 addition compared to glucose and glutamate addition. We also found that cinnamic acid and
402 stearic acid additions were associated with higher F:B gene copy ratios during the final stages of
403 incubation relative to other substrates (Fig. 2). It is possible that the addition of cinnamic acid
404 and stearic acid might have activated some specialized, but slow-growing fungal populations
405 capable of decomposing more recalcitrant components of SOC at the later stages of incubation.
406 Overall, our study reveals that both initial substrate quality and decomposer community are
407 tightly linked and interactively influence the decomposition of both substrate and soil C.

409 4.3 Pools of carbon and rates of decomposition

410
411 Modeling of C pool sizes and mineralization rates from incubation-derived data are used for
412 improved parameterization of ecosystem models. The cumulative CO₂ respiration following
413 substrate addition was best described using a double or triple pool first order exponential decay
414 model, for both amended and unamended soils (Chen et al., 2009; Farrar et al., 2012). The
415 substrate C respiration following glucose and starch addition was best fit by a triple pool model
416 (fast, intermediate and slow pools) and cinnamic acid and stearic acid additions were best fit by a
417 double pool model, i.e., fast and intermediate pools (Fig. 4). Farrar et al. (2012) also reported

418 that a triple pool model was the best fit for the glucose derived CO₂. The need for two types of
419 models for sugars vs. complex compounds indicates that initial substrate quality hypothesis hold
420 true for the decomposition of C input (Wickings et al., 2012). The native SOC-derived CO₂ data
421 was best modeled using a double pool model regardless of the type of substrate addition (Fig. 3)
422 and the length of incubation experiment could be a determinant for the lack of effect of substrate
423 type on native C pool partitioning because incubation length reflects the contribution of more
424 recalcitrant pools in the total CO₂ efflux (Schädel et al., 2013). Shorter-term incubation data is
425 often dominated by the CO₂ from more labile C fractions. Using 385 days of decomposition data,
426 Schädel et al. (2013) did not find any improvement in the fit for SOC decomposition data when a
427 three pool model was used over a two pool model, and the dominance of the third pool became
428 more evident only after 230 days of study. Scharnagl et al. (2010) reported that decomposition
429 data from a 900-day incubation experiment was sufficient in constraining all the five C pools in
430 RothC model. In our study, within 270 days only 5 to 20% of initial SOC was lost across all soils
431 and substrate addition treatments (Table 2) and it appears that 270 day incubation was not long
432 enough to constrain parameters for the third native SOC pool. This differs from our substrate C
433 modeling in which three pools were used for glucose and starch but only two pools for cinnamic
434 acid and stearic acid. These findings support the need for more long-term studies using more
435 complex substrates. Substrate-specific mineralization kinetics are useful for refining the
436 decomposition rates and pools in C cycle models.

437

438 **5 Conclusion**

439

440 This study reveals that substrate quality imparts considerable control on microbial decomposition
441 of substrates and native OC, and also calls for multiple year incubation experiments to capture

442 the dynamics of the recalcitrant fraction of the OC pool. We found that even though complex
443 substrates (cinnamic acid and stearic acid) showed an initial delay in respiration compared to
444 simpler substrates (glucose and starch), complex substrates caused enhanced mineralization of
445 SOC at later stages of incubation with a concomitant increase in fungal abundance. However, the
446 length of incubation was not long enough to fully characterize decomposition kinetics of more
447 complex substrates (cinnamic acid and stearic acid) and native SOC. This study suggests the
448 need for more detailed experiments investigating the role of substrate quality on C
449 mineralization, and the need to design experiments to capture the dynamics of both the labile and
450 recalcitrant fraction in soils. Characterizing these dynamics is critical as anthropogenically-
451 induced changes in atmospheric CO₂ concentration and N deposition are predicted to alter the
452 quality of both above ground and below ground C input to soils. Thus, understanding the control
453 of substrate chemistry or quality on soil microbial composition and function will be useful to
454 predict the future impact of climate change on SOC dynamics.

455

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464

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687 **Table 1.** Sampling locations and pre-incubation soil properties.

Descriptions	Soils			
	Mollisol	Ultisol	Andisol	Gelisol
Sampling location	Batavia, Illinois, USA	Lavras, Minas Gerais, Brazil	Krýsuvíkurheiði, Reykjanes, Iceland	Fairbanks, Alaska, USA
Organic C (g kg ⁻¹)	29.8 ± 0.50	23.2 ± 1.2	74.5 ± 0.10	20.5 ± 0.10
Total N (g kg ⁻¹)	3.00 ± 0.02	1.97 ± 0.08	7.09 ± 1.08	1.32 ± 0.02
Microbial biomass C (mg kg ⁻¹)	640 ± 35	515 ± 42	856 ± 39	48 ± 2.30
pH (1soil:2H ₂ O)	7.64 ± 0.10	5.42 ± 0.01	5.84 ± 0.01	7.03 ± 0.10
Silt (g kg ⁻¹)	570 ± 30	170 ± 20	570 ± 46	790 ± 49
Clay (g kg ⁻¹)	350 ± 15	450 ± 32	120 ± 08	130 ± 11

688 Values are mean ± standard error (n=3).

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695 **Table 2.** Cumulative soil organic C (SOC) respiration after 270 days.

Soils	Substrates	SOC respiration	
		mg C g ⁻¹ soil	% of initial C lost
Mollisol	Unamended	4.1±0.21	13.9
	Glucose	4.0±0.07	13.3
	Starch	3.9±0.15	13.1
	Cinnamic acid	4.0±0.18	13.6
	Stearic acid	4.3±0.12	14.4
Ultisol	Unamended	3.5±0.13 ^b	15.3
	Glucose	3.6±0.15 ^b	15.6
	Starch	4.0±0.10 ^{ab}	17.4
	Cinnamic acid	4.4±0.23 ^a	19.0
	Stearic acid	4.5±0.04 ^a	19.6
Andisol	Unamended	3.4±0.15 ^b	5.3
	Glucose	3.5±0.04 ^b	5.5
	Starch	3.6±0.11 ^b	5.5
	Cinnamic acid	4.6±0.21 ^a	7.2
	Stearic acid	4.4±0.04 ^a	6.8
Gelisol	Unamended	2.4±0.6 ^b	11.8
	Glucose	2.7±0.17 ^b	12.9
	Starch	2.6±0.0 ^b	12.6
	Cinnamic acid	2.9±0.11 ^{ab}	14.2
	Stearic acid	3.1±0.11 ^a	15.3

697 Each value represents mean ± standard error (n=3). Different letters indicate significant mean
 698 differences among substrate addition treatments in each soil.

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706 **Figure captions**

707 **Fig. 1.** Substrate C respiration in response to the addition of four substrates in Mollisol (A),
708 Ultisol (B), Andisol (C), and Gelisol (D). Symbols represent proportion of added substrate C
709 respired at each sampling time along with standard error bar (n=3).

710

711 **Fig. 2.** The difference in Fungal:Bacterial gene copy ratios between amended and unamended
712 treatments ($F:B_{\text{amended}} - F:B_{\text{unamended}}$) in response to the addition of four substrates in Mollisol (A),
713 Ultisol (B), Andisol (C), and Gelisol (D). * indicates that $F:B_{\text{amended}} - F:B_{\text{unamended}}$ is significantly
714 different from zero.

715

716 **Fig. 3.** Effect of substrate types on native soil organic carbon mineralization parameters: pool
717 sizes (A), mineralization rate k_1 associated with Pool 1 (B), and mineralization rate k_2 associated
718 with Pool 2 (C).

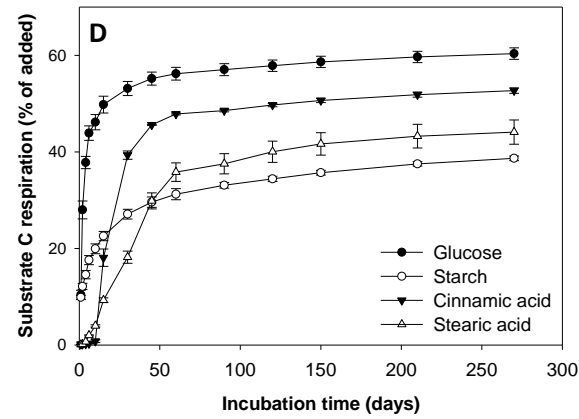
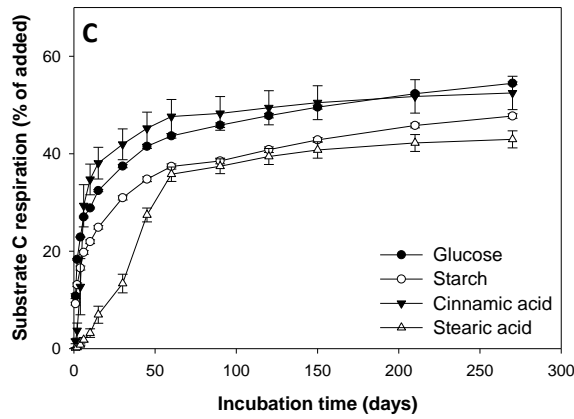
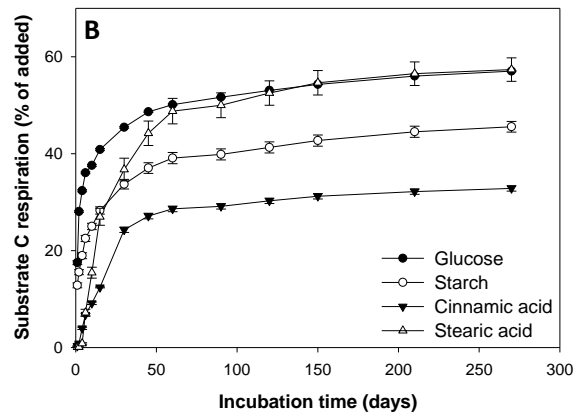
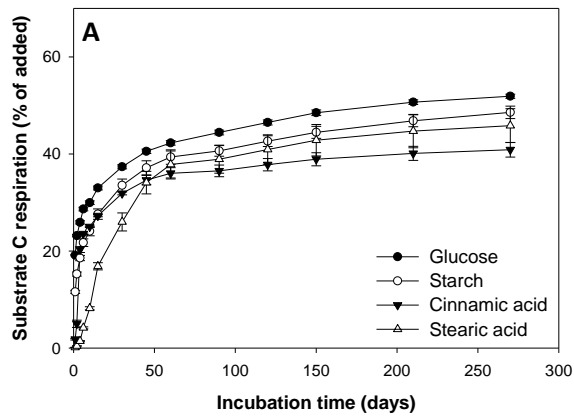
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720 **Fig. 4.** Effect of substrate types on substrate C mineralization parameters: pool sizes (A),
721 mineralization rate k_1 (B), mineralization rate k_2 (C), mineralization rate k_3 (D). Pool sizes of
722 glucose-C and starch-C respiration (fast, intermediate and slow pools) and their associated
723 mineralization rates (k_1 , k_2 and k_3) were best modeled by a triple pool model, and pool sizes of
724 cinnamic acid-C and stearic acid-C respiration (fast and intermediate pools) and their associated
725 mineralization rates (k_1 and k_2) were best modeled by a double pool model. Bars are mean \pm
726 standard error (n=3).

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732 **Fig. 1.**

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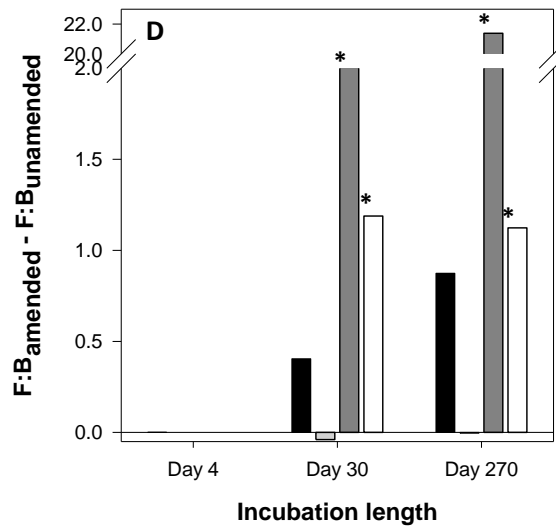
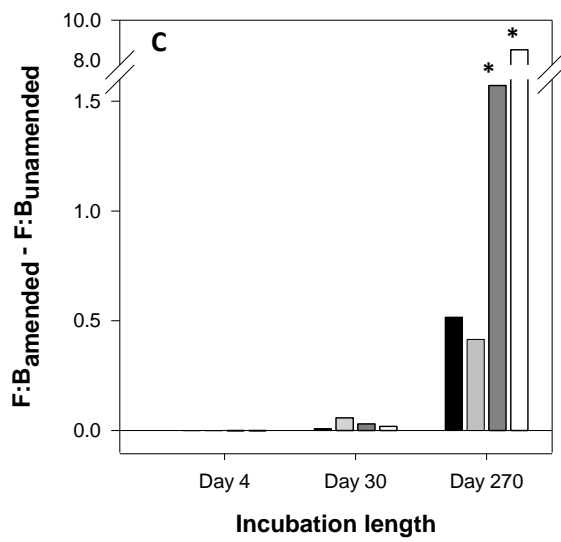
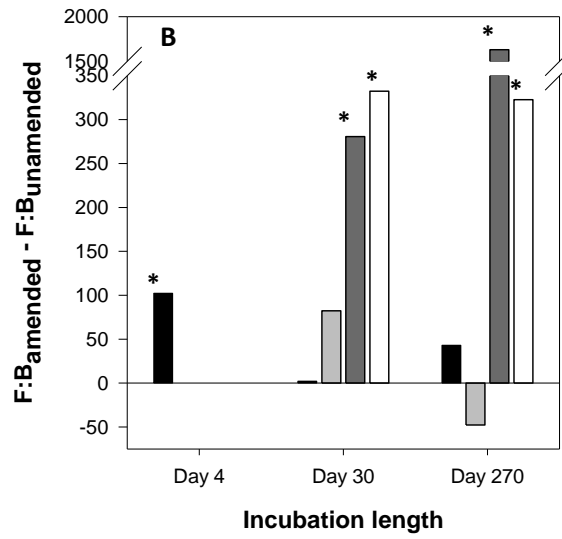
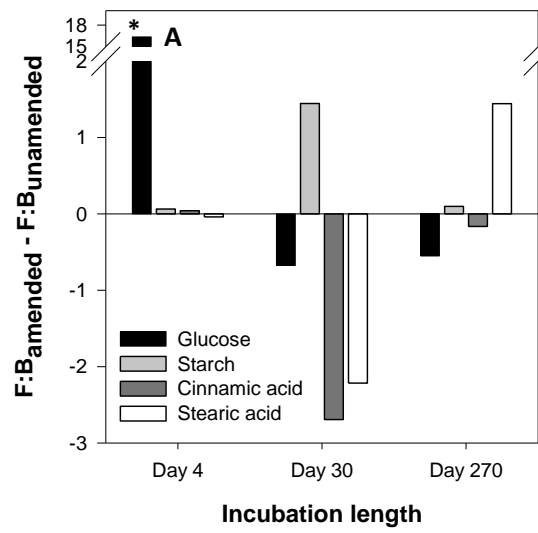
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743 **Fig. 2.**

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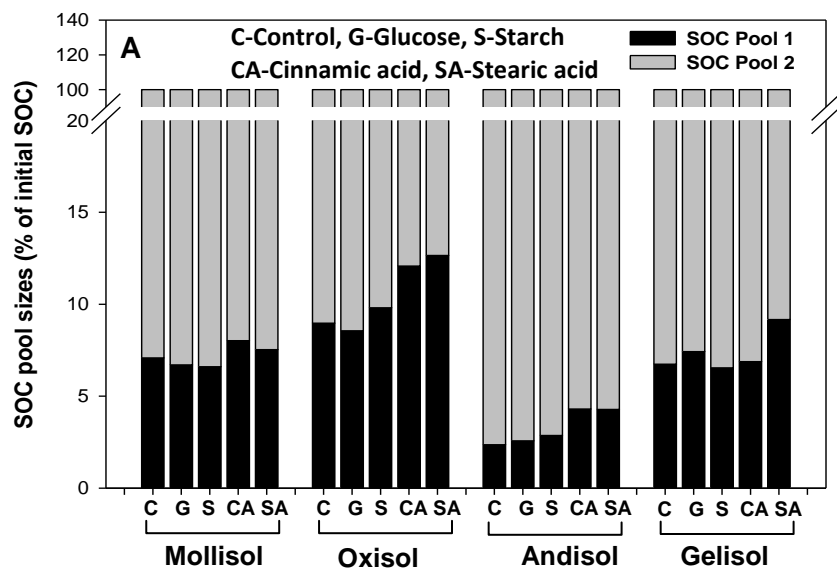
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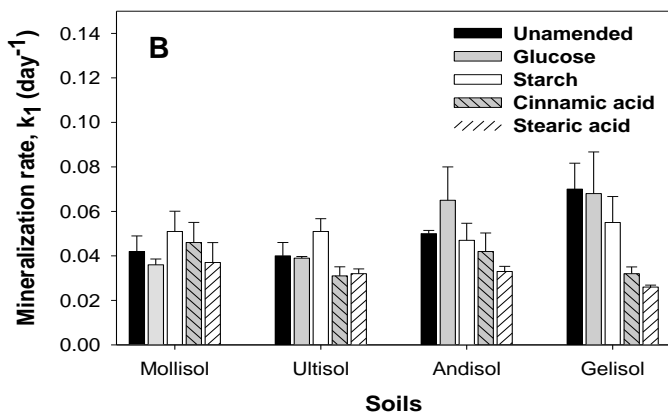
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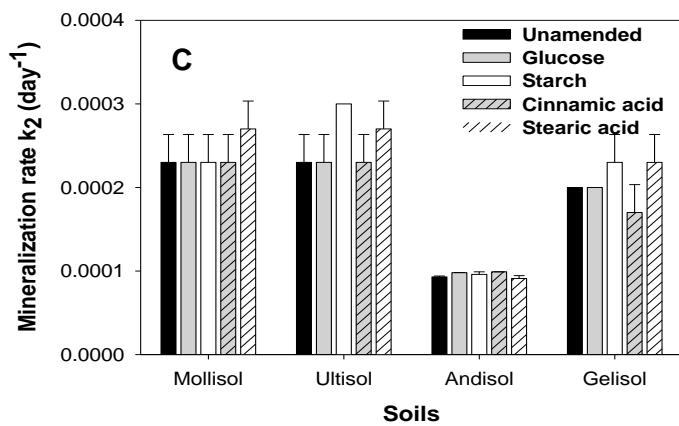
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754 **Fig. 3**

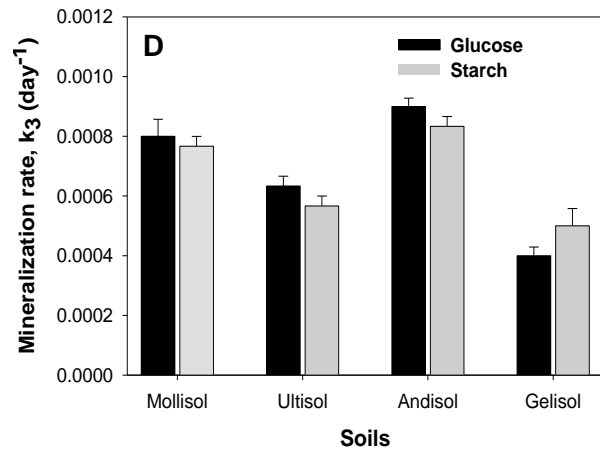
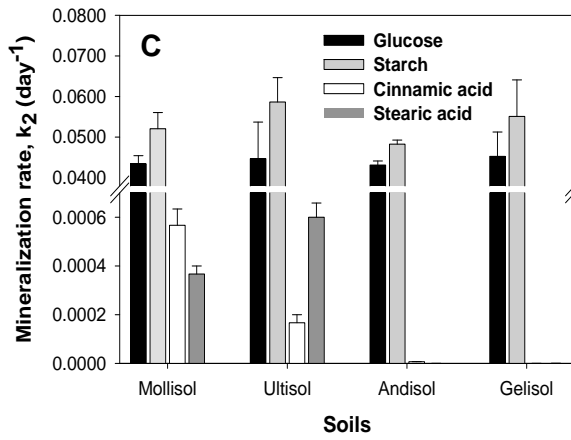
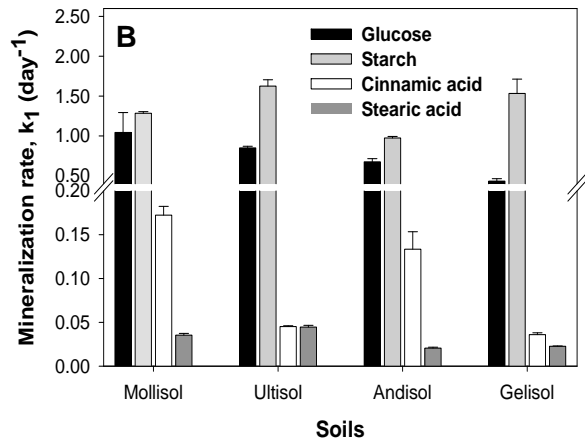
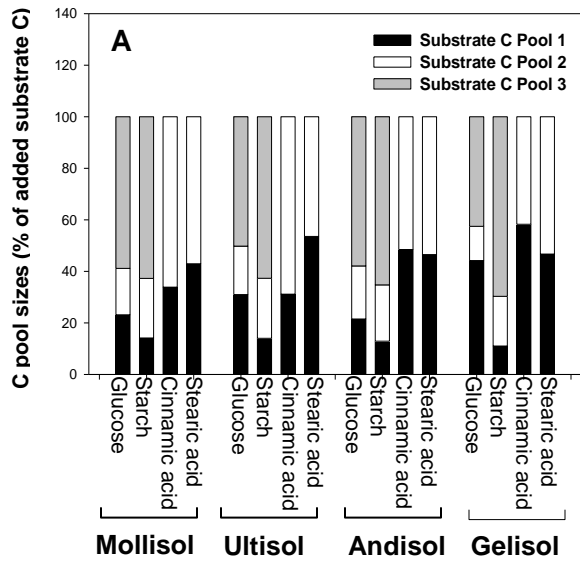


Fig. 4.