1	Enhanced viral production and virus-mediated mortality of
2	bacterioplankton in a natural iron-fertilized bloom event
3	above the Kerguelen Plateau
4	
5	Andrea Malits ^{1,2} , Urania Christaki ³ , Ingrid Obernosterer ^{4,5} and Markus G.
6	Weinbauer ^{1,2}
7	
8	
9	¹ Sorbonne Universités, UPMC Univ Paris 06, UMR 7093, LOV, Observatoire océanographique,
10	F-06230, Villefranche/mer, France
11	² CNRS, UMR 7093, LOV, Observatoire océanographique, F-06230, Villefranche/mer, France
12	³ INSU-CNRS, UMR 8187 LOG, Laboratoire d'Océanologie et des Géosciences, Université du
13	Littoral, ULCO, 32 avenue Foch, F-62930 Wimereux, France
14	⁴ Sorbonne Universités, UPMC Univ Paris 06, UMR 7621, LOMIC, Observatoire
15	océanographique, F-66650 Banyuls/mer, France
16	⁵ CNRS, UMR 7621, LOMIC, Observatoire océanographique, F-66650 Banyuls/mer, France
17	
18	
19	Correspondence to: A. Malits (<u>amalits@cadic-conicet.gob.ar</u>)
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Running title: Enhanced viral production in iron-fertilized waters

1 Abstract

2 Above the Kerguelen Plateau in the Southern Ocean natural iron fertilization sustains a large 3 phytoplankton bloom over three months during austral summer. During the KEOPS1 project 4 (KErguelen Ocean and Plateau compared Study1) we sampled this phytoplankton bloom during 5 its declining phase along with the surrounding HNLC waters to study the effect of natural iron 6 fertilization on the role of viruses in the microbial food web. Bacterial and viral abundances were 7 1.7 and 2.1 times, respectively, higher within the bloom than in HNLC waters. Viral production and virus-mediated mortality of bacterioplankton was 4.1 and 4.9 times, respectively, higher in 8 9 the bloom, while the fraction of infected cells (FIC) and the fraction of lysogenic cells (FLC) 10 showed no significant differences between environments. The present study suggests viruses to be more important for bacterial mortality within the bloom and dominate over grazing of 11 heterotrophic nanoflagellates (HNF) during the late bloom phase. As a consequence, at least at a 12 13 late bloom stage, viral lysis shunts part of the photosynthetically fixed carbon in iron-fertilized 14 regions into the dissolved organic matter (DOM) pool with potentially less particulate organic 15 carbon transfered to larger members of the food web or exported.

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2 1 Introduction

A quarter of a century ago the importance of viruses as the most abundant biological entity in the 3 4 oceans (Bergh et al., 1989) and their role in the material and energy cycles have been recognized 5 (Proctor and Fuhrman, 1990; Suttle et al., 1990). Shortly afterwards, Smith et al. (1992) 6 conducted the first study on viral distribution and their relationship to bacteria in the Southern 7 Ocean. Since then studies on viral abundance and production or infectivity in the cold high 8 latitude marine environments remained limited or are only recently accumulating (Bird et al., 9 1993;Brussaard et al., 2008b;Evans and Brussaard, 2012;Evans et al., 2009;Guixa-Boixereu et 10 al., 2002;Higgins et al., 2009;Manganelli et al., 2009;Marchant et al., 2000;Payet and Suttle, 11 2008, 2013;Smith et al., 1992;Steward et al., 1996;Strzepek et al., 2005;Weinbauer et al., 2009a). 12 These observations demonstrate that viruses are ecologically as important in these cold 13 environments as in other world's oceans.

14 Viral lysis of cells converts particulate organic matter into dissolved and colloidal organic matter, 15 reduces the carbon flow to higher trophic levels and increases the residence time of carbon and mineral nutrients in the euphotic zone (Fuhrman, 1999). By this process, called the "viral shunt" 16 17 (Wilhelm and Suttle, 1999), heterotrophic bacteria are supplied with substrate, which finally 18 increases respiration (Bonilla-Findji et al., 2008;Middelboe and Lyck, 2002). This could reduce 19 the efficiency of the biological carbon pump, i.e. the process which transforms inorganic to 20 organic carbon, part of which is then transferred to the deep ocean (Suttle, 2007). The relative 21 significance of viral lysis and protistan grazing can strongly vary on temporal and spatial scales 22 (Boras et al., 2009;Fuhrman and Noble, 1995). This has also been shown for cold marine 23 environments (Boras et al., 2010;Guixa-Boixereu et al., 2002;Steward et al., 1996;Wells and 24 Deming, 2006).

In about one-third of the World Ocean, including the sub-arctic northeast Pacific, the equatorial Pacific and the Southern Ocean, phytoplankton growth is limited by available iron resulting in excess phosphorus and nitrogen (Martin and Fitzwater, 1988). In these high nutrient-low Andrea Malits 9/10/14 2:56 P.M. Eliminado: living

chlorophyll (HNLC) regions, bacterioplankton are thought to be the key player of the « microbial
 ferrous wheel » (Kirchman, 1996), i.e. the uptake and remineralization of iron. Bacterioplankton
 contain more than twice iron per carbon units than eukaryotic phytoplankton, and they can
 thereby store up to 50% of the biogenic iron in the HNLC ocean (Tortell et al., 1996).

Viral activity has a potential impact on nutrient regeneration. Typically, nutrients released as a 5 result of viral lysis are thought to be organically complexed, which may facilitate their use by 6 7 marine plankton (Poorvin et al., 2004; Rue and Bruland, 1997). Iron released by viral lysis can 8 account for more than 10% of ambient Fe concentrations (Gobler et al., 1997) and thus 9 potentially relieve its limitation in depleted environments. Furthermore, marine viruses may serve 10 as nuclei for iron adsorption and precipitation, and thus represent a significant reservoir of iron in seawater (Daughney et al., 2004). Despite their key role, viruses are hardly included in iron 11 12 enrichment studies. These experiments were originally stimulated by the "iron hypothesis" 13 (Martin, 1990) which assigns iron a paramount role in controling ocean productivity and 14 consequently atmospheric carbon dioxide concentrations. Only two out of 13 iron fertilization 15 experiments so far performed (Secretariat of the Convention on Biological Diversity, 2009) 16 report on viral abundance and activity (Higgins et al., 2009;Weinbauer et al., 2009a). Both 17 studies from the Subartic and Southern Ocean, respectively, found that viral production was 18 significantly enhanced after iron fertilization.

19 Above the Kerguelen Plateau in the Southern Ocean, the largest HNLC ocean, a large 20 phytoplankton bloom occurs annually during austral summer. The continuous supply of Fe and 21 major nutrients from below has been shown to sustain this massive bloom (Blain et al., 2007). 22 The region off Kerguelen provides the opportunity to study natural iron fertilization in the 23 Southern Ocean and to compare it to blooms induced by mesoscale Fe additions. Within the 24 KEOPS1 project (KErguelen Ocean and Plateau compared Study1, 2005-2007), we sampled the phytoplankton bloom above the Kerguelen plateau during its late successional stage (~ 3rd month) 25 26 along with the surrounding HNLC waters. The aim of the present study was to assess the role of 27 viruses within the microbial food web affected by natural Fe fertilization and to elucidate the

- 1 possible implications for the final destiny of organic carbon. For this purpose, we measured viral
- 2 production, the fraction of infected cells (FIC), lysogeny and estimated bacterial mortality

5

3 through viral lysis in the bloom and surrounding HNLC waters.

1 2 Material and Methods

2 2.1 Description of the study site

Sampling was performed in the Indian sector of the Southern Ocean above the Kerguelen Plateau 3 4 $(49^{\circ}-53^{\circ} \text{ S and } 72^{\circ}-78^{\circ} \text{ E})$ in austral summer (18 January – 13 February 2005) on board the R/VMarion Dufresne in the framework of the project KEOPS (Blain et al., 2008). We sampled a 5 6 large phytoplankton bloom dominated by diatoms from its peak to its decline (Mosseri et al., 7 2008) (Figure 1). Satellite images dated the onset of this bloom more than two month before its 8 first visit (Blain et al., 2007). Hydrographic conditions are described in detail in Park et al. 9 (2008). Dissolved Fe concentrations in the surface mixed layer were low and similar on and off the plateau $(0.09 \pm 0.03 \text{ nM})$ but increased with depth above the plateau reaching a mean 10 11 maximum of 0.35 nM at 500m. This strong vertical gradient in combination with physical 12 features such as internal waves and tidal activity sustained the phytoplankton bloom above the 13 plateau (Blain et al., 2007).

14 2.2 Sampling strategy

15 Water was collected using General Oceanics 12L Niskin bottles mounted on a rosette with a Sea 16 Bird SBE19 plus CTD sensor for salinity, temperature and oxygen from 2-3 depths (within and 17 below the surface mixed layer) at the following stations to cover the centre and borders of each of 18 three transects (A, B, C): A3, A11, B1, B5, B11, C3 and C11 (Figure 1, Table 1). The stations A3 19 and C11 were considered as the most contrasting stations and sampled repeatedly. The first 20 sampling of station A3 (A3-1) was done during the peak of the bloom, and about two weeks later, 21 station A3 was re-sampled at a fourth visit (A3-4) during the decline of the bloom. Station B5 22 was situated within a new phytoplankton bloom above the Kerguelen plateau (Obernosterer et al., 23 2008). Station A11 was located in iron-fertilized waters. The annually occuring spring bloom 24 developped prior to our visit, explaining the low concentrations of chl a at this site (Table 1). 25 Stations B11, C3 and C11 were in HNLC waters off the Kerguelen plateau and the latter was 26 sampled twice (Table 1).

27 Total chlorophyll a (Chl a and divinyl-Chl a) was measured by High Performance Liquid

1 Chromatography (HPLC, Van Heukelem and Thomas, 2001, Uitz et al., 2009).

2 2.3 Enumeration of viruses and prokaryotes

3 Subsamples (2 mL) were fixed with glutaraldehyde (0.5% final concentration), incubated for 15-4 30 minutes at 4°C, subsequently frozen in liquid nitrogen and stored at -80°C. Within a few days 5 samples were thawed and viral particles and bacteria were stained with SYBR Green I (Molecular Probes) and quantified using a FACScalibur (Becton and Dickinson) flow-cytometer 6 7 after dilution with TE buffer (10mM Tris, 1 mM EDTA, ph = 8). For viruses an optimized 8 protocol by Brussaard (2004) was followed. Viruses and prokaryotes were determined in plots of 9 90° light scatter (SSC) and green DNA fluorescence. Differences in the green fluorescence and 10 side scatter signature in the cytometric plot allowed to separate prokaryotes with low nucleic acid 11 content (LNA) from prokaryotes with high nucleic acid content (HNA) as previously described 12 by Gasol et al. (1999). Similarly, different size classes of viruses were distinguished on the basis 13 of green fluorescence. Abundances were calculated by the used flow rate. Flow-cytometric 14 assessment of viral abundance may encompass particles other than viruses such as bacterial 15 vesicles (Biller et al., 2014). However, since bacterial and viral parameters were related 16 significantly (Table 4), a potential overestimation of viral abundances did probably not bias the 17 conclusions of the study.

18 To convert bacterial abundance (BA) to biomass we used a conversion factor of 12.4 fg C cell⁻¹
19 for oceanic prokaryotes (Fukuda et al., 1998).

20 2.3 Bacterial Production

The incorporation of ³H leucine into protein (Smith and Azam, 1992) was used to estimate the production of heterotrophic bacteria (BP). At each depth, 1.5 mL duplicate samples and a trichloroacetic acid (TCA)-killed control were incubated with a mixture of L-[4,5 -3H] leucine (Amersham, 160 Ci mmol-1) and nonradioactive leucine added at final concentrations of 7 nM and 13 nM for the upper 100 m, and 13 nM and 7 nM for the 100-200 m depth layer. Samples were incubated in the dark at the ambient temperature of the depth where samples were collected. The incubation time (2-3h) was tested to satisfy linear incorporation with time. We checked by

concentration kinetics (2.5, 5, 10, 20 and 40 nM), at 3 stations inside and outside the bloom at 5
m and 175 m depths that there was no isotopic dilution. The theoretical conversion factor of 1.55
kg of C mol-1 was used to convert leucine incorporation rates to prokaryotic carbon production
(Kirchman, 1993).

5 2.4 Viral production, the fraction of infected cells and the fraction of lysogenic

6 cells

7 Lytic viral production (VP_i) , the fraction of infected cells (FIC), induced viral production (VP_i) 8 and the fraction of lysogenic cells (FLC) were estimated using the virus reduction approach 9 (VRA, Weinbauer et al., 2009b; Wilhelm et al., 2002). The rationale behind VRA is to reduce 10 viral abundance in order to stop new viral infection. Thus, the viruses produced originate from 11 already infected cells. Briefly, bacteria from 200 mL raw seawater were concentrated using a 12 tangential flow system with a peristaltic pump (Watson-Marlow 323) equipped with a 0.2 μ m 13 cartridge (VIVAFLOW 50). To obtain virus-free seawater, the 0.2 μ m pore-size ultrafiltrate was 14 passed through a 100kDalton cartridge (VIVAFLOW 50). The bacterial concentrates were 15 brought up to the original volume with virus-free seawater and incubated in duplicate 50 mL 16 Falcon tubes in the dark at $\pm 2^{\circ}$ C in situ temperature for 24 hours. Lysogeny was estimated by adding mitomycin C (SigmaChemical Co, No. M-0503, final concentration 1 μ g mL⁻¹) to 17 18 duplicate 50mL Falcon tubes in order to induce the lytic cycle in lysogens; untreated duplicate 19 samples served as controls (Paul and Weinbauer, 2010). Subsamples (2mL) for viral and bacterial 20 abundance from each incubation were taken immediately (t_0 samples) and every 3-4 hours, fixed 21 with glutaraldehyde (0.5% final concentration), incubated for 15-30 minutes at 4°C, subsequently 22 frozen in liquid nitrogen and stored at -80°C until enumeration using a flow-cytometer as 23 described above. VP was calculated as

24 $VP_1 = (V_2 - V_1) / (t_2 - t_1)$

where V_1 and V_2 are viral abundances and t_1 and t_2 the elapsed time. Dividing the number of produced phages by an estimated burst size (BS, i.e. the number of phages released during the

(1)

- 1 lysis of a single host) yields the number of lysed cells and thus gives an estimation of FIC
- 2 (Weinbauer et al., 2002). FIC was calculated as
- 3 $FIC = 100 x ([V_2 V_1] / BS / BA$ (2)
- 4 where BA is the bacterial abundance at t_0 . The difference in phage abundance between lysogeny
- 5 treatment and the control is the number of induced phages, which is divided by BS to estimate the
- 6 fraction of lysogenic cells (FLC). FLC was calculated as
- 7 $FLC = 100 \text{ x} \left(\left[V_{MC} V_C \right] / BS / BA \right)$ (3)
- 8 where V_{MC} and V_{C} is the maximum difference in viral abundance at corresponding time points in
- 9 control and mitomycin C treatments, respectively. Calculations were performed for each replicate10 separately.
- 11 2.5 Contact rates
- 12 The rates of contact (R, number mL⁻¹ d⁻¹) between viruses and bacteria were calculated by using
- 13 the following equations (Murray and Jackson, 1992).
- $14 \qquad R = Sh \ge 2\pi d \ge D_V \ge VA \ge BA, \tag{4}$
- 15 where Sh is the Sherwood number [1.06 for a bacterial community with 10% motile cells;
- 16 (Wilhelm et al., 1998)], d is the diameter of the target, VA and BA are the abundances of viruses
- 17 and bacteria, respectively, and D_V is the diffusivity of viruses.
- 18 $D_V = k \ge T/(3 \ge \pi \ge \mu \ge d_V) = 5 \ge 10^{-8} \text{ cm}^2 \text{ s}^{-1},$ (5)
- 19 where *k* is the Boltzmann constant (1.38 x 10^{-23} J K⁻¹), *T* is the *in situ* temperature (~275 K), μ is 20 the viscosity of water (Pascal s⁻¹) and d_v is the diameter of the viral capsid (~60 nm). The contact 21 rates were divided by *in situ* bacterial abundance to estimate the number of contacts per cell on a 22 daily basis.

23 2.6 Bacterial mortality

To obtain the rate of cell lysis, viral production corrected for *in situ* bacterial abundance was divided by an estimated BS following the approach of Wells and Deming (2006), i.e. dividing the number of viruses produced during the first hours of incubation by the concomitant decline of

- 1 bacterial abundance. The number of lysed bacteria was converted into carbon by the factor of
- 2 12.4 fg C cell⁻¹ (Fukuda et al., 1998). The fraction of bacterial mortality through viral lysis
- 3 (VMM) was calculated following the Model by Binder (1999).
- 4 $VMM = FIC/LN(2) \times (1 0.186 FIC))$ (6)

5 2.7 Statistics

6 Normal distribution of data was checked using the Shapiro-Wilk W-test. Differences between

- 7 different trophic situations were analyzed by the Kruskal-Wallis test for non parametric data and
- 8 by one-way ANOVA for normally distributed data. Spearman rank correlation for non normally
- 9 distributed data was applied. Significance was considered for P < 0.05.
- 10

1 3 Results

2 3.1 Bacterial and viral abundances

From surface water down to 200m, BA was on average 1.7 fold higher within the Fe-fertilized 3 $(3.9 \times 10^5 \text{ mL}^{-1})$ than in HNLC waters $(2.4 \times 10^5 \text{ mL}^{-1})$. Kruskal-Wallis test, P < 0.0001, Table 2; 4 Figure 2). Similarly, viral abundance (VA) averaged 9.9 x 10^6 mL⁻¹ at the Fe-fertilized stations 5 and was twice as high than in the HNLC environments (4.7 x 10⁶ particles mL⁻¹, Kruskal-Wallis 6 test, P < 0.05, Table 2). VA ranged from 3.1 - 14.2 x 10⁶ mL⁻¹ with the highest values found at 7 8 the main bloom station A3 and the lowest value detected in the deep layer of the HNLC station 9 B11. Viruses were homogeneously distributed with depth at the HNLC stations. The virus to bacteria ratio (VBR) ranged from 11 to 34 and averaged 21 without significant differences 10 11 between stations or trophic situations.

12 3.2 Contact rates

13 Contact rates were significantly higher at the Fe-fertilized stations than in HNLC waters 14 (Kruskal-Wallis test, P < 0.05, Table 2). At the Fe-fertilized stations, on average 29.4 ± 11.1 15 viruses contacted a bacterial cell per day while in the HNLC waters contact rates were 14.2 ± 4.4 16 viruses cell⁻¹ d⁻¹ with the highest values at the bloom station A3 and the lowest at the HNLC 17 station B11 in accordance to the highest and lowest viral abundances, respectively (see Figure 2).

18 **3.3** Bacterial production, viral production, fraction of infected cells, lysogeny

Bacterial production ranged from 0.1 - 0.7 μ gC L⁻¹ d⁻¹ at the HNLC stations and from 0.1 - 2.5 μ gC L⁻¹ d⁻¹ at the Fe-fertilized stations (Table 2). The highest values were found throughout the depth profile of the main bloom station A3-1 and the lowest values were measured between 150 and 200m at the HNLC stations. Despite the wide range of values, BP was on average four times higher at the Fe-fertilized stations than at the HNLC stations (Kruskal-Wallis test, *P* < 0.0001, Table 2).

- 1 Initial virus abundance in the VRA was $45 \pm 25\%$ (11 88%) of *in situ* abundance. Pre-filtration 2 through 0.2μ m and 100kDa cartridge to obtain virus-free water and bacterial concentrate and
- 3 adding back resulted on average in recovery efficiencies of $26\% \pm 18\%$ (5-83%).
- 4 Lytic viral production (VP₁), corrected for *in situ* bacterial abundance averaged 59.0 x 10^6 mL⁻¹d⁻¹
- 5 in the naturally Fe-fertilized patch compared to 14.5 x 10^6 mL⁻¹ d⁻¹ in the HNCL environments.
- 6 This 4.1 fold difference was significant (Kruskal-Wallis test, P < 0.05, Table 2). Induced viral
- 7 production (VP_i) was detected in 4 out of 9 stations (3 fertilized and 1 HNLC stations, Table 3)
- 8 and averaged 44.8 \pm 44.2 x 10⁶ mL⁻¹ d⁻¹ (Table 2). VP₁ at the main bloom station A3 at 50m
- 9 increased from the first visit (15.6 x 10^6 mL⁻¹ d⁻¹) to the fourth visit (105.6 x 10^6 mL⁻¹ d⁻¹) by a
- 10 factor of 6.8, when the decline of the bloom was sampled. BS estimates ranged from 36 to 261
- 11 with mean values of 115 ± 74 in the bloom and 139 ± 77 in the HNLC waters.
- Although FIC values at the Fe-fertilized stations almost doubled that in HNLC waters, this difference between environments was not significant (Kruskal-Wallis test, Table 2). Average values for duplicate assays ranged from 4% to 47% (average: 22%) in fertilized waters and from 3% to 23% (average: 12%) in HNLC waters. Lysogenic infection of bacterioplankton could be detected only in 7 out of 15 lysogenic phage induction essays and ranged from 1-31% in fertilized waters and from 1-4% in the HNLC environment.
- 18 At the fertilized stations, on average $5.4 \pm 4.1 \times 10^5$ bacteria mL⁻¹ d⁻¹ were lysed, 5 times more
- 19 than at the HNLC stations $(1.1 \pm 0.6 \ 10^5 \text{ bacteria mL}^{-1} \text{ d}^{-1}, P < 0.05, \text{Kruskal-Wallis test}, \text{ Table 2}).$
- The resulting viral mediated loss of bacterial standing stock was on average 44 ± 24 % per day in the HNLC waters and more than twice as high at the fertilized stations although this was not significant (104 ± 76 % d⁻¹, Kruskal-Wallis test, Table 2). The fraction of bacterial mortality through viral lysis (VMM) following the model by Binder (1999) averaged 72 ± 72% in the bloom and 27 ± 19% at the HNLC sites (Kruskal-Wallis, *ns*, Table 2).

25 **3.4 Relation between the different parameters**

Spearman rank correlation coefficients ρ for chlorophyll a, viral and bacterial parameters from
HNLC and bloom stations are shown in Table 4. BA and BP correlated positively throughout

- 1 trophic situations but only in HNLC waters, BA and BP increased with chla. In the fertilized
- 2 waters VA correlated positively with BP while in HNLC waters VP increased with BA. Only in
- 3 these waters VP₁ correlated significantly and positively with the fraction of infected cells (Table

13

4 4).

1 4 Discussion

2 Viruses were the dominant mortality factor of bacteria during the late stage of a phytoplankton 3 bloom induced by natural iron fertilization in the Southern Ocean (second visit to A3) but 4 accounted for a small part of bacterial mortality within a new bloom (station B5, Table 3). Additionally, observations from the early bloom phase showed that heterotrophic nanoflagellates 5 6 (HNF) dominated the loss of BP, and viruses accounted for only 10% of bacterial mortality 7 (Christaki et al., 2014). These seasonal dynamics point to a switch from an efficient functioning 8 of the microbial food web during the onset of the phytoplankton bloom to a microbial food web 9 where organic carbon is mainly processed by the viral shunt. The increase in viral mediated 10 release of dissolved organic carbon over time has important consequences for the fate of part of 11 the photosynthetically fixed carbon and reduces its transfer to higher trophic levels and export.

12

13 4.1 Comparison of viral data with high latitude marine environments

14 Viral production rates in the present study match well the data obtained from the Australian 15 Sector of the Southern Ocean (Evans et al., 2009) and are within the range of VP rates from an iron induced bloom in the subarctic Pacific (Higgins et al., 2009). However, our VP rates are high 16 17 when compared to data from an artificial iron-fertilization experiment in the Southern Ocean 18 (Weinbauer et al., 2009a) or those from other high latitude marine environments, i.e. the Arctic 19 Sea (Steward et al., 1996;Boras et al., 2010) (Table 5). Differences between studies could be due 20 to spatio-temporal varations of VP, however, it is also conceivable that differences between 21 methods (Helton et al., 2005;Weinbauer et al., 2009a;Winget et al., 2005) have contributed to the 22 variability of reported VP data.

In the present study, the burst size averaged 128 throughout the experiments. This value is high compared to two studies from the Southern Ocean where measured BS was about 40 (Strzepek et al., 2005;Weinbauer et al., 2009a) and to a study in early spring above and off the Kerguelen plateau where BS evaluated with TEM observations varied from 6 to 88 (mean±SD, 22±15, Christaki et al., 2014). These different BS could be inherent to the study regions or due to the

used method, i.e. estimating BS by an increase in VA and a decrease of BA in the VRA (Wells
and Deming, 2006) which can result in increases of BP and thus potentially increase VP (Helton
et al., 2005;Weinbauer et al., 2009a;Winget et al., 2005). However, Steward et al. found BS as
high as 270 for areas of high productivity in the Chukchi Sea and studies from the North Sea
have reported 100 phages produced per lysed bacterium (Bratbak et al., 1992).

6

7 4.2 Viruses in HNLC waters versus a phytoplankton bloom induced by natural

8 iron fertilization

9 Viral distribution during the late stage of the phytoplankton bloom above the Kerguelen plateau 10 and its relation to the bacterial hosts (e.g. VBR) and phytoplankton biomass is extensively 11 reported, discussed and compared to existing data from similar regions in Brussard et al. (2008b). 12 During the late bloom average viral abundance at the bloom stations was twice as high as in 13 HNLC waters (Brussaard et al., 2008b), while during the early bloom viral abundance remained 14 unaffected (Christaki et al., 2014). Data from mesoscale Fe fertilization experiments showed that 15 viral abundance inside the fertilized patch were higher (Weinbauer et al., 2009a) or not 16 substantially different from outside (Higgins et al., 2009). The authors of the latter study 17 explained the lack of differences between inside and outside the fertilized patch with the time lag 18 of the microbial response to the induced bloom, since viral abundance and production were only 19 increasing at the end of their observations (day 12 after iron fertilization). This observation is in 20 line with the increase in viral abundance and activity on a seasonal scale in the Kerguelen bloom 21 (Christaki et al., 2014).

The present study observed a mature bloom and could thus track a period with a more pronounced microbial response. The 4 times higher viral production at the naturally Fe-fertilized study sites compares well to the 3 fold increase in phage production after an induced bloom through iron addition (Weinbauer et al., 2009a). Interestingly, Christaki et al. (2014) reported higher VP rates already at the early bloom stages. Thus, there is a trend of higher viral production in the iron-fertilized bloom compared to the sourrounding HNLC waters consistent with existing

data on iron fertilization (Weinbauer et al., 2009a). Complementary, within the bloom, HNF did 1 2 not seem to control enhanced bacterial production rates while in HNLC waters heterotrophic 3 HNF consumed 95% of bacterial production (Christaki et al., 2008). These studies suggest that 4 there is a switch towards viral lysis dominating in the bloom situations. More generally, this is in 5 accordance with previous studies across environments which showed viral influence to be more 6 important in more eutrophic waters (Weinbauer et al., 1993; Steward et al., 1996), particulary in 7 the cold environments such as the Arctic (Steward et al., 1996) or the Southern Ocean, where 8 Guixa-Boixereu et al. (2002) found that viruses were responsible for the entire bacterial 9 mortality. The high viral induced mortality in the bloom could also be a reason for low biomass 10 accumulation, despite the high BP. We calculated carbon release rates through viral lysis in two 11 ways; first, based on VP, and second, based on VMM related to FIC by a Model of Binder (1999) 12 (Table 6). Independent of the absolute values, which were one order of magnitude higher in the 13 former than in the latter way (Table 6), C-release through viral lysis was 5-8 times higher in the 14 Fe-fertilized than in the surrounding HNLC waters. 15 The percentages of lysogens (i.e. bacteria containing temperate viruses) were more variable in the 16 fertilized (0-31 %) than in the HNLC waters (0-4%) but not significantly different between 17 environments. Consistent with our study, Weinbauer et al. (2009a) did not find differences inside 18 and outside the iron-enriched patch during a fertilization experiment in the Southern Ocean. The 19 proportion of the lysogenized bacterial population can vary extensively, for example, from 1.5 to 20 11.4% in the Gulf of Mexico (Weinbauer and Suttle, 1996), from 4 to 38% in the Canadian 21 Arctic Shelf (Payet and Suttle, 2013) and from 0 to 100% in Tampa Bay, Florida (Williamson et 22 al., 2002). According to conceptual models, lysogeny should occur in environments, where the 23 contact rate between infective phage and hosts is too low to sustain the lytic life style (Paul et al., 24 2002), e.g. in deep sea where host abundance was low and lysogeny was highest in an across 25 system study (Weinbauer et al., 2003) or when system productivity is low (Payet and Suttle,

26 2013). Apparently, this was not the case in the present study as the fraction of lysogenic cells was

27 | not different between trophic situations. It was suggested that enhanced growth causes temperate

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- viruses to enter the lytic cycle (Wilson and Mann 1997). Both, filtration and incubation could 1 2 have stimulated bacterial production in the virus reduction approach (Weinbauer et al., 2009a) 3 and consequently induced prophages in the mitomycin C treatment controls. Additionally, it has to be stressed that mitomycin C used as an inducing agent of lysogens in the natural bacterial 4 communities may not induce all prophages and be toxic to some bacteria (Paul, 2008;Paul and 5 Weinbauer, 2010). Thus, the apparent low incidence of lysogenic infection, particularly in HNLC 6 7 waters might be an artifact. However, it could also be that the study period was not long enough 8 to induce potential changes of lysogenic infection. In addition, our study provides no evidence
- 9 10

4.3 Role of viruses for sustaining phytoplankton productivity by Fe-supply

that lysogens were induced by relieving iron addition.

12 Bacteria store about 50% of the biogenic iron in HNLC areas (Tortell et al., 1996) and the mode 13 of bacterial mortality will affect the way of Fe regeneration and bioavailability (Kirchman, 14 1996; Mioni et al., 2005; Strzepek et al., 2005). While viral lysis liberates organically complexed 15 iron, which may be assimilated rapidly, grazing mainly sets free inorganic Fe (Gobler et al., 16 1997; Poorvin et al., 2004). Assimilation studies with a model heterotrophic bacterium 17 demonstrated that Fe in the virus mediated cell lysates was more bioavailable than the 18 siderophores produced by the same cells supporting the importance of virus-mediated Fe 19 regeneration in marine surface waters (Poorvin et al., 2011). We calculated Fe release rates in 20 two ways; first, based on VP, and second, based on VMM related to FIC by a Model of Binder 21 (1999). The former resulted in average iron regeneration rates due to viral lysis of bacteria of 22 4.18 and 0.86 pMol Fe d⁻¹ in fertilized and HNLC waters, respectively, while the latter resulted in 23 more realistic values ranging from 0.03 pM d⁻¹ to 1.58 pM d⁻¹ (average: 0.42 \pm 0.49 pM d⁻¹) in iron-fertilized and from 0.004 pM d⁻¹ to 0.12 pM d⁻¹ (average: 0.05 \pm 0.05 pM d⁻¹) in HNLC 24 25 waters (Table 6). These values are similar to those found in the Southern Ocean (Evans and Brussaard, 2012) and an iron-induced bloom (ibid., Weinbauer et al., 2009a) but low compared to 26 other studies. Poorvin et al. (2004) reported Fe regeneration rates of 19.2-75.5 pM d⁻¹ in HNLC 27

waters off Peru, and Strzepek et al. (2005) found a high range over two orders of magnitude of 0.4-28 pM d⁻¹ in HNLC waters SE of New Zealand. Fe regeneration rates are calculated from viral induced bacterial loss, which is inversly related to burst size. When taking into account that the calculated burst size in the present study was five times higher than the assumed BS in the study of Poorvin et al. (2004), the values in the present study compare well to data on Fe regeneration through viral activity from artificial fertilization experiments and other environments.

8 Significantly more iron was released by viral lysis within the naturally Fe-fertilized bloom than at 9 the HNLC stations (P < 0.05, Kruskal-Wallis, Table 6). The concentration of dissolved iron in 10 the surface mixed layer on and off the Kerguelen plateau were typical for the open Southern 11 Ocean and averaged 90 ± 34 pM (Blain et al., 2007) and the estimated biogenic iron pool at the 12 main bloom station equaled 80 ± 9 pM (Sarthou et al., 2008). Taking into account the total Fe 13 demand of the producers within the bloom of 6.04 \pm 0.62 pM d⁻¹ (Sarthou et al., 2008), the 14 remobilization of iron through viral lysis above the Kerguelen plateau following the Model by 15 Binder (1999) acounts for up to 26% of the demand of the producers and this appears to be a non-16 negliglible iron source for sustaining plankton productivity.

17

18 **4.4** Implications for carbon cycling and sequestration

Enrichment of bacterial biomass and production in the naturally fertilized bloom in the present study ranged from 287 to 797 mg C m⁻² and from 23.5 mg C m⁻² d⁻¹ to 304 mg C m⁻² d⁻¹, respectively (Christaki et al., 2008). Bacterial abundance and production are often correlated with viral abundance and production. Thus, elevated bacterial activity in the (natural or induced) bloom could explain the enhanced viral abundance and production found in previous *in situ* Fe enrichment studies (Arrieta et al., 2000;Higgins et al., 2009;Weinbauer et al., 2009a).

25 The finding of higher viral lysis rates of bacteria in the sites of natural Fe fertilization, where

26 | HNF grazing could only explain a small fraction of bacterial mortality (Christaki et al., 2008) has

27 important implications for the carbon cycling. Due to enhanced viral lysis, less carbon will be

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transferred to larger members of the food web but becomes again part of the DOM pool (Middelboe et al., 1996). This viral shunt should result in elevated bacterial production and respiration, thus more CO₂ would be produced and less carbon sequestrated. Experimental studies indicate that most of the lysis products belong to the labile fraction of DOM and are consequently rapidly degraded (Weinbauer et al., 2011). By the transformation of bacterial biomass into DOM, viruses have the effect of retaining carbon and nutrients in the photic zone (Suttle, 2007). Thus, viral lysis of bacteria could short-circuit the biological pump (Brussaard et al., 2008a).

8 However, there are other possible scenarios. For example, microbial activity converts part of the 9 organic matter into recalcitrant DOM (RDOM) that is resistant to microbial utilization and can 10 persist in the interior of oceans for up to thousand of years. The detailed role of viral lysis in this 11 new concept of the microbial carbon pump (MCP) (Jiao et al., 2010) is still poorly known. 12 However, a compilation of data suggests that viral lysis increases the DOM pool and the ratio of 13 recalcitrant vs labile organic matter (Weinbauer et al., 2011). Thus, enhanced viral lysis of 14 bacteria due to Fe fertilization could result in an enhanced carbon sequestration not related to the 15 biological pump.

16 Rates of bacterial production ($[^{3}H]$ leucine incorporation) and respiration (< 0.8 μ m size-fraction) 17 were 5-6 times higher in the bloom at Station A3 than those in surrounding HNLC waters 18 indicating that heterotrophic bacteria within the bloom processed a significant portion of primary 19 production with most of it being rapidly respired (Obernosterer et al., 2008) fueling the CO₂ pool. 20 This scenario is coherent with the finding of small particulate organic carbon export fluxes to 21 depth necessary for long-term sequestration (de Baar et al., 2005; Street and Paytan, 2005), 22 despite the role of iron in regulating primary productivity. However, most in situ mesoscale iron 23 enrichment experiments so far performed in the HNLC regions did not last long enough to follow 24 the termination of the bloom (Buesseler and Boyd, 2003;Smetacek et al., 2012). In the present 25 study, we sampled a bloom in its late successional stage and could thereby track the fate of fixed 26 carbon by an iron-fertilized phytoplankton bloom. Figure 4 shows a simple sketch to highlight 27 the importance of each compartment of the microbial food web in the transfer of organic material

1 in an Fe-fertilized bloom compared to HNLC waters. Sequestration of material in viruses, 2 bacteria and dissolved matter may lead to stronger retention of nutrients in the euphotic zone in 3 systems with high viral lysis rates of bacteria, because more material remains in these small non-4 sinking forms. This could be of major importance for large-scale iron fertilization of ocean 5 regions as a means of enhancing the ability of the ocean to store anthropogenic CO_2 and mitigate 6 21^{st} century climate change.

7 5 Conclusions

8 Enhanced bacterial production following the iron-fertilized phytoplankton bloom induced a 9 switch from grazing to viral lysis as major mechanisms causing bacterial mortality. This could 10 change the carbon flow through the microbial food web. We suggest that enhanced viral lysis of 11 bacteria short-circuits the biological pump but potentially prime the microbial carbon pump.

12

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Table 1. Date, location, mixed layer depth (Zm) and physicochemical characteristics of all sampled stations.

1 2 3

Date	Station	latitude longi	tude	Water type	Zm (m)	Sampling depth (m)	Τ°C	Salinity	Chla $\mu g L^{-1}*$
1/19/05	A3-1	50°38′S 72°0)5'E	+ Fe	52	10	3.5	33.9	0.94
1/19/05	A3-1	50°38′S 72°0)5'E	+ Fe	52	50	3.3	33.9	1.72
1/19/05	A3-1	50°38'S 72°0)5'E	+ Fe	52	100	3	33.9	1.38
1/20/05	A11	49°09'S 74°0	0'E	+ Fe	44	10	3.8	33.9	0.41
1/20/05	A11	49°09′S 74°0	00'E	+ Fe	44	75	3.3	33.9	0.52
1/20/05	A11	49°09′S 74°0	00'E	+ Fe	44	200	1.6	34.1	0.21
1/26/05	C11-1	51°39′S 78°0	00'E	- Fe	73	10	1.9	33.8	0.19
1/26/05	C11-1	51°39′S 78°0	00'E	- Fe	73	80	1.6	33.8	0.29
1/26/05	C11-1	51°39′S 78°0	0'E	- Fe	73	200	1.3	34.2	0.01
1/29/05	B11	50°30′S 77°0	0'E	- Fe	59	10	2.2	33.8	0.11
1/29/05	B11	50°30′S 77°0	0'E	- Fe	59	120	0.5	33.8	0.24
1/29/05	B11	50°30'S 77°0	0'E	- Fe	59	200	0.2	34.1	0.03
2/1/05	B5	51°06'S 74°3	86'E	+ Fe	84	60	2.8	33.9	1.54
2/1/05	B5	51°06'S 74°3	86'E	+ Fe	84	100	2.6	33.9	1.39
2/2/05	B1	51°30′S 73°0	0'E	+ Fe	59	60	3.3	33.9	1.29
2/2/05	B1	51°30′S 73°0	0'E	+ Fe	59	100	2.7	33.9	1.04
2/4/05	A3-4	50°39′S 72°0)5'E	+ Fe	80	50	3.6	33.9	1.48
2/4/05	A3-4	50°39′S 72°0)5'E	+ Fe	80	150	1.7	33.9	1.54
2/6/05	C11-2	51°39′S 78°0	0'E	- Fe	20	60	1.6	33.8	0.26
2/6/05	C11-2	51°39′S 78°0	0'E	- Fe	20	100	0.6	33.9	0.20
2/9/05	C3	52°43′S 74°4	9′E	- Fe	42	60	2.5	33.9	0.19
2/9/05	C3	52°43′S 74°4	9'E	- Fe	42	100	19	33.9	0.17

4 5 6

T, temperature in ° Celsius; chl *a*, total chlorophyll *a*, +Fe, iron-fertilized, -Fe, HNLC waters * data are from Uitz et al. 2009

Table 2. Average ± SD values of viral and bacterial parameters from the iron-fertilized and

HNLC stations in the upper 200m water layer and results from one-way ANOVA for normally distributed data and Kruskal Wallis test for nonparametric data. Ranges are given in parenthesis. The average ratio between the two environments is shown and significant differences are

indicated.

Parameters	Fe-fertilized stations	HNLC stations	ratio
BA mL ⁻¹	$3.9 \pm 0.9 (1.9-5.3) \ge 10^5$	$2.4 \pm 0.7 (1.3-3.8) \ge 10^5$	1.7 ***
BP μ gC L ⁻¹ d ⁻¹	$1.1 \pm 0.7 (0.1-2.5)$	$0.3 \pm 0.2 (0.1-0.7)$	4.1 ***
VA mL ⁻¹	$9.9 \pm 3.6 (3.4-14.2) \ge 10^6$	$4.7 \pm 1.4 (3.1-7.4) \ge 10^6$	2.1 *
$VP_1 mL^{-1} d^{-1}$	59.0 ± 47.1 (9.9-117.9) x 10^6	$14.5 \pm 7.4 \ (6.0-25.6) \ x \ 10^6$	4.1 *
$VP_i mL^{-1} d^{-1}$;	$50.9 \pm 46.4 (2.8-125.5) \ge 10^6$	13.9 x 10 ⁶	3.7
FIC %	22 ± 17 (4-47)	$12 \pm 7 (3-23)$	1.8
FLC %††	$10 \pm 14 (1-31)$	$3 \pm 2 (1-4)$	4.0
Prophage replication rate mL ⁻¹ d ⁻¹ ††	$18.1 \pm 29.2 \ (0.6-61.5) \ge 10^3$	$1.0 \pm 1.2 (0.2-2.4) \ge 10^3$	18.5
R cell ⁻¹ d ⁻¹	29.4 ± 11.1 (10.3-43.0)	$14.2 \pm 4.4 (9.3-22.4)$	2.1 *
lysed bacteria mL-1 d-1	$5.4 \pm 4.1 \ (0.8-10.3) \ge 10^5$	$1.1 \pm 0.6 (0.4-2.1) \ge 10^5$	4.9 *
VMM %	72 ± 72 (8-202)	27 ± 19 (6-58)	2.6

VA, viral abundance, VP₁, lytic viral production, VP_i, induced viral production, FIC, fraction of infected cells, FLC, fraction of lysogenic cells, BA, bacterial abundance, BP, bacterial production, R, viral contacts per cell and day, VMM, virus-mediated bacterial mortality. * P < 0.05, ** P < 0.001, *** P < 0.0001 † detected in 6 out of 15 essays, only one in HNLC waters †† detected in 7 out of 15 essays

Table 3. In situ BP and viral para	neters from all	virus reduction	experiments.
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Station	Water type	Depth m	$\frac{BP}{\mu gC L^{-1} d^{-1}}$	$VP_1 = 10^6 mL^{-1} d^{-1}$	FIC %	FLC %	VMM %
A3-1	+ Fe	10	2.5	16.7	12	31	25
A3-1	+ Fe	50	1.9	15.6	6	6	12
A3-1	+ Fe	100	2.4	56.4	10	ND	19
A3-4	+ Fe	50	1.2	105.6	34	ND	106
A3-4	+ Fe	150	0.3	82.4	36	ND	115
B1	+ Fe	60	1.7	117.9	41	ND	147
B1	+ Fe	100	0.2	115.6	47	3	202
B5	+ Fe	100	1.1	11.2	4	1	8
A11	+ Fe	200	0.3	9.9	7	ND	14
B11	- Fe	10	0.2	16.3	14	ND	29
B11	- Fe	120	0.3	25.6	23	1	58
B11	- Fe	200	0.1	6.0	11	3	24
C3	- Fe	60	0.2	20.1	6	no exp	11
C3	- Fe	100	0.4	16.7	22	no exp	55
C11-1	- Fe	10	0.4	11.2	8	ND	17
C11-1	- Fe	80	0.7	9.6	9	4	17
C11-1	- Fe	200	0.1	7.5	6	ND	12
C11-2	- Fe	60	0.3	25.1	20	no exp	47
C11-2	- Fe	100	0.2	6.4	3	no exp	6

VPl, lytic viral production; FIC, fraction of infected cells; FLC, fraction of lysogenic cells; VMM, viral-mediated bacterial mortality; ND, Not detectable; no exp, no lysogen induction essay

Table 4. Nonparametric Spearman rank correlation matrix for chlorophyll *a*, bacterial and viral parameters from the fertilized (n=8-9, except for BP-BA-chla: 36-41) and HNLC stations (n=10, except for BP-BA-chla: 23-31). Bold numbers are significant r-values (* P < 0.05, ** P < 0.001, *** P < 0.0001).

	Chl a	BA	BP	VA	VP ₁
Fe-fertilized					
BA	0.209				
BP	0.243	0.633***			
VA	0.357	0.548	0.762*		
VP ₁	0.083	0.476	-0.050	0.310	
FIC	-0.183	0.333	-0.267	0.095	0.900**
HNLC					
BA	0.688*				
BP	0.380*	0.635*			
VA	0.164	0.576	0.515		
VP	0.426	0.746*	0.406	0.273	
FIC	0.168	0.304	0.310	-0.249	0.608

BA, bacterial abundance; BP, bacterial production; VA, viral abundance; VP₁, lytic viral production; FIC, fraction of infected cells

Location	Depth (m)	Method	VA (10 ⁹ L ⁻¹)	$VP(10^9L^{-1}d^{-1})$	VMM $(10^8 L^{-1} d^{-1})$	% BP	% SS d ⁻¹	Source
SO: Fe-fertilized	0-150	VRA	3.4-14.2 (9.9 ± 3.6)	9.9-117.9 (59.0 ± 47.1)	$0.8-10.3 (5.4 \pm 4.1)$	8-202 (72)	104	present study
SO: HNLC	0-200	VRA	3.1-7.4 (4.7 ± 1.4)	6.0-25.6 (14.5 ± 7.4)	$0.4-2.1 (1.1 \pm 0.6)$	6-58 (27)	44	present study
Antarctic	0-100	VDR	1-74 (13±10.4)			>100		Guixa-Boixereu et al., 2002
SO-Subantarctic	10	VRA	6.1-26	17.5-216.3	3.6-43.3	43-63	40-130	Evans et al., 2009
SO	5-200	VRA	0.5-7.6	0.4-16	0-8.7		0-72	Evans and Brussaard, 2012
SO: Fe patch	10-150	VRA	2.3-7(4.3±5.5)	0.9-3.6(1.9±0.5)		41-172 (104)*		Weinbauer et al., 2009
SO: HNLC	10-150	VRA	1.4-2.5(2.1±2)	0.3-0.8(0.6±0.1)		14–70 (39)*		Weinbauer et al., 2009
Arctic	0-10	TEM	2.5-36	0.2-4.6 (2)		2-36 (13)		Steward et al., 1996
North waters	0-200	TEM	1.36-5.55(3.3±1.6)	0.1-1.3			6-28	Middelboe et al., 2002
Arctic	0-230	VDA	1.4-4.5(2.8±1.3)	0.1-1.9	0.28-0.72			Wells and Deming, 2006
Subarctic Fe patch	0-10	TEM/VRA	40.5	30-200	90±25		7.4	Higgins et al., 2009
Subarctic outside	0-10	TEM/VRA	35.7	30-200	25.8±6.1		7.2	Higgins et al., 2009
Arctic	0-100	VRA	0.32-7.28	0.1-4.2		2-24 (9)	2-30	Boras et al., 2010
Canadian Arctic Shelf	2-56	VRA	2.7-27	0.03-7.7	0.02-4.3	31-156	1.4-29	Payet and Suttle, 2013

Table 5. Comparison of viral abundance (VA) and production (VP), virus-mediated bacterial mortality (VMM) and % loss of bacterial production (% BP) and standing stock per day (% SS d^{-1}) with literature data from other polar/subpolar environments.

SO: Southern Ocean, VDR: viral decay rates, TEM: Frequency of visibly infected cells by transmission electron microscopy, VDA: Virus dilution approach, VRA: virus reduction approach. *using BP in the VRA

Table 6. C and Fe release rates ($L^{-1} d^{-1}$) through viral lysis calculated from VP (12.4 fg C cell⁻¹, (Fukuda et al., 1998) and from FIC following the Model by Binder (1999) using bacterial iron quota of 7.5 μ Mol Fe mol C⁻¹ (Tortell et al., 1996). Averages are given in parenthesis.

	Release ba	used on VP	Release based on FIC		
	pmol Fe L ⁻¹ d ⁻¹	μ mol C L ⁻¹ d ⁻¹	pmol Fe L ⁻¹ d ⁻¹	μ mol C L ⁻¹ d ⁻¹	
Fertilized stations	$0.60-7.97 (4.18 \pm 3.15)$	$0.08 \text{-} 1.06 \ (0.56 \pm 0.42)$	$0.03-1.58 \ (0.42 \pm 0.49)$	$0.003 - 0.21 \ (0.06 \pm 0.07)$	
HNLC stations	$0.28 - 1.60 \ (0.86 \pm 0.43)$	$0.04-0.21 \ (0.11 \pm 0.06)$	$0.004 - 0.12 \ (0.05 \pm 0.05)$	$0.001 - 0.02 \ (0.01 \pm 0.01)$	
ratio	4.9*	4.9*	7.9*	7.9*	

*Values are significantly higher in the Fe-fertilized than in the HNLC stations (Kruskal-Wallis, P < 0.05)

Figure 1. Real-time satellite images of chlorophyll during the KEOPS cruise dating from the first sampling of station A3 (19/2/2005), (MODIS results provided by CSIRO marine research) and overlaid transects and sampled stations.

Figure 2. Depth profiles of bacterial (A) and viral abundance (B) in the Kerguelen study area. Full symbols indicate Fe-fertilized sites, open symbols indicate HNLC waters.

Figure 3. Lytic viral production from the Fe-fertilized (A) and HNLC (B) stations. Values are the averages of duplicates and error bar indicate the minimum and maximum values. When not visible, error bars are within the width of the line.

Figure 4. Simple sketch of the carbon and nutrient flow through the microbial food web in the Fe- fertilized (left) and HNLC waters (right). Arrow thickness represents the relative importance of factors controlling the size of each pool of the microbial food web.