

This discussion paper is/has been under review for the journal Biogeosciences (BG).  
Please refer to the corresponding final paper in BG if available.

# Vegetation heterogeneity and landscape position exert strong controls on soil CO<sub>2</sub> efflux in a moist, Appalachian watershed

J. W. Atkins<sup>1</sup>, H. E. Epstein<sup>1</sup>, and D. L. Welsch<sup>2</sup>

<sup>1</sup>University of Virginia, Department of Environmental Sciences, Charlottesville, Virginia, USA

<sup>2</sup>American Public University System, Charles Town, West Virginia, USA

Received: 6 November 2014 – Accepted: 20 November 2014 – Published: 18 December 2014

Correspondence to: J. W. Atkins (jeffatkins@virginia.edu)

Published by Copernicus Publications on behalf of the European Geosciences Union.

Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

**BGD**

11, 17631–17673, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



## Abstract

In topographically complex watersheds, landscape position and vegetation heterogeneity can alter the soil water regime through both lateral and vertical redistribution, respectively. These alterations of soil moisture may have significant impacts on the spatial heterogeneity of biogeochemical cycles throughout the watershed. To evaluate how landscape position and vegetation heterogeneity affect soil CO<sub>2</sub> efflux ( $F_{\text{SOIL}}$ ) we conducted observations across the Weimer Run watershed (373 ha), located near Davis, West Virginia, for three growing seasons with varying precipitation (2010 – 1042 mm; 2011 – 1739 mm; 2012 – 1244 mm; ~~precipitation data from BDKW2 station, Mesowest, University of Utah~~). An apparent soil temperature threshold of 11 °C at 12 cm depth on  $F_{\text{SOIL}}$  was observed in our data – where  $F_{\text{SOIL}}$  rates greatly increase in variance above this threshold. For analysis,  $F_{\text{SOIL}}$  values above this threshold were isolated and examined. Differences in  $F_{\text{SOIL}}$  among years were apparent by elevation ( $F_{4,633} = 3.17$ ;  $p = 0.013$ ) and by vegetation cover ( $F_{4,633} = 2.96$ ;  $p = 0.019$ ). For the Weimer Run watershed, vegetation exerts the major control on ~~soil CO<sub>2</sub> efflux ( $F_{\text{SOIL}}$ )~~, with the plots beneath shrubs at all elevations for all years showing the greatest mean rates of  $F_{\text{SOIL}}$  ( $6.07 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) compared to plots beneath closed-forest canopy ( $4.69 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) and plots located in open, forest gaps ( $4.09 \mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) plots. During periods of high soil moisture, we find that CO<sub>2</sub> efflux rates are constrained and that maximum efflux rates in this system occur during periods of average to below average soil water availability. ~~These findings offer valuable insight into the processes occurring within these topographically complex, temperate and humid systems, and the interactions of abiotic and biotic factors mediating biogeochemical cycles. With possible changing rainfall patterns as predicted by climate models, it is important to understand the couplings between water and carbon cycling at the watershed and landscape scales, and their potential dynamics under global change scenarios.~~

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 1 Introduction

Soil respiration ( $R_{\text{SOIL}}$ ) is a major component of the terrestrial carbon cycle (Raich and Potter, 1995; Schimel, 1995), and is 30–60 % greater than net primary productivity globally (Raich and Potter, 1995). Estimates of annual soil carbon emissions range from 68–100 Pg of carbon per year (Schlesinger, 1977; Raich and Schlesinger, 1992; Bond-Lamberty and Thomson, 2010). Temperate systems contribute approximately 20 % of the annual global  $R_{\text{SOIL}}$  (Bond-Lamberty and Thomson, 2010), but have been shown to be recent carbon sinks, averaging 0.72 Pg of C uptake per year from 1990–2007 (Pan et al., 2011).  $R_{\text{SOIL}}$  can be estimated in the field by measuring soil  $\text{CO}_2$  efflux ( $F_{\text{SOIL}}$ ) – the direct rate of  $\text{CO}_2$  crossing the soil surface over a period of time (Raich and Schlesinger, 1992).  $F_{\text{SOIL}}$  can vary spatially and temporally within and across systems as a result of the varied and complex interactions of controlling mechanisms (Drewitt et al., 2002; Trumbore, 2006; Vargas et al., 2010). The edaphic controls on  $F_{\text{SOIL}}$  at the landscape scale include soil temperature, soil moisture, root biomass, microbial biomass, soil chemistry, and soil physics (Fang et al., 1998; Davidson et al., 1998; Kang et al., 2000; Xu and Qi, 2001; Epron et al., 2004). These factors do not simply elicit additive or monotonic responses, but rather create complex responses of  $F_{\text{SOIL}}$  across spatial and temporal scales (Dilustro et al., 2005; Pacific et al., 2009).

Soil temperature is quite commonly a primary driver of  $F_{\text{SOIL}}$  (e.g. Fang and Moncrieff, 2001), and in complex terrain, temperature regimes can be mediated by elevation, slope, and aspect (Wu et al., 2013). The effects of elevation and topography on soil temperature can in turn affect carbon cycling (Schindlbacher et al., 2010) either directly or through indirect processes (Murphy et al., 1998). Soil water content (SWC) however often serves as an important secondary control on  $F_{\text{SOIL}}$ . At high SWC values,  $\text{CO}_2$  transport through the soil pore space is limited (Davidson and Trumbore, 1995; Jassal et al., 2005). Production of soil  $\text{CO}_2$  can also become limited at high SWC values due to anoxia and decreased microbial aerobic respiration (Oberbauer et al., 1992). At low SWC values,  $F_{\text{SOIL}}$  is decreased as well due to microbial desiccation

BGD

11, 17631–17673, 2014

### Soil $\text{CO}_2$ efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



and concomitantly reduced microbial activity (Van Gestel et al., 1993), resulting in decreased CO<sub>2</sub> production (Scanlon and Moore, 2000).

In topographically complex landscapes, precipitation gradients that exist as a function of elevation affect decomposition rates, CO<sub>2</sub> production, and movement of CO<sub>2</sub> through the soil (Schoor, 2001). The complex landscape structure and heterogeneity of mountain catchments also directly affect local soil moisture regimes through the lateral redistribution of soil water, adding to the spatial heterogeneity of these biogeochemical and physical processes.  $F_{SOIL}$  therefore varies across landscape positions as a function of this soil water redistribution (Riveros-Iregui and McGlynn, 2009). In subalpine forested systems for example, soil water content has been shown to be a strong driver of the spatial (Scott-Denton et al., 2003) and temporal (Pacific et al., 2008) variability of  $F_{SOIL}$ .

In addition to meteorological variables, vegetation (itself controlled by the spatial heterogeneity of micrometeorology), can influence carbon cycling within a watershed. Vegetation affects carbon cycling directly through photosynthesis (Raich and Schlesinger, 1992; Ekblad and Högberg, 2001; Högberg et al., 2001), above- and below-ground tissue allocation (Chen et al., 2013), and litter production (Chemidlin Prévost-Bouré et al., 2010). Vegetation therefore controls the quantity and quality of soil organic matter (SOM) within systems, which in part will determine decomposition rates and soil CO<sub>2</sub> production (e.g. Berg, 2000). ~~Aspects of foliar chemistry such as cellulose, lignin, and macronutrient (N, P, S) concentrations affect the rate of decomposition and the dynamics of SOM.~~ However, the role of belowground plant and microbial processes in the dynamics of SOM has become increasingly more apparent – showing that root and rhizosphere contributions to SOM are substantive (e.g. Schmidt et al., 2011). Vegetation also exerts controls on production of CO<sub>2</sub> through root respiration in the soil and through complex mycorrhizal associations that can mediate the response of soil CO<sub>2</sub> production to rain pulse events (Vargas et al., 2010). Finally, vegetation also elicits feedbacks on the abiotic aspects of a system, including the soil moisture and soil

**BGD**

11, 17631–17673, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



temperature regimes, further impacting biogeochemical cycling (Wullschleger et al., 2002; Metcalfe et al., 2011; Vesterdal et al., 2012).

Inter-annual variation in  $R_{\text{SOIL}}$  within systems can be high and exceed the inter-annual variation of net ecosystem exchange (NEE) of carbon (Savage and Davidson, 2001); this inter-annual variation can be driven in large part by the dynamics of precipitation (Raich et al., 2002). Current climate models project potentially dramatic changes in precipitation in the coming years (Kirtman et al., 2013), and presently the controls on inter-annual variation of  $R_{\text{SOIL}}$  in response to changing precipitation regimes are poorly understood at spatial scales ranging from landscapes to regions. The interactions among topography, vegetation cover, and climate are therefore an important and complicated area of study. ~~As an example, Noormets et al. (2010) found ecosystem respiration (ER) to be consistent across three years of altering precipitation regimes, with no signs of low volumetric water content (VWC) suppressing gross ecosystem production (GEP) or ER in a loblolly pine plantation in the North Carolina coastal plain.~~

Inter-annual climate variability in mountainous, subalpine catchments, however, has been shown to alter the spatio-temporal heterogeneity of carbon dynamics within those systems (Riveros-Iregui et al., 2011, 2012). In a subalpine watershed in Montana, Riveros-Iregui et al. (2012) found that areas with low upslope accumulated area (generally uplands and drier areas) showed  $F_{\text{SOIL}}$  increases during wet years, while poor-drainage areas (riparian areas) showed  $F_{\text{SOIL}}$  decreases during wet years. This resulting bidirectional response is a function of the landscape heterogeneity of the system, soil biophysics, and inter-annual climate variability (Riveros-Iregui et al., 2012). ~~Peak rates of  $F_{\text{SOIL}}$  during the growing season coincided with periods of intermediate soil moisture governed by the seasonal input of water from snowmelt in this system. As snowmelt increased, water moved from upland regions of the watershed to riparian areas lower in the watershed. During the initial stages of the growing season, rates of  $F_{\text{SOIL}}$  were suppressed due to high levels of SWC. As riparian areas began to dry down,  $F_{\text{SOIL}}$  rates increased to a maximum in early July before falling in phase with~~

Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



~~decreasing SWC. Similarly, Brito et al. (2013) found that, in a Mediterranean *Pinus canariensis* treeline forest (elevation 2070 m),  $F_{\text{SOIL}}$  was limited in the warm, dry period by low soil water potentials and also limited during the cold, wet period, by high soil water potentials.~~

5 Given the possible interactions among precipitation, topography, and vegetation, we examined how  $F_{\text{SOIL}}$  varies as a function of landscape position and vegetation cover in response to inter-annual variation in precipitation within a complex, humid watershed. To do this we used a plot-based approach with repeated measures sampling to account for spatial and temporal variation of the biophysical controls on  $F_{\text{SOIL}}$  within our study watershed. The empirical nature of this study design, coupled with the use of portable  
10 infra-red gas analyzers (IRGAs) to measure soil CO<sub>2</sub> efflux, is a robust and proven way of quantifying the seasonal dynamics of  $F_{\text{SOIL}}$  and allows for greater consideration of the spatial variability of  $F_{\text{SOIL}}$  (Riveros-Iregui et al., 2008; Riveros-Iregui and McGlynn, 2009) at the watershed scale. We attempted to answer the following questions:

- 15 1. How does  $F_{\text{SOIL}}$  respond to inter-annual variation of precipitation in a humid, complex watershed?
2. How do landscape position and vegetation heterogeneity affect  $F_{\text{SOIL}}$ , and how do they interact with inter-annual variation in precipitation?

## 2 Methods

### 20 2.1 Site description

The Weimer Run watershed (374 ha) is located in the Allegheny Mountain range in north-eastern West Virginia within the Little Canaan Wildlife Management Area near Davis, WV (39.1175, -79.4430) and is a sub-watershed of the Blackwater River, a tributary of the Cheat River. The watershed has an elevation range of 940 m  
25 (confluence of Weimer Run and the Blackwater River) to 1175 m (Bearden Knob)

BGD

11, 17631–17673, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(Fig. 1). For the climate period 1980–2010, mean annual precipitation (MAP) for the watershed was  $1450 \text{ mm yr}^{-1}$  (PRISM Climate Group). The mean daily maximum July temperature is  $18.8^{\circ}\text{C}$ , and the mean daily maximum January temperature is  $-3.9^{\circ}\text{C}$  (NCDC, Station ID DAVIS 3 SE, Davis WV). Precipitation varied during the study period, producing a relatively dry year in 2010 (1042 mm), a wet year in 2011 (1739 mm) and a mesic year in 2012 (1244 mm) (precipitation data from BDKW2 station, MesoWest, University of Utah) (Fig. 5a).

The Weimer Run watershed is adjacent to the Canaan Valley in West Virginia – which exists in a transitional zone between the Appalachian Valley and Ridge and the Appalachian Folded Plateau (Matchen, 1998). The surrounding ridge tops and the study site are underlain by Pennsylvanian sandstone from the Pottsville formation (Allard and Leonard, 1952). The over-story vegetation within the watershed is a mixed northern hardwood-coniferous forest, consisting of yellow birch (*Betula alleghaniensis*), red maple (*Acer rubrum*), red spruce (*Picea rubens*), and black cherry (*Prunus serotina*) (Allard and Leonard, 1952; Fortney, 1975). The under-story is comprised of *Rhododendron maximum*, *Kalmia latifolia*, *Osmundastrum cinnamomeum*, and *Osmunda claytoniana* (Fortney, 1975).

## 2.2 Vegetation and elevation treatments

Three elevation treatment levels were established along the north-eastern aspect of the watershed to form an elevation gradient: LOW (975 m), MID (1050 m), and HIGH (1100 m). Site elevations were determined using a digital elevation map (DEM) derived from  $1/9$  arc second elevation data from the Shuttle Radar Topography Mission (SRTM) (USGS 2006) processed with ArcGIS<sup>®</sup> software (ESRI; Redlands, CA). In order to address the effects of vegetation cover on  $F_{\text{SOIL}}$ , three vegetation cover classes were established: CANOPY – closed canopy, forest interior with no shrub layer; SHRUB – closed canopy, forest interior, with dense shrub layer; OPEN – forest gap with no canopy closure, within the forest interior. Differences among vegetation treatments were confirmed using plant area index (PAI) which was measured for each plot in June

2010 with a LAI-2000 Plant Canopy Analyzer (LI-COR Lincoln, Nebraska). PAI was strongly statistically significantly different among vegetation cover types ( $F = 13.39$ ;  $p$ -value = 0.0003). SHRUB plots were the greatest ( $3.46 \text{ m}^{-3} \text{ m}^3$ ) followed by CANOPY plots ( $2.14 \text{ m}^{-3} \text{ m}^3$ ) and then OPEN plots ( $1.75 \text{ m}^{-3} \text{ m}^3$ ) (Table A1).

At each elevation level in the watershed, three  $2 \text{ m} \times 2 \text{ m}$  plots of each vegetation class were established – for a total of 27 plots across the entire watershed (Fig. 1). One of the SHRUB replicate plots at the LOW elevation had to be removed from analysis due to inundation during the summer of 2011. Data from the remaining 26 plots were analyzed.

## 2.3 Environmental variables

### 2.3.1 Soil CO<sub>2</sub> efflux

An EGM-4 Portable Infrared Gas Analyzer (IRGA) with an attached SRC-1 Soil Chamber (PP Systems, Amesbury, MA) was used to measure soil CO<sub>2</sub> efflux rates. The EGM-4 has a measurement range of 0–2000 ppm ( $\mu\text{mol mol}^{-1}$ ) with an accuracy of better than 1% and linearity better than 1% throughout the range. The SRC-1 has a measurement range of 0–9.99  $\text{g CO}_2 \text{ m}^{-2} \text{ h}^{-1}$ . Plots were sampled approximately weekly (every 5–10 days) from the middle of May until the end of September, from 2010 to 2012. For March until mid-May, and during October and November, plots were measured approximately every two weeks (12–21 days) during times when they were snow-free.  $F_{\text{SOIL}}$  was measured 1–3 times at different locations within the plot at each measurement interval and averaged for a plot level estimation of  $F_{\text{SOIL}}$ . Plots were only sampled between 09:00 and 16:00 EST, and the sequence of plot measurements was varied to avoid a time-of-day bias in the results.

**BGD**

11, 17631–17673, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





### 2.3.2 Volumetric water content

Volumetric water content ( $\Theta_{\text{field}}$ ) was measured using a Campbell HydroSense CD 620 (Campbell Scientific) set to water content measure mode with 12 cm probes (Campbell Scientific;  $\pm 3.0\% \text{ m}^{-3} \text{ m}^3$ , with electrical conductivity  $< 2 \text{ dS m}^{-1}$ ; sampling volume using 12 cm rods  $\approx 650 \text{ cm}^3$ ). A minimum of three measurements was taken in each plot per sampling event and averaged to make a plot level estimation of  $\Theta_{\text{field}}$ .

Measurements taken by the Campbell HydroSense CD 620 have a known bias in soils where bulk density is outside of the  $1\text{--}1.7 \text{ g cm}^{-3}$  range, where organic matter is  $> 10\%$ , and where clay content is  $> 40\%$  (Campbell Scientific). In order to calibrate field measurements, a calibration procedure from Kelleners et al. (2009) was followed where  $P$ , the period, which is the square wave output from the probe in milliseconds, is converted to  $K_a$ , the relative soil permittivity (unitless).  $P$  is related to  $\Theta_{\text{field}}$  as shown in Eq. (1):

$$P = \left( -0.3385 \cdot \theta_{\text{field}}^2 \right) + (0.7971 \cdot \theta_{\text{field}}) + 0.7702 \quad (1)$$

Equation (2) converts  $P$  to  $K_a$ .

$$\sqrt{K_a} = \frac{(P - P_{\text{air}})}{(P_{\text{water}} - P_{\text{air}}) \cdot \left( \left( \sqrt{K_{\text{water}}} - 1 \right) + 1 \right)} \quad (2)$$

where  $P_{\text{air}}$  is the period in air, and  $P_{\text{water}}$  is the period in deionized water.  $P_{\text{air}}$  was calculated empirically at 0.79 ms.  $P_{\text{water}}$  was calculated at 1.37 ms following the procedure outlined in Kelleners et al. (2009) by placing the probes of the Campbell Hydrosense CD 620 in deionized water in an 18.92 L acid-washed container, with total vessel conductivity measured at  $0.47 \mu\text{mhos}$ .

Soil samples were taken in conjunction with HydroSense measurements in 2012 (depth = 12 cm, volume =  $56.414 \text{ cm}^3$ ,  $n = 37$ ), and actual VWC ( $\Theta_{\text{lab}}$ ) was calculated using Eq. (3) from Rose (2004), where  $w$  is the gravimetric water content of the soil

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



sample ( $\text{g}^{-3} \text{g}^3$ ),  $\rho_b$  is the soil bulk density ( $\text{gcm}^{-3}$ ), and  $\rho$  ( $\text{gcm}^{-3}$ ) is the density of water:

$$\theta_{\text{lab}} = \frac{w\rho_b}{\rho} \quad (3)$$

In order to calibrate field measurements of VWC ( $\Theta_{\text{field}}$ ),  $\sqrt{K_a}$  values were then regressed against  $\Theta_{\text{lab}}$  to create an Eq. (4) relating  $\sqrt{K_a}$  to  $\Theta$  ( $r^2 = 0.74$ ) such that field measurements of VWC ( $\Theta_{\text{field}}$ ) could be converted to  $\Theta$  in order to account for discrepancies in organic matter, soil bulk density, and clay content:

$$\theta = 7.0341 \cdot \left( \sqrt{K_a} \right) + 0.0806 \quad (4)$$

$\Theta$  was then converted to water-filled pore space (WFPS;  $\text{m}^{-3} \text{m}^3$ ) using the soil porosity ( $\Phi$ ;  $\text{m}^{-3} \text{m}^3$ ):

$$\text{WFPS} = \theta \cdot \Phi \quad (5)$$

WFPS provides a more mechanistic variable that takes into account the bulk density and porosity of the soil, which influence the transport and storage capacity of the soil with regard to soil  $\text{CO}_2$ .

### 2.3.3 Soil temperature

During each field sampling session, soil temperature ( $T_{\text{SOIL}}$ ;  $^{\circ}\text{C}$ ) was measured at 12 cm using a 12 cm REOTEMP Soil Thermometer (REOTEMP San Diego, CA) at a minimum of two locations within the plot. These measurements were averaged to create a plot mean temperature for each sampling event.

## BGD

11, 17631–17673, 2014

### Soil $\text{CO}_2$ efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### 2.3.4 Soils

Soil pH was determined using a 1 : 1 measure of soil (from 0–5 cm depth) with deionized water and measured with a Fieldscout Soilstik pH Meter (Spectrum Technologies, Inc. Plainfield, IL) with an accuracy of  $\pm 0.01$  pH,  $\pm 1^\circ\text{C}$ .

Soil samples were taken from 0–5, 0–12, and 0–20 cm profiles within the soil. Soil bulk density ( $\rho_s$ ), total bulk density ( $\rho_t$ ), soil particle density, and soil porosity ( $\Phi$ ) were also calculated for each sample (Grossman and Reinsch, 2002; Flint and Flint, 2002). Soil bulk density ( $\rho_s$ ) is defined as the bulk density of the soil fraction, where the soil fraction consists of soil that has been sieved to less than 2 mm and all gravel and root material have been removed. Total bulk density ( $\rho_t$ ) is defined as the absolute density of the sampled soil, including soil, roots, and gravel and is simply the sample dry mass over the sample volume. Total soil carbon and nitrogen were assessed using a NA 2500 Elemental Analyzer (CE instruments; Wigan, United Kingdom). Soil organic matter (SOM) content was estimated using loss-on-ignition at  $500^\circ\text{C}$  (Davies, 1974).

### 2.3.5 Data analysis

~~A soil temperature threshold of  $11^\circ\text{C}$  was used to parse the data, because it provided a more functional understanding than just delimiting data by growing or dormant season. Soil  $\text{CO}_2$  efflux was found to be tightly coupled to soil temperature at temperatures below  $11^\circ\text{C}$ , and highly variable at temperatures above  $11^\circ\text{C}$ . All analyses and means presented are for measurement periods where soil temperatures were  $\geq 11^\circ\text{C}$  unless otherwise noted.~~

We employed a mixed-model analysis of variance (ANOVA) with repeated measures to identify main and interactive effects of elevation and vegetation on soil  $\text{CO}_2$  efflux, soil temperature, and water-filled pore space using the proc mixed procedure in SAS 9.3 (SAS Institute, North Carolina USA). All means presented are least-squares means calculated using a Tukey-Kramer adjustment.

**BGD**

11, 17631–17673, 2014

## Soil $\text{CO}_2$ efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



To decouple the effects of soil temperature and soil moisture on  $F_{\text{SOIL}}$ , linear regressions of soil temperature against the natural-log of  $F_{\text{SOIL}}$  were done by year (2010, 2011, 2012), by vegetation cover type (OPEN, CANOPY, SHRUB), by elevation (LOW, MID, HIGH), by year and vegetation (OPEN 2010, CANOPY 2010, etc.), and by year and elevation (LOW 2010, MID 2010, etc.). The residuals from each model were then regressed against WFPS by each combination. All linear regressions use the lm function in R 3.0.1 (R Core Team 2013).

Differences in soil organic matter (SOM) among treatments were examined with a Kruskal–Wallis rank sum test using the `kruskal.test()` in R 3.0.1 (R Core Team, 2013). A two-way mixed-model ANOVA using the `proc mixed` procedure in SAS 9.3 was used to examine main and interactive effects of elevation, vegetation, and soil depth on soil bulk density and total bulk density.

### 3 Results

Exponential regression of  $F_{\text{SOIL}}$  measurements against soil temperature at 12 cm ( $T_{\text{SOIL}}$ ) (Fig. 2a) shows a positive relationship ( $r^2 = 0.316$ ;  $y = 0.829 + e^{(0.1149x)}$ ) with increases in temperature resulting in increased efflux rates. The amount of variance explained by  $T_{\text{SOIL}}$  lessens above 11 °C ( $r^2 = 0.104$ ), with  $F_{\text{SOIL}}$  measurements below 11 °C showing a much tighter relationship with temperature ( $r^2 = 0.434$ ). To explore this variance, all data above 11 °C were isolated and examined in order to parse out controls above this apparent temperature threshold for this system.

The natural log of flux measurements above 11 °C for all years were regressed against  $T_{\text{SOIL}}$  (Fig. 2b) showing a significant positive relationship with soil temperature ( $r^2 = 0.119$ ;  $y = 0.096x - 0.010$ ). From this linear model, the residuals were then regressed against WFPS. The residuals from the  $\ln(F_{\text{SOIL}})$  values above 11 °C show a significant negative relationships with WFPS (Fig. 2c) but this explains only marginally more of the variance ( $r^2 = 0.019$ ).

### 3.1 Soil CO<sub>2</sub> efflux ( $F_{\text{SOIL}}$ )

Repeated measures ANOVA analyses show no significant differences in  $F_{\text{SOIL}}$  among years when data are pooled. Significant differences among years do occur when data are parsed by elevation ( $F_{4,633} = 3.17$ ;  $p = 0.013$ ) and by vegetation ( $F_{4,633} = 2.96$ ;  $p = 0.019$ ).

Across all data above 11 °C, there was a significant effect of elevation ( $F_{2,633} = 3.44$ ;  $p = 0.032$ ), with plots at HIGH elevation sites showing the highest  $F_{\text{SOIL}}$  rates and HIGH sites statistically differing from LOW sites, with MID elevation sites not differing from either (Fig. 3a). 2010 was the only year to show a statistically significant difference in  $F_{\text{SOIL}}$  among elevation treatments within a year, with LOW elevation sites exhibiting significantly lower  $F_{\text{SOIL}}$  rates ( $F_{2,633} = 3.17$ ,  $p = 0.013$ ).

Differences among vegetation treatments were stark ( $F_{2,633} = 37.58$ ;  $p \leq 0.0001$ ). SHRUB treatments across all elevation treatments and all years had higher rates of  $F_{\text{SOIL}}$  ( $6.07 \pm 0.42 \mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$ ) than CANOPY ( $4.69 \pm 0.42 \mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$ ) or OPEN ( $4.09 \pm 0.42 \mu\text{molCO}_2 \text{m}^{-2} \text{s}^{-1}$ ) plots. This SHRUB effect was most notable during 2010, the driest year during the study, when SHRUB plots showed the highest rates of  $F_{\text{SOIL}}$  recorded during the study ( $7.48 \pm 0.674$ ). Statistical differences among vegetation treatments among years were complex. SHRUB 2010 and OPEN 2011 were uniquely different among all combinations (Fig. 3b).

### 3.2 Water-filled pore space (WFPS)

WFPS tracked well with precipitation across years, with 2010 having the lowest values of WFPS and 2011 having the highest rates of WFPS. WFPS in 2011 was significantly greater than either 2010 or 2012 ( $F_{2,633} = 17.27$ ;  $p = < 0.0001$ ) (Table 2). During 2010, when precipitation was lower than average, an apparent elevation effect on WFPS is observed, with HIGH elevation plots exhibiting significantly lower WFPS measurements than either LOW elevation or MID elevation plots (Fig. 3e). During 2011 and 2012, under extreme and moderate moisture regimes, this elevation effect is not evident.

During 2010, vegetation [treatment](#) types are not significantly different, but in 2011, when there is more moisture in the system, statistical differences among vegetation [treatments](#) are apparent, as SHRUB and CANOPY plots exhibit higher WFPS values than OPEN plots (Fig. 3f).

### 3.3 Soil temperature ( $T_{\text{SOIL}}$ )

Data for all years showed a significant effect of elevation on  $T_{\text{SOIL}}$  across elevation [treatments](#) for all data above  $11^{\circ}\text{C}$  ( $F_{2,633} = 170.76$ ;  $p \leq 0.0001$ ). LOW elevation sites were warmer ( $15.99 \pm 0.35^{\circ}\text{C}$ ), than MID sites ( $14.71 \pm 0.35^{\circ}\text{C}$ ) and HIGH ( $14.94 \pm 0.35^{\circ}\text{C}$ ) elevation sites. There was no statistical difference in soil temperature by elevation within years (Fig. 3c).

Vegetation (Fig. 3d) had a statistically significant effect on  $T_{\text{SOIL}}$  ( $F = 52.79$ ;  $p \leq 0.0001$ ). SHRUB plots were the coolest ( $14.93 \pm 0.35^{\circ}\text{C}$ ), OPEN plots the warmest ( $15.62 \pm 0.35^{\circ}\text{C}$ ), and CANOPY plots were in between ( $15.10 \pm 0.35^{\circ}\text{C}$ ). No within year comparisons were statistically significant. There were also no differences in temperature among years, when data were pooled and compared by year alone.

### 3.4 Soil physical and chemical characteristics

Soils within the Weimer Run watershed are heavily acidic, with pH ranging from 3.87–4.32 across the sampling area (Table A1). Soil bulk density ( $\rho_s$ ) from 0 to 12 cm ranges from  $0.49\text{--}1.11\text{ g cm}^{-3}$  (Fig. 4a and b), with lower values occurring beneath the shrub understory at lower elevations and higher values found in open, forest gap areas. There is an effect of elevation ( $F_{2,56} = 5.77$ ;  $p = 0.005$ ) and vegetation ( $F_{2,56} = 10.55$ ;  $p = 0.0001$ ) on  $\rho_s$  for all soil profiles (0–5, 0–12, and 0–20 cm). Elevation effects on  $\rho_s$  by soil depth are mixed, with statistical differences at 5 cm depth ( $F_{2,12} = 4.11$ ;  $p = 0.044$ ) and at 20 cm depth ( $F_{2,18} = 4.15$ ;  $p = 0.003$ ). By elevation [treatments](#), across all vegetation types,  $\rho_s$  from 0 to 12 cm is lowest at LOW elevations ( $0.65 \pm 0.08\text{ g cm}^{-3}$ ), highest at MID elevations ( $0.95 \pm 0.08\text{ g cm}^{-3}$ ), and in between at HIGH elevations

**BGD**

11, 17631–17673, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)

[◀](#)

[▶](#)

[◀](#)

[▶](#)

[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



( $0.73 \pm 0.08 \text{ g cm}^{-3}$ ). Vegetation shows significant differences at 12 cm ( $F_{2,18} = 3.60$ ;  $p = 0.048$ ) and 20 cm ( $F_{2,18} = 5.15$ ;  $p = 0.002$ ). By vegetation treatments, across all elevations,  $\rho_s$  from 0 to 12 cm is lowest in SHRUB plot ( $0.58 \pm 0.08 \text{ g cm}^{-3}$ ), highest in OPEN plots ( $0.92 \pm 0.08 \text{ g cm}^{-3}$ ), and in between at CANOPY plots ( $0.83 \pm 0.08 \text{ g cm}^{-3}$ ).

No interactive effects of elevation and vegetation were evident (Table A2).

Soil porosity from 0 to 12 cm ranges from  $0.58$  to  $0.82 \text{ m}^{-3} \text{ m}^3$  and is correlated with vegetation cover – with higher values beneath the SHRUB plots ( $0.77 \pm 0.03 \text{ m}^3 \text{ m}^{-3}$ ), medial values in CANOPY plots ( $0.68 \pm 0.03 \text{ m}^3 \text{ m}^{-3}$ ), and lower values in OPEN plots ( $0.65 \pm 0.03 \text{ m}^3 \text{ m}^{-3}$ ) (Table A5). SHRUB plots also show the highest concentrations of total soil carbon (9.35 %) significantly greater than other vegetation types ( $F = 9.79$ ;  $p = 0.0002$ ). Vegetation also influences total soil nitrogen, with SHRUB plots exhibiting higher proportions of total soil N than other plots (Table A5) ( $F = 6.36$ ;  $p = 0.0029$ ). Total soil carbon also differed by elevation, with LOW and HIGH treatments showing greater proportions of total soil carbon in samples than MID elevation sites (Table A4) ( $F = 6.28$ ;  $p = 0.0031$ ). MID level plots also showed lower proportions of total soil nitrogen than other elevation levels (Table A4) ( $F = 6.45$ ;  $p = 0.0027$ ).

Kruskal–Wallis tests show that soil organic matter (SOM) for all soil depths (0–5, 0–12, and 0–20 cm) varied significantly by vegetation ( $\chi^2 = 8.21$ ;  $p = 0.016$ ) and by soil depth ( $\chi^2 = 36.18$ ;  $p = < 0.0001$ ), but not by elevation ( $\chi^2 = 1.82$ ;  $p = 0.401$ ). Differences in SOM by vegetation treatment through the soil column were significant for the 0–5 and the 0–20 cm soil profiles (Table A4). The highest rates of SOM were found at the HIGH elevation plots (40.14 %) compared to the MID (21.73 %) and LOW elevation plots (33.03 %) (Fig. 4c). SHRUB plots (33.54 %) and CANOPY plots (33.14 %) had similar SOM values. OPEN plots were lower (27.76 %) (Fig. 4d).

## BGD

11, 17631–17673, 2014

### Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## 4 Discussion

For data above 11 °C across the three study years with varying precipitation regimes, it becomes apparent that there is an intrinsic link between the movement of carbon and water in this system in response to landscape heterogeneities and inter-annual climate dynamics. During 2010, our comparatively dry year, we see increased rates of  $F_{\text{SOIL}}$  across the watershed, but more pronounced increases in fluxes from SHRUB plots. While SHRUB plots do exhibit greater rates of fluxes than other treatments in this watershed during the course of this study, the magnitude of these fluxes is enhanced when the watershed is drier. Conversely, during 2011, the relatively wet year, vegetation-level differences in  $F_{\text{SOIL}}$  are statistically unapparent. When changing precipitation regimes are considered, along with future projections of warming and carbon dynamics, the importance of this coupling among water, carbon, and vegetation within humid watersheds cannot be understated. Changes in the distribution, variability, and amount of rainfall, as a result of climate change, are expected to have a major effect on carbon cycling (Borken et al., 2002). The magnitude of this effect, however, remains uncertain (Wu et al., 2011; Ahlström et al., 2012; Reichstein et al., 2013).

During 2010 (driest year), we see a strong effect of elevation on WFPS. During 2011 and 2012, however, there is no apparent effect of elevation on WFPS. When precipitation decreases across the watershed, as is the case during 2010, a different soil moisture regime manifests at higher elevations, with lower values of WFPS that contribute, in the case of this watershed, to increased rates of  $F_{\text{SOIL}}$ . During periods of increased precipitation, the watershed exhibits a more uniform soil moisture regime. The difference in the magnitude of carbon fluxes across elevation levels decreases during years with higher precipitation. During periods of higher precipitation and increased soil moisture, air space within the soil remains filled and transportation of  $\text{CO}_2$  through the soil is limited, resulting in decreased rates of  $F_{\text{SOIL}}$ . Production of  $\text{CO}_2$  in the soil is also decreased due to the increased incidence of anoxic conditions as a function of increased WFPS. Our LOW elevation plots were statistically similar

BGD

11, 17631–17673, 2014

### Soil $\text{CO}_2$ efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



in wetness to the MID plots, both of which were wetter than the HIGH plots during the study. The LOW elevation plots were also the warmest for each year of the study, yet exhibited the lowest rates of  $F_{\text{SOIL}}$  for the entire study period. One consideration not explicitly detailed in our study is the effect of topographic aspect on soil water redistribution as plots in our study all had an east-northeasterly aspect. Landscape positions with varying aspect can have differing soil water contents while having similar soil temperature regimes (Kang et al., 2003) that still result in varied soil carbon fluxes. Another contributor to the magnitude of carbon fluxes can be the amount of upslope accumulated area or the connectivity of varying landscape positions to flow paths within watersheds (McGlynn and Seibert, 2003; Pacific et al., 2012). During our wet year, however, we see a diminished effect of these topographic heterogeneities.

Enhanced fluxes during years of decreased precipitation suggest that soil respiration in humid mountain watersheds is strongly controlled by soil water, and to a lesser extent, soil temperature. During average and above-average precipitation years, soil respiration values are lower due to limited CO<sub>2</sub> production and/or diffusion through the soil. During years where precipitation is below average, soil respiration values increase. However, what is not considered here are the cumulative effects of inter-annual variability in precipitation. Would consecutive dry or consecutive wet years result in increases or decreases following the second year? Data from the National Climate Data Center's (NCDC) station in Canaan Valley, WV (Station ID 461393) show that precipitation in this region of WV is increasing, notably so since 1993 (Fig. 5b). This increase in precipitation appears to be driven by a notable increase in the number of extreme precipitation days (EPDs), defined here as days where precipitation exceeds 25.4 mm (Fig. 5c). While precipitation is generally increasing in the Weimer Run watershed, and similar areas across West Virginia, the year-to-year variance is increasing as well. A Breusch–Pagan test, which tests for the presence of heteroscedasticity in linear regression models, shows that NCDC precipitation data from Canaan Valley since 1970 exhibit a statistically significant increase in inter-annual variance (BP = 8.58;  $p = 0.003$ ). Meaning, the low precipitation years are trending

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



much lower than the mean, while the high precipitation years are trending much higher than the mean, with fewer overall “average” precipitation years. This increased variance appears to again be driven by the increased variance in EPDs from year-to-year (Fig. 5b and c) and has been attributed to changes in the North Atlantic Subtropical High and anthropogenic climate change (Li et al., 2011). As soils are subject to year-to-year wet/dry cycles, cumulative effects on carbon cycling and carbon fluxes are likely. It is beyond the scope of this study to answer the question posed above; however, with the observed dynamics in precipitation for the region, this may be an important line of future research. These relative extremes in rainfall amounts that occurred during this study resulted in significant differences in soil moisture regimes (measured as WFPS) across the entire watershed and among both our elevation and vegetation cover treatments (Sect. 3.2; Tables 1 and 2). During 2011, there were 34 EPDs, whereas in 2010 there were only 11 and in 2012 only 9. Precipitation also affected the variance of WFPS within the watersheds by year, as measured by the coefficient of variation (CV) with 2011 showing decreased variance of WFPS (CV = 27.85) compared to either 2010 (CV = 41.11) or 2012 (CV = 29.48). Increased precipitation and increased numbers of EPDs changes the soil moisture regime within the watershed that in turn affects CO<sub>2</sub> fluxes.

~~A noticeable shrub effect on CO<sub>2</sub> fluxes appears in our data. As noted, flux rates from SHRUB plots are typically greater than other vegetation cover types despite consistent lower soil temperatures. This effect is also exacerbated during drier years. We propose that increases in soil CO<sub>2</sub> efflux from beneath shrubs is related to the effect that shrubs have on the soils within this watershed. Soil bulk density, soil porosity, soil carbon, and other soil properties have been shown to drive the spatial variability of carbon fluxes (Jassal et al., 2004; Fiener et al., 2012; Luan et al., 2011). The effect of the soil microbial community on the temperature sensitivity of soil respiration can also be enhanced in soils with high soil C:N ratios (Karhu et al., 2014). Here we see shrubs decrease soil bulk density (Fig. 4b; Table A5) and increase soil porosity (soil porosity ( $\Phi$ ) for SHRUB plots averaged  $0.77 \text{ m}^3 \text{ m}^{-3}$  from 0 to 12 cm depth, compared~~

to  $0.65 \text{ m}^3 \text{ m}^{-3}$  for OPEN plots and  $0.68 \text{ m}^3 \text{ m}^{-3}$  for CANOPY plots; Table A5), allowing for greater diffusivity within the soil matrix, and increased transportation potential of soil  $\text{CO}_2$  through the soil. While soils under SHRUB plots have higher concentrations of SOM and soil C, soil bulk density is lower, which results in overall lower values of SOM and comparable values of soil C by volume. The increased soil porosity in soils beneath shrub cover is likely resulting in increased oxidation of labile soil C. It should be considered that SHRUB plots, to 20 cm soil depth, had the highest mean values of SOM (18.13%), higher soil C (9.35%), higher soil N (0.47%), higher C:N ratios (19.36), and lower  $\rho_s$  ( $0.39 \text{ g cm}^{-3}$ ) compared to CANOPY (SOM = 12.48%; soil C = 6.35%; soil N = 0.37%; soil C:N = 16.30) and OPEN plots (SOM = 12.48%; soil C = 5.14%; soil N = 0.31%; soil C:N = 15.76) (Table A4). The high C:N ratios for SHRUB plots indicate possibly lower amounts of available, labile carbon and possibly lower rates of decomposition than other areas of the watershed. This is corroborated by early results from a two-year litterbag experiment conducted in this watershed (Atkins et al., 2014). This indicates that root respiration contributions from shrubs may be substantive and may also be influenced by varying soil moisture and precipitation regimes.

The most dominant shrub species in this watershed is *Rhododendron maximum*, an ericaceous understory shrub that has been shown to increase SOM and soil N in forests where it is present (Boettcher and Kalisz, 1990; Wurzberger and Hendrick, 2007). *R. maximum* occurs most commonly in forest coves and on north-facing slopes with mesic to moist soil water regimes (Lipscomb and Abrams, 1990). Ericaceous litter also contributes to declines in soil fertility, lower N mineralization rates, and lower decomposition rates due to higher concentrations of foliar polyphenols (Hättenschwiler and Vitousek, 2000; DeLuca et al., 2002; Côté, 2000; Wurzberger and Hendrick, 2007). Ericaceous plants have ericoid mycorrhizae that provide a competitive advantage to breaking down organic N over ectomycorrhizae associated with many deciduous and coniferous species (Bending and Read, 1997) which leads to the inhibition of overstorey species regeneration (Nilsen et al., 2001). The areal extent of *R. maximum* has increased in some areas of southern and central Appalachia (Phillips and Murdy,

## Soil $\text{CO}_2$ efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



1985; Rollins et al., 2010; Brantley et al., 2013; Elliott et al., 2014). Shrub cover in the region is expected to continue to increase given fire suppression, lack of grazing, and forest canopy die-off from infestations (Nowacki and Abrams, 2008; Ford et al., 2011). If precipitation increases in this area in accordance with climate projections, the accompanying increase in soil moisture availability may further the expansion of *R. maximumum*. The loss of previously dominant foundational species in these systems (e.g. *Picea rubens* in West Virginia due to logging and fire in the late 1800s and early 1900s; *Tsuga canadensis* die-off from hemlock woolly adelgid across the Appalachians and eastern US) may result in possible, multiple stable-states (Ellison et al., 2005). Increase in shrub cover has the potential to further impact ecosystem fluxes and biogeochemical cycling and may contribute strongly to future forest community dynamics. However, conversely, if the variance of inter-annual precipitation continues to increase, drought years may serve as a possible control on shrub expansion. Future precipitation dynamics will be a strong determinant of the biogeochemical fluxes and cycling for this watershed, and possibly, for the region.

Our findings indicate that for this consistently wet, humid watershed, increases in precipitation may result in decreased soil water heterogeneity and decreased fluxes of carbon from the soil surface, while decreased precipitation may result in increased soil water heterogeneity and increased carbon fluxes – especially from areas with high shrub coverage. Similar studies in drier watersheds have found that increases in soil water availability largely result in increases in soil carbon fluxes. Pacific et al. (2008) showed that for the Stringer Creek watershed, a sub-alpine, montane watershed in Montana, the spatial variability of soil CO<sub>2</sub> efflux was controlled by the input of soil water driven by seasonal snowmelt. For riparian areas lower in the study watershed, soil carbon fluxes increased during the growing season in phase with increases in soil water content driven by seasonal snowmelt. Fluxes were suppressed at high levels of soil water early in the growing season, but as soil water decreased, fluxes increased. Pacific et al. (2009) further compared a wet and a dry year in the same watershed, finding that cumulative fluxes were 33% higher in riparian areas during the dry year,

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



but 8% lower at landscape positions higher in the watershed. Decreased moisture inputs resulted in significant responses in fluxes across landscape positions, but the riparian areas respond similarly to the entirety of the Weimer Run watershed, with dry years resulting in empirically validated increases in carbon fluxes. It has been shown in previous studies and research (Clark and Gilmour, 1983; Davidson et al., 2000; Sjoogersten et al., 2006; Pacific et al., 2008) that a production optimality of surface CO<sub>2</sub> efflux exists in response to soil water content such that peak rates of surface CO<sub>2</sub> efflux coincide with medial values of soil water content along a temporal transition within an elevation gradient. Our study adds the dimension of vegetation to this model, demonstrating that vegetation heterogeneity can have significant effects on surface CO<sub>2</sub> efflux within humid watersheds during periods of below-average soil water availability.

There are other possible avenues of carbon loss not considered here that may be affected by inter-annual climatic variability. It is possible that dissolved organic carbon (DOC) and dissolved inorganic carbon (DIC) fluxes from the watershed are increased during wet years due to increased flow in the system. Fluxes from these pools may be significant, but are difficult to measure and often carry a high-degree of uncertainty. DIC and DOC fluxes are highly variable spatially, coinciding with preferential flow paths within watersheds as a function of run-off (McGlynn and McDonnell, 2003; Kindler et al., 2011). Manipulative experiments have shown that simulated drought decreases DOC leaching across an elevation gradient by as much as 80–100% (Hagedorn and Joos, 2014), indicating that these fluxes are also responsive to inter-annual climate variability.

## 5 Conclusions

We completed a three-year plot-based study focusing on evaluating the effects of vegetation cover and elevation on soil carbon cycling in response to inter-annual variability in precipitation. By looking at data above 11 °C for soil temperature measured

at 12 cm depth, we were able to focus on the effects of soil moisture on carbon cycling without having to control for temperature limitation. We found that during a relatively dry year (2010; 1042 mm) the magnitude of soil carbon flux was enhanced across the watershed, but the increase was differential due to statistically greater fluxes from plots with high shrub coverage. Greater fluxes of carbon from plots with high shrub cover were due in part to decreased soil bulk density, high quantities of soil organic matter, and possible increased root respiration present beneath shrubs as compared to either closed-canopy or open-area plots. For 2011 and 2012, relatively wetter years, fluxes were decreased, and the effects of vegetation cover on the magnitude and variability of fluxes were statistically insignificant. Elevation had an effect on carbon cycling in the system by exacerbating vegetation effects during dry periods through increased effects on soil water distribution in the system. While soil water was correlated with elevation for all of our data, the effect was more pronounced during our driest year (2010) where areas higher in the watershed were much drier than lower positions. With the expected increase in precipitation as forecast by climate models and the empirical basis of increased inter-annual variance in precipitation, these findings offer important insights on the relations among landscape, vegetation, soil, and the associated biogeochemical effects for complex, humid watersheds. Given the increased likelihood of greater inter-annual variance in precipitation in the future, the coupling between carbon movement and vegetation cover is potentially quite crucial and under-considered. Further, the role of ericaceous shrubs and their future in this system are quite complex and may have profound influence on biogeochemical cycles.

## Data

Data used in this paper are available on figshare:

<http://dx.doi.org/10.6084/m9.figshare.1251229>

<http://dx.doi.org/10.6084/m9.figshare.1251228>

<http://dx.doi.org/10.6084/m9.figshare.1251201>

17652

## BGD

11, 17631–17673, 2014

### Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



*Author contributions.* H. Epstein and D. Welsch designed the experiment. J. Atkins conducted the field work and analysis. J. Atkins prepared the manuscript with input from all co-authors.

*Acknowledgements.* The authors would like to acknowledge the Canaan Valley Institute for significant financial and logistical support of this project. We would also like to acknowledge the Appalachian Stewardship Fund and the University of Virginia Department of Environmental Sciences for additional funding. Further thanks to Josh Richards, Sang Mee Ko, Jerrica Frazier, Stesha Dunker, Thomas Williams, Grace Wilkinson, Sara Taube, Trevor Klein, Virginia Mathurin, Margot Miller, and Jonathan Walter for their contributions in the field and laboratory.

## References

- Ahlström, A., Schurgers, G., Arneeth, A., and Smith, B.: Robustness and uncertainty in terrestrial ecosystem carbon response to CMIP5 climate change projections, *Environ. Res. Lett.*, 7, 044008, doi:10.1088/1748-9326/7/4/044008, 2012.
- Allard, H. A. and E. C. L.: The Canaan and Stony River Valleys of West Virginia, their former magnificent spruce forests, *Castanea*, 17, 1–60, 1952.
- Atkins, J., Epstein, H. E., and Welsch, D. L.: Leaf-litter decomposition differs by vegetation cover along an elevation gradient in a West Virginia watershed, in preparation, 2014.
- Bending, G. D. and Read, D. J.: Lignin and soluble phenolic degradation by ectomycorrhizal and ericoid mycorrhizal fungi, *Mycol. Res.*, 101, 1348–1354, 1997.
- Berg, B.: Litter decomposition and organic matter turnover in northern forest soils, *Forest Ecol. Manag.*, 133, 13–22, doi:10.1016/S0378-1127(99)00294-7, 2000.
- Boettcher, S. E. and Kalisz, P. J.: Single-tree influence on soil properties in the mountains of eastern Kentucky, *Ecology*, 1365–1372, 1990.
- Bond-Lamberty, B. and Thomson, A.: Temperature-associated increases in the global soil respiration record, *Nature*, 464, 579–582, 2010.
- Borken, W., Xu, Y. J., Davidson, E. A., and Beese, F.: Site and temporal variation of soil respiration in European beech, Norway spruce, and Scots pine forests, *Glob. Change Biol.*, 8, 1205–1216, 2002.
- Brantley, S., Ford, C. R., and Vose, J. M.: Future species composition will affect forest water use after loss of eastern hemlock from southern Appalachian forests, *Ecol. Appl.*, 23, 777–790, 2013.

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Brito, P., Trujillo, J. L., Morales, D., Jimenez, M. S., and Wieser, G.: Soil moisture overshadows temperature control over soil CO<sub>2</sub> efflux in a *Pinus canariensis* forest at treeline in Tenerife, Canary Islands, *Acta Oecol.*, 48, 1–6, 2013.

Clark, M. D. and Gilmour, J. T.: The effect of temperature on decomposition at optimum and saturated soil water contents, *Soil Sci. Soc. Am. J.*, 47, 927–929, 1985.

Chemidlin Prévost-Bouré, N., Soudani, K., Damesin, C., Berveiller, D., Lata, J. C., and Dufrière, E.: Increase in aboveground fresh litter quantity over-stimulates soil respiration in a temperate deciduous forest, *Appl. Soil Ecol.*, 46, 26–34, 2010.

Côté, L., Brown, S., Paré, D., Fyles, J., and Bauhus, J.: Dynamics of carbon and nitrogen mineralization in relation to stand type, stand age and soil texture in the boreal mixedwood, *Soil Biol. Biochem.*, 32, 1079–1090, 2000.

Davidson, E. A. and Trumbore, S. E.: Gas diffusivity and production of CO<sub>2</sub> in deep soils of the eastern Amazon, *Tellus B*, 47, 550–565, 1995

Davidson, E. A., Belk, E., and Boone, R. D.: Soil water content and temperature as independent or confounded factors controlling soil respiration in a temperate mixed hardwood forest, *Glob. Change Biol.*, 4, 217–227, 1998.

Davidson, E. A., Verchot, L. V., Cattânio, J. H., Ackerman, I. L., and Carvalho, J. E. M.: Effects of soil water content on soil respiration in forests and cattle pastures of eastern Amazonia, *Biogeochemistry*, 48, 53–69, 2000.

DeLuca, T., Nilsson, M. C., and Zackrisson, O.: Nitrogen mineralization and phenol accumulation along a fire chronosequence in northern Sweden, *Oecologia*, 133, 206–214, 2002

Dilustro, J. J., Collins, B., Duncan, L., and Crawford, C.: Moisture and soil texture effects on soil CO<sub>2</sub> efflux components in southeastern mixed pine forests, *Forest Ecol. Manag.*, 204, 87–97, 2005.

Drewitt, G. B., Black, T. A., Nestic, Z., Humphreys, E. R., Jork, E. M., Swanson, R., Ethier, G. J., Griffis, T., and Morgenstern, K.: Measuring forest floor CO<sub>2</sub> fluxes in a Douglas-fir forest, *Agr. Forest. Meteorol.*, 110, 299–317, 2002.

Eklblad, A. and Högborg, P.: Natural abundance of <sup>13</sup>C in CO<sub>2</sub> respired from forest soils reveals speed of link between tree photosynthesis and root respiration, *Oecologia*, 127, 305–308, 2001.

Ellison, A. M., Bank, M. S., Clinton, B. D., Colburn, E. A., Elliott, K., Ford, C. R., Foster, D. R., Kloeppel, B. D., Knoepp, J. D., Lovett, G. M., Mohan, J., Orwig, D. A., Rodenhouse, N. L.,



## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



Sobczak, W. V., Stinson, K. A., Stone, J. K., Swan, C. M., Thompson, J., Von Holle, B., and Webster, J. R.: Loss of foundation species: consequences for the structure and dynamics of forested ecosystems, *Front. Ecol. Environ.*, 3, 479–486, 2005.

Elliott, K. J., Vose, J. M., and Rankin, D.: Herbaceous species composition and richness of mesophytic cove forests in the southern Appalachians: synthesis and knowledge gaps, *J. Torrey Bot. Soc.*, 141, 39–71, 2014.

Epron, D., Nouvellon, Y., Rouspard, O., Mouvondy, W., Mabilia, A., Saint-André, L., Joffre, R., Jourdan, C., Bonnefond, J., Bergibier, P., and Hamel, O.: Spatial and temporal variations of soil respiration in a *Eucalyptus* plantation in Congo, *Forest Ecol. Manag.*, 202, 149–160, 2004.

Fang, C. and Moncrieff, J. B.: The dependence of soil CO<sub>2</sub> efflux on temperature, *Soil Biol. Biochem.*, 33, 155–165, 2001.

Fang, C., Moncrieff, J. B., Gholz, H. L., and Clark, K. L.: Soil CO<sub>2</sub> efflux and its spatial variation in a Florida slash pine plantation, *Plant Soil*, 205, 135–146, 1998.

Fiener, P., Dlugob, V., Korres, W., and Schneider, K.: Spatial variability of soil respiration in a small agricultural watershed – are patterns of soil redistribution important?, *Catena*, 94, 3–16, 2001.

Flint, A. L. and Flint, L. E.: Particle density, Chapter 2.2, in: *Methods of Soil Analysis: Physical Methods, Part 4*, Soil Society of America, 229–240, 2002.

Ford, C. R., Elliott, K. J., Clinton, B. D., Kloppel, B. D., and Vose, J. M.: Forest dynamics following eastern hemlock mortality in the southern Appalachians, *Oikos*, 121, 523–536, 2012.

Fortney, R. H.: The vegetation of Canaan Valley, West Virginia: a taxonomic and ecological study, *Doctoral dissertation*, West Virginia University, 1975.

Fortney, R. H. and Rentch, J. S.: Post logging era plant successional trends and geospatial vegetation patterns in Canaan Valley, West Virginia, 1945 to 2000, *Castanea*, 68, 317–334, 2003.

Grossman, R. B., Reinsch, T. G., Dane, J. H., and Topp, C.: Bulk Density and Linear Extensibility, Chapter 2.1, in: *Methods of Soil Analysis: Physical Methods, Part 4*, Soil Society of America, 201–225, 2002.

Hagedorn, F. and Joos, O.: Experimental summer drought reduces soil CO<sub>2</sub> effluxes and DOC leaching in Swiss grassland soils along an elevational gradient, *Biogeochemistry*, 117, 395–412, 2014

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



- Hättenschwiler, S. and Vitousek, P. M.: The role of polyphenols in terrestrial ecosystem nutrient cycling, *Trends Ecol. Evol.*, 15, 238–243, 2000.
- Högberg, P., Nordgren, A., Buchmann, N., Taylor, A. F., Ekblad, A., Högberg, M. N., Nyberg, G., Ottosson-Löfvenius, M., and Read, D. J.: Large-scale forest girdling shows that current photosynthesis drives soil respiration, *Nature*, 411, 6839, 789–792, 2001.
- Kirtman, B., Power, S. B., Adedoyin, J. A., Boer, G. J., Bojariu, R., Camilloni, I., Doblas-Reyes, F. J., Fiore, A. M., Kimoto, M., Meehl, G. A., Prather, M., Sarr, A., Schär, C., Sutton, R., van Oldenborgh, G. J., Vecchi, G., and Wang, H. J.: Near-term climate change: projections and predictability, in: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Boschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M., Cambridge University Press, Cambridge, UK and New York, NY, USA, 953–1028, 2013.
- Jassal, R. S., Black, T. A., Drewitt, G. B., Novak, M. D., Gaumont-Guay, D., and Nescic, Z.: A model of the production and transport of CO<sub>2</sub> in soil: predicting soil CO<sub>2</sub> concentrations and CO<sub>2</sub> efflux from a forest floor, *Agr. Forest Meteorol.*, 124, 219–236, 2004.
- Jassal, R., Black, A., Novak, M., Morgenstern, K., Nescic, Z., and Gaumont-Guay, D.: Relationship between soil CO<sub>2</sub> concentrations and forest-floor CO<sub>2</sub> effluxes, *Agr. Forest Meteorol.*, 130, 176–192, 2005.
- Jongen, M., Pereira, J. S., Aires, L. M. I., and Pio, C. A.: The effects of drought and timing of precipitation on the inter-annual variation in ecosystem-atmosphere exchange in a Mediterranean grassland, *Agr. Forest Meteorol.*, 151, 595–606, 2011.
- Kang, S., Kim, S., Oh, S., and Lee, D.: Predicting spatial and temporal patterns of soil temperature based on topography, surface covered and air temperature, *Forest Ecol. Manag.*, 136, 173–184, 2000.
- Kang, S., Doh, S., Lee, D., Lee, D., Jin, V. L., and Kimball, J. S.: Topographic and climatic controls on soil respiration in six temperate mixed-hardwood forest slopes, Korea, *Glob. Change Biol.*, 9, 1427–1437, 2003.
- Karhu, K., Auffret, M. D., Dungait, J. A., Hopkins, D. W., Prosser, J. I., Singh, B. K., Subke, J.-A., Wookey, P. A., Ågren, G. I., Sebastià, M.-T., Gouriveau, F., Bergkvist, G., Meir, P., Nottingham, A. T., Salinas, N., and Hartley, I. P.: Temperature sensitivity of soil respiration rates enhanced by microbial community response, *Nature*, 513, 81–84, 2014.

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Kelleners, T. J., Paige, G. B., and Gray, S. T.: Measurement of the dielectric properties of Wyoming soils using electromagnetic sensors, *Soil Sci. Soc. Am. J.*, 73, 1626–1637, 2009.
- Kindler, R., Siemens, J. A. N., Kaiser, K., Walmsley, D. C., Bernhofer, C., Buchmann, N., Cellier, P., Eugster, W., Gleixner, G., Grünwald, T., Heim, A., Ibrom, A., Jones, S. K., Jones, M., Klumpp, K., Kutsch, W., Larsen, K. S., Lehuger, S., Loubet, B., McKenzie, R., Moors, E., Osborne, B., Pilegaard, K., Reibmann, C., Saunders, M., Schmidt, M. W. I., Schrupf, M., Seyfferth, J., Skiba, U., Soussana, J.-F., Sutton, M. A., Tefs, C., Vowincke, B., Zeeman, M. J., and Kaupenjohann, M.: Dissolved carbon leaching from soil is a crucial component of the net ecosystem carbon balance, *Glob. Change Biol.*, 17, 1167–1185, 2011.
- Li, W., Li, L., Fu, R., Deng, Y., and Wang, H.: Changes to the North Atlantic subtropical high and its role in the intensification of summer rainfall variability in the southeastern United States, *J. Climate*, 24, 1499–1506, 2011.
- Luan, J., Liu, S., Wang, J., Zhu, X., and Shi, Z.: Rhizospheric and heterotrophic respiration of a warm-temperate oak chronosequence in China, *Soil Biol. Biochem.*, 43, 503–512, 2011.
- Matchen, D. L.: Geology of the Canaan Valley region, Geological Society of America, Southeastern Section, 47th annual meeting Charleston, West Virginia, 50, 1998.
- McGlynn, B. L. and McDonnell, J. J.: Role of discrete landscape units in controlling catchment dissolved organic carbon dynamics, *Water Resour. Res.*, 39, 1090, doi:10.1029/2002WR001525, 2003.
- McGlynn, B. L. and Seibert, J.: Distributed assessment of contributing area and riparian buffering along stream networks, *Water Resour. Res.*, 39, 1082, doi:10.1029/2002WR001521, 2003.
- Metcalfe, D. B., Fisher, R. A., and Wardle, D. A.: Plant communities as drivers of soil respiration: pathways, mechanisms, and significance for global change, *Biogeosciences*, 8, 2047–2061, doi:10.5194/bg-8-2047-2011, 2011.
- Murphy, K. L., Klopatek, J. M., and Klopatek, C. C.: The effects of litter quality and climate on decomposition along an elevational gradient, *Ecol. Appl.*, 8, 1061–1071, 1998.
- Nilsen, E. T., Clinton, B. D., Lei, T. T., Miller, O. K., Semones, S. W., and Walker, J. F.: Does *Rhododendron maximum* L. (Ericaceae) reduce the availability of resources above and belowground for canopy tree seedlings?, *Am. Midl. Nat.*, 145, 325–343, 2001.
- Noormets, A., Gavaz, M. J., McNulty, S. G., Domec, J., Sun, G., King, J. S., and Chen, J.: Response of carbon fluxes to drought in a coastal plain loblolly pine forest, *Glob. Change Biol.*, 16, 272–287, doi:10.1111/j.1365-2486.2009.01928.x, 2010.

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Nowacki, G. J. and Abrams, M. D.: The demise of fire and “Mesophication” of forests in the eastern United States, *Bioscience*, 58, 123–138, 2008.
- Oberbauer, S. F., Gillespie, C. T., Cheng, W., Gebauer, R., Serra, A. S., and Tenhunen, J. D.: Environmental effects on CO<sub>2</sub> efflux from riparian tundra in the northern foothills of the Brooks Range, Alaska, USA, *Oecologia*, 92, 568–577, 1992.
- Pacific, V. J., McGlynn, B. L., Riveros-Iregui, D., Welsch, D. L., and Epstein, H. E.: Variability in soil respiration across riparian-hillslope transitions, *Biogeochemistry*, 91, 51–70, doi:10.1007/s10533-008-9258-8, 2008.
- Pacific, V. J., McGlynn, B. L., Riveros-Iregui, D. A., Epstein, H. E., and Welsch, D. L.: Differential soil respiration responses to changing hydrologic regimes, *Water Resour. Res.*, 45, W07201, doi:10.1029/2009WR007721, 2009.
- Riveros-Iregui, D. A., McGlynn, B. L., Emanuel, R. E., and Epstein, H. E.: Complex terrain leads to bidirectional responses of soil respiration to inter-annual water availability, *Glob. Change Biol.*, 18, 749–756, 2012.
- Pan, Y., Birdsey, R. A., Fang, J., Houghton, R., Kauppi, P. E., Kurz, W. A., Phillips, O. L., Shvidenko, A., Lewis, S. L., Canadell, J. G., Ciais, P., Jackson, R. B., Pacala, S. W., McGuire, A. D., Piao, S., Rautianinen, A., Sitch, S., and Hayes, D.: A large and persistent carbon sink in the world’s forests, *Science*, 333, 988–993, 2011.
- Phillips, D. L. and Murdy, W. H.: Notes: Effects of *Rhododendron* (*Rhododendron maximum* L.) on regeneration of Southern Appalachian Hardwoods, *Forest Sci.*, 31, 226–233, 1985.
- PRISM Climate Group, Oregon State University, available at: <http://prism.oregonstate.edu>, last access: 23 March 2013, 2004.
- R Core Team: A language and environment for statistical computing, R Foundation for Statistical Computing, Vienna, Austria, available at: <http://www.R-project.org/>, 2013.
- Raich, J. W. and Potter, C. S.: Global patterns of carbon dioxide emissions from soils, *Global Biogeochem. Cy.*, 9, 23–36, 1995.
- Raich, J. W. and Schlesinger, W. H.: The global carbon dioxide flux in soil respiration and its relationship to vegetation and climate, *Tellus B*, 44, 81–99, 1992.
- Raich, J. W., Potter, C. S., and Bhagawati, D.: Interannual variability in global soil respiration, 1980–1994, *Glob. Change Biol.* 8, 800–812, 2002
- Reichstein, M., Bahn, M., Ciais, P., Frank, D., Mahecha, M. D., Seneviratne, S. I., Zscheischler, J., Beer, C., Buchmann, N., Frank, D. C., Pape, D., Rammig, A., Smith, P.,

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

[Printer-friendly Version](#)

[Interactive Discussion](#)



Thonicke, K., van der Velde, M., Vicca, S., Walz, A., and Wattenbach, M.: Climate extremes and the carbon cycle, *Nature*, 500, 287–295, 2013.

Riveros-Iregui, D. A. and McGlynn, B. L.: Landscape structure control on soil CO<sub>2</sub> efflux variability in complex terrain: Scaling from point observations to watershed scale fluxes, *J. Geophys. Res.*, 114, G02010, doi:10.1029/2008JG000885, 2009.

Riveros-Iregui, D. A., McGlynn, B. L., Epstein, H. E., and Welsch, D. L.: Interpretation and evaluation of combined measurement techniques for soil CO<sub>2</sub> efflux: discrete surface and continuous soil CO<sub>2</sub> concentration probes, *J. Geophys. Res.*, 113, G04027, doi:10.1029/2008JG000811, 2008.

Riveros-Iregui, D., McGlynn, B. L., Marshall, L., Welsch, D. L., Emanuel, R. E., and Epstein, H. E.: A watershed-scale assessment of a process soil CO<sub>2</sub> production and efflux model, *Water Resour. Res.*, 47, W00J04, doi:10.1029/2010WR009941, 2011.

Riveros-Iregui, D. A., McGlynn, B. L., Emanuel, R. E., and Epstein, H. E.: Complex terrain leads to bidirectional responses of soil respiration to inter-annual water availability, *Glob. Change Biol.*, 18, 749–756, 2012.

Rollins, A. W., Adams, H. S., and Stephenson, S. L.: Changes in forest composition and structures across the red spruce-hardwood ecotone in the central Appalachians, *Castanea*, 75, 303–314, 2010.

Rose, C.: *An Introduction to the Environmental Physics of Soil, Water and Watersheds*, University of Cambridge Press, 442 pp., 2004.

Savage, K. E. and Davidson, E. A.: Interannual variation of soil respiration in two New England forests, *Global Biogeochem. Cy.*, 15, 337–350, 2001.

Scanlon, D. and Moore, T.: Carbon dioxide production from peatland soil profiles: the influence of temperature, oxic/anoxic conditions and substrate, *Soil Sci.*, 165, 153–160, 2000.

Schimel, D. S.: Terrestrial ecosystems and the carbon cycle, *Glob. Change Biol.*, 1, 77–91, 1995.

Schaufler, G., Kitzler, B., Schindlbacher, A., Skiba, U., Sutton, M. A., and Zechmeister-Boltenstern, S.: Greenhouse gas emissions from European soils under different land use: effects of soil moisture and temperature, *Eur. J. Soil Sci.*, 61, 683–696, 2010.

Schlesinger, W. H.: Carbon balance in terrestrial detritus, *Annu. Rev. Ecol. Syst.*, 8, 51–81, 1977.

Schmidt, M. W., Torn, M. S., Abiven, S., Dittmar, T., Guggenberger, G., Janssens, I. A., Kleber, M., Kögel-Knabner, I., Lehmann, J., Manning, D. A. C., Nannipieri, P., Rasse, D. P.,

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Weiner, S., and Trumbore, S. E.: Persistence of soil organic matter as an ecosystem property, *Nature*, 478, 49–56, 2011.

Schuur, E. A. and Matson, P. A.: Net primary productivity and nutrient cycling across a mesic to wet precipitation gradient in Hawaiian montane forest, *Oecologia*, 128, 431–442, 2001.

5 Scott-Denton, L. E., Sparks, K. L., and Monson, R. K.: Spatial and temporal controls of soil respiration rate in a high-elevation, subalpine forest, *Soil Biol. Biochem.*, 35, 525–534, 2003.

Sjögersten, S., van der Wal, R., and Woodin, S. J.: Small-scale hydrological variation determines landscape CO<sub>2</sub> fluxes in the high Arctic, *Biogeochemistry*, 80, 205–216, 2006.

10 Trumbore, S.: Carbon respired by terrestrial ecosystems-recent progress and challenges, *Glob. Change Biol.*, 12, 141–153, 2006.

University of Utah, National Weather Service Forecast Office in Salt Lake City, Dataset Title: Utah MesoWest Weather Data from the University of Utah and the National Weather Service Forecast Office in Salt Lake City, available at: <http://mesowest.utah.edu/>, last access: 23 October 2014.

15 Van Gestel, M., Merckx, R., and Vlassak, K.: Microbial biomass responses to soil drying and rewetting: the fate of fast – and slow – growing microorganisms in soils from different climates, *Soil Biol. Biochem.*, 25, 109–123, 1993.

Vargas, R., Hasselquist, N., Allen, E. B., and Allen, M. F.: Effects of a hurricane disturbance on aboveground forest structure, arbuscular mycorrhizae and belowground carbon in a restored tropical forest, *Ecosystems*, 13, 118–128, doi:10.1007/s10021-009-9305-x, 2010.

20 Vesterdal, L., Elberling, B., Christiansen, J. R., Callesen, I., and Schmidt, I. K.: Soil respiration and rates of soil carbon turnover differ among six common European tree species, *Forest Ecol. Manag.*, 264, 185–196, 2012.

25 Wu, W., Tang, X. P., Guo, N. J., Yang, C., Liu, H. B., and Shang, Y. F.: Spatiotemporal modeling of monthly soil temperature using artificial neural networks, *Theor. Appl. Climatol.*, 113, 481–494, 2013.

Wu, Z., Dijkstra, P., Koch, G. W., Penuelas, J., and Hungate, B. A.: Responses of terrestrial ecosystems to temperature and precipitation change: a meta-analysis of experimental manipulation, *Glob. Change Biol.*, 17, 927–942, 2011.

30 Wullschleger, S. D., Tschaplinski, T. J., and Norby, R. J.: Plant water relations at elevated CO<sub>2</sub>-implications for water-limited environments, *Plant Cell Environ.*, 25, 319–331, 2002.

Wurzburger, N. and Hendrick, R. L.: Rhododendron thickets alter N cycling and soil extracellular enzyme activities in southern Appalachian hardwood forests, *Pedobiologia*, 50, 563–576, 2007.

5 Xu, M. and Qi, Y.: Spatial and seasonal variations of  $Q_{10}$  determined by soil respiration measurements at a Sierra Nevadan forest, *Global Biogeochem. Cy.*, 15, 687–696, 2001.

## BGD

11, 17631–17673, 2014

### Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 1.** Least-squares means of dynamic environmental variables. Error terms indicate standard error.

Year	Treatment	$F_{\text{SOIL}}$ ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ )	WFPS ( $\text{m}^3 \text{ m}^{-3}$ )	Soil temperature ( $^{\circ}\text{C}$ )
2010	LOW	4.69 ± 0.687	0.189 ± 0.014	16.29 ± 0.656
2010	MID	6.13 ± 0.691	0.184 ± 0.014	14.90 ± 0.656
2010	HIGH	6.32 ± 0.668	0.141 ± 0.014	15.30 ± 0.654
2011	LOW	4.75 ± 0.571	0.247 ± 0.012	16.61 ± 0.520
2011	MID	4.82 ± 0.561	0.250 ± 0.012	15.31 ± 0.519
2011	HIGH	4.76 ± 0.551	0.249 ± 0.012	15.54 ± 0.518
2012	LOW	4.45 ± 0.722	0.184 ± 0.014	15.08 ± 0.659
2012	MID	4.04 ± 0.702	0.206 ± 0.014	13.93 ± 0.658
2012	HIGH	4.71 ± 0.681	0.183 ± 0.014	13.98 ± 0.656
2010	OPEN	4.54 ± 0.685	0.164 ± 0.014	15.67 ± 0.656
2010	SHRUB	7.48 ± 0.674	0.187 ± 0.014	15.42 ± 0.655
2010	CANOPY	5.11 ± 0.674	0.167 ± 0.014	15.39 ± 0.655
2011	OPEN	4.02 ± 0.562	0.225 ± 0.012	16.31 ± 0.519
2011	SHRUB	5.63 ± 0.559	0.270 ± 0.012	15.38 ± 0.518
2011	CANOPY	4.68 ± 0.557	0.251 ± 0.012	15.76 ± 0.518
2012	OPEN	3.77 ± 0.698	0.188 ± 0.014	14.86 ± 0.656
2012	SHRUB	5.12 ± 0.705	0.188 ± 0.014	13.98 ± 0.658
2012	CANOPY	4.31 ± 0.697	0.198 ± 0.014	14.15 ± 0.657
	LOW	4.61 ± 0.431	0.207 ± 0.010	15.99 ± 0.356
	MID	4.99 ± 0.427	0.214 ± 0.009	14.71 ± 0.356
	HIGH	5.25 ± 0.418	0.191 ± 0.009	14.94 ± 0.355
	OPEN	4.09 ± 0.425	0.191 ± 0.009	15.61 ± 0.355
	SHRUB	6.07 ± 0.424	0.214 ± 0.009	14.93 ± 0.355
	CANOPY	4.69 ± 0.423	0.206 ± 0.009	15.10 ± 0.355
2010		5.71 ± 0.634	0.172 ± 0.013	15.50 ± 0.652
2011		4.78 ± 0.525	0.248 ± 0.011	15.82 ± 0.516
2012		4.36 ± 0.647	0.192 ± 0.013	14.36 ± 0.653

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 2.** Statistical table from repeated measures mixed-model ANOVA. For all comparisons by elevation, vegetation and year,  $n = 633$ ;  $df = 2, 633$ . For elevation by year and vegetation by year comparisons,  $n = 633$ ;  $df = 4, 633$ .

	F	<i>p</i>
Elevation		
$F_{\text{soil}}$	3.44	0.0326
WFPS (0–12 cm)	11.13	< 0.0001
Soil temp (12 cm)	170.76	< 0.0001
Vegetation		
$F_{\text{soil}}$	37.58	< 0.0001
WFPS (0–12 cm)	11.20	< 0.0001
Soil temp (12 cm)	52.79	< 0.0001
Elevation by vegetation		
$F_{\text{soil}}$	2.47	0.0436
WFPS (0–12 cm)	24.48	< 0.0001
Soil temp (12 cm)	9.55	< 0.0001
Year		
$F_{\text{soil}}$	1.40	0.2464
WFPS (0–12 cm)	17.27	< 0.0001
Soil temp (12 cm)	1.66	0.1918
Elevation by year		
$F_{\text{soil}}$	3.17	0.0134
WFPS (0–12 cm)	6.05	< 0.0001
Soil temp (12 cm)	1.02	0.3945
Vegetation by year		
$F_{\text{soil}}$	2.96	0.0192
WFPS (0–12 cm)	4.08	0.0034
Soil temp (12 cm)	5.46	0.0003

Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)**Table A1.** Least-squares means of vegetation variables and soil chemical and physical properties. Error terms indicate standard error.

Elevation	Vegetation	PAI (m <sup>3</sup> m <sup>-3</sup> )	MI (lux)	Soil pH
LOW	OPEN	1.06 ± 0.42	46 856.33 ± 2697.8	3.99 ± 0.14
LOW	SHRUB	2.01 ± 0.42	72 819.75 ± 3672.5	4.26 ± 0.14
LOW	CANOPY	1.82 ± 0.42	29 966.01 ± 1589.6	3.99 ± 0.14
MID	OPEN	1.49 ± 0.42	42 500.11 ± 3796.2	4.32 ± 0.14
MID	SHRUB	3.68 ± 0.42	19 923.95 ± 1194.9	4.11 ± 0.14
MID	CANOPY	1.54 ± 0.42	25 855.61 ± 1465.3	4.13 ± 0.14
HIGH	OPEN	2.70 ± 0.42	26 230.93 ± 1556.2	4.11 ± 0.14
HIGH	SHRUB	4.71 ± 0.42	12 060.48 ± 931.0	3.87 ± 0.14
HIGH	CANOPY	3.05 ± 0.51	20 273.25 ± 1174.5	4.17 ± 0.14
LOW		1.63 ± 0.24	49 879.7 ± 1932.9	4.08 ± 0.08
MID		2.23 ± 0.24	29 138.82 ± 1486.5	4.18 ± 0.08
HIGH		3.49 ± 0.26	19 521.56 ± 801.0	4.05 ± 0.08
	OPEN	1.75 ± 0.24	47 346.97 ± 2179.5	4.14 ± 0.08
	SHRUB	3.46 ± 0.24	26 375.92 ± 1389.7	4.08 ± 0.08
	CANOPY	2.14 ± 0.26	25 361.26 ± 852.7	4.10 ± 0.08

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table A2.** Mixed-model ANOVA results for the main and interactive effects of elevation, vegetation, and soil depth on soil bulk density ( $\rho_s$ ) and total bulk density ( $\rho_t$ ).

Treatment	Depth (cm)	Soil bulk density ( $\rho_s$ )		Total bulk density ( $\rho_t$ )	
		<i>F</i>	<i>p</i> value	<i>F</i>	<i>p</i> value
Elevation		5.77	0.0053	4.79	0.0120
Vegetation		10.55	0.0001	9.93	0.0002
Soil depth		15.70	< 0.0001	17.80	< 0.0001
Elevation*vegetation		0.40	0.8089	0.29	0.8851
Elevation*depth		1.70	0.1619	1.57	0.1951
Vegetation*depth		0.31	0.8719	0.18	0.9501
Elevation	5	4.11	0.0436	4.67	0.0316
Vegetation	5	2.72	0.1059	3.10	0.0822
Elevation*vegetation	5	1.28	0.3300	1.27	0.3342
Elevation	12	1.63	0.2228	1.17	0.3333
Vegetation	12	3.60	0.0483	3.47	0.0533
Elevation*vegetation	12	0.73	0.5856	0.66	0.6286
Elevation	20	4.15	0.0330	3.35	0.0582
Vegetation	20	5.15	0.0170	4.19	0.0321
Elevation*vegetation	20	0.30	0.8733	0.16	0.9551

## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

**Table A3.** Kruskal–Wallis rank sum test results for the effects of elevation, vegetation, and soil depth on soil organic matter (SOM %), soil C (%), soil N (%) and soil C : N ratio.

Treatment	Depth (cm)	SOM (%)		Soil C (%)		Soil N (%)		Soil C : N	
		$\chi^2$	<i>p</i> value	$\chi^2$	<i>p</i> value	$\chi^2$	<i>p</i> value	$\chi^2$	<i>p</i> value
Elevation		1.82	0.401	4.59	0.101	5.08	0.078	1.4	0.496
Vegetation		8.21	0.016	10.64	0.004	6.83	0.032	30.08	< 0.0001
Depth		36.18	< 0.0001	98.61	< 0.0001	111.28	< 0.0001	13.52	0.004
Elevation	5	0.39	0.822	10.63	0.004	11.05	0.004	6.47	0.039
Vegetation	5	8.99	0.011	5.60	0.061	4.19	0.123	12.09	0.002
Elevation	12	2.03	0.361	4.72	0.094	6.35	0.042	0.47	0.812
Vegetation	12	2.55	0.278	3.05	0.216	2.72	0.257	4.21	0.122
Elevation	20	5.72	0.057	9.29	0.009	11.68	0.002	0.64	0.724
Vegetation	20	6.14	0.046	15.28	< 0.0001	11.30	0.004	23.66	< 0.0001

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table A4.** Total soil carbon (%), total soil nitrogen (%), total soil C : N ratio and soil organic matter (SOM) (%) by all combinations of elevation, vegetation, depth levels and treatments.

Elevation	Vegetation	Depth	Soil C (%)	Soil N (%)	Soil C : N	SOM (%)
LOW	OPEN	5	14.11 ± 1.14	0.87 ± 0.069	16.12 ± 0.33	42.06 ± 7.88
LOW	OPEN	12	13.89 ± 2.47	0.67 ± 0.065	18.69 ± 1.84	25.79 ± 8.92
LOW	OPEN	20	7.43 ± 0.87	0.42 ± 0.042	15.42 ± 0.45	15.26 ± 4.71
LOW	SHRUB	5	29.62 ± 1.37	1.53 ± 0.063	19.02 ± 0.22	53.44 ± 8.21
LOW	SHRUB	12	26.13 ± 1.20	1.23 ± 0.060	21.42 ± 0.78	18.20 ± 2.38
LOW	SHRUB	20	10.92 ± 0.55	0.54 ± 0.027	20.19 ± 0.18	22.46 ± 3.31
LOW	CANOPY	5	19.41 ± 1.07	0.96 ± 0.041	20.19 ± 0.51	34.83 ± 2.46
LOW	CANOPY	12	20.47 ± 2.39	1.05 ± 0.112	18.64 ± 0.30	68.45 ± 3.42
LOW	CANOPY	20	5.96 ± 0.49	0.36 ± 0.023	14.85 ± 0.42	14.32 ± 2.90
MID	OPEN	5	11.11 ± 0.85	0.66 ± 0.044	15.57 ± 0.35	33.13 ± 6.91
MID	OPEN	12	14.47 ± 0.63	0.81 ± 0.034	17.89 ± 0.30	17.41 ± 2.61
MID	OPEN	20	3.71 ± 0.06	0.23 ± 0.002	15.99 ± 0.12	7.06 ± 0.44
MID	SHRUB	5	15.62 ± 1.51	0.75 ± 0.060	19.97 ± 0.42	29.17 ± 7.44
MID	SHRUB	12	18.79 ± 1.06	0.90 ± 0.041	20.00 ± 0.40	21.86 ± 2.03
MID	SHRUB	20	4.42 ± 0.13	0.25 ± 0.005	17.37 ± 0.20	10.26 ± 0.94
MID	CANOPY	5	17.32 ± 2.60	0.92 ± 0.125	17.90 ± 0.50	43.67 ± 9.03
MID	CANOPY	12	19.47 ± 1.05	1.02 ± 0.041	18.51 ± 0.43	23.57 ± 5.61
MID	CANOPY	20	6.59 ± 0.35	0.37 ± 0.019	17.81 ± 0.16	9.43 ± 0.56
HIGH	OPEN	5	27.22 ± 2.86	1.30 ± 0.120	19.14 ± 1.21	65.76 ± 4.91
HIGH	OPEN	12	23.51 ± 2.44	1.18 ± 0.093	18.75 ± 0.79	32.76 ± 8.74
HIGH	OPEN	20	4.63 ± 0.26	0.28 ± 0.006	15.80 ± 0.54	9.95 ± 1.01
HIGH	SHRUB	5	48.12 ± 1.17	2.12 ± 0.007	22.65 ± 0.48	80.80 ± 3.38
HIGH	SHRUB	12	28.13 ± 3.15	1.34 ± 0.117	19.88 ± 0.65	44.01 ± 5.95
HIGH	SHRUB	20	11.76 ± 0.39	0.60 ± 0.019	19.55 ± 0.13	21.67 ± 2.28
HIGH	CANOPY	5	24.27 ± 3.25	1.22 ± 0.166	20.16 ± 0.45	54.29 ± 7.70
HIGH	CANOPY	12	29.59 ± 3.40	1.47 ± 0.128	19.22 ± 1.01	37.40 ± 16.8
HIGH	CANOPY	20	6.52 ± 0.19	0.39 ± 0.009	16.42 ± 0.12	13.69 ± 0.38
LOW		5	21.61 ± 0.55	1.15 ± 0.026	18.38 ± 0.14	43.44 ± 9.26
LOW		12	20.21 ± 0.81	0.99 ± 0.035	19.45 ± 0.30	38.94 ± 13.5
LOW		20	8.13 ± 0.22	0.44 ± 0.010	16.87 ± 0.15	17.35 ± 5.01
MID		5	14.17 ± 0.49	0.76 ± 0.022	17.57 ± 0.16	35.33 ± 10.2
MID		12	17.70 ± 0.32	0.91 ± 0.013	18.84 ± 0.13	20.95 ± 4.88
MID		20	4.80 ± 0.07	0.88 ± 0.003	16.99 ± 0.06	8.91 ± 1.10
HIGH		5	31.46 ± 1.25	1.48 ± 0.053	20.36 ± 0.34	66.95 ± 8.95
HIGH		12	26.90 ± 0.93	1.32 ± 0.035	19.29 ± 0.24	38.14 ± 11.2
HIGH		20	7.90 ± 0.15	0.44 ± 0.006	17.40 ± 0.11	15.10 ± 3.09
	OPEN	5	16.30 ± 0.58	0.89 ± 0.026	16.68 ± 0.20	46.98 ± 10.91
	OPEN	12	16.99 ± 0.54	0.89 ± 0.022	18.33 ± 0.24	25.26 ± 9.05
	OPEN	20	5.144 ± 0.16	0.31 ± 0.007	15.76 ± 0.11	10.76 ± 3.9
	SHRUB	5	27.84 ± 0.90	1.35 ± 0.039	20.06 ± 0.14	54.47 ± 13.6
	SHRUB	12	23.02 ± 0.62	1.10 ± 0.024	20.28 ± 0.18	28.03 ± 7.57
	SHRUB	20	9.134 ± 0.17	0.47 ± 0.008	19.06 ± 0.07	18.13 ± 4.12
	CANOPY	5	20.05 ± 0.74	1.02 ± 0.035	19.36 ± 0.17	44.27 ± 9.64
	CANOPY	12	21.92 ± 0.63	1.12 ± 0.027	18.70 ± 0.17	43.86 ± 13.7
	CANOPY	20	6.353 ± 0.11	0.37 ± 0.005	16.30 ± 0.10	12.48 ± 2.42
LOW			17.01 ± 0.19	0.88 ± 0.009	18.10 ± 0.05	33.03 ± 22.0
MID			14.09 ± 0.13	0.74 ± 0.006	18.06 ± 0.03	21.73 ± 17.1
HIGH			19.52 ± 0.24	0.97 ± 0.010	18.55 ± 0.05	40.14 ± 27.2
	OPEN		13.36 ± 0.15	0.73 ± 0.007	16.94 ± 0.05	27.76 ± 22.6
	SHRUB		20.16 ± 0.19	0.99 ± 0.009	19.68 ± 0.03	33.54 ± 24.3
	CANOPY		16.33 ± 0.18	0.84 ± 0.008	18.03 ± 0.04	33.14 ± 23.9
		5	21.14 ± 0.26	1.07 ± 0.011	18.55 ± 0.06	48.58 ± 23.4
		12	20.73 ± 0.20	1.04 ± 0.008	19.11 ± 0.06	32.21 ± 21.2
		20	6.93 ± 0.05	0.38 ± 0.002	17.09 ± 0.03	13.79 ± 7.80

Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

[Title Page](#)

[Abstract](#)

[Introduction](#)

[Conclusions](#)

[References](#)

[Tables](#)

[Figures](#)



[Back](#)

[Close](#)

[Full Screen / Esc](#)

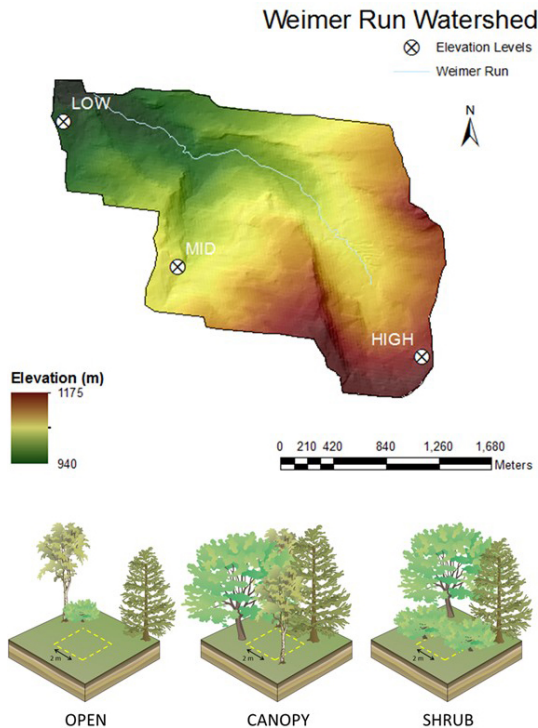
[Printer-friendly Version](#)

[Interactive Discussion](#)



**Table A5.** Least-squares means of soil porosity ( $\Phi$ ) and soil bulk density ( $\rho$ ) for the 0–12 cm soil depth used in calculating WFPS.

Elevation	Vegetation	$\Phi$ (m <sup>3</sup> m <sup>-3</sup> )	$\rho$ (g cm <sup>-3</sup> )
LOW	OPEN	0.68 ± 0.05	0.83 ± 0.14
LOW	SHRUB	0.81 ± 0.05	0.49 ± 0.14
LOW	CANOPY	0.75 ± 0.05	0.64 ± 0.14
MID	OPEN	0.58 ± 0.05	1.11 ± 0.14
MID	SHRUB	0.72 ± 0.05	0.72 ± 0.14
MID	CANOPY	0.61 ± 0.05	1.01 ± 0.14
HIGH	OPEN	0.69 ± 0.05	0.81 ± 0.14
HIGH	SHRUB	0.79 ± 0.05	0.55 ± 0.14
HIGH	CANOPY	0.68 ± 0.05	0.83 ± 0.14
LOW		0.75 ± 0.03	0.65 ± 0.08
MID		0.64 ± 0.03	0.95 ± 0.08
HIGH		0.72 ± 0.03	0.73 ± 0.08
	OPEN	0.65 ± 0.03	0.92 ± 0.08
	SHRUB	0.77 ± 0.03	0.58 ± 0.08
	CANOPY	0.68 ± 0.03	0.83 ± 0.08



**Figure 1.** (Above) Weimer Run watershed (374 ha) with elevation levels indicated on map. (Below) Conceptual diagram showing vegetation treatments. Images courtesy of the Integration and Application Network, University of Maryland Center for Environmental Science (ian.umces.edu/symbols/).

Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.

Title Page

Abstract Introduction

Conclusions References

Tables Figures

◀ ▶

◀ ▶

Back Close

Full Screen / Esc

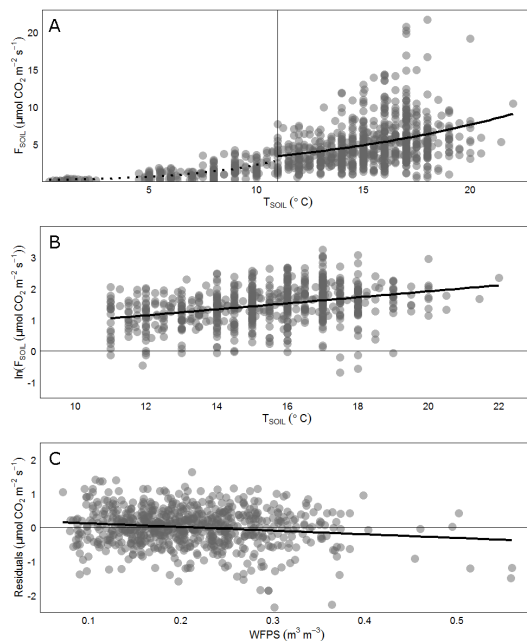
Printer-friendly Version

Interactive Discussion



Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.



**Figure 2.** (a) Soil CO<sub>2</sub> efflux ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) against soil temperature ( $^{\circ}\text{C}$ ) at 12 cm with data split at 11  $^{\circ}\text{C}$ . For all data  $R^2 = 0.3163$ . For flux rate values below 11  $^{\circ}\text{C}$ ,  $R^2 = 0.434$ , for flux rate values above 11  $^{\circ}\text{C}$ ,  $R^2 = 0.104$ . (b) Natural log of soil CO<sub>2</sub> efflux ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) against soil temperature ( $^{\circ}\text{C}$ ) at 12 cm for all data above 11  $^{\circ}\text{C}$ . For flux rate values below 11  $^{\circ}\text{C}$ ,  $R^2 = 0.1188$ ,  $p = \ll 0.0001$ . (c) Residuals of the natural log of soil CO<sub>2</sub> efflux ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ) against water-filled pore space (0–12 cm) for all data above 11  $^{\circ}\text{C}$ .  $R^2 = 0.0208$ ,  $p = \ll 0.0001$ .

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

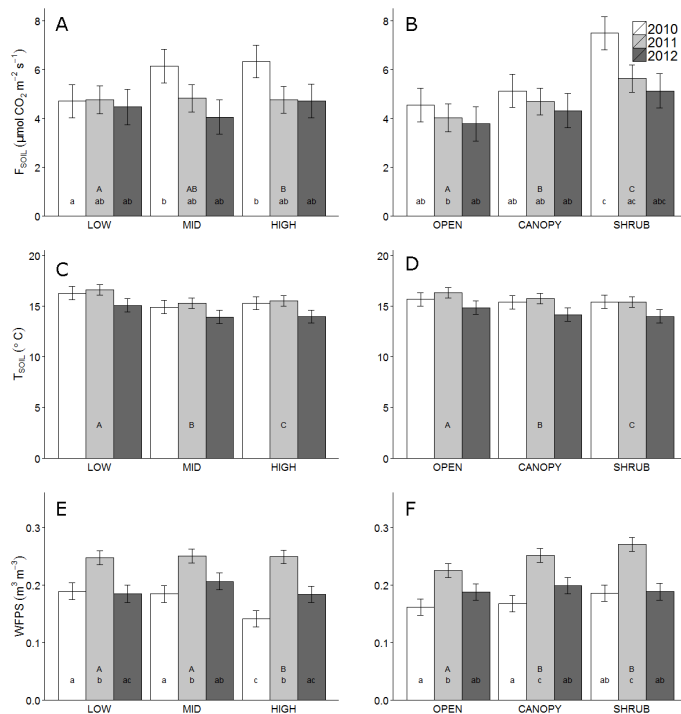
Interactive Discussion





Soil CO<sub>2</sub> efflux in an Appalachian watershed

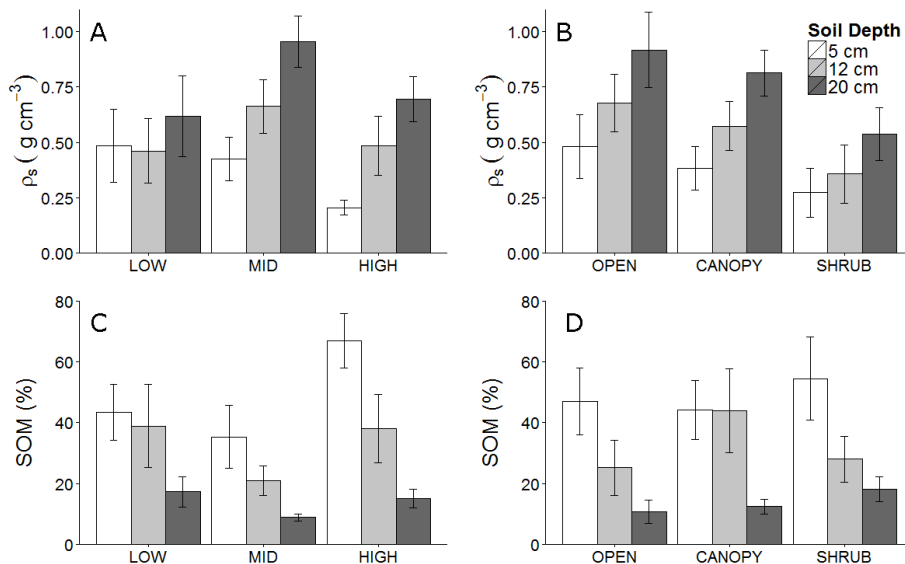
J. W. Atkins et al.



**Figure 3.** (a, c, e) Least-squares means of soil CO<sub>2</sub> efflux ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ); WFPS ( $\text{m}^3 \text{ m}^{-3}$ ); and soil temperature at 12 cm ( $^{\circ}\text{C}$ ) by elevation. (b, d, e) Least-squares means of soil CO<sub>2</sub> efflux ( $\mu\text{mol CO}_2 \text{ m}^{-2} \text{ s}^{-1}$ ); WFPS ( $\text{m}^3 \text{ m}^{-3}$ ); and soil temperature at 12 cm ( $^{\circ}\text{C}$ ) by vegetation. Capital letters indicate difference between elevation treatments and lower case letters indicate differences between vegetation treatments. \* year interactions. Bars indicate standard error. Colors indicate sampling year.

Soil CO<sub>2</sub> efflux in an Appalachian watershed

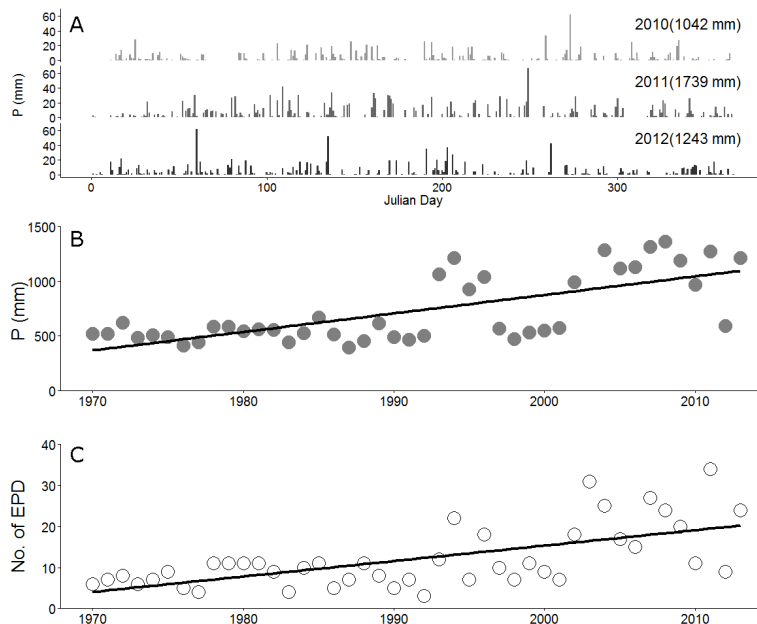
J. W. Atkins et al.



**Figure 4.** (a, c) Means of soil bulk density ( $\text{g cm}^{-3}$ ) and soil organic matter (%) by elevation treatment. (b, d) Means of soil bulk density ( $\text{g cm}^{-3}$ ) and soil organic matter (%) by vegetation treatment. Bars indicate standard error. Colors indicate soil depth profiles.

Soil CO<sub>2</sub> efflux in an Appalachian watershed

J. W. Atkins et al.



**Figure 5.** (a) Hyetographs for 2010, 2011, and 2012 from the Bearden Knob weather station located within the Weimer Run watershed (BDKW2 MesoWest; University of Utah). Precipitation totals by year are indicated within each graph and are in  $\text{mm yr}^{-1}$ . (b) Precipitation for the years 1970–2013 ( $\text{mm yr}^{-1}$ ) from NCDC station Canaan Valley, WV (461393). Linear regression shows that mean annual precipitation is increasing by  $17.88 \text{ mm yr}^{-1}$  ( $r = 0.697$ ;  $F_{1,42} = 39.74$ ;  $r^2 = 0.474$ ;  $p \leq 0.0001$ ). The year to year variance in precipitation is also increasing (BP = 8.58;  $p = 0.003$ ). (c) Number of extreme precipitation days (EPD) per year (defined as days where total precipitation exceeded 25.4 mm per day). The number of EPDs are increasing by 0.38 days per year ( $r = 0.637$ ;  $F_{1,42} = 28.69$ ;  $r^2 = 0.392$ ;  $p \leq 0.0001$ ). The variance is also increasing (BP = 11.12;  $p \leq 0.0001$ ).