

1 **Increase in soil organic carbon by agricultural intensification in Northern China**

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3 Y. Liao<sup>1</sup>, W. L. Wu<sup>1</sup>, F. Q. Meng<sup>1</sup>, P. Smith<sup>2</sup>, R. Lal<sup>3</sup>

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5 <sup>1</sup> *College of Resources and Environmental Sciences, China Agricultural University,*  
6 *Beijing 100193, China*

7 <sup>2</sup> *Institute of Biological & Environmental Sciences, University of Aberdeen, 23 St*  
8 *Machar Drive, Aberdeen, AB24 3UU, UK*

9 <sup>3</sup> *Carbon Management and Sequestration Center, The Ohio State University, 2021*  
10 *CoffeyRoad, Columbus, OH43210, USA*

11  
12 *Correspondence to: F. Q. Meng (mengfq@cau.edu.cn)*

13  
14 **Abstract.** Agricultural intensification has contributed greatly to the sustained food  
15 supply of China's 1.3 billion population over the 30 year period during 1982-2011.  
16 Intensification has several and widely recognized negative environmental impacts  
17 including depletion of water resources, pollution of water bodies, greenhouse gas  
18 emissions and soil acidification. However, there have been few studies over this  
19 period on the impacts of intensification on soil organic carbon (SOC) at the regional  
20 level. The present study was conducted in Huantai county, a typical intensive farming  
21 region in Northern China, to analyze the temporal dynamics of SOC influenced by  
22 climate and farming practices. The results indicate that from 1982 to 2011, SOC  
23 content and stock in the 0-20 cm layer of the cropland increased from  $7.8 \pm 1.6$  to  
24  $11.0 \pm 2.3 \text{ g} \cdot \text{kg}^{-1}$  (41%) and  $21 \pm 4.3$  to  $33.0 \pm 7.0 \text{ Mg} \cdot \text{ha}^{-1}$  (54%), respectively. The  
25 SOC stock (0-20 cm) of the farmland for the entire county increased from 0.75 to 1.2  
26 Tg (59%). Correlation analysis revealed that incorporation of crop residues  
27 significantly increased SOC, while increase in the mean annual temperature decreased  
28 the SOC level. Therefore, agricultural intensification has increased crop productivity  
29 and contributed to SOC sequestration in Northern China. In the near future, more  
30 appropriate technologies and practices must be developed and implemented for a  
31 maintenance or enhancement of SOC in this region and elsewhere in Northern China,

32 that also reduce non-CO<sub>2</sub> greenhouse gas emissions, since the climate benefit from the  
33 additional SOC storage is estimated to be smaller than the negative climate impacts of  
34 N<sub>2</sub>O from N fertilizer additions.

35

36 **Keywords:** Soil organic carbon, agricultural production intensification, crop residue  
37 incorporation, nitrogen fertilizer, temperature, Northern China

38

### 39 **1 Introduction**

40 Increasing soil organic matter (SOM) storage in arable lands can ensure the  
41 sustained supply of nitrogen (N) and other nutrients to crop growth and maintain  
42 appropriate soil quality such as aeration, permeability, water-holding capacity and  
43 nutrient preserving capacity (Smith et al., 2012). Globally, accumulation of SOM or  
44 soil organic carbon (SOC) stock in arable lands, which contributes to the mitigation of  
45 greenhouse effect and a concomitant improvement in soil fertility (Matson et al., 1997;  
46 Sainju et al., 2009), may be achieved by a range of improved farming practices. These  
47 practices include adoption of high-yielding crop varieties, balanced fertilization, crop  
48 residue incorporation, no-till (NT) or reduced tillage, optimal irrigation, high cropping  
49 intensity (Matson et al., 1997; Kucharik, et al., 2001). For instance, agricultural soils  
50 in US had a carbon sink capacity of 1.3-21.2 Tg C annually from 1982 to 1997, due to  
51 land use, NT, higher cropping intensity etc. (Eve et al., 2002; Ogle et al., 2003). The  
52 rate of increase in SOC stock in Canada was 5.7 Tg C yr<sup>-1</sup> during 1991-2001 (Vanden  
53 Bygaart et al., 2004). In The Netherlands, the SOC content of arable land increased by  
54 about 0.08 g kg<sup>-1</sup> yr<sup>-1</sup> between 1984 and 2004 (Reijneveld et al., 2009). Benbi and  
55 Brar (2009) reported that SOC in the Punjab state of India increased from 2.9 to 4.0 g  
56 kg<sup>-1</sup> (38%) between 1981 and 2006, largely resulting from irrigation, optimal  
57 fertilization, and an increase in crop productivity.

58 Northern China is one of the most important agricultural regions, producing 60-80%  
59 and 35-40% of nation's wheat (*Triticum aestivum* L.) and maize (*Zea mays* L.),  
60 respectively (NBSC, 2014). Similar to other regions in Northern China, Huantai  
61 county has experienced the agricultural intensification process including a high

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64 cropping intensity (200%, winter wheat, summer maize), high fertilizer rate (500-600  
65 kg N ha<sup>-1</sup> yr<sup>-1</sup>), frequent irrigation and tillage, and an increasing ratio of crop residues  
66 incorporation since the 1980s (Shi et al., 2013; Kong et al., 2014). Agricultural  
67 intensification increased crop yield within a short period. For example, Huantai  
68 county achieved production of >15 Mg grains (wheat + maize) ha<sup>-1</sup> yr<sup>-1</sup> for all of its  
69 farmland since 1990. The effect of agricultural intensification on increasing crop  
70 yields has been well documented, however, most of the research done on SOC  
71 sequestration in agricultural soils is confined to long-term plot-scale experiments.  
72 Studies available at national or region levels were within short period, especially in  
73 Northern China (<20 years), where is characterized by low levels of SOM.

74 We collected three-decades of data of climate, farm management and crop yield  
75 from Huantai county, to: 1) analyze the evolution of SOC at the regional level from  
76 1982 to 2011, and 2) establish the cause-effect relationship between the driving forces  
77 and SOC change. The results derived from this study may contribute to improved  
78 farm management for the long-term sustainable agricultural development in the  
79 intensive farming of Northern China and elsewhere.

80

## 81 2 Materials and methods

### 82 2.1 Study area

83 Huantai county, with an area of 509 km<sup>2</sup>, is located in Northern China  
84 (36°51'50"-37°06'00" N latitude and 117°50'00"-118°10'40" E longitude). It is  
85 characterized by a warm temperate continental monsoon climate, with annual average  
86 temperature of 13.4 °C and annual precipitation of 604 mm. The rainfall occurs  
87 mainly in June, July and August; with the annual frost-free season of about 198 days.  
88 Slope gradient of the landscape is low in northern and high in southern regions, with  
89 an average altitude of 6.5-29.5 m, falling gently from southwest to northeast. The  
90 main soil types, according to the US soil classification system, include Haplustalfs,  
91 Aquepts and Vertisols. The household contract responsibility system was  
92 implemented in Huantai county in 1980, and land productivity has increased  
93 significantly with an increase in fertilizer input, frequent irrigation and adoption of

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110 high-yielding varieties. In 1990, Huantai county became the first *Dun Liang County*  
111 which achieved high productivity of one Mg of grain per mu for the whole county (1  
112 mu=1/15 ha) in China. Vegetable production has also been intensified in the county  
113 since 1990s.

114

## 115 2.2 Data Collection

116 The soil, climate and farming data were collected between 2011 to 2013 from the  
117 sources as described below:

118 - *The 2<sup>nd</sup> National Soil Survey*: The 2<sup>nd</sup> National Soil Survey in China was  
119 conducted during 1981-1983, and was undertaken in 1982 for Huantai county. The  
120 survey collected and analyzed soil samples for genesis, physical, chemical and  
121 biological properties. The SOM content was tested using the potassium dichromate  
122 titrimetric method (Jankauskas et al., 2006). The SOM data in this study was obtained  
123 for 258 soil samples (0-20 cm), together with the corresponding GPS location  
124 coordinates.

125 - *Annual Soil Fertility Survey*: The annual SOM data of farmland soil (0-20 cm)  
126 were collected from the Soil Fertility Survey, a program undertaken by the county  
127 agricultural extension since 1987. Every year, samples from the topsoil layer were  
128 collected after the autumn harvest and analyzed for pH, SOM and nutrient contents.  
129 Soil sampling in the survey was stratified according to the division of administrative  
130 villages. Each village had at least one composite soil sample, representing 6.7-33.3 ha  
131 of farmland. The number of soil samples was from 199 (in 2003) to 3637 (in 2007)  
132 and the average was about 800 each year. The SOM content was determined by the  
133 potassium dichromate titrimetric method (Jankauskas et al., 2006). The data for 1988,  
134 1991, 2000, 2001, 2004, 2005 and 2010 were for each town as the survey was only  
135 done at the township level. As there were no data for the period 1983 to 1986, it was  
136 assumed that a steady change in SOC occurred during this period. Thus, the average  
137 SOC content for the whole county was obtained by interpolation based on the data  
138 from 1982 and 1987.

139 - *Soil Sampling and Measurement Program*: In September 2011, a soil survey

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141 was implemented for the whole county. Soil samples were obtained from 0-20 cm  
142 depth of the farmland (including cropland and vegetable land), in an evenly  
143 distributed 2 × 2 km grid. Each sample was composited from three collection points.  
144 The GPS location was recorded for each sampling point. Soil was ground and passed  
145 through a 0.15 mm sieve. For the SOC analysis, 2-3 g of the soil sample was weighed  
146 into the beaker, and then 20 mL HCl of 0.5 mol·L<sup>-1</sup> was added for acidification and  
147 removal of carbonates. Samples were then transferred into a rotary oscillator  
148 (Ronghua, HY-B) to shake for 30 min. After standing for 12 h, deionized water was  
149 used to remove the acid from soil samples until neutral pH was obtained. Soil samples  
150 were dried in the oven at 60 °C and weighed. The SOC content was determined by a  
151 C and N Elemental Analyzer (Thermo EA flash 1112). As acidification led to soil  
152 mass loss, the SOC data were corrected using Eq. (1):

153

$$154 \quad W_{soil2} = \frac{m_{soil1} \times W_{soil1}}{m_{soil2}} \quad (1)$$

155

156 Where,  $W_{soil2}$  is the corrected SOC content,  $m_{soil1}$  is the soil mass after  
157 acidification,  $W_{soil1}$  is the SOC content after acidification and  $m_{soil2}$  is the soil mass  
158 before acidification. The comparative study of SOC content determined by Elemental  
159 Analyzer and potassium dichromate titrimetric method indicated that these two  
160 methods are comparable and the differences in between are not significant (Wang et  
161 al., 2014).

162 - *Climate and farm management*: Data were obtained for land use (e.g. area under  
163 crops, vegetables and urban use), grain and straw yields of wheat and maize, nitrogen  
164 fertilizer rate, and the rate of straw incorporation from the Huantai agricultural  
165 yearbook of 1982 - 2011. Climate data for temperature and precipitation were  
166 obtained from the China Climatic Data Center, National Meteorological Information  
167 Center, CMA (<http://cdc.cma.gov.cn>). The temperature and precipitation data was  
168 daily value for the county level.

169 - *Calculation of SOC density and stock*: The Van Bemmelen conversion factor,

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172 namely that SOM contains 58% carbon (Page et al., 1982), was used to compute SOC  
173 content in (Eq. (2)). The SOC density and stock were computed by Eq. (3) and Eq. (4),  
174 respectively:

175

$$176 \quad SOC = SOM \times 0.58 \times 10 \quad (2)$$

$$177 \quad SOCD = SOC \times \gamma \times H \times 10^4 \quad (3)$$

$$178 \quad SOCS = SOCD \times S \times 10^{-9} \quad (4)$$

179

180 where, SOC is the soil organic carbon content ( $\text{g kg}^{-1}$ ), SOM is the soil organic  
181 matter content (%), *SOCD* is the SOC density ( $\text{kg ha}^{-1}$ ),  $\gamma$  is the soil bulk density (BD,  
182  $\text{g cm}^{-3}$ ), H is the thickness (m) of soil layer (0-20 cm), *SOCS* is the SOC stock (Tg)  
183 for the whole county, and S is the farmland area (ha). Soil BD values were  
184 interpolated over years from measured values taken in 1982 and 2011. Average  
185 number of cropland and vegetable land soil samples was 800 and 79, respectively. For  
186 the calculation of SOCS, we assumed that the SOC content of land for industrial or  
187 residential use was maintained at the same level after the farmland was converted.

### 188 2.3 Data analysis

189 The Kolmogorov-Smirnov test was used in the SPSS Statistics 17.0 package to  
190 determine if SOC content followed a normal distribution. Central tendency, dispersion  
191 degree and distribution characteristics of SOC data were calculated. Pearson  
192 correlation analysis and/or partial correlation analysis were conducted between SOC  
193 content and driving factors including mean annual temperature, mean annual  
194 precipitation, grain yield, nitrogen fertilizer rate and straw C incorporation. A  
195 multivariate regression model was developed to account for the impact of these  
196 factors on temporal SOC change in Huantai county.

197

## 198 3 Results

### 199 3.1 Evolution of SOC content and density between 1982 and 2011

200 Figure 1 presents the dynamic changes of topsoil SOC content and density for the

201 cropland (winter wheat-summer maize) from 1982 to 2011. The mean increase in  
202 cropland from 1982 to 2011 was  $7.8 \pm 1.6 \text{ g kg}^{-1}$  to  $11.0 \pm 2.3 \text{ g kg}^{-1}$  for SOC content  
203 and,  $21.4 \pm 4.3 \text{ Mg ha}^{-1}$  to  $33.0 \pm 7.0 \text{ Mg ha}^{-1}$  for SOC density, with rates of increase of  
204 41% and 54%, respectively. The mean SOC content of vegetable land increased  
205 similarly as cropland, i.e., from  $7.8 \pm 1.6 \text{ g kg}^{-1}$  to  $11.0 \pm 2.8 \text{ g kg}^{-1}$  (data not shown).  
206 This trend indicates that the rate of increase in SOC content and density of cropland  
207 (0-20 cm) in Huantai county since the early 1980s was  $0.11 \text{ g kg}^{-1} \text{ yr}^{-1}$  and  $0.40 \text{ Mg}$   
208  $\text{ha}^{-1} \text{ yr}^{-1}$ , respectively. The growth of SOC density was significantly related to  
209 increasing SOC content, but also supported by the increase in soil BD in the 0-20 cm  
210 layer of the farmland ( $1.4 \text{ g cm}^{-3}$  in 1982 to  $1.5 \text{ g cm}^{-3}$  in 2011).

211

### 212 | 3.2 Change of SOC stock in Huantai county from 1982 to 2011

213 An adjustment in the local agricultural sector altered the land use between 1982  
214 and 2011 in Huantai county. The farmland area in 1982 was 35204 ha, of which more  
215 than 99% was under winter wheat - summer maize cropping (cropland). Due to the  
216 expansion of vegetable production since 1990s, land used for vegetable production  
217 increased to about 20% by the early 2000s, and has remained constant until 2011 at  
218 | about 6000 ha. The average SOC content of vegetable land is not significant **different**  
219 from that of cropland ( $11.0 \text{ vs. } 11.0 \text{ g kg}^{-1}$  in 2011 for whole county). Some farmland  
220 was converted to construction use with the expansion of industry and urban land uses.  
221 Of the 31% reduction of farmland area between 1982 and 2011, 16% was used for  
222 vegetable farming and 15% converted to construction land.

223 Although area under cropland decreased from 35,204 ha in 1982 to 24,343 ha in  
224 2011, the SOC stock of cropland (0-20 cm) increased from  $0.75 \pm 0.15$  to  $0.80 \pm 0.17$   
225 Tg C (an increase by 6.7%, Fig. 2). When the SOC stock in vegetable land and  
226 construction land (converted from cropland and for industrial and residential use. The  
227 area increased every year as urbanization took place) was also included, total SOC  
228 stock of the farmland in the Huantai county was estimated to be 1.2 Tg C in 2011,  
229 | with a total increase of 59% compared to the  $0.75 \pm 0.15$  Tg C in 1982 (Fig. 2).\_

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### 232 3.3 The cause-effect relationship governing change in SOC level

233 Among the natural/climate forces which can influence SOC level, mean air  
234 temperature in Huantai county increased within the period of 1982 to 2011, with a  
235 relationship represented by the regression equation  $y = 0.073x + 12.2$  ( $R^2 = 0.67$ ,  
236  $P < 0.0001$ , Fig. 3a). However, there was no significant change in precipitation over  
237 this period. Nitrogen fertilizer input for farmland was  $400 \text{ kg N ha}^{-1} \text{ yr}^{-1}$  in the 1980s  
238 and peaked at about  $600 \text{ kg N ha}^{-1} \text{ yr}^{-1}$  in 1994, followed by a decline to  $500 \text{ kg N ha}^{-1}$   
239  $\text{yr}^{-1}$  in 2011. There were also significant increases in grain yield over this period ( $R^2 =$   
240  $0.63$ ,  $P < 0.0001$ , Fig. 3b). The carbon input between 1982 and 1987 was estimated at  
241 about  $800$  to  $1000 \text{ kg C ha}^{-1} \text{ yr}^{-1}$ , mainly through roots. However, crop production  
242 experienced a rapid growth after 1988, and crop residues (mainly wheat straw) were  
243 returned to the soil, leading to a significant increase in the input of biomass-C,  
244 especially the aboveground C. As much as 70% of maize straw was also incorporated  
245 to the farmland from 2007 onward, leading to a total C input of  $> 8000 \text{ kg ha}^{-1} \text{ yr}^{-1}$  ( $R^2 =$   
246  $0.90$ ,  $P < 0.0001$ , Fig. 3c).

247 Correlation and regression analyses were performed between SOC of cropland  
248 and driving factors, i.e., temperature, precipitation, crop (wheat and maize) yield, N  
249 fertilizer rate and C input from crop residues (Table 1). There was a highly significant  
250 correlation ( $P < 0.01$ ) between SOC content and temperature, crop yield, C input from  
251 crop residues, with correlation coefficients ( $r$ ) of 0.55, 0.79 and 0.91, respectively.  
252 The correlation between SOC and N fertilizer rate was also significant ( $r = 0.38$ ,  
253  $P < 0.05$ ). However, there was no significant correlation between SOC content and the  
254 mean annual precipitation.

255 A partial correlation analysis was conducted to determine the relationship  
256 between SOC content of cropland with any one major driving factor, as the effect of a  
257 set of controlling random variables removed (Table 2). The data indicated a highly  
258 significant and positive correlation between SOC content and the C input from crop  
259 residues ( $r = 0.80$ ,  $P < 0.0001$ ), but a negative correlation between SOC content and  
260 annual mean temperature ( $r = -0.42$ ,  $P = 0.027$ ). The weak positive correlation between



261 SOC content and N fertilizer rate ( $r=0.03$ ,  $P=0.86$ ), and weak negative correlation  
262 between SOC and crop yield ( $r=-0.08$ ,  $P=0.70$ ), indicated that N fertilizer and  
263 increasing crop yield did not contributed to the augment of SOC in Huantai from 1982  
264 to 2011. As the effect of C input from crop residues was removed during partial  
265 correlation, the rising temperature during the past 3 decades significantly decreased  
266 the cropland SOC content in Huantai, or promoted the SOC decomposition. In  
267 addition to the partial correlation analysis, a multivariate regression model was also  
268 developed as follows:

269 
$$Y=12.0 - 0.31 \text{ Temperature} - 0.0003 \text{ Precipitation} - 0.0006 \text{ Nitrogen} + 0.0005 \text{ C}$$
  
270 
$$\text{input} - 0.0001 \text{ Yield} \quad (r^2=0.88)$$

271 where, Y is the SOC content in  $\text{g kg}^{-1}$ . The multi-variate regression analysis also  
272 confirmed the positive impact of the input of crop residue carbon and negative impact  
273 of the increase in temperature.

274

## 275 4 Discussion

### 276 4.1 Driving factors for SOC accumulation

277 The SOC level of farmland was influenced by climate (mainly temperature and  
278 precipitation) and farming practices, including crop residue incorporation, N fertilizer  
279 use, crop yield etc. (Khan et al. 2007; Ladha et al. 2011).

280 *Climate factors:* Climate warming may increase the rate of SOM decomposition,  
281 while the effect of precipitation on SOC is mostly known to be positive because in  
282 general SOC increases with increase in precipitation (Post et al., 1982). Whereas air  
283 temperature in the Huantai region has significantly increased since the 1980s, the  
284 precipitation has not. There is a significant positive correlation between the  
285 temperature and SOC content (Table 1), however the partial correlation analysis  
286 showed that the correlation was negative ( $r=-0.42$ ,  $p=0.03$ , Table 2), indicating that  
287 SOC in Huantai county was enhanced by factors other than temperature and  
288 precipitation. Indeed, it may take much longer (~50 yrs) to observe the effect of  
289 climate change on SOC level (De Bruijn et al., 2012), indicating a strong need for  
290 long-term research.

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298 *Nitrogen fertilizer*: In general, N is the most limiting nutrient in crop production  
299 systems (Robertson and Vitousek, 2009). It promotes the production of crop dry  
300 matter (and therefore C input to the soil) while chemically stabilizing C in the soil,  
301 thereby potentially increasing soil C storage (Paustian et al., 1997). A high input of N  
302 fertilizer was a prominent feature of farming in Huantai county, where average N  
303 fertilizer rate increased from 400 kg N ha<sup>-1</sup> yr<sup>-1</sup> in the 1980s to 600 kg N ha<sup>-1</sup> yr<sup>-1</sup> in  
304 the 1990s. However, the rate of N fertilizer gradually declined and stabilized at 500  
305 kg N ha<sup>-1</sup> yr<sup>-1</sup> in the 2010s. Such a trend is attributed to the increased use of machinery  
306 in agriculture and the extension of formula fertilization techniques adopted in the  
307 region. The strategy was to balance the N fertilization rate in consideration of the high  
308 grain output (>15 Mg grain ha<sup>-1</sup> yr<sup>-1</sup>). However, the rate of SOC increase was still  
309 high during this period, as indicated by a weak positive correlation between N  
310 fertilizer rate and SOC level over the 30-year period (Table 1). The vegetable  
311 production in Huantai county is open-field and there is less organic fertilizer input  
312 | (3.5 t ha<sup>-1</sup> yr<sup>-1</sup>) compared with other regions like Shouguang (>15 t ha<sup>-1</sup> yr<sup>-1</sup>) in China,  
313 SOC increase due to organic fertilizers is not significant and at the same level as  
314 cropland (11.0 vs. 10.9 g kg<sup>-1</sup> in 2011).

315 A few studies (Khan et al., 2007; Mulvaney et al., 2009) have reported that even  
316 the long-term input of a massive amount of residue-C and synthetic N fertilizer do not  
317 sequester SOC; this was not the case with intensification of cropland in Northern  
318 China. A major factor lies in the concentration of the principal parameter (i.e., SOM)  
319 which was significantly lower in cropland soils of Northern China than those of the  
320 U.S. Corn Belt (9 vs. 25 g kg<sup>-1</sup>, Ludwig et al., 2011) or countries of the western  
321 Europe, because large areas of farmlands in Northern China were affected by  
322 saline-alkaline processes prior to 1970s. The initial low crop productivity in the  
323 temperate region where N fertilization rate was low resulted in higher SOC level in  
324 Northern China with increasing biomass input achieved by higher input of N. Similar  
325 trends have been reported by other studies from around the world (Song et al., 2005;  
326 Alvarez, 2005). Nonetheless, it is important to point out that any excessive application  
327 of mineral N not only increases the production cost, but also exacerbates the negative

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330 environment effects, including nitrate contamination of ground and surface waters and  
331 N<sub>2</sub>O emission to the atmosphere (Triberti et al., 2008). Carbon sequestration effects  
332 should be valued in a systematic approach (Wang and Cao, 2011). Indeed, the climate  
333 benefit from the additional SOC storage is smaller than the climate damage caused by  
334 N<sub>2</sub>O from N fertilizer additions (N<sub>2</sub>O emissions are estimated (conservatively) to be  
335 ~57 kt CO<sub>2</sub>-eq. yr<sup>-1</sup>, assuming the smallest farmland area of 24343 ha and current  
336 stable N application rate of 500 kg N ha<sup>-1</sup> yr<sup>-1</sup> using the IPCC Tier 1 default Emission  
337 Factor of 0.01; soil C storage over the period is equivalent to 54 kt CO<sub>2</sub>-eq. yr<sup>-1</sup>). If  
338 the indirect emissions (ammonia volatilization, leaching and runoff) are considered,  
339 the environmental effect of nitrogen fertilizer application will be more serious.

340 Balanced fertilization should be widely promoted, for optimization of the  
341 integrated economic benefits and ecosystem services. It is important to understand  
342 that increasing input of N fertilizer increased SOC only when crop residues were  
343 returned to the soil. There may have been either no, or only a slight increase in SOC  
344 level, if the aboveground crop residues were removed or burnt (Alvarez, 2005) so N  
345 fertilization in itself is not a suitable strategy to increase SOC, particularly considering  
346 the over-riding effects of N<sub>2</sub>O emissions from the N fertilizer.

347 *Grain yield:* Grain yield in Huantai county increased from 7200 kg ha<sup>-1</sup> yr<sup>-1</sup> in  
348 1982 to 16117 kg ha<sup>-1</sup> yr<sup>-1</sup> in 2011 (an increase of 124%). The highly significant  
349 correlation (r=0.79, P<0.01) between SOC content and grain yield indicates the  
350 importance of SOM to achieving high crop productivity, and *vice versa* (Pan et al.,  
351 2009). Indeed, the interdependence between crop yield and SOM content is widely  
352 recognized (Lal, 2002; 2013). Increase in the SOC pool of 1 Mg C ha<sup>-1</sup> in the root  
353 zone can increase annual food production by 30-50 million Mg in developing  
354 countries (Lal, 2013). Meta-analysis indicated that crop yield greatly increased by  
355 crop residue retention (P<0.001), particularly in upland China (Liu et al., 2014). Qiu  
356 et al. (2009) estimated that for every increase of 1 g C kg<sup>-1</sup> of SOC in Huantai county,  
357 grain yield could increase by 454 kg ha<sup>-1</sup>. In Northern China, beneficial effects of crop  
358 yield increase will be enhanced as the strategy of returning crop residues is extended

359 to other regions of Northern China.

360 *Carbon input from crop residues:* Carbon input is one of the most efficient factors  
361 for the accretion of SOC, which is also confirmed by the highly significant correlation  
362 between SOC content and C input from crop residue incorporation ( $r=0.80$ ,  $P<0.0001$ ,  
363 Table 2). Similar results have also been reported by other scientists (Freibauer et al.,  
364 2004). Smith et al. (2005; 2012) reported that input of crop residues could attain the  
365 highest rate of C sequestration ( $0.7 \text{ Mg C ha yr}^{-1}$ ) in comparison with that of merely  
366  $0.2 \text{ Mg C ha}^{-1} \text{ yr}^{-1}$  with the input of mineral N fertilizer. In our study, there are two  
367 SOC increase stages, one early (1987-1992) and one late (2007-2012), which the early  
368 one coincides nicely with increased wheat residue incorporation and the late one with  
369 the increased maize residue incorporation. Again, it highlights the significance of crop  
370 straw input for the building of soil organic matter. In comparison with data from other  
371 countries like India ( $\sim 1 \text{ Mg C ha}^{-1} \text{ yr}^{-1}$ , Srinivasarao et al., 2014) or US ( $\sim 4.2 \text{ Mg C}$   
372  $\text{ha}^{-1} \text{ yr}^{-1}$ , Johnson et al., 2006), the input of residue-C in Northern China is much  
373 higher ( $\sim 8 \text{ Mg C ha}^{-1} \text{ yr}^{-1}$ , Fig. 3c), for maintaining a rapid rate of increase of SOM.  
374 Thus, rate of input of residue-C is the principal determinant of the rate of increase of  
375 SOM: the annual rate of SOM increase being lower in India, intermediate in the U.S.  
376 and high in the North China Plains. Since retention of the entire amount of residues of  
377 maize in North China started from 2007/2008, the accretion of SOC is projected to  
378 continue for another two to three decades until the mid-2040s. It should also be noted  
379 that in our study, soil bulk density from 1983 to 2011 was interpolated and this could  
380 give rise to errors in the SOC storage estimation.

381 Miao et al. (2011) reported that the significance of N fertilizer and crop residues  
382 incorporation to the maintenance and increase of the SOM. Retention of residues  
383 (wheat and maize) in conjunction with appropriate rate of N fertilization have been  
384 properly implemented in Northern China where the antecedent levels SOM are much  
385 lower than those in northern America and western Europe. Therefore, a judicious  
386 continuation of these practices will continue to accumulate SOM for a long time to  
387 come. However, similar trends may not occur under all situations. For example, Khan  
388 et al. (2007) explained that after attaining a steady state, it is unlikely that SOC will

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392 continue to increase, and may even decline with continuous use of synthetic N,  
393 because of the enhanced activities of heterotrophic soil microorganisms in using  
394 crop-derived residues or SOM. In the context of Northern China, therefore, it is likely  
395 that the increase in SOC level will continue because of the improved crop productivity  
396 and retention of crop residues, but it will eventually attain a new equilibrium. Some  
397 other SOC-enhancing practices including application of residue-based animal manure  
398 derived from the same land unit, contribute further to SOC sequestration. In addition,  
399 higher level of mechanization during the agricultural intensification process may  
400 increase the soil BD (about 7% in our study), and this also contributed to the increase  
401 of SOC stock and should not be ignored when quantifying of farmland carbon  
402 sequestration.

403

#### 404 4.2 Comparison of SOC accretion through agricultural intensification in different 405 regions of China

406 Agriculture in China has grown rapidly over the three decades since 1982,  
407 primarily because of the Household Contract Responsibility System and adoption of  
408 the open policy. Further, the intensification process achieved some economic and  
409 environmental benefits (Firbank et al., 2013). In the case of Northern China, however,  
410 there were also problems with increases in soil compaction and water pollution over  
411 the period of three decades.

412 Among all agronomic regions, Northern China registered the highest rate of SOC  
413 sequestration. In contrast, however, the SOC level has declined in Northeast China  
414 since the 1980s (Table 3). Northeast China, one of the few world regions  
415 characterized by the black soil (Phaeozems) and cold climate, has high antecedent  
416 SOC content ( $\sim 20 \text{ g kg}^{-1}$  in uncultivated soils). Thereafter, cultivation and intensive  
417 farming after the 1980s have increased the rate of decomposition of SOM along with  
418 a low input of organic materials and biomass-C. The SOC level in Northern China  
419 was extremely low (5 to  $10 \text{ g kg}^{-1}$ ) in comparison with the soils of Northeast China  
420 and other regions. This trend can be explained by the fact that Northern China has a  
421 long history of low crop productivity, long dry season (Stockmann et al., 2013), sandy

422 soil texture and low input of organic amendments. Application of N fertilizer and  
423 retention of crop residues has increased the SOC level more in dry lands of Northern  
424 China than in those of paddy soils of the lowlands (Liu et al., 2014). Huantai county  
425 and other agronomic regions (except Northeast China) in China also witnessed  
426 increase in SOC level, mainly attributed to the practices of fertilization (synthetic and  
427 organic), increase in crop yield, retention of crop residues, adoption of conservation  
428 tillage and use of organic amendments (Table 3). The 0-30 cm layer of soils of  
429 cropland in China have gained SOC at the rate of  $17\text{--}28 \text{ Tg C yr}^{-1}$  between 1980 and  
430 2000 (Huang et al., 2010), which is similar to the estimates of  $25\text{--}37 \text{ Tg C yr}^{-1}$   
431 reported by Lal (2002).

432 Agricultural intensification in China is an ongoing process, and is progressively  
433 evolving over time. For instance, since 2012, maize residues from some cropland have  
434 been harvested by Huantai farmers for use as cattle feed and the eventual return of the  
435 animal manure to cropland. It is possible that the efficiency of SOM accretion through  
436 animal manure is higher than that of returning maize straw (Wilhelm et al., 2007),  
437 which may result in yet another period of SOM accretion at the regional level with  
438 proper dissemination of this technology (Ladha et al., 2011). Similar to northern  
439 China, other important grain production region like Midwest US also experienced the  
440 stage of SOC accumulation, although the practices (residue management, non- or  
441 reduced tillage and crop rotation) are different (Ogle et al., 2003; Jelinski and  
442 Kucharik, 2009). Adoption of conservation agriculture (NT or minimum tillage) may  
443 be another option for SOC sequestration. However, its applicability and efficiency  
444 need to be validated through long-term research.

445

## 446 5 Conclusions

447 The study of the impact of agricultural intensification on SOC content and stock  
448 was conducted in the Huantai county, which is a representative region in Northern  
449 China. The farmland SOC stock of the whole county increased by more than 50%  
450 over three decades from 1982-2011. Among several improved farming practices,  
451 retention of crop residues strongly contributed to the restoration of SOC, but there

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456 was no synergistic effect between N fertilization rate and crop yield on increase in  
457 SOC. The SOC content decreased with increase in mean annual temperature. The  
458 temporal change in SOC was significantly influenced by the evolution of the practice  
459 of retention of crop residues through implementation of some local farming policies.  
460 The data support the conclusion that agricultural intensification may both increase  
461 crop productivity and enhance some ecosystem services, such as SOC sequestration in  
462 croplands of Northern China. However, current farming practice (e.g., retention of  
463 crop residues) may not always linearly increase SOC over time, indicating a strong  
464 need for a long-term research. Furthermore, there is also a need to explore other  
465 options such as the application of manure through integration of crop and animal  
466 production. Research on the use of animal manure within the region is a priority,  
467 because of its multiple benefits for grain production, the economy and ecosystem  
468 services such as SOC sequestration.

469

#### 470 **Author contribution**

471 Fanqiao Meng, Wenliang Wu and Yan Liao designed the experiments and Yan Liao  
472 carried them out. Yan Liao and Pete Smith performed the calculations and data  
473 analysis. Yan Liao, Fanqiao Meng, Pete Smith and Rattan Lal prepared the  
474 manuscript with contributions from all co-authors.

475

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630  
631

632 Table 1 Correlation analysis between SOC content of cropland and driving factors in  
 633 Huantai county

	SOC	Temperature	Precipitation	Crop yield	N fertilizer rate	C input from crop residues
SOC	1					
Temperature	0.55**	1				
Precipitation	0.30	0.09	1			
Crop yield	0.79**	0.62**	0.30	1		
N fertilizer rate	0.38*	0.01	0.48**	0.55**	1	
C input from crop residues	0.91**	0.73**	0.35	0.89**	0.43*	1

634 \*\* means a highly significant correlation at the level of  $p < 0.01$ ; \* means a significant  
 635 correlation at the level of  $p < 0.05$

636  
 637

638 Table 2 Partial correlation analysis between SOC content of cropland and driving  
 639 factors in Huantai county

Control variables	Partial correlation coefficient between SOC and the driving factor*	
Crop yield, C input from crop residues	SOC vs. Temperature	-0.42 (p=0.03, df=26)
Crop yield, C input from crop residues	SOC vs. N fertilizer rate	0.03 (p=0.86, df=26)
C input from crop residues, N fertilizer rate, Temperature	SOC vs. Crop yield	-0.08 (p=0.70, df=25)
N fertilizer rate, Temperature, crop yield	SOC vs. C input from crop residues	0.80 (p<0.0001, df=25)

640 \*All of the partial correlation coefficients are 2-tailed.

641

642

643 Table 3 Topsoil (0-20 cm) SOC change from 1980s to 2000s in different agronomic  
 644 regions of China\*

Region	Site	Soil type (FAO)	Climate	Period	SOC	SOCD	SOCS	Literature
					%	Mg ha <sup>-1</sup> yr <sup>-1</sup>	Tg yr <sup>-1</sup>	
Our study					41	0.40	0.15	
North China	Quzhou	Fluvisols	Warm, semi-arid, temperate monsoon	1980-2000	31			Liu, et al., 2005 M 1/23/2015 15:59 已删除: ~
	Daxing	Fluvisols	Warm, semi-humid, temperate monsoon	1981-2000	33			Hu, et al., 2006 M 1/23/2015 15:59 已删除: ~
	Hebei	Calcaric Cambisol	Warm, semi-arid, temperate monsoon	1984-2004		0.34	2.2	Xi, et al., 2013 M 1/23/2015 15:59 已删除: ~
	Henan	Fluvisols	Warm, semi-humid, temperate monsoon	1984-2004		0.32	2.8	Xi, et al., 2013 M 1/23/2015 15:59 已删除: ~
Northeast China	Luancheng	Calcaric Cambisol	Warm, semi-arid, temperate monsoon	1979-2000	38			Zhang et al., 2003 M 1/23/2015 16:03 已删除: ~ M 1/23/2015 16:01 已删除: 0.65 M 1/23/2015 16:02 已删除: ~ M 1/23/2015 16:02 已删除: ~
	Liaoning	Chernozems	Semi-humid, temperate continental monsoon	1984-2004		-0.57	-2.9	Xi, et al., 2013 M 1/23/2015 16:02 已删除: ~ M 1/23/2015 16:02 已删除: ~
	Jilin	Albic luvisols	Semi-humid temperate monsoon	1985-2005		-0.81	-7.2	Xi, et al., 2013 M 1/23/2015 16:01 已删除: 1984-2004 M 1/23/2015 16:02 已删除: ~
	Heilongjiang	Phaeozems	Cold temperate monsoon	1986-2006		-0.70	-5.5	Xi, et al., 2013 M 1/23/2015 16:02 已删除: ~ M 1/23/2015 15:59 已删除: ~
Northwest China	Heilongjiang	Phaeozems	Cold temperate monsoon	1982-2002	-14			Yu et al., 2013 M 1/23/2015 15:59 已删除: ~
	Yining	Calcistoll	Arid temperate continental	1981-2001	-9.3			Hou et al., 2013 M 1/23/2015 15:59 已删除: ~
	Akesu	Calcaric	Arid, warm	1982-2001	10			Li et al., 2002

		fluvisols	temperate continental monsoon						
	Huangshui	Eutric cambisols	Arid, warm temperate continental monsoon	1981-2001	28				Chen et al, 2003
	Zhangye	Eutric cambisols	Arid, warm temperate continental monsoon	1982-2003		0.2			Zhang et al., 2009
East China	Jiangsu	Fluvisols/cambisols	Northern subtropical humid monsoon	1980-2000	34				Yu et al., 2003
	Jinhua	Haplic alisols	subtropical humid monsoon	1982-2002	2.6				Xie et al., 2003
	Xuzhou	Fluvisols	Warm, semi-humid, temperate monsoon	1981-2001		0.1~0.5			Zhang et al., 2009
Central China	Hunan	Haplic alisols	Humid, mid-subtropical monsoon	1984-2004		0.17	0.6		Xi, et al., 2013
	Hubei	Eutric cambisols	Humid, mid-subtropical monsoon	1984-2004		0.31	1.2		Xi, et al., 2013
	Zhengzhou	Fluvisols	Warm, semi-humid, temperate monsoon	1982-2003	35				Fu et al., 2004
	Taoyuan	Haplic alisols	Humid, mid-subtropical monsoon	1979-2003	32				Liu et al., 2006
	Jiangnan	Eutric cambisols	Humid, mid-subtropical monsoon	1984-2004		0.62			Xi et al., 2013
	Hunan	Haplic alisols	Humid, mid-subtropical monsoon	1984-2004		0.35			Xi et al., 2013
Southern China	Hainan	Ferralic cambisols	Tropical monsoon	1984-2004		-0.16	-0.5		Xi, et al., 2013

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Binyang	Haplic acrisols	Tropical monsoon	1981-2001	19	Liu et al., 2006
Pixian	Haplic acrisols	Sub-tropical monsoon	1981-2002	9.0	Wei et al., 2004
Yucheng	Haplic acrisols	Sub-tropical monsoon	1981-2002	19	Hu et al., 2004
Meitan	Haplic acrisols	Sub-tropical monsoon	1980-2001	67	Ding et al., 2002

665 \*only the literatures with more than 20 years of experimental interval and more than 30  
666 observation/sampling points were included.

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678 Figure captions

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680 Figure 1 Evolution of SOC content and density for cropland from 1982 to 2011 in  
681 Huantai County. Error bars are highlighted only for the years which soil survey is  
682 undertaken.

683

684 Figure 2 Change of SOC stock for farmland (cropland, vegetable land, farmland  
685 converted to construction land) from 1982 to 2011 in Huantai County.

686

687 Figure 3 Regression analysis of driving factors with the years in Huantai. a)  
688 temperature vs. year, b) yield vs. year, and c) carbon input vs. year.





