

## ***Interactive comment on “Glacial meltwater and primary production as drivers for strong CO<sub>2</sub> uptake in fjord and coastal waters adjacent to the Greenland Ice Sheet” by L. Meire et al.***

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### GENERAL COMMENTS

The authors report a nice data-set of DIC, TA and pCO<sub>2</sub> in a Greenland fjord that covers seasonality, and show that this system acts as a sink for atmospheric CO<sub>2</sub>. The drivers of the CO<sub>2</sub> dynamics are investigated with a model and the authors conclude that part of the CO<sub>2</sub> sink is related to mixing of meltwater and seawater (a purely thermodynamic effect) but that the larger part is related to NEP (although this is somewhat understated, while the effect of meltwater could be overstated).

### MAJOR COMMENTS

In their model, the authors used data of TA and DIC measured on melted glacier ice as  
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the freshwater (FW) endmember. There are some problems in this approach :

\* As stated in page 17929 only 60% of the freshwater relates to glacier meltwater and 34% relates to surface runoff. It is unlikely that TA and DIC values for surface runoff are the same as those from the melted glacier ice. Hence, the model should have used two FW end-members.

\* 6% of the freshwater comes from precipitation. Is this precipitation directly on the fjord surface ? or is this precipitation on the watershed ? How can it be distinguished from surface runoff ? Isn't possible to account this in model (adding an addition flux in the mass balance for direct rain on the fjord surface) ?

\* Shouldn't groundwater inputs somehow also be included in the model ?

\* The glacier ice does not simply melt and mix with seawater. There's in fact some quite complex water infiltration dynamics (e.g. Olichwer et al. 2012) and weathering below the glacier (Wadham et al. 2010; Graly et al. 2014), so that the FW end-members of TA and DIC are in really different from those used in the model. This in fact appears somehow in the Y-intercept of the linear regressions of DIC and TA vs salinity in the fjord (page 17938) that are different from the FW values (50 vs 159 for TA and 80 vs 61 for DIC). In the end, the CO<sub>2</sub> content of Greenland “glaciers” is in fact quite high (Ryu & Jacobson 2012) and not below saturation as stated in page 17942.

P 17935 L 11 : so no seasonality was accounted for the SW end-member ? I expect that in such a shallow system DIC increases in bottom shelf waters after the bloom organic matter sinks.

I would also expect an enrichment of bottom water DIC in the fjord from remineralization of organic matter that sediments from the surface to the bottom of the fjord. So that DIC concentration transported from bottom to surface (F flux) should be different in each box and increase from Zone 1 to Zone 3.

Is there a possibility that coccolithophores also bloomed in the studied fjord ? Coc-

colithophores are common in high-latitude fjords (Schei 1975; Kristiansen et al. 1994) and this affects CO<sub>2</sub> dynamics (Purdie & Finch, 1994).

I suggest that the authors bank their raw data at CDIAC and mention this at the end of the material & methods section.

P 17943 : The effect of sea-ice meltwater on CO<sub>2</sub> dynamics has been shown before in several studies (I picked a recent one by Bates et al. 2014 but there are others), so this is not entirely a complete conceptual revolution of CO<sub>2</sub> ocean dynamics, as suggested. Although this might be the first time it is shown for glacier meltwater in a fjord.

Did the surface freeze at some point in the fjord during the annual cycle (sea-ice formation) ? If so, shouldn't this be somehow accounted in the mass balance and the air-sea CO<sub>2</sub> fluxes ?

#### SPECIFIC COMMENTS

P 17927 L 7 : add the months and years of the cruises.

P 17927 L 20 : is this the global uptake in oceans or oceans+land ?

P 17928 L 5 : might be useful to cite some papers that show this fingerprint on other biogeochemical variables in principle more studied than CO<sub>2</sub>. Salinity, Nutrients, chlorophyll-a.

P 17928 + 17941 : There's a mix of units that should all be uniform : tC/month/km<sup>2</sup>, gC/m<sup>2</sup>/yr, gC/m<sup>2</sup>/d

P 17928 L 11 : "exact mechanisms (...) not well understood". Please rephrase: ocean CO<sub>2</sub> dynamics isn't rocket science or quantum physics. There are a handful of thermodynamic and biological processes that control CO<sub>2</sub> in the oceans. Whether they've been evaluated and quantified in every single part of the ocean is another matter, and indeed there's some work left.

P 17928 : defined DIC and pCO<sub>2</sub> abbreviations first time they're used.

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P 17930 : please provide some information on the precision and accuracy of the pCO<sub>2</sub> measurements, and comment on the calibration of the instrument (how and how often ?). Paper by Fietzek et al. 2014 might be useful.

P 17930 : define SD

P 17931 : define ICES

P 17931 L 23 : in the dark ?

P 17931 L 23 : Incubations are quite long. Is this standard procedure ? I thought incubations for bacterial production were shorter, e.g. 2 h

P 17932 L 10 : There's increasing evidence that the Wanninkhof & McGillis (1999) parameterization provides over-estimates at high wind speeds, e.g. Ho et al. (2006). This is likely due to strong errors at high wind speeds when the eddy-covariance apparatus is thrown around with ship movement. The use of the Wanninkhof & McGillis (1999) parameterization is a problem in the present study since the fluxes computed at wind speeds of 30 m/s (as stated) will be sky high.

P 17934 L 2 : this equation should be added in Table 1. As it stands, in table 1, there are 7 unknowns (Q1, Q2, Q3, F1, F2, F3, Qg) for 6 equations which is puzzling if you do not carefully read all the M&M, and you'll miss how this was solved.

P 17937 : "strikingly" is a self-evaluation. Let the reader decide what' striking in your work.

P 17937 L 5 : Here and in numerous other places in the ms (17939, 17942, 17943, 17944). The correct phrasing is : "surface waters are undersaturated in CO<sub>2</sub> with respect to atmospheric equilibrium" or "pCO<sub>2</sub> in surface waters is below atmospheric pCO<sub>2</sub>". The mix of both "pCO<sub>2</sub> undersaturation" or "pCO<sub>2</sub> is undersaturated" is not correct.

P 17930 L L 7 : Are these relations for all of the data (all stations and all cruises) ?

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if not please specify. Also, these plots might be included in the paper. Finally, isn't the fact that DIC is conservative contradictory with the fact that NEP seems to control DIC dynamic (Fig. 10) and lead to large temporal changes of DIC (Fig. 10) ? This is probably related to the large salinity variation that obscures the NEP effect on DIC. But some rewording could be useful. Plotting DIC & TA versus Salinity will help to discuss this.

P 17942 : In page 17935 TAFW is 50, and here it's 54  $\mu\text{M}$

P 17943 : One way to also demonstrate this is to plot DIC, TA and pCO<sub>2</sub> in surface waters as a function salinity, and compute the theoretical pCO<sub>2</sub> change along the salinity gradient using the two extreme salinity values (DIC and TA end-members) (this gives the the thermodynamic effect of salinity change on constants and on pCO<sub>2</sub> in a closed system, e.i. no air-sea CO<sub>2</sub> exchange).

L 17944 : At such low salinities (8) the choice of dissociation constants becomes a problem, as well as the total boron value (computed from salinity), so this should be detailed in the Material and methods.

#### REFERENCES

Bates et al. 2014 Sea-ice melt CO<sub>2</sub>-carbonate chemistry in the western Arctic Ocean: meltwater contributions to air-sea CO<sub>2</sub> gas exchange, mixed-layer properties and rates of net community production under sea ice, *Biogeosciences*, 11, 6769-6789, [www.biogeosciences.net/11/6769/2014/](http://www.biogeosciences.net/11/6769/2014/) doi:10.5194/bg-11-6769-2014

Fietzek et al. 2014: In situ Quality Assessment of a Novel Underwater pCO<sub>2</sub> Sensor Based on Membrane Equilibration and NDIR Spectrometry. *J. Atmos. Oceanic Technol.*, 31, 181-196.

Graly et al. 2014 Chemical weathering under the Greenland Ice Sheet, *Geology*, 42, 551-554

Ho et al. (2006), Measurements of air-sea gas exchange at high wind speeds in the  
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Southern Ocean: Implications for global parameterizations, *Geophys. Res. Lett.*, 33, L16611, doi:10.1029/2006GL026817

Kristiansen et al. 1994. An *Emiliana huxleyi* dominated bloom in-Samnangerfjorden, Western Norway: importance of hydrography and nutrients. *Sarsia* 79, 357-368.

Olichwer et al. 2012 Chemical composition of groundwaters in the Horsund region, southern Spitsbergen, *Hydrology Research*, doi:10.2166/nh.2012.075

Purdie, D. A. and Finch, M. S.: Impact of a coccolithorid bloom on dissolved carbon dioxide in sea water enclosures in a Norwegian fjord, *Sarsia*, 79, doi:10.1080/00364827.1994.10413569, 1994.

Ryu & Jacobson 2012 , CO<sub>2</sub> evasion from the Greenland Ice Sheet: A new carbon-climate feedback, *Chemical Geology*, Volumes 320-321, Pages 80-95

Schei, B. 1975. Coccolithophorid distribution and ecology in coastal waters of North Norway. *Norwegian Journal of Botany*, 22, 217-225

Wadham et al. 2010 Biogeochemical weathering under ice: Size matters, *Global Biogeochem. Cycles*, 24, GB3025, doi:10.1029/2009GB003688

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