Estimates of micro-, nano-, and picoplankton contributions to particle export in the northeast Pacific

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Abstract

The contributions of micro-, nano-, and picoplankton to particle export were estimated from measurements of size-fractionated particulate ^{234}Th, organic carbon, and phytoplankton indicator pigments obtained during five cruises between 2010 and 2012 along Line P in the subarctic northeast Pacific Ocean. Sinking fluxes of particulate organic carbon (POC) and indicator pigments were calculated from ^{234}Th–^{238}U disequilibria and, during two cruises, measured by sediment trap at Ocean Station Papa. POC fluxes at 100 m ranged from 0.65 – 7.95 mmol m^{-2} d^{-1}, similar in magnitude to previous results at Line P. Microplankton pigments dominate indicator pigment fluxes (averaging 69±19% of total pigment flux), while nanoplankton pigments comprised the majority of pigment standing stocks (averaging 64±23% of total pigment standing stock). Indicator pigment loss rates (the ratio of pigment export flux to pigment standing stock) point to preferential export of larger microplankton relative to smaller nano- and picoplankton. However, indicator pigments do not quantitatively trace particle export resulting from zooplankton grazing, which may be an important pathway for the export of small phytoplankton. These results have important implications for understanding the magnitude and

mechanisms controlling the biological pump at Line P in particular, and more generally in oligotrophic gyres and high-nutrient, low-chlorophyll regions where small phytoplankton represent a major component of the autotrophic community.

1 Introduction

Phytoplankton community structure exerts an important influence on the strength and efficiency of the biological pump (Michaels and Silver, 1988; Boyd and Newton, 1999; Thibault et al., 1999; Brew et al., 2009; Lomas and Moran, 2011). Small nano- and picoplankton dominate the phytoplankton community in the oligotrophic gyres and high-nutrient, low-chlorophyll (HNLC) oceanographic regions. It has traditionally been thought that small phytoplankton represent a relatively small fraction of the downward flux of particulate organic carbon (POC) relative to larger phytoplankton, such as diatoms, which are generally thought to contribute disproportionately to POC export (e.g., Michaels and Silver, 1988). Recent studies have challenged this idea, suggesting that small phytoplankton contribute significantly to POC export, possibly through aggregation and incorporation into fecal pellets (Richardson and Jackson, 2007; Amacher et al., 2009; Stukel and Landry, 2010; Lomas and Moran, 2011; Stukel et al., 2013). A better understanding of the controls on the relative importance of small phytoplankton in POC export is needed to refine our understanding of the magnitude and mechanisms controlling the biological pump, particularly as recent climate models predict an expansion of the oligotrophic gyres where small cells dominate (Irwin et al., 2006; Polovina et al., 2008; Morán et al., 2010).

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Ocean Station Papa (OSP, 50°N, 145°W), the site of one of the longest-running ocean time-series, is located in the northeast Pacific Ocean in one of three major HNLC regions. Previous attempts to resolve the apparent paradox of low phytoplankton biomass and high nitrate concentrations at OSP concluded that a bottom–up control related to iron limitation is most important for large phytoplankton (Muggli et al., 1996; Harrison, 2006; Marchetti et al., 2006), while microzooplankton grazing exerts a strong top–down control on pico- and nanoplankton (Landry et al., 1993; Harrison et al., 1999; Rivkin et al., 1999). Primary production at the stations proximal to the coast on Line P (P4 & P12) is not iron-limited and diatom blooms are typically observed in spring and late summer (Boyd and Harrison, 1999; Thibault et al., 1999). At the offshore stations (including OSP) the phytoplankton community is dominated by cells <5-

µm and the seasonal variability of primary production is relatively low (~25 mmol C m^{-2} d^{-1} in winter and ~67 mmol C m^{-2} d^{-1} in summer) (Boyd and Harrison, 1999; Thibault et al., 1999; Choi et al., 2014). In contrast to the low variability in primary production, POC export recorded by moored sediment traps at OSP exhibits a stronger seasonal cycle with fluxes at 200 m depth ranging from ~0.4 mmol C m^{-2} d^{-1} in winter to ~2.4 mmol C m^{-2} d^{-1} in summer (Timothy et al., 2013). The average annual sediment trap POC flux at OSP (1.4 ± 1.1 mmol C m^{-2} d^{-1}) is nearly five times lower than the annual net community production (ANCP) at OSP (6.3 ± 1.6 mmol C m^{-2} d^{-1}), suggesting that the majority of organic carbon export is due to active transport by zooplankton and/or dissolved organic carbon (DOC) export (Timothy et al., 2013; Emerson, 2014).

This study builds upon prior investigations of phytoplankton community composition and export production along Line P by examining the distributions of organic carbon, phytoplankton indicator pigments, and ^{234}Th in three particle size-fractions. Sinking fluxes of POC and indicator pigments from the upper waters (~100 m) were calculated from the ^{234}Th–^{238}U disequilibrium and, during two cruises, measured at OSP using free-floating sediment traps. A comparison of indicator pigment fluxes with the respective standing stocks suggests that microplankton (20 – 200-µm) make up a higher percentage of particle export than biomass, whereas pico- and nano plankton (0.2 – 2-µm and 2 – 20-µm) make up a lower percentage of particle export than biomass.

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2 Methods

2.1 Study location

Sample collection was conducted at five stations along Line P (P4, P12, P16, P20, and P26 (OSP)) during cruises aboard the *CCGS John P. Tully* in August 2010, February 2011, June 2011, February 2012, and June 2012 (Fig. 1, Table 1). Line P is located at the southern edge of the Alaskan Gyre, and the prevailing winds and surface currents are west-east (Bograd et al., 1999). Because precipitation and continental run-off exceed evaporation, a permanent halocline exists at ~100 m impeding deep winter mixing. In addition, a seasonal thermocline forms at ~50 m in spring and shoals to ~20 m in summer (Freeland et al., 1997; Thibault et al., 1999; Freeland, 2013; Timothy et al., 2013).

2.2 Net primary production by ^{14}C incubation

Rates of net primary production (NPP) were determined following the protocols outlined in Lomas et al. (2012). Samples were collected with Niskin bottles from seven depths in the euphotic zone corresponding to 1, 5, 9, 17, 33, 55, and 100% of surface irradiance. Three 'light' bottles, a single 'dark' bottle, and a single initial (T_0) bottle were each spiked with ~10 µCi $NaH^{14}CO_3$. A sub-sample to confirm total added activity was removed from the T_0 bottle at each light depth and immediately added to an equal volume of β-phenylethylamine. Bottles were incubated under simulated in situ conditions, using neutral density screening to mimic light levels at the depth of sample collection, in an on-deck incubator for ~24 hours. After incubation, 125 mL sub-samples from each light and dark bottle were filtered through an Ahlstrom 151 (0.7-µm nominal pore size) and a Whatman Track Etch 5-µm filter and rinsed with 10% HCl. Samples were counted on a Perkin Elmer TriCarb 2900LR ~48 h after the addition of 5 mL of Ultima Gold (Perkin Elmer, USA) scintillation cocktail.

2.3 Water column ^{234}Th

Total ^{234}Th (dissolved + particulate) analysis followed the procedures outlined in Bauman et al. (2013). Briefly, samples (4 L) were collected by Niskin bottle at 12 depths (surface to ~500 m) and spiked with ^{230}Th to monitor Th recovery. Samples were then treated with 7-8 drops of concentrated NH_4OH solution, followed by 25 µL of 0.2 M $KMnO_4$, and finally with 11.5 µL of 1.0 M $MnCl_2$ to form a MnO_2 precipitate that quantitatively scavenges Th (Benitez-Nelson et al., 2001; Buesseler et al., 2001; van der Loeff et al., 2006). After 1 hour, samples were vacuum filtered onto 25 mm glass microfiber filters (GM/F, 1-µm nominal pore size) that were frozen for later analysis in the shore-based laboratory. To prepare samples for counting, filters were dried at 50°C for ~24 hours, mounted on acrylic planchets, and covered with aluminum foil. To quantify ^{234}Th, the beta emission of ^{234m}Pa (E_{max} = 2.19 MeV; $t_{1/2}$ = 1.2 min) was counted using a RISØ National Laboratory low-background beta detector (Roskilde, Denmark). Each sample was counted four times over a period of approximately six half-lives, with the first count made at least 10 days after collection to allow for the decay of short-lived isotopes, and the final count used to quantify background levels. Data were fitted to the ^{234}Th decay curve to calculate the decay-corrected activity at the time of sample collection. Following the ^{234}Th analysis, Th was radiochemically purified and ^{230}Th was measured by alpha particle

emission in order to determine scavenging efficiency. Small-volume scavenging efficiencies were found to be >90%. ^{238}U activities were calculated from salinity using the relationship ^{238}U = 0.07081 x S (‰) (Chen et al., 1986)

2.4 Water column POC, Chl a, and indicator pigments

Water samples for POC, Chl *a*, and phytoplankton indicator pigments were collected from the same depths in the photic zone as for NPP samples. Suspended POC was measured on 1 L seawater samples filtered onto pre-combusted Ahlstrom 151 filters and frozen at -20°C until analysis. Samples were dried at 60°C in a drying oven, fumed in a desiccator containing concentrated hydrochloric acid for 24 h to remove inorganic carbonates, and dried again at 60°C. Samples were then analyzed on an EA-440 Analyzer (Exeter Analytical, Inc., Chelmsford, MA) (Pike and Moran, 1997). Chl *a* samples were analyzed using the methods outlined in Lomas et al. (2012). Separate samples (~0.2 L) were filtered onto Ahlstrom 151 and 5-µm Whatman Track Etch polycarbonate filters and frozen at -20°C until analysis. Samples were then extracted in 5 mL of 90% acetone for 24 h at -20°C and analyzed using a calibrated TD-700 fluorometer.

Indicator pigment samples were collected on separate Ahlstrom 151 filters and stored at -80°C until analysis by high-performance liquid chromatography (HPLC) at the Bermuda Institute of Ocean Sciences in the Bermuda Atlantic Time-series Study Laboratory (Knap et al., 1997). Fucoxanthin (FUCO), peridinin (PER), 19'-hexanoyloxyfucoxanthin (HEX), 19'-butanoyloxyfucoxanthin (BUT), alloxanthin (ALLO), total chlorophyll *b* (TChl *b*), and zeaxanthin (ZEA) were analyzed as indicator pigments based on their correspondence to particular phytoplankton taxonomic groups. Indicator proportion factors (PFs) were calculated to further analyze the size-distribution of the phytoplankton community (Hooker et al., 2005; Lomas and Moran, 2011). The sum of FUCO and PER concentrations was used to determine the microplankton proportion factor (mPF), while the sum of HEX, BUT, ALLO, and TChl *b* was used to determine the nanoplankton proportion factor (nPF), and ZEA was used to determine the picoplankton proportion factor (pPF) (Hooker et al., 2005; Lomas and Moran, 2011). Hooker et al. (2005) included TChl *b* in pPF, but because *Prochlorococcus* is not found in the study region, it was assumed in this study that any Chl *b* would be found in cells (e.g., chlorophytes and euglenophytes) in the nanoplankton size-class.

159 **2.5 In situ pump sampling**

160 Large-volume in situ pumps (Challenger Oceanic Systems and Services, UK and McLane
161 Scientific, Falmouth, MA) were deployed for approximately four hours at depths of 30, 50, 100,
162 150, and 200 m. Each pump sampled 100 – 1000 liters to collect size-fractionated particles, with
163 seawater passing sequentially through 53-µm, 10-µm, and 1-µm Nitex screens. Particles were
164 resuspended by ultrasonication in 0.7-µm prefiltered seawater and filtered onto separate pre-
165 combusted GF/F filters for parallel analysis. Indicator pigment samples were stored at -80°C
166 until analysis by high-performance liquid chromatography (HPLC) at the Bermuda Institute of
167 Ocean Sciences in the Bermuda Atlantic Time-series Study Laboratory (Knap et al., 1997).
168 Filters for analysis of POC and ^{234}Th were frozen at -20°C until analysis. A sub-sample (~30%
169 by weight) was cut with acetone-cleaned stainless steel scissors from each ^{234}Th filter for POC
170 analysis, and these sub-samples were dried and fumed with concentrated HCl as described
171 above. POC was then measured using a CE 440 CHN Elemental Analyzer (Exeter Analytical,
172 Inc., Chelmsford, MA). The ^{234}Th filter subsample was dried at 60°C in a drying oven and
173 counted on a RISØ beta detector as noted above.
174
175 **2.6 Sediment trap sampling**

176 Surface-tethered particle interceptor traps (PITS) with cylindrical tubes (KC-Denmark,
177 Silkeborg, Denmark) were deployed for ~3 days at station P26 during the June 2011 and June
178 2012 cruises to collect particles at the depths of 30, 50, 100, 150, and 200 m. Due to limited
179 wire-time and other cruise constraints it was not possible to deploy sediment traps at any other
180 stations sampled as part of this study. The trap design and sampling procedure is described in
181 Baumann et al. (2012). Four tubes (72 mm diameter, 450 mm length) were used at each depth,
182 and tubes were filled with non-poisoned, 0.4-µm filtered brine (S = ~85 ‰) prior to deployment.
183 Upon recovery trap brines were combined, particles were re-suspended and filtered onto pre-
184 combusted GF/F filters, and swimmers were removed. Filters were stored frozen and later
185 analyzed for POC, ^{234}Th, and indicator pigments as described above.
186
187 **3 Results**

188 **3.1 Hydrography and NPP**

Depth sections of temperature and density anomaly (sigma-t) were generated using results from all CTD casts for a given cruise to improve horizontal data resolution (Fig. 2). The seasonal change in water temperature is largely confined to the upper ~100 m. Surface temperatures in August 2010 were ~14°C, while during the February cruises, surface temperatures were slightly cooler offshore (~6°C) than inshore (~8°C). During the June cruises, inshore temperatures were warmer (~10 – 12°C) while offshore temperatures remained relatively cool (~8°C). Density anomaly did not vary greatly between cruises below ~100 m. During the winter, a pool of less dense water (density of 1023 – 1025 kg m^{-3}) was observed toward the coast (east of ~126°W). During the June cruises, this pool was observed extending west to ~130°W and during August 2010, it extended out to OSP (145°W). These data follow the expected seasonal pattern of a well-mixed water column in winter and increasing stratification moving from spring to summer.

Total NPP and >5-µm size-fractionated NPP values were trapezoidally integrated over the euphotic zone (Table 2). A maximum total NPP of 91.9 mmol m^{-2} d^{-1} was measured at station P26 during June 2011, whereas the lowest value of 12.4 mmol m^{-2} d^{-1} was measured at station P26 during February 2012. These values agree to within a factor of two with the seasonal averages reported by Boyd and Harrison (1999). A maximum >5-µm NPP of 39.6 mmol m^{-2} d^{-1} was at station P4 during June 2012 and a minimum of 2.2 mmol m^{-2} d^{-1} was measured at station P12 in February 2012.

3.2 Small- and large-volume POC concentrations

Suspended POC concentrations from Niskin bottle samples collected in the photic zone range from 1.1 – 7.1 µmol L^{-1}. POC concentrations were generally lowest at the base of the photic zone, though decreasing concentrations with depth were not observed at all stations (Table S1). The highest suspended POC concentrations were measured at station P4 during all cruises. POC concentrations were also measured in three size-fractions of particles collected with large-volume in situ pumps (Table S2). Concentrations of each size-fraction tended to decrease with depth and were typically less than 0.5 µmol L^{-1} at all depths. One exception was at station P26 during February 2011 when POC concentrations at 30 m were between 1.8 and 2.9 µmol L^{-1} for all size-fractions.

219 The concentrations of POC collected using small-volume and large-volume methods
220 often do not agree for samples collected at the same location and depth (Gardner, 1977; Moran et
221 al., 1999; Liu et al., 2005; Liu et al., 2009). As reported in these previous studies, POC
222 concentrations measured by large-volume in situ pumps (summed for all size-fractions) are
223 significantly (ANOVA, $p < 0.05$) less than small-volume POC measurements from the same
224 station and similar depth (Fig. 3a). Explanations put forth to account for this discrepancy include
225 DOC adsorption to filters, pressure effects on particle retention in pump samples, the collection
226 of zooplankton by Niskin bottles but not pumps, and particle washout from pump filters (Moran
227 et al., 1999; Liu et al., 2005; Liu et al., 2009). In this study, the smallest pump size-fraction
228 was collected using a 1-µm Nitex screen, not a GF/F, resulting in the pumps missing the portion
229 of the POC on particles between 0.7- and 1-µm, which may further contribute to the difference
230 observed between the two methods. Lomas and Moran (2011) reported that sonication of in situ
231 pump samples to resuspend particles from the Nitex screens had no significant effect on
232 measured POC concentrations.
233
234 **3.3 Particulate ^{234}Th and POC/^{234}Th ratios**
235 Size-fractionated particulate ^{234}Th activities in samples collected by in situ pump
236 generally decrease with depth, and are typically less than 0.1 dpm L^{-1} (Table S2). As with in situ
237 pump POC concentrations, station P26 during February 2011 is an exception, with values
238 exceeding 0.1 dpm L^{-1} for all size fractions at 30 m and throughout most of the water column for
239 the 1 – 10-µm fraction. Size-fractionated POC/^{234}Th ratios (Fig. 4, Table S2) are less than ~6
240 µmol dpm^{-1} for all size-classes at most stations, with higher values measured at stations P4 and
241 P12 in February 2012 and P4 in June 2012. POC/^{234}Th ratios tend to decrease or remain constant
242 with depth, with one exception at station P12 during February 2012 where the maximum
243 POC/^{234}Th was at 100 m for all size fractions. Also, the POC/^{234}Th ratio does not vary greatly
244 between size-fractions (Fig. 4) as was observed in Speicher et al. (2006) and Brew et al. (2009).
245 The accuracy of ^{234}Th as a tracer of POC export depends on the assumption that ^{234}Th
246 and POC are sinking on the same particles, and therefore sinking at the same rate (Moran et al.,
247 2003; Smith et al., 2006; Speicher et al., 2006; Burd et al., 2007; Brew et al., 2009). A high
248 degree of correlation between the size-fractionated distributions of ^{234}Th and POC (Fig. 4) along
249 Line P provides evidence in support of this assumption. All correlations were statistically

significant ($p < 0.05$) and imply a strong coupling between particulate ^{234}Th and POC for all cruises. In addition, the clustering of data for the different size-fractions of particles (Fig. 4) indicates that in February 2012 the 10 – 53-µm size class contained the highest percentage of POC and particulate ^{234}Th, while the >53-µm size class contained the lowest percentage. In June 2012, the 1 – 10-µm size class had the lowest percentage of POC and particulate ^{234}Th while both the 10 – 53-µm and the >53-µm fractions contained higher percentages (Fig. 4).

3.4 Small-volume Chl a and indicator pigments

Concentrations of total Chl *a* and >5-µm Chl *a* measured by fluorometer (Table S1) were trapezoidally integrated over the photic zone to determine respective standing stocks. During August 2010, the >5-µm fraction accounted for >30% of the Chl *a* at all stations, with a maximum of 50% at station P26. During the other four cruises, the >5-µm size-fraction generally accounted for <30% of the total Chl *a*, except at station P26 in February 2012 and station P4 in June 2012. Previous studies have reported that larger cells are more abundant at stations closer to the coast (Boyd and Harrison, 1999), though this was not always apparent. The highest >5-µm percentage of Chl *a* was measured at station P26 during August 2010, June 2011, and February 2012. Phytoplankton indicator pigments and Chl *a* concentrations in samples from the euphotic zone samples were also measured by HPLC (Table S1). HPLC and fluorescence Chl *a* concentrations generally agreed to within a factor of two, and the correlation between the two measurements was statistically significant ($p < 0.05$) (Fig. S1). The correlation between the sum of the indicator pigment concentrations and the Chl *a* concentration was statistically significant ($p < 0.05$) and roughly 1:1, suggesting that the indicator pigments examined in this analysis accounted for most of the phytoplankton biomass (Fig. S2). Furthermore, the correlation between the >5-µm fraction of Chl *a* and mPF is statistically significant ($p < 0.05$), suggesting that this PF is a reasonable representation of that size-fraction of the phytoplankton community. Profiles of indicator pigment concentrations were trapezoidally integrated over the photic zone to quantify standing stocks (Table 3). FUCO was the most abundant microplankton pigment, and HEX was the most abundant nanoplankton pigment at most stations. Indicator pigment PFs (Fig. 5, Table S3) reveal that the phytoplankton community was typically dominated by nanoplankton, although at P4, and to a lesser extent at P20 in June 2012, microplankton pigments made up the bulk of the sample (~86% and ~52% respectively).

3.5 Large-volume size-fractionated Chl a and indicator pigments

Size-fractionated Chl *a* and indicator pigment concentrations were also measured by in situ pump (Table S4). Chl *a* was once again strongly correlated in a roughly 1:1 ratio with the sum of the indicator pigments ($p < 0.05$) (Fig. S3). The highest Chl *a* concentrations were measured in the 10 – 53-µm fraction during all cruises. In February 2012, the >53-µm fraction generally had the lowest concentrations, while in June 2012 and June 2011 the lowest concentrations were generally in the 1 – 10-µm fraction.

Ideally, small-volume and large-volume concentrations of Chl *a* and indicator pigments should agree for samples collected at the same station and depth, but this was not observed in this study (Fig. 3). Although differences between small- and large-volume measurements of POC have been reported (Gardner, 1977; Moran et al., 1999; Liu et al., 2005; Liu et al., 2009), few studies have compared Niskin bottle and in situ pump measurements of indicator pigments (Lomas and Moran, 2011). Relative to bottle samples, the pump samples indicate higher concentrations of microplankton pigments FUCO and PER and lower concentrations of ZEA and TChl *b*, which are pigments associated with pico- and nanoplankton (Fig. 3b-d). Large-volume pump and small-volume bottle measurements of the nanoplankton indicator pigments HEX, BUT, and ALLO generally agree within a factor of two (Fig. 3b-d). Given the small size of ZEA-containing *Synechococcus* and TChl *b*-containing chlorophytes and prasinophytes, it is likely that many of these cells pass through the 1-µm Nitex screen which would lead to under-sampling by the pumps (Liu et al., 2005). Bottles may undersample large, rare cells because the small volume might not be a statistically representative sample (Lomas and Moran, 2011). Furthermore, larger cells may settle below the spigot of the Niskin bottles, leading to a further bias against the collection of large cells (Gardner, 1977; Gundersen et al., 2001). Pumps sample higher concentrations of Chl *a* than bottles (Fig. 3a) at stations with high concentrations of Chl *a*, but when Chl *a* concentrations are low (<200 ng L^{-1}), the pumps tend to undersample relative to the bottles.

Given these sampling differences, it is important to note that although the total concentrations (summed for all size-fractions) measured by the in situ pumps may be inaccurate, it is still possible that the >53-µm fraction accurately represents the composition of sinking particles. The disruption of loosely-bound aggregates during collection by the pumps could

cause an error in the >53-μm fraction, but this is considered unlikely due to the presence of nanoplankton (and in some cases picoplankton) pigments in this fraction. Furthermore, a recent study in the Sargasso Sea employed a similar methodology and also found picoplankton pigments in three particle size-classes, each >10-μm (Lomas and Moran, 2011).

Indicator pigment PFs calculated for the size-fractionated particles (Table S3) and plotted against depth (Figs. 6-8) reveal that while the overall indicator pigment concentrations vary with depth and across size-fractions, the PFs do not exhibit a systematic pattern of variation across size classes, depths, or seasons. The picoplankton pigment ZEA typically represents <10% of the total indicator pigments for all size classes. Microplankton pigments dominated samples at station P4 in February 2012 and June 2012, with mPFs typically exceeding 0.5 and 0.8, respectively, for each cruise. In addition, mPFs were high at station P26 during these times, with values generally exceeding 0.5 (Figs. 7-8). Nanoplankton pigments dominated at station P12 in February 2012 cruise with nPFs exceeding 0.5 for most samples. As with the small volume samples, FUCO was usually the most abundant microplankton pigment while HEX was usually the most abundant nanoplankton pigment (Table S4).

3.6 Total ^{234}Th, ^{234}Th/^{238}U activity ratios, and ^{234}Th fluxes

Total (dissolved + particulate) ^{234}Th activities, ^{238}U activities, and ^{234}Th/^{238}U activity ratios are listed in Table S5. Depth sections of these ^{234}Th/^{238}U activity ratios (Fig. 2d) indicate that areas of low ^{234}Th/^{238}U are prevalent in spring and summer and corresponding to periods known to have high particle export in this region (Wong et al., 1999; Timothy et al., 2013). ^{234}Th fluxes (P_{Th}) were calculated using these ^{234}Th/^{238}U results and a 2-D steady-state model of the radiochemical balance for ^{234}Th in the upper ocean,

$$\frac{\partial A_{Th}}{\partial t} = A_U \lambda_{Th} - A_{Th}\lambda_{Th} - P_{Th} + K_h \frac{\partial^2 A_{Th}}{\partial^2 x} + U_h \frac{\partial A_{Th}}{\partial x} \quad (1)$$

where A_U is the activity of ^{238}U, λ_{Th} is the ^{234}Th decay constant, A_{Th} is the activity of ^{234}Th, P_{Th} is the vertical flux of ^{234}Th on sinking particles, K_h is the eddy diffusion coefficient, and U_h is the current velocity (Coale and Bruland, 1985; Charette et al., 1999). Assuming a steady-state

($\partial A_{Th}/\partial t = 0$) over several weeks to months, and that the diffusive flux of ^{234}Th is small relative to advection and can therefore be ignored, the vertical flux of ^{234}Th (in dpm m^{-2} d^{-1}) is defined by,

$$P_{Th} = \int_0^z \left[\lambda_{Th}(A_U - A_{Th}) + U_h \frac{\partial A_{Th}}{\partial x}\right] dz \quad (2)$$

where z is the depth of the water column over which the flux is measured. In this study, the gradient of thorium ($\partial A_{Th}/\partial x$) was only measured in the east-west direction (along Line P). Therefore, x is the east-west distance across which the gradient will be measured and U_h is the east-west current velocity. Current velocities determined from 5-year seasonal averages of surface drifter data (available from Fisheries and Oceans Canada) were found to be 6±4 cm s^{-1} for the February cruises, 4±2 cm s^{-1} for the June cruises, and 5±3 cm s^{-1} for the August cruise. These values agree well with the ~10 cm s^{-1} value reported by McNally, (1981) and used by Charette et al., (1999). Given that the currents in the region generally flow west-east, and with no data at stations north and south of Line P, the north-south transport of ^{234}Th by advection had to be assumed to be negligible. At stations P12, P16, and P20, the ^{234}Th gradient was measured between the adjacent stations. For stations P4 and P26 (at either end of Line P), the gradient of ^{234}Th was determined from the adjacent station assuming a linear change extended beyond the measured transect.

^{234}Th fluxes (P_{Th}) calculated using the 2-D model are within 5% of fluxes determined using a steady-state 1-D model that ignores advection (Fig. S4). This indicates that, under these assumptions, the vertical flux of ^{234}Th on sinking particles is the dominant transport term. Consistent with previous studies, ^{234}Th fluxes at all stations were higher during the August and June cruises than during the February cruises (Fig. 9a) (Charette et al., 1999). Also, ^{234}Th fluxes did not exhibit a consistent trend along Line P.

3.7 ^{234}Th-derived POC fluxes

The POC/^{234}Th ratio in the >53-µm size-class and P_{Th} for a given depth horizon were used to calculate POC fluxes (P_{POC}) (Fig. 9). In most cases, P_{POC} decreases with depth, although in some cases, the maximum P_{POC} in a given profile occurs at 50 or 100 m. P_{POC} fluxes at 100 m range from 0.65 – 7.95 mmol m^{-2} d^{-1}; they are generally higher in summer than winter, and

highest at station P4, consistent with previous studies at Line P (Charette et al., 1999; Wong et al., 1999; Timothy et al., 2013).

The ratio of P_{POC} flux to NPP, referred to as the *ThE*-ratio, is an estimate of efficiency of the biological pump (Buesseler, 1998). *ThE*-ratios determined using P_{POC} fluxes at the base of the photic zone (Table 2, Fig. 10) are similar to those reported by Charette et al. (1999), and are also in line with an annual average e-ratio determined using average sediment trap POC fluxes (Wong et al., 1999) and annual average NPP (Harrison, 2002) (Fig. 10).

3.8 Sediment trap ^{234}Th and POC fluxes

The particle fluxes of both ^{234}Th and POC fluxes determined by the PITS traps (F_{Th} and F_{POC} respectively) generally decrease with depth (Table 4). F_{Th} was higher in June 2012 than in June 2011, though there was no clear difference between the two cruises for F_{POC}. A comparison of the F_{Th} with the P_{Th} from corresponding stations and depths indicates that the F_{Th} is consistently higher than the P_{Th}, though usually not by more than a factor of two. F_{POC} is also consistently higher than P_{POC}, though again not by more than a factor of two (Fig. 11a). The POC/^{234}Th ratios of particles caught in sediment traps (Table 9) tend to be slightly higher (generally within a factor of 2) than the ratio of particles sampled by pumps at the corresponding station and depth.

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3.9 ^{234}Th-derived and sediment trap pigment fluxes

Sinking fluxes of Chl *a* (P_{Chla}) and indicator pigments ($P_{Pigment}$) were calculated from P_{Th} and the Pigment/^{234}Th ratio measured on >53-µm particles. Chl *a* and indicator pigment fluxes (Table 3, Fig. 12a-c) are generally highest at station P4 and decrease moving offshore. The highest indicator pigment fluxes were typically observed for microplankton pigments (FUCO and PER) whereas the lowest were observed for the picoplankton pigment ZEA (Table 3, Fig. 12a-c). It is important to note that the differences between fluxes of different pigments at a given station are determined by the pigment ratio on the >53-µm particles and are independent of P_{Th}.

Sediment trap pigment fluxes ($F_{Pigment}$) were typically lower than $P_{Pigment}$ (Table 3, Fig. 11b). The maximum sediment trap fluxes of Chl *a* and most indicator pigments were determined at 50 m in June 2011 and at 30 m in June 2012 (Table 3). For both deployments the deepest fluxes were generally the lowest, presumably due to the progressive degradation of sinking

404 phytoplankton and resulting loss of pigments. Chl *a* and indicator pigment fluxes were generally
405 higher in June 2011 than in June 2012, which is the opposite of the trend observed for F_{Th}.
406 Pigment PFs determined for material captured by the PITS traps do not vary greatly with
407 depth, suggesting that the quality of material sinking to depth is similar to that in the surface
408 water, despite the general decrease of material (Figs. 6 and 8). Microplankton PFs are higher for
409 trap samples than for bottle samples but not as high as for pump samples, while nPFs and pPFs
410 are higher for trap samples than for pump samples but lower than for bottle samples.
411
412 **4 Discussion**
413 The results presented in this study build on previous investigations of export production
414 in the northeast Pacific by providing estimates of the relative contributions of different
415 phytoplankton size-classes to particle export. A comparison of indicator pigment standing stocks
416 determined from small-volume samples and $P_{Pigment}$ fluxes suggests that while nanoplankton
417 represented the bulk of phytoplankton biomass (68±24% of pigment standing stock, averaged for
418 all stations and cruises), microplankton dominated the flux of pigmented material (69±19% on
419 average) (Table 3, Fig. 12). Sediment trap pigment fluxes indicate a lower, but still substantial,
420 relative contribution of microplankton to export, with microplankton pigments making up 47%
421 and 33% of the total sediment trap indicator pigment flux in June 2011 and June 2012
422 respectively, as compared to 81% and 85% of total $P_{Pigment}$ fluxes. Though nano- and
423 picoplankton did not form the majority of the algal aggregate flux, their 29±19% contribution is
424 significant and similar to contributions reported by Lomas and Moran (2011) for cyanobacteria
425 and nano-eukaryotes in the Sargasso Sea.
426 Indicator pigment loss rates determined from both $P_{Pigment}$ fluxes and sediment trap
427 pigment fluxes imply that microplankton are exported more efficiently than nano- or
428 picoplankton (Table 3, Fig. 12d-f). Loss rates of pigments, estimated as the ratio of $P_{Pigment}$
429 fluxes to pigment standing stock, averaged (for all cruises) 8±12% for microplankton pigments,
430 1±2% for nanoplankton pigments and 0.6±1% for picoplankton pigments. These results suggest
431 that export of large cells by direct sinking of algal aggregates is more efficient than the export of
432 small cells by the same pathway. Sediment trap loss rates for microplankton were also higher
433 than those for nano- and picoplankton, further indicating preferential export of microplankton.
434 Even though differences between bottle and pump samples may exaggerate the extent to which

large cells dominate export, sediment trap loss rates support and confirm the preferential export of large cells by algal aggregation.

In contrast to the trends observed for pigment fluxes and loss rates, the low variability of pump indicator pigment PFs with depth (Figs. 6-8) does not appear to indicate preferential export of microplankton. Furthermore, the presence of nano- and picoplankton pigments in the >53-µm size-fraction and in samples below the mixed layer suggests that nano- and picoplankton are incorporated into aggregates and that some of these aggregates are exported from the surface ocean. If large cells were being preferentially exported, microplankton pigments would be expected to make up a larger percentage of total pigments in samples below the mixed layer than in samples from the mixed layer, but this is not observed in the results of this study. It is possible that some of this discrepancy can be attributed to differences between bottle and pump samples. Because cells <1-µm in size can pass through the 1-µm Nitex screens used in the pumps, the sum of the pump size-fractions does not accurately reflect the community composition in the euphotic zone, and may miss a change in indicator pigment PFs with depth. In addition, the under-sampling of large cells by Niskin bottles may lead to an underestimate of microplankton standing stocks, and thus and overestimate of microplankton loss rates.

These pigment fluxes are likely lower estimates of the total contribution of each phytoplankton group to particle export. The use of indicator pigments as tracers of phytoplankton export only accounts for the direct sinking of healthy, ungrazed cells, because grazing degrades the indicator pigments to an analytically undetectable form (Head and Harris, 1992; Strom et al., 1998; Thibault et al., 1999). Indirect export (via grazing) is thought to be an important pathway for picoplankton export in the HNLC Equatorial Pacific (Richardson et al., 2004; Stukel and Landry, 2010). Given that grazing has been shown to control the biomass of small phytoplankton in the northeast Pacific (Landry et al., 1993; Harrison et al., 1999; Rivkin et al., 1999), indirect export may also be a significant pathway for small cell export in this region. Because this pathway is not accounted for by the methodology employed in this study, the results presented here may underestimate the export of small phytoplankton, which may be less likely to sink directly.

Although grazing and fecal pellet export were not directly measured in this study, a comparison of sediment trap and pump measurements of Chl *a*, indicator pigments, and POC, suggests that zooplankton fecal pellets may be an important component of POC export at OSP, at

least in spring (Fig. 11). While F_{POC} fluxes are higher than the corresponding P_{POC} fluxes, $F_{Pigment}$ fluxes are lower than $P_{Pigment}$ fluxes, indicating that the material captured by the sediment traps is enriched in carbon and depleted in Chl *a* and indicator pigments relative to that sampled by the pumps. Because the trap brine was not poisoned, zooplankton grazing and cell degradation in the trap tube may also have contributed to some loss of pigments over the ~3 day deployment of the PITS traps. However, the collection of carbon-rich and pigment-depleted fecal pellets by the traps but not by the pumps, which do not quantitatively sample fecal pellets (Lomas and Moran, 2011), could also explain these observations. This latter explanation is consistent with the results presented in Thibault et al. (1999), which indicate that fecal pellet export is 3 to 6 times greater than algal aggregate export at Line P.

5 Conclusions

New estimates of phytoplankton indicator pigment loss rates calculated from both ^{234}Th-derived and sediment trap pigment fluxes suggest that large cells are preferentially exported at Line P. Specifically, microplankton pigments on average made up 69±19% of the total pigment flux, but only 32±24% of pigment standing stock (determined from small-volume samples), whereas nano- and picoplankton pigments on average formed 31±19% of pigment flux in spite of representing 68±24% of the standing stock. These results are consistent with traditional food web models (Michaels and Silver, 1988; Legendre and Le Fèvre, 1995) that suggest nano- and picoplankton are underrepresented in particle flux relative to their contribution to phytoplankton biomass; they also lend support to the conclusions of Choi et al. (2014). However, the methods employed in this study do not quantitatively account for export via zooplankton fecal pellets, which could be significant for small phytoplankton as they are controlled by grazing in this region (Landry et al., 1993; Harrison et al., 1999; Rivkin et al., 1999; Thibault et al., 1999). Furthermore, the determination of pigment loss rates also required a comparison between small- and large-volume samples, and the inherent differences of these sampling techniques likely led to an overestimation of the microplankton contribution to algal aggregate export. Therefore, it is possible that all sizes-classes of phytoplankton contribute to POC export in approximate proportion to their contribution to NPP as predicted by Richardson and Jackson (2007).

This study, conducted in a subarctic HNLC region, contributes to the ongoing discussion of small cell export that has largely focused on tropical and subtropical regions (Richardson et

al., 2004; Richardson et al., 2006; Richardson and Jackson, 2007; Stukel and Landry, 2010; Lomas and Moran, 2011). In particular, these results suggest that nano- and picoplankton may contribute significantly to POC export in this subarctic HNLC region, even if they are not as efficiently exported as larger microplankton. If large phytoplankton drive more efficient POC export in the northeast Pacific as suggested by this study, it could have important implications for understanding the biological pump. It has been proposed that decreasing winter mixed layer depths (Freeland et al., 1997; Freeland, 2013) and variations of macronutrient concentrations linked to shifts in climate regime (Pena and Varela, 2007) in the northeast Pacific could lead to shifts in the phytoplankton community composition. This study suggests that such changes in phytoplankton community composition could significantly affect the efficiency of the biological pump, and in turn, the cycling of carbon. While the results indicate that shifts in community composition favoring larger phytoplankton could lead to more efficient particle export, they do not indicate that shifts favoring smaller phytoplankton would lead to a shutdown of POC export as suggested by some previous studies (e.g., Michaels and Silver, 1988), but merely that the export of POC could be less efficient.

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Table 1: Cruise dates and sample collection along Line P

Cruise Dates	P4	P12	P16	P20	P26
2010-14	Total Th	Total Th	Total Th	Total Th	Total Th

Aug. 2010 (8/19/10 - 8/31/10)		WC Pig	WC Pig	WC Pig	
2011-01 Feb. 2011 (2/9/11 - 2/15/11)		Total Th WC Pig	Total Th WC Pig	Total Th WC Pig	Total Th
2011-26 June 2011 (6/4/11 - 6/16/11)	Total Th WC Pig	Total Th WC Pig	Total Th WC Pig	Total Th WC Pig	Total Th Part. Th WC Pig Part. Pig Traps
2012-01 Feb. 2012 (2/7/12 - 2/19/12)	Total Th Part. Th WC Pig Part. Pig	Total Th Part. Th WC Pig Part. Pig	Total Th	Total Th	Total Th Part. Th WC Pig Part. Pig
2012-12 June 2012 (5/23/12 - 6/7/12)	Total Th Part. Th WC Pig Part. Pig	Total Th Part. Th WC Pig Part. Pig	Total Th Part. Th WC Pig Part. Pig	Total Th Part. Th Part. Pig	Total Th Part. Th Part. Pig Traps

Table 2: Total net primary production (NPP) and >5 µm size-fractionated NPP determined from simulated in situ incubations. ^{234}Th-derived POC flux (P_{POC}) and sediment trap POC flux ($Trap_{POC}$) determined at the base of the photic zone and the corresponding *ThE*-ratios (P_{POC}/NPP) and trap *e*-ratios ($Trap_{POC}$/NPP).

Cruise	Station	Integration Depth (m)	Total NPP (mmol m^{-2} d^{-1})	>5 µm NPP (mmol m^{-2} d^{-1})	P_{POC} (mmol m^{-2} d^{-1})	$Trap_{POC}$ (mmol m^{-2} d^{-1})	*ThE*-ratio	Trap *e*-ratio
Feb. 2011	P20	77	36.64	3.26				
June 2011	P26-D	83	105.14	13.67	2.94	5.91	0.03	0.06
	P26-R	85	78.75	12.98	2.75	5.91	0.03	0.08
Feb. 2012	P4	50	27.91	3.58	7.29		0.26	
	P12	95	34.56	4.58	4.65		0.13	
	P26	75	23.41	5.22	0.31		0.01	
June 2012	P4	103	82.36	39.55	7.95		0.10	
	P12	164	40.24	4.16	2.12		0.05	
	P20	115	57.84	4.10	0.54		0.01	
	P26	60	49.45	9.28	2.96	6.55	0.06	0.13

Table 3: Chl *a* and indicator pigment standing stocks determined by integrating small volume pigment concentrations (determined by HPLC) across the photic zone, pigment fluxes (^{234}Th and PITS-derived) measured at the base of the photic zone, and pigment loss rates, or the percent of the surface concentration represented by those fluxes. Pigment standing stocks are in mg m^{-2} and pigment fluxes are in mg m^{-2} d^{-1}.

Cruise	Station	Depth	Chl *a*	FUCO	PER	HEX	BUT	ALLO	Chl *b*	ZEA
Aug. 2010 (2010-14)	P12	Surface (1-75 m)	23.918	3.498	0.375	7.705	1.165	0.220	4.038	1.435
	P16	Surface (1-75 m)	14.165	1.288	0.340	6.010	1.018	0.065	2.588	0.165
	P20	Surface (1-75 m)	19.040	3.138	0.398	6.298	1.453	0.065	2.620	0.188
Feb. 2011 (2011-01)	P12	Surface (1-65 m)	30.122	2.848	0.379	5.630	2.431	0.838	7.133	0.922
	P16	Surface (1-95 m)	16.230	1.286	0.202	5.728	1.726	0.161	4.439	1.643
	P20	Surface (1-77 m)	55.053	5.207	0.689	18.064	6.697	1.116	11.435	4.516
June 2011 (2011-26)	P4	Surface (1-72 m)	29.791	2.635	0.127	10.619	2.663	0.720	5.836	5.234
	P12	Surface (1-90 m)	26.115	5.060	0.085	11.988	3.263	0.498	2.665	3.063
	P16	Surface (1-105 m)	22.088	4.044	0.104	11.390	2.195	0.181	2.612	1.569
	P20	Surface (1-70 m)	19.421	4.423	0.197	8.132	1.913	0.166	2.090	1.129
	P26	Surface (1-84 m)	29.376	7.239	0.184	10.532	4.406	0.232	3.723	2.663
		Flux at 100 m	0.765	0.474	0.036	0.059	0.0002	0.016	0.028	0.018
		% Flux	2.605	6.548	19.762	0.564	0.004	6.686	0.753	0.658
		Trap (150 m)	0.125	0.056	0.027	0.049	0.014	-	0.017	0.015
		% Flux	0.424	0.767	14.879	0.466	0.311	-	0.461	0.545

Feb. 2012	P4	Surface (1-38 m)	22.684	3.765	-	4.592	1.434	0.917	3.781	0.280
(2012-01)		Flux at 50 m	3.283	1.863		0.811	0.122			
		% Flux	14.471	49.468	-	17.668	8.537	-	-	-
	P12	Surface (1-38 m)	11.003	1.425	0.116	5.606	1.894	0.017	1.915	0.500
		Flux at 100 m	0.046	0.020	0.000	0.014	0.005	0.000	0.000	0.000
		% Flux	0.415	1.381	0.000	0.254	0.249	0.000	0.000	0.000
	P26	Surface (1-38 m)	12.161	2.092	1.218	2.923	1.615	0.137	0.902	0.228
		Flux at 100 m	0.380	0.251	0.035	0.046	0.038	0.000	0.014	0.045
		% Flux	3.126	11.999	2.898	1.581	2.373	0.000	1.524	19.919
June 2012	P4	Surface (1-103 m)	21.313	31.420	-	5.192	-	-	-	-
(2012-12)		Flux at 200 m	1.076	0.919	0.047	0.126	-	-	-	0.036
		% Flux	5.047	2.926	-	2.435	-	-	-	-
	P12	Surface (1-164 m)	27.677	5.967	-	22.445	6.552	-	-	-
		Flux at 200 m	0.051	0.047	-	0.075	0.010	-	0.025	-
		% Flux	0.185	0.787	-	0.335	0.156	-	-	-
	P16	Surface (1-66 m)	12.830	8.722	-	17.321	4.238	-	0.942	0.777
		Flux at 100 m	0.312	0.319	0.045	0.044	0.007	-	-	-
		% Flux	2.431	3.662	-	0.252	0.174	-	-	-
	P20	Surface (1-115 m)	18.344	33.038	-	13.892	-	-	13.090	3.538
		Flux at 100 m	0.016	0.016	0.004	-	0.002	-	0.005	0.001
		% Flux	0.088	0.049	-	-	-	-	0.036	0.033
	P26	Surface (1-60 m)	14.024	1.977	-	13.572	2.018	-	4.969	2.768
		Flux at 100 m	0.255	0.304	-	-	0.029	-	0.025	-
		% Flux	1.821	15.359	-	-	1.437	-	0.507	-
		Trap (100 m)	0.055	0.025	0.006	0.041	0.004	-	0.009	0.008
		% Flux	0.393	1.243	-	0.304	0.190	-	0.179	0.288

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Table 4: ^{234}Th and POC fluxes and POC/^{234}Th ratios measured by the PITS traps.

Depth (m)	Days In-situ	^{234}Th flux (dpm m^{-2} d^{-1})	POC flux (mmol m^{-2} d^{-1})	POC/^{234}Th ratio (µmol dpm^{-1})
June 2011 P26				
30	3.32	3192 ± 117	15.3 ± 0.4	4.8 ± 0.2
50	3.32	2909 ± 92	10.1 ± 0.3	3.5 ± 0.1
100	3.32	2256 ± 94	5.9 ± 0.2	2.6 ± 0.1
150	3.32	1928 ± 79	5.0 ± 0.2	2.6 ± 0.1
200	3.32	2281 ± 97	8.5 ± 0.3	3.7 ± 0.2
June 2012 P26				
30	2.82	3999 ± 206	14.7 ± 0.4	3.7 ± 0.2
50	2.82	5485 ± 290	13.5 ± 0.5	2.5 ± 0.2
100	2.82	3154 ± 192	6.5 ± 0.2	2.1 ± 0.1
150	2.82	2151 ± 135	5.5 ± 0.2	2.5 ± 0.2
200	2.82	3959 ± 129	5.0 ± 0.2	1.3 ± 0.1

748 Figure 1. Map showing the Line P stations sampled in this study.
749
750 Figure 2. Temperature (°C), Sigma-t (kg m^{-3}), and $^{234}Th/^{238}U$ activity ratio distributions along
751 Line P cruises in August 2010, February 2011, June 2011, February 2012, and June 2012.
752
753 Figure 3. Comparison of small-volume Niskin bottle and large-volume in situ pump
754 measurements of a) POC, b) picoplankton indicator pigments, c) nanoplankton indicator
755 pigments, d) microplankton pigments. Niskin bottle measurements are lower than pump
756 measurements for microplankton pigments, and higher for nanoplankton pigments and POC.
757
758 Figure 4. a) $POC/^{234}Th$ ratios on 1 – 10-µm particles and on 10 – 53-µm particles plotted against
759 the $POC/^{234}Th$ ratio on >53-µm particles. Fractional distributions of POC and particulate ^{234}Th
760 are plotted for three size-classes of particles. The percentage of total POC associated with each
761 particle size-class is plotted against the percentage of total particulate ^{234}Th for samples collected
762 at stations on Line P during b) June 2011, c) February 2012, and d) June 2012. The correlation
763 coefficient (r^2) and the slope of the linear regression (m) are shown for each cruise.
764
765 Figure 5. Pigment Proportion Factors (PF) for each phytoplankton size-class plotted as a
766 function of sample depth at stations sampled on Line P during the five cruises in the study. All
767 data were collected from Niskin bottles.
768
769 Figure 6. Pigment PF for each phytoplankton size group plotted as a function of sample depth
770 and particle size-class at stations sampled on Line P in June 2011. Size-fractioned data are pump
771 data. Sediment trap PF's are also included.
772
773 Figure 7. Pigment PF for each phytoplankton size group plotted as a function of sample depth
774 and particle size-class at stations sampled on Line P in February 2012. Size-fractioned data are
775 pump data.
776

Figure 8. Pigment PF for each phytoplankton size group plotted as a function of sample depth and particle size-class at stations sampled on Line P in June 2012. Size-fractioned data are pump data. Sediment trap PF's are also included where available.

Figure 9. Depth profiles of a) ^{234}Th fluxes (P_{Th}) determined using the 2-D model, b) POC/^{234}Th ratios on >53 μm particles, and c) ^{234}Th-derived POC fluxes (P_{POC}) at stations on Line P during the five cruises in this study.

Figure 10. Net primary production (NPP) plotted against ^{234}Th-derived POC fluxes (P_{POC}) for stations along Line P in this study. The slopes of the dashed lines represent *ThE*-ratios. For reference NPP and P_{POC} values determined by Charette et al. (1999) for winter, spring and summer are included, along with annual average NPP and sediment trap POC fluxes (at 200 m) reported in Harrison (2002) and Wong et al. (1999) respectively.

Figure 11. a) Comparison of sediment trap POC fluxes and ^{234}Th-derived POC fluxes, and b) a comparison of sediment trap Chl *a* and total indicator pigment fluxes and ^{234}Th-derived pigments fluxes at OSP during June 2011 and June 2012.

Figure 12. a-c) ^{234}Th-derived indicator pigment fluxes determined using the Pigment/^{234}Th ratio on >53-μm particles plotted for micro-, nano-, and picoplankton pigments. d-f) Indicator pigment standing stocks plotted against indicator pigment fluxes for micro-, nano-, and picoplankton pigments. The slopes of the dashed lines indicate pigment loss rates. g-i) The contribution to total pigment standing stock plotted against the contribution to total pigment flux for micro-, nano-, and picoplankton pigments. Data points above the 1:1 line indicate preferential export by direct sinking and points below the 1:1 line indicate disproportionately low export by direct sinking relative to biomass contributions.

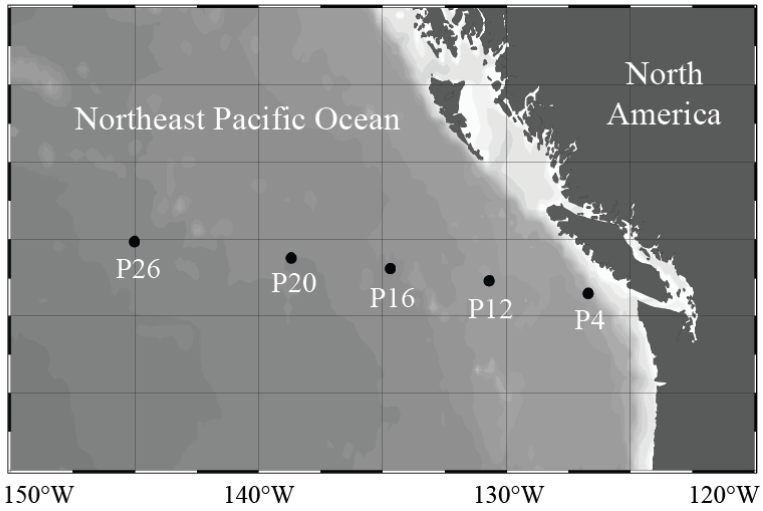

54°N
50°N
46°N
150°W
140°W
130°W
120°W
Northeast Pacific Ocean
North
America
P26
P20
P16
P12
P4

Fig. 1.

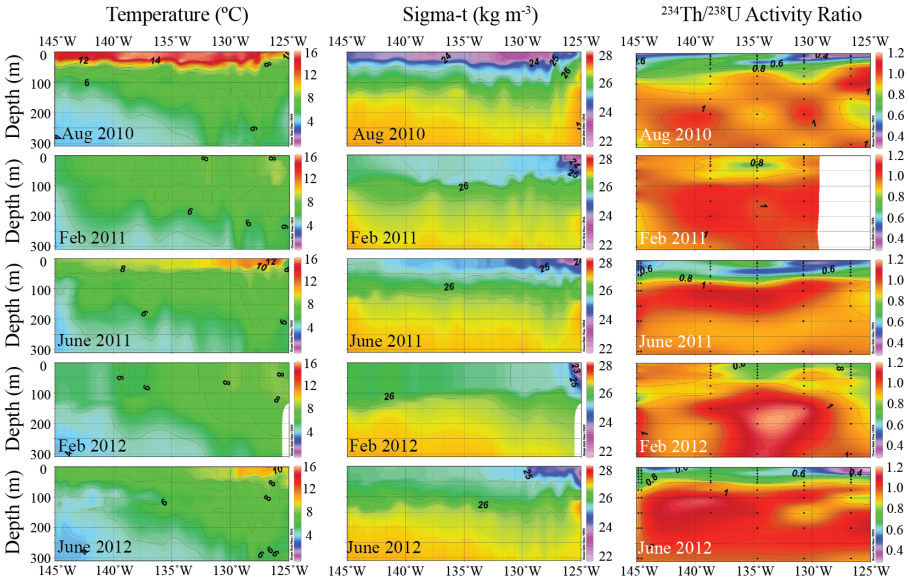

Temperature (ºC)
Sigma-t (kg m-3)
234Th/238U Activity Ratio
Depth (m)
Aug 2010
Feb 2011
June 2011
Feb 2012
June 2012
145˚W
140˚W
135˚W
130˚W
125˚W

Fig. 2.

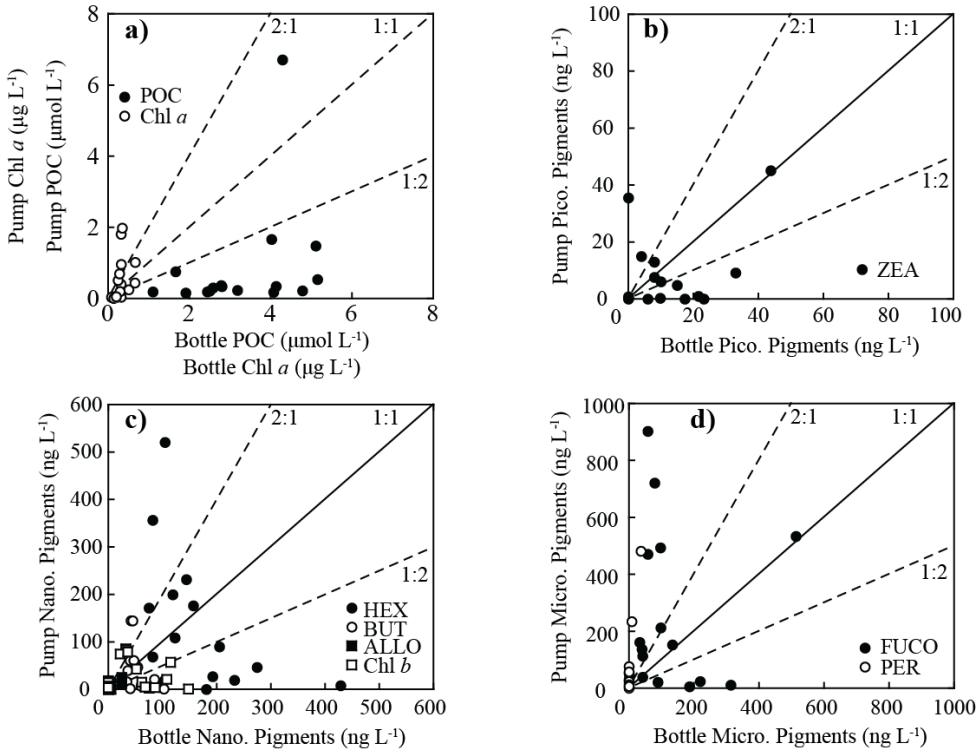

a)
2:1
1:1
1:2
POC
Chl a
Pump Chl a (µg L-1)
Pump POC (µmol L-1)
Bottle POC (µmol L-1)
Bottle Chl a (µg L-1)
b)
ZEA
Pump Pico. Pigments (ng L-1)
Bottle Pico. Pigments (ng L-1)
c)
HEX
BUT
ALLO
Chl b
Pump Nano. Pigments (ng L-1)
Bottle Nano. Pigments (ng L-1)
d)
FUCO
PER
Pump Micro. Pigments (ng L-1)
Bottle Micro. Pigments (ng L-1)

Fig. 3.

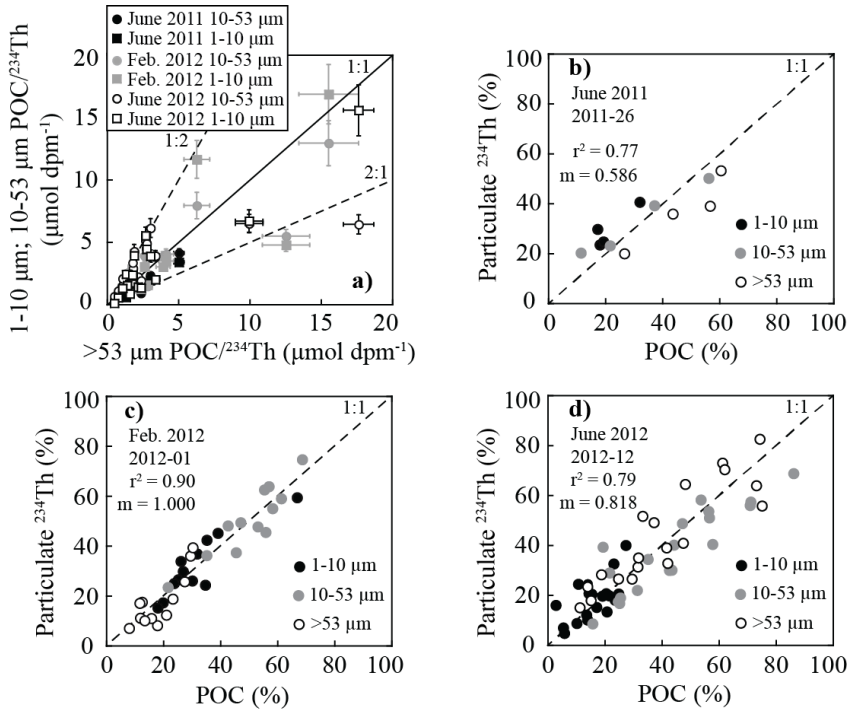
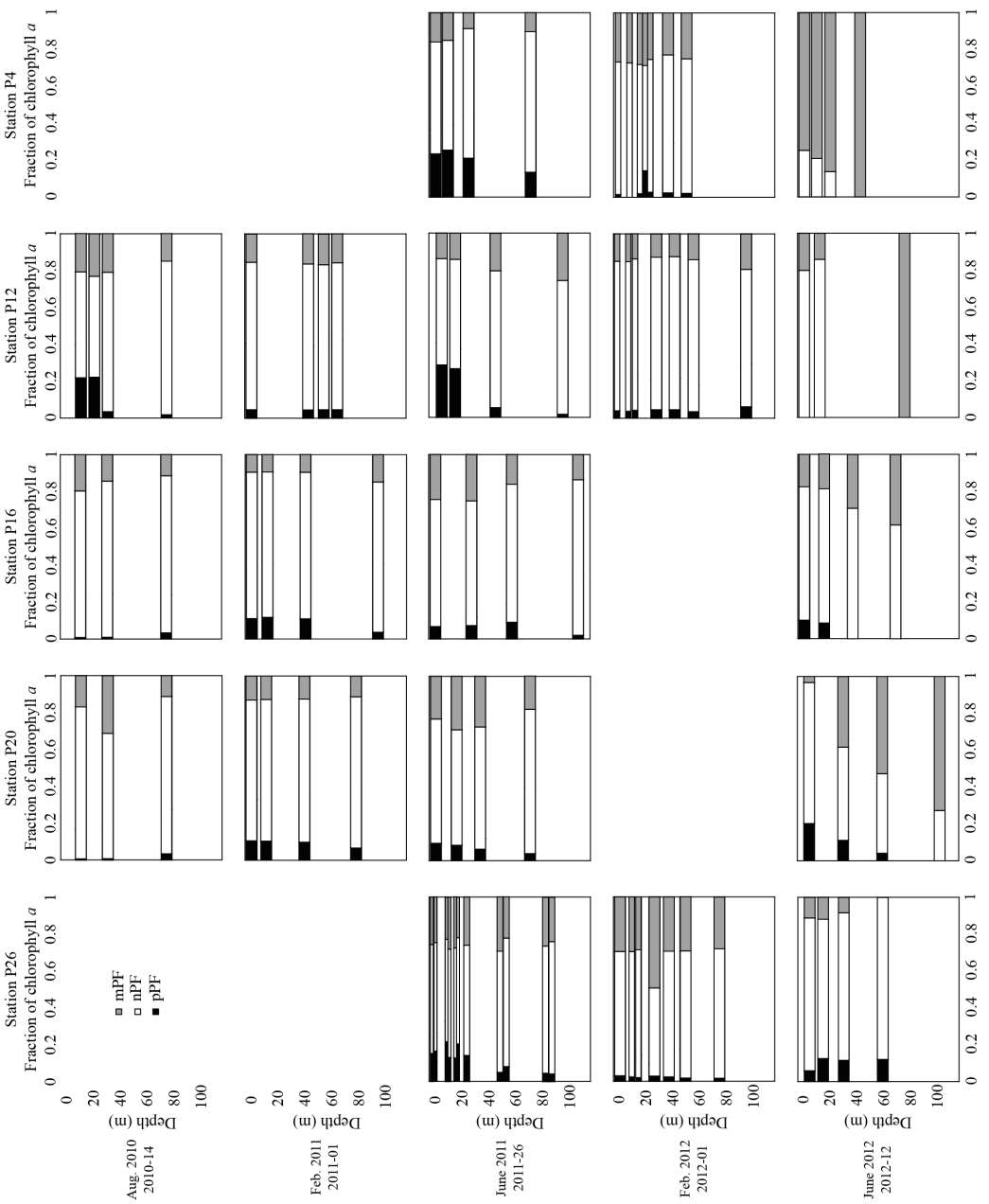

June 2011 10-53 µm
June 2011 1-10 µm
Feb. 2012 10-53 µm
Feb. 2012 1-10 µm
June 2012 10-53 µm
June 2012 1-10 µm
1:1
1:2
2:1
a)
1-10 µm; 10-53 µm POC/234Th (µmol dpm-1)
>53 µm POC/234Th (µmol dpm-1)
b)
June 2011
2011-26
r2 = 0.77
m = 0.586
Particulate 234Th (%)
POC (%)
1-10 µm
10-53 µm
>53 µm
c)
Feb. 2012
2012-01
r2 = 0.90
m = 1.000
d)
June 2012
2012-12
r2 = 0.79
m = 0.818

Fig. 4.



858
859
860 Fig. 5.

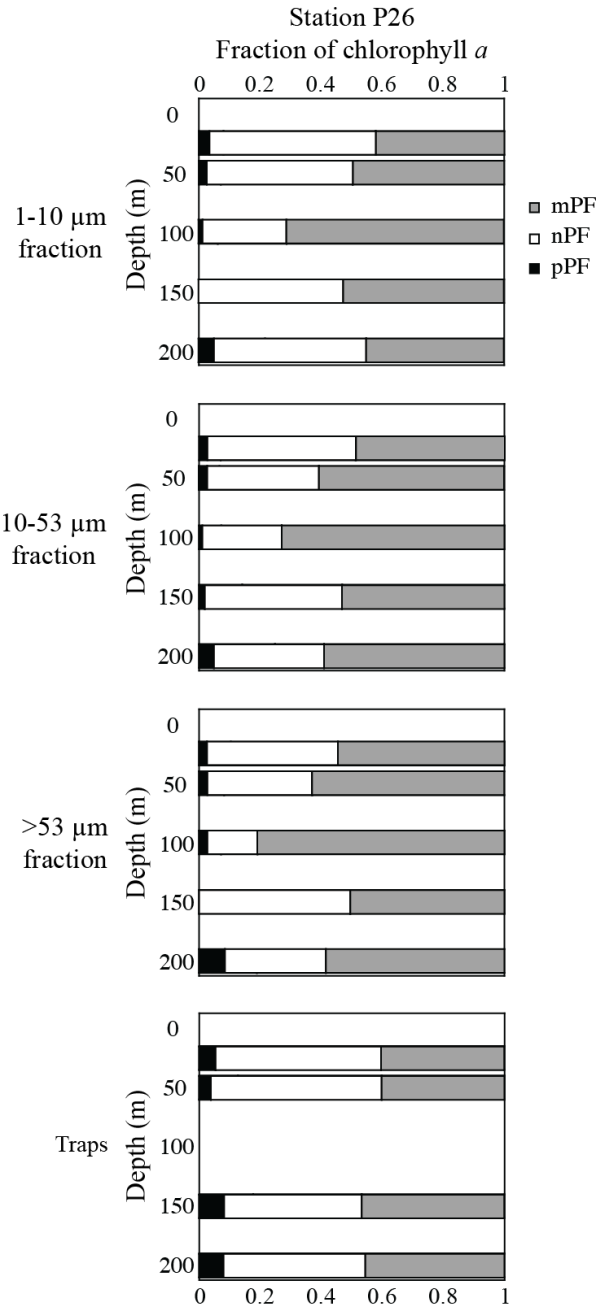

Station P26
Fraction of chlorophyll *a*
0 0.2 0.4 0.6 0.8 1
1-10 µm fraction
10-53 µm fraction
>53 µm fraction
Traps
Depth (m)
0
50
100
150
200
mPF
nPF
pPF

Fig. 6.

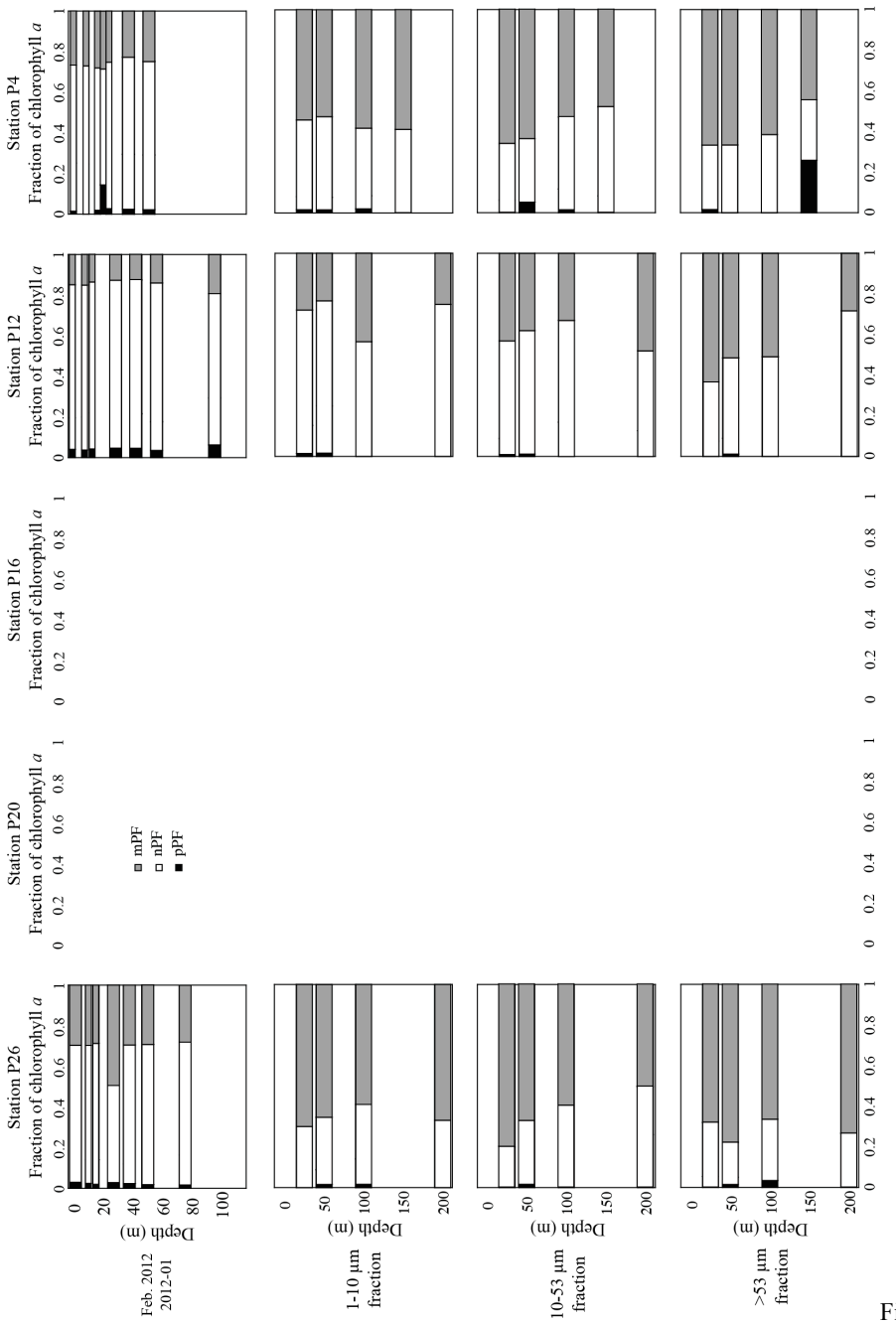

Station P26
Station P20
Station P16
Station P12
Station P4
Fraction of chlorophyll a
mPF
nPF
pPF
Depth (m)
Feb. 2012
2012-01
1-10 µm
fraction
10-53 µm
fraction
>53 µm
fraction

Fig. 7.

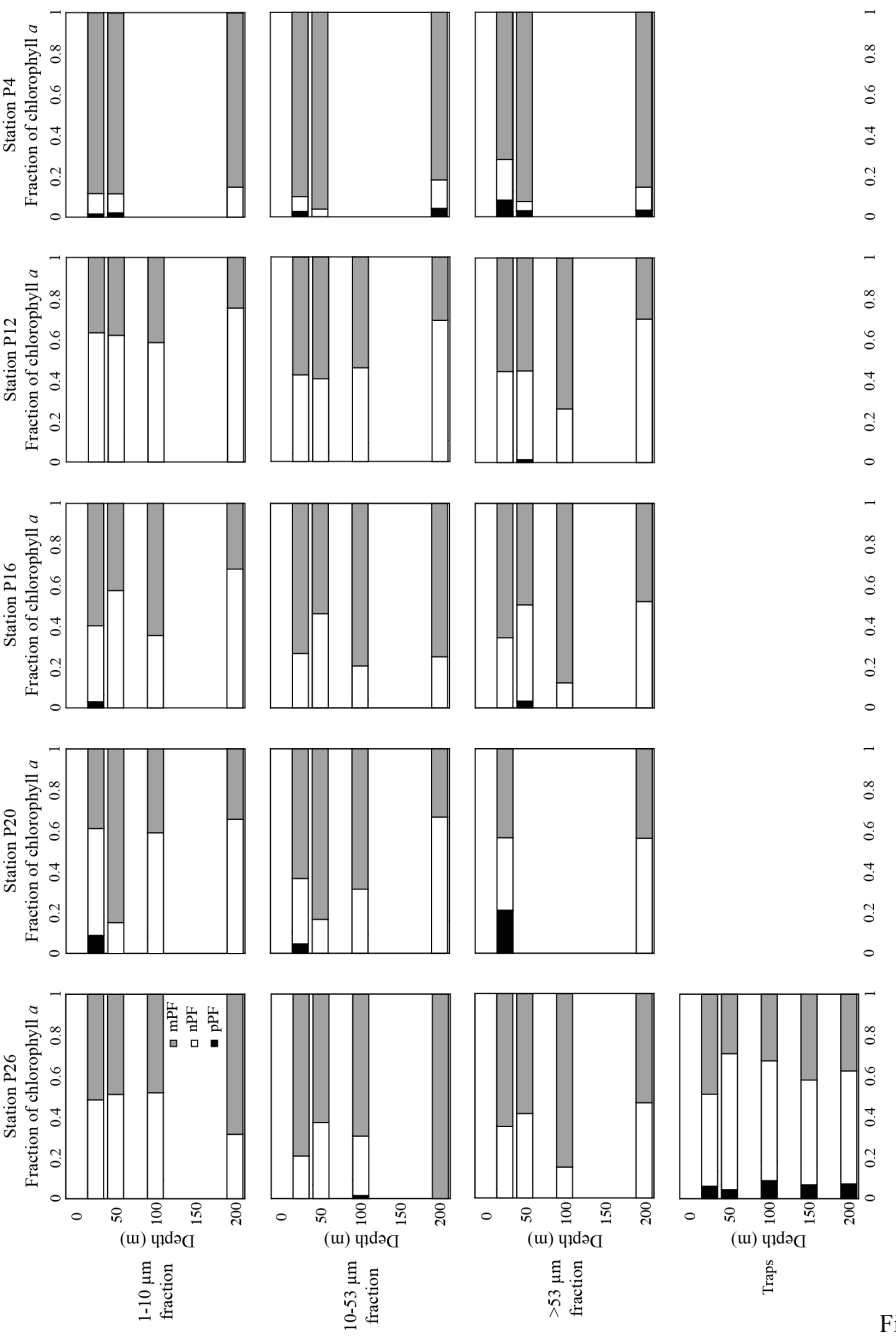

Station P26
Station P20
Station P16
Station P12
Station P4
Fraction of chlorophyll a
0 0.2 0.4 0.6 0.8 1
mPF
nPF
pPF
Depth (m)
0 50 100 150 200
1-10 µm fraction
10-53 µm fraction
>53 µm fraction
Traps

Fig. 8.

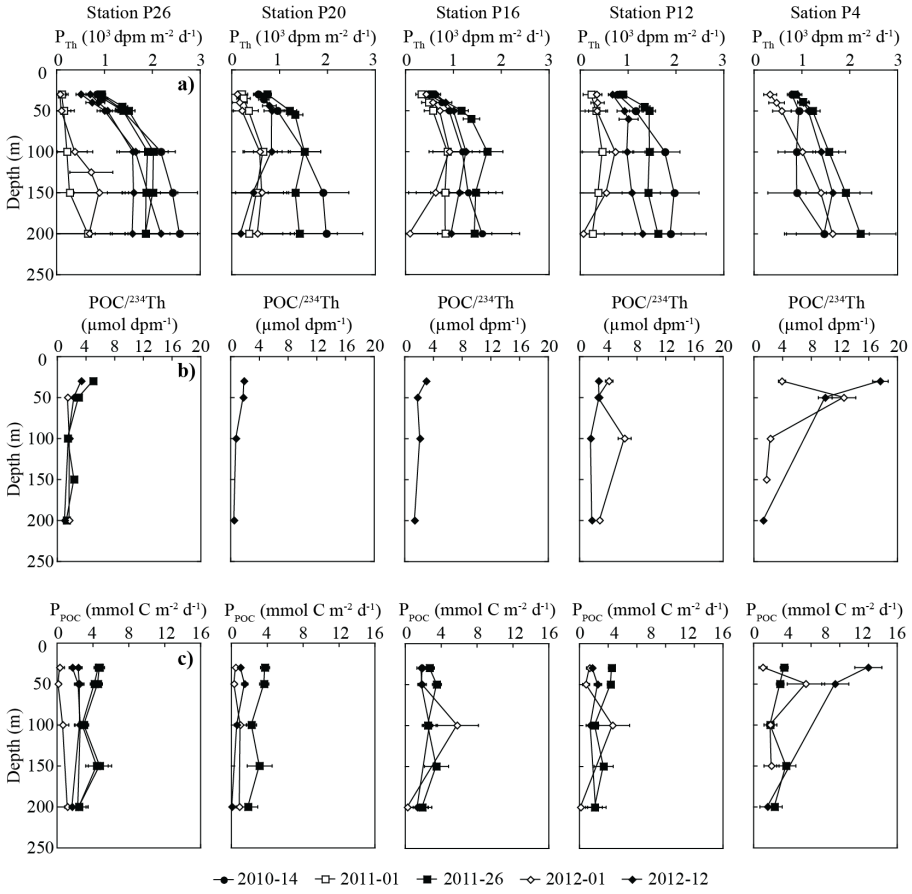

Station P26
Station P20
Station P16
Station P12
Sation P4
P_Th (10³ dpm m⁻² d⁻¹)
a)
b)
c)
Depth (m)
POC/²³⁴Th (µmol dpm⁻¹)
P_POC (mmol C m⁻² d⁻¹)
2010-14
2011-01
2011-26
2012-01
2012-12

Fig. 9.

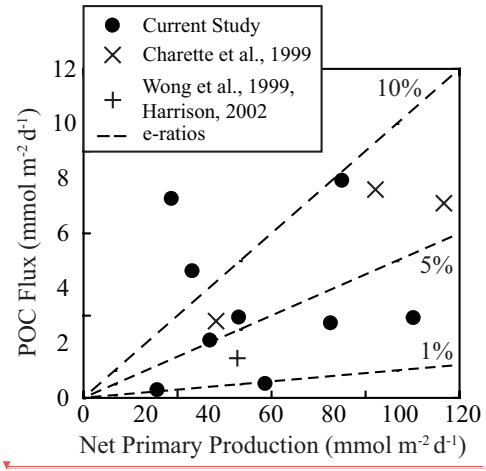
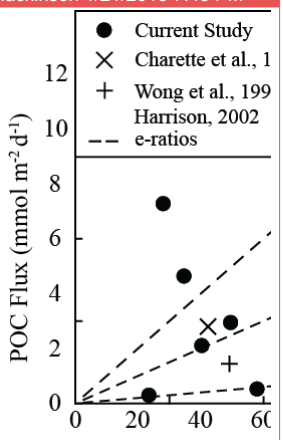


Fig. 10.

Brendan Mackinson 1/21/2015 7:45 PM



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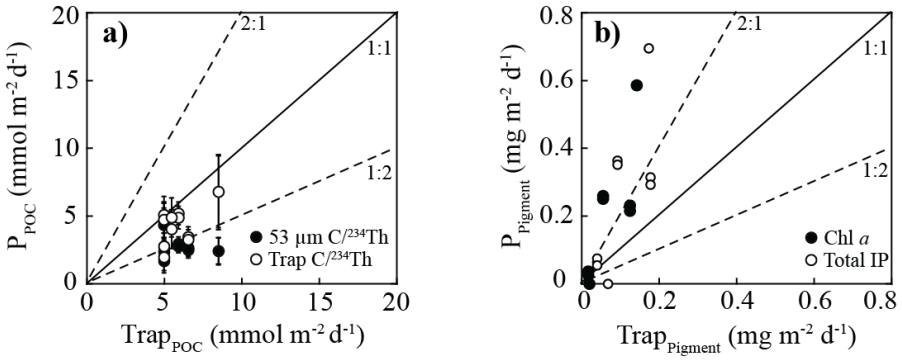
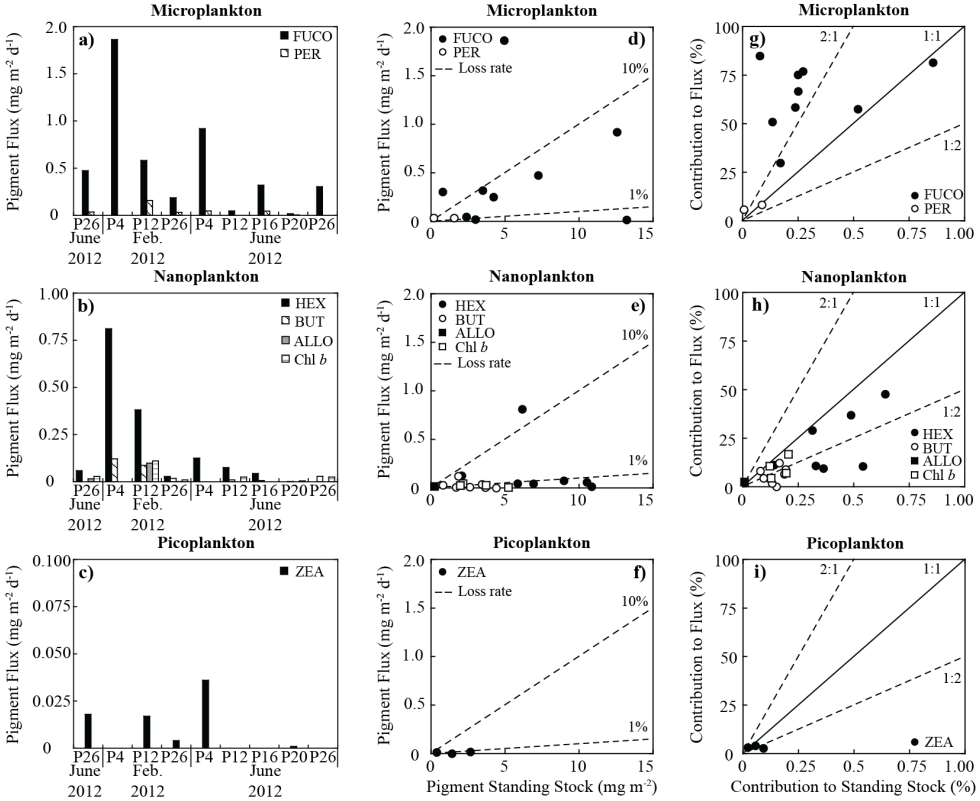

a)
2:1
1:1
1:2
53 µm C/234Th
Trap C/234Th
P POC (mmol m-2 d-1)
Trap POC (mmol m-2 d-1)
b)
2:1
1:1
1:2
Chl a
Total IP
P Pigment (mg m-2 d-1)
Trap Pigment (mg m-2 d-1)

Fig. 11.



913
914
915 Fig. 12.