Thank you very much for taking into consideration the reviewers' comments. I consider that your manuscript has been improved significantly. So we are very pleased to inform you that your manuscript has been accepted for publication in Biogeosciences. Before it is published, please consider my technical corrections below:

L88: a "surface" chl

We have corrected as suggested. (L88)

L89: Did the authors use Aqua/MODIS data? Otherwise, Terra or both?

We used Aqua/MODIS data. We have added the information at L89.

L174–175 plus Supporting Information L10: what does the meaning of 98"+" atom%? Did the authors assume 98 atom% 15N for the calculation of nitrogen fixation rates?

The gas provided from SI Science Co. is labeled 98+ atom%. We assume 98 atom% for the calculation. To avoid confusion, "+" has been deleted from the revised manuscript.

L210–217: What software(s) did the authors use for the nMDS and ANOSIM analyses?

The nMDS and ANOSIM analyses were performed using PRIMER 6 software. We have added this sentence at L218.

L275: phosphate "concentrations" ... "were" ...

We have corrected as suggested. (L276)

L286: dissolved iron "concentrations" ... "were" ...

We have corrected as suggested. (L287)

L341: Remove the hyphen between "spp." and "bloom".

We have removed as suggested.

L381: Remove a space between "that" and "in".

We have removed as suggested.

L454: Remove a comma immediately after "plot".

We have removed as suggested.

L564: Start a new line from "Jickells, T.D...".

We have changed as suggested.

Supporting Information L60: Remove a comma immediately after "2015".

We have removed as suggested.

Legend in Fig. S3: KT-07-21

We have corrected as suggested.

Why is Trichodesmium abundant in the Kuroshio? 1 T. Shiozaki^{1,2}, S. Takeda^{1,3}, S. Itoh², T. Kodama^{1,4}, X. Liu^{1,5}, F. Hashihama⁶, K. $\mathbf{2}$ Furuya¹ 3 [1]{Department of Aquatic Bioscience, Graduate School of Agricultural and Life 4 Sciences, The University of Tokyo, Tokyo, 113-8657, Japan } $\mathbf{5}$ 6 [2]{Atmosphere and Ocean Research Institute, The University of Tokyo, Chiba, $\overline{7}$ 277-8564, Japan} [3] {Faculty of Fisheries, Nagasaki University, Nagasaki, 852-8521, Japan } 8 [4] [Japan Sea National Fisheries Research Institute, Fisheries Research Agency, 9 10 Niigata, 951-8121, Japan} [5]{College of Ocean and Earth Sciences, Xiamen University, Xiamen, 361005, 11 China} 1213[6] {Department of Ocean Sciences, Tokyo University of Marine Science and Technology, Tokyo, 108-8477, Japan} 14

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17 Abstract

The genus Trichodesmium is recognized as an abundant and major diazotroph in the 1819 Kuroshio, but the reason for this remains unclear. The present study investigated the abundance of *Trichodesmium* spp. and nitrogen fixation together with concentrations 20of dissolved iron and phosphate in the Kuroshio and its marginal seas. We performed 2122the observations near the Miyako Islands, which form part of the Ryukyu Islands, 23situated along the Kuroshio, since our satellite analysis suggested that material transport could occur from the islands to the Kuroshio. Trichodesmium spp. bloomed 24 $(>20,000 \text{ filaments } L^{-1})$ near the Miyako Islands, abundance was high in the Kuroshio 2526and the Kuroshio bifurcation region of the East China Sea, but was low in the Philippine Sea. The abundance of Trichodesmium spp. was significantly correlated 2728with the total nitrogen fixation activity. The surface concentrations of dissolved iron 29(0.19–0.89 nM) and phosphate (<3–36 nM) were similar for all of the study areas, 30 indicating that the nutrient distribution could not explain the spatial differences in 31Trichodesmium spp. abundance and nitrogen fixation. Numerical particle-tracking 32experiments simulated the transportation of water around the Ryukyu Islands to the Kuroshio. Our results indicate that Trichodesmium growing around the Ryukyu 33 34Islands could be advected into the Kuroshio.

1. Introduction

37	The Kuroshio is a western boundary current in the North Pacific Ocean that
38	originates in the North Equatorial Current and bifurcates to the east of the Philippines.
39	The main stream of the Kuroshio enters the East China Sea (ECS) northeast of Taiwan,
40	flows out through the Tokara Strait, and runs along the Japanese islands of Shikoku
41	and Honshu. While the Kuroshio and its adjacent waters are characterized by highly
42	oligotrophic conditions, phytoplankton and zooplankton communities in the Kuroshio
43	are distinct compared to those from adjacent waters (McGowan, 1971). McGowan
44	(1971) suggested that some plankton species are delivered by the Kuroshio to the
45	north from the equatorial region.
46	The abundance of the cyanobacterial genus Trichodesmium in the Kuroshio is
47	much higher than that in neighboring seas (Marumo and Asaoka, 1974). Because
48	Trichodesmium is a major nitrogen fixer in the Kuroshio, it is believed to be the key
49	genus for understanding the Kuroshio ecosystem (Chen et al., 2008, 2014; Shiozaki et
50	al., 2014a). Nevertheless, the factors controlling the distribution of Trichodesmium in
51	this region are poorly understood. Marine nitrogen fixation is thought to be regulated
52	by the supply of iron and phosphorus (Mahaffey et al., 2005), and Trichodesmium
53	thrives in iron-rich oligotrophic regions (Moore et al., 2009; Shiozaki et al., 2010,

54	2014b). A major source of iron in the ocean is atmospheric dust deposition (Jickells et
55	al., 2005; Mahowald et al., 2009). Modeling studies indicate that dust deposition in
56	the western North Pacific decreases exponentially from the continental shelf to the
57	Philippine Sea (Jickells et al., 2005; Mahowald et al., 2009), and hence, deposition is
58	not as high in the Kuroshio as in the adjacent waters. As for phosphorus limitation,
59	iron-enhanced nitrogen fixation causes phosphorus depletion, and the nitrogen
60	fixation is consequently limited by phosphorus (Mather et al., 2008). The phosphate
61	distribution has been examined in this study region using a conventional colorimetric
62	method, and the surface phosphate concentration in the Kuroshio has been reported to
63	be as low as that in the Philippine Sea (Chen, 2008). Therefore, the distinct high
64	abundance of Trichodesmium in the Kuroshio is probably not explained by nutrient
65	and trace metal concentrations; however, distributions of dissolved iron and phosphate
66	at the nanomolar level have not been well studied in this region (Obata et al., 1997;
67	Shiozaki et al., 2010; Kodama et al., 2011).
68	Nitrogen fixation by Trichodesmium has recently also been found to be active
69	around oceanic islands; New Caledonia Island, Efate Island, Fiji Island, Tahiti Island,
70	and Northern Mariana Islands (Shiozaki et al., 2010, 2013, 2014c; Lin et al., 2011).
71	Furthermore, these studies demonstrated that abundant Trichodesmium is delivered by

 noted in the western Pacific warm pool and western South Pacific, it can also occur in and around the Kuroshio and may contribute to the distribution of <i>Trichodesmium</i> in this region. In the present study, we simultaneously determined <i>Trichodesmium</i> abundance and bulk water nitrogen fixation together with concentrations of dissolved iron and phosphate at the nanomolar level in the Kuroshio and its marginal seas. In addition, we conducted intensive observations around the Miyako Islands section of the 		the current to areas that are remote from the Islands. Although this phenomenon was
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80 km resolution during summer between July 2003 and September 2009. We defined 81 summer as July through September. The bloom frequency for each pixel was 82 calculated from the ratio of counts in which chl *a* was >0.15 mg m⁻³ to the total 83 counts in which chl *a* was detected.

To examine the current field, geoelectrokinetograph and ship-mounted acoustic Doppler current profiler (ADCP) data from the uppermost layer for the summers between 1953 and 2008 were obtained from the Japan Oceanographic Data Center (http://www.jodc.go.jp). Regridding, removal of anomalous values, and smoothing of the dataset were performed as described by Isobe (2008).

99

100 2.2. Cruise observations

Experiments were conducted during summer on-board the R/V *Tansei-maru* (KT-06-21, September 9–17, 2006; KT-07-22, September 5–13, 2007; KT-09-17, September 8–13, 2009; KT-10-19, September 4–12, 2010) and the T/V *Nagasaki-maru* (242, July 19–28, 2007) (Fig. 1a, Table S1). The stations during the KT-06-21, KT-07-22, and *Nagasaki-maru* 242 cruises were divided into three areas based on the temperature-salinity diagram (see Fig.2 of Shiozaki et al., 2011): the ECS, Kuroshio, and Philippine Sea. During the KT-09-17 cruise, we conducted

108	experiments around the Miyako Islands which were distinguished from the other three
109	areas. During the KT-10-19 cruise, we performed observations in the ECS, the
110	Kuroshio, and around the Miyako Islands (Liu et al., 2013).
111	
112	2.2.1. Light intensity, hydrography, nutrients, and chl a
113	Water samples for all of the experiments, with the exception of determination of
114	the dissolved iron concentration, were collected using an acid-cleaned bucket and
115	Niskin-X bottles. The depth profile of light intensity was determined immediately
116	before the water sampling using a light sensor (during the KT-06-22, KT-07-21,
117	KT-09-17, and KT-10-19 cruises) or an empirical equation (during the Nagasaki-maru
118	242 cruise) (Shiozaki et al., 2011). Temperature and salinity profiles to a depth of 200
119	m were obtained using a conductivity, temperature, and depth (CTD) sensor. Mixed
120	layer depth (MLD) was defined as the depth at which the sigma-t increased by 0.125
121	from its value at a depth of 10 m. Water samples for nitrate+nitrite (N+N) and
122	phosphate were collected from 0, 10, 30, 40, 50, 60, 70, 80, 90, 100, 125, 150, and
123	200 m, and from depths at given light intensities. At all of the stations, the N+N and
124	phosphate concentrations were determined at the nanomolar level using a
125	supersensitive colorimetric system consisting of an AutoAnalyzer II (Technicon) and

126	Liquid Waveguide Capillary Cells (World Precision Instruments, USA) (Hashihama et
127	al., 2009). The detection limits of N+N and phosphate were both 3 nM. When the
128	concentration was greater than 0.1 μ M, it was determined by conventional methods
129	using a TRAACS 2000 autoanalyzer (Bran:Luebbe, UK). In addition to the
130	observations at the stations, temperature, salinity, and the in vivo chl fluorescence of
131	the surface water were monitored continuously during the cruises by a
132	thermosalinograph (Ocean Seven, Idronaut, Italy) and a fluorometer (Minitracka,
133	Chelsea, UK).

135 **2.2.2. Dissolved iron**

Water was sampled to estimate the dissolved iron concentration from 0.5-m depth 136 during the KT-06-21 and KT-07-22 cruises and from 10-m depth during the KT-09-17 137138cruise using an acid-cleaned Teflon bellows pump (AstiPure PFD2; Saint-Gobain) with Teflon tubing (inner diameter = 12 mm). The water was filtered through an 139acid-cleaned 0.22 µm pore filter (Millipak100; Millipore) connected to the in-line of 140 141the Teflon tubing with a Teflon connector. Filtered seawater was collected in a 125 142mL low-density polyethylene (LDPE) bottle (Nalgene, Nalge Nunc International), which had been washed using following technique: the sample bottles were 143

144	sequentially cleaned by soaking in 5% alkali detergent for at least 2 days, in 4 N HCl
145	for at least 1 day, in 0.3 N metal analysis-grade HNO3 at 60°C overnight, and finally,
146	in Milli-Q water at 60°C overnight. After rinsing with Milli-Q water, the bottles were
147	dried in a laminar flow space and stored in double plastic bags. The filtrate samples
148	were acidified to a pH <1.7 with trace-metal-grade HCl (Tamapure AA-100; Tama
149	Chemicals) in a Class-100 clean-air bench, and stored at room temperature for more
150	than 1 year.
151	The dissolved iron concentration was determined using an automatic Fe(III) flow
152	injection analytical system (Kimoto Electric Co., Ltd.) using a chelating resin
153	pre-concentration and chemiluminescence detection method (Obata et al., 1993). A
154	buffer solution of 10 M formic acid and 2.4 M ammonium formate was added to the
155	samples. The sample pH was adjusted to 3.0 with 20% ammonium hydroxide
156	(NH ₄ OH; Tamapure AA-10; Tama Chemicals) immediately prior to analysis. The
157	detection limit of this method was 0.05 nM. The SAFe reference standards S1 and D2
158	were measured during the course of sample analysis, and the results were within the
159	range of the published consensus values: S1 = 0.097 \pm 0.043 nM and D2 = 0.91 \pm
160	0.17 nM (Johnson et al., 2007).

162 2.2.3. Nitrogen fixation and abundance of *Trichodesmium* spp.

163	Samples for the incubation experiments were collected vertically at all of the
164	stations, except at Sts. T0621, GN-3, and T0905, where samples were only collected
165	from the surface. All samples were collected in duplicate in acid-cleaned 4.5-L
166	polycarbonate bottles. During the Nagasaki-maru 242 cruise, water samples were
167	collected from four different depths corresponding to 100%, 25%, 10%, and 1% of the
168	surface light intensity. During the other cruises, samples were collected from a depth
169	of 50% surface light intensity. Samples at 100% surface light intensity were collected
170	from 0 m during all of the cruises, except during the KT-10-19 cruise in which the
171	samples were collected from a depth of 5 m. The bulk water nitrogen fixation activity
172	was determined based on primary production using a dual isotopic ($^{15}N_2$ and ^{13}C)
173	technique (Shiozaki et al., 2009). After ¹³ C-labeled sodium bicarbonate (99 atom%
174	13 C; Cambridge Isotope Laboratories) was added to each bottle, 2 mL of 15 N ₂ gas (98
175	+ atom% ¹⁵ N; SI Science Co. Japan) was injected directly into the incubation bottles
176	through a septum using a gastight syringe. The bottles were covered with
177	neutral-density screens to adjust the light level and incubated for 24 h in an on-deck
178	incubator cooled by flowing surface seawater for 24 h. We determined the nitrogen
179	fixation activity using the ${}^{15}N_2$ gas bubble addition method (Montoya et al., 1996).

180	This method is believed to underestimate the nitrogen fixation rate relative to the ${\rm ^{15}N_2}$
181	gas dissolution method (Mohr et al., 2010). The start time of incubation in this study
182	varied at each station (Table S1). Considering daily periodicity of nitrogen fixation in
183	each diazotroph (Zehr, 2011) and the time to reach equilibration of the $^{15}\mathrm{N}_2$ gas
184	bubble with seawater (>12 h, Mohr et al., 2010), the level of underestimation could
185	vary at each station. Meanwhile, the level of underestimation is thought to be low in
186	Trichodesmium dominant water because Trichodesmium can float to the top of the
187	bottle and directly use the added ${}^{15}N_2$ in the bubble method (Großkopf et al., 2012).
188	Although the bias of underestimation could not be estimated from the results in this
189	study, the actual nitrogen fixation rate could be higher than the obtained rate.
190	A recent study demonstrated that commercial $^{15}N_2$ gas could be contaminated by
191	¹⁵ N-labeled nitrate and ammonium (Dabundo et al., 2014). We tested the
192	contamination in $^{15}\mathrm{N}_2$ gas produced by SI Science Co., Ltd., which was used (from
193	different batch numbers) in the present study (see Supporting Information). Briefly,
194	the ${}^{15}\mathrm{N}_2$ gas was dissolved in aged subtropical surface water, and concentrations of
195	nitrate, nitrite, and ammonium at the nanomolar levels were determined using
196	supersensitive colorimetric systems. The results showed that there were no significant
197	differences between the control and samples to which ${}^{15}N_2$ had been added (Fig. S1),

suggesting that the contamination of nitrate, nitrite, and ammonium in the ${}^{15}N_2$ gas was insignificant (Supporting Information).

200 Water samples were collected for microscopic analysis at all light depths during the Nagasaki-maru 242 and KT-07-21 cruises, and only from the surface during the 201 202 KT-06-22, KT-09-17, and KT-10-19 cruises. The samples were fixed using acidified Lugol's solution. Trichodesmium spp. were counted using the Utermöhl method under 203204 inverted microscope observation. Trichodesmium greater than ca. 300 µm in length were counted as 1 filament and shorter lengths were counted as 0.5 filaments. In 205addition, phytoplankton other than Trichodesmium spp. were identified from the 206 207 samples obtained during the KT-09-17 cruise.

208

209 2.3. Statistical analysis of environmental variables

We used non-metric multi-dimensional scaling (nMDS) to investigate the spatial differences in the environmental variables that could influence *Trichodesmium* growth and bulk water nitrogen fixation; temperature, mixed layer depth, nitrate, dissolved iron, and phosphate. The environmental variables were transformed by $log_{10}(x + 1)$ prior to analysis. A dissimilarity/similarity matrix between stations was constructed using the Bray-Curtis index. The nMDS was used to visualize similarities in the environmental variables among the stations. An Analysis of Similarity (ANOSIM)
was used to test the differences in the environmental variables among the stations._
<u>The nMDS and ANOSIM analyses were performed using PRIMER 6 software.</u>

219

220 2.4. Numerical experiments

Numerical particle-tracking experiments were conducted to investigate the 221222 transport of water masses at the surface from areas around the Miyako Islands in the summer season from 2003 to 2009. Surface velocity data were derived from the 223224FRA-JCOPE2 reanalysis product (Miyazawa et al., 2009), which is an eddy-resolving (1/12°) ocean model combined with three-dimensional variational data assimilation 225(satellites, ARGO floats, and shipboard observations), and is one of the most reliable 226models for the region around Japan for the above time period. The method of tracking 227particles was basically the same as in Itoh et al. (2009), but we did not include the 228random walk for simplicity. The release points of particles were selected at the surface 229230of the model grid points around the coastal waters of the Miyako Islands. We assumed 231that the particles did not increase, die, or sink from the surface during the experiments. To focus on transport during the summer season (July-September), particles were 232released one month before the summer (June 1) and were tracked until September 30. 233

234	To examine differences in the output depending on the start time within the same
235	year, we also performed experiments starting on June 1, 11, and 21, and July 1 in
236	2009. The ratio of particles that reached areas downstream of the Tokara Strait
237	(hereafter Area K) (Fig. 7), including the particles' entrainment to the Kuroshio, to
238	total particles released from the Miyako Islands was computed in all experiments. It
239	should be noted that these experiments contained the following two uncertainties.
240	First, the distribution of <i>Trichodesmium</i> around the islands, which strongly influences
241	the destinations of particles, was not able to be determined in advance.
242	Trichodesmium is known to aggregate and not to occur uniformly in the ocean
243	(Capone et al., 1997). Second, the model cannot reproduce the current very close to
244	the islands. If a water mass very near the islands was delivered to the open ocean by
245	tide and/or river plumes that were not considered in the model, seaward dispersion of
246	particles was likely underestimated.

248 **3. RESULTS**

249 3.1. The Kuroshio path and bloom frequency

The average surface current field indicated that the main stream of the Kuroshio flowed along the continental shelf in the ECS, and then passed to the south of the

252	Kyushu and Shikoku Islands (Fig. 1b). In addition, the Kuroshio branch bifurcated
253	northward at 25°N and 30°N at the continental shelf. Hence, all of the stations in the
254	ECS were subject to the influence of the Kuroshio. While the northeastward stream of
255	the Kuroshio was prominent in this region, smaller-scale flows and circulations were
256	observed in the areas around and to the southeast of the Ryukyu Islands. In the west of
257	the main stream of the Kuroshio, because the average chl a was over 0.15 mg m ⁻³ (Fig.
258	S2), the frequency of chl <i>a</i> values >0.15 mg m ⁻³ was high (Fig. 1b). In contrast, the
259	bloom frequency in the east of the main stream of the Kuroshio differed from the
260	distribution of the average chl a; algal blooms occurred frequently in the Ryukyu
261	Islands. Around the Miyako Islands, water of high bloom frequency was located to the
262	west of the islands, extending to the north.
263	
264	3.2. Region-wide environmental conditions, <i>Trichodesmium</i> spp., and
265	nitrogen fixation
266	The sea surface temperature (SST) ranged from 25.1–30.5°C at all of the stations
267	(Table S1), and there were no significant differences among the areas (p >0.05,
268	Tukey's honestly significant difference [HSD] test). The MLD varied from 12-60 m
269	at all of the stations, and was relatively deep around the Miyako Islands compared to

270	the other areas (Table S1). The surface N+N concentration varied between <3 and 42
271	nM, except around the Miyako Islands (Shiozaki et al., 2010, 2011) (Table S1). The
272	highest surface N+N concentration (374 nM) was observed at St. T0904 where
273	upwelling occurred (see below). No significant difference in the surface N+N was
274	observed among the four areas (p >0.05, Tukey's HSD test). The surface phosphate
275	concentration varied between <3 and 36 nM at all of the stations (Fig. 2a). The
276	phosphate concentrations at the surface and within the MLD was-were not
277	significantly different among the four areas (p >0.05, Tukey's HSD test). There was a
278	greater increase in the phosphate concentrations below 40-50 m in the ECS compared
279	to the other areas (Fig. 3a–d). Furthermore, the phosphate concentrations below 40–50
280	m near the Miyako Islands were higher than those in the Kuroshio and the Philippine
281	Sea, which were depleted down to 100 m, except at St. T1004 located near the
282	continental shelf. The N/P (= N+N/phosphate) ratio at the surface varied from 0.28 to
283	6.40 except at St. T0904 (N/P = 16.3) (Table S1), and no significant differences were
284	observed among the four areas ($p > 0.05$, Tukey's HSD test). The surface dissolved
285	iron concentration ranged from 0.19 to 0.89 nM at all of the stations (Fig 2b), with no
286	significant spatial differences among the four areas (p >0.05, Tukey's HSD test). The
287	surface dissolved iron concentrations at Sts. T0622 and T0907 was were elevated to

288	0.83 nM and 0.89 nM, respectively, with lower salinity water than in the adjacent
289	waters (salinity data are shown in Fig. 4a and Kodama et al., 2011). The nMDS
290	showed that the environmental variables at all stations were the same at the >80%
291	similarity level and were >90 % similar excepting station T0904 (Fig. 5). The
292	ANOSIM indicated no significant differences among the stations ($p > 0.05$).
293	The abundance of Trichodesmium spp. was highest at the surface at almost all of
294	the stations during the Nagasaki-maru 242 and KT-07-21 cruises (Fig. S3). The
295	surface Trichodesmium spp. abundances were positively correlated with the
296	depth-integrated abundances ($r^2 = 0.51$, $p < 0.05$) (Fig. 6a). Thus, the surface
297	abundance was used to discuss the geographical distribution of Trichodesmium spp.
298	The Trichodesmium spp. abundance at the surface varied widely, and there was no
299	significant difference among the four areas $(p > 0.05, \text{Tukey's HSD test})$.
300	Trichodesmium spp. were observed at all of the stations in the Kuroshio and around
301	the Miyako Islands, whereas they were not always observed in the ECS and the
302	Philippine Sea (Fig. 2c). The average surface abundance in the Philippine Sea was the
303	lowest among all of the areas (Table 1). The highest abundance of Trichodesmium spp.
304	(>20000 filaments L^{-1}) was observed near the Miyako Islands at St. T0906, where
305	they bloomed (see below). Tuft-shaped colonies were found at Sts. T0706, T0723,

CK-10, and T0906. The nitrogen fixation rate was highest in the upper 25% light 306 depth, and decreased with increasing depth at all of the stations (Fig. 3e-h). The 307surface rates were positively correlated with the depth-integrated rates ($r^2 = 0.79$, $p < 10^{-10}$ 308 0.05) (Fig. 6b), suggesting that the distribution of nitrogen fixation was indexed by 309 310 the surface activity. Surface and depth-integrated nitrogen fixation ranged from 0.54 to 62 nmol N $L^{-1} d^{-1}$ and from 29.5 to 753 µmol N $m^{-2} d^{-1}$, respectively (Fig. 2d and 311 312Table S1). Surface nitrogen fixation in the Philippine Sea was significantly lower than that in the Kuroshio (p < 0.05, *t*-test). 313The surface abundance of *Trichodesmium* spp. in the entire study area was 314 positively correlated with the nitrogen fixation rate at the surface ($r^2 = 0.80$; p < 0.05315 $[r^2 = 0.52; p < 0.05 \text{ if the datum taken at the Trichodesmium-bloom station T0906 is}]$ 316 excluded]) (Fig. 6c), suggesting that they significantly contributed to nitrogen fixation 317318 in the study region. However, active nitrogen fixation occurred in the ECS where Trichodesmium abundance was low, and hence, the other diazotrophs could also be 319important for nitrogen fixation. 320

321

322 3.3. Observation around the Miyako Islands during the KT-09-17 cruise

323 The SST was lower to the northwest of the Miyako Islands than in adjacent

324	waters, and chl a was enriched in the same location (Fig. 4b,c). Therefore, the
325	enhanced productivity was probably due to nutrient supply by upwelling. This
326	upwelling generally occurs in the lee of islands (Hasegawa et al., 2009), suggesting
327	that there was a northward current during the cruise. The surface salinity was lower
328	east of the Miyako Islands than in the surrounding waters (Fig. 4a). The absence of
329	any large river on the east side of Miyako-jima Island and the separation of low
330	salinity water from the island suggest that the low salinity was caused by rainfall.
331	St. T0904 was located near the upwelling water; its SST of 29.0°C was lowest and its
332	surface N+N concentration of 374 nM was highest among all of the stations. However,
333	the N+N concentration at St. T0904 at the surface was higher than that at the
334	subsurface (an approximate depth of 50 m; Fig. S4), indicating that St. T0904 was not
335	located in the middle of the upwelling. At St. T0904, the surface phosphate
336	concentration was also highest (23 nM) and the N/P ratio (=16.3) was higher than the
337	Redfield ratio. With the exception of the surface at St. T0904, the phosphate
338	concentration was low ($<3-9$ nM) in the upper 50 m, with no noticeable variation
339	among the stations (Fig. 2a). The dissolved iron concentration varied between 0.19
340	and 0.89 nM at the surface (Fig. 2b). The highest dissolved iron concentration was
341	observed at St. T0907.

342	During the same cruise, we encountered a Trichodesmium sppbloom at St.
343	T0906 (Fig. 2c), which had colored water at the surface. The abundance of
344	Trichodesmium spp. at St. T0906 was >20,000 filaments L^{-1} , which was far higher
345	than that at other stations (2–102 filament L^{-1}). The nitrogen fixation rate at the
346	surface (61.9 nmol N $L^{-1} d^{-1}$) of this station was more than 30-fold that just below the
347	surface, and was the highest among all of the stations (Fig. 3h). The diatom
348	abundance was markedly higher at St. T0904 than that at the other stations.
349	Cylindrotheca closterium was the most numerically dominant diatom (59%), followed
350	by Navicula spp. (23%) and Nitzschia spp. (13%). C. closterium was not detected at
351	the other stations, indicating that the high chl a induced by the island wake effect
352	mainly consisted of diatoms.
353	
354	3.4. Numerical simulation
355	As the Kuroshio generally flows along the continental slope north of the Miyako
356	Islands (Fig. 1b), particles around the Miyako Islands were not transported along the
357	typical path of the Kuroshio to the northeast, especially at their initial stages (Fig. 7a).
358	Some particles migrated around the Miyako Islands, or turned south after they passed
359	the Tokara Strait. Nevertheless, the particles delivered to Area K east of the Tokara

360	Strait increased as time elapsed, and the ratio of particles delivered to Area K to the
361	total released particles ranged from 13–56% (30 \pm 16%) by day 120 in 2003–2009
362	(Fig. 7b). The year-to-year variations in the ratio are mainly due to influences of
363	mesoscale eddies as partly seen in the particle trajectories in Fig. 7a, and likely
364	occurred over relatively short time scales (shorter than the seasonal time scale). This
365	is supported by another series of experiments in which particles were released on June
366	1, 11, and 21, and July 1 in 2009, which yielded ratios of 6.2–38% ($22 \pm 13\%$) by day
367	120 (Fig. S5).

369 **4. DISCUSSION**

370 4.1. Distribution of phosphate and dissolved iron concentrations

Phosphate concentrations were consistently low within the MLD in all of the studied areas, and the maximum abundance of *Trichodesmium* spp. and total nitrogen fixation activity generally occurred near the surface, suggesting that the phosphate conditions for surface *Trichodesmium* spp. and other diazotrophs were similar among all of the areas. Furthermore, with the exception of St. T1004 located near the continental shelf, the vertical distribution of phosphate in the Kuroshio was analogous to that in the Philippine Sea. Therefore, at least in the oceanic region of the two areas, phosphate availability for *Trichodesmium* spp. and the other diazotrophs was similar
throughout the water column.

380	The surface distribution of the dissolved iron concentration demonstrated no
381	significant variation among the areas. The dissolved iron concentration (0.19-0.89
382	nM) was higher than that -in the western North Pacific subtropical region (0.15-0.4
383	nM) (Brown et al., 2005). Obata et al. (1997) demonstrated that the vertical
384	distribution of the dissolved iron concentration in the ECS showed two peaks (at the
385	surface and in the deep water), suggesting that aerial dust significantly contributes to
386	the high dissolved iron concentration at the surface in all of our study areas. In
387	accordance with our results, previous modeling studies estimated the amount of dust
388	deposition to be similar in all four areas (Jickells et al., 2005; Mahowald et al., 2009).
389	Therefore, iron availability for Trichodesmium spp. and the other diazotrophs was also
390	likely similar across all of the study areas. Iron can be supplied from deep water to the
391	surface by mixing processes (Johnson et al., 1999). However, if this were the case, the
392	nitrate concentration would be expected to increase simultaneously at the surface
393	(Johnson et al., 1999), and we observed no noticeable elevation in N+N in any of the
394	areas, except at St. T0904. High concentrations of dissolved iron (>0.8 nM)
395	corresponded with low salinity at Sts. T0622 and T0907, suggesting that wet

deposition was an important process for iron supply. Dry deposition could also be
important since the iron-enriched water at Sts. T0601 and T0715 did not correspond
with low salinity.

Satellite data analysis indicated that there was a "pipeline" of material transport 399 from the Miyako Islands to the Kuroshio, and this was supported by numerical 400 401 simulations. According to the hypothesis of Marumo and Asaoka (1974), the growth 402 of Trichodesmium in the Kuroshio could be maintained by the supply of iron and phosphorus from the islands situated along the Kuroshio, and the Miyako Islands 403 were considered a possible nutrient source to the Kuroshio. Hence, assuming this 404 hypothesis to be valid, the iron and phosphate concentrations near the Miyako Islands 405(especially in our observed area) would be expected to be higher than those in the 406 other areas. However, we observed no significant difference in the iron and phosphate 407408 concentrations among the four areas. This suggested that there was no detectable washout of iron and phosphorus from the Miyako Islands during our observations, or 409 410 that diazotrophs and other phytoplankton exhausted the nutrient supply close to the 411 islands.

412

413 4.2. Factors controlling the distributions of *Trichodesmium* spp. and

414 nitrogen fixation

415	Although there was no statistically significant difference in Trichodesmium spp.
416	abundance among the study areas probably because the data were limited and the
417	variation was large, Trichodesmium spp. were always observed in the Kuroshio and
418	were abundant at most stations. Furthermore, at St.CK-10 in the East China Sea which
419	is in the Kuroshio branch current, a high abundance of Trichodesmium spp. was
420	observed. On the other hand, Trichodesmium spp. abundance in the Philippine Sea
421	tended to be lower than in the other areas. Such Trichodesmium distribution was also
422	reported in the previous study (Marumo and Asaoka, 1974). The present study also
423	showed lower surface nitrogen fixation in the Philippine Sea compared to that in the
424	Kuroshio ($p < 0.05$, t-test). Previous studies demonstrated that Trichodesmium spp.
425	flourished in some regions of the subtropical ocean where the iron levels were high
426	(Moore et al., 2009; Shiozaki et al., 2014b), which can be attributed to the high iron
427	requirement of Trichodesmium spp. for their growth compared to other diazotrophs
428	and non-diazotrophs (Kustka et al., 2003; Saito et al., 2011). Therefore, the
429	distribution of Trichodesmium spp. in the study area was expected to be associated
430	with the dissolved iron concentration at the surface. Furthermore, the iron-enhanced
431	active nitrogen fixation causes phosphorus depletion, and is consequently limited by

432	phosphorus (Mather et al., 2008). No significant differences in surface iron and
433	phosphate were observed among the study areas, which cannot explain the
434	distribution of <i>Trichodesmium</i> spp. and nitrogen fixation in the study region.
435	Johnson et al. (1999) reported that the iron supply increased around the
436	continental shelf because re-suspension from the bottom to the euphotic zone
437	becomes significant. However, in the continental shelf of the ECS, the abundance of
438	Trichodesmium spp. and nitrogen fixation were low (Marumo and Asaoka, 1974;
439	Zhang et al., 2012). Zhang et al. (2012) suggested that the low nitrogen fixation in the
440	continental shelf was attributable to mixing processes and the influence of the
441	Changjiang River. Turbulence near the sea floor influences the surface water in the
442	shallower bottom region (Matsuno et al., 2006), and Zhang et al. (2012) suggested
443	that the physical disturbance reduces diazotrophy since diazotrophs including
444	Trichodesmium favor calm seas. Furthermore, the water in the continental shelf of the
445	ECS is strongly influenced by the Changjiang River. The N/P ratio of the Changjiang
446	River plume is significantly higher than the Redfield ratio, which results in
447	phosphorus limitation, and can contribute to the low nitrogen fixation (Zhang et al.,
448	2012). In the present study, despite the fact that the surface phosphate concentration
449	was low throughout the study areas, the N/P ratio was generally lower than the

450	Redfield ratio, suggesting that biological production was limited by the availability of
451	nitrogen compared to phosphate (Moore et al., 2008, 2013). Furthermore, the
452	insignificant difference in MLD among the ECS, the Kuroshio, and the Philippine Sea
453	(p >0.05; Tukey HSD test) indicated similar vertical mixing conditions. Therefore, the
454	environmental variables related to nitrogen fixation only slightly differed as
455	demonstrated by the nMDS plot.
456	In our study, we found a Trichodesmium spp. bloom near the Miyako Islands.
457	Recent studies demonstrated that Trichodesmium spp. thrived near oceanic islands
458	(Shiozaki et al., 2010, 2014c; Dupouy et al., 2011). Given that some aspect of the
459	environment around the islands increases Trichodesmium spp. abundance and that
460	they are transported from the islands to the Kuroshio, this can explain why the
461	Trichodesmium distribution was not estimated from environmental variables.
462	Accordingly, the low abundance of Trichodesmium spp. in the Philippine Sea was
463	likely due to the low density of islands. Furthermore, higher nitrogen fixation in the
464	Kuroshio than in the Philippine Sea might be explained in the same manner.
465	Trichodesmium is a major nitrogen fixer in the Kuroshio (Chen et al., 2008, 2014;
466	Shiozaki et al., 2014a), and our results showed that the bulk water nitrogen fixation
467	was positively correlated with Trichodesmium abundance.

468	The numerical simulation demonstrated that released particles from the Miyako
469	Islands were generally transported to the northeast and flowed along the Kuroshio
470	during summer between 2003 and 2009. Thus, if Trichodesmium increases and active
471	nitrogen fixation usually occurs around the Miyako Islands, the water would be
472	delivered to the Kuroshio. Furthermore, we performed additional particle tracking
473	experiments whose particle release points were set at major islands in the Ryukyu
474	Islands (Amami Islands, Okinawa Main Island, and the Ishigaki Islands) (Figs. S6 and
475	S7). The results demonstrated that the particles released from the other islands of the
476	Miyako Islands were also delivered to the Kuroshio, with some exceptions. Based on
477	the calculations for 2003–2009, 13–56% (30 \pm 16%) of particles released from the
478	islands reached Area K by day 120 (Fig. S7).
479	Studies on nitrogen fixation around islands in the study region are fairly limited
480	(Liu et al., 2013), and the present study is the first report of a Trichodesmium bloom
481	around islands in the area. The Miyako Islands are surrounded by reefs, and studies
482	have shown that Trichodesmium blooms can be associated with reef environments
483	(Bell et al., 1999; McKinna et al., 2011). However, the factors causing the
484	Trichodesmium blooms around islands are not well understood (Shiozaki et al.,
485	2014c). Further studies are required to identify which characteristics of the near island

486 environment are important for the growth and/or accumulation of *Trichodesmium* and487 other diazotrophs.

488

489 5. CONCLUSIONS

We hypothesize that the high abundance of *Trichodesmium* spp. and active 490 nitrogen fixation in the Kuroshio were ascribable not to the unique nutrient 491 492 environment, but rather to the supply of Trichodesmium spp. and other diazotrophs from the surrounding islands. The Ryukyu Islands would not be the only islands with 493 abundant *Trichodesmium* spp., as *Trichodesmium* spp. also flourish in the upstream 494 Kuroshio near Luzon Island (Chen et al., 2008). Therefore, the abundance of 495Trichodesmium spp. would be generally increased around islands situated along the 496 Kuroshio, and the abundant Trichodesmium spp. would likely be transported to the 497498 mainstream of the Kuroshio. Trichodesmium is a major diazotroph in the Kuroshio (Chen et al., 2008, 2014; Shiozaki et al., 2014a), and diazotrophy in the Kuroshio is 499500considered to influence the nutrient stoichiometry in the North Pacific (Shiozaki et al., 5012010). Thus, our results indicate that phenomena around the islands located along the Kuroshio are important for determining the partial nitrogen inventory in the North 502Pacific. 503

Author Contributions 505

506	T.S., S.T., S.I., and K.F. designed the experiment and T.S., S.T., T.K., X.L., F.H., and
507	K.F. collected the samples at sea. T.S. determined nitrogen fixation and abundance of
508	Trichodesmium spp. during the KT-06-21, KT-07-21, KT-09-17, and Nagasaki-maru
509	242 cruises, and X.L. did during the KT-10-19 cruise. T.S. analyzed datasets of
510	satellite and climatological current field. S.T. analyzed concentration of dissolved iron.
511	S.I. performed numerical experiments. T.K. and F.H. determined nutrient
512	concentration. T.S. prepared the manuscript with contributions from all co-authors.
513	

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Area	<i>Trichodesmium</i> [*] [filaments l ⁻¹]	N_2 fixation [µmolN L ⁻¹ d ⁻¹]	Temperature [*] [°C]	MLD [m]	NO ₃ ⁻ +NO ₂ ^{-*,†} [nM]	PO ₄ ^{3-*,†} [nM]	DFe [*] [nM]
East China Sea	21±58	170±140	28.5±1.2	24±12	19±11	15±9	0.76±0.18
Kuroshio	43±33	199±142	29.4±0.81	27±8	9±8	15±7	0.45±0.13
Philippine Sea	8±8	58.3±25.1	29.4±0.1	23±3	8±3	14±19	0.51±0.25
Miyako Islands	3019±8478	201±274	29.3±0.3	40±12	61±128	8±7	0.38 ± 0.24

678 fixation and its related parameters in the four representative study areas.

679 * values in surface water

[†]When the concentration was below the detection limit (3 nM), we assumed a concentration of 3 nM to

681 calculate the mean.



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Figure 1. (a) Sampling stations during the KT-06-21 (circles), KT-07-22 (inverted 685triangles), KT-09-17 (crosses), KT-10-19 (diamonds), and 242 (triangles) cruises. 686Symbols of stations located in the East China Sea, the Kuroshio, the Philippine Sea, 687 688 and near the Miyako Islands are indicated in purple, light blue, orange, and black, respectively. (b) Climatological surface current fields during summer (1953–2008) 689from geoelectrokinetograph measurements and ship-mounted ADCP data. The 690 background contour represents the percentage of chlorophyll a of >0.15 mg m⁻³ 691 during summer between 2003 and 2009. Dashed lines indicate 200 m isobaths. 692



Figure 2. Distribution of (a) phosphate, (b) dissolved iron, (c) *Trichodesmium* spp., and (d) nitrogen fixation at the surface. The parameters in the small boxes indicate results from the KT-09-17 cruise. The areas of the circles are proportional to the concentration, abundance, or activity.



Figure 3. Vertical profiles of phosphate and nitrogen fixation in the East China Sea (a
and e), the Kuroshio (b and f), the Philippine Sea (c and g), and the Miyako Islands (d
and h).





Figure 4. Surface (a) salinity, (b) temperature, and (c) chlorophyll a during the





713 Figure 5. nMDS ordination of sampling stations with environmental variables



Figure 6. Relationships (a) between surface and depth-integrated *Trichodesmium* spp.

abundance, (b) between surface and depth-integrated nitrogen fixation rates, and (c)

519 between *Trichodesmium* spp. abundance and nitrogen fixation rate at the surface.



Figure 7. (a) Trajectories of particles released from points around the Miyako Islands
on June 1, 2003–2009. (b) The ratio of particles delivered to Area K to the total
released particles.