

Abstract

Continuous permafrost zones with well-developed polygonal ice-wedge networks are particularly vulnerable to climate change. Thermo-mechanical erosion can initiate the development of gullies that lead to substantial drainage of adjacent wet habitats. How vegetation responds to this particular disturbance is currently unknown but has the potential to strongly disrupt function and structure of Arctic ecosystems. Focusing on three major gullies of Bylot Island, Nunavut, we aimed at estimating the impacts of thermo-erosion on plant community changes. We explored over two years the influence of environmental factors on plant species richness, abundance and biomass in 62 low-centered wet polygons, 87 low-centered disturbed polygons and 48 mesic environment sites. Gullying decreased soil moisture by 40 % and thaw-front depth by 10 cm in the center of breached polygons within less than five years after the inception of ice wedge degradation, entailing a gradual yet marked vegetation shift from wet to mesic plant communities within five to ten years. This transition was accompanied by a five times decrease in graminoid above-ground biomass. Soil moisture and thaw-front depth changed almost immediately following gullying initiation as they were of similar magnitude between older (> 5 years) and recently (< 5 years) disturbed polygons. In contrast, there was a lag-time in vegetation response to the altered physical environment with plant species richness and biomass differing between the two types of disturbed polygons. To date (ten years after disturbance), the stable mesic environment cover has still to be reached. Our results illustrate that wetlands are highly vulnerable to the process of thermo-erosion, which can drive landscape transformation on a relative short period of time for High Arctic perennial plant communities (five to ten years). Such succession towards mesic plant communities can have substantial consequences on the food availability for herbivores and methane emissions of Arctic ecosystems.

1 Introduction

Warming in the Arctic is occurring twice as fast as the global average (USGCRP, 2009; New et al., 2011; NOAA, 2014). Perennially frozen ground (permafrost) has consequently warmed by 2 °C over the last 20 to 30 years (Christiansen et al., 2010; Romanovsky et al., 2010), and there is now evidence of a decrease in both permafrost area extent across the northern hemisphere and permafrost thickness at the local scale (Beilman and Robinson, 2003; Payette et al., 2004; Couture and Pollard, 2007; Smith, 2011).

Permafrost is tightly associated with biophysical components such as air temperatures, soil conditions, surface water, groundwater, snow cover and vegetation (Jorgenson et al., 2010; Sjöberg, 2015). Permafrost impedes water to drain to deeper soil layers and maintains a perched water table and saturated soils which favors the existence of wetlands (Woo, 2012; Natali et al., 2015). Permafrost degradation that would increase subsurface drainage and reduce the extent of lakes and wetlands at high latitudes (Avis et al., 2011; Jorgenson et al., 2013; Beck et al., 2015) would thus have major consequences on ecosystem structure and function (Collins et al., 2013; Jorgenson et al., 2013). It would also strongly influence variations of active layer depths (Wright et al., 2009; Shiklomanov et al., 2010; Gangodagamage et al., 2014), as illustrated by long-term monitoring sites throughout the circumpolar North (Tarnocai et al., 2004; Nelson et al., 2008; Smith et al., 2009; Shiklomanov et al., 2010).

Several forms of ground and massive ice can be found within permafrost (Rowland et al., 2010), especially ice wedges in regions where winter temperatures enable thermal contraction cracking (Fortier and Allard, 2005; Ewertowski, 2009; Kokelj et al., 2014; Jorgenson et al., 2015b; Sarrazin and Alard, 2015). Continuous permafrost zones with well-developed polygonal ice-wedge networks are particularly vulnerable to climate change because ice wedges are usually found near the top of permafrost (Smith et al., 2005;

Jorgenson et al., 2006; Woo et al., 2008; Vonk et al., 2013). In these regions, thawing permafrost can result in ground ice erosion and displacement of sediments, carbon and nutrients by drainage (Rowland et al., 2010; Godin et al., 2014; Harms et al., 2014). This thermo-erosion process has especially been observed across North-America (Grosse et al., 2011), in Siberia (Günther et al., 2013) and in the Antarctic Dry Valleys (Levy et al., 2008). On Bylot Island in Nunavut, thermo-mechanical erosion by water has initiated permafrost tunneling and the development of gully networks in aeolian, organic and colluvial depositional environments of nearly 158 000 m² (Fortier et al., 2007; Godin and Fortier, 2012a; Godin et al., 2014; Veillette et al., 2015). A fine-scale spatio-temporal monitoring of the six largest gullies showed that their development rate ranged from 14 to 25 m.yr⁻¹ going up to 80 m.yr⁻¹ during their inception (Godin and Fortier, 2012b), leading to substantial changes in the drainage network and increasing eroded area throughout the valley (Godin et al., 2014). To date, the contribution of thermal-erosion and lateral erosion processes in permafrost feedbacks to climate is still poorly documented (Turetsky et al., 2015; Fortier et al., 2015).

Many observational and experimental studies have documented shifts in tundra plant community structure and plant species productivity in response to warming temperatures (Jonsdottir et al., 2005; Hudson and Henry, 2010; Epstein et al., 2013; Naito and Cairns, 2015). Substantial work has also been devoted to the vegetation response to disturbances such as landslides associated with permafrost degradation (Ukrainitseva et al., 2003; Walker et al., 2009; Cannone et al., 2010). In contrast, little is known about how thermo-erosion gullying affects plant community structure and plant species abundance. Yet, this information is urgently needed as vegetation plays an important role in structuring Arctic ecosystems and regulating permafrost response to climate change (Jorgenson et al., 2010; Gauthier et al., 2011; Legagneux et al., 2012). Wetlands serve as preferred grounds for Arctic herbivores

such as snow geese (Gauthier et al., 1996; Massé et al., 2001; Doiron et al., 2014). They are also expected to produce more methane compared to shrub-dominated areas (Olefeldt et al., 2013; Nauta et al. 2015; Treat et al., 2015).

The present study aimed at examining plant community patterns following thermo-erosion gullyng. Bylot Island, where geomorphological and ecological processes in response to climate change have been monitored for over two decades (Allard, 1996; Fortier and Allard, 2004; Gauthier et al., 2013), offered a unique opportunity to specifically assess the response of wetlands to gullyng. The following questions were addressed: (1) to what extent thermo-erosion gullyng modifies environmental conditions of low-centered wetland polygons? (2) how do plant communities cope with these geomorphological changes, i.e. do we observe shifts in plant diversity, abundance and productivity?

2 Methods

2.1 Study area

This study took place in the Qarlikturvik valley of Bylot Island, Nunavut, Canada (73°09'N, 79°57'W; Fig. 1a). Bound to the North and South by plateaus < 500 m asl, it connects C-79 and C-93 glaciers to the Navy Board Inlet sea via a proglacial river. The sampling sites were specifically located in the valley floor (*ca* 65 km²), which is characterized by a low-centered polygon landscape that has resulted from ice wedge development and sediment accumulation during the late Holocene (Ellis and Rochefort, 2004; Fortier and Allard, 2004; Ellis et al., 2008). Two baseline vegetation types can be recognized. Wetlands, often associated with intact low-centered polygons, represent *ca* 23 % of the valley area (Hughes et al., 1994) and are dominated by sedges (*Carex aquatilis*, *Eriophorum angustifolium*, *Eriophorum scheuchzeri*), grasses (*Dupontia fisheri*, *Pleuropogon sabinei*; Gauthier et al., 1995) and fen mosses (*Drepanocladus* spp.; Ellis et al., 2008; Pouliot et al., 2009). Mesic environments,

such as low-centered polygon rims, gently sloping terrain and hummocky tundra, support a more diverse group of species including *Salix* spp., *Vaccinium uliginosum*, *Arctagrostis latifolia*, *Poa arctica* and *Luzula confusa* with *Aulacomnium* spp. as dominant moss species (Zoltai et al., 1983). As a result, the valley houses many herbivores such as snow geese (in summer) and brown and collared lemmings (Gauthier et al., 1995; Gruyer et al., 2008), thus representing a critical environment for tundra food web (Gauthier et al., 2011; Legagneux et al., 2012).

2.2 Field sites

Our work was specifically conducted around three gullies that were selected among the 36 identified in the valley (Godin and Fortier, 2012b). These gully networks have originated from snowmelt water infiltration into cavities of the frozen active layer and the subsequent formation of underground tunnels that have ended up collapsing (Fortier et al., 2007; Godin and Fortier, 2010). The gullies R08p and R06, respectively 835 and 717 m long, are characterized by ongoing thermo-erosion (Fortier et al., 2007; Godin and Fortier, 2012b) whilst the gully RN08, 180 m long, has not been active in recent years. A total of 197 sampling sites were randomly selected around the three gullies (Table 1; Fig. 1b) and classified into one of four categories (referred hereafter as habitats) that represented the two baseline vegetation types (wet and mesic) as well as increasing levels of disturbance related to thermo-erosion. The disturbed habitats were sorted via a visual assessment of the low-centered polygon rim integrity coupled with a recent close monitoring of drainage system development along the gullies (Fortier et al., 2007; Godin and Fortier, 2012a,b). The habitats were defined as follows: (i) intact low-centered wet polygons (n = 62) that were not affected by gullying. Their elevated rims enclose a central depression that retains snow cover during winter and is flooded by snowmelt water during spring (Woo and Young, 2006; Minke et al.,

2007). These polygons are hydrologically independent, with their water content representing the balance between precipitation inputs (snow and rain) and evapotranspiration outputs (Fortier et al., 2006); (ii) less than five-year (recently) disturbed polygons (n = 44), located along the most recent sections of the gullies, with partially degraded rims, incomplete drainage and heterogeneous water content; (iii) more than five-year (older) disturbed polygons (n = 43), with heavily breached rims adjacent to the gullies and substantial or complete drainage; (iv) mesic environments (n = 48), with distinct heterogeneous mesic vegetation. They are found on the rim of polygons and in adjacent areas and are generally not induced by thermo-erosion gullying but rather dependent of local sedimentary and hydrological dynamics.

2.3 Environmental condition monitoring

Daily precipitation was recorded with a manual rain gauge throughout summer 2010 at the base camp, located 700 m west of the gully R08p (Gauthier et al., 2010). Soil (top 10 cm) moisture was recorded at the center of each sampling site using ECH₂O EC-5 moisture sensors (accuracy of ± 3 % VWC, resolution of ± 0.1 % VWC) connected to Em5b dataloggers (Decagon Devices, WA, USA). Measurements were carried out in 2010, both July 5 (early season) and July 30 (late season) along the gullies R08p and R06, and July 30 only along the gully RN08. Thaw-front depth was recorded at the center of each sampling site using a graduated steel rod driven into the thawed active layer. The data collection spanned two years with measurements in July 2009 and 2010 in the polygons of the gullies R08p and R06 and in July 2010 in those situated along the gully RN08. One measure was taken in each of the wet and disturbed sampled polygons whereas three random measurements were conducted in each mesic environment site because of the heterogeneity of this type of habitat. In this case, a mean per site was calculated prior to analyses.

2.4 Plant community characterization

Species richness and abundance were determined at each site in July 2009 or 2010 using three randomly placed 70 cm x 70 cm quadrats that were vertically photographed at ca 1.3 m from the ground (see detailed protocols in Chen et al., 2010 and the IPY CiCAT, 2012; Fig. 2). Abundances of vascular plants, lichens, mosses, *Nostoc* spp., fungi, cryptogamic crust, bare ground, litter, vascular plant standing dead, standing water, signs of grubbing and goose feces were evaluated as cover percentages using photography analyses (Perreault, 2012). Daubenmire cover abundance classes (Daubenmire, 1959) were used on each quadrat picture overlain by a 7-cm grid to evaluate species cover as the projection on the ground of all species above-ground parts.

Five sampling sites per habitat were also randomly selected along the gullies R08p and R06 to measure above-ground biomass of graminoid species. At each site, an enclosure of 1 m x 1 m was made of chicken wire 30 cm high and supported by wooden stakes at each corner (see Gauthier et al., 1995). Enclosures were set up in early July 2012 to avoid any significant grazing by geese. Above-ground biomass was harvested inside the enclosures near peak production in early August 2012 using random grids of 25 cm x 25 cm for wet and recently disturbed polygons and of 50 cm x 20 cm for older disturbed polygons and mesic environments. Two different grid sizes were used because of the difference in structure of the vegetation (herbaceous vs. shrubs) associated with the habitat heterogeneity (Legagneux et al., 2012). All graminoids present in the random grids were cut to a standard height, i.e. at an average of 1 cm below the moss surface (Gauthier et al., 1995; Doiron et al., 2014), and live biomass was sorted as follows: *Carex aquatilis*, *Eriophorum angustifolium* and *Eriophorum scheuchzeri* (Cyperaceae); *Luzula arctica* and *Luzula confusa* (Juncaceae); *Anthoxanthum arcticum*, *Arctagrostis latifolia*, *Dupontia fisheri* and *Festuca brachyphylla* (Poaceae).

Biomass was then oven-dried at 65°C until constant dry weight and further weighed to ± 0.0001 g using an electronic weighing scale.

2.5 Statistical analyses

Differences in soil moisture, thaw-front depth and graminoid above-ground biomass among habitats were tested with a generalized linear mixed model (procedure MIXED, REML method in SAS, version 9.4, SAS Institute, Cary, NC, USA). Soil moisture, thaw-front depth as well as date or year of measurements and the interaction terms were treated as fixed factors and gully as a random factor. Type III sums of squares were used for the calculation of fixed effect *F* statistics while random effects were assessed using a log likelihood ratio test from the full and reduced models (Littell et al., 2006). Post hoc contrasts were performed to ascertain specific differences between habitats at $\alpha < 0.05$ using the LSMEANS statement and Bonferroni adjustment. Canonical Correspondence Analyses (CCA) were conducted to test unimodal relationships between habitats and environmental variables (ter Braak, 1986; Zuur et al., 2007) using the ‘vegan’ package (Oksanen et al., 2015) in R 3.1.3 (R development core team, 2015). Two matrices were elaborated: one of 197 sites x 65 taxa using their mean cover per site, and another of 197 sites x 8 environmental characteristics comprising the following continuous variables: soil moisture, thaw-front depth, litter cover, bare ground, grubbing, vascular plant standing dead, standing water and goose feces. Soil moisture measurements obtained on July 30 2010 were used in the analyses.

3 Results

3.1 Environmental conditions

In 2009, above-average spring temperatures led to a rapid snowmelt (June 16) while summer was one of the driest on record (Gauthier et al. 2009). In 2010, despite a relatively warm

spring (0.26°C above normal), the high snow pack at the end of the winter (41.6 cm on May 31) delayed the snowmelt in the lowlands to June 28, which was a week later than normal. Summer was characterized by warm and sunny conditions as well as below-average precipitations (cumulative rain of 84 mm vs long-time average of 92 mm; Gauthier et al. 2010). However, the 36 mm received in only five days in mid-July 2010 significantly increased soil moisture between the two monitoring dates in all habitats (significant date effect: $df = 1$, $F = 88.99$, $P < .001$; Fig. 3). There was overall a significant difference in soil moisture among habitats ($df = 3$, $F = 79.86$, $P < .001$), which was associated to differences between wet polygons and the other habitats (5 July: $df = 3$, $F = 33.41$, $P < .001$, 30 July: $df = 3$, $F = 47.36$, $P < .001$; Fig. 3, Table A1). Throughout the summer, soil moisture was approximately 40 % higher in wet polygons compared to disturbed polygons and mesic environments. We also found a significant difference in thaw-front depth between wet polygons and the other habitats (2009: $df = 3$, $F = 21.30$, $P < .001$, 2010: $df = 3$, $F = 33.86$, $P < .001$; Fig. 4, Table A1). Thaw-front depth of wet polygons was approximately 10 cm deeper than in disturbed polygons and mesic environments for both years of monitoring (35-36 vs 25-27 cm). The two types of disturbed polygons thus showed similar soil moisture and thaw-front depth despite their differing time since disturbance inception. Results on the environmental conditions of the 197 sampling sites were obtained regardless of the gully location (non-significant gully effect: $df = 2$, $LLR = 1.6$, $P = 0.21$ for soil moisture, $df = 2$, $LLR = 0.0$, $P = 1.0$ for thaw-front depth).

3.2 Plant community characterization

A total of 18 vascular plant families encompassing 59 species were sampled throughout the study (Table A2). The greatest species richness was found in polygons that were disturbed at least five years ago and where both hydrophilic and mesic species were present (Table 2). The

transition from wet polygons to mesic environments was accompanied by significant changes in vascular plant community composition, especially with the decline in Cyperaceae and Poaceae cover and the emergence of Salicaceae species (Table 2). *Carex aquatilis* and *Dupontia fisheri* were respectively present in 100 % and 93 % of the wet polygons sampled where they accounted for 52 % and 26 % of the total vascular plant cover. They were found in only 47 % and 16 % of mesic environments accounting for 9 % and 0.45 % of the total vascular plant cover. In contrast, *Salix arctica* and *Arctagrostis latifolia*, which were found in approximately half of the wet polygons accounting for 3 % of the total vascular plant cover, were present in 98 % of the mesic environments where they respectively accounted for 50 % and 14 % of the total vascular plant cover. Differences among habitats were also noted in non-vascular taxa. Abundance of lichens such as *Cladonia* spp., *Stereocaulon* spp. and *Peltigera* spp. increased in polygons disturbed for at least five years and mesic environments (Table 2). Mosses were mostly found living in wet polygons and mesic environments and dried (i.e. dead) in disturbed polygons (Table 2), while a significant species turn-over was observed between habitats with wetland-dominant *Drepanocladus* species replaced by *Aulacomnium* species in mesic environments.

Moreover, we observed vegetation changes through the decline of graminoid above-ground biomass which varied significantly among habitats ($df = 3$, $F = 11.59$, $P < 0.001$; Fig. 5a; Table A1). Graminoid biomass was nearly five times greater in wet than in mesic environments (29.2 vs 5.9 g.m⁻²) and decreased twofold between < 5-year disturbed and > 5-year disturbed polygons (22.3 vs 9.3 g⁻²; Fig. 2 and 5a). Differences were mainly driven by the decline of hydrophilic species, i.e. *Carex aquatilis*, *Eriophorum scheuchzeri*, *Anthoxanthum arcticum* and *Dupontia fisheri*, between wet and mesic habitats (28.8, 19.7, 3.6 and 2.5 g.m⁻² in wet, < 5-year disturbed, > 5-year disturbed and mesic polygons, respectively; Fig. 5b). Above-ground biomass of *Luzula* spp., *Arctagrostis latifolia* and *Festuca*

brachyphylla was contrastingly nine times greater in mesic than in wet habitats (3.30 vs 0.35 g.m⁻²) and twenty-four times greater in > 5-year disturbed than in < 5-year disturbed polygons (5.61 vs 0.23 g.m⁻²).

3.3 Relationships between plant communities and environmental variables

The first two axes of the Canonical Correspondence Analysis retained 14 % of the vegetation data variance and 80 % of the vegetation-environment relationship variance (Table 3). Five of the eight environmental variables tested were significant within the canonical model ($P < .05$, 999 permutations), but only three – litter cover, thaw-front depth and soil moisture - showed high correlations with the canonical axes (Table 4). Thaw-front depth and soil moisture were strongly related to the first axis, while litter cover was mainly associated with the second axis (Table 4). Altogether, these variables discriminated well the four studied habitats of the Qarlikturvik valley. Wet polygons were mainly related to high soil moisture and substantial thaw-front depth whilst mesic environments were associated with greater litter cover (Fig. 6). The gradual vegetation transition was also observed from < 5-year to > 5-year disturbed polygons following the soil moisture shift in these habitats (Fig. 6).

4 Discussion

Sustainability of wetlands at high latitudes essentially relies on perennial frozen ground that prevents drainage and allows wet soil conditions (Woo and Young, 2006; Ellis et al., 2008). However, snowmelt water run-off through ice-wedge polygon landscapes can initiate thermal erosion of the permafrost and the development of gullies (Fortier et al., 2007; Godin and Fortier, 2014). We showed here that permafrost gullying significantly altered wetlands by changing the original polygons' microtopography, and decreasing soil moisture and thaw-front depth of disturbed polygons along the gullies. Vegetation was sensitive to this process,

and mesic environment plant species gradually replaced hydrophilic species within five to ten years, although the full transition hasn't been reached yet. This vegetation turn-over can have substantial consequences on wildlife biology, permafrost stabilization and ecosystem-level greenhouse gas emissions (Blok et al., 2010; Doiron et al., 2014; Jorgenson et al., 2015b; McEwing et al., 2015).

4.1 Transient environmental conditions

The initiation of thermo-erosion gullyng led to a significant decrease in soil moisture and thaw-front depth of breached polygons. Both older and recently disturbed polygons had similar soil moisture and thaw-front depth while differing in time since disturbance, which shows that the change in polygon environmental conditions after permafrost disturbance was rapid. The decrease in soil moisture following polygon rim erosion is consistent with what has been previously observed in gullied areas (Seppälä 1997; Poesen et al., 2003; Godin and Fortier, 2012a; Godin and Fortier, 2014; Harms et al., 2014; Godin and Fortier, 2015) and concurs with a modeling analysis showing that the transformation of low-centered to high-centered polygon landscape following ice wedge degradation is accompanied by a significant alteration in the water balance partitioning (Liljedahl et al., 2012). In our study, all types of polygons were recharged by snowfall and summer rainfall, yet disturbed habitats had lower soil moisture than wet polygons and a thorough examination of moisture evolution throughout an entire summer showed that soil moisture of breached polygons was significantly more variable than that of wet polygons at both intra- and inter-polygonal scales (Godin et al., 2015). Given that soil moisture is an important driver of plant community composition (Muc et al., 1989; Dagg and Lafleur, 2011), it is no surprise that we observed a shift in vegetation following changes in moisture regime.

Decreasing soil moisture in the center of disturbed polygons came with decreasing

thaw-front depth, which was expected given that active layer thickness is closely related to soil moisture (Nelson et al., 1999; Hinzman et al., 2005; Minke et al., 2009; Wright et al., 2009; Gangodagamage et al., 2014). This result however contrasts with the active layer thickening generally observed in response to climate warming (Tarnocai et al., 2004; Woo et al., 2007; Akerman and Johansson 2008; Smith et al., 2009; Nauta et al., 2015), and this is likely due in part to ground surface subsidence and drainage which follows ice-rich permafrost thawing (Shiklomanov et al., 2013) and in part to snow accumulation patterns (Godin et al., 2015). Within five years of drainage, thaw-front depth in disturbed polygons decreased by 37 % compared to that in wet polygons. This is mainly explained by the greater heat transfer from water and higher thermal conduction rates in wetter polygons that provide substantial melt energy to the frost table (Nelson et al., 1997; Hinzman et al., 2005; Wright et al., 2009; Romanovsky et al., 2010). This effect is also sharpened by the low thermal conductivity of drier moss carpets in disturbed habitats (Wright et al., 2009) and altered local snow conditions within the polygons adjacent to the gullies (Godin et al., 2015).

4.2 Vegetation changes

Overall, the floristic composition of our sampling sites is in line with previous field surveys conducted in the same area (Gauthier et al., 1996; Duclos, 2002; Doiron, 2014). The presence of *Carex aquatilis*, *Eriophorum scheuchzeri* and *Dupontia fisheri* characterizes well the typical vegetation of Arctic wetlands (Jorgenson et al., 2013; Sandvik and Odland, 2014; Lara et al., 2015) whilst that of *Arctagrostis latifolia*, *Luzula* and *Salix spp* are common features of Arctic mesic environments (Audet et al., 2007; Sjogersten et al., 2008). Disturbed polygons were the most diverse habitats given that they offered a middle-range state between wet and mesic conditions where hydrophilic species were still present while mesic environment ones had successfully established.

The development of gullies in the Qarlikturvik valley and the subsequent drainage of adjacent low-centered polygons have led within five to ten years to a gradual change in plant communities with vegetation of disturbed polygons leaning toward a new equilibrium, that of mesic environments. Mesic environment species such as *Luzula* and *Salix spp* have established or increased in cover following the decrease in soil moisture and thaw-front depth and replaced hydrophilic Cyperaceae and Poaceae. The secondary succession pioneered here by the gullying process in disturbed polygons follows the directional-species replacement model examined by Svoboda and Henry (1987). However, by occurring within five to ten years, it has been remarkably more rapid than what is usually documented for the High Arctic where perennial plant communities are largely resistant to disturbance (Hollister et al., 2005; Jonsdottir et al., 2005; Hudson and Henry, 2010) and succession dynamics are slow due to short growing seasons and low summer temperatures (Svoboda and Henry, 1987). For instance, plant cover of northeastern Alaska changed little over a twenty-five year period despite a significant rise in summer temperatures (Jorgenson et al., 2015a). This gradual yet rapid species replacement has been triggered in our system by the hydrological and thermal shift caused by gullying and favored by the mosaic of wet and mesic habitats allowing for a substantial species pool with both vegetative and sexual reproduction.

In the canonical ordination analysis, the soil moisture gradient discriminated wet polygons from the other habitats as well as recently disturbed from older disturbed habitats. The 37 % decrease in soil moisture between wet and disturbed polygons represents a drastic change of conditions for plant communities and is of similar magnitude than what has been documented in Alaskan drying wetlands as a result of increasing temperatures (Klein et al., 2005). The strong influence of soil moisture in separating plant community types at high latitudes has indeed been well documented (Hinzman et al., 2005; Daniëls and de Molenaar, 2011; Daniëls et al., 2011; Sandvik and Odland, 2014). Four other variables significantly

influenced the distinction among habitats: (i) thaw-front depth discriminated habitats in the same direction than soil moisture with a 30 % decrease in disturbed polygons and mesic environments compared to wet polygons, which was expected since these two factors are closely related (see section 4.1); (ii) litter cover separated mesic environments from the others, which may be explained by increased organic matter related to greater shrub abundance in mesic environments (Zamin et al., 2014); (iii) vascular plant standing dead separated wet and recently disturbed polygons from the other habitats, which can be explained by the senescence of Cyperaceae tillers that are highly abundant at these locations (Fig. 5); (iv) goose feces were mainly associated with older disturbed and mesic environments. While this may suggest a higher use of these habitats by geese, the slower degradation of feces in dryer habitats cannot be ruled out; this has yet to be tested.

The shift in vegetation composition in disturbed polygons was accompanied by significant changes in biomass. Above-ground biomass of graminoids was the greatest in wet polygons, which is concordant with the fact that wetlands are the most productive habitats of forage plants in the Arctic (Sheard and Geale, 1983; Duclos, 2002; Doiron, 2014). It gradually decreased in disturbed polygons as conditions became closer to those of mesic environments. Compared to the immediate change in environmental conditions, we nonetheless observed a lag-time in vegetation response to thermo-erosion related disturbance as graminoid biomass differed significantly between recently and older disturbed polygons. In our study, graminoid above-ground biomass of wet polygons was 35 % lower than what Cadieux et al. (2008) found via a long-term plant monitoring on Bylot Island (45.2 g/m^2), and 62.7 % lower than what Gauthier et al. (2012) measured in the most productive wetlands of the Qarlikturvik Valley ($78.4 \pm 10.5 \text{ g/m}^2$). These contrasts may be explained by varying species composition and to a lesser extent by earlier plant harvesting in our case. Indeed, while we focused on wet polygons dominated by *Carex aquatilis*, Cadieux et al. (2008) and Gauthier et al. (2012)

worked on wet polygons dominated by *Dupontia fisheri* and *Eriophorum scheuchzeri*. Because our study was part of a large scale multisite project on wetland carrying capacity for snow geese (Legagneux et al., 2012; Doiron, 2014), we only focused on forage plant (i.e. graminoids) biomass and did not sample forbs or shrubs. Since above-ground biomass of graminoids account for more than 90 per cent of vascular plant biomass in wetlands (Gauthier et al., 1995), we provide here an accurate estimate of the total above-ground biomass that can be found in these habitats. However, the total above-ground biomass in mesic environments was probably underestimated. For instance, biomass of shrubs and forbs respectively ranged between 22 and 48 g.m⁻² and between 6 and 20 g.m⁻² in the mesic environments adjacent to our study area (E. Lévesque, *unpublished data*). Overall, total above-ground biomass in wetlands and mesic environments is of similar order of magnitude (50.5 g.m⁻² ± 2.8 SE in wetlands and 44.2 g.m⁻² ± 6.8 SE in mesic tundra for the period 2007-2009; Legagneux et al., 2012).

4.3 Impacts on ecosystems

It is likely that the replacement of hydrophilic plants by mesic vegetation will severely impact wildlife biology. The Qarlikturvik valley of Bylot Island represents an important summer habitat for greater snow geese (Legagneux et al., 2012). It is well documented that this species mostly relies on wetlands for food resources (Gauthier et al., 1995, 2011), especially because graminoids are easily digested due to their low fiber concentration and rich nutritive elements (Sedinger and Raveling, 1989; Manseau and Gauthier, 1993; Audet et al., 2007). For instance, geese removed respectively 40 % and 31 % of the total annual production of *Dupontia fisheri* and *Eriophorum scheuchzeri* during the period 1990-2007 (Cadieux et al., 2008). It remains to quantify the extent to which gullying alters wetland carrying capacity. In addition, the presence of ponds in wetlands provides geese refuges from predators such as arctic fox

(Hughes et al., 1994; Lecomte et al., 2009), and their disappearance might also change predator-prey interactions.

Effects of gully-induced vegetation changes may finally be visible on variations of greenhouse gas emissions. There is evidence for a strong vegetation control on methane emission from wetlands (Olefeldt et al., 2013; Lara et al., 2015; McEwing et al., 2015; Tveit et al., 2015). In wet polygonal tundra of Northern Siberia, Kutzbach et al. (2004) found for instance that dense *Carex aquatilis* stands emitted more methane than sites with low *Carex* densities. Overall, wetland and lake expansion are thought to increase methane emission but also carbon storage (Myers-Smith, 2005; Nauta et al., 2015; Treat et al., 2015; Bouchard et al., 2015). We can therefore expect that the reverse transition from wet to mesic environments observed within our low-centered polygon landscape would lead to reduced methane emission and increased CO₂ emission through enhanced decomposition. However, no general pattern on ecosystem responses to decreased water table position and subsequent gas emissions has emerged to date (see Grosse et al., 2011 for review). It will thus be crucial to determine in the near future the specific evolution of *Salix* and *Luzula spp* primary production in mesic environments in order to accurately predict the effects of wetland retreat on methane and soil organic carbon cycles.

5 Conclusions

This study illustrates that changes in the hydrological and thermal regimes following the inception of thermo-erosion gullies boost landscape transformation from wet to mesic habitats, providing evidence that permafrost disturbance is a critical component of ecosystem modification at high latitudes. Ecological studies should consequently start using an approach that integrates disturbed permafrost monitoring if one wants to more efficiently document climate change effects on arctic terrestrial ecosystems. In addition, our latest field observations showed that hydrology and thaw regimes of breached polygons have yet to reach equilibrium with new conditions. Similarly, vegetation remains in transition given that, ten years after disturbance, the cover of dominant shrubs and mesic bryophytes in disturbed polygons is still lower than in adjacent mesic environments. It is currently not possible to predict how long these species would take to out compete declining species and cryptogamic crust and reach a new mesic environment equilibrium. This current state stresses the importance of long-term permafrost-vegetation dynamic monitoring. In addition, while surface warming is expected to alter the ground's thermal regime (Couture and Pollard, 2007), more work should be devoted to the feedback effects of plant communities and vegetation succession on thermal and mechanical stabilization dynamics of disturbed permafrost terrains. This is especially needed since plant community differences between disturbed and intact sites can last several centuries (Cray and Pollard, 2015).

436 **Table A1.** Sample sizes and means (\pm SE) of soil moisture, thaw-front depth and graminoid above-ground biomass in the four habitats studied in
437 the Qarlikturvik valley of Bylot Island, Nunavut. Letters refer to post hoc contrasts that were carried out using the LSMEANS statement and
438 Bonferroni adjustment (procedure MIXED, SAS 9.4). For a given variable, different letters indicate significant differences among habitats.
439

| Habitats | Soil moisture 5 Jul 2010 (%) | | Soil moisture 30 Jul 2010 (%) | | Thaw-front depth Jul 2009 (cm) | | Thaw-front depth Jul 2010 (cm) | | Above-ground biomass (g/m ²) | |
|-----------------------------|---------------------------------|-----------------|----------------------------------|-----------------|-----------------------------------|-----------------|-----------------------------------|-----------------|---|-----------------|
| | <i>n</i> | Mean \pm SE | <i>n</i> | Mean \pm SE | <i>n</i> | Mean \pm SE | <i>n</i> | Mean \pm SE | <i>n</i> | Mean \pm SE |
| Wet polygons | 53 | 36.3 \pm 1.7a | 54 | 48.0 \pm 1.3a | 52 | 36.0 \pm 0.9a | 62 | 34.6 \pm 0.8a | 5 | 29.2 \pm 5.2a |
| < 5-year disturbed polygons | 39 | 21.9 \pm 1.2b | 39 | 31.0 \pm 1.5b | 39 | 25.5 \pm 1.5b | 44 | 25.0 \pm 1.1b | 5 | 22.3 \pm 1.4a |
| > 5-year disturbed polygons | 33 | 22.0 \pm 1.1b | 33 | 31.2 \pm 1.2b | 32 | 25.3 \pm 1.1b | 43 | 24.9 \pm 0.8b | 5 | 9.3 \pm 4.0b |
| Mesic environments | 46 | 19.6 \pm 1.3b | 50 | 28.4 \pm 1.4b | 44 | 27.6 \pm 1.1b | 48 | 26.5 \pm 0.7b | 5 | 5.9 \pm 2.6b |

440

441 **Table A2.** List of the vascular plant species inventoried in the Qarlikturvik valley, Bylot Island, Nunavut during the 2009 and 2010 field seasons.
 442 Species names are from the integrated taxonomic information system (ITIS 2011).
 443

| Family | Species |
|------------------|--|
| Asteraceae | <i>Erigeron uniflorus</i> var. <i>eriocephalus</i> (Vahl) Boivin, <i>Taraxacum hyparcticum</i> Dahlst., <i>Tephroseris palustris</i> (L.) Reichenbach |
| Brassicaceae | <i>Cardamine bellidifolia</i> L., <i>Cardamine nymanii</i> Gand., <i>Draba cinerea</i> Adams, <i>Draba corymbosa</i> R. Br., <i>Draba glabella</i> Pursh, <i>Draba lactea</i> Adams, <i>Draba nivalis</i> Lilj., <i>Draba</i> sp. ^a , <i>Eutrema edwardsii</i> R. Br. |
| Caryophyllaceae | <i>Cerastium alpinum</i> L. s. lat., <i>Minuartia rubella</i> (Wahlenb.) Hiern., <i>Silene involucrata</i> (Cham. & Schlecht.) Bocquet subsp. <i>involucrata</i> , <i>Silene uralensis</i> (Rupr.) Bocquet subsp. <i>uralensis</i> , <i>Stellaria longipes</i> Goldie subsp. <i>longipes</i> |
| Cyperaceae | <i>Carex aquatilis</i> var. <i>minor</i> Boott, <i>Carex marina</i> Dewey, <i>Eriophorum angustifolium</i> Honck subsp. <i>angustifolium</i> , <i>Eriophorum scheuchzeri</i> Hoppe s. lat./ <i>russeolum</i> subsp. <i>leiocarpum</i> Novos. |
| Equisetaceae | <i>Equisetum arvense</i> L. s. lat. |
| Ericaceae | <i>Cassiope tetragona</i> var. <i>tetragona</i> (L.) Don |
| Fabaceae | <i>Astragalus alpinus</i> var. <i>alpinus</i> L. <i>Oxytropis maydelliana</i> ssp. <i>Melanocephala</i> (Hook.) Porsild |
| Juncaceae | <i>Juncus biglumis</i> L., <i>Luzula arctica</i> Blytt, <i>Luzula confusa</i> Lindebl. |
| Onagraceae | <i>Epilobium arcticum</i> Sam. |
| Papaveraceae | <i>Papaver</i> sp (including <i>P. dahlianum</i> , <i>P. cornwalisense</i> , <i>P. lapponicum</i> ssp. <i>occidentale</i>) |
| Poaceae | <i>Alopecurus borealis</i> Trin., <i>Anthoxanthum arcticum</i> Veldkamp, <i>Arctagrostis latifolia</i> subsp. <i>latifolia</i> (R. Br.) Griseb., <i>Deschampsia brevifolia</i> R. Br., <i>Dupontia fisheri</i> R. Br., <i>Festuca brachyphylla</i> ssp. <i>Brachyphylla</i> Schult. & Schult. f., <i>Phippsia algida</i> (Sol.) R. Br., <i>Pleuropogon sabinei</i> R. Br., <i>Poa arctica</i> R. Br. s. lat. |
| Polygonaceae | <i>Bistorta vivipara</i> (L.) Delarbre, <i>Oxyria digyna</i> (L.) Hill |
| Pyrolaceae | <i>Pyrola grandiflora</i> Radius |
| Ranunculaceae | <i>Ranunculus hyperboreus</i> Rottb., <i>Ranunculus nivalis</i> L. |
| Rosaceae | <i>Dryas integrifolia</i> Vahl subsp. <i>integrifolia</i> , <i>Potentilla hyparctica</i> Malte subsp. <i>hyparctica</i> |
| Salicaceae | <i>Salix arctica</i> Pall., <i>Salix herbacea</i> L., <i>Salix reticulata</i> L., <i>Salix richardsonii</i> Hook. |
| Saxifragaceae | <i>Chrysosplenium tetrandrum</i> (Lund) Th. Fries, <i>Micranthes foliolosa</i> (R. Br.) Gornall, <i>Micranthes hieraciifolia</i> (Waldst. & Kit.) Haworth, <i>Micranthes nivalis</i> (L.) Small, <i>Saxifraga cernua</i> L., <i>Saxifraga cespitosa</i> L., <i>Saxifraga hirculus</i> L., <i>Saxifraga oppositifolia</i> L. |
| Scrophulariaceae | <i>Pedicularis langsдорffii</i> subsp. <i>arctica</i> (R. Br.) Pennell, <i>Pedicularis sudetica</i> subsp. <i>albolabiata</i> Hult. |

444
 445 ^a *Draba* sp. was not counted as a distinct species because it is probably one of the five species already mentioned.

Acknowledgements

The authors are grateful to the Inuit community of Pond Inlet and to Parks Canada-Sirmilik National Park, Centre d'études nordiques (CEN) and Dr Gilles Gauthier (Université Laval) for the access to the field camp during summers 2009, 2010 and 2012. We also thank Alexandre Guertin-Pasquier, Etienne Godin, Jonathan Lasnier, Stephanie Coulombe and Coralie Henry-Brouillette for their fieldwork support as well as Alexandre Moreau, Stephan Ouellet, Noémie Boulanger-Lapointe for their help with statistical analyses. This project was funded by the International Polar Year program of the Government of Canada, Fonds Québécois de la Recherche sur la Nature et les Technologies (FQRNT), Natural Sciences and Engineering Research Council of Canada (NSERC) Discovery Frontiers grant "Arctic Development and Adaptation to Permafrost in Transition" (ADAPT), Network of Centers of Excellence of Canada ArcticNet, Northern Scientific Training Program (NSTP), NSERC CREATE Training Program in Northern Environmental Sciences (EnviroNorth) and Groupe de Recherche en Biologie Végétale (GRBV) of Université du Québec à Trois-Rivières. Essential logistic support was provided by Polar Continental Shelf Program (Natural Resources Canada).

462 **References**

- 463 Ackerly, D.: Functional strategies of chaparral shrubs in relation to seasonal water deficit and
464 disturbance, *Ecol. Monogr.*, 74, 25-44, 2004.
- 465 Akerman, H. J. and Johansson, M.: Thawing permafrost and thicker active layer in sub-arctic
466 Sweden, *Permafrost Periglac.*, 19, 279-292, 2008.
- 467 Allard, M.: Geomorphological changes and permafrost dynamics: key factors in changing
468 arctic ecosystems. An example from Bylot Island, Nunavut, Canada, *Geosci. Can.*, 23,
469 205–212, 1996.
- 470 Audet, B., Lévesque, E. and Gauthier G.: Seasonal variation in plant nutritive quality for
471 Greater Snow Goose goslings in mesic tundra, *Can. J. Bot.*, 85, 457-462, 2007.
- 472 Avis, C. A., Weaver, A. J. and Meissner, K. J.: Reduction in areal extent of high-latitude
473 wetlands in response to permafrost thaw, *Nat. Geosci.*, 4, 444-448, 2011.
- 474 Beck, I., Ludwig, R., Bernier, M., Lévesque, E. and Boike, J.: Assessing permafrost
475 degradation and land cover changes (1986-2009) using remote sensing data over Umiujaq,
476 sub-Arctic Québec, *Permafrost Periglac.*, DOI: 10.1002/ppp.1839, 2015.
- 477 Beilman, D. W. and Robinson, S. D.: Peatland permafrost thaw and landform type along a
478 climate gradient, in: *Proceedings, Eighth International Conference on Permafrost*, edited by:
479 Phillips, M., Springman, P. M. and Arenson, S. U., A. A. Balkema, Zurich, Switzerland, 1,
480 61-65, 2003.
- 481 Billings, W. D. and Peterson, K. M: Vegetational change and ice-wedge polygons through the
482 thaw-lake cycle in Arctic Alaska, *Arct. Alp. Res.*, 12, 413–432, 1980.
- 483 Blok, D., Heijman, M. M. P. D., Schaepman-Strub, G., Kononov, A. V., Maximov, T. C. and
484 Berendse, F.: Shrub expansion may reduce summer permafrost thaw in Siberia tundra,
485 *Glob. Change Biol.*, 16, 1296-1305, 2010.
- 486 Bouchard, F., Laurion, I., Preskienis, V., Fortier, D., Xu, X. and Whiticar, M. J.: Modern to
487 millennium-old greenhouse gases emitted from freshwater ecosystems of the Eastern
488 Canadian Arctic, *Biogeosciences Discuss.*, 12, 11661-11705, doi:10.5194/bgd-12-11661-
489 2015, 2015.
- 490 Britton, A. J. et al.: Biodiversity gains and losses: evidence for homogenization of Scottish
491 alpine vegetation, *Biol. Conserv.*, 142, 1728-1739, 2009.
- 492 Cadieux, M.-C., Gauthier, G., Gagnon, C. A., Lévesque, E., Bêty, J. and Berteaux, D.:
493 Monitoring the environmental and ecological impacts of climate change on Bylot Island,
494 Sirmilik National Park, Université Laval, 113 pp., 2008.
- 495 Cannone, N., Lewkowicz, A. G. and Guglielmin, M.: Vegetation colonization of permafrost-
496 related landslides, Ellesmere Island, Canadian High Arctic, *J. Geophys. Res.*, 115, G04020,
497 doi:10.1029/2010JG001384, 2010.
- 498 Cray, H. A. and Pollard, W. H.: Vegetation recovery patterns following permafrost
499 disturbance in a Low Arctic setting: case study of Herschel Island, Yukon, Canada, *Arct.*
500 *Antarct. Alp. Res.*, 47, 99-113, 2015.
- 501 Chen, Z. H., Chen, W. J. and Leblanc, S. G. and Henry, G. H. R.: Digital photograph analysis
502 for measuring percent plant cover in the Arctic, *Arctic*, 63, 315–326, 2010.
- 503 Christiansen, H. H., Eltzelmueller, B., Isaksen, K., Juliussen, H., Farbrøt, H., Humlum, O.,
504 Johansson, M., Ingeman-Nielsen, T., Kristensen, L., Hjort, J. et al.: The thermal state of
505 permafrost in the Nordic Area during the International Polar Year 2007-2009, *Permafrost*
506 *Periglac.*, 21, 156-181, 2010.
- 507 Collins, Knutti, R., Arblaster, J., Dufresne, J.-L., Fichet, T., Friedlingstein, P., Gao, X.,
508 Gutowski, W. K., Johns, T., Krinner, G. et al.: Long-term climate change: Projections,
509 commitments and irreversibility, in: *Climate change 2013: the physical basis*, Contribution
510 of working group I to the Fifth Assessment Report of the Intergovernmental Panel on

511 Climate Change, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S.
 512 K., Boschung, J., Nauels, A., Xia, Y., Bex, V. and Midgley, P. M., Cambridge University
 513 Press, Cambridge, United Kingdom, 2013.
 514 Couture, N. J. and Pollard, W. H.: Modelling geomorphic response to climatic change, *Clim.*
 515 *Change*, 85, 407–431, 2007.
 516 Dagg, J. and Lafleur, P.: Vegetation community, foliar nitrogen, and temperature effects on
 517 tundra CO₂ exchange across a soil moisture gradient, *Arct. Antarct. Alp. Res.*, 43, 189-197,
 518 2011.
 519 Daniëls, F. J. A. and de Molenaar, J. G.: Flora and vegetation of Tasiilaq, formerly
 520 Angmagssalik, southeast Greenland: a comparison of data between around 1900 and 2007,
 521 *Ambio*, 40, 650-659, 2011.
 522 Daniëls, F. J. A. et al.: Vegetation change in southeast Greenland? Tasiilaq revisited after 40
 523 years, *J. Appl. Veg. Sci.*, 14, 230-241.
 524 Daubenmire, R.: A canopy-coverage method of vegetation analysis, *Northwest Sci.*, 33,
 525 43–64, 1959.
 526 Doiron, M.: Impacts des changements climatiques sur les relations plantes-herbivores dans
 527 l'Arctique, PhD thesis, Université Laval, 123 pp., 2014.
 528 Doiron, M., Gauthier, G. and Lévesque, E.: Effects of experimental warming on nitrogen
 529 concentration and biomass of forage plants for an arctic herbivore, *J. Ecol.* 102, 508-517,
 530 2014.
 531 Duclos, I.: Milieux mésiques et secs de l'Île Bylot, Nunavut (Canada): caractérisation et
 532 utilisation par la Grande Oie des Neiges. MSc thesis, Université du Québec à Trois-
 533 Rivières, Trois-Rivières, 113 pp., 2002.
 534 Ellis, C. J. and Rochefort L.: Century-scale development of polygon-patterned tundra
 535 wetland, Bylot Island (73° N, 80° W), *Ecology*, 4, 963-978, 2004.
 536 Ellis, C. J. and Rochefort L.: Long-term sensitivity of a High Arctic wetland to Holocene
 537 climate change, 94, 441-454, *J. Ecol.*, 2006.
 538 Ellis, C. J., Rochefort L., Gauthier G. and Pienitz, R.: Paleoecological evidence for transitions
 539 between contrasting landforms in a polygon-patterned High Arctic wetland, *Arct. Antarct.*
 540 *Alp. Res.*, 40, 624–637, 2008.
 541 Epstein, H. E., Myers-Smith, I. and Walker, D. A.: Recent dynamics of arctic and sub-arctic
 542 vegetation, *Environ. Res. Lett.*, 8, 015040, 6 pp., 2013.
 543 Ewertowski, M.: Ice-wedge pseudomorphs and frost-cracking structures in Weichselian
 544 sediments, Central-West Poland, *Permafrost Periglac.*, 20, 316-330, 2009.
 545 Fortier, D. and Allard, M.: Late Holocene syngenetic ice-wedge polygons development, Bylot
 546 Island, Canadian Arctic Archipelago, *Can. J. Earth Sci.*, 41, 997–1012, 2004.
 547 Fortier, D. and Allard, M.: Frost-cracking conditions, Bylot Island, Eastern Canadian Arctic
 548 archipelago, *Permafrost Periglac.*, 16, 145-161, 2005.
 549 Fortier, D., Allard, M. and Pivot, F.: A late-Holocene record of loess deposition in ice-wedge
 550 polygons reflecting wind activity and ground moisture conditions, Bylot Island, eastern
 551 Canadian Arctic, *Holocene*, 16, 635-646, 2006.
 552 Fortier, D., Allard, M. and Shur, Y.: Observation of rapid drainage system development by
 553 thermal erosion of ice wedges on Bylot Island, Canadian Arctic Archipelago, *Permafrost*
 554 *Periglac.*, 18, 229–243, 2007.
 555 Fortier, D., Godin, E., Lévesque, E., Veillette, A. and Lamarque, L. J.: Thermal erosion of
 556 ice-wedge polygon terrains changes fluxes of energy and matter of permafrost ecosystems,
 557 American Geophysical Union fall meeting, San Francisco, United States of America, 2015.
 558 Gangodagamage, C., Rowland, J. C., Hubbard, S. S., Brumby, S. P., Liljedahl, A. K.,
 559 Wainwright, H., Wilson, C. J., Altmann, G. L., Dafflon, B., Peterson, J. et al.:

- Extrapolating active layer thickness measurements across Arctic polygonal terrain using LiDAR and NDVI data sets, *Water Resour. Res.*, 50, doi:10.1002/2013WR014283, 2014.
- Gauthier, G., Bêty, J., Cadieux, M.-C., Legagneux, P., Doiron, M., Chevallier, C., Lai, S., Tarroux, A. and Berteaux D.: Long-term monitoring at multiple trophic levels suggests heterogeneity in responses to climate change in the Canadian Arctic tundra, *Phil. Trans. R Soc. B*, 368: 20120482, 12 pp., 2013.
- Gauthier, G., Cadieux, M.-C., Lefebvre, J., Bêty, J. and Berteaux D.: Population study of Greater Snow Geese and its nesting habitat on Bylot Island, Nunavut, in 2009: A progress report, 31 pp., 2009.
- Gauthier, G., Cadieux, M.-C., Lefebvre, J., Bêty, J. and Berteaux D. Population study of Greater Snow Geese and its nesting habitat on Bylot Island, Nunavut, in 2010: A progress report, 32 pp., 2010.
- Gauthier, G., Cadieux, M.-C., Lefebvre, J., Bêty, J. and Berteaux D. Population study of Greater Snow Geese and its nesting habitat on Bylot Island, Nunavut, in 2012: A progress report, 36 pp., 2012.
- Gauthier, G., Berteaux D., Bêty, J., Tarroux, A., Therrien, J.-F., McKinnon, L., Legagneux, P. and Cadieux, M.-C.: The tundra food web of Bylot Island in a changing climate and the role of exchanges between ecosystems, *Ecoscience*, 18, 223-235, 2011.
- Gauthier, G., Hughes, R. J., Reed, R., Beaulieu, J. and Rochefort, L.: Effect of grazing by greater snow geese on the production of graminoids at an arctic site (Bylot Island, NWT, Canada). *J. Ecol.*, 83, 653–664, 1995.
- Gauthier, G., Rocheford, L. and Reed, A.: The exploitation of wetland ecosystems by herbivores on Bylot Island, *Geosci. Can.*, 23, 253–259, 1996.
- Godin, E. and Fortier, D.: Geomorphology of thermo-erosion gullies – case study from Bylot Island, Nunavut, Canada, in: *Proceedings, 6th Canadian Permafrost Conference and 63rd Canadian Geotechnical Conference*, Calgary, 2010.
- Godin, E. and Fortier, D.: Fine scale spatio-temporal monitoring of multiple thermo-erosion gullies development on Bylot Island, eastern canadian archipelago, in: *Proceedings, Tenth International Conference on Permafrost (TICOP)*, Salekhard, Russia, 7 pp., 2012a.
- Godin, E. and Fortier, D.: Geomorphology of a thermo-erosion gully, Bylot Island, Nunavut, Canada, *Can. J. Earth Sci.*, 49, 979-986, 2012b.
- Godin, E., Fortier, D. and Coulombe, S.: Effects of thermo-erosion gullying on hydrologic flow networks, discharge and soil loss, *Environ. Res. Lett.*, 9, 10 pp., doi:10.1088/1748-9326/9/10/105010, 2014.
- Godin, E. Fortier, D. and Lévesque, E.: Nonlinear thermal and moisture dynamics of high Arctic wetland polygons following permafrost disturbance, *Biogeosciences Discuss.*, 12, 11797-11831, doi:10.5194/bgd-12-11797-2015, 2015.
- Grosse, G., Harden, J., Turetsky, M., McGuire A. D., Camill, P., Tarnocai, C., Frolking, S., Schuur, E. A. G., Jorgenson, T., Marchenko, S. et al.: Vulnerability of high-latitude soil organic carbon in North America to disturbance, *J. Geophys. Res.*, 116, G00K06, doi:10.1029/2010JG001507, 2011.
- Gruyer, N. G., Gauthier, G. and Berteaux, D.: Cyclic dynamics of sympatric lemming populations on Bylot Island, Nunavut, Canada, *Can. J. Zool.*, 86, 910-917, 2008.
- Günther, F., Overduin, P. P., Sandakov, A. V., Grosse, G. and Grigoriev, M. N.: Short- and long-term thermo-erosion of ice-rich permafrost coasts in the Laptev Sea region, *Biogeosciences*, 10, 4297-4318, 2013.
- Harms, T. K., Abbott, B. W. and Jones, J. B.: Thermo-erosion gullies increase nitrogen available for hydrologic export, *Biogeochemistry*, 117, 299-311, 2014.
- Hinzman, L. D., Bettes, N. D., Bolton, W. R., Chapin III, F. S., Dyurgerov, M. B., Fastie, C. L., Griffith, B., Hollister, R. D., Hope, A., Huntington, H. P. et al.: Evidence and

- implications of recent climate change in northern Alaska and other arctic regions, *Clim. Change*, 72, 251–298, 2005.
- Hollister, R. D., Webber, P. J. and Tweedie, C. E.: The response of Alaskan arctic tundra to experimental warming: differences between short- and long-term responses, *Glob. Change Biol.*, 11, 525–536, 2005.
- Hudson, J. M. G. and Henry, G. H. R.: High Arctic plant community resists 15 years of experimental warming, *J. Ecol.*, 98, 1035–1041, 2010.
- Hughes, R. J., Reed, A. and Gauthier, G. Space and habitat use by greater snow goose broods on Bylot Island, Northwest Territories. *J. Wildlife Manage.*, 58, 536–545. 1994.
- IPY, International Polar Year: Climate Change Impact on Canadian Arctic Tundra, Protocols for taking vertical photographs and field measurements of plant height for biomass and LAI estimation in Arctic tundra (for vegetation height < 0.5 m), <http://ipytundra.ca/protocols>, 2012.
- Jonsdottir, I. S., Magnusson, B., Gudmunsson, J., Elmarsdottir, A. and Hjartarson, H.: Variable sensitivity of plant communities in Iceland to experimental warming, *Glob. Change Biol.*, 11, 553–563, 2005.
- Jorgenson, J. C., Reynolds, M. K., Reynolds, J. H. and Benson, A-M.: Twenty-five year record of changes in plant cover on tundra of northeastern Alaska, *Arct. Antarct. Alp. Res.*, 47, 785–806, 2015a.
- Jorgenson, M. T., Harden, S., Kanevskiy, M., O'Donnell, J., Wickland, K., Ewing, S., Manies, K., Zhuang, Q., Shur, Y., Striegl, R. and Koch, J.: Reorganization of vegetation, hydrology and soil carbon after permafrost degradation across heterogeneous boreal landscapes, *Environ. Res. Lett.*, 8, 13 pp., doi:10.1088/1748-9326/9/10/105010, 2013.
- Jorgenson, M. T., Kanevskiy, M., Shur, Y., Moskalenko, N., Brown, D. R. N., Wickland, K., Striegl, R. and Koch, J.: Role of ground ice dynamics and ecological feedbacks in recent ice wedge degradation and stabilization, *J. Geophys. Res. Earth Surf.*, 120, 2280–2297, doi:10.1002/2015JF003602, 2015b.
- Jorgenson, M. T., Romanosky, V., Harden, S., Shur, Y., O'Donnell, J., Schuur E. A. G., Kanevskiy, M. and Marchenko, S.: Resilience and vulnerability of permafrost to climate change, *Can. J. For. Res.*, 40, 1219–1236, 2010.
- Jorgenson, M. T., Shur, Y. T. and Pullman, E. R.: Abrupt increase in permafrost degradation in Arctic Alaska, *Geophys. Res. Lett.*, 33, L02503, doi:10.1029/2005GRL024960, 2006.
- Kokelj, S. V., Lantz, T. C., Wolfe, S. A., Kanigan, J. C., Morse, P. D., Coutts, R., Molina-Giraldo, N. and Burn, C. R.: Distribution and activity of ice wedges across the forest-tundra transition, western Arctic Canada, *J. Geophys. Res. Earth Surf.*, 119, doi:10.1002/2014JF003085, 2014.
- Kutzbach, H. J., Wagner, D. and Pfeiffer, E. M.: Effect of microrelief and vegetation on methane emission wet polygonal tundra, Lena Delta, Northern Siberia. *Biogeochemistry*, 69, 341–362, 2004.
- Lara, M. J., McGuire, A. D., Euskirchen, E. S., Tweedie, C. E., Hinkel, K. M., Skurikhin, A. N., Romanovsky, V. E., Grosse, G., Bolton, W. R. and Genet, H.: Polygonal tundra geomorphological change in response to warming alters future CO₂ and CH₄ flux on the Barrow Peninsula, *Glob. Change Biol.*, 21, 1634–1651, 2015.
- Lecomte, N., Gauthier, G. and Giroux, J.-F.: A link between water availability and nesting success mediated by predator-prey interactions in the Arctic, *Ecology*, 90, 465–475, 2009.
- Legagneux, P., Gauthier, G., Berteaux, D., Bêty, J., Cadieux, M.-C., Bilodeau, F., Bolduc, E., McKinnon, L., Tarroux, A., Therrien, J.-F., Morissette, L. and Krebs, C. J.: Disentangling trophic relationships in a High Arctic tundra ecosystem through food web modeling, *Ecology*, 93, 1707–1716, 2012.
- Levy, J. S., Head, J. W. and Marchant, D. R.: The role of thermal contraction crack polygons

- in cold-desert fluvial systems, *Antarct. Sci.*, 20, 565-579, 2008.
- Liljedahl, A. K., Hinzman, L. D. and Schulla, J.: Ice-wedge polygon type controls low-gradient watershed-scale hydrology, in: *Proceedings, Tenth International Conference on Permafrost (TICOP)*, Salekhard, Russia, 6 pp., 2012.
- Littell, R. C., Milliken, G. A., Stroup, W. W., Wolfinger, R. D. and Schabenberger, O.: *SAS[®] for Mixed Models*, 2nd edn, SAS Institute Inc, Cary, NC, 2006.
- Manseau, M. and Gauthier, G.: Interactions between greater snow geese and their rearing habitat, *Ecology*, 74, 2045–2055, 1993.
- Massé, H., Rochefort, L. and Gauthier, G.: Carrying capacity of wetland habitats used by breeding greater snow geese, *J. Wildlife Manage.*, 65, 271–281, 2001.
- McEwing, K. R., Fisher, J. P. and Zona, D.: Environmental and vegetation controls on the spatial variability of CH₄ emission from wet-sedges and tussock tundra ecosystems in the Arctic, *Plant Soil*, 338, 37-52, 2015.
- Minke, M., Donner, D., Karpov, N. S., de Klerk, P. and Joosten, H.: Distribution, diversity, development and dynamics of polygon mires: examples from Northeast Yakutia (Siberia), *Peatlands Int.*, 1, 36–40, 2007.
- Minke, M., Donner, D., Karpov, N. S., de Klerk, P. and Joosten, H.: Patterns in vegetation composition, surface height and thaw depth in polygon mires in the Yakutian Arctic (NE Siberia): a microtopographical characterisation of the active layer, *Permafrost Periglac.*, 20, 357–368, 2009.
- Muc, M., Freedman, B. and Svoboda, J.: Vascular plant communities of a polar oasis at Alexandra Fiord (79°N), Ellesmere Island, *Can. J. Bot.*, 67, 1126-1136, 1989.
- Myers-Smith, I. H.: Carbon exchange and permafrost collapse: implications for a changing climate, MSc thesis, University of Alaska Fairbanks, Fairbanks, 66 pp., 2005.
- Myers-Smith, I. H., Warden, J. W., Wilkening, M., Fuller, C. C., McGuire, A. D. and Chapin III, F. S.: Wetland succession in a permafrost collapse: interactions between fire and thermokarst, *Biogeosciences*, 5, 1273-1286, 2008.
- Naito, A. T. and Cairns, D. M.: Patterns of shrub expansion in Alaskan arctic river corridors suggest phase transition, *Ecol. Evol.*, 5, 87-101, 2015.
- Natali, S. M., Schuur, E. A. G., Mauritz, M., Schade, J. D., Celis, G., Crummer, K. G., Johnston, C., Krapek, J., Pegoraro, E., Salmon, V. G. and Webb, E. E.: Permafrost thaw and soil moisture driving CO₂ and CH₄ release from upland tundra, *J. Geophys. Res. Biogeosci.*, 120, doi:10.1002/2014JG002872, 2015.
- Nauta, A. L., Heijmans, M. M. P. D., Blok, D., Limpens, J., Elberling, B., Gallagher, A., Li, B., Petrov, R. E., Maximov, T. C., van Huissteden, J. and Berendse, F.: Permafrost collapse after shrub removal shifts tundra ecosystem to a methane source, *Nature Clim. Change*, 5, 67-70, 2015.
- Nelson, F. E., Shiklomanov, N. I. and Mueller, G. R.: Variability of active-layer thickness at multiple spatial scales, North-central Alaska, U.S.A, *Arct. Antarct. Alp. Res.*, 31, 179-186, 1999.
- Nelson, F. E., Shiklomanov, N. I., Hinkel, K. M. and Brown J.: Decadal results from the Circumpolar Active Layer Monitoring (CALM) program, *Proceedings, Ninth International Conference on Permafrost*, edited by: Hinkel, K. M. and Kane, D. L., 1273-1280, University of Alaska Press, Fairbanks, Alaska, 2008.
- New, M., Liverman, D., Schroder, H. and Anderson, K.: Four degrees and beyond: the potential for a global temperature increase of four degrees and its implications, *Phil. Trans. R. Soc. A*, 369, 6-19, 2011.
- NOAA (National Oceanic and Atmospheric Administration), Arctic Research Program: Arctic Report Card, edited by: Jeffries, M. O., Richter-Menge, J. and Overland, J. E., <http://www.arctic.noaa.gov/reportcard>, 2014.

- Oksanen, J., Blanchet, F. G., Kindt, R., Legendre, P., O'Hara, B., Simpson, G. L., Solymos, P., Stevens, M. H. M. and Wagner, H.: *vegan*: Community ecology package, R package version 2.2-1, 2015.
- Olefeldt, D., Turetsky, M. R., Crill, P. M. and McGuire, A. D.: Environmental and physical controls on northern terrestrial methane emissions across permafrost zones, *Glob. Change Biol.*, 19, 589-603, 2013.
- Payette, S., Delwaide, A., Caccianiga, M. and Beauchemin, M.: Accelerated thawing of subarctic peatland permafrost over the last 50 years, *Geophys. Res. Lett.*, 31, L18208, 2004.
- Pauli, K. et al.: High Arctic vegetation after 70 years: a repeated analysis from Svalbard, *Pol. Biol.*, 33, 635-639, 2012.
- Perreault, N.: Impact de la formation de ravins de thermo-erosion sur les milieux humides, Ile Bylot, Nunavut, Canada, MSc thesis, Université du Québec à Trois-Rivières, 104 pp., 2012.
- Poesen, J., Nachtergaele, J., Verstraeten, G. and Valentin, C.: Gully erosion and environmental change: importance and research needs, *Catena*, 50, 91-133, 2003.
- Pouliot, R., Rochefort, L. and Gauthier, G.: Moss carpets constrain the fertilizing effects of herbivores on graminoid plants in arctic polygon fens, *Botany*, 87, 1209-1222, 2009.
- R Development Core Team: R: a language and environment for statistical computing. R Foundation for Statistical Computing, Vienna, Austria, ISBN 3-900051-07-0, URL <http://www.R-project.org>, 2015.
- Ross, L. C. et al.: Biotic homogenization of upland vegetation: patterns and drivers at multiple spatial scales over five decades, *J. Veg. Sci.*, 23, 755-770, 2012.
- Rowland, J. C., Jones, C. E., Altmann, G., Bryan, B., Crosby, B. T., Geernaert, G. L., Hinzman, L. D., Kane, D. L., Lawrence, D. M., Mancino, A., Marsh, P., McNamara, J. P., Romanosky, V. E., Toniolo, H., Travis, B. J., Trochim, E. and Wilson C. J.: Arctic landscapes in transition: responses to thawing permafrost, *Eos Trans. Am. Geophys. Union*, 91, 229-236, 2010.
- Romanovsky, V. E., Smith, S. L., and Christiansen, H. H.: Permafrost thermal state in the polar Northern Hemisphere during the international polar year 2007–2009: a synthesis, *Permafrost Periglac.*, 21, 106-116, doi:10.1002/ppp.689, 2010.
- Sandvik, S. M. and Odland, A.: Changes in alpine snowbed-wetland vegetation over three decades in norther Norway, *Nordic J. Bot.*, 32, 377-384, 2014.
- Sarrazin, D. and Allard, M.: The thermo-mechanical behavior of frost-cracks over ice wedges: new data from extensometer measurements, 68th Canadian Geotechnical Conference and 7th Canadian Permafrost Conference, Quebec City, Canada, 2015.
- Sedinger, J. S., White, R. G., Mann, F. E., Burris, F. A. and Kedrowski, R. A.: Apparent metabolizability of alfalfa components by yearling Pacific Black Brant, *J. Wildl. Manage.*, 53, 726-734, 1989.
- Seppälä, M.: Piping causing thermokarst in permafrost, Ungava Peninsula, Quebec, Canada, *Geomorphology*, 20, 313–319, 1997.
- Sheard, J. W. and Geale, D. W.: Vegetation studies at Polar Bear Bass, Bathurst Island, N.W.T., I. Classification of plant communities, *Can. J. Bot.*, 61, 1618-1636, 1983.
- Shiklomanov, N. I., Streletskiy, D. A., Nelson, F. E., Hollister, R. D., Romanovsky, V. E., Tweedie, C. E., Bockheim, J. G. and Brown, J.: Decadal variations of active-layer thickness in moisture-controlled landscapes, Barrow, Alaska, *J. Geophys. Res.*, 115, G00I04, 2010.
- Shiklomanov, N. I., Streletskiy, D. A., Little, J. D. and Nelson, F. E.: Isotropic thaw subsidence in undisturbed permafrost landscapes, *Geophys. Res. Lett.*, 40, doi:10.1002/2013GL058295, 2013.

- Sjöberg, Y.: Linking water and permafrost dynamics, PhD thesis, Stockholm University, 27 pp., 2015.
- Smith, S.: Trends in permafrost conditions and ecology in northern Canada: Ecosystem Status and Trends 2010, Technical Thematic Report No. 9, Canadian Councils of Resource Ministers, Ottawa, ON, iii + 22 p, 2011.
- Smith, S. L., Burgess, M. M., Riseborough, D. and Nixon, F. M.: Recent trends from Canadian Permafrost Thermal Monitoring network sites, *Permafrost Periglac.*, 16, 19-30, 2005.
- Smith, S. L., Wolfe, S. A., Riseborough, D. W. and Nixon, F. M.: Active-layer characteristics and summer climatic indices, Mackenzie Valley, Northwest Territories, Canada, *Permafrost Periglac.*, 20, 201-220, 2009.
- Tarnocai, C., Nixon, F. M. and Kutny, L.: Circumpolar-Active-Layer-Monitoring (CALM) sites in the Mackenzie Valley, Northwestern Canada, *Permafrost Periglac.*, 15, 141-153, 2004.
- Treat, C. C., Natali, S. M., Ernakovitch, J., Iversen, C. M., Lupascu, M., McGuire, A. D., Norby, R. J., Chowdhury, T. R., Richter, A., Santruckova, H. et al.: A pan-Arctic synthesis of CH₄ and CO₂ production from anoxic soil incubations, *Glob. Change. Biol.*, doi:10.1111/gcb.12875, 2015
- Turetsky, M., McGuire, A. and Olefeldt, D.: Assessing the contributions of thermokarst and thermal erosion in permafrost feedbacks to climate, American Geophysical Union fall meeting, San Francisco, United States of America, 2015.
- Tveit, A. T., Urich, T., Frenzel, P. and Svenning, M. M.: Metabolic and trophic interactions modulate methane production by Arctic peat microbiota in response to warming, *Proc. Nat. Ac. Sci.*, doi:10.1073/pnas.1420797112, 2015.
- Ukrainitseva, N. G., Streletskaya, I. D., Ermokhina, K. A. and Yermakov, S. Yu.: Geochemical properties of plant-soil-permafrost systems on landslide slopes, Yamal, Russia, in: *Proceedings, Tenth International Conference on Permafrost*, Zurich, Switzerland, 1149-1154, 2003.
- USGCRP: Global Climate Change Impacts in the United States, edited by: Karl, T. R., Melillo, J. M. and Peterson, T. C., United States Global Change Research Program, Cambridge University Press, New York, NY, 2009.
- ter Braak, C. J. F.: Canonical correspondence analysis: a new eigenvector technique for multivariate direct gradient analysis, *Ecology*, 67, 1167–1179, 1986.
- Veillette, A., Fortier, D. and Godin, E.: Contrasting patterns of thermo-erosion gullies formed in syngenetic ice wedge polygonal terrains on Bylot Island, eastern Canadian Arctic: case studies from three different sedimentary environments, 68th Canadian Geotechnical Conference and 7th Canadian Permafrost Conference, Quebec City, Canada, 2015.
- Vonk, J. E., Mann, P. J., Dowdy, K. L., Davydova, A., Davydov, S. P., Zimov, N., Spencer, R. G. M., Bulygina, E. B., Eglinton, T. I. and Holmes, R. M.: Dissolved organic carbon loss from Yedonna permafrost amplified by ice wedge thaw, *Environ. Res. Lett.*, 8, 9 pp., doi:10.1088/1748-9326/8/3/035023, 2013.
- Walker, D. A., Leibman, M. O., Epstein, H. E., Forbes, B. C., Bhatt, U. S., Raynolds, M. K., Comiso, J. C., Gubarkov, A. A., Khomutov, A. V., Jia, G. J., Kaarlejärvi, E., Kaplan, J. O., Kumpula, T., Kuss, P., Matyshak, G., Moskalenko, N. G., Orekhov, P., Romanovsky, V. E., Ukrainitseva, N. G. and Yu, Q.: Spatial and temporal patterns of greenness on the Yamal Peninsula, Russia: interactions of ecological and social factors affecting the Arctic normalized difference vegetation index, *Environ. Res. Lett.*, 4, 16 pp., doi:10.1088/1748-9326/4/4/045004, 2009.
- Woo, M.-K.: *Permafrost hydrology*, Springer, New York, NY, 564 pp., 2012.

809 Woo, M.-K. and Young, K. L.: High Arctic wetlands: their occurrence, hydrological
810 characteristics and sustainability, *Journal of Hydrology*, 320, 432–450, 2006.

811 Woo, M.-K., Kane, D. L., Carey, S. K. and Yang, D.: Progress in permafrost hydrology in the
812 new millennium, *Permafrost Periglac.*, 19, 237–254, 2008.

813 Woo, M.-K., Mollinga, M. and Smith, S. L.: Climate warming and active layer thaw in the
814 boreal and tundra environments of the Mackenzie Valley, *Can. J. Earth Sci.*, 44, 733–743,
815 2007.

816 Wright, N., Hayashi, M. and Quinton, W. L.: Spatial and temporal variations in active layer
817 thickness and their implication on runoff generation in peat-covered permafrost terrain,
818 *Water Resources Res.*, 45, W05414, 2009.

819 Zamin, T. Z., Brett-Harte, M. S. and Grogan, P.: Evergreen shrubs dominate responses to
820 experimental summer warming and fertilization in Canadian mesic low arctic tundra, *J.*
821 *Ecol.*, doi: 10.1111/1365-2745.12237, 2014.

822 Zoltai, S. C., McCormick K. J. and Scotter, G. W.: A natural resource survey of Bylot Island
823 and adjacent Baffin Island, Northwest Territories, Parks Canada, Ottawa, ON, 176 pp.,
824 1983.

825 Zuur, A. F., Ieno, E. N. and Smith, G. M.: *Analysing ecological data*, edited by: Gail, M.,
826 Krickeberg, K., Samet, J., Tsiatis, A. and Wong, W., Springer, New York, 672 pp., 2007.

Table 1. Repartition of the studied polygons per habitat and per gully.

| Gullies | Habitats | | | | Total |
|---------|--------------|-----------------------------|-----------------------------|--------------------|------------|
| | Wet polygons | < 5-year disturbed polygons | > 5-year disturbed polygons | Mesic environments | |
| R08p | 30 | 22 | 22 | 21 | 95 |
| R06 | 23 | 19 | 12 | 27 | 81 |
| RN08 | 9 | 3 | 9 | 0 | 21 |
| Total | 62 | 44 | 43 | 48 | 197 |

Table 2. Species richness, family total cover and species mean cover of vascular taxa as well as mean cover of non-vascular taxa in each of the four habitats sampled at Bylot Island, Nunavut. Mean species richness is given for sampled areas of 49 dm² (70 x 70 cm quadrats). Numbers in brackets denote the number of species inventoried in each family. << = cover < 0.01 %; < = cover < 0.1 %.

| | Wet polygons (n = 62) | Disturbed polygons | | Mesic environments (n = 48) |
|---------------------------------|-----------------------------|-----------------------|-----------------------|-----------------------------------|
| | | < 5 years (n = 44) | > 5 years (n = 43) | |
| Vascular taxa | | | | |
| Total plant species richness | 36 | 37 | 54 | 47 |
| Mean plant species richness | 8 | 11 | 17 | 15 |
| Plant species cover (%) | | | | |
| Asteraceae (3) | << | < | < | << |
| Brassicaceae (8) | < | < | 0.3 | < |
| Caryophyllaceae (5) | << | 0.1 | 1.0 | 0.2 |
| Cyperaceae (4) | 17.4 | 18.5 | 7.0 | 1.9 |
| <i>Carex aquatilis</i> | 14.7 | 17.8 | 6.4 | 1.9 |
| <i>Carex marina</i> | 0 | << | 0 | 0 |
| <i>Eriophorum angustifolium</i> | 0.5 | 0.4 | 0.5 | 0 |
| <i>Eriophorum scheuchzeri</i> | 2.1 | 0.4 | << | < |
| Equisetaceae (1) | < | < | < | < |
| Ericaceae (1) | 0 | 0 | << | < |
| Fabaceae (2) | << | << | 0.2 | 0.1 |
| Juncaceae (3) | < | 0.2 | 3.4 | 3.0 |
| <i>Luzula arctica</i> | << | < | 0.4 | 1.0 |
| <i>Luzula confusa</i> | < | 0.1 | 3.0 | 2.0 |
| Onagraceae (1) | << | 0 | 0 | 0 |
| Papaveraceae (1) | 0 | 0 | << | << |
| Poaceae (9) | 8.7 | 5.0 | 2.7 | 3.4 |
| <i>Anthoxanthum arcticum</i> | 0.5 | < | << | << |
| <i>Arctagrostis latifolia</i> | 0.6 | < | 0.3 | 3.0 |
| <i>Dupontia fisheri</i> | 7.4 | 4.8 | 1.8 | 0.1 |
| <i>Festuca brachyphylla</i> | << | < | 0.4 | < |
| Polygonaceae (2) | << | << | < | < |
| Pyrolaceae (1) | 0 | 0 | << | 0 |
| Ranunculaceae (2) | 0 | 0 | << | << |
| Rosaceae (2) | 0 | 0 | < | < |
| Salicaceae (4) | 1.2 | 0.9 | 3.4 | 12.6 |
| <i>Salix arctica</i> | 1.2 | 0.9 | 3.1 | 10.8 |
| <i>Salix herbacea</i> | 0 | 0 | < | 0.1 |
| <i>Salix reticulata</i> | << | << | 0.2 | 1.6 |
| <i>Salix richardsonii</i> | << | << | << | << |
| Saxifragaceae (8) | < | 0.2 | 1.5 | < |
| Scrophulariaceae (2) | 0.7 | 0.3 | < | < |
| Non-vascular taxa | | | | |
| Lichens | < | < | 1.9 | 6.6 |
| Live mosses | 53.1 | 26.6 | 15.9 | 44.6 |
| <i>Drepanocladus</i> spp. | 39.9 | 21.4 | 5.1 | 0 |
| <i>Aulacomnium</i> spp. | 7.5 | 3.2 | 1.4 | 39.3 |
| Dried mosses | 2.9 | 29.6 | 18.2 | 0.9 |
| Fungi | < | 0.2 | 0.1 | < |
| <i>Nostoc</i> spp. | < | < | << | << |
| Cryptogamic crust | << | < | 14.6 | 5.1 |

Table 3. Information of the Canonical Correspondence Analysis. CCA-1: first canonical axis; CCA-2: second canonical axis.

| | CCA-1 | CCA-2 |
|---|--------|--------|
| Eigenvalues | 0.2208 | 0.1910 |
| Cumulative % of vegetation data | 7.32 | 13.60 |
| Cumulative % of vegetation-environment relationship | 42.8 | 79.8 |
| Total inertia | 3.0172 | |

Table 4. Canonical correspondence analysis for vegetation in the Qarlikturvik valley of Bylot Island, Nunavut. CCA-1: first canonical axis; CCA-2: second canonical axis. Statistically significant values ($P < 0.05$) after 999 permutations are shown in bold. Standing dead represents the cover of dead attached vascular plants.

| Environmental variables | df | χ^2 | F | P | CCA-1 | CCA-2 |
|-------------------------|-----|----------|---------|--------------|---------|---------|
| Litter cover | 1 | 0.1523 | 11.4495 | 0.008 | 0.0206 | -0.8681 |
| Thaw-front depth | 1 | 0.1768 | 13.2854 | 0.001 | -0.8608 | -0.1436 |
| Soil moisture | 1 | 0.0429 | 3.2206 | 0.002 | -0.7339 | 0.2424 |
| Standing dead | 1 | 0.0714 | 5.3688 | 0.01 | -0.2393 | 0.2588 |
| Bare ground | 1 | 0.0065 | 0.4889 | 0.629 | 0.0207 | -0.0522 |
| Standing water | 1 | 0.0090 | 0.6759 | 0.534 | -0.2365 | 0.0830 |
| Grubbing | 1 | 0.0240 | 1.8003 | 0.097 | -0.1051 | 0.1983 |
| Goose feces | 1 | 0.0329 | 2.4717 | 0.028 | 0.4769 | -0.0772 |
| Residual | 188 | 2.5014 | | | | |

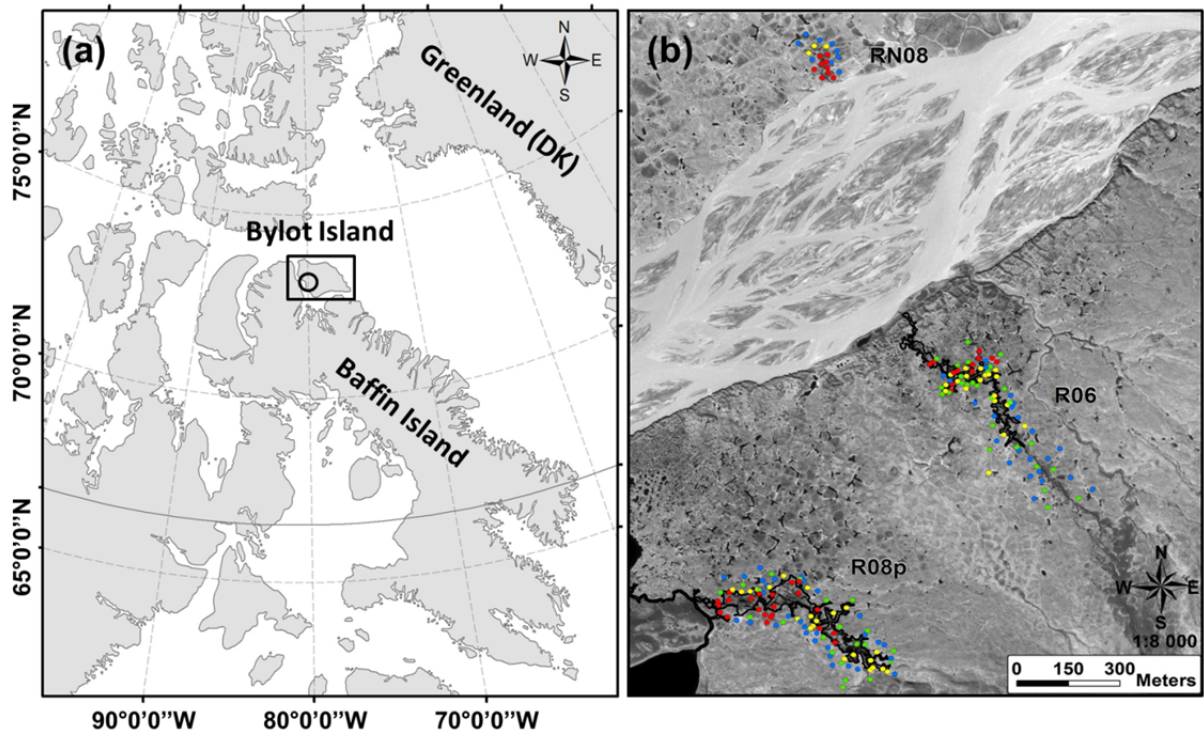


Figure 1. Location of the study area. (a) The Qarlikturvik valley (black circle) of Bylot Island, Nunavut (black rectangle). (b) the 197 sites sampled along three gullies (wet polygons (n = 62): blue circles; < 5-year disturbed polygons (n = 44): yellow circles; > 5-year disturbed polygons (n = 43): red circles; mesic environments (n = 48): green circles).

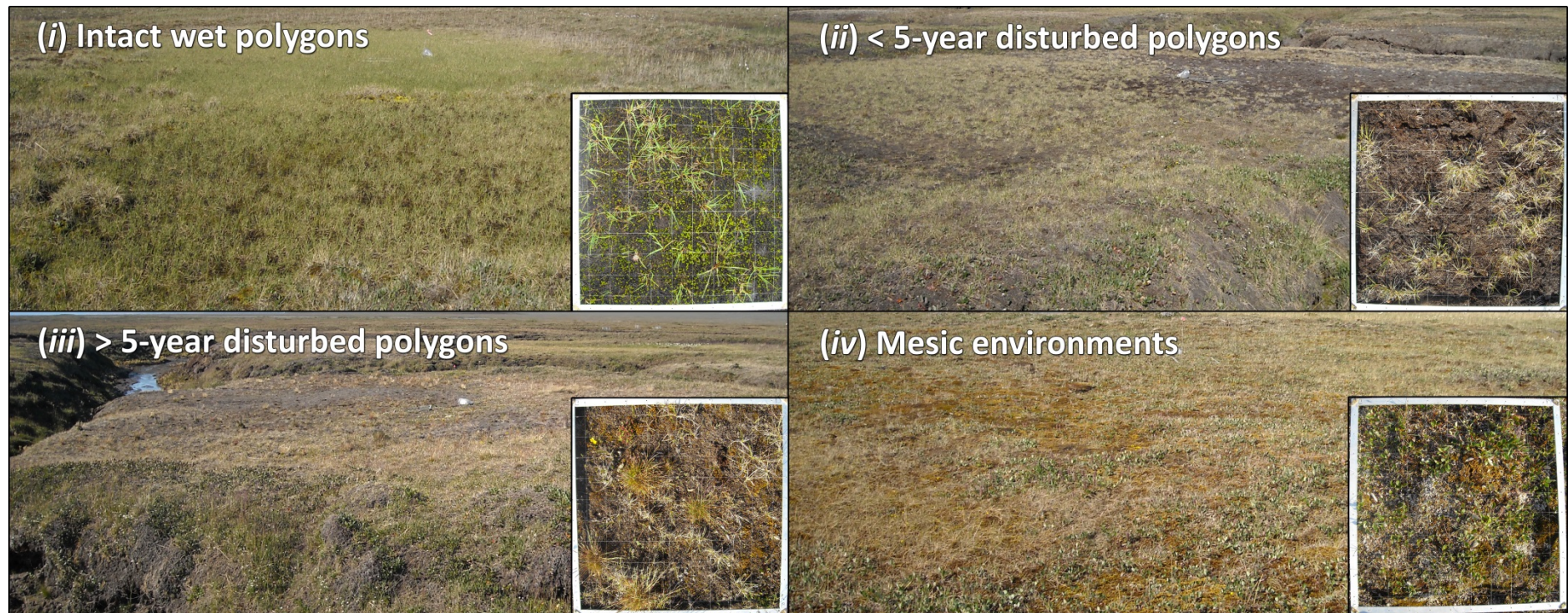


Figure 2. Illustration of the four habitats studied in the Qarlikturvik valley of Bylot Island, Nunavut. The close view at the bottom right of each picture represents the 70 cm x 70 cm quadrats that were used to determine species richness and abundance in each sampling site.

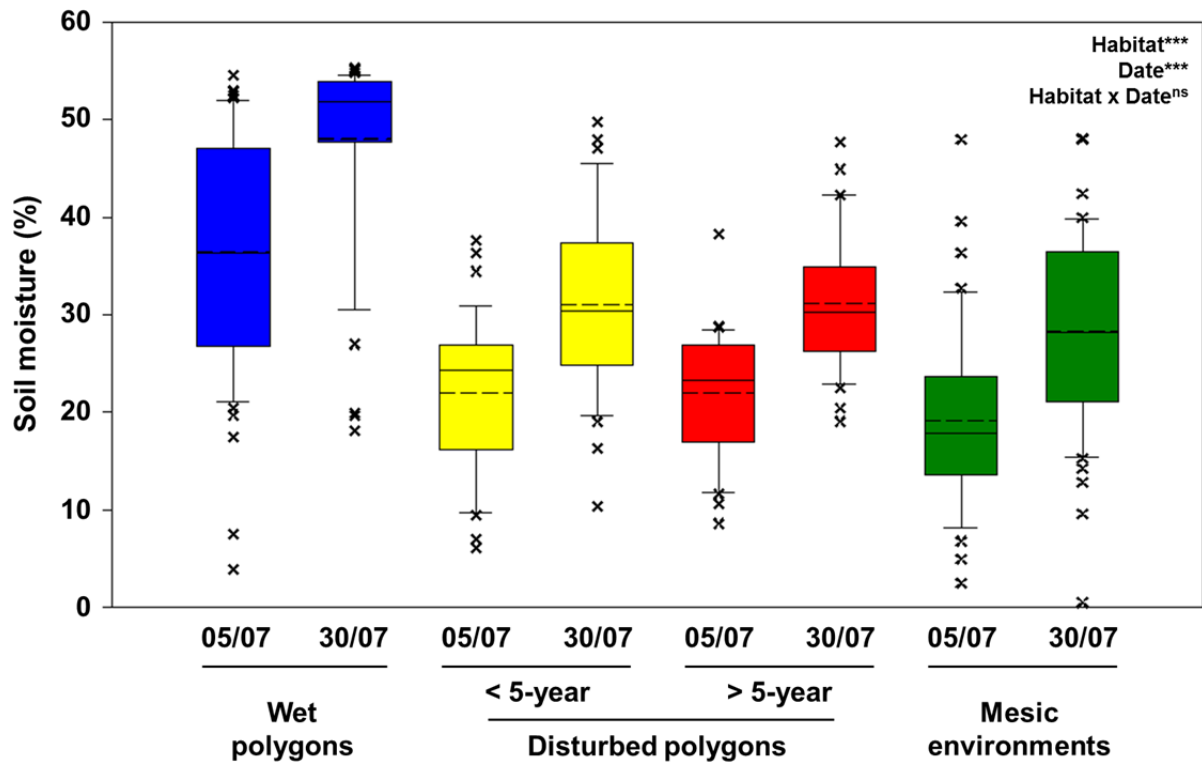


Figure 3. Soil moisture monitored early and late July 2010 in the four habitats studied in the Qarlikturvik valley of Bylot Island, Nunavut. 10th percentiles, lower quartile, median, mean (dash line), upper quartile and 90th percentiles are shown. See Table A1 for sample sizes and post-hoc contrasts. *** $P < .001$, ns: statistically non-significant effect.

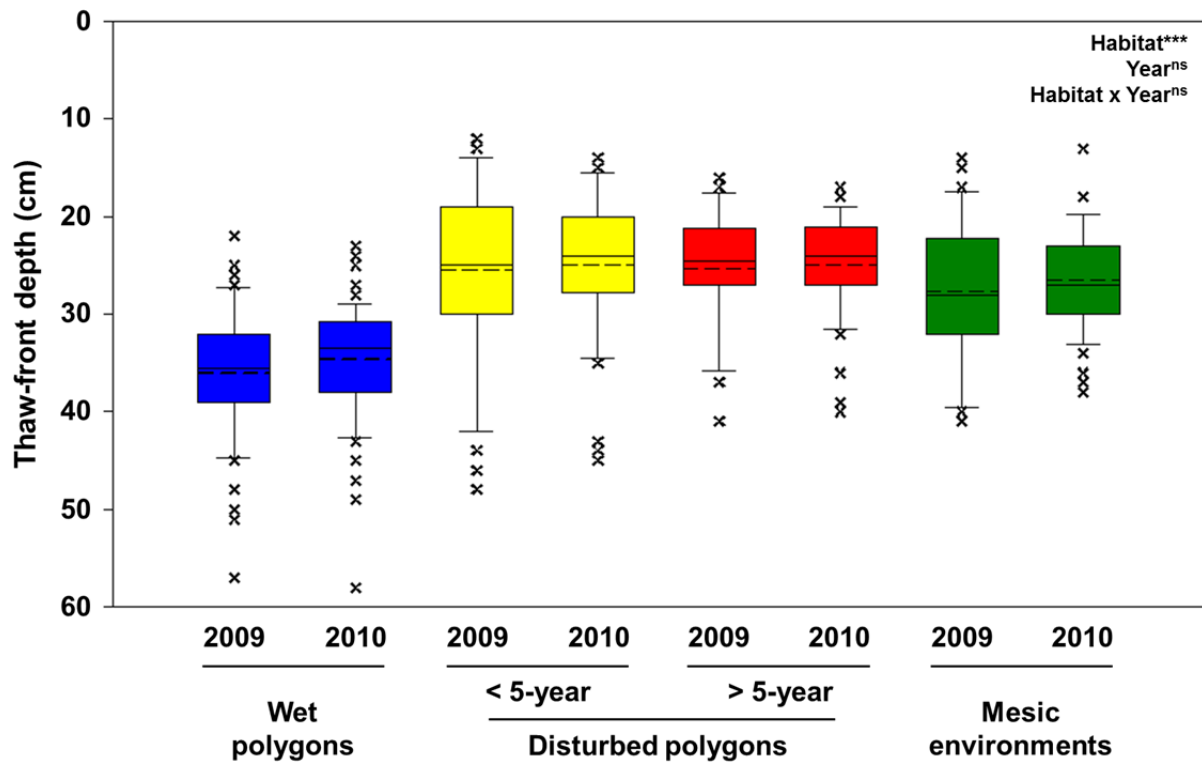


Figure 4. Thaw-front depth monitored in July 2009 and 2010 in the four habitats studied in the Qarlikturvik valley of Bylot Island, Nunavut. 10th percentiles, lower quartile, median, mean (dash line), upper quartile and 90th percentiles are shown. See Table A1 for sample sizes and post-hoc contrasts. *** $P < .001$, ns: statistically non-significant effect.

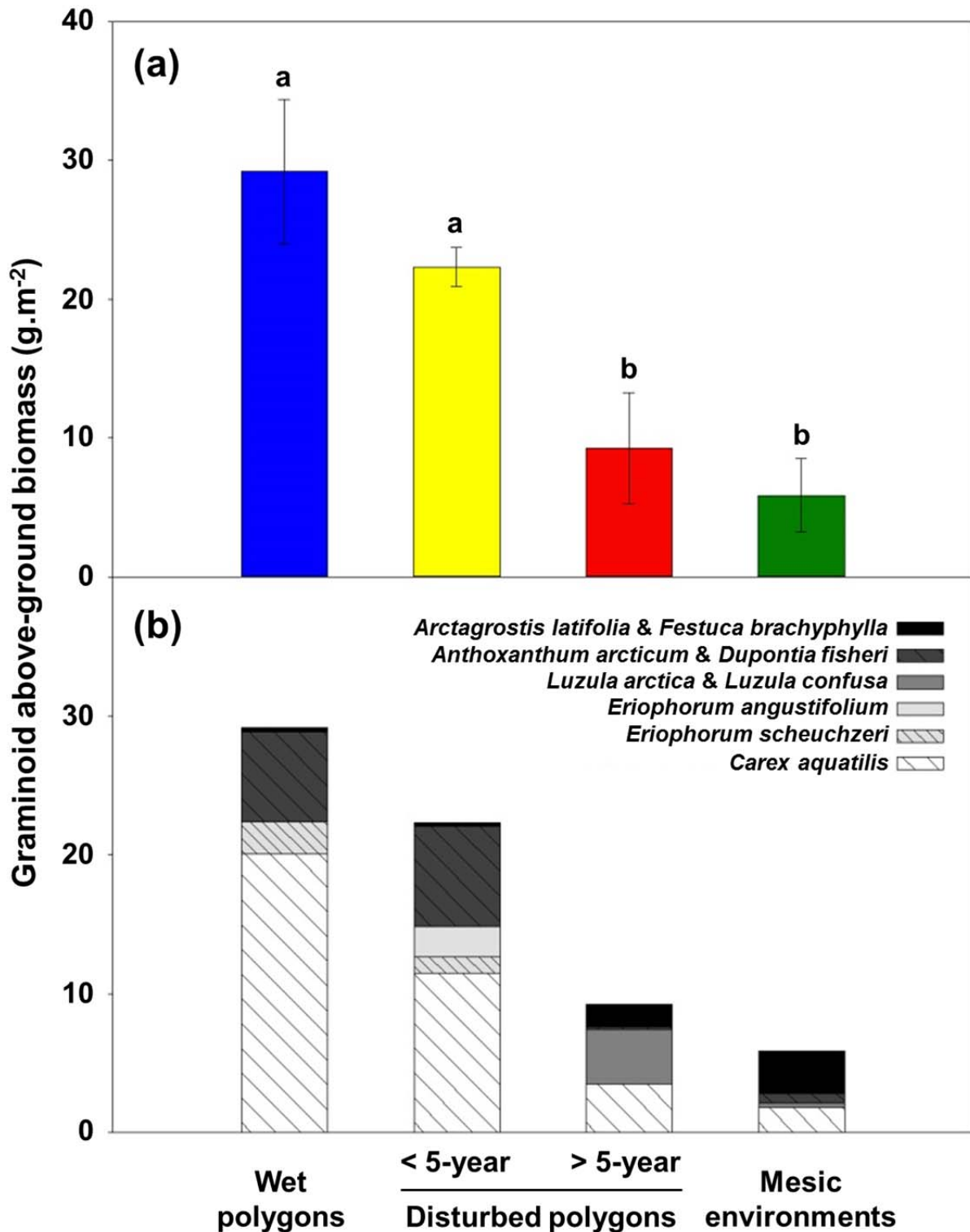


Figure 5. Above-ground biomass of graminoids growing in the four habitats studied in the Qarlikturvik valley of Bylot Island, Nunavut. (a) Mean total biomass (\pm SE), and (b) mean biomass per species. In panel (b), dashes represent hydrophilic species. $n = 5$ per habitat. Species aforementioned belong to the following families: Cyperaceae (*Carex aquatilis*, *Eriophorum angustifolium* and *Eriophorum scheuchzeri*), Juncaceae (*Luzula arctica* and *Luzula confusa*), and Poaceae (*Anthoxanthum arcticum*, *Arctagrostis latifolia*, *Dupontia fisheri* and *Festuca brachyphylla*).

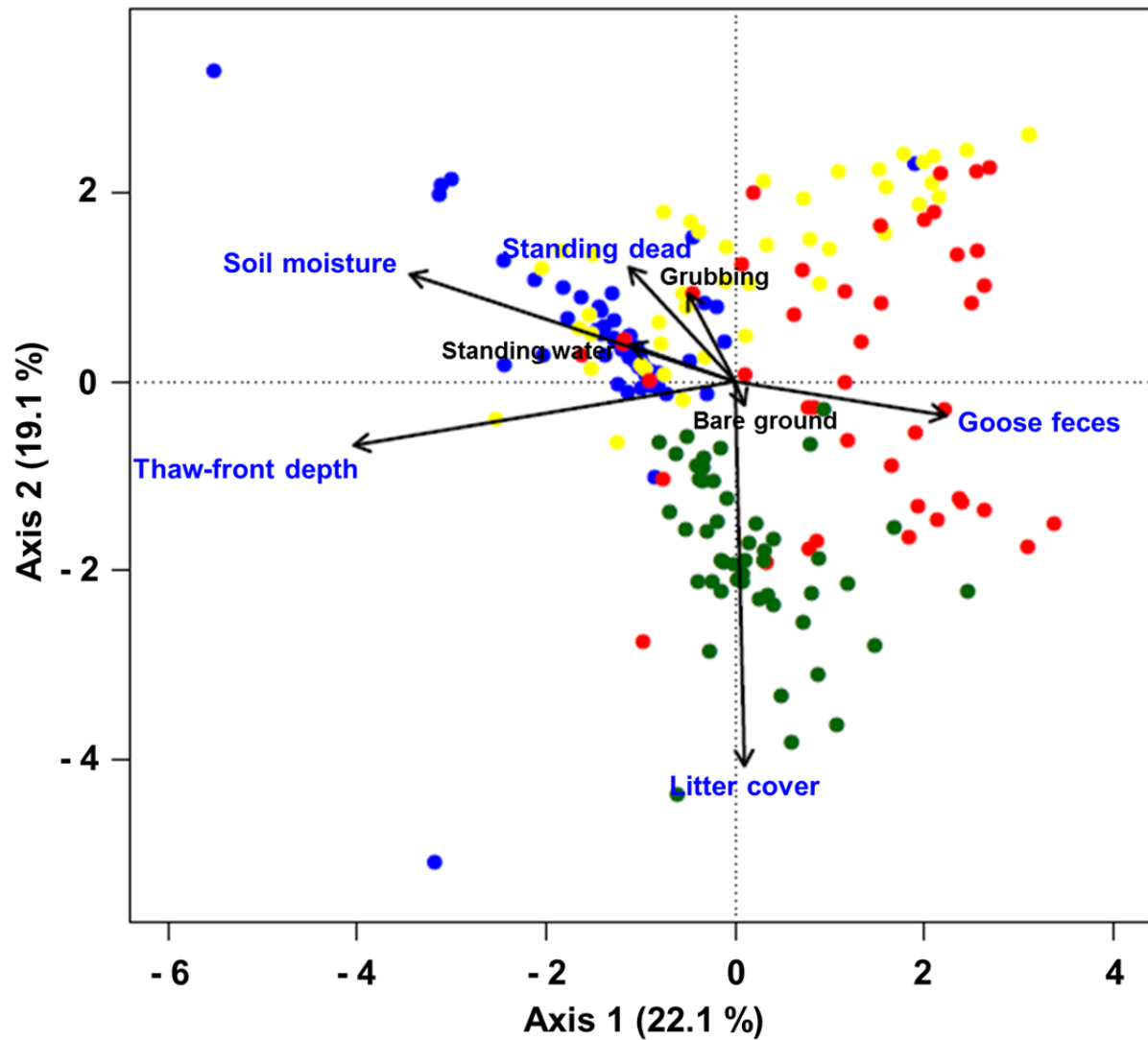


Figure 6. Canonical correspondence analysis (CCA) ordination of the 197 sites sampled in the Qarlikturvik valley of Bylot Island, Nunavut. Wet polygons (n = 62): blue circles; < 5-year disturbed polygons (n = 44): yellow circles; > 5-year disturbed polygons (n = 43): red circles; mesic environments (n = 48): green circles. Blue font indicates statistically significant environmental variables (see Table 4). Standing dead represents the cover of dead attached vascular plants.