

EUROSPEC: At the interface between remote sensing and ecosystem CO₂ flux measurements in Europe

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Abstract

Resolving the spatial and temporal dynamics of gross primary productivity (GPP) of terrestrial ecosystems across different scales remains a challenge. Remote sensing is regarded as the solution to upscale point observations conducted at the ecosystem level, using the eddy covariance (EC) technique, to the landscape and global levels. In addition to traditional vegetation indices, the photochemical reflectance index (PRI) and the emission of solar-induced chlorophyll fluorescence (SIF), now measurable from space, provide a new range of opportunities to monitor the global carbon cycle using remote sensing. However, the scale mismatch between EC observations and the much coarser satellite-derived data complicates the integration of the two sources of data. The solution is to establish a network of *in situ* spectral measurements that can act as bridge between EC measurements and remote sensing data. *In situ* spectral measurements have been already conducted for many years at EC sites, but using variable instrumentation, setups, and measurement standards. In Europe in particular, *in situ* spectral measurements remain highly heterogeneous. The goal of EUROSPEC Cost Action ES0930 was to promote the development of common measuring protocols and new instruments towards establishing best practices and standardization of these measurements. In this review we describe the background and main tradeoffs of *in situ* spectral measurements, review the main results of EUROSPEC Cost Action, and discuss the future challenges and opportunities of *in situ* spectral measurements for improved estimation of local and global estimates of GPP over terrestrial ecosystems.

Keywords: Carbon balance, Chlorophyll Fluorescence, COST Action, Flux Measurements, ICOS, Optical Sampling, Proximal Sensing, Remote Sensing, Spectral Reflectance, Greenness, SpecNet, FluxNet, Vegetation Indices.

1. Introduction

Accurate quantification of carbon fluxes across space and time is of primary importance to climate scientists, land use managers, and policymakers (Beer et al., 2010; Ciais et al., 2014; Joos et al., 2001). Carbon budgets can be estimated with high accuracy at the ecosystem level (e.g. Clement et al., 2012; Grace et al., 2006; Zanutelli et al., 2015) but global estimates of Gross Primary Productivity (GPP) and carbon balance in terrestrial ecosystems still have high levels of uncertainty (Alton, 2013; Balzarolo et al., 2014; Beer et al., 2010; Enting et al., 2012; Jung et al., 2011; Keenan et al., 2012; Piao et al., 2013). The primary method used to measure the net flux of carbon dioxide (CO₂) between the Earth's surface and the atmosphere is eddy covariance (EC) (Aubinet et al., 2000; Baldocchi, 2008; Goulden, 1996). The EC technique has dramatically improved our understanding of inter and intra annual variations in the carbon fluxes at the ecosystem level (Baldocchi, 2008). However, upscaling these local observations to the landscape and regional level remains a challenge. Because EC measurements require sites with moderate to low slopes, there is a bias towards certain topography and ecosystem types (Göckede et al., 2008). In addition, the footprint of EC measurements is not constant and varies with wind direction and speed, measurement height, and vegetation structure (Schmid, 2002; Vesala et al., 2008). Although footprint size can be controlled to some extent with tower height, the resulting data may still represent different vegetation communities depending on time-varying wind direction. Footprint analysis (e.g. Kormann and Meixner, 2001) is required to cope with this source of variability that complicates the interpretation of flux data. Despite these limitations, the global number of active flux sites exceeds 500 and is constantly increasing (Schimel et al., 2015). The question remains as to how to better upscale these point measurements to the landscape, regional and global scale.

Given that most of the factors affecting carbon fluxes have strong spatial and temporal components it is difficult to envisage upscaling without the use of remote sensing data, the only means to provide regular and spatially continuous observations of the Earth surface. One of the most widely applied approaches to assimilate remotely sensed data is to estimate GPP through a light use efficiency (LUE) model (Monteith, 1972; Reichstein et al., 2014; Ruimy et al., 1994):

$$GPP = PAR \times fAPAR \times LUE \quad (1)$$

where GPP is expressed as a function of the incident photosynthetically active radiation (PAR), the fraction of this PAR that is eventually absorbed by vegetation (fAPAR), and the efficiency by which absorbed PAR is used to assimilate atmospheric CO₂, termed the light use efficiency (LUE) (Hilker et al., 2008a). The fAPAR in Eqn 1 is a function of canopy chlorophyll content or green

1 biomass (i.e. more chlorophyll results in more absorption). It is important to note that fAPAR in
2 Eqn 1 corresponds to green fAPAR, in contrast to total canopy fAPAR where both photosynthetic
3 and non-photosynthetic elements such as wood contribute to PAR absorption. Green fAPAR has
4 been widely estimated using reflectance-based vegetation indices as proxy, notably the Normalized
5 Difference Vegetation Index (NDVI) derived from red and near-infrared (NIR) reflectance (Rouse
6 et al., 1973; Tucker, 1979). These vegetation indices correlate better with green fAPAR than with
7 total fAPAR because their spectral formulation can significantly discriminate green from non-green
8 elements (Gamon et al. 1995). However, canopy structural factors, background properties, or sun-
9 target-sensor geometry can complicate the estimation of green fAPAR with vegetation indices (Di
10 Bella et al. 2004; Gamon 2015, Knyazikhin et al.2013).Vegetation indices (VIs) have been
11 successfully used to track seasonal dynamics in GPP in ecosystems characterized by strong seasonal
12 dynamics in green biomass such as croplands, grasslands, broadleaf forests (Gitelson et al., 2006,
13 2008, 2012; Harris and Dash, 2010; Peng and Gitelson, 2012; Rossini et al., 2010).

14 In ecosystems dominated by evergreen species, the seasonal variation in GPP can be strongly
15 controlled by LUE in addition, or instead of, fAPAR (e.g. Garbulsky et al., 2008; Gamon, 2015).
16 The LUE term is usually estimated as the product of the potential maximum LUE (ϵ_{\max}) and an
17 environmental scalar (f) expressing the influence of one or several environmental stress factors
18 constraining ϵ_{\max} . For example, in the MODIS GPP product (Running et al., 2004), LUE is
19 estimated by using a plant functional type or biome-dependent ϵ_{\max} , and climate variables
20 describing the environmental scalar f (i.e. air temperature and vapour pressure deficit). The problem
21 of this approach is that maximum LUE depends only on vegetation type (Xiao et al., 2004a,b; Zhao
22 et al., 2005), while inter-seasonal variability due to plant phenology and photosynthetic dynamics
23 (Lagergren et al., 2005) is only considered via the instantaneous effect of the environmental scalar,
24 which cannot reproduce the slow response dynamics of vegetation.

25 Importantly, LUE generates optical signatures that can be measured with optical remote sensing
26 instruments mounted on airborne or satellite platforms. These signatures are the photochemical
27 reflectance index (PRI), and the emission of solar-induced chlorophyll-a fluorescence (SIF) by
28 vegetation. The PRI uses the reflectances at 531 and 570nm to capture the temporal dynamics in
29 LUE via variations in the xanthophyll-cycle pigments and the relative ratio of carotenoids to
30 chlorophyll found in foliage (Gamon et al. 1992; Porcar-Castell et al. 2012; Wong and Gamon
31 2015). The PRI has been successfully estimated from sensors mounted on multiple platforms
32 including towers, aircrafts and satellites (e.g. Drolet et al., 2005; Garbulsky et al., 2008; Nichol et
33 al. 2002). In contrast, SIF are photons of red/far-red light (660-800nm) emitted during the first steps

1 of photosynthesis. Accordingly, the emission of SIF is expected to depend on both fAPAR and LUE
2 (Porcar-Castell et al. 2014). Despite the challenge of measuring SIF, due to the small intensity of
3 the signal relative to that of the reflected light, recent technical advances make it now possible to
4 estimate SIF from towers, aircrafts and satellites (Frankenberg et al., 2011; Guanter et al. 2007;
5 Joiner et al., 2011; Porcar-Castell et al., 2014; Rossini et al., 2010; 2015). Overall, the growing
6 number of satellite missions with enhanced capacity to retrieve PRI and SIF at increasing spatial
7 and temporal resolutions (e.g. Clevers and Gitelson, 2013; Frankenberg et al., 2014; Guanter et al.
8 2015; Guan et al., 2015) open up new possibilities to improve carbon models and upscale EC data.

9 Noting that the relationship between NDVI and fAPAR tends to saturate at high canopy densities
10 (Myneni and Williams, 1994; Olofsson and Eklundh, 2007) also other approaches have been used to
11 estimate vegetation carbon uptake. For example, the Enhanced Vegetation Index (EVI) (Huete et
12 al., 2002) efficiently describes the seasonal variability in GPP across both dense and sparse
13 vegetation canopies (Schubert et al., 2010, 2012; Sims et al., 2006; Sjöström et al., 2011; Xiao et
14 al., 2004a, 2004b; Xiao et al., 2010). More recently, the plant phenology index (PPI) (Jin and
15 Eklundh, 2014) has been shown to be linearly related to green leaf area index (LAI), and better
16 related to seasonal GPP variations than NDVI and EVI of coarse-resolution MODIS data at
17 northern latitudes. This illustrates the value of investigating the relationship between carbon uptake
18 and spectral information in flux footprint areas beyond the LUE model depicted in Eqn 1.

19 Integrating satellite and EC data into large scale carbon models is not straightforward. The spatial
20 mismatch between EC measurements and coarser grid-cell information in models and most satellite-
21 derived remote sensing data adds significant uncertainty (Chen et al., 2012; Oren et al., 2006). The
22 large viewing angle of many satellite products, e.g. MODIS, results in ill-defined and variable
23 footprint areas leading to additional geometric uncertainties (e.g. Tan et al., 2006). Furthermore,
24 airborne and space-borne data needs to be corrected for atmospheric absorption and scattering
25 effects (Karpouzli et al., 2003; Richter, 2011), a process that again can add further uncertainty
26 (Drolet et al., 2005; Hilker et al., 2009). All these physical limitations could be substantially
27 reduced by including *in situ* long-term spectral measurements to the network of EC flux sites
28 (Gamon et al., 2010; Hilker et al., 2009).

29 Coordinated *in situ* spectral measurements require a network of stable sensors that follow the same
30 measurement standards, calibration protocols, and have traceable technical specifications to allow
31 across-site comparisons. Following the example of the EC Fluxnet community
32 (<http://fluxnet.ornl.gov/>), SpecNet (www.specnet.info) was set to cover the needs for networking

1 and standardization of optical measurements across flux sites (Gamon et al., 2006, 2010). Although
2 the geographical coverage of SpecNet collaborators is constantly increasing the network was until
3 very recently strongly biased towards North American sites. The European community of scientists
4 conducting optical measurements remained highly dispersed and heterogeneous. It is in this context
5 that the EUROSPEC COoperation in Science and Technology (COST) Action ES0903 ([http://cost-](http://cost-es0903.fem-environment.eu/)
6 [es0903.fem-environment.eu/](http://cost-es0903.fem-environment.eu/)) originated in 2009 and operated until 2013. European COST Actions
7 are four year projects aimed at promoting trans-national cooperation and networking among
8 scientists and engineers across Europe and beyond (see www.cost.eu for further details).

9 The goal of EUROSPEC COST was to promote the development of common measuring protocols
10 and new instruments for *in situ* spectral measurements, bringing together scientists and industries in
11 order to increase the reliability, value and cost-efficiency of such measurements. This was done so
12 that field-installed spectral sensors could be used as “bridge” between the EC and optical remote
13 sensing communities.

14 The Action was divided in four Working Groups (WG): WG1, network and state-of-the-art
15 characterization. The goal was to characterize the variability of spectral measurements and methods
16 being used across flux sites in Europe. WG2, intercomparison and standardization of instruments.
17 The goal was to characterize the sources of variability between sensors, methods and protocols.
18 WG3, New instruments. The goal was to promote the development of new instruments that better
19 match sensor design, specifications, cost and purpose. And WG4, upscaling methods. The goal was
20 to evaluate challenges and tools to upscale point observations to the footprint area and beyond.

21 The main objective of this review is to contextualize and synthesize the accomplishments made
22 during EUROSPEC and to identify a number of challenges and opportunities for the near future.
23 We review the main outcomes from EUROSPEC. We describe the background of *in situ* long-term
24 spectral measurements and their main tradeoffs, followed by presenting the main results of each
25 EUROSPEC WG and by a final discussion on future challenges and opportunities of these
26 measurements.

27 **2. *In situ* long-term spectral measurements: principles and trade-offs**

28 Remote sensing measurements can be collected from platforms that may operate at variable
29 distance from the Earth’s surface: from satellites for regional-global extent measurements, to field
30 spectrometers mounted on top of towers for close-range *in situ* measurements. In between these two
31 scales are airborne platforms including piloted and unpiloted aircraft, kites and blimps that can

1 measure at multiple scales depending on height. In EUROSPEC we focused on long-term *in situ*
2 optical measurements conducted from EC towers.

3 There are a number of important differences between close-range *in situ* measurements and the
4 traditional remote sensing from aircraft or satellites. *In situ* measurements, sometimes referred to as
5 proximal sensing, are conducted at short distances and are to a large extent free from the
6 atmospheric absorption and scattering effects that affect traditional remotely sensed data (Cheng et
7 al., 2006; Meroni et al., 2009; Thenkabail et al., 2002). *In situ* measurements can be used to track
8 the spectral properties of individual biological elements (leaves, shoots, plants, homogeneous
9 canopies) while traditional remote sensing tends to measure at coarser scales where multiple
10 species, soil and non-vegetated areas may contribute to the measured signals. Most importantly, *in*
11 *situ* measurements can provide data at high-temporal resolution, something that cannot be
12 accomplished with traditional remote sensing. All these characteristics make *in situ* measurements
13 ideal to study and disentangle the link between optical signals and carbon flux dynamics, as well as
14 for calibrating and validating satellite data and atmospheric correction algorithms (Brook and Ben-
15 Dor, 2015; Czapla-Myers et al., 2015; Hilker et al., 2009).

16 *In situ* spectral measurements involve the measurement of the down-welling (incoming) and up-
17 welling (both reflected and emitted) radiation fluxes from the Earth surface. These measurements
18 can be conducted with variable setup and approaches and the optimal solution will depend on the
19 purpose, characteristics of the site and amount of resources available.

20 **2.1 Single vs Dual Field of View (SFOV vs DFOV)**

21 Measurements of down-welling and up-welling radiation can be carried out either in sequence
22 (when a single sensor or spectrometer is used), or simultaneously (when two separate sensors or
23 spectrometers are used) (Fig. 1). These are also addressed as single beam/field-of-view (SFOV) or
24 dual beam/Field of view (DFOV) configurations, respectively, and have their own advantages and
25 disadvantages (See Table 1). A SFOV system is generally configured with a single sensor (or
26 spectroradiometer) and will be generally cheaper to set up than a DFOV using two sensors. Having
27 a single sensor means also that there is no need to inter-calibrate the sensor pair. However, long-
28 term and unattended measurements with a SFOV system face the challenge of automating a single
29 sensor/spectrometer to measure both down-welling and up-welling radiation. This automation
30 usually involves moving parts (e.g Meroni et al., 2011; Sakowska et al., 2015), which may become
31 a problem for long-term field operation under certain environments and entail a time delay between
32 up-welling and down-welling measurements, which in turn may generate noisy data under cloudy

1 conditions. Similarly, DFOV systems have also associated advantages and disadvantages. Because
2 radiometric measurements are temperature-sensitive (Saber et al., 2011), DFOV based on two
3 spectrometers are particularly sensitive to temperature. A practical solution is to keep the two
4 sensors at constant temperature, e.g. by housing them in a temperature-controlled enclosure (e.g.
5 (Drolet et al., 2014), but this might not be always possible due to power limitations. Also, regular
6 intercalibration of the two sensors will be essential in DFOV measurements (see e.g. Anderson et
7 al., 2006; Gamon et al., 2015; Jin and Eklundh, 2015). Additionally, long-term measurements with
8 a DFOV may be constrained by aging-dependent degradation of the two sensor heads (see Section
9 3.2.4). These limitations were partly overcome with new DFOV systems developed during
10 EUROSPEC that include a single spectrometer (see Section 3.3). The advantage of a DFOV system
11 is that it guarantees quasi simultaneous measurements of down-welling and up-welling radiances,
12 each within a few hundred milliseconds of the other and may be easier to automate because do not
13 require moving parts to shift from up-welling to down-welling measurements. Importantly, systems
14 such as the Piccolo and SIF-Sys (developed during EUROSPEC and described in section 3) share
15 the benefits from both SFOV and DFOV systems as they include a single spectrometer but make
16 use of bifurcated fibre optics to sample two fields of view (Fig. 1).

17 **2.2 Multispectral vs Hyperspectral Sensors**

18 Spectral information can be acquired at different spectral resolution, which depends on the sampling
19 intervals (discrete or continuous) and the width of the spectral bands (Fig. 2). Accordingly, spectral
20 measurements can be classified into multispectral or hyperspectral. Multispectral sensors measure a
21 limited number of spectral bands, from two to five bands found for example in the SKYE SKR-
22 1800 or 1860 sensor series (Skye Instruments Ltd, UK) the Decagon-SRS series (Decagon devices
23 Inc, WA, USA) or the Cimel five-band sensors (Cimel Electronique, FR), or up to 16 bands found
24 in the Cropscan MSR16R (Cropscan Inc. MN, USA) (Balzarolo et al., 2011; Sakowska et al.,
25 2014). The bandwidth of these sensors (in terms of full width at half the maximum response,
26 FWHM) is in the order of 10 nm or greater and the sampling across a specific spectral range is
27 typically discrete (Fig. 2). These sensors are typically manufactured using optical filters, light
28 emitting diodes (LEDs) and photodiode detectors (Norton, 2010; Ryu et al., 2010). These sensors
29 are characterized by relatively low cost (from a few hundred to a few thousand euros/dollars), ease
30 of maintenance, weather-proof design, and low power consumption. Hence, they are useful and
31 affordable instruments for deployment at flux tower sites for extended periods of time. In addition,
32 their relatively low cost allows mounting of several sensors in different positions to study spatial
33 heterogeneity. Multispectral sensors can be deployed to measure a number of vegetation indices

(e.g. NDVI or PRI) to track and study vegetation phenology and seasonality. They can also be used to produce satellite calibration and validation data, provided that their spectral configuration can be related to that of the spaceborne sensor.

In contrast, hyperspectral sensors (more often addressed as spectrometers, or spectroradiometers when they are radiometrically calibrated) can measure hundreds of spectral bands, often 250 or more, with bandwidths usually less than 10 nm full width at half maximum (FWHM) and sampling intervals from less than 1 nm to 10 nm depending on configuration (Fig. 2). The obvious advantage of hyperspectral sensors is that they can resolve more detailed features of the vegetation (Milton et al., 2009) and serve to estimate parameters that require higher spectral resolution, such as the emission of chlorophyll fluorescence (Meroni et al., 2009). Moreover, since hyperspectral information can be resampled to coarser spectral resolutions, data from hyperspectral systems can be flexibly convoluted to match spectral bands of different remote sensors (Olsson et al., 2011) increasing its value as a source of satellite calibration and validation data. In addition, hyperspectral data can be used to mine new spectral band combinations to match different ecosystem variables (e.g. Balzarolo et al., 2015; Heiskanen et al., 2013; Inoue et al., 2008; le Maire et al., 2008; Milton et al., 2009; Tagesson et al., 2015; Wang et al., 2011; Yao et al., 2010). Milton and coworkers (2009) presented an extensive review of how hyperspectral proximal sensing, or field spectroscopy, has developed and listed the most commonly available field spectrometers. Spectrometers are very complex opto-electro-mechanical instruments and tend to be expensive, from a few thousand Euros/dollars for optical benches measuring in the VNIR, to several tenths of thousands for field instruments measuring in the VNIR and the SWIR. The common limitation of all these spectrometers is that they are not designed for unattended or long-term field operation. Accordingly, users need to build their own weatherproof housing, power supply, automatic datalogging, and control units (see next section). As a result, the overall cost of these user-made systems is difficult to quantify because in addition to off-the-shelf components they involve plenty of in-house skilled technician hours. Field spectrometers are also more susceptible to physical damage (due to their inherent complexity), and are more difficult and expensive to automate for continuous or periodic logging applications. In addition, these systems tend to be considerably larger and heavier than their multispectral counterparts, presenting a structural challenge to their deployment on flux towers. Despite these limitations, the number of such measurements is rapidly increasing (Drolet et al., 2014; Huber et al., 2014; Pacheco-Labrador and Martín, 2015; Rossini et al., 2012; Sakowska et al., 2015).

2.3 Instrument Configurations: hemispherical-conical, bi-conical and bi-hemispherical

Reflectance factors relate the radiant flux reflected by a target surface to the radiant flux incident on it and they can be measured using different instrument configurations (see Schaepman-Strub et al., 2006 for a full mathematical explanation of the different factors and terms). Three main instrument configurations have been applied to *in situ* field measurements to quantify incoming and reflected radiation and estimate reflectance factors: bi-conical, hemispherical-conical and the bi-hemispherical configurations (Fig. 1).

Hemispherical-conical measurements use a foreoptic diffuser assembly, designed to have a cosine response at changing solar zenith angle to estimate down-welling irradiance, and a conical foreoptic for upwelling measurements which can be installed at nadir or off-nadir. The hemispherical-conical configuration lends itself to both multispectral and hyperspectral measurements from flux towers.

Bi-conical measurements rely on a diffuse white reference panel, typically of Spectralon® (Labsphere Inc., NH, USA), reflecting down-welling solar radiant flux, normally viewed from nadir through a fixed angularly limited (conical) field-of-view foreoptic, to provide the reference measurement. The potential limitation of using a reference panel is that it needs to be kept clean and stable over time which may become a challenge in the field due to particle deposition (but see Sakowska et al., 2015 in Section 3). In practice, because both direct and diffuse light contribute to the up-welling signal reflected from the reference panel when measuring under field conditions, field data obtained in a bi-conical instrument configuration can be used to derive hemispherical-conical reflectance factors (HCRF) (Schaepman-Strub et al., 2006).

Bi-hemispherical measurements use a foreoptic diffuser to capture both down-welling and up-welling irradiance. Bi-hemispherical measurements require a nadir-view installation for both sensor heads and have the great advantage of enabling the sampling of a wider area. The main limitation of this configuration is that while the hemispherical-conical measurements can be taken by observing the canopy at nadir or off nadir, all viewing directions (both nadir and off-nadir) contribute to the bi-hemispherical measurements (Meroni et al., 2011). For this reason, bi-hemispherical measurements tend to be more sensitive to variations in illumination geometry compared to hemispherical-conical measurements collected with a nadir view particularly for large illumination zenith angles.

1 **3. EUROSPEC Main Results**

2 **3.1. Network and State-of-the-art characterization**

3 When research questions become global, as in the case of global carbon cycle monitoring,
4 networking becomes a key methodological element to ensure consistent implementation of
5 measuring protocols, data sharing and management. One of the objectives of EUROSPEC was to
6 contribute to build up a distributed European spectral sampling network to foster data sharing in
7 order to better understand relationships between optical responses of vegetation and carbon cycle.

8 *In situ* long-term spectral measurements at flux towers are still accomplished with instruments that
9 are of variable design and performance, use different configurations, are installed with contrasting
10 geometries, and are conducted with different calibration or quality assessment regimes. In
11 EUROSPEC, we conducted a detailed review of proximal sensing measurements based on the
12 responses to a questionnaire obtained from groups working at 40 flux tower sites in Europe
13 including two sites from Africa and Australia (Balzarolo et al., 2011). *In situ* measurements
14 included SFOV and DFOV systems; bi-hemispherical or hemispherical-conical configurations, and
15 both multispectral and hyperspectral sensors. The study portrayed the lack of consensus on what are
16 the most suitable proximal sensing systems and methods to support EC measurements. No
17 standards were being applied in terms of system performance (e.g. non-linearity in response, signal-
18 to-noise ratios, and cosine response of down-welling radiant flux foreoptics); measurement
19 geometries (e.g. hemispherical vs conical and their combinations); different foreoptic field-of-
20 views; installation geometry (e.g. nadir or off-nadir; height of sensor above target surface), or
21 calibration regimes (e.g. regularly calibrated by manufacturer, calibrated *in situ* or even not
22 regularly calibrated) (Balzarolo et al., 2011). The lack of regular calibrations was presented as a
23 fundamental limitation to overcome in order to produce high quality data, reliability of time series
24 analysis and to enable inter-comparison of results between network sites, e.g. Integrated Carbon
25 Observation System (ICOS) sites (<https://www.icos-ri.eu/>). At the end of EUROSPEC, the need to
26 standardize these measurements still exists.

27 Balzarolo et al. (2011) suggested also that two possible levels of instrumentation could be
28 considered. The first; termed the Basic Standard, would include only multispectral broadband
29 sensors to estimate selected vegetation indices. The second, termed Advanced Standard, would
30 augment these multispectral sensors sites with hyperspectral sensor systems. The question remains
31 as to what specific instruments and sensors would be more appropriate in each case. Anderson et al.
32 (2013) conducted a field intercomparison experiment to assess the reproducibility of measurements

collected by different sensors used at flux tower sites. The analysis showed that lower-cost spectroradiometer systems performed similarly to more costly models and suggested that cost-effective and accurate measurements in the PAR range can also be acquired using lower-cost instrumentation. Similar conclusions were obtained by Harris et al. (2014) when they compared the performance of lower-cost multispectral sensors with a reference spectroradiometer to estimate the photochemical reflectance index (PRI). Another conclusion of these studies was the importance of characterizing sensor properties to allow inter-comparison of results between sensors and sites (see next section). Further long-term field instrument intercomparisons will be needed before final conclusions can be drawn from these studies.

EUROSPEC managed to establish an active network including scientists from 28 countries from Europe and beyond. This network remains active under a new COST Action (OPTIMISE-ES1309) and under the umbrella of SpecNet. Together we hope to continue promoting the standardization and implementation of optical measurements across flux sites.

3.2 Sources of variability

Limited consideration had been given to the comparability of spectral measurement protocols and systems before EUROSPEC (but see Castro-Esau et al., 2006; Pfitzner et al., 2011). Because the use of the same type of sensor in all sites is not realistic, and perhaps neither desirable as it would undermine the development of new sensors, we need to examine the factors that influence the variability in the data collected with different sensors. In EUROSPEC we dealt with a number of these factors: linearity of spectrometer response, impact of cosine diffusers and reference panels, effect of sensor FOV, and temporal stability of measurements and calibrations.

3.2.1 Linearity of spectrometer response

Linearity refers to the linear relationship between the signal generated by a radiation sensor and the impinging light power. Any dependence of this relationship on third or more variables leads to a systematic error in the measurements which needs to be characterized and corrected. Pacheco-Labrador et al. (2014) and Pacheco-Labrador and Martín (2014) assessed the linearity of one of the commercial field portable spectroradiometers, the Unispec-DC (PP Systems) currently used in unattended systems at EC sites (Hilker et al., 2010; Pacheco-Labrador and Martín, 2015), finding that both the gray level measured and also the integration time had an effect on linearity. They showed that non-linearity could be a significant problem in hyperspectral proximal sensing, especially for *in situ* and long-term unattended measurements. The impact of gray level-dependent

1 non-linearities may be significant when estimating narrow band indices, such as the photochemical
2 reflectance index (PRI), and therefore cannot be left uncharacterized and uncorrected (Pacheco-
3 Labrador and Martin, 2014). The impact of non-linearity can be minimized avoiding the most non-
4 linear region of the dynamic range. In turn, non-linearities related with the integration time affected
5 also the characterization of other instrumental artifacts (Pacheco-Labrador et al., 2014). This
6 dependence, previously reported in cameras (Ferrero et al., 2006) but not in field
7 spectroradiometers, is significant when the integration time is close to the readout time of the sensor
8 (i.e. a photodiode continues to collect photons during the time when the signal is being processed
9 (the readout time), producing an extra signal that is added to that obtained during the integration
10 time). Despite that integration-time dependent non-linearities have been characterized only in the
11 Unispec-DC (Pacheco-Labrador and Martín, 2015) it would be recommendable to avoid integration
12 times close to the instrument readout time, unless the integration-time dependent non-linearity has
13 been characterized.

14 **3.2.2 Cosine diffusers and reference panels**

15 Irradiance (i.e. downwelling radiant flux) can be measured with cosine corrected foreoptics pointed
16 vertically up, or with a diffuse white reference panel (see Fig. 1). Importantly, the materials,
17 calibration status, and method selected to measure irradiance may have an impact on the result.
18 Biggs et al. (1971) highlight the need for a properly designed foreoptic to avoid spectral variations
19 caused by changing sun azimuth and zenith angles. Malthus and Mac Lellan (2010) demonstrated
20 that the material selected for the foreoptic diffuser can significantly affect the spectra with
21 angular/wavelength dependencies and a poor cosine response. Similarly, they reported that the
22 performance of cosine diffusers in the Short Wave Infrared (SWIR) tends to be very poor with high
23 signal attenuation above 1400 nm (Malthus and Mac Lellan, 2010), and suggested that a diffuse
24 reference panel will provide a better cosine response than a cosine-corrected foreoptics.
25 Importantly, because reference panels present an angular-dependent time degradation at such
26 wavelengths, especially when used in the field (e.g. Anderson et al., 2002; Georgiev and Butler,
27 2007) (Fig. 3), recalibration and regular maintenance is essential. Recalibration requires a dedicated
28 laboratory facility as demonstrated by Georgiev et al. (2011) for the SWIR region of the solar
29 spectrum. In turn, Labsphere, the manufacturers of Spectralon® provide guidance on how to clean
30 reference panels. Consequent care, careful cleaning and recalibration of reference panels are
31 essential to minimise error propagation and uncertainties when conducting spectral measurements.

1 Selection of reference panel material is also very important. Manufacturers of reference panels for
2 spectroscopy such as LabSphere (Spectralon ®) or SphereOptics (Zenith polymer ®) use a sintered
3 fluoropolymer manufactured to have a very high reflectance, possibly in excess of 96% and
4 approximate a Lambertian reflectance across the 400 nm to 2,500 nm spectral region. Alternatively,
5 low cost PTFE sheets (i.e. Teflon) can be purchased at lower cost. However, PTFE sheets have
6 lower reflectance, approximately 80%, and have higher specular reflection. Also, because PTFE
7 sheets are not manufactured to be used as ‘references’ there may be variability between individual
8 sheets and wavelength dependent reflectances may be unknown. Overall, PTFE sheets are not
9 recommended as field spectroscopy reference standards.

10 Similarly, the material and design of cosine receptors affect the estimation of hemispherical
11 reference factors (Malthus and MacLellan, 2010) and consequently, the indices derived from them.
12 Therefore, significant and unquantified uncertainties will be introduced when comparing data from
13 sites that used different cosine receptors or sites characterized by a different range of Solar Zenith
14 Angles (SZA). For example, for SZA greater than 60° variation in reflectance factors obtained with
15 different cosine receptors can exceed 20%. This effect is reduced when considering normalized VIs,
16 and it depends on the spectral distance between bands selected for the VI calculation (Julitta, 2015).
17 Again, characterization of the properties of cosine diffusers and regular maintenance/replacement
18 should be included inside the measurement routine.

19 **3.2.3 The fields-of-view of field spectrometers and multispectral sensors**

20 Field spectroscopists normally assume that the Earth surface sampled by a non-imaging
21 spectrometer with a limited FOV foreoptic is spatially delimited by the solid angle specified by the
22 manufacturer, and that the response across the surface delimited by the FOV is the same for all
23 points inside the given FOV (Castro-Esau et al., 2006; Ferrier et al., 2009; Murphy et al., 2005;
24 Nichol and Grace, 2010). In practice, the spectral response within FOV of a field spectrometer is
25 not constant (i.e. certain areas within the FOV contribute more to the signal than others)
26 (MacArthur et al., 2012, Eklundh et al. 2011), this can be determined by the viewing angle and the
27 instrument’s Directional Response Function (DRF) (CIE, 1987) which can be characterized. The
28 DRF will be affected by both the internal design of the spectrometer (e.g. open path or fiber optic
29 transfer to individual detectors) and the foreoptics used. When measurements of heterogeneous
30 surfaces were simulated using the measured DRFs, significant differences were found between
31 simulated reflectance factors and those expected from the manufacturers’ specifications (Mac
32 Arthur et al., 2012). Even when less optically complex spectrometers, measuring only across the

1 VNIR region are considered, the Earth surface sampled is not necessarily that inferred from the
2 manufacturers' specified FOV included solid angle (Caras et al., 2011). The manufacturers of some
3 spectrometers now offer optical elements within their foreoptic mounts to defocus the foreoptics
4 and thereby homogenize the light received (e.g. the ASD FS pistol grip "scrambler"), or have
5 improved the optical components used to minimize 'chromatic aberrations' and heterogeneities and
6 again, homogenize the light received prior to it being distributed to the detectors (e.g. SVC HR-
7 2014i spectrometers). Therefore, the spectrometers' response should be more closely represented by
8 a Gaussian or Cauchy response, albeit centre-weighted, with all areas within the FOV represented in
9 the integrated measurement. These limitations affect the estimation of reflectance factors measured
10 from heterogeneous Earth surfaces (Mac Arthur et al. 2013) because the sample area is ill defined
11 and unknown but systematic sampling errors appear. In contrast, multispectral field sensors
12 normally comprise of individual foreoptics/detector assemblies for each spectral band and
13 subsequently have less complex optical paths than their hyperspectral counterparts, and each sensor
14 can be more reasonably assumed to have a center-weighted and Cauchy response, though this
15 response is also affected by the viewing angle of the instrument (Eklundh et al., 2011). For a more
16 detailed discussion of the FOV and DRF of field spectrometers and multispectral sensors we refer
17 readers to Mac Arthur et al. (2012) and Eklundh et al. (2011), respectively.

18 **3.2.4 Temporal stability of measurements and calibrations**

19 The temporal stability of the measurements and the calibrations are essential factors to be
20 considered when conducting long-term *in situ* spectral measurements. Factors such as diurnal or
21 seasonal fluctuations in temperature, gradual particle deposition onto optical parts (e.g cosine
22 diffusers or reference panels), or any other processes causing a temporal drift in the functioning of
23 will interact with the measured signals and calibrations. In turn, the impact of these factors will
24 depend on the signals we are measuring and the instrumentation we use. For example, because the
25 impact of these factors may be wavelength-dependent it may interfere with the estimation of
26 reflectance indices. Similarly, in DFOV systems constructed around two sensors, the differential
27 impact of these factors in each sensor may also introduce significant errors. Unfortunately, the
28 quantitative characterization of these sources of variability, and the establishment of a set of
29 recommendations, remains a key question after EUROSPEC and clearly requires further attention.
30 We briefly introduce the topic and present some indicative data that we hope will help the reader to
31 understand the importance of temporal stability.

1 Stability issues can be grouped around two points: The temporal stability of the calibration or cross
2 calibration of a sensor pair; and thermal stability of the measurements.

3 1) Temporal stability of the calibration/cross calibration. Sensor calibration against a source of
4 known spectral and radiometric properties is needed to derive radiometric units and control for
5 spectral shifts in sensor response. Similarly, cross calibration of two sensors (e.g. Gamon et al.
6 2015, Jin and Eklundh 2015) is essential for deriving reflectance factors using two different sensors
7 (e.g. DFOV systems) and to control for between-sensor variability. Importantly, particle deposition,
8 component aging, or partial damage of sensor components such as optical fibers, may cause a
9 change in these calibrations which we need to detect, quantify and correct for. For example, the
10 temporal degradation of the white reference panel becomes a critical issue in systems such as the
11 ASD-WhiteRef (See Section 3.3), which thanks to the system design was found to be insignificant
12 (maximum of 2% differences at 400nm) over the measuring period (Sakowska et al. 2015). In the
13 absence of additional information, the general recommendation is to start with an intensive
14 calibration/cross-calibration scheme and adjust the frequency later on when the stability of the
15 calibrations for the specific field conditions is known. Key questions that the user should consider
16 are: what is the temporal drift in calibration for the specific sensors and measuring conditions? what
17 is the impact of this drift on the resulting signal/indices? and what is the optimal calibration/cross
18 calibration frequency?

19 2) Thermal stability. Changes in temperature may have an impact on both the intensity and the
20 spectral information of the measured signal. Accordingly, characterizing the temperature stability of
21 a spectral system and its impact on the signal we seek to measure, is a critical step when designing
22 and deploying *in situ* spectral measurements. For example, the radiometric response to temperature
23 in silicon diodes is more pronounced in the NIR compared to the visible. Saber et al. (2011)
24 characterized the percentage change in the response a spectrometer relative to its response at the
25 calibration temperature (20 °C) and found a variation of -0.13 % / °C at 30 °C that was constant
26 between 400 and 700 nm but increased to +0.2 % / °C at 30 °C at 1050 nm. Similarly, Pacheco-
27 Labrador et al. (2014) characterized the effect of temperature on the signal and the spectral
28 calibration of two sensors in a DFOV system based on a pair of Unispec DC spectroradiometers, for
29 the range of temperatures between 13.9 °C and 46.1 °C, finding higher variation in the NIR,
30 between -10 % to 21% (relative to 30 °C) compared to variations between -1 % to + 4 % below 750
31 nm. Clearly, the spectral component of this temperature dependencies do not only affect the
32 quantification of radiometric quantities, but also the comparison of the quantities measured at
33 different wavelengths. This is especially critical when estimating VIs or solar-induced chlorophyll

1 fluorescence where two or more spectral bands are combined. In these cases, seasonal changes in
2 temperature could, if not properly addressed and corrected, mask the physiological component of
3 the signal or even generate spurious dynamics (e.g. Pacheco-Labrador et al. 2014). Key questions
4 that the user should consider are: How do changes in temperature affect sensor/spectrometer
5 output? what is the impact on the resulting signal/indices? What are the wavelength/sensor
6 dependent thermal responses? If needed, how to control and correct for thermal stability for the
7 selected application?

8 **3.3. Development of instrumentation for continuous field measurements**

9 Conducting *in situ* long-term spectral measurements in the field is not straightforward. In addition
10 to a number of logistic and infrastructural requirements, long-term field measurements require
11 instrumentation specially designed and conceived for the task. One of the goals of EUROSPEC was
12 to identify the main requirements of such sensors and to promote the development of new dedicated
13 instrumentation. As part of these activities we organized a Science-Industry Interaction Meeting
14 where EUROSPEC scientists got together with representatives of the “spectrometry” industry
15 sector. A number of general requirements for field optical sensors were identified and are
16 summarized in Box 1. In addition, industry representatives raised the issue of how to cover the non-
17 recurring engineering costs associated with instrument development. The possibility of establishing
18 partnerships and seeking funding for joint collaborative projects between science and industry was
19 suggested, particularly to produce prototypes for new instruments.

20 Four different hyperspectral systems were identified during EUROSPEC for continuous proximal
21 sensing from EC towers in Europe (Fig. 4): 1) A temperature-controlled spectrometer system for
22 continuous and unattended measurements of canopy spectral radiance and reflectance (UNIEDI
23 System) developed by the University of Edinburgh (Drolet et al., 2014) that has been operating at
24 the FluxNet Hyytiälä site (<http://fluxnet.ornl.gov/site/447>) in Southern Finland since March 2010.
25 2)The Multiplexer Radiometer Irradiometer (MRI) developed by the Remote Sensing of
26 Environmental Dynamics Laboratory, Dipartimento di Scienze dell’Ambiente e del Territorio e di
27 Scienze della Terra, Università degli Studi Milano- Bicocca (Italy) and deployed for relatively short
28 periods (weeks to months) in the context of different projects (Bresciani et al., 2013; Cogliati et al.,
29 2015); 3) the HyperSpectral Irradiometer (HSI) also developed by the previous group which has
30 operated in the field from 2009 to 2011 (Meroni et al., 2011; Rossini et al., 2012, 2014); and 4) the
31 AMSPEC-MED system, a version of the automated, multiangular spectroradiometer system
32 AMSPEC II (Hilker et al., 2010) modified by the Environmental Remote Sensing and Spectroscopy

Box 1. General Requirements for *in situ* long-term optical sensors

- Waterproof (should withstand direct rain)
- Robust design: external parts withstanding tension
- Avoidance of holes and cavities (perfect place for birds and insect nests)
- Attachments for easy field installation
- Minimum payload (threshold depending on application)
- Minimum size (threshold depending on application)
- Low power consumption (threshold depending on site infrastructure)
- Maximize long-term stability of optical parts (i.e. filters, diffusers, etc.) to minimize recalibration frequency
- For global networks: Operating temperature range matching the wide thermal distribution of terrestrial plant species, from -50°C to 50°C. For local measurements: temperature range matching local variation.
- High thermal stability
- Linear sensor response
- High signal-to-noise ratio
- Optimal cosine directional response function
- Logic user interface and easy to program systems
- Remote access to data and system control

1

2 Laboratory (SpecLab), Spanish National Remote Sensing (CSIC) and the Centro de Estudios
3 Ambientales del Mediterráneo (CEAM) in Spain. The system has been operating at Las Majadas
4 Fluxnet site in Spain (<http://fluxnet.ornl.gov/site/440>) since August 2013 (Pacheco-Labrador and
5 Martín, 2015).

6

7 The first three systems are based on commercially available spectrometers from Ocean Optics,
8 relatively low cost and compact optical benches housed in temperature controlled environments and
9 operated by dedicated software. The main difference between each of these systems lies in their
10 design.

11 The UNIEDI system (Fig. 4A) has a hemispherical-conical configuration, and is a DFOV system
12 that uses a pair of spectrometers (Ocean Optics, USB2000+) with one spectrometer measuring up-
13 welling radiance through a FOV limited foreoptic (24.8 deg) and the other spectrometer measuring
14 down-welling irradiance using a cosine corrected diffuser. One limitation of this system is that,
15 despite that both spectrometers are kept at constant temperature using a temperature control system
16 (Drolet et al., 2014), regular intercalibration of the two spectrometers is essential to account for

1 sensor-specific time- and temperature-dependent drift in their radiometric capabilities.
2 Intercalibration can be done using a calibration lamp and a dark setting either *in situ* or in the
3 laboratory. In fact, Anderson et al (2006) showed that for calibrating field-based spectra collected
4 with a DFOV spectrometer, a field-derived intercalibration function provides the most accurate
5 results.

6 The MRI system has also a hemispherical-conical configuration, but it is a SFOV system with a
7 single spectrometer. A commercially available optical multiplexer is used to switch the input to the
8 spectrometer from down-welling to up-welling radiant flux. Irradiance can be measured through a
9 fibre connected to either a cosine corrected diffuser or an up-looking integrating sphere foreoptic.
10 Up-welling radiance is measured through a bare optical fibre with a FOV of 25 deg. The S-FLUOR
11 has been more recently developed in collaboration with the Forschungszentrum Jülich GmbH based
12 on the MRI design (see Fig 4B). The major technical improvements are an overall compact design
13 and the integration of the cooling system within the instrument box (Cogliati et al., 2015). The S-
14 FLUOR has been used in the years from 2012 to 2014 as reference instrument for the HyPlant
15 (Specim, Finland) airborne fluorescence imager (Rossini et al., 2015). The HSI system (Fig. 4C)
16 has a bi-hemispherical configuration and a SFOV and uses of a rotating optic with a cosine response
17 to measure the down-welling and the up-welling radiant flux. Because both the MRI and HSI are
18 SFOV they present a time delay (at best multiple seconds) between the down-welling and the up-
19 welling measurement. These delays may add some measurement uncertainties due to changes in sky
20 conditions between the individual measurements, particularly under cloudy or overcast conditions.
21 The AMSPEC-MED system is based on a commercial Unispec dual channel VIS-NIR
22 spectroradiometer (PP-Systems, Amesbury, MA, USA) equipped with a motor driven pan-tilt unit
23 that allows measuring up-welling radiance in a range of zenithal and azimuthal angles. Similar to
24 the UNIEDI, the system is a DFOV system and, therefore, cross calibration between spectrometers
25 is performed regularly using a Spectralon[®] panel. Because the system is operated with solar panels,
26 temperature control is not possible due to power restrictions. Instead, temperature sensitivity of
27 each of the spectrometers and its impact on the resulting hemispherical-conical reflectance factors
28 was characterized in the laboratory and used in signal post-processing (Pacheco-Labrador and
29 Martín, 2015). Note that power constraints are not system dependent but rather site specific,
30 depending on power availability and site temperature range.

31 In an attempt to address some of the limitation of the systems reviewed above and based on
32 discussions between groups during EUROSPEC, three new approaches were developed (Fig. 5):

1 1) The Piccolo system, developed by the UK Natural Environment Research Council (NERC) Field
2 Spectroscopy Facility (FSF) Geoscience, University of Edinburgh, is based on a DFOV
3 hemispherical-conical configuration with a cosine corrected foreoptic to capture down-welling
4 radiant flux and a configurable up-welling channel to capture up-welling radiant flux. The up-
5 welling foreoptic can either be fitted with a view angle limited foreoptic or with another cosine
6 corrected receptor to enable a bi-hemispherical measurement approach to be adopted (Fig. 1). The
7 novelty of this system is the use of low weight components for decreased weight, and the use of
8 bifurcated fibre optic with electronic shutters for decreased time delay between up and down-
9 welling measurements (Mac Arthur et al., 2014). In addition, as both light inputs can be closed at
10 the same time, the systems' dark current (inherent electrical noise) can be recorded and used in post
11 processing. The Piccolo system is currently undergoing service life cycle testing and will be field
12 trialled in a number of flux towers in the near future. In addition, the low weight and DFOV mode
13 of this system makes it compatible with unmanned aerial vehicles (UAV) applications, opening a
14 new range of research possibilities. 2) A similar configuration has been adopted in the SIF-Sys
15 (Burkart et al., 2015) developed by the Forschungszentrum Jülich GmbH. The system hosts a low
16 cost and small size spectrometer (STS-VIS, Ocean Optics, Inc., Dunedin, US) and uses also a
17 bifurcated optical fibre with optical shutters to split the optical signal between two channels: one
18 channel pointing to a white reference panel to measure the down-welling radiant flux and the down-
19 looking channel measuring the radiant flux up-welling from the vegetation. SIF-Sys is specifically
20 intended to measure SIF and, for this reason, it is equipped with a LED emitting at the wavelength
21 of SIF (at 760 nm). The LED is placed in the instrument down-looking FOV and it is used as a
22 reference to assess the uncertainty of passive SIF retrieval in field conditions. SIF-Sys has been
23 tested in dedicated field experiments and will be installed at flux towers for long term and
24 unattended data collection in the near future. 3) The ASD-White Ref system (Sakowska et al., 2015)
25 is an automated system designed for continuous acquisition of measurements using an ASD
26 FieldSpec spectroradiometer. The WhiteRef system was developed by the Forests and
27 Biogeochemical Cycles Research Group, Sustainable Agro-Ecosystems and Bioresources
28 Department, Research and Innovation Centre—Fondazione Edmund Mach, San Michele all'Adige,
29 together with the Institute of Biometeorology—National Research Council, Firenze in Italy, and the
30 contribution of NERC Field Spectroscopy Facility, School of Geosciences, University of
31 Edinburgh, and has been deployed in a grassland site in the Viote del Monte Bondone in Northern
32 Italy. The main advantage of this system is the possibility to scan in the VNIR and SWIR regions
33 (350 nm to 2,500 nm) using a popular and commercially available spectrometer. The system is
34 SFOV and measures in a hemispherical-conical configuration with a FOV of 25 deg. A novelty of

1 the WhiteRef system is that both, reference and vegetation target radiances, are measured by
2 automatically sliding a white reference panel under the fibre optic. To protect the WR panel from
3 light, dust, rain, insects and adverse weather conditions, the WR is kept inside a waterproof box and
4 ejected only during the measurements. Each acquisition is preceded by a reading of a dedicated
5 wetness sensor signal, and in case of rainfall or dew the reference measurements are not conducted.
6 In addition, to remove eventual dust/insects from the measurement surface, the WR panel is sprayed
7 with compressed air during each ejection and insertion phase.

8 **3.4. Upscaling optical data and fluxes from the footprint to the landscape level**

9 *In situ* spectral measurements are essential for the successful upscaling of optical and flux data
10 across space and time. In particular, the temporal match between *in situ* spectral measurements and
11 flux data facilitates the characterization, modelling and validation of their linkage. Spatial and
12 temporal scales are tightly connected with each other and neither temporal or spatial upscaling can
13 be fully accomplished without giving attention to the other. Considering that the temporal link
14 between optical and flux data can covered with *in situ* spectral measurements the main question is
15 probably that of upscaling these signals across space from the footprint to the landscape level (Fig.
16 6).

17 In the process of integrating remote sensing data with flux measurements an assumption is
18 commonly made: the match between flux footprint and image pixel (e.g. Beer et al. 2010,
19 Tramontana et al 2015). The same assumption can be used between flux data and *in situ* spectral
20 measurements. However, a number of factors related to footprint variability, pixel heterogeneity,
21 the BRDF properties of the surface, and the geometry of the measurements can momentarily or
22 systematically decouple optical and flux data adding noise or bias to their relationship.

23 Despite of efforts to orientate the FOV of *in situ* spectral measurements to cover the dominant
24 footprint of EC measurements (e.g. using footprint modelling techniques), the flux footprint will
25 still differ from that of optical measurements most of the time due to footprint variability. Most flux
26 sites are located in places with homogeneous vegetation where footprint variability is not expected
27 to decouple flux and optical data. For example, accurate modelling of the flux footprint did not
28 improve the predictive power of optical data to estimate GPP in a Mediterranean savanna (Pacheco-
29 Labrador et al. 2015) or in a subalpine grassland (Vescovo et al. 2015). However, the mismatch can
30 be relevant in sites with heterogeneous vegetation like agricultural land, ecotones, or sites with
31 adjacent patches of vegetation. In these sites, characterization of the area of interest and footprint
32 modelling will be critical for the successful implementation of data driven models, e.g. the light use

1 efficiency model introduced in Eqn. 1. For example, when estimating GPP in an agricultural area
2 using MODIS data and a footprint model, Gelybó et al (2013) were able to reduce the RMS error by
3 28% compared to non-footprint weighted values. Dealing with the effect of optical vs flux footprint
4 mismatch is challenging from a point of view of tower based measurements. One of the conclusions
5 from EUROSPEC was that new tools are needed to characterize these scale issues more precisely.
6 One of them is the use of small and relatively affordable UAVs or remotely piloted aircrafts (RPAs)
7 on which light weight spectrometers, both multi and hyperspectral, and cameras can be deployed.
8 Hyperspectral imaging systems onboard of aircraft or unmanned aerial vehicles (UAV) can for
9 example provide high spatial resolution imagery enabling the identification of pure species pixels
10 within the flux footprint (Zarco-Tejada et al., 2013a). The flexibility, maneuverability, and capacity
11 to view the same target from different heights allows to study the impact of footprint variability
12 and to bridge *in situ* spectral measurements with coarser satellite or airborne data, facilitating their
13 interpretation and un-mixing (Fig. 6). For example, the availability of pure pixels can be used to
14 investigate the effect of aggregating different species or land-cover classes on the resulting
15 hyperspectral signal (Zarco-tejada et al., 2013b). As reported in Gamon et al. (2015) or Whitehead
16 K. and Hugenholtz (2014), the cost effectiveness of UAV platforms make them a valid solution to
17 address footprint variability. Two UAV-based statistical sampling approaches are possible to
18 systematically address footprint variability: i) with no previous knowledge a regular grid might be
19 recommended, whereas ii) if the spatial patterns of vegetation are already known, a stratified
20 sampling for different vegetation types might be more efficient. Overall, the systematic optical
21 sampling of the footprint/pixel area can serve to characterize the different sources of error when
22 upscaling from *in situ* spectral measurements to the satellite pixel level. These topics have just
23 started to be addressed as low cost UAVs and proper instrumentation are becoming available. The
24 technology is relatively under-explored in the context of flux scaling studies, but there are a
25 growing number of papers that comment on the utility of UAVs for fine-scale sensing of landscape
26 ecology and vegetation parameters (e.g. Dandois and Ellis, 2013).

27 Data from *in situ* and remote sensing measurements are also affected by the structure of the canopy
28 and the geometry of the observation and illumination *per se* (Jones and Vaughan, 2010). The reason
29 is that photons hitting a surface are preferentially scattered (or reflected) in given directions
30 depending on the properties of the surface. This can be characterized by the BRDF of the surface
31 (Nicodemus et al., 1977; Roberts, 2001). In other words, if we would measure an ideal plant canopy
32 with constant fAPAR and LUE, our sensors would still register diurnal and seasonal variations in
33 vegetation parameters due to variations in solar elevation and azimuth. This is particularly relevant

when comparing seasonal time series of optical data which may have been acquired with significantly different sun elevations (e.g. in boreal latitudes). Accordingly, knowledge on the BRDF properties of the surface under examination becomes essential to correct for these geometry effects.

The BRDF can be quantified and investigated by mounting sensors on pan-tilt heads (e.g. Huber et al., 2014), by deploying a number of (low-cost) sensors with different fixed off-nadir positions, or by using the UAV systems discussed above. Hilker et al. (2007, 2010), presented a hyperspectral system capable of quantifying and measuring these effects, the AMSPEC and AMSPEC II systems (Automated Multiangular Spectro-radiometer for Estimation of Canopy reflectance). The AMSPEC system (see previous chapter) is a DFOV system that samples hemispherical-conical reflectance factors at different observation angles from the canopy surrounding the tower. Multiangular measurements are used to retrieve the BRDF and can be used to normalize observations to the same viewing and illumination conditions. Data acquired by AMSPEC system over forest stands in North America showed how optical indices, such as the PRI, can be influenced by view angle and shadow fractions (Hilker et al., 2008b). Moreover, retrieved BRDF estimates allows mimicking off-nadir observations of remote sensors and provides a top-of-canopy reference for atmospheric corrections (Hilker et al., 2009).

4. Future challenges and opportunities

Quantifying and modelling the spatiotemporal dynamics of the carbon cycle remains a key goal in climate change and global biogeochemistry research. Global questions call for global initiatives to provide sensible data at the global scale. Flux tower networks such as FluxNet (<http://fluxnet.ornl.gov/>) and other long-term monitoring infrastructures such as ICOS (<https://www.icos-ri.eu/>) or NEON (<http://www.neoninc.org/>) are responding to these needs by ensuring a long-term and an increasing flow of global carbon flux and ecological data. Simultaneously, the increasing number of current and planned satellite missions warrants an equally increasing flow of remotely sensed data (e.g. Venus, Sentinel-2 and 3; OCO-2, FLEX), offering improved geographical coverage, as well as temporal, spatial, and spectral resolutions. However, our capacity to capitalize on these space developments depends very much on how well we can relate the resulting satellite data to ground observations of ecosystem processes, such as photosynthetic carbon assimilation.

EC measurements provide good temporal resolution of carbon fluxes at the ecosystem level but they are limited by spatial resolution and coverage. In contrast, remote sensing data provides good to

1 moderate spatial resolution and coverage but are limited by temporal resolution. The
2 complementarity and synergy between these two sources of data is clear but their integration
3 remains a challenge due to scale mismatch. We need a Rosetta stone to help us translate and link the
4 information from these two sources of data: something that can be done only via *in situ* spectral
5 measurements. On one hand, *in situ* spectral measurements can provide the same optical indices
6 than satellites, serving as a landmark to interpret, calibrate and validate remotely sensed data
7 products (i.e. we can establish a link between satellite data and ground optical data). On the other
8 hand, because data from *in situ* spectral measurements has comparable temporal resolution and
9 relatively similar biological footprint to that of EC measurements, they can be used to develop
10 quantitative models that associate the two signals.

11 Overall, there is a clear need to establish a global network of sites with standardized and
12 coordinated *in situ* spectral measurements to facilitate the integration of remotely sensed data and
13 EC data towards improving the global monitoring of the carbon cycle. In addition, such network is
14 also needed to calibrate and validate satellite data products, and to resolve and avoid problems that
15 appear when inferring ecosystem properties directly from satellite data, such as the “spurious
16 amazon green-up” (Morton et al., 2014; Soudani et al., 2014); or the controversy around the remote
17 sensing of foliar-nitrogen (Knyazikhin et al., 2013; Townsend et al., 2013).

18 The EUROSPEC Cost Action was a starting point for the organization of the European community
19 of scientists working with *in situ* spectral measurements. We identified many areas that still need
20 further work and perhaps the main conclusion of EUROSPEC was to realize that we need more
21 projects such as EUROSPEC. As a continuation, a new COST Action (ES1309) “Innovative optical
22 tools for proximal sensing of ecophysiological processes (OPTIMISE)
23 (<http://optimise.dcs.aber.ac.uk/>) was recently begun that expands the work of EUROSPEC to
24 include UAVs, “smart” spectral data storage systems and to go in-depth into the measurement and
25 interpretation of multi-scale chlorophyll fluorescence data.

26 Despite that regional level networking projects such as the EUROSPEC and OPTIMISE COST
27 Actions, AusCover (<http://www.auscover.org.au/>), or EcoSIS
28 (<http://labs.russell.wisc.edu/townsend/tag/ecosis/>) are important, we need also activities and
29 networking at the global level. SpecNet is an excellent platform that could be used to accomplish
30 this coordination goal and liaise with national and regional projects. SpecNet could be also used to
31 share information, know-how, data, general guidelines on measurement and calibration protocols,
32 and challenges between scientists, but also including industry stakeholders. This is perhaps the

1 fastest and most effective way, in terms of costs and results, to promote standardization. As long as
2 the information remains disperse and the global network links remain weak, independent groups
3 will continue to adopt different solutions for *in situ* spectral measurements without following a set
4 of general guidelines. This is perhaps the main risk behind *in situ* spectral measurements in the near
5 future.

6 The following challenges and opportunities were identified during EUROSPEC:

7 1) Need to compile information on best-practices for *in situ* spectral measurements. Information on
8 what to purchase, how to install, maintain, calibrate, analyze, and store the data from *in situ* spectral
9 measurements is to some extent available from a number of studies conducted as part of
10 EUROSPEC or by other groups (see e.g. Anderson et al., 2011; Balzarolo et al., 2011; Gamon et al.,
11 2015; Harris et al., 2014; Jin and Eklundh, 2015). These types of studies will most likely continue
12 to appear in the near-future. However, a major up-to-date synthesis effort is urgently needed to
13 provide a comprehensive treatise on such measurements. This would facilitate the different phases
14 of decision-making by site PIs and promote standardization within relevant networks such as ICOS
15 and FLUXNET.

16 2) Quantifying and dealing with uncertainty. Measurement uncertainty is instrument- and
17 environment-specific. Accordingly, characterization of sensor performance and quantification of
18 measurement uncertainty is crucial to produce accurate data (Anderson et al., 2011; Castro-Esau et
19 al., 2006; Jung et al., 2012). Anderson et al (2011) have demonstrated that laboratory-derived
20 measurement uncertainties do not present a useful means of quantifying all uncertainties in field
21 spectroscopy. Laboratory measurements can serve to define features such as signal-to-noise ratio,
22 noise equivalent radiance and linearity, but these uncertainties are added to by complexities of the
23 hemispherical illumination environment experienced in the field. Clearly, the optimal way to
24 characterize measurement uncertainty is to do so in the conditions that typify the measurement
25 scenario. Protocols for systematic measurement uncertainty characterizations in the field should be
26 adopted in the future.

27 3) Need for characterization and calibration. Networks of research sites engaged in optical sampling
28 should follow an instrument characterisation and calibration scheme to ensure direct result inter-
29 comparison (Anderson et al., 2013; Balzarolo et al., 2011). Optical sensors could for instance be
30 characterized and calibrated against a common standard in a central laboratory prior to field
31 deployment then tested annually to monitor change or degradation. In addition, portable
32 calibration/verification standards could be rotated periodically around sites to conduct validation

1 measurements across space and time. Cross calibration of sensors in DFOV systems is also critical
2 and should be accomplished regularly. Calibration frequency will depend on signal drift rate which
3 may be instrument- and climate-dependent. Accordingly, it would seem logical to characterize and
4 adjust calibration demands to each site and instrument, for example by calibrating at high
5 frequencies during the first measuring season and adjusting later on depending on signal drift rate.

6 4) Need for a ‘smart’ data repository and information access portal. Spectral data are time intensive
7 to collect but their analysis is even more time consuming. In turn, most spectral data collections
8 remain poorly documented which greatly reduces their use for data sharing, if not nullifying it.
9 There is an urgent need for a spectral information system that: (a) establishes a data pool that can
10 hold spectral data collected from various instruments, providing them in an easily accessible and
11 generic form, and (b) includes metadata that is standardised to a degree that allows data selections
12 to answer new science questions. Such databases have been developed (Bojinski et al., 2003;
13 Ferwerda, 2006; Hueni and Tuohy, 2006), but their adoption by the spectroscopy community has
14 been slow. Currently, there are only a few available and persisting spectral information systems, the
15 most prominent one being the open source SPECCHIO (Hueni et al., 2009). SPECCHIO has seen
16 many upgrades over time with a large contribution by the Australian National Data Service (Hueni
17 et al., 2012) and support from EUROSPEC. The challenges for the future are numerous, but most
18 pressing appears the issue of automated data quality and metadata standards. SPECCHIO is
19 currently being further developed under the new COST Action OPTIMISE .

20 5) Upscaling issues: BRDF and footprint analysis. Scaling up *in situ* spectral measurements to those
21 acquired from airborne and satellite platforms and linking this optical data to that of EC
22 measurements remains an issue (Ju et al., 2005; Simic et al., 2004; Wu and Li, 2009). BRDF
23 effects, footprint variability, and scale mismatch are factors that constrain our capacity to link and
24 upscale remotely sensed and EC data. The rapid advances in UAV technology have opened new
25 opportunities to deal with these challenges. Micro-hyperspectral field spectrometers and imaging
26 sensors can now be mounted on UAVs. These measurements can serve to retrieve the BRDF of
27 challenging Earth surfaces, such as forest canopies, to measure footprint optical variability, or to
28 sample the same target at different heights facilitating the treatment of the scale issue. Deployment
29 of specialized instruments on board UAVs with a view to collecting narrowband multispectral or
30 hyperspectral data will constitute a step-change in scientific understanding of the connection
31 between spectral data and multiple ecosystem processes. Investigating this potential is again one of
32 the goals of the new COST Action OPTIMISE.

6) Permanent platform for communication and information dissemination. The need for and the potential of a permanent channel for cross-talk between research communities as well as between scientists and industry stakeholders was identified. Such a channel could be for example a moderated mailing list (e.g. Fluxnet type) or making use of other social media. This platform would provide the opportunity to share known-how and best-practices between users helping to promote standardization. In addition, it would also promote collaboration between research groups as well as between scientists and industry stakeholders, which in turn might foster the development of new instruments.

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1 **Table 1.** Advantages and disadvantages of different approaches and configurations for *in situ*
2 spectral measurements.

3

	Advantages	Disadvantages
Method of data acquisition		
SFOV	<ul style="list-style-type: none"> • Only one instrument, lower cost • No need for intercalibration 	<ul style="list-style-type: none"> • Non-simultaneous up-welling/down-welling measurements. • Need to automate the system for target and reference measurements. • Either temperature control, or characterization of the temperature sensitivity of the instrument and post-processing, are needed to acquire consistent time series of radiometric measurements
DFOV	<ul style="list-style-type: none"> • Quasi-simultaneous up-welling/down-welling measurements • No need to automate for target and reference measurements. 	<ul style="list-style-type: none"> • Typically two instruments, higher cost (but see Piccolo and SIF-Sys) • Need for regular intercalibration • Either temperature control, or characterization of the temperature sensitivity of each of the instruments and post-processing, are needed both for derivation of reflectance factors and to acquire consistent time series of radiometric measurements More sensitive to component degradation effects (each unit degrading at different rate)
Spectral Resolution		
Multispectral	<ul style="list-style-type: none"> • Low cost • Low weight and power consumption 	<ul style="list-style-type: none"> • Fixed wavelengths/ no post-purchase flexibility • Spectral calibration is difficult
Hyperspectral	<ul style="list-style-type: none"> • High versatility • Possibility to estimate several 	<ul style="list-style-type: none"> • Higher cost • More sensible to spectral shifts and

	<p>indices (including future indices) and resample to specific spectral bands.</p> <ul style="list-style-type: none"> • Increased range of possibilities for the inversion of radiative transfer models 	<p>miscalibration</p> <ul style="list-style-type: none"> • Higher power consumption (either due to automation or temperature control units)
Configuration		
Hemispherical-conical	<ul style="list-style-type: none"> • Easier comparison with airborne or satellite measurements • Small sampling area (allows studying smaller units: shoots, single canopies) 	<ul style="list-style-type: none"> • Small sampling area (depending on FOV), requires taller towers for increased sampling area. • Poor cosine response of cosine receptor and Spectralon in the SWIR • If using Spectralon (SFOV): more expensive and delicate than cosine receptor.
Bi-hemispherical	<ul style="list-style-type: none"> • Wider sampling area 	<ul style="list-style-type: none"> • More prone to BRDF effects of vegetation • Difficult for comparison with airborne or satellite measurements if sun zenith angle is large • Reflectance measurements not comparable with nadir observations • Poor cosine response of the receptors in the SWIR

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1 **Figure Legends**

2 Fig. 1. Main instrument configurations adopted for *in situ* spectral measurements.

3 Fig. 2. Different sampling strategies: discrete vs continuous. Spectral resolution denoted by
4 bandwidth and sampling intervals. The narrower the bandwidth and the shorter the interval between
5 bands the higher will be the resolution at which spectral features can be resolved. Note that sensor
6 responses are here represented as Gaussian just for simplicity reasons but typical sensor responses
7 may present different shapes.

8 Fig. 3. Spectralon® panel reflectance before (dotted line) and after cleaning (solid line). The panel
9 had been ‘lightly’ and carefully used in the field for one season. Data provided by C. MacLellan,
10 NERC/NCEO Field Spectroscopy Facility, GeoScience, U. of Edinburgh.

11 Fig. 4. Hyperspectral systems in use before and during EUROSPEC. The UNIEDI System (A), the
12 S-FLUOR/MRI System (B), the HSI System in which both the up- and down-welling
13 measurements are hemispherical (C) and the AMSPEC-MED System where the spectrometer
14 measuring up-welling radiance can measure at multiple zenithal and azimuthal angles (D). Down-
15 welling hemispherical foreoptics (1), Up-welling conical-FOV foreoptics (2) and the box housing
16 the spectrometers (3) are marked in the respective panels.

17 Fig. 5. Hyperspectral systems developed during EUROSPEC. The Piccolo system (A) and the SIF-
18 Sys (B) use a single spectrometer to measure two fields of view by means of a bifurcated fibre
19 optics. The White-Ref system (C) uses a sliding white reference panel to measure irradiance. See
20 further details in the text.

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22 Fig. 6. Upscaling of optical and flux data across space and time. While long-term *in situ* spectral
23 measurements help us establish a link between optical and flux data across time, new tools like
24 UAVs are still needed to facilitate the spatial upscaling from the footprint to the landscape level.

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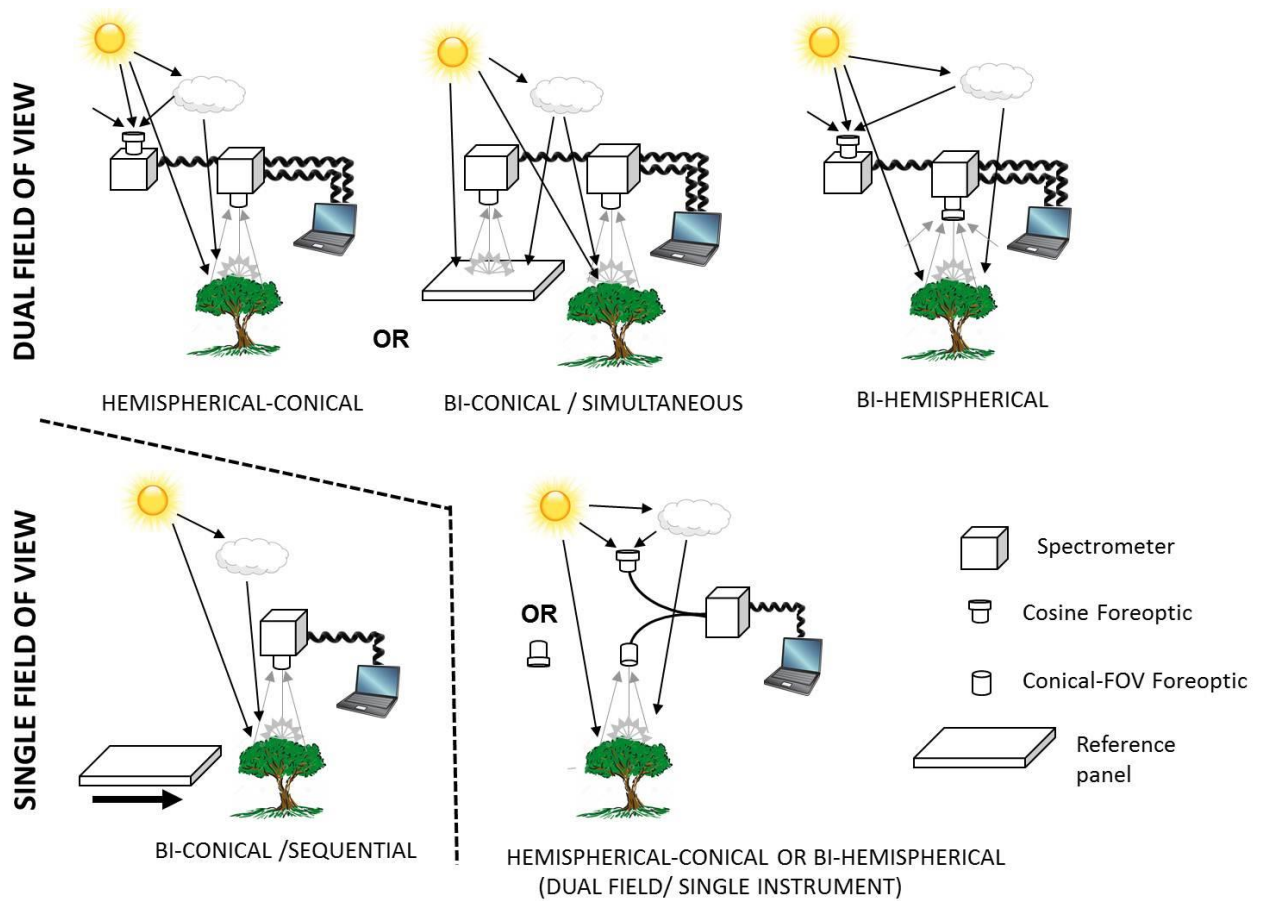


Figure 1.

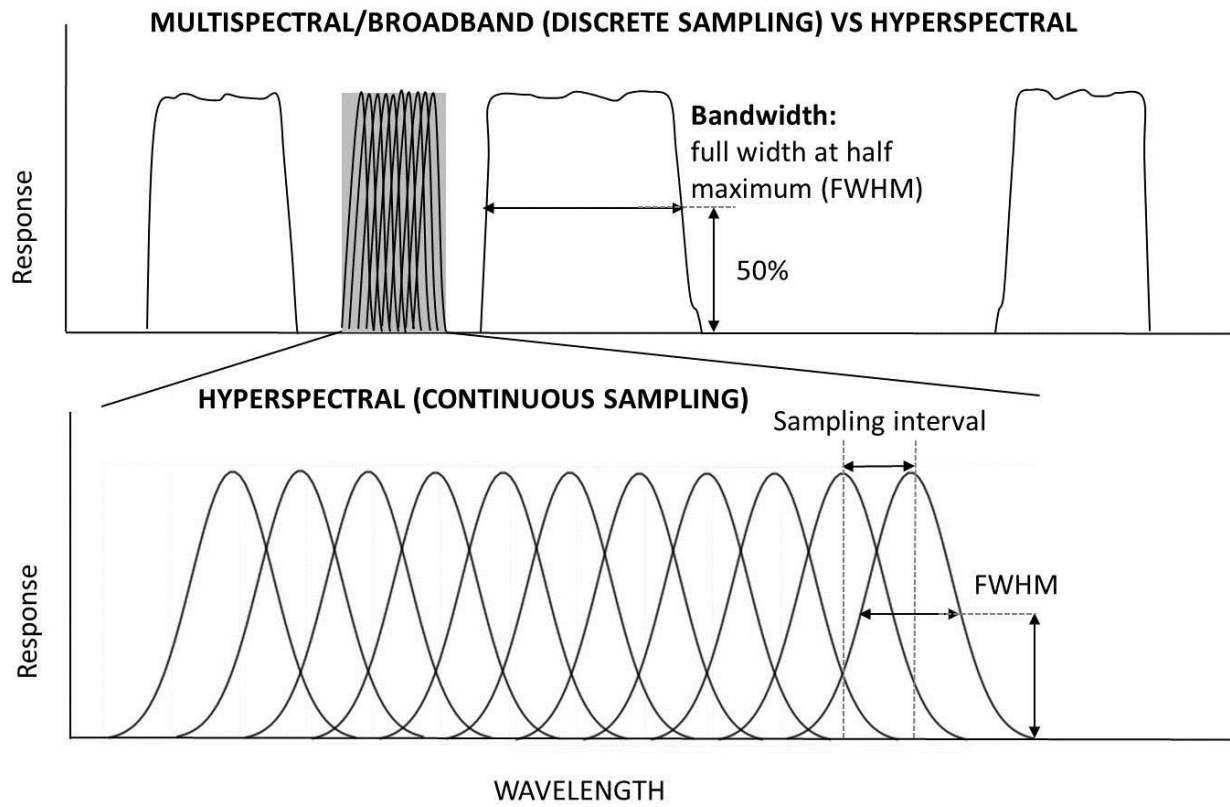
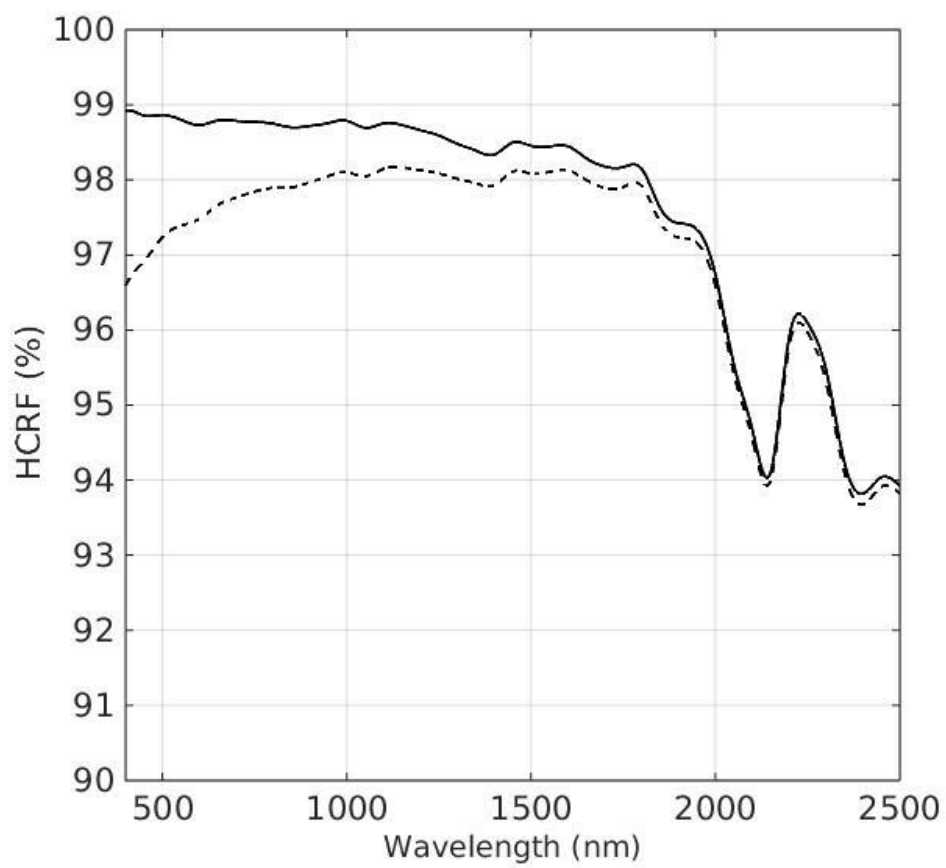


Figure 2.



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2 Figure 3.

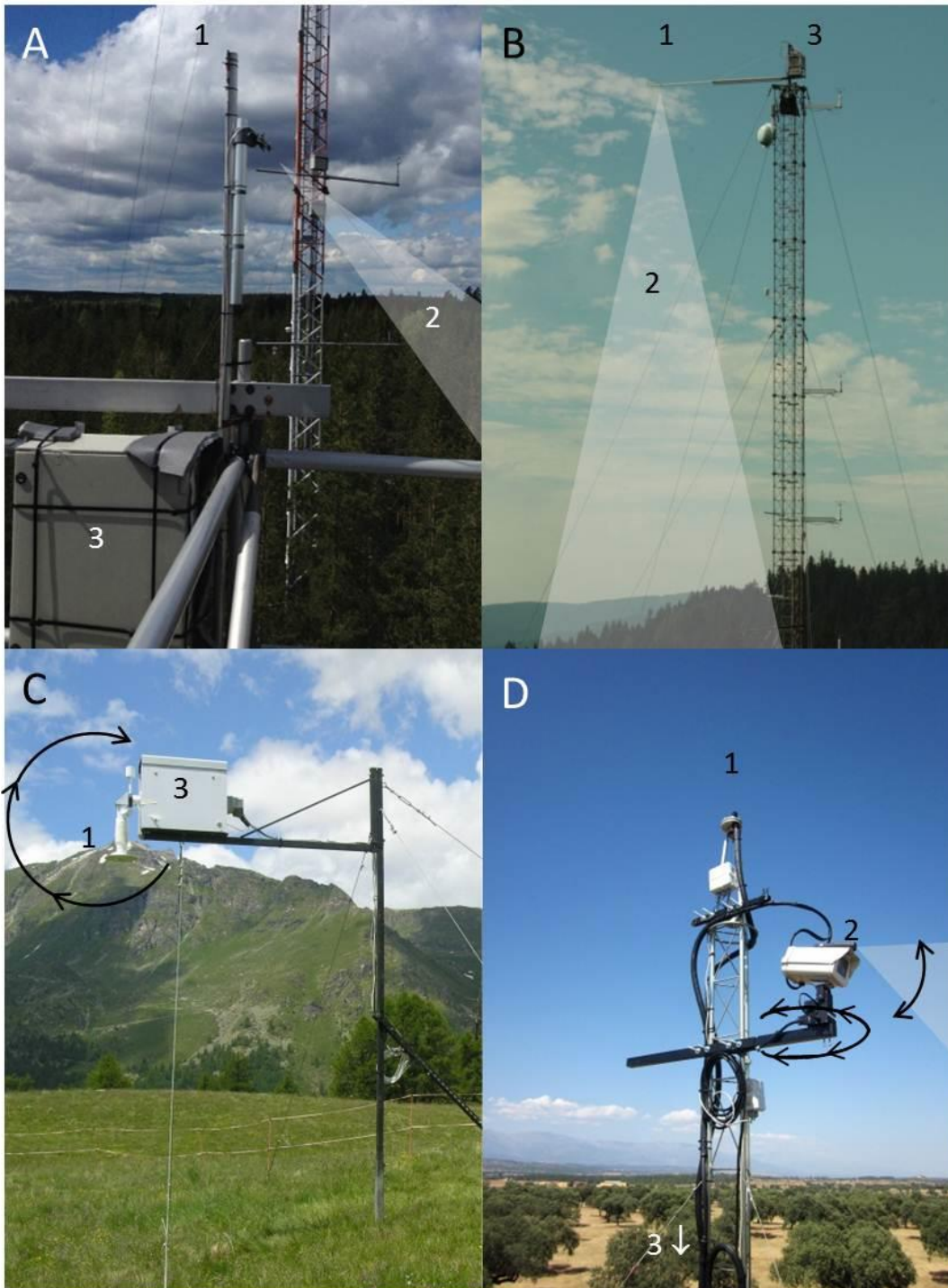
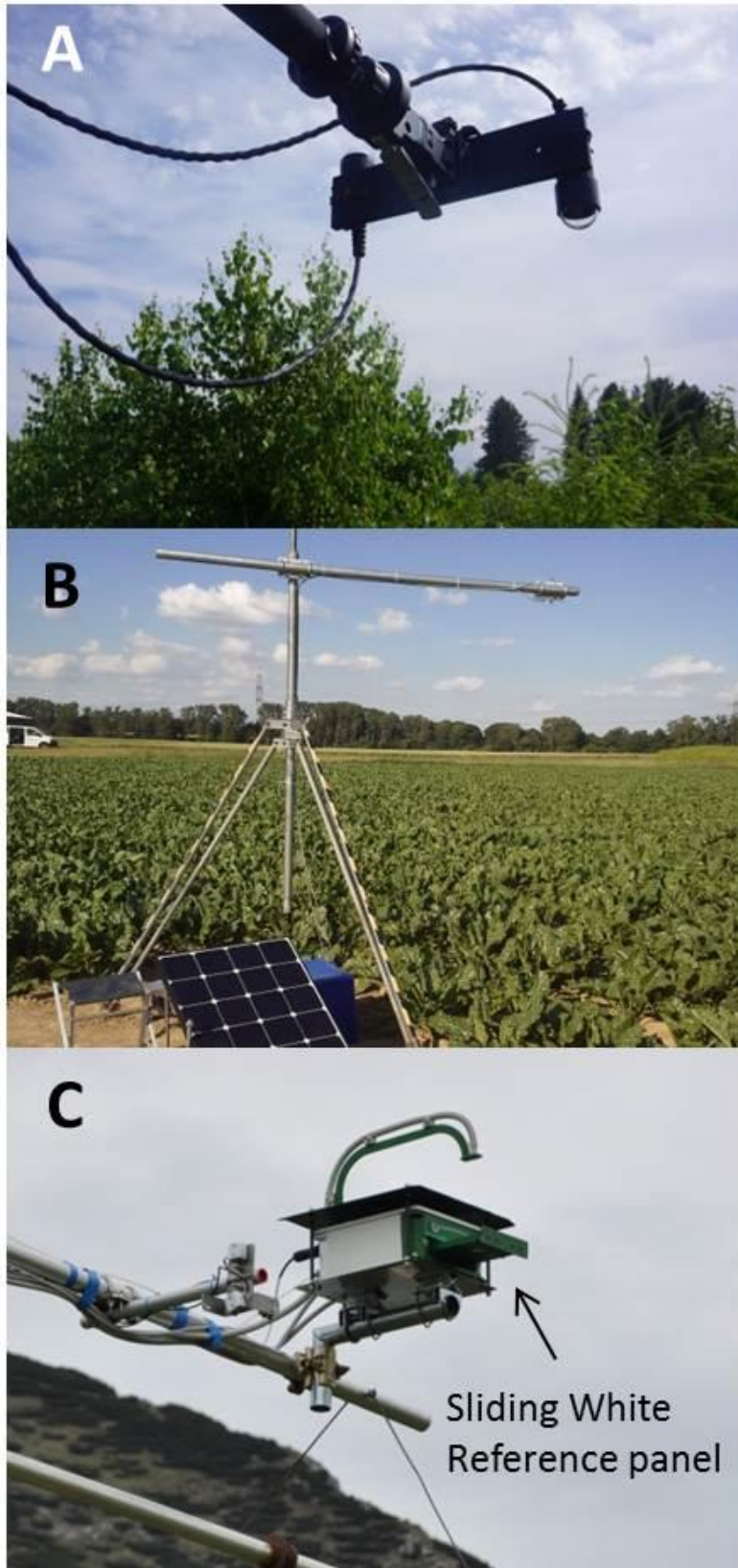
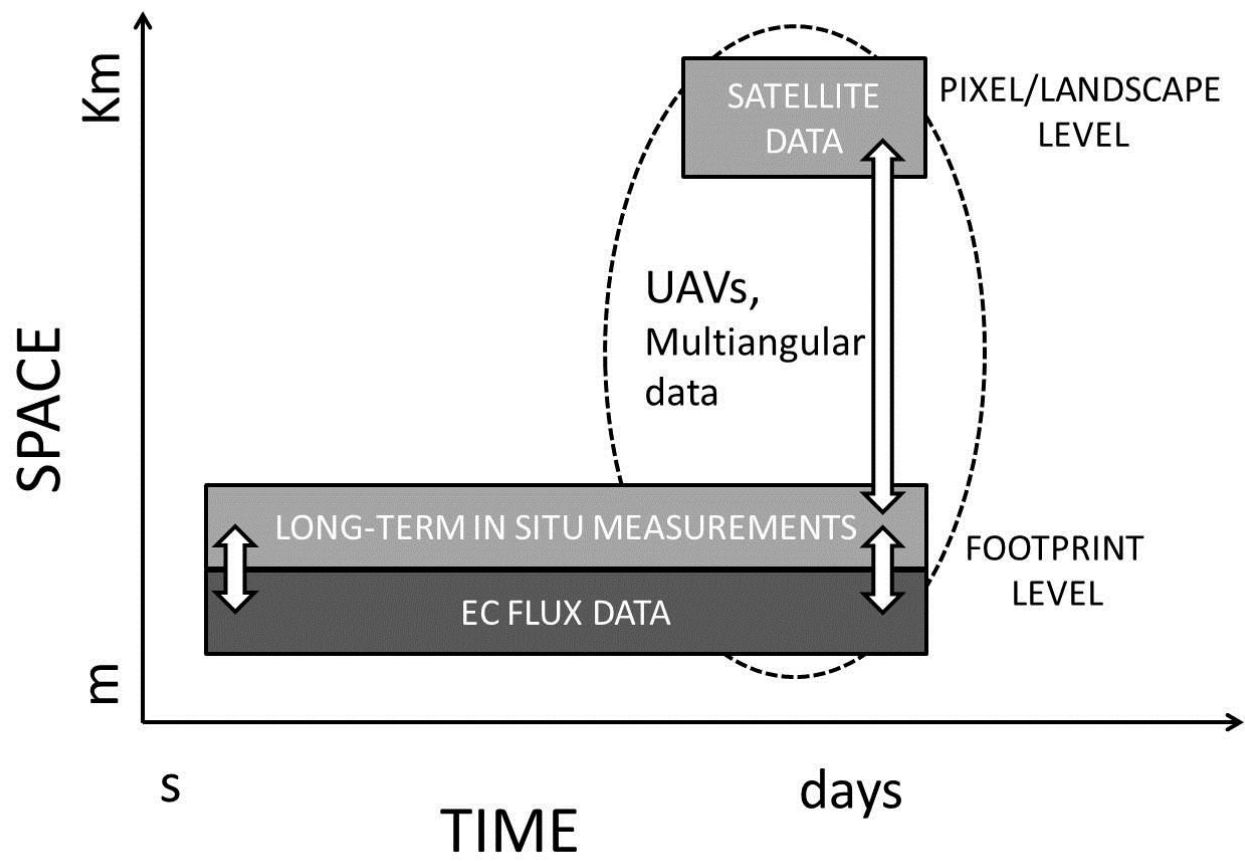


Figure 4.



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2 Figure 5.
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2 Figure 6.
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