1	Southwestern	Tropical Atlantic c	oral growth res	ponse to atmosp	heric circulation changes
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2 induced by ozone depletion in Antarctica

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- 4 Heitor Evangelista^{a*}, Ilana Wainer^b, Abdelfettah Sifeddine^{c,d}, Thierry Corrège^e, Renato C. Cordeiro^c,
- 5 Saulo Lamounier^a, Daniely Godiva^{a,c}, Chuan-Chou Shen^f, Florence Le Cornec^d, Bruno Turcq^d, Claire
- 6 E. Lazareth^d and Ching-Yi Hu^f
- 7
- ^a LARAMG/IBRAG/Uerj. Pav. HLC, Subsolo. Rua São Francisco Xavier 524, Maracanã Rio de
- 9 Janeiro. 20550-013. Brazil. Phone and fax 55 21 2334 0133.

- ^b Universidade de São Paulo USP/IO/Departamento de Oceanografia Física. Praça do Oceanográfico
- 12 191. 05508-120 São Paulo, SP, Brazil.
- 13
- 14 ^c Departamento de Geoquímica, Instituto de Química UFF. Outeiro de São João Batista s/n, Centro –
- 15 Niterói. 24020-007. Brazil.
- 16
- ^d IPSL/LOCEAN, UPMC/CNRS/IRD/MNHN, Centre IRD France Nord, 32 avenue Henri Varagnat,
- 18 93143 Bondy cedex, France.
- 19
- ^e Université de Bordeaux . UMR CNRS 5805 EPOC. Allée Geoffroy St Hilaire, 33615 Pessac cedex,
- 21 France.
- 22
- ¹^f High-Precision Mass Spectrometry and Environment Change Laboratory (HISPEC), Department of
- 24 Geosciences. National Taiwan University. Roosevelt Rd., Taipei 10617, Taiwan ROC
- 25
- 26
- 27
- 28 *: to whom correspondence should be addressed (E-mail: evangelista.uerj@gmail.com)
- 29
- 30 Primary Research Article

31 Abstract

33	Recent Southern Hemisphere (SH) atmospheric circulation, predominantly driven by stratospheric
34	ozone depletion over Antarctica, has caused changes in climate across the extra-tropics. Here we
35	present evidences that the Brazilian coast (Southwestern Atlantic) may have been impacted from both
36	winds and sea surface temperature changes derived from this process. Skeleton analysis of massive
37	coral species living in shallow waters off Brazil are very sensitive to air-sea interactions, and seem to
38	record this impact. Growth rates of Brazilian corals show a trend reversal that fits the ozone depletion
39	evolution, confirming that ozone impacts are far reaching and potentially affect coastal ecosystems in
40	tropical environments.
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46	Keywords: ozone depletion; corals; climate change; South Atlantic; Antarctica; Brazil
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61 **1. Introduction**

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63 Sensitivity of coral species to change in sea surface temperature (SST) has become a focus of 64 attention in the context of modern climate change (Coles and Jokiel, 1978; Howe and Marshall, 2002; 65 Cantin et al., 2010). Response to thermal stress ranges from growth rate alteration to bleaching and 66 eventually death. SST has long been recognized as an important environmental driver for coral growth 67 pattern and calcification (Lough and Barnes, 1997). In a scenario of a warming planet, it is expected 68 that until 2050, ~95% of global reefs will be under pressure from thermal stress and that only 15% will 69 remain in areas of adequate aragonite saturation for proper coral growth (Burke, 2011). Brazilian coral 70 species have high level of endemism and are predominantly found in shallow waters, with the 71 scleractinian coral Siderastrea stellata (Verril, 1868) being one of the main reef builders. It is a colonial, 72 massive, and symbiotic coral (Lins-de-Barros and Pires, 2007) with a spatial distribution that encloses 73 all Brazilian reefs and coral communities from nearly 0°S to 23°S (Castro and Pires, 2001). For 74 Siderastrea stellata, that often grows in shallow water in reef tidal pools (where temperatures normally 75 range between 25°C and 31°C), optimum calcification has been observed at temperatures between 76 28°C and 30°C. For this particular species, the aragonite saturation seems to play a less relevant role 77 than SST for calcification, especially when SST reaches ~26°C (da Silva et al., 2009).

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79 Several factors can induce change in SST that will ultimately impact coral communities. One of them is 80 the change in wind circulation patterns. Several authors have reported a close relationship between 81 wind-stress and SST (Lindzen and Nigam, 1987; Hashizume et al., 2001; Chelton et al., 2007) through 82 modulation of surface heat flux and upper ocean mixing. In general, warmer SST can be associated 83 with deeper boundary layers and higher wind stress (Cayan, 1992; O'Neill et al., 2010). The Southern 84 Hemisphere mid-to-high latitude circulation has undergone marked changes in wind dynamics over the 85 past few decades. One of the most pronounced changes is the poleward displacement of the Southern 86 Hemisphere westerly jet, which has been accompanied by a poleward shift and intensification of 87 oceanic circulation (Thompson et al., 2000; Hartmann et al., 2000). These changes have been 88 attributed in part to external factors such as increased air temperature due to greenhouse gases 89 (GHG), but also to the Southern Annular Mode (SAM) variability enhanced by the Antarctic ozone 90 depletion. The polar stratospheric ozone depletion is also believed to affect the entire Southern

91 Hemisphere, resulting in a broadening of the Hadley cell and a poleward extension of the subtropical 92 dry zone (Polvani et al., 2011). Trends in the Southern Hemisphere tropospheric circulation indicate 93 that the strengthening of the westerly winds is linked to ozone loss (Thompson et al., 2011). Earlier 94 studies (Sexton, 2001; Thompson and Solomon, 2002) postulated that the Antarctic ozone depletion 95 was the primary cause for tropospheric circulation changes in the Southern Hemisphere that occurred 96 in the late 1970s. Another impact of polar ozone depletion on subtropical regions was examined using 97 The Canadian Middle Atmosphere Model (CMAM) and the National Center for Atmospheric Research 98 (NCAR) Community Atmospheric Model (CAM3). Kang et al. (2011) showed that the ozone depletion 99 area evolution has caused a poleward shift of the extratropical circulation, resulting in substantially 100 increase in subtropical precipitation in austral summer. According to Thompson et al. (2011), the 101 influence of the ozone hole on the Southern Annular Mode has led to a range of significant summertime 102 surface climate changes not only over Antarctica and the Southern Ocean, but also over New Zealand, 103 Patagonia, and southern regions of Australia. All in all, ozone depletion appears to have large and far-104 reaching impacts and to potentially be an important player in the Southern Hemisphere climate system. 105 106 One question arising is whether changes in surface winds and SST, triggered in the South Polar region, 107 could affect the Tropical South Atlantic coastal ecosystems. Considering that warming of surface waters 108 may have implications over coral metabolism and health, potentially compromising the sustainability of 109 coral reefs, we here combined measurements of coral growth rate and climate-oceanography modeled 110 data to investigate wind driven impacts in costal environments, especially over the highly sensitive coral 111 communities living in Abrolhos National Park of Brazil (17º25' to 18º09' S; 038º33' to 039º05' W), the

112 most important coral-reef site in the Southwestern Atlantic.

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114 **2. Materials and Methods**

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116 2.1. Coral sampling

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- 118 Massive coral skeleton cores were retrieved within the scope of the Brazil-France projects RECORD
- 119 (REconstructing the Climate from cORal Drilling), LMI PALEOTRACES and the IRD-CNPq
- 120 (CLIMPAST). Corals were drilled using a custom-made pneumatic corer with a ~5 cm diameter barrel.

121 Samples were collected in the National Park of Abrolhos, State of Bahia/Brazil. A Siderastrea stellata, 122 labelled CS1, was sampled at Chapeirão do Sueste (17°57'S; 038°38'W) on March 5th, 2007, water 123 depth was 8.5 m and core length is 53 cm, Figure 2a. The sample labeled UFBA was retrieved from a 124 colony of Favia leptophylla, at Abrolhos (17°52'S; 039°38'W) in the winter of 1977 at approximately 5 m 125 water depth. The core is 28.5 cm long and is hosted at the Department of Geology of the Universidade 126 Federal da Bahia. The chronology of this coral core is described in Evangelista et al. (2007), Figure 2b. 127 A Favia leptophylla coral labelled P1 was sampled in the vicinity of CS1, at Chapeirão do Pierre, on March 4th, 2007; core length is 51 cm, Figure 2c. In the laboratory, CS1 and P1 cores were washed with 128 129 Milli-Q water by sonication and then cut in half with a circular saw. One half was sectioned to produce a 130 1 cm thick slab that was washed again before drying at 40° C.

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132 2.2. Coral skeleton chronology

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134 Sections of coral cores were X-rayed to generate a chronology based on yearly density band counting 135 (Knutson et al., 1972). X-ray images were digitalized and their original contrast was slightly improved in 136 order to enhance the recognition of high and low density growth bands. Band counting took the year 137 (and the season) of sampling as reference for chronology. In order to validate the chronology based on 138 X-ray radiography, we dated the CS1 core with the U/Th technique, as proposed by Shen et al. (2008). 139 For coral skeletons, the U/Th dating method is a robust absolute chronological tool due to the high 140 precision mass spectrometric technology and to the commonly high uranium levels (in the order of ppm) 141 in coral skeletons. Four sub-samples of ~0.1 g were taken from the core with approximately equal 142 spacing from the top to the base. Age corrections were calculated using an estimated radionuclide ²³⁰Th/²³²Th ratio of 4±2 ppm. All radiometric analyses were carried out using a SF-ICP-MS at the High-143 144 precision Mass Spectrometry and Environmental Change Laboratory (HISPEC), Department of 145 Geosciences, National Taiwan University (NTU) (Shen et al., 2008; Shen et al., 2010). Figure 1 show 146 the comparison of both methods.

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Figure 1 - Comparison of chronologies obtained from X-ray radiography (right) and U/Th radiometry
 (left) for a coral core of *Siderastrea stellate* (CS1) from Abrolhos National Park/Brazil.

152 2.3. Instrumental data and model outputs

153 Coral growth data were compared to the extent of the ozone depleted area, the wind stress, the zonal 154 wind, the SST, and the PDO (Pacific Decadal Oscillation). The extent of the ozone depleted area over 155 Antarctica (in million km²) is defined as the maximum daily area in October of each year between 1979 156 and 2008. Data were obtained from NASA Ozone Watch (ozonewatch.gsfc.nasa.gov). Ozone depletion 157 over Antarctica was discovered in the mid 80's (Farman et al., 1985) and was attributed to the positive 158 trend in atmospheric halogenated hydrocarbons released by human activities (Albritton & Kuijpers 159 1999; Solomon 1999). Gradual ozone depletion recover occurred just from the begging of the XXI 160 century (almost 10 years after the implementation of the Montreal Protocol in 1989 that controls 161 emissions of ozone depleting substances). Annual mean wind stress data used here is from ERA-40 162 (Uppala et al., 2005) and zonal wind from NCEP-NCAR reanalysis. SST and sea level were provided by 163 NCEP-NCAR reanalysis and the Simple Ocean Data Assimilation - SODA (Carton & Giese 2008) 164 product. These parameters spanned the full existing database since 1948. PDO data is available at 165 http://www.jisao.washington.edu/aao/. PDO is a robust, recurring pattern of ocean-atmosphere climate 166 variability centered over the mid-latitude Pacific basin, which influences a significant part of the globe,

167 especially South America. It is a long-lived (at decadal scale) El Niño-like pattern of the Pacific climate.

168 This parameter was considered here, since previous works have detected El Niño-like signals on

169 growth rate of tropical Atlantic at Abrolhos, Evangelista et al. (2007).

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171 **3. Results and discussions**

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173 The three coral cores collected in Abrolhos exhibit a decline in annual growth rate towards the 174 present (Figure 2a,2b,2c). Corals used in this study were found in a healthy condition, with no 175 sign of bleaching at the time of sampling. To blend data from the 3 coral cores, we used a 176 simple Gaussian z-score transformation (Figure 2d). Growth anomaly from individual coral cores 177 exhibits a transition from positive to negative approximately after the mid-60's and the 70's, a 178 trend highlighted by the z-score compilation. Figure 2d indicates that the decline in coral growth 179 is guite coincident with SST anomaly for Abrolhos, Figure 2e, from negative to positive above 180 average. Other works have suggested that mean annual SSTs throughout the tropics and 181 subtropics have increased between 0.4° and 1°C in the past four decades (Kleypas et al., 182 2008). For Abrolhos, the mean annual SST changed from ~24.8 °C to ~25.8 °C between 1948 183 and 2006 (NCEP-NCAR reanalysis). Evidence of sharp decline in coral growth rate since the 184 late 1970's was also reported in comparable latitudes away from South America, like that 185 reported for genus Porites from 44 reefs in Australia (Lough et al., 2002). In a subsequent work, 186 these authors have expanded their database to include 5 other reef sites (Lough, 2008). Their 187 average time series depicts a sharp drop in the calcification rate $(q/cm^2/y)$ starting in the late 188 1970's. The transition of both SST and coral growth were concomitant to zonal wind changes 189 around Antarctica as depicted in Figure 2a.





Figure 2 - Time variability and linear trends of annual coral growth anomaly for samples a (CS1), b

193 (Abrolhos/UFBA) and c (P1); (d) z-score for a , b and c ; (e) SST anomaly for Abrolhos National

194 Park/Brazil based on NCEP-NCAR reanalysis.

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196 $\,$ The 70's was a decade when major climatic processes, largely involving the ocean, shifted phases. For

197 example, there was a negative to positive phase transition of the Pacific Decadal Oscillation (PDO),

198 which is associated with more frequent El Niño events (Miller et al., 1994). The atmospheric dynamics

around the Antarctic continent also experienced important changes. The westerly winds increased by

200 the end of the 70's (Polvani et al., 2011). This intensification was coeval with the positive trend of the

201 Antarctic Oscillation (evidenced by the AAO Index) that is the dominant pattern of non-seasonal 202 tropospheric circulation variations south of 20 °S and is an observation-based Southern Hemisphere 203 Annular Mode index used to derive the zonal wind means from 40 °S to 65 °S (Marshall, 2003). The 204 positive phase of AAO is generally associated with stronger cyclones at high southern latitudes (Pezza 205 and Ambrizzi, 2003). The intensification of the westerlies in conjunction with the positive trend of the 206 AAO has been documented in observations, reanalysis, and climate models simulations from the mid-207 1960's to present (Thompson and Solomon, 2002; Gillett and Thompson, 2003; Baldwin et al., 2007). 208 New climatic observations have provided consistent information on the influence of the Antarctic Ozone 209 depletion in the intensification of the westerlies (Albritton and Kuijpers, 1999; Solomon, 1999). 210 Atmospheric temperature shifts concomitant with ozone depletion have been measured at several 211 coastal Antarctic stations from meteorological balloon profiles (Randel and Wu, 1999). One 212 consequence of this lowering in temperature is an increase of the cold-core low pressure area over 213 Antarctica, the Antarctic circumpolar vortex - ACV. The intensification of the ACV may contribute to 214 strengthening the polar-to-subtropics pressure and the air temperature gradient between Central 215 Antarctica and the surroundings affecting the westerlies. Known impacts are the poleward shift of the 216 mid-latitudes jet and of the Hadley cell that ultimately will affect tropical SST. Figure 3 shows our 217 calculations (based on NCEP-NCAR) of main parameters related to ozone-induced climatology and 218 oceanography.





Figure 3 - Calculations based on NCEP-NCAR reanalysis for : (a) the zonal winds around Antarctica (before and after 1979), previously published in 2013 (Cataldo *et al.* 2013); (b) sea level pressure difference (also before and after 1979, but with emphasis to the Southern Atlantic); (c) correlation of wind stress and the ozone depletion area (arrow lengths correspond to r-Pearson values); (d) ozone depletion area; and (e) the zonal wind anomalies calculated to Abrolhos site.

226 Southern tropospheric wind changes associated with ozone depletion have been largely 227 discussed in the literature (Gillett and Thompson, 2003; Shindell and Schmidt, 2004; Baldwin et 228 al., 2007; Son et al., 2008, 2009; Purich and Son, 2012). What is new here is the way winds and 229 Tropical Atlantic SSTs are impacted by ozone depletion. Based on the NCEP-NCAR reanalysis, 230 the tropical Brazilian coast presumably experienced since 1979 increased zonal winds (from 231 East to West) as depicted in Figure 3e. Taking into account the correlations observed between 232 ozone depletion area, Figure 3d, and wind stress (represented here by the length of the arrows 233 at Figure 3c), it points to an impact of the winds over the latitudinal band that encloses the 234 tropical South Atlantic sector, from Africa to the Brazilian coast, reaching the Abrolhos site. A 235 potential impact of that is the piling up of warm waters against the Brazilian continental shelf

236 that would result in increasing SST. The U.K. Meteorological Office Hadley Centre SST 237 climatology (the best spatial and temporal resolved database which dates back to 1870, Rayner 238 et al., 2003) shows a clear moderate stepwise positive change of SST since 1980 on the 239 Brazilian continental shelf (Belking, 2009). An average SST change of ~1 °C was found for the 240 period 1981-2006 with respect to 1957-1980. Long tide gauge time series (calibrated with 241 satellite altimetry) at Cananéia (25°S) and Ilha Fiscal Stations (22°S), both on the Brazilian 242 tropical coast, exhibit sea level increase of ~100 mm between 1976 and 2005 with respect to 243 1954-1975 and of ~160 mm between 1971 and 2004 with respect to 1963-1970, respectively 244 (Dalazoana et al., 2005). Figure 3b, based on NCEP, also depict a positive anomaly at Abrolhos 245 before and after 1979. These instrumental data combined with the high positive correlation 246 found between ozone depletion area and SST in the Southwestern Atlantic sector, as shown in 247 Figure 4, indicate that an active mechanism links Antarctica and the tropics.

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Figure 4 - (a) Correlation map (SST x PDO) and; (b) PDO (Pacific Decadal Oscillation)

anomaly since 1948.

253 In addition, the PDO is a significant parametrization when considering the 1970's climate shift. Several 254 studies postulate a great contribution of the PDO to this climate shift, particularly in the Northern 255 Hemisphere (e.g.: Mantua et al., 1997). Nevertheless, according to Agosta and Compagnucci (2008), in 256 the regional context of Southern South America and adjacent Southern Atlantic, changes in the basic 257 atmospheric circulation of the late 1970's climate transition cannot be solely attributed to the PDO or 258 associated El Niño-like variability. In Figure 4 we show the correlation map between SST and PDO. The 259 result indeed shows an impact of the PDO in the South Atlantic with significant correlation coefficients 260 mostly occurring at a latitude band from ~35 °S to 65 °S. The fact here is that significant correlation 261 coefficients found in Figure 5 (ozone depletion area and SST) are observed where coral communities 262 live in the Brazilian coastal line, while the higher correlations with PDO is confined to subtropics and 263 central Atlantic. Since r-Pearson values between SST and ozone at the Brazilian coral site varied from 264 +0.6 to + 0.7, the explained variance of SST (~40%) due to the climate-induced effects (changes in 265 atmospheric circulation) of the ozone depletion is much higher than the PDO influence at that site. 266



268 **Figure 5 -** Correlation map between ozone depletion area in Antarctica and Southwestern Atlantic SST.

270 Both processes (depletion of stratospheric ozone in Antarctica and GHG - Greenhouse Gases 271 emissions) account for reduced surface pressure over the high latitudes accompanying increase of 272 surface pressure at mid-latitudes. This is associated with the meridional temperature gradient and the 273 position of mid-latitude upper level jet that modulates the tropospheric circulation (winds) and ultimately 274 impact the SST spatial distribution. The near-surface ocean temperature is forced by winds, radiation, 275 and freshwater fluxes. The ocean then impacts the atmosphere via latent, sensible, and radiative heat 276 losses that are dependent on SST and fundamentally on the wind-stress. Since SST is closely related 277 to the mixing layer variability, SST variations are intimately connected with the heat budget of the mixed 278 layer (McPhaden and Hayes, 1991; Chen et al., 1994; Wang and McPhaden, 2000; Foltz et al., 2003). 279 Furthermore, atmospheric circulation anomalies have been shown to precede the development of 280 basin-wide SST patterns for the tropical Atlantic (Nobre and Shukla, 1996). 281 282 4. Conclusions 283 284 Records of coral growth anomaly at Abrolhos site/Brazil evidenced changes (from positive to negative 285 growth rate anomaly) concomitant to ozone area evolution which in turn was highly correlated to SST 286 increases at coral living sites. Herein we point the ozone depletion area as a potential causal 287 parameter. This teleconnected process is of concern to the regional coastal environments, since it 288 would constitute an additional forcing in the SST increase on a scenario of expected global warming 289 due to GHG. Ozone levels in the stratosphere are expected to recover by the end of the century 290 (Perlwitz et al., 2008), and that should theoretically weaken westerly winds (Arblaster and Meehl, 2006;

291 Turner et al., 2009) and contribute to a trend reversal in zonal wind and SST anomalies.

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Author contributions:

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A. All co-authors provided significant contributions to the discussions and the final version of the text;

B. Prof. Heitor Evangelista, Dr. Abdelfettah Sifeddine and Dr. Bruno Turcq were the leaders of the

297 RECORD project that generated the coral database;

298 C. Dr. Ilana Wainer provided the modeling in the manuscript;

299	D. Dr. Thierry Corrège, Dr. Daniely Godiva and Dr. Renato Campello Cordeiro made the coral drillings
300	at the Brazilian coast. Dr. Thierry Corrège made the final revision of the text and significatively
301	improved that;
302	E. Dr. Florence Le Cornec and Dr. Claire E. Lazareth were responsible by Sr/Ca and U/Ca
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329	References
330	
331	Agosta, E.A and Compagnucci, R. H.: The 1976/77 Austral summer climate transition e_ects
332	on the atmospheric circulation and climate in southern South America, J. Climate, 21, 4365-
333	4383, 2008.
334	
335	Albritton, D. and Kuijpers, L.: Synthesis of the Reports of the Scientific, Environmental E_ects,
336	and Technology and Economic Assessment Panels of the Montreal Protocol: a Decade of
337	Assessments For Decision Makers Regarding the Protection of the Ozone Layer: 1988-
338	1999, UNEP/Earthprint, Nairobi, Kenya, 1999.
339	
340	Arblaster, J. and Meehl, G.: Contributions of external forcings to southern annular mode trends,
341	J. Climate, 19, 2896–2905, 2006.
342	
343	Baldwin, M., and Dameris, M.: Climate-ozone connections, Chapter 5, Scientific Assessment
344	of Ozone Depletion: 2006, World Meteorological Organization Global Ozone Research and
345	Monitoring Project, Report No. 50, Geneva, WHO, 5.1-5.49, 2007.
346	
347	Belkin, I. M.: Rapid warming of large marine ecosystems, Prog. Oceanogr., 81, 207–213, 2009.
348	
349	Burke, L.: Reefs at Risk Revisited, World Resources Institute, Washington DC, 2011.
350	
351	Cantin, N., Cohen, A., Karnauskas, K., Tarrant, A., and McCorkle, D.: Ocean warming slows
352	coral growth in the central red sea, Science, 329, 322-325, doi:10.1126/science.1190182,
353	2010.
354	
355	Castro, C. B. and Pires, D. O.: A bleaching event on Brazilian coral reef, Revista Brasileira de
356	Oceanografia, 47, 87–90, 1999.
357	
358	Castro, C, B. and Pires, D. O.: Brazilian coral reefs: what we already know and what is still

359 missing, B. Mar. Sci., 69, 357–371, 2001.

- 360
- 361 Cataldo, M., Evangelista, H., Simões, J.C., Godoi, R.H.M., Simmonds, I., Hollanda, M.H.,
- 362 Wainer, I., Aguino, E.A., Van Grieken, R. : Mineral dust variability in central West Antarctica
- 363 associated with ozone depletion. Atmospheric Chemistry and Physics, 13, 2165 2175, 2013.
- 364
- 365 Cayan, D.: Latent and sensible heat flux anomalies over the northern oceans: driving the sea
- 366 surface temperature, J. Phys. Oceanogr., 22, 859–881, 1992.
- 367
- 368 Chelton, D., Schlax, M., Samelson, R., and de Szoeke, R.: Global observations of large oceanic

369 eddies, Geophys. Res. Lett., 34, L15606, doi:10.1029/2007GL030812, 2007.

- 370
- 371 Chen, D., Busalacchi, A., and Rothstein, L.: The roles of vertical mixing, solar radiation, and
- 372 wind stress in a model simulation of the sea surface temperature seasonal cycle in the trop-
- 373 ical Pacific Ocean, J. Geophys. Res, 99, 20345–20359, 1994.
- 374
- 375 Coles, S. and Jokiel, P.: Synergistic e_ects of temperature, salinity and light on the hermatypic
- 376 coral montipora verrucosa, Mar. Biol., 49, 187–195, 1978.
- 377
- 378 Corrège, T.: Sea surface temperature and salinity reconstruction from coral geochemical
- 379 tracers, Palaeogeogr. Palaeocl., 232, 408–428, 2006.
- 380
- 381 Dalazoana, R., Luz, R. T., and Freitas, S. R. C.: Estudos do NMM a partir de séries temporais
- 382 maregráficas e de altimetria por satélite visando a integração da rede vertical brasileira ao
- 383 SIRGAS, Revista Brasileira de Cartografia, 57, 140–153, 2005.
- 384
- 385 da Silva, L., Júnior, G., and Amaral, F.: Estudo do branqueamento do coral (Siderastrea stellata
- 386 verrill, 1868) da praia de Porto de Galinhas-PE, in: Jornada de Pesquisa, Ensino e Extensão,
- 387 Annals of JEPEX, Recife, 2009.
- 388

- 389 Evangelista, H., Godiva, D., Sifeddine, A., Leão, Z.M.A.N., Rigozo, N.R., Segal, B., Ambrizzi, T.,
- 390 Kampel, M., Kikuchi, R., and Le Cornec, F.: Evidences linking ENSO and coral growth in the
- 391 Southwestern-South Atlantic. Climate Dynamics, 29: 869–880, 2007.
- 392
- 393 Farman, J. C., Gardiner, B. G., and Shanklin, J. D.: Large losses of total ozone in Antarctica
- 394 reveal seasonal CIO*x*=NO*x* interaction, Nature, 315, 207–210, 1985.
- 395
- 396 Foltz, G., Grodsky, S., Carton, J., and McPhaden, M.: Seasonal mixed layer heat budget of the
- 397 tropical Atlantic ocean. J. Geophy. Res, 108, 3146, doi:10.1029/2002JC001584, 2003.
- 398
- 399 Gillett, N. and Thompson, D.: Simulation of recent Southern Hemisphere climate change,
- 400 Science, 302, 273–275, doi:10.1126/science.1087440, 2003.
- 401
- 402 Hartmann, D., Wallace, J., Limpasuvan, V., Thompson, D., and Holton, J.: Can ozone depletion
- 403 and global warming interact to produce rapid climate change?, P. Natl. Acad. Sci. USA, 97,
- 404 1412–1417, doi:10.1073/pnas.97.4.1412, 2000.
- 405
- 406 Hashizume, H., Takeuchi, K., Xie, S., and Liu, W.: Local and remote atmospheric response to
- 407 tropical instability waves a global view from space, J. Geophys. Res., 106, 173–185, 2001.
- 408
- 409 Howe, S. and Marshall, A.: Temperature e_ects on calcification rate and skeletal deposition in
- 410 the temperate coral, *Plesiastrea versipora* (Lamarck), J. Exp. Mar. Biol. Ecol., 275, 63–81,
- 411 2002.
- 412
- Kang, S. M., Polvani, L. M., Fyfe, J. C., and Sigmond, M.: Impact of polar ozone depletion on
 subtropical precipitation, Science, 332, 951–954, 2011.
- 415
- 416 Kleypas, J. A., Danabasoglu, G., and Lough, J. M.: Potential role of the ocean thermostat in
- 417 determining regional di_erences in coral reef bleaching events, Geophys. Res. Lett., 35,
- 418 L03613, doi:10.1029/2007GL032257, 2008.

420	Knutson, D. W., Buddemeier, R. W., and Smith, S. V.: Coral chronologies: seasonal growth
421	bands in reef corals, Science, 177, 270–272, 1972.
422	
423	Le Cornec, F. and Corrège, T.: Determination of uranium to calcium and strontium to calcium
424	ratios in corals by Inductively Coupled Plasma Mass Spectrometry, J. Anal. Atom. Spectrom.,
425	12, 969–973, 1997.
426	
427	Lindzen, R. and Nigam, S.: On the role of sea surface temperature gradients in forcing low-level
428	winds and convergence in the tropics, J. Atmos. Sci., 44, 2418-2436, 1987.
429	
430	Lins-de-Barros, M. and Pires, D. O.: Comparison of the reproductive status of the scleractinian
431	coral Siderastrea stellata throughout a gradient of 20_ of latitude, Braz. J. Oceanogr., 55,
432	67–69, 2007.
433	
434	Lough, J.: Shifting climate zones for australia's tropical marine ecosystems, Geophys. Res.
435	Lett., 35, L14708, doi:10.1029/2008GL034634, 2008.
436	
437	Lough, J. and Barnes, D.: Several centuries of variation in skeletal extension, density and
438	calcificationin massive porites colonies from the Great Barrier Reef: a proxy for seawater
439	temperatureand a background of variability against which to identify unnatural change, J. Exp.
440	Mar. Biol. Ecol., 211, 29–67, 1997.
441	
442	Lough, J., Barnes, D., and McAllister, F.: Luminescent lines in corals from the Great Barrier
443	Reef provide spatial and temporal records of reefs a_ected by land runo_, Coral Reefs, 21,
444	333–343, 2002.
445	
446	Mantua, N. J., Zhang, Y., Wallace, J. M., and Francis, R. C.: A Pacific interdecadal climate
447	oscillation with impacts on salmon production, B. Am. Meteorol. Soc., 78, 1069–1079, 1997.

448	Marshall, G	a.: Trends in	the southern	annular mod	de from	observations	and reanaly	ysis, J.
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449 Climate, 16, 4134–4143, 2003.

450

- 451 McPhaden, M. and Hayes, S.: On the variability of winds, sea surface temperature, and surface
- 452 layer heat content in the western Equatorial Pacific, J. Geophys. Res., 96, 3331–3342, 1991.
- 453 Miller, A. J., Cayan, D. R., Barnett, T. P., Graham, N. E., and Oberhuber, J. M.: The 1976–77
- 454 climate shift of the Pacific Ocean, Oceanography, 7, 21–26, 1994.
- 455
- 456 Nobre, P. and Shukla, J.: Variations of sea surface temperature, wind stress, and rainfall over
- 457 the tropical Atlantic and South America, J. Climate, 9, 2464–2479, 1996.
- 458
- 459 O'Neill, L., Chelton, D., and Esbensen, S.: The e_ects of sst-induced surface wind speed and
- 460 direction gradients on midlatitude surface vorticity and divergence, J. Climate, 23, 255–281,
- 461 2010.
- 462
- 463 Perlwitz, J., Pawson, S., Fogt, R., Nielsen, J., and Ne_, W.: Impact of stratospheric ozone hole
- 464 recovery on antarctic climate, Geophys. Res. Lett., 35, L08714, doi:10.1029/2008GL033317,
- 465 2008.
- 466
- 467 Pezza, A. B. and Ambrizzi, T.: Variability of Southern Hemisphere cyclone and anticyclone
 468 behavior: further analysis, J. Climate, 16, 1075–1083, 2003.
- 469
- 470 Polvani, L., Waugh, D., Correa, G., and Son, S.: Stratospheric ozone depletion: the main driver
- 471 of 20th century atmospheric circulation changes in the Southern Hemisphere, J. Climate, 24,
- 472 795–812, 2011.

473

474 Purich, A. and Son, S.: Impact of Antarctic ozone depletion and recovery on Southern

- 475 Hemisphere precipitation, evaporation and extreme changes, J. Climate, 25, 3145–3154, 2012.
- 476
- 477 Randel, W. J. and Wu, F.: Cooling of the Arctic and Antarctic Polar Stratospheres due to ozone

478 depletion, J. Climate, 12, 1467–1479, 1999.

- 479
- 480 Shen, C.-C., Li, K.-S., Sieh, K., Natawidjaja, D., Cheng, H., Wang, X., Edwards, R. L., Lam,
- 481 D. D., Hsieh, Y.-T., Fan, T.-Y., Meltzner, A. J., Taylor, F. W., Quinn, T. M., Chiang, H.-W., and
- 482 Kilbourne, K. H.: Variation of initial 230Th=232Th and limits of high precision U-Th dating of
- 483 shallow-water corals, Geochim. Cosmochim. Ac., 72, 4201–4223, 2008.
- 484
- 485 Shen, C.-C., Kano, A., Hori, M., Lin, K., Chiu, T.-C., and Burr, G. S.: East Asian monsoon
- 486 evolution and reconciliation of climate records from Japan and Greenland during the last
- 487 deglaciation, Quaternary Sci. Rev., 29, 3327–3335, 2010.
- 488
- 489 Solomon, S.: Stratospheric ozone depletion: a review of concepts and history, Rev. Geophys.,
- 490 37, 275–316, 1999.
- 491
- 492 Son, S., Polvani, L. M., Waugh, D. W., Akiyoshi, H., Garcia, R., Kinnison, D., Pawson,
- 493 S., Rozanov, E., Shepherd, T. G., and Shibata, K.: The impact of stratospheric
- 494 ozone recovery on the Southern Hemisphere westerly jet, Science, 320, 1486–1489,
- 495 doi:10.1126/science.1155939, 2008.
- 496
- 497 Son, S., Tandon, N., Polvani, L., and Waugh, D.: Ozone hole and Southern Hemisphere climate
- 498 change, Geophys. Res. Lett., 36, L15705, doi:10.1029/2009GL038671, 2009.
- 499
- 500 Son, S. W., Gerber, E. P., Perlwitz, J., Polvani, L. M., Gillett, N. P., Seo, K. H., Eyring, V.,
- 501 Shepherd, T. G., Waugh, D., Akiyoshi, H., Austin, J., Baumgaertner, A., Bekki, S., Braesicke,
- 502 P., Brühl, C., Butchart, N., Chipperfield, M. P., Cugnet, D., Dameris, M., Dhomse, S., Frith, S.,
- 503 Garny, H., Garcia, R., Hardiman, S.C., Jöckel, P., Lamarque, J. F., Mancini, E., Marchand,
- 504 M., Michou, M., Nakamura, T., Morgenstern, O., Pitari, G., Plummer, D. A., Pyle, J., Rozanov,
- 505 E., Scinocca, J. F., Shibata, K., Smale, D., Teyssèdre, H., Tian, W., and Yamashita, Y.: The
- 506 impact of stratospheric ozone on Southern Hemisphere circulation changes: a multimodel
- 507 assessment, J. Geophys. Res., 115, D00M07, doi:10.1029/2010JD014271, 2010.

509	Sexton, D.: The e	_ect of stratospheric oz	one depletion on the	phase of the	Antarctic oscillation,
-----	-------------------	--------------------------	----------------------	--------------	------------------------

- 510 Geophys. Res. Lett., 28, 3697–3700, 2001.
- 511
- 512 Shindell, D. and Schmidt, G.: Southern Hemisphere climate response to ozone changes and
- 513 greenhouse gas increases, Geophys. Res. Lett., 31, L18209, doi:10.1029/2004GL020724,
- 514 2004.
- 515
- 516 Thompson, D. and Solomon, S.: Interpretation of recent Southern Hemisphere climate change,
- 517 Science, 296, 895–899, doi:10.1126/science.1069270, 2002.
- 518
- 519 Thompson, D., Wallace, J., and Hegerl, G.: Annular modes in the extratropical circulation, Part
- 520 II: trends, J. Climate, 13, 1018–1036, 2000.
- 521
- 522 Thompson, D., Solomon, S., Kushner, P. J., England, M. H., Grise, K. M., and Karoly, D. J.:
- 523 Signatures of the Antarctic ozone hole in Southern Hemisphere surface climate change,
- 524 Nat. Geosci., 4, 741–749, doi:10.1038/ngeo1296, 2011.
- 525
- 526 Turner, J., Comiso, J. C., Marshall, G. J., Lachlan-Cope, T. A., Bracegirdle, T., Maksym, T.,
- 527 Meredith, M. P., Wang, Z., and Orr, A.: Non-annular atmospheric circulation change induced
- 528 by stratospheric ozone depletion and its role in the recent increase of Antarctic sea ice extent,
- 529 Geophys. Res. Lett., 36, L08502, doi:10.1029/2009GL037524, 2009.
- 530
- 531 Wang, W. and McPhaden, M.: The surface-layer heat balance in the equatorial Pacific Ocean,
- 532 Part II: interannual variability, J. Phys. Oceanogr., 30, 2989–3008, 2000.
- 533
- 534 Zazulie, N., Rusticucci, M., and Solomon, S.: Changes in climate at high southern latitudes:
- a unique daily record at Orcadas spanning 1903–2008, J. Climate, 23, 189–196, 2010.
- 536
- 537