Primary production in the Chukchi Sea with potential effects of freshwater content

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Abstract

The *in situ* primary production rates and various environmental variables were investigated in the Chukchi Sea during the RUSALCA expedition, which was conducted in 2012, to identify the current status of primary production. A ^{13}C-^{15}N dual tracer technique was used to measure the daily primary production rates, which ranged from 0.02 to 1.61 g C m^{-2} d^{-1} (mean ± SD = 0.42 ± 0.52 g C m^{-2} d^{-1}). The primary production rates showed large regional differences, with the southern region (0.66 ± 0.62 g C m^{-2} d^{-1}) producing approximately five times as much as the northern region (0.14 ± 0.10 g C m^{-2} d^{-1}), which was primarily due to the differences in phytoplankton biomasses induced by regional nutrient conditions. The primary production rates in the Chukchi Sea were averaged using data acquired during the three different RUSALCA expeditions (2004, 2009, and 2012) as 0.33 g C m^{-2} d^{-1} (SD = 0.40 g C m^{-2} d^{-1}), which was significantly lower than previously reported rates. In addition to strong seasonal and interannual variations in primary production, recent decreases in the concentrations of major inorganic nutrients and chlorophyll *a* could be among the reasons for the recent low primary production in the Chukchi Sea because the primary production is mainly affected by nutrient concentration and phytoplankton biomass. The nutrient inventory and primary production appear to be largely influenced by the freshwater content (FWC) variability in the region due to the significant relationships between FWC, nitrate concentrations (r = 0.54, $p < 0.05$) and primary production rates (r = 0.56, $p < 0.05$). Moreover, we found highly significant relationships between the nutrient levels and the primary production rates (r =

0.75, $p < 0.001$). In conclusion, the primary production in the Chukchi Sea is primarily controlled by nutrient availability which is strongly related to the FWC variability. Our results imply that the predicted increase in freshwater accumulation might cause a decrease in primary production by lowering the nutrient inventory in the euphotic zone of the Chukchi Sea.

Keywords:

Phytoplankton, Primary production, Chukchi Sea, Freshwater content, Arctic

1. Introduction

Over recent years, the Arctic Ocean has undergone drastic changes in the extent and thickness of sea ice (Stroeve et al., 2008; Comiso et al., 2008; Kwok et al., 2009; Overland and Wang, 2013). The continuing loss of sea ice may result in changes to various physical and chemical environmental conditions in the Arctic Ocean. For example, the loss in sea ice cover allows more sunlight to enter the surface layer of the Arctic Ocean, which results in a longer growing season for phytoplankton growth (Arrigo et al., 2008; Ardyna et al., 2014). Stroeve et al. (2014) reported that the arctic melt season has lengthened at a rate of 5 days decade^{-1} from 1979 to 2013, due to later autumn freeze-up. In accordance with their findings, Ardyna et al. (2014) documented the development of a second bloom in the Arctic Ocean during the fall, which coincides with the delayed freeze-up and the increased exposure to wind stress.

However, the loss in sea ice can cause an increase in the input of freshwater (McPhee et al., 2009). In fact, the freshwater volumes in the Canada and Makarov Basins increased by of 8500 km^3 in 2008 due to increased sea ice melting and river discharge (McPhee et al., 2009; Rabe et al., 2011). This phenomenon can enhance the stratification in the upper ocean (Yamamoto-Kawai et al., 2009) and consequently reduce vertical mixing, thereby preventing nutrient inputs from deep waters to the euphotic zone. In fact, McLaughlin and Carmack (2010) found a deepening of the nutricline due to the accumulation of surface freshwater in the Canada Basin.

In the Chukchi Sea as inflow shelf, there was an increased volume flux of 50% in 2011 (~ 1.1Sv) relative to 2001 (~ 0.7 Sv), which was accompanied by increases in heat and freshwater fluxes (Woodgate et al., 2012). Though the volume flux may vary both seasonally and annually under the influence of the local wind fields, the recent increases in freshwater fluxes in the region may have important implications for phytoplankton in terms of nutrient availability for their growth (Woodgate et al., 2005a, 2005b, 2006). Thus, it is important to identify how phytoplankton respond to these environmental changes in the region in terms of production and/or community structure. According to Li et al. (2009), the phytoplankton community has changed under the freshening and stratifying condition in the Canada Basin. Notably, the abundance of small phytoplankton (< 2 µm) has increased, whereas the abundance of large phytoplankton (2-20 µm) has decreased. Yun et al. (2014) also found that compared with previous reports, the small phytoplankton were more abundant on the Chukchi Sea shelf, which is dominated by low nutrients and freshening conditions. Therefore, the changes in recent phytoplankton production under the rapidly changing environmental conditions need to be monitored because the changes in phytoplankton production could have important implications for understanding ecosystem changes in the Arctic Ocean.

In order to understand climate and ecosystem change in the Pacific Arctic Ocean which is a region summer sea ice cover was declining dramatically (Crane and Ostrovskiy, 2015), the RUSALCA (Russian-American Long-term Census of the Arctic) expedition, which is a joint US-Russian research program, started from 2004 as multidisciplinary investigations in the Bering and Chukchi Seas. Three RUSALCA expeditions (2004, 2009, and 2012) provided a good opportunity for continuous measurements of the primary production in the entire Chukchi Sea, including the Exclusive Economic Zone of the Russian Federation. The 2004 RUSALCA expedition was conducted from 8 to 24 August, 2004 (Lee et al., 2007). The 2009 RUSALCA expedition was executed from 1 to 30 September, 2009 (Yun et al., 2014). The 2012 RUSALCA expedition was carried out from 27 August to 16 September, 2012. This study is part of the 2012 RUSALCA expedition.

In this study, we addressed the regional characteristics of primary production by examining the main driving factors responsible for the regional variability in the Chukchi Sea based on measurements taken in 2012. In addition, we investigated the recent trends in primary production in the Chukchi Sea based on the results of the three RUSALCA expeditions (2004, 2009, and 2012) in the Chukchi Sea. Finally, we emphasized the potential effects of freshwater accumulation on the primary production in the Chukchi Sea because changing amounts and distributions of freshwater content could lead to changes in the primary production rates.

2. Materials and methods

2.1. Study area and sampling

The RUSALCA expedition in 2012 was conducted onboard the Russian vessel *Professor Khromov* in the Chukchi Sea from 27 August to 16 September. The study area was comprised of several sections between the Bering Strait and the vicinity of Herald Canyon (Fig. 1). To understand the regional characteristics of primary production, the study area was divided into two geographic regions (northern/southern) following Yun et al. (2014). The northern region consisted of stations in the vicinity of Herald Canyon (CEN and HC sections) (Fig. 1). The stations in the Chukchi South and Cape Lisburne (CS and CL sections) were included in the southern region. Most of the bathymetric depths in the entire study area were quite shallow, with a mean of 55 m (SD = ± 11 m). Between the production stations, the depth of euphotic zone from the surface to 1% light depth varied between 20 and 46 m, with a mean of 29 ± 10 m (Table 1).

Oceanographic/biological samples were taken from a total of 54 conductivity-temperature-depth (CTD) stations. The vertical profiles of water temperature and salinity were obtained using a Sea-Bird model SBE911*plus* CTD profiler. Water samples were collected with a stainless-steel rosette sampler that was equipped with 21 10-liter bottles at every CTD station. The data from the previous RUSALCA expeditions (in 2004 and 2009) were included to understand the recent trends in primary production in the Chukchi Sea.

2.2. Physical and chemical variables

The stratification index of the water column ($\Delta\sigma_t$) (in kg m^{-3}) was determined as the difference in $\Delta\sigma_t$ values between the surface and the bottom depth according to Yun et al. (2014). The surface mixed layer

(Z_m) was defined as the depth at which the density (sigma-*t*) gradient was 0.05 kg m^{-3} higher than the surface density, as in Coupel et al. (2015). The depth of the euphotic zone (Z_{eu}) in this study was defined as the depth receiving 1% of the surface PAR value, as in Lee et al. (2007) and Yun et al. (2014), and was obtained from a Biospherical QSP-2300 PAR sensor (Biospherical Instruments Inc.) that was lowered with the CTD/rosette sampler. The nitracline (Z_{nit}) was determined as the depth at which the nitrate gradient was greater than 0.1 µM m^{-1} according to the definition of Coupel et al. (2015).

2.3. Fresh Water Content (FWC)

To assess the surface water freshening, the freshwater content (FWC) was calculated following Carmack et al. (2008):

$$FWC = \int_{z_{lim}}^{0} (1 - S(z)/S_{ref}) dz$$

where S and S_{ref} are the *in situ* and reference salinities, respectively, and Z_{lim} is the depth where S equals S_{ref} (34.8 on the practical salinity scale). We used a reference salinity of 34.8 following Aagaard and Carmack (1989) to computing freshwater since it has been considered as the mean salinity for the Arctic Ocean.

2.4. Nutrient concentration measurements

The discrete water samples used in measuring the nutrient concentrations were obtained from 5 to 9 different depths depending on the water depths. The dissolved inorganic nutrient concentrations (nitrite+nitrate, ammonium, phosphate, and silicate) were analyzed onboard immediately after collection using an automated nutrient analyzer (ALPKEM RFA model 300) following the method of Whitledge et al. (1981).

2.5. Chlorophyll a concentration measurements

The water samples used for measuring the chlorophyll *a* concentration were obtained from 4 to 7 different depths at most stations. The water samples were filtered through Whatman GF/F filters (24 mm), and the filters were then kept frozen until analysis in the laboratory. The filters were subsequently extracted in a 3:2 mixture of 90% acetone and DMSO in a freezer for 24 h, followed by centrifugation (Shoaf and Lium, 1976). The chlorophyll *a* concentrations were measured using a Turner Designs model 10-AU fluorometer, which was calibrated using commercially available preparations of purified chlorophyll *a* (Turner Designs, USA). The methods and calculations used to determine the chlorophyll *a* concentrations followed the procedure of Parsons et al. (1984).

2.6. In situ primary production measurements

The water samples used to measure primary production were collected at six photic depths (100, 50, 30, 12, 5, and 1% penetration of the surface irradiance, PAR). At 11 selected morning stations, the *in situ* primary productions of phytoplankton were measured using a ^{13}C-^{15}N dual tracer technique (Lee and Whitledge, 2005; Lee et al., 2007). This method could be useful for distinguish the relative importance of nitrate and ammonium as nitrogen sources for the cell and population (Dugdale and Goering, 1967). We followed the same analytical procedure of Lee et al. (2007) and Yun et al. (2014) to the measure primary production to consistently compare the primary production levels determined in the three studies. Briefly, heavy isotope-enriched (98-99%) carbon ($NaH^{13}CO_3$), nitrate ($K^{15}NO_3$), and ammonium ($^{15}NH_4Cl$) substrates were inoculated in polycarbonate bottles (1 L) and then incubated on deck in a large polycarbonate incubator cooled with running surface seawater under natural light conditions. After approximately 4 to 5 h of incubation , all samples were filtered using pre-combusted (450°C, 4 h) glass fiber filters (Whatman GF/F; diameter = 25 mm). After HCl fume treatment, the samples were sent to the Alaska Stable Isotope Laboratory of the University of Alaska, Fairbanks, USA. The abundances of ^{13}C and ^{15}N and the total amounts of particulate organic carbon (POC) and nitrogen (PON) were determined using a Thermo Finnigan Delta+XL mass spectrometer. Finally, the carbon and nitrogen production rates were calculated based on Hama et al. (1983) and Dugdale and Goering (1967), respectively.

3. Results

3.1. Physical conditions

The surface temperature (T_{sur}) varied from -2 to 9 °C in the study area in 2012 (Fig. 2a). The higher temperatures were found in the eastern side of the southern Chukchi Sea due to the strong influence of the Alaskan Coastal Water (warmer and less saline). The freezing temperatures were observed in the vicinity of the Herald Canyon and gradually decreased toward the northward. At the surface, the salinity varied between 21 and 33 psu. The surface salinity (S_{sur}) was considerably lower in the southwestern side compared with the northeastern side of the southern Chukchi Sea (Fig. 2b). The stratification index ($\Delta\sigma_t$) in the study area ranged from 0.7 to 9.7 kg m^{-3}, with a mean of 3.8 ± 2.2 kg m^{-3}. The stratification in the southern region was higher than in the northern region (Fig. 2c). The general distribution of the stratification index was similar to that of surface salinity because it tended to be high in areas where surface salinity was low. The surface mixed layer (Z_m) was thinner than 15 m over the entire study area (Fig. 2d). In the study area, the depths of nitracline (Z_{nit}) ranged from 2.5 m to 35 m (Fig. 2e), with a mean nitracline depth of 12.8 ± 7.7 m.

3.2. Nutrient distribution

Since the mean depths of euphotic zone in this study was about 30 m, the distribution of ambient nutrient concentrations integrated from surface to 30 m of the water column is shown in Fig. 3. The inventory of nitrite+nitrate ranged from 21.51 to 355.43 mmol m^{-2}, whereas the ammonium inventory ranged from 15.36 to 109.51 mmol m^{-2} (Figs. 3a and 3b). High nitrite+nitrate inventory that exceeded 200 mmol m^{-2} were observed at the center of the CL section (Fig. 3a). The inventories of these nutrients in the southern region (134.15 ± 98.41 mmol m^{-2} for nitrite+nitrate and 61.22 ± 20.55 mmol m^{-2} for ammonium, respectively) were approximately two times higher than their inventories in the northern region (75.01 ± 52.01 mmol m^{-2} for nitrite+nitrate and 40.49 ± 20.69 mmol m^{-2} for ammonium) (see Table 2). The inventory of phosphate in the study area was fairly uniform, with a mean of 24.03 ± 8.30 mmol m^{-2} (Fig. 3c). The silicate inventory was generally higher in the southern region than in the northern region (Fig. 3d).

3.3. Chlorophyll a concentration

The distribution of the chlorophyll *a* concentration in the upper 30 m (i.e., mean depth of euphotic zone in this study) of the entire study area is shown in Fig. 4. High chlorophyll *a* concentrations of over 80 mg m^{-2} were observed in the western side of the CL section (from st. CL5 to st. CL8), and low chlorophyll *a* concentrations were shown in the western side of the CS section (Fig. 4). The highest concentration (286.4 mg m^{-2}) was obtained at station CL8. Over the entire study area, the mean chlorophyll *a* concentration integrated from the surface to 30 m was 42.7 mg m^{-2} (SD = ± 57.4 mg m^{-2}). The average concentrations were 21.7 mg m^{-2} (SD = ± 19.6 mg m^{-2}) and 54.5 mg m^{-2} (SD = ± 67.7 mg m^{-2}) for the northern and southern regions, respectively.

3.4. Primary production rates

Overall, the hourly carbon production rates integrated over the euphotic zone from six light depths ranged from 1.1 to 108.6 mg C m^{-2} h^{-1}, with a mean of 27.7 mg C m^{-2} h^{-1} (SD = 34.7 mg C m^{-2} h^{-1}). The highest primary production rates were found at station CL8 (108.6 mg C m^{-2} h^{-1}) followed by station CL5A (82.1 mg C m^{-2} h^{-1}) (Fig. 5). In the northern region, the carbon production rates ranged from 1.1 to 18.7 mg C m^{-2} h^{-1}, with a mean of 9.0 mg C m^{-2} h^{-1} (SD= ± 6.4 mg C m^{-2} h^{-1}). In comparison, the average rates in the southern region were approximately five times higher than the average rates in the northern region (43.3 ± 41.7 mg C m^{-2} h^{-1}).

The vertically integrated nitrate production rates ranged from 0.14 to 18.77 mg NO_3 m^{-2} h^{-1}, with a mean of 2.72 mg N m^{-2} h^{-1} (SD = ± 5.51 mg N m^{-2} h^{-1}), whereas the ammonium production rates ranged from 1.16 mg NH_4 m^{-2} h^{-1} to 16.16 mg NH_4 m^{-2} h^{-1}, with a mean of 4.66 mg NH_4 m^{-2} h^{-1} (SD = ± 4.38 mg NH_4 m^{-2} h^{-1}) (Fig. 6). The total nitrogen (nitrate+ammonium) production rates ranged from 1.31 mg N m^{-2} h^{-1} to 34.94 mg N m^{-2} h^{-1}, with a mean of 7.38 mg N m^{-2} h^{-1} (SD = ± 9.71 mg N m^{-2} h^{-1}). At most stations except for stations of CL8 and CS8R, the ammonium production rates were generally higher than the nitrate production rates (Fig. 6). The average nitrate production rate was 0.41 mg NO_3 m^{-2} h^{-1} (SD = ± 0.51 mg NO_3 m^{-2} h^{-1}) in the northern region, whereas the average nitrate production rate for the southern region was 4.64 mg NO_3 m^{-2} h^{-1} (SD = ± 7.13 mg NO_3 m^{-2} h^{-1}). In comparison, the average ammonium production rates for the northern and southern regions were 2.56 mg NH_4 m^{-2} h^{-1} (SD = ± 1.74 mg NH_4

m^{-2} h^{-1}) and 6.41 mg NH_4 m^{-2} h^{-1} (SD = ± 5.28 mg NH_4 m^{-2} h^{-1}), respectively.

3.5. Statistical analysis of environmental variables according to geographic regions

One-way analysis of variance (ANOVA) was performed to assess significant regional differences in the environmental and biological variables of the two geographic regions (i.e., northern and southern). One-way ANOVA revealed significant regional differences for some of the environmental and biological variables in the study area (Table 2). The temperature and salinity of the surface were significantly different due to the effects of various water masses in the region. The stratification also exhibited a significant regional variability due to the higher accumulation of freshwater in the southern region ($p < 0.05$). However, the mean mixed layer depths were not significantly different, with means of 7.6 m (SD= ± 2.8 m) and 8.4 m (SD = ± 2.4 m) for the northern and southern regions, respectively (Table 2). The mean depths of nitracline were similar between the regions, although there were differences between the stations. The ambient nutrient concentrations of the upper 30 m showed highly significant differences, with higher concentrations in the southern region, although the phosphate concentration was not significantly different between the regions (Fig. 3 and Table 2). In addition, the chlorophyll *a* concentrations were significantly different ($p < 0.05$), with a value that was approximately two times higher in the southern region than in the northern region.

3.6. FWC distribution

To understand the potential effects of recent changes in the FWC on the primary production in the Chukchi Sea, the FWC data obtained from the three RUSALCA expeditions were used for a comparison. In 2012, the FWC in the study area were ranged from 2.1 to 8.5 m, with a mean of 4.5 m (SD = 1.2 m) (Fig. 7a). The strongest freshwater accumulation was observed in the western side of the CS section and north of the Herald Canyon (FWC = 6.7-8.5 m), whereas the lowest freshwater accumulation was observed at the center of the CL section in the southern region (FWC = 2.8-3.7 m) (Fig. 7a). The FWC in 2009 ranged from 2.6 to 11.8 m, with a mean of 5.1 m (SD = ± 1.7 m) (Fig. 7b). The mean value in 2009 was a little higher than that of 2012 due to the high accumulation of FWC from the East Siberian Sea and the region north of Herald Canyon (Fig. 7b). In 2009, the FWC in the southern region was evenly distributed with an accumulation of below 6 m. In 2004, the mean FWC was 4.7 ± 1.3 m and ranged from 2.0 to 9.9 m (Fig. 7c). Unlike the observations from 2012 and 2009, the FWC in the southern region in 2004 indicated a low accumulation in the western side and a progressive increase in FWC toward the eastern side (Fig. 7c).

4. Discussion

4.1. Regional carbon and nitrogen production rates in 2012

In this study, there were large differences in the carbon and nitrogen production rates the between southern and northern regions (Figs. 5, 6 and Table 2). The average rate of carbon production in the southern region was about five times higher than that of the northern region (Fig. 5 and Table 2). Similarly, the total nitrogen (nitrate+ammonium) production rates were approximately four times higher in the southern region than in the northern region (Fig. 6). In particular, the regional differences were much higher for the nitrate production rate than the ammonium production rate (Table 2). We also found that the carbon production rates obtained from all of the RUSALCA expeditions (2004, 2009 and 2012) showed highly significant differences between the regions ($p < 0.001$, n = 43) (data not shown).

The regional differences in phytoplankton production rates may have resulted from the different environmental conditions, as revealed the statistical analysis (Table 2). Especially, the different nutrient conditions and thereby different phytoplankton biomasses may be an important reason for the regional differences in the production rates of phytoplankton, since there was a positive relationship between the ambient nutrient concentrations (nitrate) and the chlorophyll *a* concentrations integrated from surface to 30 m (r = 0.6468, $p < 0.0001$, n = 41) (Fig. 8a). Moreover, we found that the carbon, nitrate and ammonium production rates were significantly correlated with the chlorophyll *a* concentration (r = 0.9234, r = 0.9641 and r = 0.9798, $p < 0.0001$, n=11, respectively) (Fig. 8b). Even though the regional differences in temperature are quite similar to that in primary production rates, there was no significant relationship between temperature and primary production rates in this study. According to Gosselin et al. (1997), the latitudinal variability in the phytoplankton production and biomass were primarily regulated by changes in the surface ice cover and the depths of the surface mixed layer, which determine the amount of light available to the phytoplankton in the water column. However, this was not the case in our study, as the mixed-layer depths were not significantly different between the southern and northern regions of the Chukchi Sea (Table 2).

The production/biomass ratio (P/B ratio), which was calculated by dividing the daily carbon production rate (mg C m^{-2} d^{-1}) by the integrated chlorophyll *a* concentration (mg chl m^{-2}), in the southern region (9.61 ± 4.26 mg C (mg chl-a)$^{-1}$$d^{-1}$) was somewhat higher than the P/B ratio in the northern region (5.46 ± 1.27 mg C (mg chl-a)$^{-1}$$d^{-1}$). This result indicated better carbon production efficiency by the phytoplankton in the southern region. Therefore, the regional differences in the primary production rates may have been affected by different production efficiencies in addition to the different phytoplankton biomasses induced under different nutrient conditions.

4.2. Primary production rate in 2012 compared to the previous RUSALCA expeditions

Based on a 15-hour photo period in the Chukchi Sea (Hansell and Goering 1990; Lee et al. 2007; Yun et al. 2014) and the hourly carbon production rates measured in this study, in 2012, the daily carbon production rates integrated from the surface to 1 % light depth ranged from 0.02 to 1.61 g C m^{-2} d^{-1}. The daily carbon production rate in 2012 (mean ± SD = 0.42 ± 0.52 g C m^{-2} d^{-1}), which was averaged from the values from all the productivity stations, was quite similar to the daily carbon production rate of 2004 (mean ± SD = 0.41 ± 0.53 g C m^{-2} d^{-1}) reported by Lee et al. (2007). The production rates (mean ± SD = 0.26 ± 0.24 g C m^{-2} d^{-1}) obtained in 2009 and presented by Yun et al. (2014) were significantly lower than those from 2012 and 2004, which is believed to be due to the different sampling times among the three cruises because the seasonal variation in primary productivity is quite large in this region (Springer and McRoy 1993; Wang et al. 2005; Hill et al., 2013). These differences in the primary production rates obtained by the three cruises also may have been due to interannual variations in primary productivity in the Chukchi Sea, as Hirawake et al. (2012) used satellite remote sensing data obtained from 2002 to 2010 to show that the Chukchi Sea experiences strong interannual variation in August and September.

In 2012, the average daily carbon production rates were 0.66 g C m^{-2} d^{-1} (SD = ± 0.62 g C m^{-2} d^{-1}) in the southern region and 0.14 g C m^{-2} d^{-1} (SD = ± 0.10 g C m^{-2} d^{-1}) in the northern region. By comparison, the average daily carbon production rates in the southern and northern regions were 0.57 g C m^{-2} d^{-1} (SD = ± 0.64 g C m^{-2} d^{-1}) and 0.16 g C m^{-2} d^{-1} (SD = ± 0.18 g C m^{-2} d^{-1}) in 2004, respectively, and 0.38 g C m^{-2} d^{-1} (SD = ± 0.26 g C m^{-2} d^{-1}) and 0.14 g C m^{-2} d^{-1} (SD = ± 0.16 g C m^{-2} d^{-1}) in 2009, respectively. From the regional comparisons, we found that the pattern of primary production in the Chukchi Sea is largely different depending on regions. The primary production rates in the northern region were consistently low, since the regionally low nutrient conditions and phytoplankton biomass. Thus, they were not largely changed among the three cruises. In contrast, the primary production rates for the southern region were considerably variable among the three cruises, although they including seasonal and interannual variations. Since this study revealed that the nutrient is an important factor in controlling primary production, the recent change in primary production for the southern region could be induced by changes in nutrient conditions in the region. The changes in freshwater inputs in the region may have been closely related to the nutrient and primary production variability (detailed in section 4.3).

4.3. The effects of FWC on the nutrients and primary production in the southern Chukchi Sea

FWC plays an important role in determining the nutrient distribution/inventory and, therefore, the availability of nutrients for phytoplankton growth in the Arctic Ocean. Coupel et al. (2015) showed that the strong freshening of the Canada Basin resulted in the deepening of the nitracline, which had a negative impact on primary production. In addition, Yun et al. (2014) reported that the low primary production rate in the Chukchi Sea could be due to the decreases in the nutrient and chlorophyll *a* concentrations that resulted from the increased input of fresh waters. In 2012, we found that the freshwater had strongly accumulated in the western side of the southern Chukchi Sea and especially in the CS section (Fig. 7a) due to an inflow of fresh Siberian Coastal Water or sea ice meltwater. This could have resulted in the low primary production rates observed in the western region and the CS section of the southern Chukchi Sea (Fig. 5). In contrast, relatively high production rates were observed in the center of the CL section, the region with the lowest accumulation of freshwater (Figs. 5 and 7a). The strong inflow of Siberian Coastal Water from the East Siberian Sea into the Chukchi Sea was also found in 2009, though it was not detected in 2004 (Figs. 7b and 7c). These inputs of freshwater presumably influenced the nutrient reservoir and its replenishment from deeper layers by altering stratification of the water column (Coupel et al., 2015), eventually driving the observed changes in primary production in the region. Based on data obtained from southern region during three cruises, we found that FWC had a significant negative effect on the nitrate concentrations (r = 0.5363, $p < 0.05$) and primary production rates (r = 0.5645, $p < 0.05$) (Figs. 9a and 9b). As a result, the primary production rates in the Chukchi Sea could be highly significantly correlated with the nitrate concentrations (r = 0.7482, $p < 0.001$) (Fig. 9c). Therefore, we might conclude that the primary production in the Chukchi Sea could be primarily controlled by nutrient availability related to FWC variability, as reported in previous studies conducted in different regions of the Arctic Ocean (Tremblay and Gagnon, 2009; Tremblay et al., 2002, 2006, Coupel et al., 2015). However, the influence of ocean circulations should be examined further because the ocean circulation such as pacific inflow and Beaufort Gyre can redistribute the amount of freshwater (Giles et al., 2012), eventually leading to regional differences in FWC (Giles et al., 2012; Morison et al., 2012). Additionally, we need to consider the local wind field, as the spatial distribution of FWC is largely dependent on the wind and is controlled by atmospheric pressure patterns (Anderson et al., 2011).

4.4. Current status of the primary production in the Chukchi Sea

To understand the recent status of primary production in the Chukchi Sea, the *in situ* measurements of primary production in the region in recent years were plotted with those from the previous studies in decades ago (Fig. 10). The average carbon production rate from the three RUSALCA cruises in the Chukchi Sea was 0.33 g C m^{-2} d^{-1} (SD = 0.40 g C m^{-2} d^{-1}). In addition, Hill and Cota (2005) reported that the mean daily production rate during the initial ice breakup was 0.8 g C m^{-2} d^{-1} in 2002. The daily production rates obtained by Lee et al. (2012) and Lee et al. (2013) were 0.54 and 0.86 g C m^{-2} d^{-1}, respectively (Fig. 10). Even though the different sea ice condition could affect difference in the productivity, these recent measurements of primary production (Hill and Cota 2005; Lee et al., 2007, 2012 and 2013; Yun et al., 2014; This study) showed significantly lower rates compared with the previously reported rates from the region (Hameedi, 1978; Korsak, 1992; Zeeman, 1992) (*t*-test, $p < 0.01$). Moreover, it is obviously shown a decreasing trend of primary production (r = 0.7689, $p < 0.01$) (Fig. 10). This is very interesting because primary production could be expected to increase in the region due to the increased light availability to the phytoplankton. For example, based on satellite ocean color data, Arrigo et al. (2008) found large increases in the annual net primary production on the continental shelves of the Chukchi Sea as well as Siberian and Laptev seas due to increased open water areas and longer growing seasons. However, the *in situ* measurements of primary production in recent years (Hill and Cota 2005; Lee et al., 2007, 2012 and 2013; Yun et al., 2014; This study) have shown consistently lower primary production rates compared to those in previous studies.

The strong seasonal and interannual variation in the region could be suggested for reason causing the low primary production, as discussed above. Hill et al. (2013) found that the seasonal variations in primary production in the southern Chukchi Sea peaked in July and then progressively declined in August and September. In fact, the lowest primary production rates given by Yun et al. (2014) were obtained from the late summer season (i.e., from 1 to 30, September, 2009) compared with the rates found in the present study (from 30 August to 14 September, 2012) or in Lee et al. (2007) (from 11 to 22 August, 2004). In comparison, previous studies (Hameedi, 1978; Korsak, 1992; Zeeman, 1992) included the measurements obtained from July to August (Fig. 10). However, their measurements just starting from the end of July were mostly done during August (Korsak, 1992; Zeeman, 1992). Although recent measurements from the three RUSALCA cruises (2004, 2009 and 2012) may not have reflected the highest values (i.e., July) of primary production, the measurements from Hill and Cota (2005) or Lee et al. (2012 and 2013) include the values in the mid-July and early August. Therefore, the recent low rates of primary production might be reflected by decreasing trend rather than results of seasonal and interannual variations.

More plausible reason for the recent low primary production in the Chukchi Sea could be due to the decreased concentrations of nutrients and chlorophyll *a*. According to Whitledge and Lee (unpublished data), in recent years, there have been significant decreases of 30-50% in nutrient concentrations and approximately 40% in the integrated chlorophyll *a* concentration in the Bering Strait and the Chukchi Sea. Based on the significant relationships between primary production and the nutrient and FWC (discussed in section 4.3), the recent decrease in nutrient and chlorophyll a concentrations may have been closely related to the changes in freshwater inputs in the region. According to Serreze et al. (2006), there was recently larger import of freshwater through the Bering Strait compared with previous estimates. Therefore, the recent decreases in the concentrations of major inorganic nutrients and chlorophyll *a* may have resulted in lower primary production rates in the Chukchi Sea.

Recently, the freshwater content in the Arctic Ocean, which includes river discharge, pacific water inflow through the Bering Strait, sea ice melt water and net precipitation (Jones et al., 2008), has increased over the past few decades. If the increased freshwater content in the Chukchi Sea are continuously observed, the Chukchi Sea might have become less productive region compared with previous decades.

5. Conclusions

This study reported the regional characteristics of primary production in the Chukchi Sea and recent trend of primary production based on *in situ* measurements. The different nutrient conditions and phytoplankton biomass could be an important reason for the regional differences in the production rates of phytoplankton. Based on comparison between previous studies in decades ago and recent measurements, we found that recent primary production in the Chukchi Sea showed a decreasing trend. The changes in freshwater inputs in the region may have been closely related to the nutrient and primary production variability. Although Coupel et al. (2015) reported that the recent freshening of the Arctic

Ocean does not significantly affect primary production in the Chukchi shelf based on comparison with measurements in the deep Canada Basin, our results showed that the freshwater variability in the Chukchi Sea has had a large influence on the recent changes in primary production by controlling the nutrient inventory. If the increased freshwater inflow persists, the primary production in the region will considerably decrease, ultimately resulting in changes in the regional characteristics of primary production. However, a large interannual variability of primary production remains despite the statistical significance observed in this study. Therefore, more measurements under various environmental conditions are needed to better understand the recent variations in the primary production in the Chukchi Sea. In particular, there could be some changes in the phytoplankton community structures because the smaller cells benefit more than the larger cells under increased freshening conditions (Li et al., 2009).

Acknowledgments

We thank the captain and crew of the *Professor Khromov* for their outstanding assistance during the cruise. We gratefully acknowledge the physical oceanographers from Woods Hole Oceanographic Institute (M. Swartz, S. Mills and J. Pietro), who provided the CTD data. We also appreciate NOAA for supporting the RUSALCA program. This study was financially supported by the “2015 Post-Doc. Development Program” of Pusan National University, Korea. This work was also supported by grants from the Korea-Polar Ocean in Rapid Transition (K-PORT; PM14040) program funded by the Ministry of Oceans and Fisheries, Korea.

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Figure Legends

Fig. 1. Locations of sampling stations during the 2012 RUSALCA expedition in the Chukchi Sea. The primary production rates were measured at the stations identified by blue circles. The st. CS8R location represents a revisit to st. CS8.

Fig. 2. The distribution of surface temperature [T_{sur}][^oC] (a), surface salinity [S_{sur}] (b), stratification index [$\Delta\sigma_t$][kg m^{-3}] (c), surface mixed layer depth [Z_m][m] (d), and nitracline depth [Z_{nit}][m] (e) during the 2012 RUSALCA .

Fig. 3. The distributions of the integrated concentrations of ambient nitrite+nitrate [NO_2+NO_3][mmol m^{-2}] (a), ammonium [NH_4][mmol m^{-2}] (b), phosphate [PO_4][mmol m^{-2}] (c), and silicate [SiO_4][mmol m^{-2}] (d) from surface to 30 m during the 2012 RUSALCA.

Fig. 4. The chlorophyll *a* concentrations integrated from surface to 30 m [mg m^{-2}] during the 2012 RUSALCA.

Fig. 5. Hourly carbon uptake rates [mg C m^{-2} h^{-1}] integrated from the surface to 1% light depth during the 2012 RUSALCA.

Fig. 6 Hourly nitrate and ammonium uptake rates [mg N m^{-2} h^{-1}] integrated from the surface to 1% light depth during the 2012 RUSALCA.

Fig. 7. The distributions of Fresh Water Content (FWC in m) in the Chukchi Sea in 2012 (a), 2009 (b) and 2004 (c).

Fig. 8. Relationships between nitrate concentrations (mmol m^{-2}) and chlorophyll a concentrations (mg m^{-2}) integrated from surface to 30 m (a) (n = 41); chlorophyll a concentrations (mg m^{-2}) and daily carbon (g C m^{-2} d^{-1}) and nitrogen production rate (mg N m^{-2} d^{-1}) over the euphotic zones (b) (n = 11). All data obtained

during the 2012 RUSALCA.

Fig. 9. Relationships between FWC (m) and nitrate concentrations (mmol m^{-2}) (a); FWC (m) and daily primary production rate (g C m^{-2} d^{-1}) (b); nitrate concentrations (mmol m^{-2}) and daily primary production rate (g C m^{-2} d^{-1}) (c). All data obtained from southern region during the three RUSALCA cruises.

Fig. 10. A recent trend of primary production based on *in situ* carbon uptake measurements (^{13}C or ^{14}C) in the Chukchi Sea. All Chukchi Sea represents southern and northern Chukchi Sea.

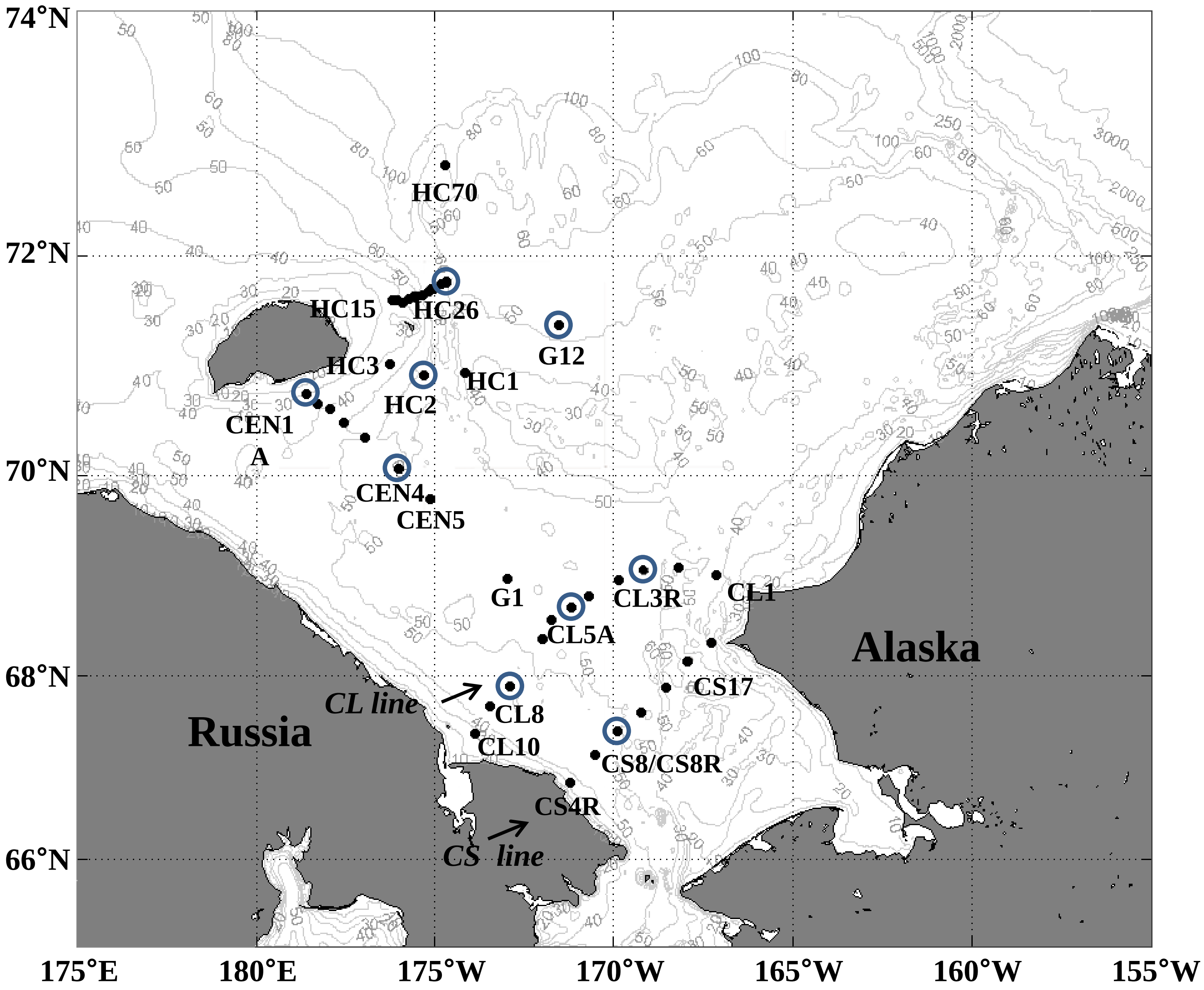
Table 1 Location, water depth (m) and euphotic depth (Z_{eu}) for primary productivity stations in the Chukchi Sea in 2012.

Region	Station	Date (mm/dd/yr)	Location		Depth (m)	Z_{eu} (m)
			Latitude (°N)	Longitude (°W)		
Northern	CEN4	09/05/12	69.9828	-175.6857	63	34
	CEN1A	09/06/12	70.7085	-178.2988	38	20
	HC2	09/07/12	70.9000	-175.0127	74	36
	HC26	09/08/12	71.7878	-174.3945	55	46
	G12	09/11/12	71.3980	-171.2597	55	46
Southern	CS8	08/30/12	67.4312	-169.6030	51	24
	CS17	09/01/12	68.2983	-167.0418	40	22
	CL5A	09/02/12	68.6407	-170.9423	59	20
	CL3R	09/12/12	69.0048	-168.9000	57	26
	CL8	09/13/12	67.8692	-172.5482	53	24
	CS8R	09/14/12	67.4312	-169.6030	51	26

Table 2 Summary of one-way analysis of variance (ANOVA) for environmental variables in two geographic regions of the Chukchi Sea in 2012. The mean values (ranges in parentheses) and their significant differences (> or <) between northern and southern regions are given for surface temperature (T_{sur}), surface salinity (S_{sur}), stratification index ($\Delta\sigma_t$), surface mixed layer depth (Z_m), nitracline depth (Z_{nit}), fresh water content (FWC), nitrite+nitrate concentration of the upper 30 m (NO_2+NO_3), ammonium concentration of the upper 30 m (NH_4), phosphate concentration of the upper 30 m (PO_4), silicate concentration of the upper 30 m (SiO_4) and chlorophyll *a* concentration of the upper 30 m (Chl-a). *: $p < 0.05$, **: $p < 0.01$, ***: $p < 0.001$, ns: not significant. n = 52. Also given are mean and range values for carbon production (CP), nitrate production (NP) and ammonium production (AP) (all mg C or N m^{-2} d^{-1}). n=11.

Variables	**Northern**		**Southern**
T_{sur} (°C)	0.62 (-1.33 ~ 4.13)	<***	3.89 (1.60 ~ 8.53)
S_{sur}	29.27 (27.30 ~ 32.04)	>*	27.48 (21.48 ~ 32.35)
$\Delta\sigma_t$ (kg m^{-3})	3.15 (0.79 ~ 5.34)	<*	4.47 (0.71 ~ 9.71)
Z_m (m)	7.6 (4.0 ~ 14.0)	ns	8.4 (4.0 ~ 14.0)
Z_{nit} (m)	13.0 (2.5 ~ 30.0)	ns	12.6 (2.5 ~ 35.0)
NO_2+NO_3 (mmol m^{-2})	75.01 (21.51 ~ 218.22)	<*	134.15 (21.82 ~ 355.43)
NH_4 (mmol m^{-2})	40.49 (15.36 ~ 86.93)	<**	61.22 (28.54 ~ 109.51)
PO_4 (mmol m^{-2})	22.19 (5.43 ~ 34.26)	ns	25.95 (8.30 ~ 43.57)
SiO_4 (mmol m^{-2})	245.49 (104.79 ~ 800.49)	<***	410.86 (129.17 ~ 669.94)
Chl-a (mg/m^{-2})	21.7 (2.2 ~ 69.3)	<*	54.5 (3.1 ~ 286.4)
CP (mg C m^{-2} d^{-1})	134.7 (16.3 ~ 280.7)		649.1 (151.3 ~ 1628.9)

NP (mg N m^{-2} d^{-1})	6.1 (2.2 ~ 19.9)	69.7 (4.5 ~ 281.6)
AP (mg N m^{-2} d^{-1})	38.4 (17.4 ~ 83.6)	96.2 (45.0 ~ 242.4)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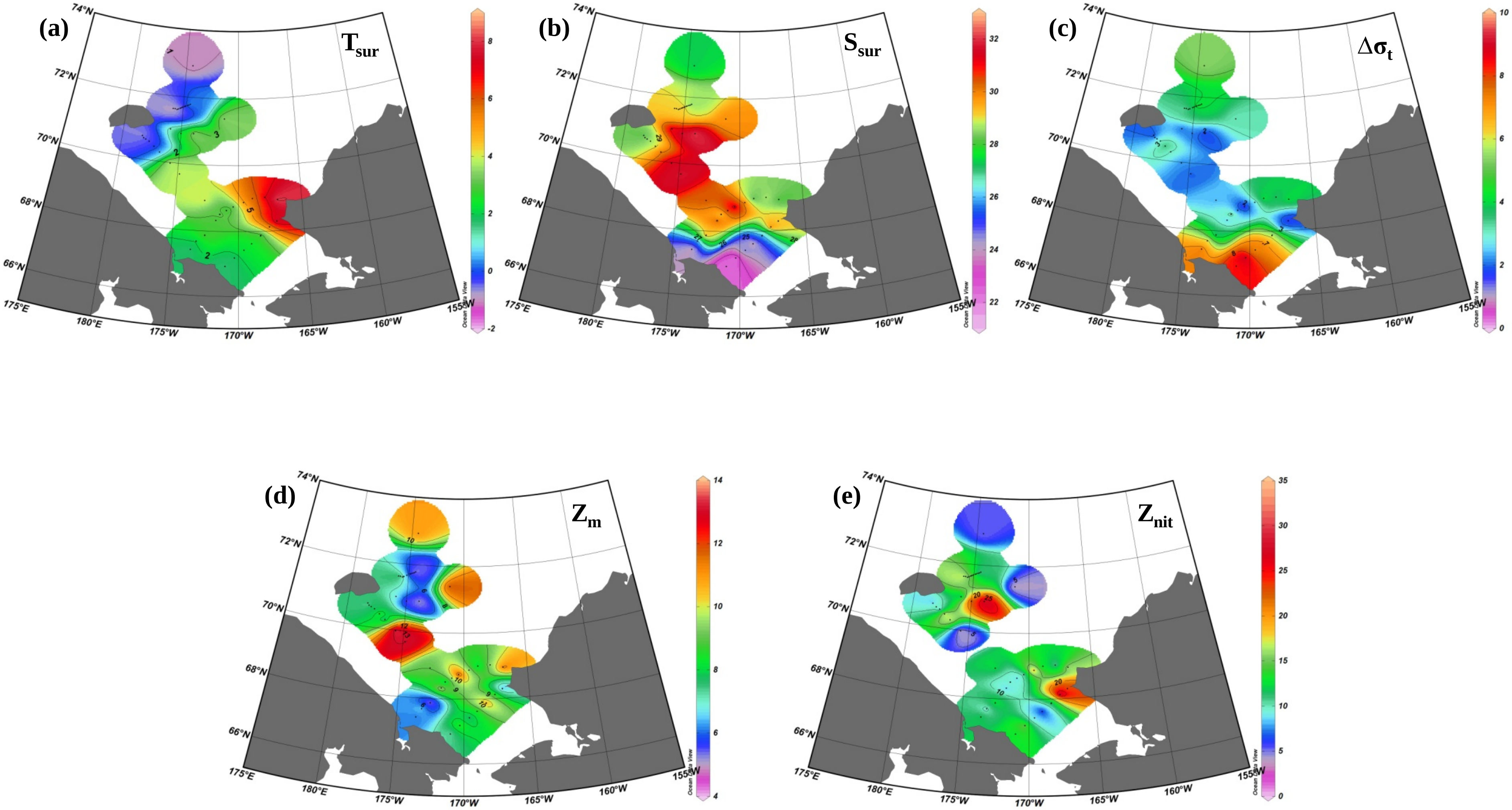
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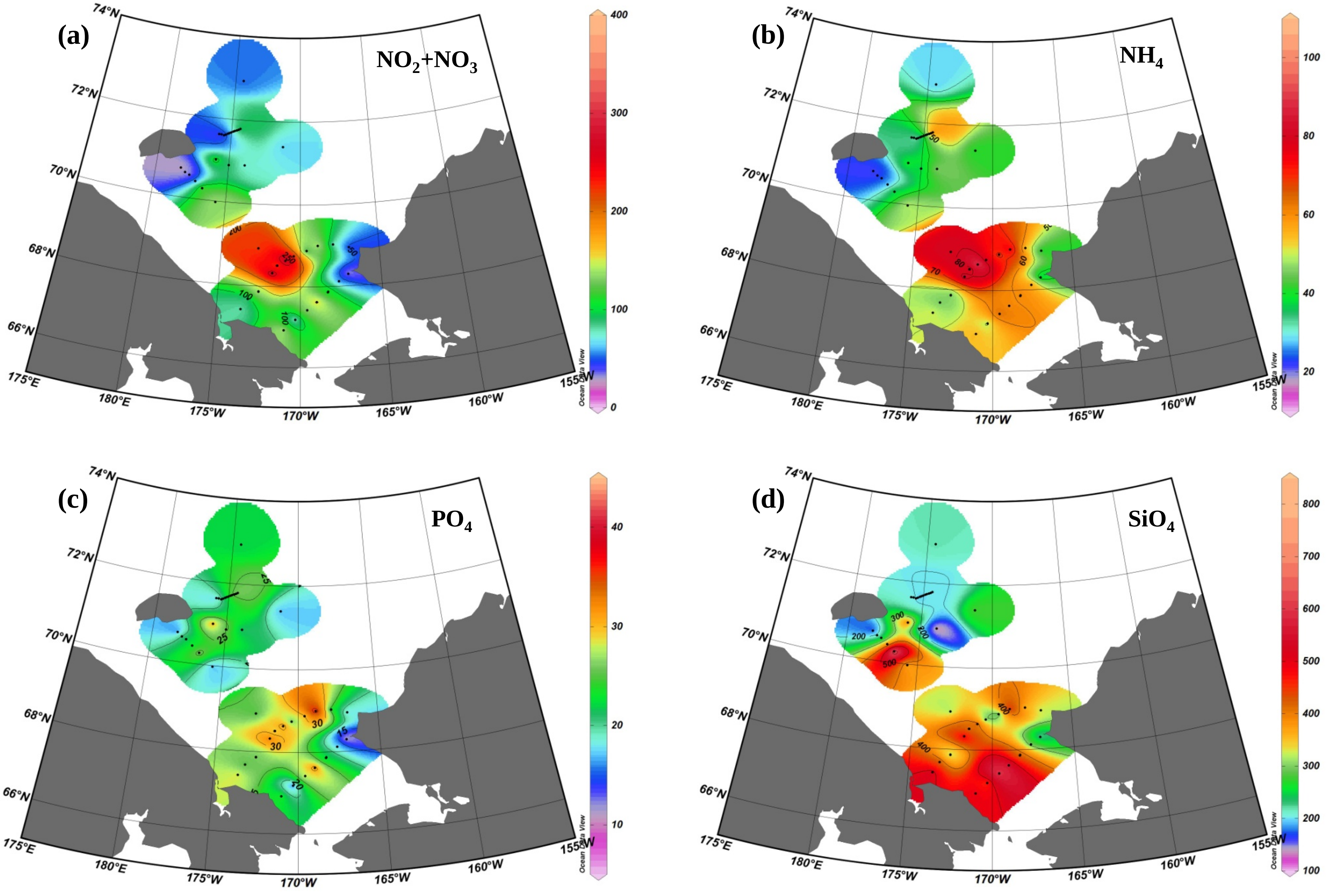
Ocean Data View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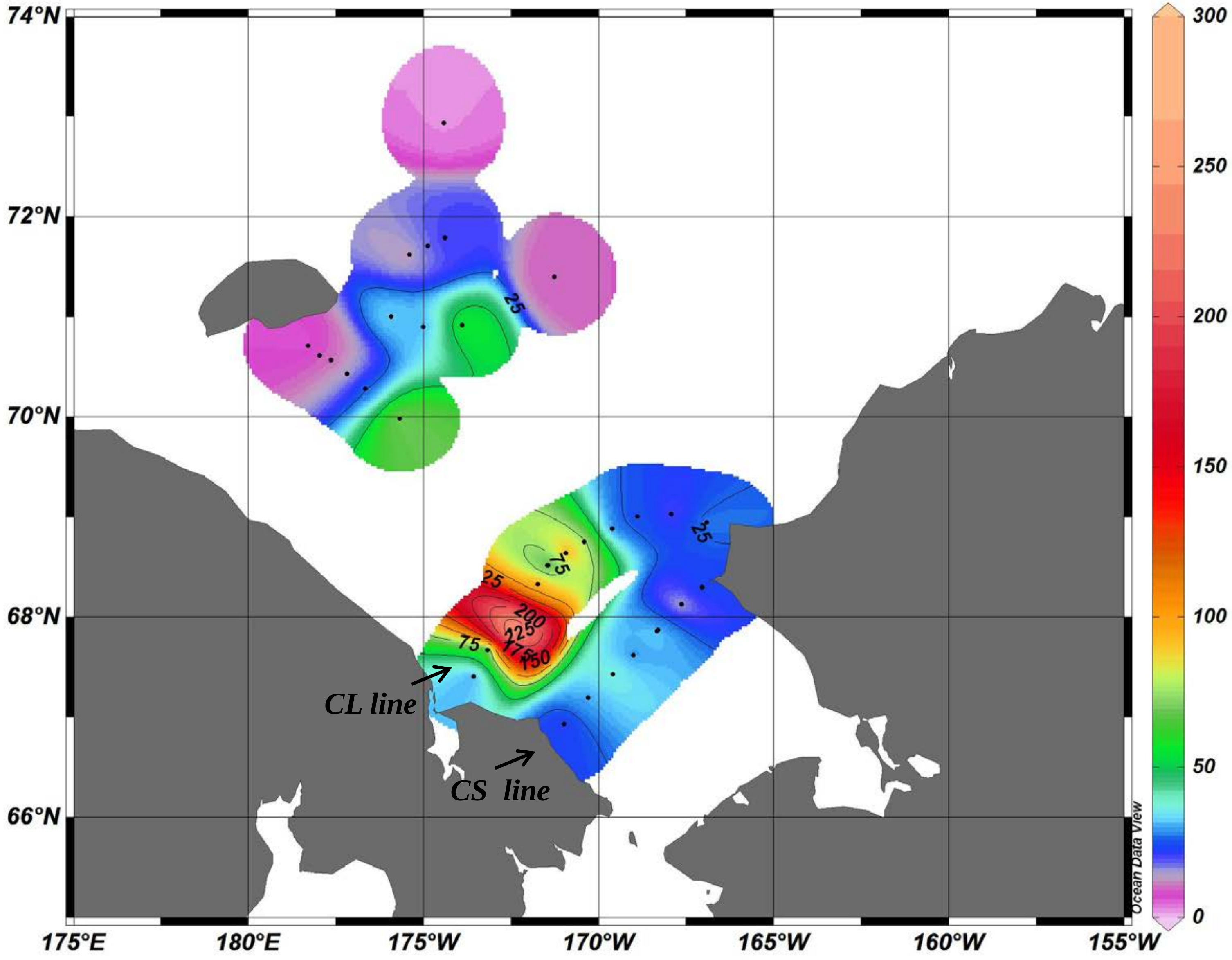
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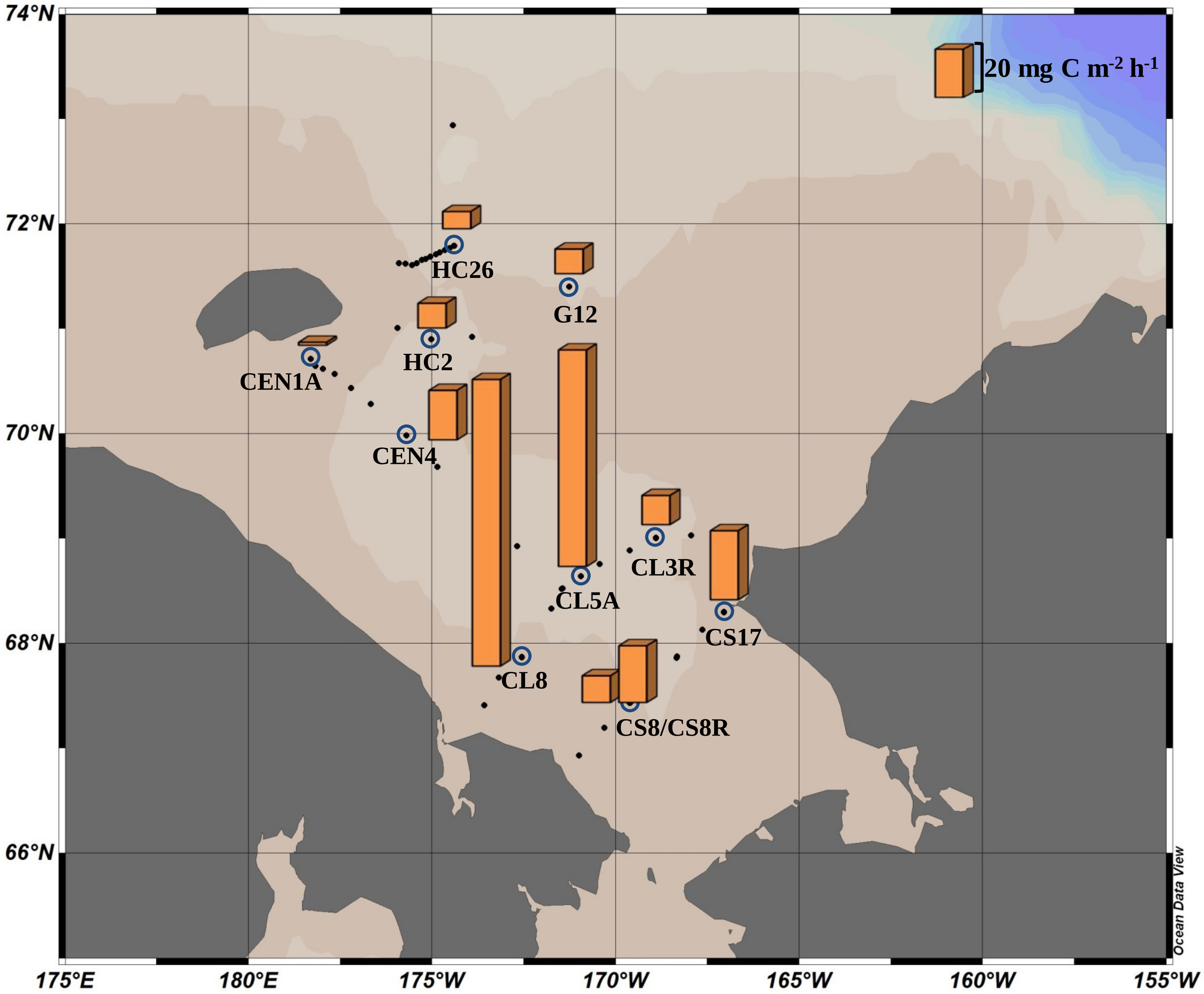
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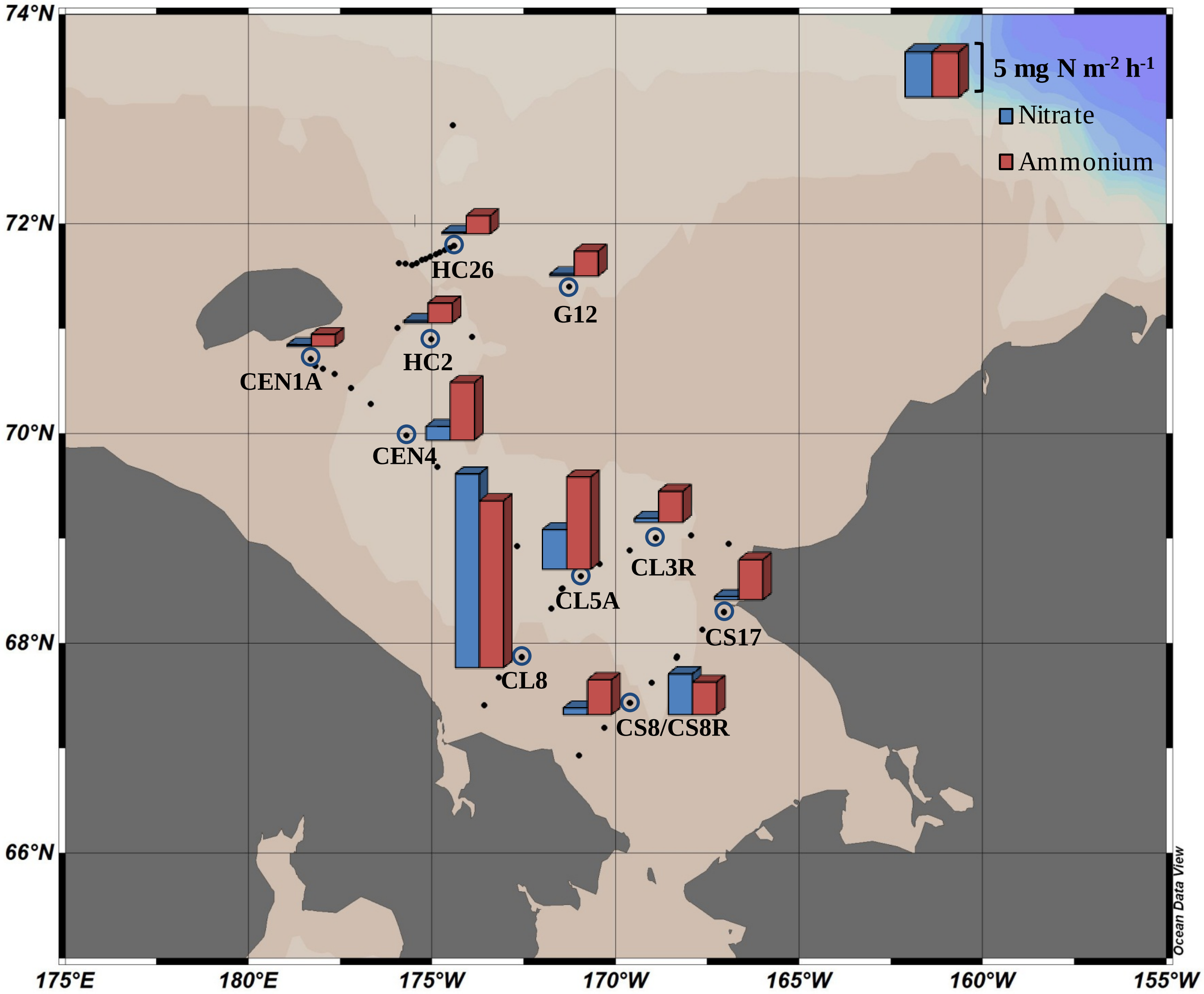
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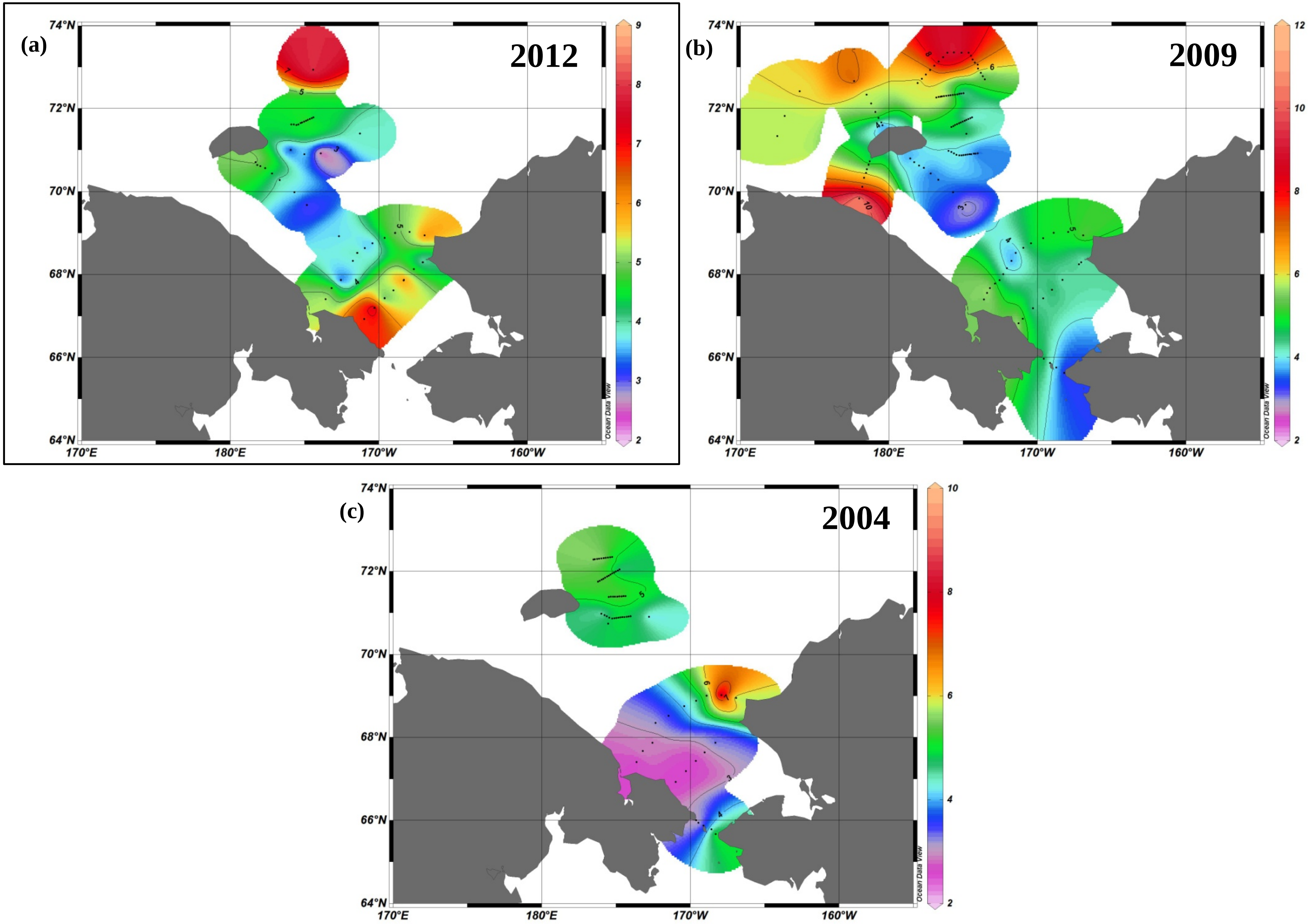
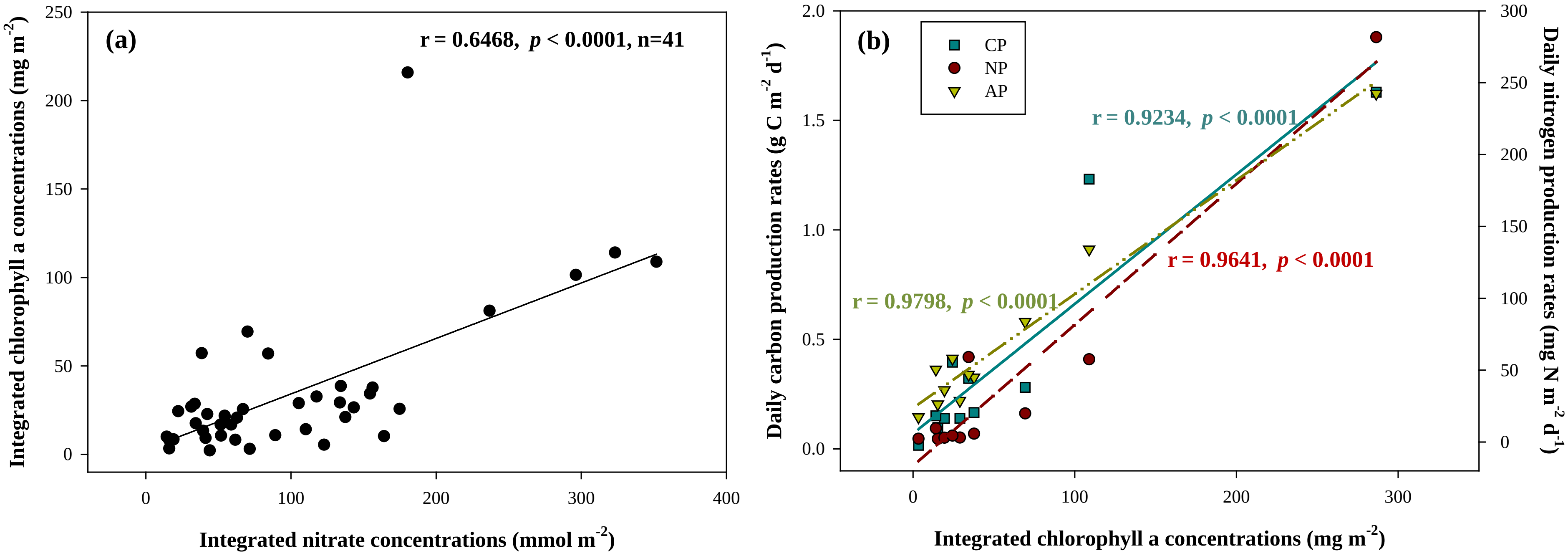
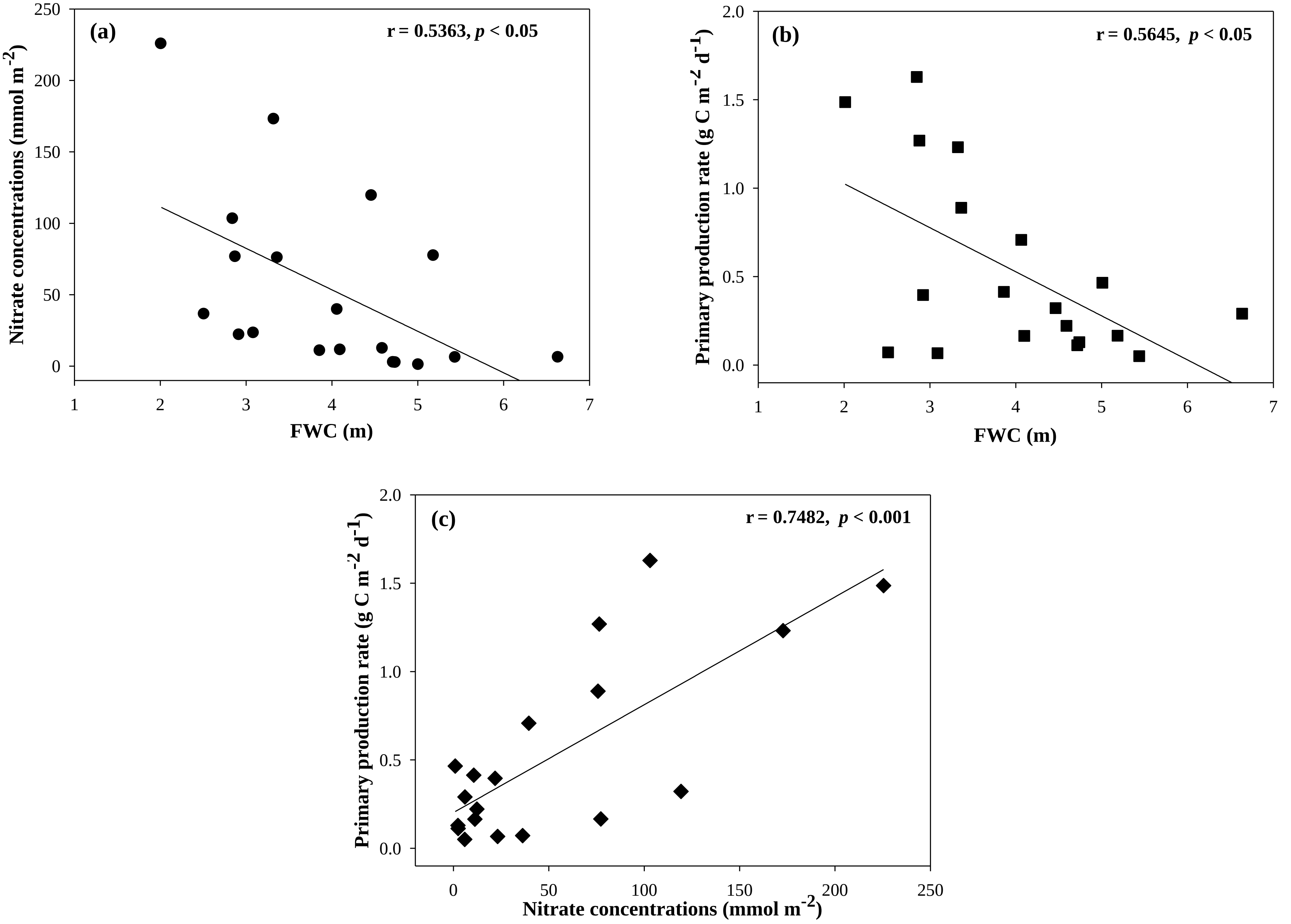
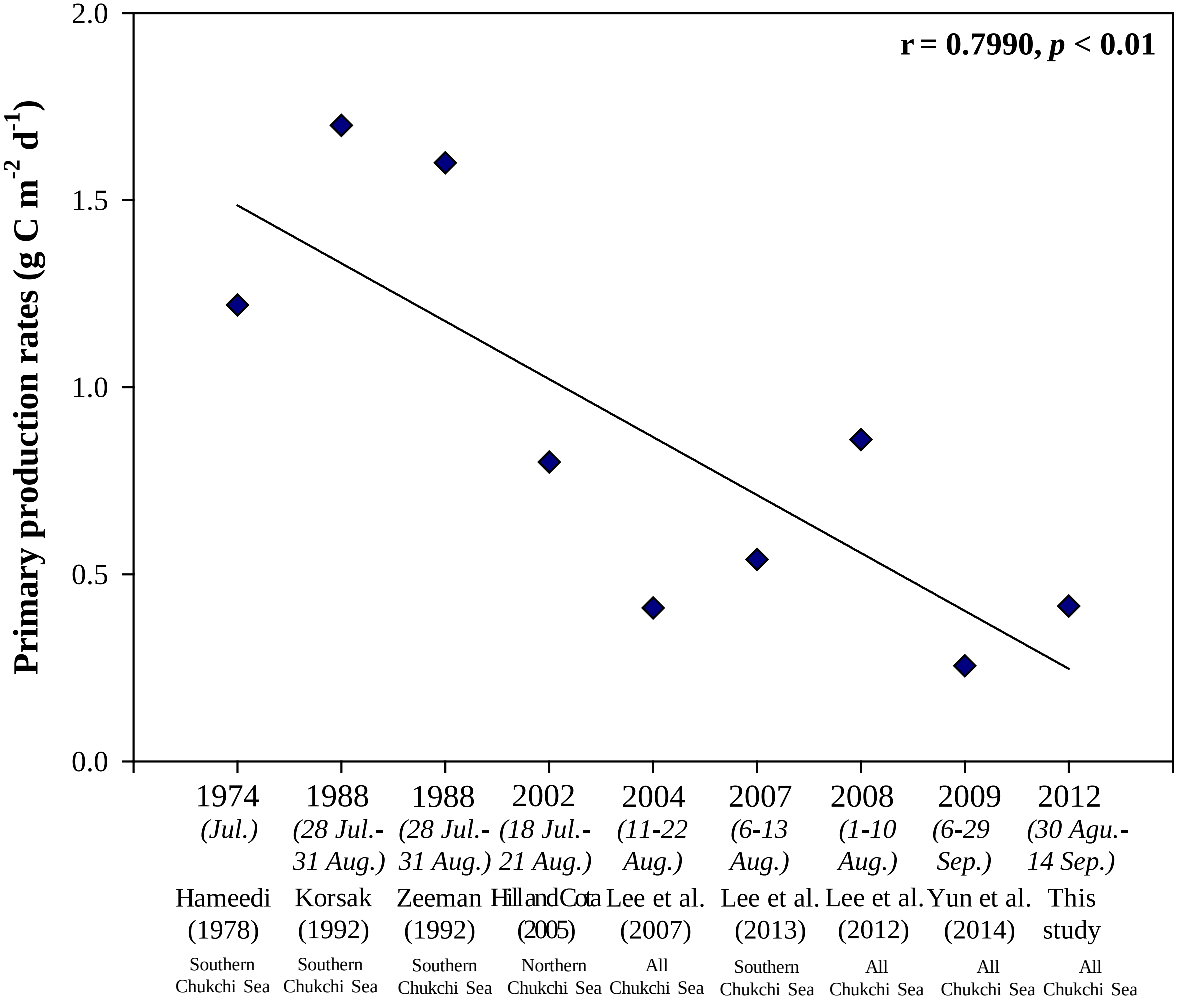
CS8/CS8R
Ocean Data View


(a)
2012
(b)
2009
(c)
2004
74°N
72°N
70°N
68°N
66°N
64°N
170°E
180°E
170°W
160°W
9
8
7
6
5
4
3
2
12
10
8
6
4
2
10
8
6
4
2
Ocean Data View


(a)
r = 0.6468, p < 0.0001, n=41
Integrated chlorophyll a concentrations (mg m-2)
250
200
150
100
50
0
0
100
200
300
400
Integrated nitrate concentrations (mmol m-2)
(b)
CP
NP
AP
r = 0.9234, p < 0.0001
r = 0.9641, p < 0.0001
r = 0.9798, p < 0.0001
Daily carbon production rates (g C m-2 d-1)
2.0
1.5
1.0
0.5
0.0
Daily nitrogen production rates (mg N m-2 d-1)
300
250
200
150
100
50
0
0
100
200
300
Integrated chlorophyll a concentrations (mg m-2)


(a)
r = 0.5363, $p < 0.05$
Nitrate concentrations (mmol m^{-2})
250
200
150
100
50
0
1
2
3
4
5
6
7
FWC (m)
(b)
r = 0.5645, $p < 0.05$
Primary production rate (g C m^{-2} d^{-1})
2.0
1.5
1.0
0.5
0.0
1
2
3
4
5
6
7
FWC (m)
(c)
r = 0.7482, $p < 0.001$
Primary production rate (g C m^{-2} d^{-1})
2.0
1.5
1.0
0.5
0.0
0
50
100
150
200
250
Nitrate concentrations (mmol m^{-2})


r = 0.7990, p < 0.01
Primary production rates (g C m^{-2} d^{-1})
2.0
1.5
1.0
0.5
0.0
1974 (Jul.) Hameedi (1978) Southern Chukchi Sea
1988 (28 Jul.-31 Aug.) Korsak (1992) Southern Chukchi Sea
1988 (28 Jul.-31 Aug.) Zeeman (1992) Southern Chukchi Sea
2002 (18 Jul.-21 Aug.) Hill and Cota (2005) Northern Chukchi Sea
2004 (11-22 Aug.) Lee et al. (2007) All Chukchi Sea
2007 (6-13 Aug.) Lee et al. (2013) Southern Chukchi Sea
2008 (1-10 Aug.) Lee et al. (2012) All Chukchi Sea
2009 (6-29 Sep.) Yun et al. (2014) All Chukchi Sea
2012 (30 Agu.-14 Sep.) This study All Chukchi Sea