We would like to thank both referees for their critical and constructive comments to our manuscript. Their comments helped to significantly improve the quality and clarity of the manuscript. We hope that our answers and revisions are sufficient to accept this work for publication in Biogeosciences. Please find our responses to each of the individual comments below.

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Referee # 1 Dr. Riemann

Received and published 22 September 2015

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10 Review of Gier et al. 2015. The paper concerns N2 fixation and sulfate reduction (SR) in sediments below OMZ waters off Peru. The work demonstrates an interesting coupling 11 between N2 fixation and SR, as also suggested by nifH gene analyses. Moreover, the study 12 indicates that organic matter load and sulfide are major drivers of N2 fixation. The paper 13 contributes to the compiling data on factors regulating diazotrophy and specifically to the 14 15 rather limited number of studies from sediments. The paper is generally well written, clear, and to the point. My points of criticism are overall minor, but should improve the readability 16

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and clarity of the paper. 18

- 1. The wording should be changed at several places in the abstract. The current version 19 20 seems to indicate that rates were measured in water, and not just in sediments. For
- instance line 6: "measured in OMZ mid-waters"; line 8: "Benthic N2 fixation profiles" etc. 21
- Please, make sure the reader cannot be misled to believe that water samples were 22
- 23 analyzed.
- The wording in the abstract regarding the measurements has been changed according to 24
- the referee's suggestions. 25
- 2. P1. I. 11. Define nifH genes 26
- A definition regarding the *nifH* gene has been added. 27 28
- 29 3. P1. I14. Delete "various"
- "Various" has been deleted. 30

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- 4. P6. I1. "These bacteria..."
- Changed. 33 34

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5. P6. I10-14. Unclear where this information comes from The author information (Dale et al., 2015) has been added. 36

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6. P7, I16-22. It would be good to reduce the overall length of the manuscript. This section could be easily reduced. Most readers will know the principle of acetylene reduction. We thank the reviewer for this suggestion. We reduced the method part regarding the

description of the acetylene reduction assay. 41

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7. p8, l5. Specify whether samples were analyzed onboard or stored somehow. Samples were analyzed onboard and this information has been added.

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- 8. P8, I13. OK, but why were they expressed as NA. Isn't that just confusing? If keeping it 46 as NA, then please explain why. 47
- As both referees pointed out that it is confusing to have nitrogenase activity (NA) and N2 48 49 fixation in the manuscript, values were recalculated for N₂ fixation and all figures, tables and text were changed accordingly and we now only refer to N₂ fixation. 50

- 52 9. P10, I2. Please, specify how many sequences were obtained per sample. Also, describe negative controls and whether they were blank. 53
- The information regarding the sequences and the negative controls has been added. 54

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10. P10, I14. How can you in the description of your sediments cite literature which is published before this sampling was carried out? This is your Results section – you should describe your results, not those of others.

Thanks for noticing. We agree with the referee and deleted this citation from the results part.

11. P10, I18. Redundant, described 3-4 lines higher up.

The sentence has been deleted.

12. P13. It should be evident from the text why the authors are interested in looking at C/N ratios. It is not enough to address that later in the discussion. Likewise, it should be explained why data on DIC flux are reported (Fig. 4), also how this was measured is unclear to me.

Information on why we looked at the C/N ratios and DIC values, as well as how DIC was measured has been added.

13. P14, I8. Rephrase. A novel clade cannot belong to anything. It may be related to something...

The sentence has been rephrased.

14. P15. L5-6. Again, this sounds like water samples. Please, rephrase Rephrased.

15. P15, I6. "Sometimes both depth profiles revealed similar trends". Clarify what is meant by depth profiles.

Clarified

16. P15, I8. "were"

86 Corrected.

17. P15, I21. What does "this study" refer to?

"This study" referred to the citation in the sentence before. The sentence was changed to make this clear.

92 18. P. 15, I28. "SR bacteria were..."

93 Corrected.

 19. P16, I11-15. Needs work. That samples have a "certain diversity" is not informative. Unclear what "these results" refer to (line 13). Farnelid et al. did not sample an OMZ (line 15).

The paragraph has been rephrased and the citation Farnelid et al. has been removed.

20. P17, I10-11. Weird and unclear sentence. Please, revise or remove.

101 The sentence has been removed.

21. P17, I20-28. I have not understood the point with the DIC fluxes. Please, make this clearer here as well as earlier in the manuscript.

As stated at comment number 12, information on the DIC fluxes has been added.

22. P20, I7-8. Sentence is out of context. Please, clarify the point or remove.

108 The sentence has been removed.

23. Figure 1, text. Please, define MUC.

Has been defined.

24. Figure 6, text. Delete "expressed". Clarify whether the sizes of the triangles are proportional to the number of sequences within each triangle. Moreover, indicate on the figure how many clones the triangles etc represent.

Expressed has been deleted. The sizes of the triangles should not be used for quantification. To make it clearer, all triangles were changed to the same size and the information how many clones each triangle represent has been added inside the triangles.

167 Referee # 2 Dr. Ionescu

168 Received and published: 11 November 2015

The paper by Gier et al discusses N fixation in oxygen minimum zones in marine sediments (specifically off the coast of Peru). The study suggests a link between sulfate reduction and N fixation in these environments and supports this previously mentioned hypothesis by rates measurements and phylogenetic data. This paper adds to our understanding regarding diazotrophy in sediments as well as highlights our gap in knowledge on the matter by showing that not all patterns can be explained by the presented data. The paper is generally well written with some exceptions where the English can be improved and the wording can be phrased in a more accurate manner.

The manuscript was cross-checked by an English speaker.

I tried to highlight these places in the comments below. Additionally as stated below the figures are not suited to the page size used by the journal and hence are often not readable.

We tried to improve the readability and clarity of the figures.

Page 14408 line 4 – The definition of formalin is an aqueous solution of 37% (m:v) formaldehyde. Hence 37 % formalin would mean 13 % formaldehyde. I guess this is not what the authors meant. To avoid misunderstandings, I suggest using 37% formaldehyde solution.

We agree with the referee and changed the information according to his suggestion.

Page 14408 line 5 – The acetylene reduction assay should not be used for longer than 48 h. Some consider this to be too long as well. The reason is that the saturation of the enzyme with acetylene leads to a lack of fixed N and reduction in cell viability and accordingly N-fixation (See for examples Seitzinger and Garber, 1987 MEPS 37 and references therein).

We agree with the referee and we are aware that incubation with acetylene can lead to a potential lack of fixed N, however to the best of our knowledge this is the standard method used for the determination of N_2 fixation in sediments (15N rate determinations are not feasible in sediments as incubation times would need to be several weeks to months to achieve a signal above the natural 15N sediment background). We have added in a recent citation (Bertics et al., 2013) that describes the method in further detail and we point towards this limitation of the method in the manuscript.

Page 14408 line 14. If you have converted the NA from C2H4 reduction to N fixation, why do the graphs in Fig 3 still discuss C2H4. While the value of 3 is not fixed for all environments it is indeed widely used. If you used it you can now refer to N2 rather than C2H4.

As both referees pointed out that it is confusing to have nitrogenase activity (NA) and N_2 fixation in the manuscript, values were recalculated for N_2 fixation and all figures, tables and text were changed accordingly and we now only refer to N_2 fixation.

Page 14409 line 27: 1 µl of BSA is not very informative as we don't know the concentration of the stock solution nor the reaction volume.

The information has been added.

216 Page 14410 line 25: No need for "The" in "The St. 9".

217 Changed.

Page 14411 line 3: "The deepest St. 10" means that there are several stations named St. 10 and this is the deepest of them. I suggest "The deepest station (10; 1025 m)..."

221 Or "St. 10 (the deepest; 1-25 m) ..."

222 Changed.

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224 Page 14411 line 11: Erase "The" in "The St. 4 and 6".

225 Corrected.

Page 14411 line 16: The shallowest St 1 - see my previous comment about the deepest St 226 227

228 Corrected.

> Page 14412 line 2: "Sediment depth profiles of N2 fixation activity are expressed in nitrogenase activity (NA), i.e. without the conversion factor of 3 C2H4: 1 N2" - Why convert in some cases (integrated rates) and not everywhere. Either you trust the conversion factor or you don't - no need to confuse the reader. Providing N2 fixation rates also allows for direct comparison with other studies. Please change this.

> As both referees pointed out that it is confusing to have nitrogenase activity (NA) and N2 fixation in the manuscript, values were recalculated for N₂ fixation and all figures, tables and text were changed accordingly and we now only refer to N₂ fixation.

Page 14412 line 9: In all cases so far you used the abbreviation St. even when several stations where mentioned why here the full word stations.

241 Corrected.

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Page 14412 line 8-10: The choice of sentence structure is not clear - Simply state: NA and SR rates where high (or highest) at the shallow St.... and lowest at deep St...

Changed.

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Page 14412 line 11 - page 14413 line 13: This section is messy and hard to follow. For example, St 1 has its own paragraph while the other stations are mentioned in a single paragraph. I also find this section too detailed. I believe you should only highlight the important things from the figures and not literally describe the graphs.

The paragraph has been shortened and only highlights from the graphs are specified. We hope this improves the clarity of this section.

Page 14413 line 15: The rate conversion was done from C2H4 to N2 and not to N (same in Fig. 4). Also the units (mmol) is missing.

256 Corrected.

257 258

Page 14413 line 25, 27, 28: mmol N2

Corrected.

Page 14414 line 7: Instead of "three novel clades and seven novel clades..." write "three and seven novel clades were detected, respectively".

Changed.

263 264 265

Page 14414 line 15: For the sake of correctness add: for a "known" Vibrio species...

266 Corrected. 267

Page 14416 line 21: The term heterotrophic N2 fixation is a bit obscure as autotrophy refers 268 to carbon. If the authors refer to N2 fixation by heterotrophs this should be stated in such a 269 270 271

The term heterotrophic has been clarified.

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275 276 Page 14416 line 23: The integrated N2 fixation rate and the Corg concentration clearly showed similar trends. Nevertheless, the use of the word "correlated" requires a statistical measure which I believe was not provided. Either provide such data (which should be straight forward) or rephrase the sentence to address the similarity in trends.

We agree with the referee and have rephrased the sentences accordingly.

279 Page 14417 line 22. Fig 5 should be Fig 4.

280 Corrected.

282 Figures:

Fig 2 – The figure is probably designed to cover and entire page (A4 or Letter). However, this is not the format used by this journal. Hence he printed figure is not readable. Online viewing requires as well magnification to 250 % for clear reading. Consider splitting into two panels spanning two pages.

The final format of Biogeosciences is letter format, hence the Fig. will be printed on a full page.

Fig. 3 – A similar problem as above with the addition of long text as the axis title. This cannot be read at 100% magnification on a screen or print.

The figure, as well as the axis title has been changed and the fonts were increased.

Fig. 4. As stated before I believe the correct unit is mmol N2 and not mmol N. Fonts need to be increased.

We agree with the referee and changed the unit. Also the fonts were increased.

Fig. 5. The same comment as above. Additionally, the yellow line and text are hardly visible.

The whole figure and all fonts have been increased, the yellow line has been darkened and the unit was changed accordingly.

Fig. 6. Needless to say that this is useless in print or at standard screen viewing. The fonts need to be larger. Sequences from this study should be bold. The shaded frames should be positioned in the background of the tree and not above it as they hide the text. Consider cutting the tree into two sections on two pages.

We agree with the referee and tried our best to increase the quality of the whole figure. The sequences from this study have been increased and were made bold. The shaded frames were changed to a transparent design for a better visibility. We considered cutting the tree into two sections, however this would make a direct comparison and association of the sequences more difficult for the reader and therefore we decided to show the tree on one page.

333 Nitrogen fixation in sediments along a depth transect through the Peruvian oxygen minimum zone 334 Jessica Gier^{1*}, Stefan Sommer¹, Carolin R. Löscher², Andrew W. Dale¹, Ruth A. Schmitz², 335 Tina Treude^{1,3*} 336 337 ¹ GEOMAR Helmholtz Centre for Ocean Research Kiel, Germany 338 ² Institute for Microbiology, Christian-Albrechts-University Kiel, Germany 339 ³Present address: University of California, Los Angeles, Department of Earth, Planetary & Space 340 Sciences and Department of Atmospheric & Oceanic Sciences, USA 341 342 *Correspondence: jgier@geomar.de, ttreude@g.ucla.edu 343

Abstract

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362 363 Benthic nitrogen (N2) fixation and sulfate reduction (SR) were investigated in the Peruvian oxygen minimum zone (OMZ). Sediment samples, retrieved by a multiple corer were retrieved by a multiple corer taken at six stations (70 - 1025 m water depth) along a depth transect at 12°S, covering anoxic and hypoxic bottom water conditions. Benthic N₂ fixation was detected at all sites using the acetylene reduction assay, with high rates measured in OMZ mid-waters between the 70 m and 253 m and lowest lower N2 fixation rates at greater depthbelow 253 m down to 1025 m water depth. SR rates were decreasing decreased with increasing water depth., with highest rates at the shallow site. Benthic N₂ fixation and SR depth profiles in sediments showed similar qualitative trends largely overlapped with SR depth profiles, suggesting a coupling of that both processes are coupled. The pPotential of benthic N₂ fixation by SR-sulfate-reducing bacteria was verified by the molecular analysis of nifH genes. Detected nifH sequences, i.e., the key functional gene for N2 fixation, encoding for the nitrogenase enzyme, clustered with sulfate-reducing SR-bacteria that have been demonstrated to fix N2 in other benthic environments. Depthintegrated rates of benthic N2 fixation and SR showed no direct correlation along the 12°S transect, suggesting that the benthic diazotrophs in the Peruvian OMZ are being is controlled by additional various environmental factors such as. The organic matter and free sulfideavailability and the presence of sulfide appear to be major drivers for benthic diazotrophy. It was further found that N2 fixation in OMZ sediments was not inhibited by

high ammonium concentrations. N_2 fixation rates in OMZ sediments were similar to rates measured in other organic-rich sediments. Overall, this work_study_improves our knowledge on fixed_N sources and_in_marine_sediments_and_contributes_to_a_better understanding of N cycling in OMZ sediments_oxygen deficient environments.

1. Introduction

Only 6 % of nitrogen (N) in seawater is bioavailable (Gruber, 2008). This bioavailable N is mainly present in the form of nitrate (NO_3^-) , whereas the large pool of available atmospheric dinitrogen gas (N_2) is only available for N_2 fixing microorganisms (diazotrophs). Therefore, N is often controlling limits the marine productivity (Ward & Bronk, 2001; Gruber, 2008) and the largest this limitation makes N_2 fixation the dominant source of bioavailable N (i.e. ammonium (NH_4^+)) in the marine environment is N_2 fixation (Falkowski

et al., 1998; Strous et al., 1999; Brandes & Devol, 2002).

To date, the quantitative contribution of diazotrophs in the marine N cycle remains unclear and numerous estimates of global sources and sinks of global N have exist, lead ing to an unbalanced budget with deficits of around 200 Tg N yr⁻¹ (Gruber, 2004; Brandes et al., 2007; Capone & Knapp, 2007; Codispoti, 2007). In most studies, oceanic N sinks are either estimated to be higher than oceanic N sources, suggesting that This suggests that either previous determination of N₂ fixation rates determinations have been underestimated (Montoya et al., 1996; Codispoti, 2007) (Großkopf et al., 2012) or that N loss processes are overestimated (Codispoti, 2007). But also almost b However, also balanced budgets such as exist that calculated ~265 Tg N yr⁻¹ for N sources and ~275 Tg N yr⁻¹ for N sinks exist (Gruber, 2004). These Bbudget discrepancies illustrate that the current knowledge on diazotrophys and the marine N cycle is still limited.

Latest Recent investigations argue that N₂ fixation in the water column cannot be totally attributed to phototrophic cyanobacteria, but that also heterotrophic prokaryotes contribute—a substantially part—(Riemann et al., 2010; Farnelid et al., 2011; Dekaezemacker et al., 2013; Löscher et al., 2014; Fernandez et al., 2015)—similar to marine benthic habitats. This relation—was shown for the Peruvian oxygen minimum zone (OMZ), where proteobacterial clades were—dominateding—and_with_heterotrophic_diazotrophs—mainly

occurred, indicating that cyanobacterial diazotrophs are of minor importance in this area (Löscher et al., 2014).

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Pelagic N₂ fixation has been studied mostly in the oligotrophic surface oceans, but it was not until the past decade that also benthic habitats began to received more attention (Fulweiler et al., 2007; Bertics et al., 2010; Bertics et al. 2013). Most studies on benthic N₂ fixation focused on coastal environments (Capone et al., 2008 and references therein). For example, subtidal sediments in Narragansett Bay (Rhode Island) were found to switch from being a net sink in the form of denitrification to being a net source of bioavailable N by N2 fixation, caused by a decrease of organic matter deposition to the sediments (Fulweiler et al., 2007). Shallow brackish-water sediments off the Swedish coast revealed benthic N2 fixation along with a diverse diazotrophic community (Andersson et al., 2014). N2 fixation The nitrogenase activity was positively influenced by a variety of environmental factors, such as salinity and dissolved inorganic nitrogen, while wave exposure had a negative influence. Recent work revealed that benthic N2 fixation is often linked to sulfatereducing (SR) bacteria., e.g., For instance, bioturbated coastal sediments showed enhanced N₂ fixation activity mediated by sulfate-reducing SR—bacteria, adding new dissolved inorganic N to the system (Bertics et al., 2010; Bertics & Ziebis, 2010). Further coupling of N₂ fixation to SR was found observed in organic-rich sediments of the seasonal hypoxic Eckernförde Bay (Baltic Sea) (Bertics et al., 2013), as well as in the sub-tidal, heterotrophic sediments of Narragansett Bay (Rhode Island, USA) (Fulweiler et al., 2013). Several sulfatereducing SR-bacteria carry the functional gene marker for N₂ fixation, the nifH gene for encoding the nitrogenase enzyme (Sisler & ZoBell, 1951; Riederer-Henderson & Wilson, 1970; Zehr & Turner, 2001) and were shown to actively fix N2 in culture experiments (Riederer-Henderson & Wilson, 1970). Therefore, we need to better understand SR bacteria and their potential to fix N in the environment. However, information on sulfatereducing bacteria and their contribution to N2 fixation in the environment is rather sparse and makes this one of the remaining questions to be solved.

So far, the distribution of benthic N_2 fixation and its relevance for N cycling in the Peruvian (OMZ), defined by dissolved oxygen < 20 μ mol kg⁻¹ (Fuenzalida et al., 2009)_z are unknown. The shelf and the upper slope in the Peruvian OMZ represent recycling sites of dissolved inorganic N with dissimilatory NO_3 reduction to NH_4 being the dominant process driving in

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the benthic N cycle (Bohlen et al., 2011). This process is mediated by the filamentous sulfide-oxidizing *Thioploca* bacteria (Schulz, 1999; Schulz & Jørgensen, 2001). Along with dissimilatory NO_3^- reduction to NH_4^+ , also-benthic denitrification by foraminifera between 80 and 250 m water depth occurs in the Peruvian OMZ (Glock et al., 2013). These authors calculated a potential NO_3^- flux rate of 0.01 to 1.3 mmol m^{-2} d⁻¹ via this pathway and suggested that foraminifera could be responsible for most of the benthic denitrification.

The high input of labile organic carbon to the Peruvian OMZ sediments (Dale et al., 2015) and subsequent SR should support favor benthic N₂ fixation. Sulfate-reducing SR bacteria could considerably contribute to N₂ fixation in these organic-rich OMZ sediments, given that several sulfate-reducing SR bacteria (e.g. Desulfovibrio spp. (Riederer-Henderson & Wilson, 1970; Muyzer & Stams, 2008)) carry the genetic ability to fix N₂, and provide an important bioavailable N source for non-diazotrophic organisms (Bertics et al., 2010; Sohm et al., 2011; Fulweiler et al., 2013). We therefore hypothesize a possible coupling of N₂ fixation and SR in sediments off Peru. The aim of the present study was the to identifyication and quantifyication of benthic N₂ fixation along a depth transect through the Peruvian OMZ, together with potentially coupled SR. Additionally, the identification of bacteria facilitating these processes will help to understand should shed light into the diazotrophic community structure of inhabiting these sediments. The overall knowledge gained is useful will be used to better constrain benthic N cycling in OMZs and to improve our knowledge on sources and sinks of fixed N.

2. Materials and Methods

2.1 Study area

The most extensive OMZ worldwide developed is found in the eastern tropical south Pacific ocean at the Central Peruvian coast (Kamykowski & Zentara, 1990). The Peruvian OMZ ranges between 50 m and 700 m water depth with oxygen (O_2) concentrations below the detection limit in the mid-waters (Stramma et al., 2008). The mean water depth of the upper OMZ boundary deepens during intense El Niño Southern Oscillation years and can reach a depth of 200 m (Levin et al., 2002) with oxygenation episodes reaching concentrations of up to 100 μ M O_2 (Gutiérrez et al., 2008). O_2 concentrations (Fig. 1, Tab. 1) off Peru are affected modulated by coastal trapped waves (Gutiérrez et al., 2008), trade

winds (Deutsch et al., 2014) or and subtropical-tropical cells (Duteil et al., 2014), and can vary on monthly to interannual time-scales (Gutiérrez et al., 2008).

At 12°S, the OMZ extends from water depths between 50 and 550 m (Dale et al., 2015) (Fig. 1). During our field work, Bbottom water O₂ concentrations varied greatly with water depth and were below the detection limit (5 μM) at stations from 70 m to 407 m water depth. Bottom water O₂ increased from to 19 μM at 770 m water depth to and 53 μM at 1025 m water depth, indicating the lower boundary of the OMZ (Dale et al. 2015). Between 70 m and 300 m water depth, the sediment surface was colonized by dense filamentous mats of sulfur-oxidizing bacteria, presumably of the genera Mari Tthioploca spp (Gutiérrez et al., 2008; Mosch et al., 2012). This These bacteria are able to glide up to 1 cm h⁻¹ through the sediment in order to feedaccess on hydrogen sulfide (Fossing et al., 1995; Jørgensen & Gallardo, 1999; Schulz, 1999). Sediments at the lower boundary (770 m and 1025 m) of the OMZ were shown to have host a variety of macrofaunal organisms e.g. ophiuroids, gastropods, and crustaceans (Mosch et al., 2012).

The 12°S region is in the center of <u>an</u> extensive upwelling <u>zone</u> and features high primary productivity (Pennington et al., 2006). Sediments at 12°S have higher rates of particulate organic carbon <u>accumulation</u> (2-5 times) compared to other continental margins and a high carbon burial efficiency <u>at deep stations</u>, indicating <u>high preferential</u> preservation of organic matter in <u>sediments below</u> the Peruvian OMZ (Dale et al., 2015). The shelf (74 m) of the Peruvian OMZ is characterized by high sediment<u>ation accumulation</u> rates of 0.45 cm yr⁻¹, while <u>mid-waters and below the OMZ show</u> rates between 0.07 and 0.011 cm yr⁻¹ <u>were found in OMZ mid-waters and below the OMZ, additionally. <u>s</u>Sediment porosity was high at the shelf stations and in OMZ mid-waters (0.96 – 0.9) and was lowest (0.74) at the deepest 1024 m station (Dale et al., 2015).</u>

2.2 Sampling

Sediment samples were taken in January 2013, at six stations (70, 144, 253, 407, 770, and 1025 m) at 12°S along a depth transect at 12°S in the OMZ off Peru (Fig. 1) during an expedition on RV Meteor (M92). January represents austral summer, i.e. the low upwelling season in this area (Kessler, 2006). Samples were retrieved using a TV-guided multiple corer (MUC) equipped with seven core liners. The core liners had a length of 60 cm and an inner diameter of 10 cm. Location, water depth, temperature, and O₂ concentration (from

Dale et al. 2015) at the six sampling stations are listed in Table 1. Retrieved cores for microbial rate measurements were immediately transferred to cold rooms (4-9 °C) for further processing.

2.3 Geochemical analyses

Porewater analysis and the determination of sediment properties and geochemical data have been previously described in detail by Dale et al. (2015). In short, the first core was subsampled under anoxic conditions using an argon-filled glove bag, to preserve redox sensitive constituents. NH_4^+ and sulfide concentrations were analyzed on a Hitachi U2800 UV/VIS spectrophotometer using standard photometric procedures (Grasshoff et al., 1999), while sulfate (SO_4^{2-}) concentrations were determined by ion chromatography (Methrom 761).

The second replicate core was sampled to determine porosity by the weight difference of the fresh sediment subsamples before and after freeze-drying. The pParticulate organic carbon and particulate organic nitrogen contents were analyzed using a Carlo-Erba element analyzer (NA 1500).

2.4 Benthic nitrogenase activity-nitrogen fixation

At each of the six stations, one MUC core was sliced in a coldrefrigerated container (9°C) in 1-cm intervals from 0-6 cm, in 2-cm intervals from 6-10 cm, and in 5-cm intervals from 10-20 cm. The acetylene reduction assay (Capone, 1993; Bertics et al. 2013) was applied, to quantify nitrogenase activity (NA). This application is an *ex situ* method, based on the reduction of acetylene (C_2H_2) to ethylene (C_2H_4) by the nitrogenase enzyme, which reduces other small triple bond molecules, like acetylene (Lockshin & Burris 1965; Dilworth, 1966). The temporal increase of C_2H_4 in samples can be measured by flame ionization gas chromatography (Hardy et al. 1968; Stewart et al. 1967). Thereby, the amount of C_2H_2 reduced to C_2H_4 serves as an indication for N_2 fixation rates. To convert from nitrogenase activity to N_2 fixation, a conversion factor of 3 C_2H_4 :1 N_2 was applied (Patriquin & Knowles, 1972; Donohue et al., 1991; Orcutt et al., 2001; Capone et al., 2005)—was applied, which was previously used to measure N_2 fixation in sediments (Welsh et al., 1996; Bertics et al., 2013).

Serum vials (60 mL) were flushed with N_2 and filled with 10 cm³ sediment from each sampling depth (triplicates). The samples were flushed again with N_2 , crimp sealed with

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butyl stoppers and injected with 5 mL of C_2H_2 to saturate the nitrogenase enzyme. Serum vials were stored in the dark and at 9 °C, which reflected the average *in situ* temperature along the transect (compare with Tab. 1). Two sets of triplicate controls (10 cm³) were processed for every station. Sediment was collected from each core liner from 0-5 cm, 5-10 cm, and from 10-20 cm and placed in 60 mL serum vials. One set of controls was used to identify natural C_2H_4 production; without the injection of acetylene, and the second control set was fixed with 1 mL formalin (37.5%)—formaldehyde solution.

The increase of C_2H_4 in each sediment slice was measured onboard over one week (in total 5 time points, including time zero) using gas chromatography (Hewlett Packard 6890 Series II). From each serum vial, a 100 μ l headspace sample was injected into the gas chromatograph and the results were analyzed with the HP ChemStation gas chromatograph software. The gas chromatograph was equipped with a packed column (Haye SepT, 6 ft, 3.1 mm ID, Resteck) and a flame ionization detector. The carrier gas was helium and the combustion gases were synthetic air (20 % O_2 in O_2) and hydrogen. The

Sediment depth profiles were expressed in NA. To convert from NA to N₂ fixation, a conversion factor of 3 C₂H₄:1 N₂ for the integrated rates was applied. This conversion factor is based on comparisons between the C₂H₂ reduction assay and ¹⁵N incubations (Patriquin & Knowles, 1972; Donohue et al., 1991; Orcutt et al., 2001; Capone et al., 2005) and was previously used to measure N₂ fixation in sediments (Welsh et al., 1996; Bertics et al., 2013). Standard deviation for depth profiles was calculated from three replicates per sediment depth and error bars for standard deviation of for integrated N₂ fixation were calculated from three integrated rates per station.

column had a temperature of 75°C and the detector temperature was 160°C.

2.5 Sulfate reduction rates

 One MUC core per station was used for determination of SR activity. First, two replicate push cores (length 30 cm, inner diameter 2.6 cm) were subsampled from one MUC core. The actual push core length varied from 21 - 25 cm total length. Then, 6 μ l of the carrier-free $^{35}SO_4^{2-}$ radio tracer (dissolved in water, 150 kBq, specific activity 37 TBq mmol⁻¹) was injected into the replicate push cores in 1-cm depth intervals according to the whole-core injection method (Jørgensen, 1978). The push cores were incubated for ~12h at 9°C. After incubation, bacterial activity was stopped by slicing the push core into 1-cm intervals and

Feldfunktion geändert Feldfunktion geändert transferring each sediment layer into 50 mL plastic centrifuge tubes filled with 20 mL zinc acetate (20% w/w). Controls were done in triplicates from different depths and first fixed with zinc acetate before adding the tracer. Rates for SR were determined using the cold chromium distillation procedure according to Kallmeyer et al. (2004).

It should be mentioned that the yielded SR rates have to be treated with caution due to long (up to 3 half-life times of 35 S) and unfrozen storage. Storage of SR samples without freezing has recently been shown to result in the re-oxidation of 35 S-sulfides (Røy et al., 2014). In this reaction, FeS is converted to ZnS. The released Fe²⁺ reacts with O₂ and forms reactive Fe(III). The Fe(III) oxidizes ZnS and FeS, which are the major components of the total reduced inorganic sulfur species, resulting in the generation of SO_4^{2-} and hence an underestimation of SR rates. However, because all SR samples in the present study were treated the same way, we trust the relative distribution of activity along sediment depth profiles and recognize potential underestimation of absolute rates.

2.6 nifH gene analysis

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Core samples for DNA analysis were retrieved from the six stations and were sliced in the same sampling scheme as described for the NAbenthic N2 fixation. Approximately 5 mL sediment from each depth horizon was transferred to plastic whirl-paks® (Nasco, Fort Atkinson, USA), frozen at -20 °C and transported back to the home laboratory. To check for the presence of the nifH gene, DNA was extracted using the FastDNA® SPIN Kit for Soil (MP Biomedicals, CA, USA) following the manufacturer's instructions with a small modification. Sample homogenization was done in a Mini-BeadbeaterTM (Biospec Products, Bartlesville, USA) for 15 seconds. PCR amplification, including primers and PCR conditions, was done as described by Zehr et al. (1998), using the GoTaq kit (Promega, Fitchburg, USA) and additionally 1 µL bovine serum albumin BSA (20 mg mL⁻¹ (Fermentas)). The TopoTA Cloning® Kit (Invitrogen, Carlsbad, USA) was used for cloning of PCR amplicons, according to the manufacturer's protocol. Sanger sequencing (122 nifH sequences) was performed by the Institute of Clinical Molecular Biology, Kiel, Germany- For the sampling sites 70 m, 144 m, 253 m, 407 m, 770 m, and 1025 m water depth the number of obtained sequences was 22, 24, 24, 13, 18, and 21, respectively. Negative controls were performed using the PCR mixture as described without template DNA; no amplification was detected. Sequences were ClustalW aligned in MEGA 6.0 (Tamura et al., 2007), and a maximum likelihood tree

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was constructed on a 321 bp-base pair fragment and visualized in iTOL (Letunic & Bork, 2007, 2011). Reference sequences were obtained using BlastX on the NCBI database. (Sequence submission being in Progress). Sequences were submitted to Genbank (Accession numbers: KU302519 - KU302594).

3. Results

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3.1 Sediment properties

Although sediment description and porewater sampling was done down to the bottom of the core, the focus here is on sediments from 0-20 cm where-NA_benthic N₂ fixation was investigated.

Sediments at the shelf station (St.) 1 (70 m) were black between 0 - 1 cm and then olive green until 20 cm. Only a few metazoans (polychaetes) were observed in the surface sediment. The sediment surface was colonized by dense filamentous mats of sulfuroxidizing Mari Tthioploca spp. (Gutiérrez et al., 2008; Mosch et al., 2012). These bacteria reached down to a sediment depth of 36 cm in the sediment cores. The sediment at on the outer shelf St. 4 (144 m) was dark olive green from 0 - 13 cm and dark grey until 20 cm. At the sediment surface and in MUC cores, Thioploca spp. was visible. At St. 6 (253 m), which was within the OMZ <u>core</u>, sediment appeared dark olive green between 0-17 cm and olive green with white patches between 17 - 20 cm. At this station, Mari#thioploca spp. was abundant. Uniquely, surface sediments (0 - 3 cm) at St. 8 (407 m), consisted of a fluffy, dark olive-green layer mixed with white foraminiferal ooze. This layer also contained cmsized phosphorite nodules with several perforations (ca. 1 - 3 mm in diameter). Below 2 cm, the sediment consisted of a dark olive green, sticky clay layer. No Thioploca mats were found at St. 8. The St. 9 (770 m) was below the OMZ. Sediments were brown to dark olive green with white dots particles between 0 - 12 cm and appeared brown to olive green without white dots particles below this depth. Organisms such as anemones, copepods, shrimps and various mussels were visible with the TV-guided MUC and in sediment cores. The deepest St. $\frac{10}{(10)}$ 1025 m) had dark olive green sediment from 0 – 20 cm and black patches from 17 - 20 cm. The sediment was slightly sandy and was colonized with polychaete tubes at the surface and organisms that were also present at St. 9. For further sediment core descriptions see also Dale et al. (2015).

Geochemical porewater profiles of NH₄⁺, SO₄², sulfide, organic carbon content, and organic 608 C/N ratio between 0 – 20 cm of at the six stations are shown in Fig 2. In all cores, NH_4^+ 609 concentrations increased with sediment depth. The highest NH_4 concentration was 610 reached at St. 1 (70 m), increasing from 316 µM in the upper cmat the sediment surface to 611 2022 μM at 20 cm. The St. 4 and 6 showed intermediate NH₄⁺ concentrations between 300 612 μ M and 800 μ M at 20 cm, respectively. At St. 8 (407 m) the NH_4^+ concentration increased 613 from 0.7 μM in at the surface to 107 μM at 20 cm. The two deep stations (St. 9 and 10) had 614 the lowest NH_4^+ concentrations with 33 μ M and 22 μ M at 20 m sediment depth, 615 616 respectively. The SO_4^{2-} concentrations remained relatively constant in the surface sediments of along the 617 transect. Only at the shallowest-St. 1, a decrease from 28.7 µM in the surface layer to 19.4 618 μM at 20 cm was observed. Along with the decrease in SO_4^{2-} , only St. 1 revealed 619

considerable porewater sulfide buildupaccumulation. Sulfide increased from 280 µM in at the surface sediment to 1229 μM at 20 cm.

Organic carbon content decreased with increasing sediment depth at St. 1 (70 m), 9 (770 m), and 10 (1025 m). The highest surface organic carbon content (~15 wt%) was found at St. 6, whereas Tthe lowest surface organic carbon content (~2.6 wt%) was detected at the deep St. 10. The average (0 - 20 cm) organic carbon value (Fig. 5) increased from St. 1 to St. 6 (15 \pm 1.7 wt%) and decreased from St. 6 to the lowest value at St. 10 (2.4 \pm 0.4 wt%). C/N ratios, as a proxy for the freshness of the organic matter, increased with increasing sediment depth (Fig. 5). The lowest benthic-surface C/N ratio (6.2) was measured at the

3.2 Benthic nitrogen fixation and sulfate reduction (SR)

for the estimation of integrated rates (0 - 20 cm).

shallow St. 1, while the highest surface C/N ratio (11) was found at St. 10.

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For an straighforwardeasy comparison of SR rates with benthic N₂ fixation NA-only the sediment depths between 0 – 20 cm are considered. Sediment depth profiles are expressed as in nitrogenase activity (NA)N2 fixation, i.e. that is, without the conversion factor of 3 C₂H₄:1 N₂ to achieve actual N₂ fixation rates. The conversion to N₂ fixation was applied only

Highest N₂ fixation NA-and SR rates were detected in the surface sediments (0 – 5 cm) and both rates tended to decrease with increasing sediment depth (Fig. 3). While N₂ fixation NA and SR rates were high at the shallower stations St. 1, 4, and 6 (70 m, 144 m, 253 m) and,

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NA and SR rates were lowest and lowest at the three-deeper stations St. 8 - 10 (407 m, 770
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        m, 1025m).
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        At St. 1, N2 fixation NA- and SR rates showed different trends in the top layer of the cores,
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         but depth profiles were more aligned below. While Although St. 1 had the highest SR rates
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        of all sites, reaching 248 nmol SO_4^{2-} cm<sup>-3</sup> d<sup>-1</sup> at 0 – 1 cm, N_2 fixation NA-was not highest at
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        this station. Only St. 1 had considerablely porewater sulfide concentrations and a decrease
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        of SO_4^{2-} concentration with increasing sediment depth, as well as the highest NH_4^+
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        concentrations throughout the core.
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         At St. 4 (144 m), both N<sub>2</sub> fixation NA-and SR revealed peaks close to the surface. N<sub>2</sub> fixation
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         NA-decreased from 3.5 \pm 0.6 nmol C<sub>2</sub>H<sub>4</sub>-cm<sup>-3</sup>-d<sup>-1</sup> to 0.9 \pm 0.08 nmol C<sub>2</sub>H<sub>4</sub>-cm<sup>-3</sup>-d<sup>-1</sup>-between
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        0-8 cm and increased below 8 cm<sub>.</sub>, reaching 2.2 ± 1.2 nmol C_2H_4-cm<sup>-3</sup> d<sup>-1</sup> at 20 cm. This
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        increase was not observed in SR rates, which were highest in the surface (181 nmol SO<sub>4</sub><sup>2-</sup>
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        cm<sup>-3</sup> d<sup>-1</sup>) and decreaseding towards the bottom of the core. St. 6 (253 m) had the highest
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        N_2 fixation NA of all stations. After decreasing from 6.6 \pm 0., with 7 nmol C_2H_4 rates of 4.0 \pm
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        0.5 nmol N<sub>2</sub> cm<sup>-3</sup> d<sup>-1</sup> in the surface cm to 1.7 \pm 0.2 nmol C<sub>2</sub>H<sub>4</sub>-cm<sup>-3</sup> d<sup>-1</sup> in 6 - 8 cm, NA
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         increased to 2.5 ± 2.2 nmol C<sub>2</sub>H<sub>4</sub> cm<sup>-3</sup> d<sup>-1</sup> with a peak at 10 - 15 cm. Although N<sub>2</sub> fixation
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        NA-and SR had corresponding depth profiles, the highest SR rate of all stations was not
         detected at St. 6-(18 nmol SO<sub>4</sub> 2-cm<sup>-2</sup>-d<sup>-1</sup>). Very low N<sub>2</sub> fixation NA-rates were measured at
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        St. 8 (407 m) (0.<del>77-5</del> \pm 0.<del>37-25</del> nmol \frac{\text{CN}_2\text{H}_4}{\text{cm}^{-3}} d<sup>-1</sup> in the surface), as well as very low SR
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        rates (0 - 4.3 nmol SO_4^{2-} cm<sup>-3</sup> d<sup>-1</sup>). This station was unique due to the presence of
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        foraminiferal ooze, phosphorite nodules and a sticky clay layer below 2 cm. Here, NA was
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         extremely low below 2 cm, not exceeding 0.09 ± 0.04 nmol C<sub>2</sub>H<sub>4</sub> cm<sup>-3</sup> ·d<sup>-1</sup>. The N<sub>2</sub> fixation
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        NA and SR rates showed a peak at 5 cm and at 7 cm, respectively. At St. 9 (770 m) N<sub>2</sub>
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         fixation NA-was low in the surface and at 20 cm sediment depth, with a peak in activity at 4
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         -5 cm (\frac{1.20.8}{2.00} \pm 0.0812 nmol \frac{6}{10} nmol \frac{6}{10} cm<sup>-3</sup> d<sup>-1</sup>). At St. 10 (1025 m), \frac{8}{10} fixation NA-rates were
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        low throughout the sediment core, not exceeding ranging between 0.23-16 ± 0.023 nmol
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         \text{CN}_2\text{H}_4 \text{ cm}^{-3} \text{ d}^{-1}. in surface sediments and 0.06 \pm 0.01 nmol \text{C}_2\text{H}_4 \text{ cm}^{-3} \text{ d}^{-1} in 10 - 15 cm. In
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         accordance with this observation, tThis site had the lowest organic carbon content
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throughout the core (between 2.6 wt% at the surface and 1.9 wt% at 20 cm), as well as low

 NH_4^+ concentrations. At St. 9 (below 9 cm depth) and St. 10 (entire core) SR rates were

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below detection, which could point either to the absence of SR or to the complete loss of total reduced inorganic sulfur due to the long, unfrozen storage (see methods).

Integrated N₂ fixation (0 – 20 cm) increased from St. 1 to St. 6, with the highest rate (0.4 \pm $0.06 \text{ N}_2 \text{ m}^{-2} \text{ d}^{-1}$) at St. 6 (253 m), and decreased from St. 6 (407 m) to St. 10 (1025 m) (Fig. 4). Integrating SR rates over 0 to 20 cm sediment depth, Integrated SR rates (0 to 20 cm) ranged from ~4.6 mmol SO_4^{2-} m⁻² d⁻¹ at St. 1 to below detection $\frac{1}{2}$ m $\frac{1}{2}$ m $\frac{1}{2}$ at St. 9 (Fig. 4). Overall, integrated SR rates decreased with increasing water depth. Integrated N₂ fixation rates and SR were almost in general inversely correlated between St. 1 and St. 6, and - Overall, N₂ fixation rates followed the organic carbon content from St. 1 to St. 6 (70 – 253 m) (Fig. 5). Both parameters had the highest value at St. 6. This pattern did not hold was not conform with for the relatively lower integrated SR rate at St. 6. The C/N ratio, averaged over 20 cm, increased with increasing water depth (Fig. 5). Regarding the three deep stations, the lowest integrated N_2 fixation rate (0.008 \pm 0.002 N_2 m⁻² d⁻¹) was detected at St. 8 (407 m). Also the integrated SR rate was low at this site (~0.46 mmol SO₄²⁻ m⁻² d⁻¹). At St. 9 and 10 (770 and 1025 m), integrated N₂ fixation had low rates of was low at $0.05 \pm 0.005 \, \text{N}_2 \, \text{m}^{-2} \, \text{d}^{-1}$ and $0.01 \pm 0.001 \, \text{N}_2 \, \text{m}^{-2} \, \text{d}^{-1}$, respectively and also integrated SR rates were also lowest at St. 9 (770 m). From St. 8 to 10 a decrease of integrated N₂ fixation and SR together with the average organic carbon content was detected.

 $\underline{\text{No activity was detected } Inin} \text{ controls for } N_2 \text{ fixation and } SR \underline{\text{ no activity was detected}}.$

3.3 Molecular analysis of the nifH gene

NifH gene sequences were detected at all six sampling sites and clustered with Cluster I proteobacterial sequences and Cluster III sequences as defined by Zehr & Turner (2001) (Fig. 6). In Cluster I and Cluster III, three novel clades and seven novel clades were detected, respectively. In general, most of the novel previously unidentified clades belong to uncultured bacteria. One distinct novel clade was found for the St. 1 – 6. Furthermore, several clades consisting of different stations were found. No Cluster I cyanobacterial nifH sequences were detected and no potential PCR contaminants were present (Turk et al., 2011). In this study, detected sequences clustered with sulfate-reducing SR-bacteria, such as Desulfovibrio vulgaris (Riederer-Henderson & Wilson, 1970; Muyzer & Stams, 2008) and Desulfonema limicola (Fukui et al., 1999). One cluster (OMZ 144 m) belonged was closely related to Vibrio diazotrophicus (Guerinot et al., 1982), which has the unique property for

a known Vibrio species to perform N₂ fixation and -which was found previously in the water column of the OMZ off Peru (P7 M773 28) (Löscher et al., 2014). The other organisms with which OMZ sequences clustered belonged to the genera of bacteria using fermentation, namely Clostridium beijerincki (Chen, 2005), and to the genera of iron-reducing bacteria, namely Geobacter bemidjiensis (Nevin et al., 2005). In addition, several sequences were phylogenetically related to an uncultured bacterium from the Eastern Tropical South Pacific (KF151591.1) and a gamma proteobacterium (Zehr & Turner, 2001) (TAS801) from the Pacific Ocean (AY896428.1).

4. Discussion

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4.1 Coupling of benthic nitrogen fixation and sulfate reduction

Based on the high organic matter input to Peruvian sediments underneath the OMZ we hypothesized a presence of N2 fixation and it's coupling to sulfate reduction (SR). We confirmed the presence of N2 fixation NA-in sediments at all sampled stations along the depth transect between 70 and 1025 m water depth. This activity was generally enhanced, where SR peaked and sometimes both activity depth profiles revealed similar trends. However, while peaks in SR where very pronounced, maximum N2 fixation NA-showed a much broader distribution over depth. This discrepancy indicates that N2 fixation might be partly coupled to processes other than SR (see nifH discussion below). But it should be kept in mind that the N₂ fixation NA-and SR were determined in replicate MUC cores, which had a sampling distance of up to 50 cm, depending on the locationwhere of the cores liners were situated in the instrumentmultiple corer. Nonetheless, it appears that Ithe observed N₂ fixation NA-is therefore not directly fuelled by the observed SR activity. Trends might vary naturally. We are also aware of potential microbial community shifts driven by the addition of C_2H_2 (Fulweiler et al., 2015). However, a community shift would be expected to cause rather an underestimation of absolute N2 fixation rates. Further, incubation with acetylene can lead to a potential lack of fixed N; however, to the best of our knowledge this is the standard method used for the determination of N2 fixation in sediments (Bertics et al., 2013).

The more surprising finding is that integrated rates of N_2 fixation NA and SR showed opposite trends at the three shallowest stations, pointing to potential environmental control mechanisms (see 5.2).

The coupling between N₂ fixation and SR has been previously suggested for coastal sediments off California, where (Bertics & Ziebis, 2010). In this study N₂ fixation significantly decreased when SR was inhibited (Bertics & Ziebis, 2010). Different studies confirmed that sulfate-reducing SR bacteria, such as D-esulfovibrio vulgaris can supply organic-rich marine sediments with bioavailable N through N₂ fixation (Welsh et al., 1996; Nielsen et al., 2001; Steppe & Paerl, 2002; Fulweiler et al., 2007; Bertics et al., 2013; Fulweiler et al., 2013). –Fulweiler et al. (2013) conducted a study in sediments of the Narrangaset Bay and found several nifH genes related to sulfate-reducing SR bacteria, such as Desulfovibrio spp., Desulfobacter spp. and Desulfonema spp., suggesting that sulfate-reducing SR bacteria are were the dominant diazotrophs.

The *nifH* gene sequences obtained in our study strongly indicated the genetic capability of sulfate reducers in the Peruvian sediments to conduct N₂ fixation. They clustered with the <u>sulfate-reducing SR-</u>bacteria *Desulfovibrio vulgaris*, which is a confirmed diazotroph (Sisler & ZoBell 1951; Riederer-Henderson & Wilson 1970), as well as *Vibrio diazotrophicus*, which recently clustered with sequences from the Peruvian OMZ water column (Fernandez et al., 2011; Löscher et al., 2014). Sequences taken from the seasonally hypoxic Eckernförde Bay in the Baltic Sea also clustered with *Desulfovibrio- vulgaris* (Bertics et al., 2013), suggesting a major involvement of <u>SR-sulfate-reducing</u> bacteria in N₂ fixation in organic-rich sediments underlying OMZs. Interestingly, we detected several new *nifH* gene clusters in the Peruvian OMZ that have not been identified yet (Fig. 6). These findings suggest certain diversity among the benthic diazotrophic community and a possible coupling of N₂ fixation also to processes other than SR, which might explain some of the discrepancies between the two activities (see above). These results add to the growing evidence that "heterotrophic" N₂ fixation is dominant in the Peruvian OMZ (Farnelid et al., 2011; Fernandez et al., 2011; Löscher et al., 2014).

The molecular analysis further indicates that not all of the benthic diazotrophs are known sulfate-reducing SR—organisms. Therefore, a possible—coupling of N₂ fixation also to processes other than SR is likelypossible, which might explain some of the discrepancies

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between N₂ fixation and SR_{the two} activityies (see above). Other relevant processes may include the usage of reduced carbon compounds as previously suggested for diazotrophic

organisms in the water column of the Peruvian OMZ Dekaezemacker et al., 2013; Löscher

762 et al., 2014).

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4.2 Environmental factors potentially controlling benthic N₂ fixation

The observed differences between integrated N_2 fixation and SR along the depth transect indicate potential environmental factors that are—controlling the extent of benthic N_2 fixation, which will be discussed in the following section.

4.2.1 Organic matter quantity and quality

A major driver for microbial processes such as SR and "heterotrophic" N₂ fixation by potentially heterotrophic organism is the availability of the organic material (Jørgensen, 1983; Howarth et al., 1988; Fulweiler et al., 2007). Integrated N2 fixation and average organic carbon content correlated showed similar trends along the Peruvian OMZ depth transect (Fig. 5). Thus, organic matter availability appears to be a major factor controlling N₂ fixation at this study site. Low N₂ fixation rates were previously shown to be related to low organic matter content in slope sediments in the Atlantic Ocean (Hartwig & Stanley, 1978). This pattern is supported by the study of Bertics et al. (2010), which showed that burrow systems of the bioturbating ghost shrimp Neotrypaea californiensis can lead to enhanced organic matter availability in deeper sediment layers, resulting in high rates of N2 fixation. However, high organic matter availability does not always result in enhanced N2 fixation rates. Subtidal sediments in the Narragansett Bay were found to switch from being a net sink via denitrification to being a net source of bioavailable N via N2 fixation (Fulweiler et al., 2007). This switch from N sink to N source was caused by a decrease of organic matter deposition to the sediments, which was in turn triggered by low primary production in the surface waters. Especially this switch is an interesting feature, showing us that there are still major gaps in our understanding of benthic N₂ fixation.

Besides quantity also the quality of organic matter in sediments is a major factor influencing microbial degradation processes (Westrich & Berner, 1984). In the Peruvian OMZ sediments, the average C/N ratio increased with water depth indicating that the shallow stations received a higher input of fresh, labile organic material compared to the deeper stations. Similar trends were reported for a different depth transect off Peru (Levin

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et al., 2002). However, an increase of the C/N ratio with depth would suggest highest integrated N₂ fixation rate at the shallowest St. 1 (70 m), which however is not in line with our observation that shows an increase in rate from St. 1 (70) to St. 6 (253 m) (Fig. 5). Similarly, DIC fluxes measured using benthic chambers at the same stations can be used as an indicator for organic matter degradation rates were at the same stations during the expedition by (Dale et al., (2015). The DIC flux did not correlate with integrated N₂ fixation rates, but instead roughly followed the pattern of SR rates along water depth (Fig. 45). The highest integrated SR rate and DIC flux was were found at St. 1 (70 m), whereas the lowest integrated SR rate and DIC flux was found occurred at St. 10 (1025 m). Assuming that SR is largely responsible for organic matter remineralization in i.e. DIC fluxes, in the sediments below the OMZ (Bohlen et al., 2011; Dale et al. 2015), the difference between integrated SR and DIC flux is expected to mainly represent the underestimated fraction, which likely resulted from the the long duration of unfrozen storage of the samples (see methods).

4.2.2 Ammonium

Interestingly, the highest N₂ fixation was measured in sediments colonized by the sulfuroxidizing and nitrate-reducing filamentous bacteria Mari Thioploca spp. (Schulz, 1999; Schulz & Jørgensen, 2001; Gutiérrez et al., 2008; Salman et al., 2011; Mosch et al., 2012). <u>Mari</u> $_{t}$ <u>hioploca</u> facilitates dissimilatory NO₃ reduction to NH₄, which preserves fixed N in the form of NH₄⁺ in the environment (Kartal et al., 2007). OMZ sediments off Peru are generally rich in NH₄⁺ (Bohlen et al., 2011). This co-occurrence of Thioploca Marithioploca and N₂ fixation was puzzling since high concentrations of NH₄⁺, could inhibit N₂ fixation (Postgate, 1982; Capone, 1988; Knapp, 2012). It remains questionable why microorganisms should fix N₂ in marine sediments, when reduced N species are abundant. Some doubt remains as to the critical NH₄⁺ concentration that inhibits N₂ fixation and whether the inhibitory effect is the same for all environments (Knapp, 2012). For example, NH₄⁺ concentrations up to 1000 μM did not fully suppress benthic N₂ fixation in a hypoxic basin in the Baltic Sea (Bertics et al., 2013), indicating that additional environmental factors must control the distribution and performance of benthic diazotrophs (Knapp, 2012). We observed high porewater NH_4^+ concentrations at the shallow St. 1 with 316 μM at the sediment surface (0 - 1 cm) increasing to 2022 μM at 20 cm (Fig. 2), while no inhibition of N_2 fixation was found. ThoughHowever, we cannot exclude that a partial suppression occurred. Inhibition experiments of N_2 fixation with NH_4^+ have been conducted in several environments with different findingsresults. For example, benthic N_2 fixation was measured in the Carmens River estuary (New York) with ambient and was still abundant at $\frac{2800 \ \mu M}{NH_4^+}$ concentrations of $\frac{2800 \ \mu M}{NH_4^+}$ (Capone, 1988). In general, these studies suggested that the impact of NH_4^+ on N_2 fixation is more complex than previously thought and poorly understoodhitherto hardly known.

One explanation for why diazotrophs still fix N under high NH₄⁺ concentrations could be that bacteria try to preserve the intracellular redox state by N₂ fixation functioning as an excess for electrons, particularly with a deficient Calvin–Benson–Bassham pathway, as it was shown for photoheterotrophic nonsulfur purple bacteria (Tichi & Tabita, 2000). Previous studies on benthic environments propose that the organic carbon availability can reduce an inhibition of N₂ fixation by abundant NH₄⁺ (Yoch & Whiting, 1986; McGlathery et al., 1998). In the study of Yoch & Whiting (1986), it was shown that enrichment cultures of *Spartina alterniflora* salt marsh sediment showedreacted with different N₂ fixation inhibition stages on for different organic matter species. Another explanation could be that microniches, depleted in NH₄⁺ exist between the sediment grains, which we were unable to track with the applied porewater extraction techniques (Bertics et al., 2013). Such microniches wehre are found in the form of localized organic matter hot spots (Brandes & Devol, 2002; Bertics & Ziebis, 2010), and could also occur for supply NH₄⁺.

4.2.3 Sulfide

 Sulfide is a known inhibitor for many biological processes (Reis, et al., 1992; Joye & Hollibaugh, 1995) and could potentially affect N_2 fixation (Tam et al., 1982). The shallow St. 1 was the only station with sulfide in the porewater, reaching 280 μ M in surface sediments and 1229 μ M in 20 cm (Fig. 2). The presence of relatively high concentrations of sulfide might explain why N_2 fixation was lower at St. 1 compared to St. 6, despite the higher quality, i.e. lower C/N ratio, of organic matter at this station. Because SR rates were highest at St. 1 (Fig. 4), we exclude direct inhibition on SR, although the effect has generally been reported (Postgate, 1979; McCartney & Oleszkiewicz, 1991). Interactions of sulfide with benthic N_2 fixation have so far not been investigated, and hence we can therefore not rule out a partial inhibition of N_2 fixation by sulfide.

4.2.4 Oxygen

Dissolved O₂ can have a considerable influence on N₂ fixation, because of due to the O₂ sensitivity of the key enzyme nitrogenase (Postgate, 1998; Dixon & Kahn, 2004). Bioturbating and bioirrigating organisms can transport O₂ much deeper into sediments than molecular diffusion (Orsi et al., 1996; Dale et al., 2011). In coastal waters, the bioturbation and bioirrigation activity of ghost shrimps was found to reduce N₂ fixation, when sediments were highly colonized by these animals (Bertics et al., 2010). While bottom water O₂ concentrations in the Peruvian OMZ were below the detection limit at the St. 1 to 8 (70 m to 407 m), thereby mainly excluding benthic macrofauna, O₂ concentrations increased to levels above 40 µM at St. 10 (1025 m) where, supporting a diverse bioturbating and bioirrigating benthic macrofauna community was observed (Mosch et al. 2012). Accordingly, this station revealed some of the lowest N₂ fixation activity. We are, however, unable to decipher whether O₂, low organic matter content, and/or the low C/N ratio was responsible for this low activity. Furthermore, several marine diazotrophs have developed strategies to protect the nitrogenase from O₂ (Jørgensen, 1977).

4.3 Comparison of benthic N₂ fixation in different environments

We compiled a list of N_2 fixation rates from different marine environments to gain an overview of the magnitude of N_2 fixation rates measured in the Peruvian OMZ sediments (Tab. 2). We found that N_2 fixation rates from the Peruvian sediments exceed those reported for open ocean sediments (2800 m) (Howarth et al., 1988), bioturbated coastal lagoon sediment (Bertics et al., 2010) and sediments >200 m water depth (Capone, 1988). The highest integrated N_2 fixation rate determined in our study (0.4 mmol N m⁻² d⁻¹, St. 6) closely resembles highest rates found in salt marsh surface sediments (0.38 mmol N m⁻² d⁻¹) and Zostera estuarine sediments (0.39 mmol N m⁻² d⁻¹) (Capone, 1988). Further, our rates were characterized by a similar range of N_2 fixation rates that were previously measured in an organic-rich hypoxic basin in the Baltic Sea (0.08 - 0.22 mmol N m⁻² d⁻¹, Bertics et al., 2013). Different to the above examples, our N_2 fixation rates were 8.5 times lower compared to shallow (< 1 m) soft-bottom sediment off the Swedish coast (Andersson et al., 2014) and 17 times lower than coral reef sediments (Capone, 1988). However, in these environments, phototrophic cyanobacterial mats contributed to benthic N_2 fixation. Given the dark incubation, N_2 fixation of the present study seems to be attributed to

heterotrophic diazotrophs, which is additionally confirmed by the *nifH* gene analysis, where none of the sequences clustered with cyanobacteria (Fig. 6).

5. Summary

 To the best of our knowledge, this is the first study combining N_2 fixation and SR rate measurements together with molecular analysis in OMZ sediments. We have shown that N_2 fixation occurred throughout the sediment and that elevated activity often overlapped with peaks of SR. The molecular analysis of the *nifH* gene confirmed the presence of heterotrophic diazotrophs at all sampling sites. Sequences clustered with sulfate-reducing SR-bacteria, such as *Desulfovibrio vulgaris*, which is a known diazotroph in sediments. In combination, our results suggest that N_2 fixation and SR were coupled to a large extend, but that additional coupling to other metabolic pathways cannot be ruled out completely. The major environmental factor controlling benthic diazotrophs in the OMZ appears to be the organic matter content. -Sulfide was identified as a potential inhibitor for N_2 fixation. We further found no inhibition of N_2 fixation by high NH_4^+ concentrations, highlighting gaps in our understanding of the relationship between NH_4^+ availability and the stimulation of N_2 fixation. N_2 fixation rates determined in the Peruvian OMZ sediments were in the same range of other organic-rich benthic environments, underlining the relation between organic matter, heterotrophic activity, and N_2 fixation.

Author contribution

- J. G. and T. T. collected samples and designed experiments. J. G. performed nitrogen fixation experiments and T. T. conducted sulfate reduction experiments. S. S. and A. W. D. measured porosity, DIC, organic carbon content and C/N. J. G., T. T., C. R. L. and S. S. analyzed the data. J. G. and C. R. L. performed PCR assay and sequence molecular analysis.
- J. G. prepared the manuscript with contributions from all co-authors and T. T. supervised the work.

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Figure captions

Fig. 1. Cross-section of dissolved O_2 concentrations (μ M) along the continental margin of the Peruvian OMZ at 12°S. The vertical lines represent CTD cast for O_2 measurement during the cruise M92. Stations 1 to 10 for <u>MUC-multicorer (MUC)</u> sampling are indicated by station numbers according to Dale et al. (2015).

Fig. 2: Biogeochemical porewater profiles in MUC cores from sampling stations along the 12°S depth transect. Graphs show NH_4^+ (μ M), $SO_4^{2^-}$ (mM), sulfide (μ M), organic carbon content (C_{org} , wt%) and the C/N ratio (molar). Information about bottom water O_2 concentrations (BW O_2 , μ M) is provided at the right margin.

 Fig. 3: Sediment profiles of N_2 fixation nitrogenase activity (NA, nmol $C_2H_4-N_2$ cm⁻³ d⁻¹, average of three replicates) and sulfate reduction rates (SR, nmol SO_4^{2-} cm⁻³ d⁻¹, two replicates (R1 and R2)) from 0 - 20 cm at the six stations. The upper x-axis represents the N_2 fixation NA, while the lower x-axis represents the SR. Error bars indicate standard deviation of N_2 fixation NA.

 Fig. 4: Integrated nitrogen fixation (mmol N m $^{-2}$ d $^{-1}$, grey bars, average of three replicates) and integrated sulfate reduction (mmol SO₄ $^{2-}$ m $^{-2}$ d $^{-1}$, green bars, two replicates) from 0 - 20 cm, including dissolved inorganic carbon (DIC, mmol m $^{-2}$ d $^{-1}$, red curve from Dale et al., (2015)) and bottom water O₂ (μ M, blue curve) along the depth transect (m). Error bars indicate standard deviation of N₂ fixation.

Fig. 5: Integrated nitrogen fixation (mmol $N_2 \text{ m}^{-2} \text{ d}^{-1}$, grey bars, average of three replicates), average organic carbon content (C_{org} , wt%, orange curve) and the average C/N ratio (molar, yellow curve) from 0-20 cm along the depth transect (m). Error bars indicate standard deviation.

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Fig. 6: Phylogenetic tree of expressed-nifH genes based on the analysis of 120 sequences from the six sampling stations between 70 and 1025 m water depth. Novel detected clusters consisting of several sequences from the same sampling depth are indicated by grey triangles. Reference sequences consist of the alternative nitrogenase anfD, anfG, anfK. Cluster III sequences as defined by Zehr and Turner (2001) are highlighted in blue, Cluster I cyanobacterial sequences are highlighted in green and Cluster I proteobacterial sequences are highlighted in orange. The scale bar indicates the 10% sequences divergence. Sequences marked with an asterisk represent potential PCR contaminated products, with novel clusters distant from those clusters. Sequences determined in this study are termed OMZ plus the corresponding water depth.

1219 Tables

Tab. 1: Sampling deployments, including station number according to Dale et al. (2015), core ID, sampling date and coordinates. Water depth (m) recorded by the ship's winch and bottom water temperature (°C) and bottom water O_2 concentration (μ M; bdl=below detection limit (5 μ M)) measured by the CTD.

Station	Core ID	Date (2013)	Latitude (S)	Longitude (W)	Depth (m)	Temp. (°C)	O ₂ (μM)
1	MUC 13	January 11	12°13.492′	77°10.511′	70	14	bdl
4	MUC 11	January 09	12°18.704′	77°17.790′	144	13.4	bdl
6	MUC 6	January 07	12°23.322'	77°24.181′	253	12	bdl
8	MUC 23	January 15	12°27.198′	77°29.497′	407	10.6	bdl
9	MUC 17	January 13	12°31.374′	77°35.183′	770	5.5	19
10	MUC 28	January 19	12°35.377′	77°40.975′	1025	4.4	53

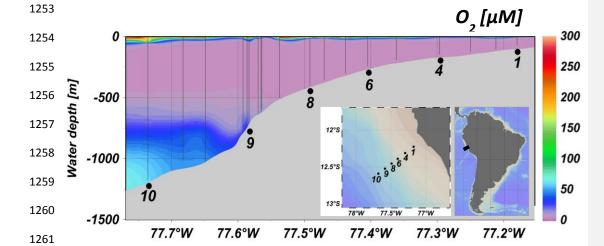
Tab. 2: Integrated rates of nitrogen fixation (mmol m^{-2} d^{-1}) in the Peruvian OMZ sediments from this study compared to other marine benthic environments. Only the highest and lowest integrated rates are shown, as well as the integrated sediment depth (cm) and the method used (ARA=acetylene reduction assay, MIMS=membrane inlet mass spectrometry).

Benthic Environment	N-fixation (mmol N m ⁻² d ⁻¹)	Depth of integration (cm)	Method	Reference
	(IIIIIOTTE III u)	integration (cm)		
PERU OMZ	0.08 - 0.4	0 – 20	ARA	This study
Coastal Region				
Baltic Sea, hypoxic basin	0.08 - 0.22	0 – 18	ARA	Bertics et al., 2013
Bioturbated coastal lagoon	0.8 - 8.5	0 – 10	ARA	Bertics et al., 2010
Brackish-water sediment	0.03 - 3.4	0 – 1	ARA	Andersson et al., 2014
Coral reef sediment	6.09 (± 5.62)	-	-	Capone 1983
Eelgrass meadow sediment	0.15 - 0.39	0 – 5	ARA	Cole and McGlathery, 2012
Eutrophic estuary	0 – 18	0 – 20	MIMS	Rao and Charette, 2012
Mangrove sediment	0 - 1.21	0 – 1	ARA	Lee and Joye, 2006
Salt marsh surface sediment	0.38 (± 0.41)	-	-	Capone 1983
Subtidal sediment	0.6 - 15.6	0 - 30	MIMS	Fulweiler et al., 2007
Zostera estuarine sediment	0.39	-	-	Capone 1983
OPEN OCEAN				
Atlantic ocean (2800 m)	0.00008	-	-	Howarth et al., 1988
< 200 m sediments	0.02 (± 0.01)	-	-	Capone 1983
Mauritania OMZ	0.05 - 0.24	0 – 20	ARA	Bertics and Treude, unpubl

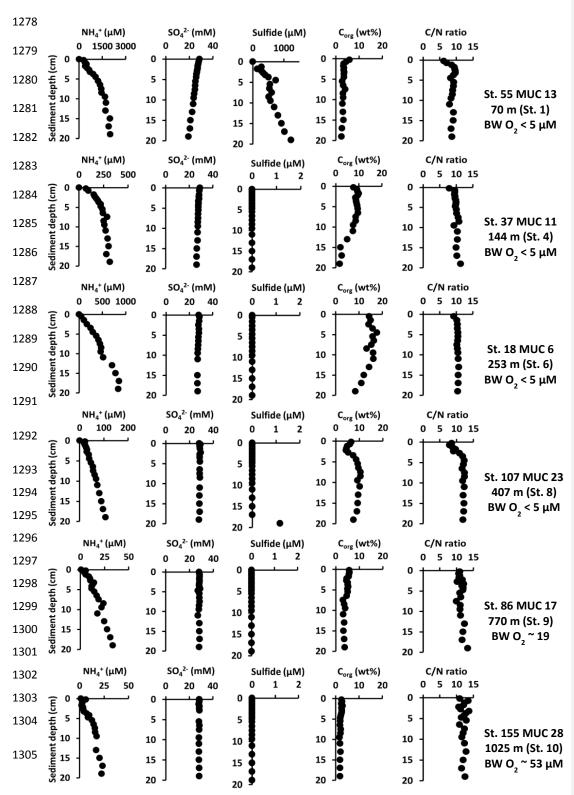
1249 Figures

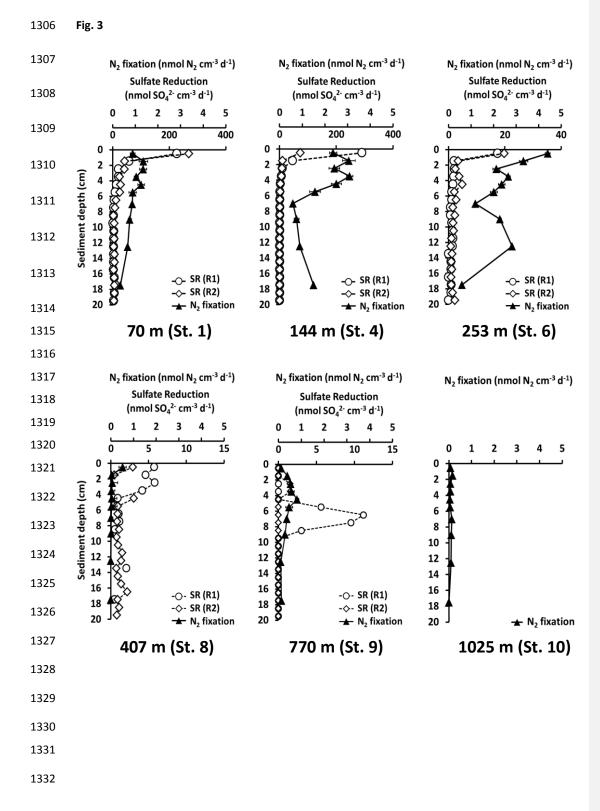
Fig. 1

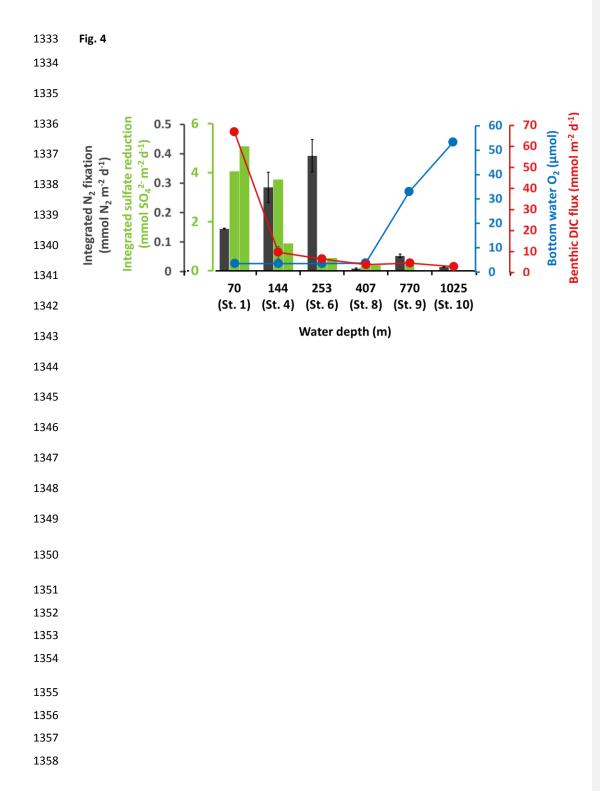


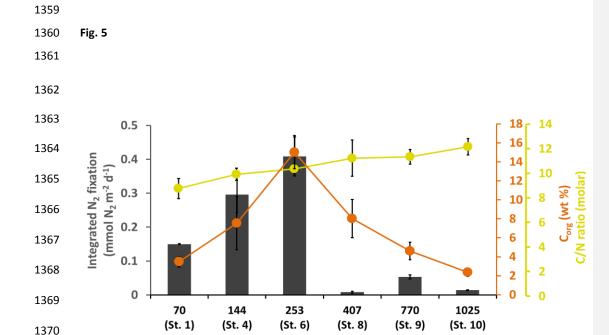












(St. 6)

(St. 8)

Water depth (m)

(St. 9)

Fig. 6

