

Response to Referee #1:

1. We agree, deleted the sentence and replaced it by the following: “How did long-term N and PK fertilization in absence of the respective other influence SOC stocks and how did pedo-climatic factors affect those responses?”

2. We agree that the sample preparation was not described in sufficient detail. A) soils were air-dried and sieved to 2 mm. B) Unfortunately, there is no good technical documentation on that issue. However, we know that two of the earlier methods (LOI and Walkley-Black) have been used in the past for SOC analysis. In our case, we are looking at relative differences between treatments. Therefore, the possible bias induced by methodological considerations should be insignificant or very small. C) Samples with pH (H₂O) exceeding 6.7 were pre-treated with 2 M HCl to remove carbonates. We complemented the M+M section accordingly.

3. We have now added the 1NPK, 2NPK and 3NPK treatments as a “positive reference” for the reduced fertilisation treatments. However, the focus of this manuscript should remain the comparison of only PK vs. only N. Therefore we have only added those treatments in Figure 1 and present the results in the results section shortly as follows: “The effect of NPK fertilisation on NPP was synergistic, thus higher as the sum of the increases after PK and N fertilisation alone (Fig. 1 (a)). This strong increase in C input did also positively affect SOC stocks.”

4. We agree that the discussion was relatively weak regarding the explanatory variables. But due to the inconsistency in the models (different explanatory variables), the discussion of these models is not straightforward. However, we now added the following sentences: “Surprisingly, SOC dynamics of the PK_{0N} fertilised treatments were independent from relative yield increase and thus carbon input, while the variability in SOC dynamics in 1N_{0PK} was partly explained by differences in relative yield increase. This fits the observation, that carbon dynamic in the PK_{0N} treatments was decoupled from aboveground NPP, while it was positively correlated with aboveground NPP in the N_{0PK} treatments.”

5. The presentation of Figure 2 was chosen to better visualise that the SOC/yield relation is completely opposing between the two fertilisation treatments. As such, the graph carries a clear message. High regression coefficients (although only for 3 points) is “nice to have”, but not extremely important for the message. We have now added error bars in both directions.

6. We did not understand the suggestion with the datapoint size, since all datapoints represent a combination of the explanatory variables used. We however understand the criticism that the figure might not be of much value and that the statistical results could also be presented in a table. It might be a matter of taste, but we think that visualising the model fits, although it might be somewhat abstract, helps the reader to directly get an idea of how good the used variables could explain the data and how the different experiments distribute around the 1:1 line. I myself prefer graphs, and the manuscript is not overloaded with graphs (only 2 others). Therefore we do not think that it is absolutely necessary to change it into a table.

Specific comments:

1. We agree and added a reference

2. The reference was missing. We now added it to the reference list.

3. We changed this accordingly.
4. We changed this accordingly.
5. We changed this into SOC stock changes.
6. We agree that this sentence belongs to M+M and also added the N and NPK treatments, which have not been counted in this sentence (which was from an earlier manuscript version). The sentence now reads: “A total of 88 pairs of unfertilised vs. fertilised was investigated, with 3PK_{ON} only being present in 8 experiments, while all other treatments were present in all 10 experiments.” We also had to change the next sentence, which now reads: “As expected, we observed an average increase in yields of 1% (1PK_{ON}), 10% (2PK_{ON}) and 15% (3PK_{ON}) (Fig. 1A) following PK fertilisation.”
7. We changed that accordingly.

Response to Referee #2:

We agree that increases in P availability can also increase SOC decomposition in N-rich soils. One indication for that might be the results of Cleveland et al. (2006), who found higher respiration when N+P was added as compared to N alone. We therefore added the following sentence at the end of the hypothesis 4 section: “Furthermore, Cleveland et al. (2006) found higher respiration when adding glucose with N+P as compared to the addition of glucose and N. This does indicate, that the stimulating effect of P on heterotrophic respiration is not restricted to N-poor soils.”

The experimental design of this study does however hamper such an investigation. We have checked SOC stock changes with increasing PK fertilization also with N (and different N levels) as a reference (instead of taking the unfertilized treatment as a reference). However, the combination of N+PK did always strongly increase net primary production and thus carbon inputs, which did most likely override the effects of PK on decomposition (because we find higher stocks under NPK than under N). Therefore we cannot confirm this hypothesis with the data we have. We only found this net negative effect of PK fertilization when compared to the unfertilized controls.

We agree that the paper was lacking some thoughts about the implications of our findings. We therefore added the following sentence to the conclusions: “It (the finding) might be helpful to model SOC stock changes after sudden stoichiometric imbalances in ecosystems as induced by fires, atmospheric deposition, or enhanced soil P extraction by legumes.” We do not think that it makes much sense to develop even more on this aspect in the discussion because that issue lies beyond the scope of the manuscript.

1 Phosphorus fertilisation under nitrogen limitation can deplete soil
2 carbon stocks - Evidence from Swedish meta-replicated long-term
3 field experiments

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20 Abstract

21 Increasing soil organic carbon (SOC) in agricultural soils can mitigate atmospheric CO₂
22 concentration and also contribute to increase soil fertility and ecosystem resilience. The role of
23 major nutrients on SOC dynamics is complex, due to simultaneous effects on net primary
24 productivity (NPP) that influence crop residue carbon inputs and on the rate of heterotrophic
25 respiration (carbon outputs). This study investigated the effect on SOC stocks of three different
26 levels of phosphorus and potassium (PK) fertilisation rates in the absence of nitrogen fertilisation
27 and of three different levels of nitrogen in the absence of PK. This was done by analysing data
28 from 10 meta-replicated Swedish long-term field experiments (>45 years). With N fertilisation,
29 SOC stocks followed yield increases. However, for all PK levels, we found average SOC losses
30 ranging from $-0.04 \pm 0.09 \text{ Mg ha}^{-1} \text{ yr}^{-1}$ (ns) for the lowest to $-0.09 \pm 0.07 \text{ Mg ha}^{-1} \text{ yr}^{-1}$ ($p=0.008$) for
31 the highest application rate, while crop yields as a proxy for carbon input increased significantly
32 with PK fertilization by 1, 10 and 15%. We conclude that SOC dynamics are mainly output-
33 driven in the PK fertilised regime but mostly input-driven in the N fertilised regime, due to the
34 much more pronounced response of NPP to N than to PK fertilisation. It has been established
35 that P rather than K is the element affecting ecosystem carbon fluxes, where P fertilisation has
36 been shown to: i) stimulate heterotrophic respiration, ii) reduce the abundance of arbuscular
37 mycorrhizal fungi and iii) decrease crop root:shoot ratio, leading to higher root-derived carbon
38 input. The higher export of N in the PK fertilised plots in this study could iv) have led to
39 increased N mining and thus mineralisation of organic matter. More integrated experiments are
40 needed to gain a better understanding of the relative importance of each of the above-mentioned
41 mechanisms leading to SOC losses after P addition.

42

43 1. Introduction

44 Maintaining or increasing soil organic carbon (SOC) stocks in agricultural soil is important for
45 ecosystem resilience, soil fertility and sustainable crop production. Sequestration of SOC is also
46 an important climate mitigation option (~~Lugato et al., 2006~~Lal, 2003); Lugato et al., 2006). The
47 annual export of plant biomass demands a sufficient nutrient supply, mainly of nitrogen (N),
48 phosphorus (P) and potassium (K). To cope with steadily increasing global food demand and
49 limitations to spatial expansion of agricultural area, the optimal balance between production and
50 sustaining soil fertility has to be found. The magnitude and dynamics of SOC stocks are
51 determined by the ratio between inputs and outputs of carbon. It is a well-established fact that
52 mineral (NPK) fertiliser has a strong positive effect on net primary production (NPP) and thus on
53 carbon inputs (Kätterer et al. 2012). In three global meta-analyses, the increase in topsoil SOC
54 following mineral nitrogen fertilisation was reported to range from 3.5 to 8% (Alvarez, 2005;
55 ~~Ladha et al., 2011; Lu et al., 2011~~). Mineral fertilisation is therefore generally recommended as a
56 suitable management practice for SOC sequestration (Lugato et al., 2006). In many long-term
57 experiments with differing fertiliser treatments, SOC stocks have been well explained by
58 measured or estimated NPP (~~e.g. Kätterer et al. 2013~~)(Kätterer et al., 2012). However, the effect
59 of N fertilisation on the corresponding C output by heterotrophic respiration and the mechanisms
60 involved are still not fully understood. Nitrogen fertilisation has been shown to decrease
61 respiratory C losses, especially in combination with a labile C source amendment. This is mostly
62 explained by the “N mining”-theory, which predicts accelerated microbial decomposition of
63 more recalcitrant, energy-poor organic matter in the search of nitrogen under N deficiency.
64 Energy for this break-down is thereby derived from labile C, which in turn cannot efficiently be
65 used for biosynthesis (Bradford et al., 2008; Craine et al., 2007; Milcu et al., 2011; Schimel and

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66 Weintraub, 2003). Cleveland and Liptzin (2007) reported a globally well constrained microbial
67 biomass C:N:P ratio of 60:7:1, which indicates stoichiometric ~~constrains~~constraints for microbial
68 growth. Recently, Murphy et al. (2015) observed that specific N-rich compounds of soil organic
69 matter were primed after glucose addition. This observation can be interpreted as selective N-
70 mining. The opposite effect, i.e. acceleration of C mineralisation by N addition, has also been
71 reported (Allen and Schlesinger, 2004; Milcu et al., 2011; Reed et al., 2011), indicating that N
72 can be a limiting factor for microbial breakdown. To date, studies about the effect of mineral
73 fertiliser on SOC have mainly focused on N or combined N, P and K fertiliser, because nitrogen
74 has the most obvious effects on the carbon cycle and NPK is the most common fertiliser
75 combination used in commercial agriculture. However, it has recently been shown that P or PK
76 fertilisation can have stimulating effects on heterotrophic respiration (Cleveland and Townsend,
77 2006; Cleveland et al., 2002; Fisk et al., 2015; Nottingham et al., 2015). These studies were
78 conducted in forest ecosystems, mostly on P-limited tropical soils, short-term and biased towards
79 carbon output, since changes in carbon input after fertilisation are more difficult to measure in
80 forest systems and the assessment of net effects on SOC requires long-term observation. The
81 response of SOC to P and K fertilisation of arable soils is not well documented, although these
82 elements are considered to be the most important plant nutrients after nitrogen. As such, PK
83 fertiliser application does usually have a positive effect on NPP, which should hypothetically
84 ~~counterbalances~~counterbalance eventual losses via increased heterotrophic respiration. “P
85 mining” at the cost of C, as the equivalent to “N mining” under P limited conditions has not been
86 observed in soils (Craine et al., 2007). The Swedish long-term Soil Fertility Experiments were
87 established between 1957 and 1966 (Carlgren and Mattsson, 2001). These experiments are
88 located across the country and cover a wide range of climatic and pedological conditions. They

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89 are unique because all of them have an almost identical experimental design consists of a
90 combination of different N and PK levels. Due to this meta-replication and the length of the
91 experiments (>45 years) they have provided good opportunities to examine different hypotheses
92 under *in situ* conditions throughout several decades (Simonsson et al., 2009; Williams et al.,
93 2013). The effect of PK without N or N without PK, as examples for two extremely opposing
94 nutrient regimes, has not been assessed. This comparison has the potential to provide powerful
95 insights on long-term nutrient effects on SOC dynamics. The main question asked was: ~~Do~~How
96 did long-term N and PK ~~fertilizer applications~~fertilization in absence of the respective other
97 ~~deviate in their effects on~~influence SOC stocks ~~regarding magnitude and direction?~~how did
98 pedo-climatic factors affect those responses?

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100 2. Materials and Methods

101 2.1 Dataset

102 The dataset covered 10 sites from the ongoing Swedish long-term soil fertility experiments
103 (Table 1), all located in central and southern Sweden (with five sites in each region). A total of
104 16 combinations of mineral fertiliser applications (NPK), including four different N levels and
105 four different PK levels, are being compared in two crop rotations (i.e. with or without manure
106 application) with two replicates in a randomised split block design. The N levels (applied as
107 ammonium nitrate before 1989 and as Nitro Chalk afterwards) are 0, 50, 100, and 200 kg for the
108 southern Swedish sites and 0, 41, 82, 125 kg for central Sweden and will be referred to as 0N,
109 1N, 2N and 3N in the text. The PK levels (P applied as mono superphosphate before 1994 and as
110 triple superphosphate afterwards and K applied as potassium chloride) are 0 kg, replacement of

111 the harvested PK, replacement plus 15 kg P and 40 kg K and replacement plus 30 kg P and 80 kg
112 K for the southern Swedish sites, and 0 kg, replacement, replacement plus 20 kg P and 50 kg K
113 and replacement plus 30 kg P and 80 kg K for the central Swedish sites. These respective levels
114 are referred to as 0PK, 1PK, 2PK and 3PK in the text. In this study, we analysed the effect on
115 SOC of the four PK levels with no N fertiliser (Unfertilised control, 1PK_{0N}, 2PK_{0N}, 3PK_{0N}) and
116 the four N levels with no PK fertiliser (Unfertilised control, 1N_{0PK}, 2N_{0PK}, 3N_{0PK}) in the rotation
117 without farmyard manure. As a positive reference, we also assessed the 1NPK, 2NPK, 3NPK
118 treatments. A total of 88 pairs of unfertilised vs. fertilised was investigated, with 3PK_{0N} only
119 being present in 8 experiments, while all other treatments were present in all 10 experiments.
120 The four-year rotation investigated in the southern Swedish experiments consisted of barley,
121 oilseed rape, winter wheat and sugar beet, while the six-year rotation investigated in central
122 Sweden consisted of barley, oats, oilseed rape, winter wheat, oats and winter wheat. Initial soil
123 characteristics and climate data are presented in Table 2. A detailed description of the soil profile
124 and mineralogical composition of the soil at the sites is provided in numerous publications
125 (Carlgren and Mattsson, 2001). A summary is provided in Table 2. The soils are sampled
126 regularly to a depth of 20 cm (plough layer), air-dried and initially sieved to 2 mm. Initial plot-
127 wise carbon data are available for all experiments. Samples with pH (H₂O) exceeding 6.7
128 were treated with 2 M HCl to remove carbonates. Total carbon and nitrogen was determined by
129 dry combustion using an elemental analyser (LECO-CNS-200, St. Joseph, MI, USA), P and K
130 were both extracted using two different methods, according to Swedish standards: 0.1M NH₄-
131 lactate and 0.4M HOAc adjusted to pH 3.75 and subsequently analysed using an inductively
132 coupled plasma apparatus (ICP-AES Optima 5300, Waltham, MA, USA). The initial carbon
133 values were measured either with loss on ignition or the Walkley-Black wet oxidation, while

134 there is no good technical documentation on that issue. However, due to the fact that we were
135 investigating differences between treatments, a possible bias induced by methodological
136 considerations should be insignificant or very small.

137 2. 2 Calculation of carbon stocks and dynamics

138 The SOC stocks were calculated by multiplying soil bulk density, sampling depth and SOC
139 concentration. To account for the fact that equal soil masses should be compared within one
140 experiment (Ellert and Bettany, 1995), we used the initial average bulk density value for all
141 treatments, which mathematically equals a depth adjustment to obtain the same soil mass from
142 individual bulk density values (Poeplau et al., 2011). Bulk density values are reported for all
143 experiments (Table 1) and did not change significantly over time (data not shown). The
144 difference in SOC stocks between the unfertilised control and the respective treatment was then
145 determined. The N and PK treatment effects on SOC stocks ($\Delta\text{SOC}_{\text{treatment}}$) was calculated
146 treatment-wise from the difference between initial ($\Delta\text{SOC}_{\text{initial}}$) and final ($\Delta\text{SOC}_{\text{final}}$) SOC stock
147 in treatment and control plots:

$$148 \Delta\text{SOC}_{\text{treatment}} = \Delta\text{SOC}_{\text{final}} - \Delta\text{SOC}_{\text{initial}} \quad (\text{Eq. 1})$$

149 The southern Swedish experiments were initiated in 1957 but the experiment was first divided
150 into subplots as they exist today in 1962, which was the first year in which plot-wise soil
151 sampling was conducted. Thus for these five sites, 1962 was the reference year used to calculate
152 $\Delta\text{SOC}_{\text{treatment}}$. The difference in annual net flux of carbon ($\Delta\text{Cflux}_{\text{net}}$) was calculated by dividing
153 $\Delta\text{SOC}_{\text{treatment}}$ by the number of years between the first and the latest sampling. This represents the
154 mean annual SOC stock change rate in a certain treatment compared with the unfertilised
155 control. As a proxy for carbon inputs (Bolinder et al., 2007), we calculated relative differences in

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156 annual yields between any considered fertilizer treatment and the unfertilised control, and
157 averaged those differences over the whole period.

158 2.3 Statistics

159 Linear mixed effect models were used to assess whether any of the fertiliser treatments had a
160 significant effect on yields and SOC ~~stocks~~stock changes across sites. Treatment was used as the
161 fixed effect, while site was used as a random effect. We calculated the model twice for both
162 fertiliser regimes (N and PK fertilised soils), either using all levels as individual treatments or
163 using levels 1-3 together as one treatment (fertilised) testing it against the unfertilised control.
164 Multiple linear regression models were used to explain the observed variability in the response to
165 PK fertilisation with different explanatory variables for all three PK levels and all N levels
166 separately. These were: clay content, sand content, average initial C, N, $P_{\text{available}}$, $K_{\text{available}}$
167 contents, stoichiometric ratios of those elements, average initial soil pH, average losses of
168 available phosphorus and potassium in the unfertilised control between the first and last
169 sampling, final differences in available phosphorus and potassium between the unfertilised
170 control and the fertilised treatment, final differences in soil pH between the control and the
171 fertilised treatment and relative yield increase after fertilization. Due to the high number of
172 potential explanatory variables, multi-model inference (R package MuMIn) was used for a-priori
173 model selection and model averaging based on the Akaike Information Criterion (AIC). Error
174 ranges shown in the text and in diagrams are standard deviations. All statistics were performed
175 using the R software.

176

177 3. Results

178 ~~A total of 28 pairs of unfertilised and PK fertilised treatment were investigated, with 1 PK and 2~~
179 ~~PK being present at all 10 sites and 3 PK only at 8 sites.~~ As expected, we observed an average
180 increase in yields of 1% (1PK_{0N}), 10% (2PK_{0N}) and 15% (3PK_{0N}) (Fig. 1A) following PK
181 fertilisation. This increase was significant for 2PK_{0N} (p=0.012), 3PK_{0N} (p=0.007) and for all
182 pairs together (p=0.002). An equal, yet more pronounced response of NPP to fertilisation was
183 observed for the different N levels (30 pairs) with significant increases of 45%, 80% and 86% for
184 1N_{0PK}, 2N_{0PK} and 3N_{0PK} respectively.

185 Despite these expected positive effects on yields under both fertiliser regimes, we found
186 contrasting responses of SOC stocks. In the PK fertilised soils, we detected a relative SOC stock
187 depletion over time for 22 pairs out of 28 pairs (Figure 1B), which was statistically significant
188 (p=0.004). All experiments have been running for more than 40 years and since we accounted for
189 the initial differences in SOC between experimental plots, we considered that any influence of
190 prior land use would be negligible. Only at one site, Fors, no decrease in SOC as a consequence
191 of PK fertilisation was observed at all. On average, the negative effect of PK fertilisation
192 increased with PK level: With 1PK_{0N}, SOC stocks decreased by 2.0±4.3 Mg ha⁻¹ or 0.04±0.09
193 Mg ha⁻¹yr⁻¹ (ns); with 2PK_{0N} they decreased by 2.3±5.8 Mg ha⁻¹ or 0.05±0.12 Mg ha⁻¹yr⁻¹ (ns);
194 and with 3PK_{0N} they significantly (p=0.008) decreased by 4.2±3.3 Mg ha⁻¹ or 0.09±0.07 Mg ha⁻¹
195 yr⁻¹ (Fig. 2). SOC stocks under PK fertilisation alone did thus develop directly disproportional to
196 yields, which is in contrast to the observations in the N fertilised soils (Fig. 1, 2). Although SOC
197 stock changes were not significant after long-term N fertilisation and even 1N_{0PK} did ~~even~~ lead to
198 slight losses of SOC, the trend of increasing yields with increasing N level was ~~traced~~tracked by
199 the trend in SOC stocks (Fig. 2).

200 The effect of NPK fertilisation on NPP was synergistic, thus higher as the sum of the increases
201 after PK and N fertilisation alone (Fig. 1 (a)). This strong increase in C input did also positively
202 affect SOC stocks.

203 Observed annual SOC stock changes at different PK_{0N} levels were partly explained by the
204 variables considered, although the degree of explanation varied across PK levels. The changes in
205 the 1PK_{0N} treatments were best explained by mean annual temperature (MAT), average initial
206 SOC concentration and average initial C:P_{available} ratio, whereby all three variables were
207 negatively correlated with annual SOC stock changes. The variables used explained 92% of the
208 observed variation ($R^2=0.92$) (Fig. 3). Losses of SOC were thus most pronounced at sites with
209 high SOC concentration, a relatively warm climate and low available P. The effect of 2PK_{0N}
210 fertilisation on annual SOC stock changes was only significantly correlated with average initial
211 soil pH (Fig. 3) ($R^2=0.23$), while that of 0N 3PK was significantly correlated with mean annual
212 precipitation (MAP) and initial soil carbon, achieving 70% explanation.

213 In contrast to the PK_{0N} fertiliser levels, we found significant explanatory power of relative yield
214 increases on SOC stock change in the 1N_{0PK} levels, which indicates the link between carbon
215 input and carbon stock dynamic. Furthermore, 1N_{0PK} was negatively influenced by the C:P_{available}
216 ratio in the soil, with more negative changes in P-poor soils adding up to a total explanation of
217 39% of the observed variance. C:P_{available} ratio was also the only variable, which significantly
218 explained some of the variability in observed SOC responses to 2N_{0PK} ($R^2=0.34$, Fig. 3).

219 4. Discussion

220 In most agricultural long-term experiments assessing the effects of mineral fertilisers on SOC,
221 the focus is on the application rates or sources of nitrogen (Lugato et al., 2006; Malhi et al.,

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222 1997), because they might be more directly related to applied agronomy issues. The majority of
223 existing studies did observe increased yields and SOC stocks with increased N or NPK
224 fertilization, leading to the perception that carbon input is the major driver for SOC dynamics in
225 agricultural systems (Christopher and Lal, 2007). This pattern has been similarly observed in our
226 study. Our finding that phosphorus and potassium (PK) in the absence of nitrogen fertilisation
227 can have a negative net effect on SOC stocks, significantly so at the highest rate of PK
228 fertilisation, has not been highlighted before. However, in several studies, P fertilisation was one
229 among many other investigated fertiliser treatments without receiving much attention in the
230 respective discussion. In those long-term experiments, mixed responses of P or PK fertilisation
231 on SOC stocks have been observed. Some studies showed losses of SOC (Wyngaard et al., 2012;
232 Yan et al., 2013; Zhu et al., 2007), while others found gains (Shao and Zheng, 2014; Zhengchao
233 et al., 2013). The effect of P or PK fertilisation on yields was mostly slightly positive, which is in
234 line with our findings for the Swedish sites. It is however a new observation that SOC stock
235 changes do not always seem to follow changes in NPP and annual C inputs to soil, as the major
236 source of SOC. This requires explanation. We considered four different mechanisms, each of
237 which could explain parts of this observed depletion in SOC stocks under presumed increased
238 carbon inputs:

- 239 1. *Fertilisation with P stimulates heterotrophic respiration.* A strong microbial response to
240 P addition has been observed in several studies, mostly conducted in P-limited tropical
241 forests (Cleveland and Townsend, 2006; Nottingham et al., 2015), but also in northern
242 hardwood forests (Fisk et al. 2015). Craine et al. (2007) tested the effect of N and P
243 addition on mineralisation for a wide range of different plant materials in different soils
244 and found that C mineralisation decreased with N availability, while P fertilisation

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245 increased short- and long-term C mineralisation. They concluded that basic
246 stoichiometric decomposition theory needs to be revised and carbon cycling models
247 restructured to take into account the effect of P on SOC turnover. It is not entirely
248 understood, why N and P additions show opposing effects on microbial activity, but
249 shifts in microbial community structure could play an important role (Ramirez et al.,
250 2012). Recently, Hartman and Richardson (2013) found that the metabolic quotient of
251 microbes (i.e., respiration per biomass) increased significantly following the addition of
252 inorganic P, which could indicate that microbial activity in soil is largely controlled by P
253 availability. In addition to microbial metabolism, the availability of P and N has also
254 shown to trigger the resource allocation to exo-enzymes that are breaking down organic
255 matter (Sinsabaugh and Moorhead, 1994). Extractable phosphorus has been a stronger
256 predictor for C mineralization than land use (Strickland et al., 2010). Cleveland and
257 Townsend (2006) reported an *in situ* increase in CO₂ efflux of 37% after P fertilisation
258 and an *in vitro* 12% higher respiration in the P fertilised than in the unfertilised soil after
259 adding glucose. Fisk et al. (2015) reported 28-122% higher CO₂ efflux in the P-fertilised
260 soil than in the unfertilised soil, depending on added labile C substrates. Considering
261 these observations, it is not an unrealistic hypothesis that an increase in respiration could
262 be the most important mechanism underlying the observed changes in SOC stocks.
263 Furthermore, the average relative phosphorus demand of microbial biomass corresponds
264 to C:P ratio of 60:1 and a global average C:P ratio of 186:1 in soils is globally well
265 constrained (Cleveland and Liptzin, 2007). In our study, the average C:P_{available} ratio in
266 the unfertilised control soils was 706:1. Thus, the stimulating effect of P on microbial

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267 activity might override the stimulating effect on plant growth, leading to net losses of
268 SOC.

269 2. *Colonisation of roots with arbuscular mycorrhizal fungi (AMF) is reduced when easily*
270 *available P is applied to soil* (Grant et al., 2005; Nagahashi et al., 1996). The main role of
271 AMF hyphae is to increase root uptake capacity for nutrients, including P, by enlarging
272 the surface area of the root system (Pearson and Jakobsen, 1993). In return, the fungi
273 receive carbon assimilates from the plant in the form of carbohydrates. It has been shown
274 that plants can provide up to 17% of their photosynthetic carbon to the fungi (AMF,
275 among other mycorrhiza types) (Hobbie and Hobbie, 2006). Clemmensen et al. (2013)
276 identified root-associated fungi as important regulators of ecosystem carbon dynamics. In
277 addition to the higher C input to soil through greater AMF colonisation, AMF-derived
278 carbon (glomalin) may lead to the formation of more stable aggregates, protecting
279 organic matter (Miller and Jastrow, 2000) and thereby preserving the SOC pool (Grant et
280 al., 2005; Nagahashi et al., 1996; Rillig et al., 2001). Therefore, a decrease in AMF
281 colonisation as a consequence of PK fertilisation can potentially have negative effects on
282 SOC stocks.

283 3. *Fertilisation with P decreases plant root:shoot ratio*. A deficiency in major nutrients (N,
284 P and K) generally leads to an increased root:shoot ratio (Wilson, 1988). Under N
285 deficiency, a typical increase in root:shoot ratio is in the range of 15 to 50% (Hansson et
286 al., 1987; Welbank et al., 1973). There is also evidence that P fertilisation can decrease
287 the root:shoot ratio of crop plants (Ericsson, 1995; Marschner et al., 1996) compared with
288 unfertilised soil. This change in root:shoot ratio occurs because nutrient deficiency forces
289 crops to develop larger root systems than under conditions of sufficient nutrient supply

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290 (Marschner et al., 1996). Moreover, root-derived carbon has been found to contribute
291 more to the stable soil carbon pool than shoot-derived carbon (Kätterer et al., 2011).
292 Carbon inputs to the soil are usually yield-based estimates assuming stable carbon
293 allocation coefficients to different plant parts across fertiliser regimes. This is common
294 practice in soil carbon modelling, although there are indications that it might be
295 erroneous . The observed increase in yield for the 0N PK treatments which was 1, 10 and
296 15% for 1PK_{0N}, 2PK_{0N} and 3PK_{0N}, respectively lead to increased C input to the soil when
297 assuming stable ratios of plant carbon allocation. However, relative to the unfertilised
298 control, if a certain shift in root:shoot ratio did occur, PK fertilisation could actually have
299 caused a decrease in total NPP. Thus, despite higher aboveground NPP, the smaller root
300 inputs and the associated qualitative shift of the total carbon input might have
301 significantly contributed to the observed SOC losses after PK_{0N} fertilisation. However,
302 the disproportional trend of yields and SOC stocks seems unrealistic to be explained by
303 root:shoot ratio shift. Unchanged SOC stocks over all PK_{0N} levels could be expected, but
304 most likely not decreases.

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305 4. *Stronger N mining leads to higher mineralisation of C.* In the 0N PK-fertilised plots with
306 higher yields, more N is lost from the system through export by harvest. Such a negative
307 effect of PK fertilisation on soil N has been observed by Glendining et al. (1997). This
308 could have increased the N deficiency in the soil, leading to stronger “N mining” by
309 microbes, a process in which nutrients are mobilised via decomposition of more stable
310 organic matter (Fontaine et al., 2004). This might lead to loss of C from this pool and is
311 in line with the findings of Kirkby et al. (2011) and (2013), who found that SOC
312 sequestration relies on the availability of nutrients and that the stoichiometric ratio of

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313 C:N:P:S of stable soil organic matter is globally well constrained. We did not find any
314 negative effect of PK fertilisation on SOC stocks in the presence of nitrogen fertilisation
315 ~~(data not shown)~~. This is in support of the hypothesis that the decreasing SOC stocks we
316 observed were rather induced by a indirect response to PK fertilisation. However, the
317 estimated surplus of N extraction in the fertilised plots (with an average crop C:N ratio of
318 70) accounted for only 2-20% of the total soil N loss (average soil C:N ratio of 11) after
319 0N PK fertilisation. Consequently, this mechanism was probably only of minor
320 importance in this study. Furthermore, Cleveland et al. (2006) found higher respiration
321 when adding glucose with N+P as compared to the addition of glucose and N. This does
322 indicate, that the stimulating effect of P on heterotrophic respiration is not restricted to
323 N-deficient soils.

324 The relative importance of each of the first three mechanisms for SOC dynamics in the PK
325 fertilised but N deficient regime remains speculative, since CO₂ efflux, abundance of AMF and
326 belowground biomass production were not measured. However, the observed explanatory power
327 of available C, available P, soil pH and the climate parameters MAT and MAP provide support
328 for mechanisms (1) and (2), i.e. increased soil respiration and less AMF colonisation due to 0N
329 PK fertilisation. Temperature, moisture, substrate availability and nutrient availability are key
330 drivers for microbial activity. Surprisingly, SOC dynamics of the PK_{0N} fertilised treatments were
331 independent from relative yield increase and thus carbon input, while the variability in SOC
332 dynamics in 1N_{0PK} was partly explained by differences in relative yield increase. This fits the
333 observation, that carbon dynamic in the PK_{0N} treatments was decoupled from aboveground NPP,
334 while it was positively correlated with aboveground NPP in the N_{0PK} treatments. Furthermore,
335 Clark et al. (1999) found twice the amount of AMF biomass in a soil with pH 4 compared with a

336 soil of pH 5, which is in line with the more pronounced SOC losses we observed in low pH soils.
337 Attributing the losses completely to changes in heterotrophic respiration, as developed above,
338 appeared realistic when compared with the strong stimulations observed in other studies.
339 However, the tendency of SOC losses after 1N 0PK fertilisation might support the hypothesis of
340 significantly altered root:shoot ratio, which has been equally observed for N and P deficiencies
341 (Wilson, 1988). However, these losses could also be explicable by P mining due to a higher
342 extraction of P in the N fertilised soils.

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343 Craine et al. (2007) provided evidence for the N mining theory, but did not find any indications
344 for P mining. This could potentially be related to the fact, that soil N is mainly stored in organic
345 matter, while considerable fractions of soil P are also stored in anorganic forms, thus P mining
346 would be less related to the break-down of organic matter. However, the fact that P availability
347 did significantly trigger SOC responses to N fertilisation (Fig. 3, higher C:P_{available} ratio led to
348 less increase or more decrease in SOC) could denote that P mining did play a role in the studied
349 experiments to some extent. The positive effect of N on NPP was much more pronounced than
350 the effect of PK, which was most likely enough to overshadow any nutrient control on SOC
351 decomposition. Differences between the two fertiliser regimes on carbon output or any of the
352 above mentioned mechanisms is thus not possible to disclose in the present dataset. It can
353 however be concluded, that SOC dynamics were input-driven under N fertilisation but P
354 deficiency, but output-driven under PK fertilisation with N deficiency.

355 In this study, we were unable to separate the effect of P from a potential effect of K, since the
356 two elements were applied in combination. Previously, Shao and Zheng (2014) found a slightly
357 positive effect of K fertilisation on SOC stocks, whereas Yan et al. (2013) found no changes in
358 yield and SOC stock compared with the unfertilised soil. Several studies have shown that K has

359 an opposing, but less pronounced, effect on root:shoot ratio than P (Ericsson, 1995; Hackett,
360 1968). Van Cleve and Moore (1978) found strongly increasing soil respiration with N and P
361 fertilisation, but slightly decreasing soil respiration with K fertilisation. These studies indicate
362 that the negative effect of PK_{0N} fertilisation we observed in the Swedish experiments is most
363 likely related to P fertilisation, while K could have even counterbalanced this effect to a certain
364 degree. A pH effect on SOC decomposition due to different fertiliser regimes (especially due to
365 K fertilisation) could be excluded, since no systematic difference in pH was observed across
366 treatments. This might be due to the fact that the soils were repeatedly limed, according to local
367 agricultural practices. We cannot anticipate any other mechanism following K fertilisation in our
368 study that could have led to SOC losses. However, it is recognized that the combined fertilisation
369 of P and K used in this study does not allow a clear separation of the individual effects of these
370 two elements. In the few available studies in which specifically P effects were studied, P was
371 however also applied in combination with K as KH₂PO₄ (Cleveland and Townsend, 2006; Craine
372 et al., 2007).

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373 5. Conclusions

374 For soils receiving no N fertiliser, a negative effect of phosphorus and potassium (PK)
375 fertilisation on soil organic carbon (SOC) stocks was observed in nine out of the 10 meta-
376 replicated Swedish long-term experiments studied. To our knowledge, it is the first study that
377 shows a significant depletion of SOC stocks after PK fertilisation in the absence of N
378 fertilisation. The processes involved might be diverse and are certainly not well studied, in
379 contrast to the rising awareness of the importance of soil nutrients for net ecosystem carbon
380 fluxes (Reed et al., 2011). According to the literature, of these two elements, P rather than K is
381 likely to have the highest impact on ecosystem carbon fluxes. ~~These findings require~~ This finding

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382 requires attention, since a positive effect of N on NPP and SOC stock changes is widely
383 accepted. It might be helpful to predict SOC stock changes after sudden stoichiometric
384 imbalances in ecosystems as induced by fires, atmospheric deposition, or enhanced soil P
385 extraction by legumes. Many studies focus on certain aspects of SOC cycling, such as soil
386 respiration or the contribution of arbuscular mycorrhizal fungal (AMF) to the total SOC pool.
387 The results obtained from these studies are obviously helpful in identifying potential processes
388 involved in soil responses, e.g. to certain management practices. However, as long as the relative
389 contribution of each process remains unknown, it is difficult to refine our holistic knowledge on
390 SOC dynamics. The observed negative effect of PK_{0N} fertilisation, with its diverse potential
391 causes, calls for a more integrated approach for studying SOC dynamics.

392

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398

399 References

400 Allen, A. S. and Schlesinger, W. H.: Nutrient limitations to soil microbial biomass and activity in loblolly
401 pine forests, *Soil Biology and Biochemistry*, 36, 581-589, 2004.
402 Alvarez, R.: A review of nitrogen fertilizer and conservation tillage effects on soil organic carbon storage,
403 *Soil Use and Management*, 21, 38-52, 2005.

404 Bolinder, M., Janzen, H., Gregorich, E., Angers, D., and VandenBygaart, A.: An approach for estimating
405 net primary productivity and annual carbon inputs to soil for common agricultural crops in Canada,
406 *Agriculture, ecosystems & environment*, 118, 29-42, 2007.

407 Bradford, M. A., Fierer, N., and Reynolds, J. F.: Soil carbon stocks in experimental mesocosms are
408 dependent on the rate of labile carbon, nitrogen and phosphorus inputs to soils, *Functional Ecology*, 22,
409 964-974, 2008.

410 Carlgren, K. and Mattsson, L.: Swedish Soil Fertility Experiments, *Acta Agriculturae Scandinavica, Section*
411 *B — Soil & Plant Science*, 51, 49-76, 2001.

412 Christopher, S. F. and Lal, R.: Nitrogen Management Affects Carbon Sequestration in North American
413 Cropland Soils, *Critical Reviews in Plant Sciences*, 26, 45-64, 2007.

414 Clark, R. B., Zeto, S. K., and Zobel, R. W.: Arbuscular mycorrhizal fungal isolate effectiveness on growth
415 and root colonization of *Panicum virgatum* in acidic soil, *Soil Biology and Biochemistry*, 31, 1757-1763,
416 1999.

417 Clemmensen, K. E., Bahr, A., Ovaskainen, O., Dahlberg, A., Ekblad, A., Wallander, H., Stenlid, J., Finlay, R.
418 D., Wardle, D. A., and Lindahl, B. D.: Roots and Associated Fungi Drive Long-Term Carbon Sequestration
419 in Boreal Forest, *Science*, 339, 1615-1618, 2013.

420 Cleveland, C. C. and Liptzin, D.: C: N: P stoichiometry in soil: is there a “Redfield ratio” for the microbial
421 biomass?, *Biogeochemistry*, 85, 235-252, 2007.

422 Cleveland, C. C. and Townsend, A. R.: Nutrient additions to a tropical rain forest drive substantial soil
423 carbon dioxide losses to the atmosphere, *Proceedings of the National Academy of Sciences*, 103, 10316-
424 10321, 2006.

425 Cleveland, C. C., Townsend, A. R., and Schmidt, S. K.: Phosphorus limitation of microbial processes in
426 moist tropical forests: Evidence from short-term laboratory incubations and field studies, *Ecosystems*, 5,
427 680-691, 2002.

428 Craine, J. M., Morrow, C., and Fierer, N.: MICROBIAL NITROGEN LIMITATION INCREASES
429 DECOMPOSITION, *Ecology*, 88, 2105-2113, 2007.

430 Ellert, B. H. and Bettany, J. R.: Calculation of organic matter and nutrients stored in soils under
431 contrasting management regimes, *Can. J. Soil Sci.*, 75, 529-538, 1995.

432 Ericsson, T.: Growth and shoot: root ratio of seedlings in relation to nutrient availability. In: *Nutrient*
433 *Uptake and Cycling in Forest Ecosystems*, Springer, 1995.

434 Fisk, M., Santangelo, S., and Minick, K.: Carbon mineralization is promoted by phosphorus and reduced
435 by nitrogen addition in the organic horizon of northern hardwood forests, *Soil Biology and Biochemistry*,
436 81, 212-218, 2015.

437 Fontaine, S., Bardoux, G., Abbadie, L., and Mariotti, A.: Carbon input to soil may decrease soil carbon
438 content, *Ecology Letters*, 7, 314-320, 2004.

439 Glendining, M. J., Poulton, P. R., Powlson, D. S., and Jenkinson, D. S.: Fate of ¹⁵N-labelled fertilizer
440 applied to spring barley grown on soils of contrasting nutrient status, *Plant Soil*, 195, 83-98, 1997.

441 Grant, C., Bittman, S., Montreal, M., Plenchette, C., and Morel, C.: Soil and fertilizer phosphorus: Effects
442 on plant P supply and mycorrhizal development, *Canadian Journal of Plant Science*, 85, 3-14, 2005.

443 Hackett, C.: A STUDY OF THE ROOT SYSTEM OF BARLEY, *New Phytologist*, 67, 287-299, 1968.

444 Hansson, A.-C., Pettersson, R., and Paustian, K.: Shoot and root production and nitrogen uptake in
445 Barley, with and without nitrogen fertilization, *J. Agron. Crop Sci.*, 158, 163-171, 1987.

446 Hartman, W. H. and Richardson, C. J.: Differential nutrient limitation of soil microbial biomass and
447 metabolic quotients (qCO₂): is there a biological stoichiometry of soil microbes?, *PLoS one*, 8, e57127,
448 2013.

449 Hobbie, J. E. and Hobbie, E. A.: ¹⁵N IN SYMBIOTIC FUNGI AND PLANTS ESTIMATES NITROGEN AND
450 CARBON FLUX RATES IN ARCTIC TUNDRA, *Ecology*, 87, 816-822, 2006.

451 [Kätterer, T., Bolinder, M., Berglund, K., and Kirchmann, H.: Strategies for carbon sequestration in](#)
452 [agricultural soils in northern Europe, Acta Agriculturae Scandinavica, Section A–Animal Science, 62, 181-](#)
453 [198, 2012.](#)

454 [Kätterer, T., Bolinder, M. A., Andrén, O., Kirchmann, H., and Menichetti, L.: Roots contribute more to](#)
455 [refractory soil organic matter than above-ground crop residues, as revealed by a long-term field](#)
456 [experiment, Agriculture, Ecosystems & Environment, 141, 184-192, 2011.](#)

457 Kirkby, C. A., Kirkegaard, J. A., Richardson, A. E., Wade, L. J., Blanchard, C., and Batten, G.: Stable soil
458 organic matter: A comparison of C:N:P:S ratios in Australian and other world soils, Geoderma, 163, 197-
459 208, 2011.

460 Kirkby, C. A., Richardson, A. E., Wade, L. J., Batten, G. D., Blanchard, C., and Kirkegaard, J. A.: Carbon-
461 nutrient stoichiometry to increase soil carbon sequestration, Soil Biology and Biochemistry, 60, 77-86,
462 2013.

463 ~~[Kätterer, T., Bolinder, M. A., Andrén, O., Kirchmann, H., and Menichetti, L.: Roots contribute more to](#)~~
464 ~~[refractory soil organic matter than above-ground crop residues, as revealed by a long-term field](#)~~
465 ~~[experiment, Agriculture, Ecosystems & Environment, 141, 184-192, 2011.](#)~~

466 Ladha, J. K., Reddy, C. K., Padre, A. T., and van Kessel, C.: Role of nitrogen fertilization in sustaining
467 organic matter in cultivated soils, J. Environ. Qual., 40, 1756-1766, 2011.

468 [Lal, R.: Global potential of soil carbon sequestration to mitigate the greenhouse effect, Critical Reviews](#)
469 [in Plant Sciences, 22, 151-184, 2003.](#)

470 Lu, M., Zhou, X., Luo, Y., Yang, Y., Fang, C., Chen, J., and Li, B.: Minor stimulation of soil carbon storage
471 by nitrogen addition: a meta-analysis, Agriculture, ecosystems & environment, 140, 234-244, 2011.

472 Lugato, E., Berti, A., and Giardini, L.: Soil organic carbon (SOC) dynamics with and without residue
473 incorporation in relation to different nitrogen fertilisation rates, Geoderma, 135, 315-321, 2006.

474 Malhi, S. S., Nyborg, M., Harapiak, J. T., Heier, K., and Flore, N. A.: Increasing organic C and N in soil
475 under bromegrass with long-term N fertilization, Nutrient Cycling in Agroecosystems, 49, 255-260, 1997.

476 Marschner, H., Kirkby, E. A., and Cakmak, I.: Effect of mineral nutritional status on shoot-root
477 partitioning of photoassimilates and cycling of mineral nutrients, Journal of Experimental Botany, 47,
478 1255-1263, 1996.

479 Milcu, A., Heim, A., Ellis, R., Scheu, S., and Manning, P.: Identification of General Patterns of Nutrient and
480 Labile Carbon Control on Soil Carbon Dynamics Across a Successional Gradient, Ecosystems, 14, 710-719,
481 2011.

482 Miller, R. and Jastrow, J.: Mycorrhizal fungi influence soil structure. In: Arbuscular mycorrhizas:
483 physiology and function, Springer, 2000.

484 Murphy, C. J., Baggs, E. M., Morley, N., Wall, D. P., and Paterson, E.: Rhizosphere priming can promote
485 mobilisation of N-rich compounds from soil organic matter, Soil Biology and Biochemistry, 81, 236-243,
486 2015.

487 Nagahashi, G., Douds Jr, D. D., and Abney, G. D.: Phosphorus amendment inhibits hyphal branching of
488 the VAM fungus Gigaspora margarita directly and indirectly through its effect on root exudation,
489 Mycorrhiza, 6, 403-408, 1996.

490 Nottingham, A. T., Turner, B. L., Stott, A. W., and Tanner, E. V.: Nitrogen and phosphorus constrain labile
491 and stable carbon turnover in lowland tropical forest soils, Soil Biology and Biochemistry, 80, 26-33,
492 2015.

493 Pearson, J. N. and Jakobsen, I.: The relative contribution of hyphae and roots to phosphorus uptake by
494 arbuscular mycorrhizal plants, measured by dual labelling with ³²P and ³³P, New Phytologist, 124, 489-
495 494, 1993.

496 Poelau, C., Don, A., Vesterdal, L., Leifeld, J., Van Wesemael, B., Schumacher, J., and Gensior, A.:
497 Temporal dynamics of soil organic carbon after land-use change in the temperate zone - carbon
498 response functions as a model approach, Glob. Change Biol., 17, 2415-2427, 2011.

499 Ramirez, K. S., Craine, J. M., and Fierer, N.: Consistent effects of nitrogen amendments on soil microbial
500 communities and processes across biomes, *Glob. Change Biol.*, 18, 1918-1927, 2012.

501 Reed, S., Vitousek, P., and Cleveland, C.: Are patterns in nutrient limitation belowground consistent with
502 those aboveground: results from a 4 million year chronosequence, *Biogeochemistry*, 106, 323-336,
503 2011.

504 Rillig, M., Wright, S., Nichols, K., Schmidt, W., and Torn, M.: Large contribution of arbuscular mycorrhizal
505 fungi to soil carbon pools in tropical forest soils, *Plant Soil*, 233, 167-177, 2001.

506 Schimel, J. P. and Weintraub, M. N.: The implications of exoenzyme activity on microbial carbon and
507 nitrogen limitation in soil: a theoretical model, *Soil Biology and Biochemistry*, 35, 549-563, 2003.

508 Shao, X.-h. and Zheng, J.-w.: Soil Organic Carbon, Black Carbon, and Enzyme Activity Under Long-Term
509 Fertilization, *Journal of Integrative Agriculture*, 13, 517-524, 2014.

510 Simonsson, M., Hillier, S., and Öborn, I.: Changes in clay minerals and potassium fixation capacity as a
511 result of release and fixation of potassium in long-term field experiments, *Geoderma*, 151, 109-120,
512 2009.

513 Sinsabaugh, R. L. and Moorhead, D. L.: Resource allocation to extracellular enzyme production: A model
514 for nitrogen and phosphorus control of litter decomposition, *Soil Biology and Biochemistry*, 26, 1305-
515 1311, 1994.

516 Strickland, M. S., Callahan Jr, M. A., Davies, C. A., Lauber, C. L., Ramirez, K., Richter Jr, D. D., Fierer, N.,
517 and Bradford, M. A.: Rates of in situ carbon mineralization in relation to land-use, microbial community
518 and edaphic characteristics, *Soil Biology and Biochemistry*, 42, 260-269, 2010.

519 Van Cleve, K. and Moore, T. A.: Cumulative Effects of Nitrogen, Phosphorus, and Potassium Fertilizer
520 Additions on Soil Respiration, pH, and Organic Matter Content¹, *Soil Sci. Soc. Am. J.*, 42, 121-124, 1978.

521 Welbank, P. J., Gibb, M. J., Taylor, P. J., and Williams, E. D.: Root growth of cereal crops. , 26-66 pp.,
522 1973.

523 Williams, A., Börjesson, G., and Hedlund, K.: The effects of 55 years of different inorganic fertiliser
524 regimes on soil properties and microbial community composition, *Soil Biology and Biochemistry*, 67, 41-
525 46, 2013.

526 Wilson, J. B.: A review of evidence on the control of shoot: root ratio, in relation to models, *Annals of*
527 *Botany*, 1988. 433-449, 1988.

528 Wyngaard, N., Echeverría, H. E., Rozas, H. R. S., and Divito, G. A.: Fertilization and tillage effects on soil
529 properties and maize yield in a Southern Pampas Argiudoll, *Soil and Tillage Research*, 119, 22-30, 2012.

530 Yan, X., Zhou, H., Zhu, Q. H., Wang, X. F., Zhang, Y. Z., Yu, X. C., and Peng, X.: Carbon sequestration
531 efficiency in paddy soil and upland soil under long-term fertilization in southern China, *Soil and Tillage*
532 *Research*, 130, 42-51, 2013.

533 Zhengchao, Z., Zhuoting, G., Zhouping, S., and Fuping, Z.: Effects of long-term repeated mineral and
534 organic fertilizer applications on soil organic carbon and total nitrogen in a semi-arid cropland, *Eur. J.*
535 *Agron.*, 45, 20-26, 2013.

536 Zhu, P., Ren, J., Wang, L., Zhang, X., Yang, X., and MacTavish, D.: Long-term fertilization impacts on corn
537 yields and soil organic matter on a clay-loam soil in Northeast China, *Journal of Plant Nutrition and Soil*
538 *Science*, 170, 219-223, 2007.

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546 Table 1: The 10 experimental sites used in this study with coordinates [$^{\circ}$ N / $^{\circ}$ E], starting year
547 (starting year in this study), year of last sampling, time span between first and last sampling
548 [years] and length of the rotation [years].

Site	Coordinates	Start of experiment	Last sampling	Time span	Rotation Length
Fjärdingslov	55 40 / 13 23	1957 (1962)	2011	49	4
Orup	55 82 / 13 50	1957 (1962)	2011	49	4
Örja	55 88 / 12 87	1957 (1962)	2011	49	4
S. Ugglarp	55 63 / 13 43	1957 (1962)	2007	45	4
Ekebo	55 98 / 12 87	1957 (1962)	2011	49	4
Vreta Kloster	58 50 / 15 50	1966	2007	45	6
Högåsa	58 50 / 15 45	1966	2007	45	6
Bjertorp	58 23 / 13 13	1966	2007	45	6
Kungsängen	59 83 / 17 67	1963	2011	49	6
Fors	60 33 / 17 48	1963	2011	49	6

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555 Table 2: Basic soil parameters and climate data for the 10 study sites at the beginning of the
556 experiments: clay and sand content [%], soil organic carbon content [%], soil pH, bulk density [g
557 cm⁻³], mean annual temperature (MAT) [°C] and mean annual precipitation (MAP) [mm].

Site	Clay	Sand	SOC	pH	Bulk density	MAT	MAP
Fjärdingslov	14	62	1.4	7.9	1.66	7.7	550
Orup	12	59	2.4	6.6	1.51	8.3	769
Örja	23	52	1.1	7.8	1.72	8.3	593
S. Ugglarp	12	63	1.5	6.7	1.50	7.7	686
Ekebo	18	47	3.1	6.9	1.44	8.2	622
Vreta Kloster	48	8	2.1	6.2	1.43	6.4	527
Högåsa	7	78	2.4	6.7	1.38	6.4	527
Bjertorp	30	16	2.2	6	1.37	6.5	593
Kungsängen	56	4	2.1	7.1	1.31	6	543
Fors	18	24	2.2	7.7	1.49	5.5	613

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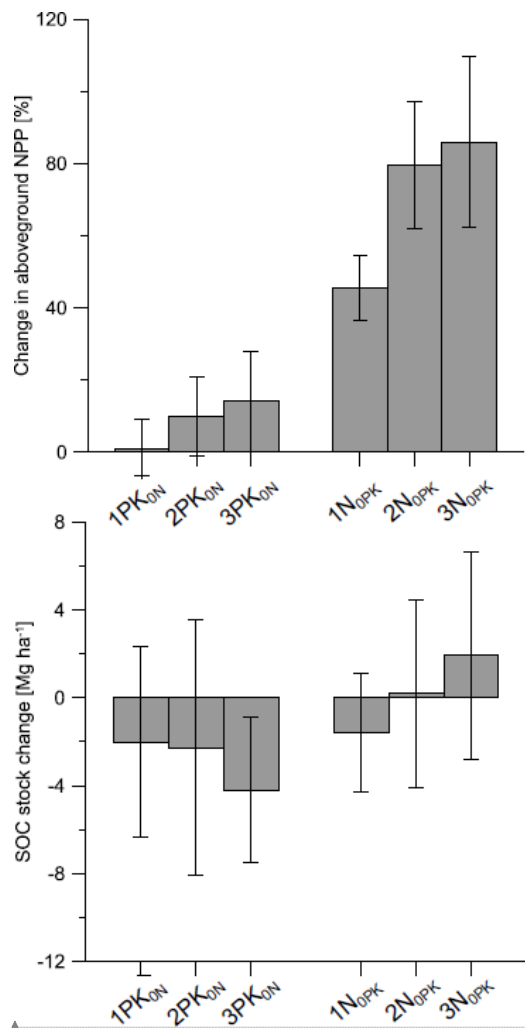
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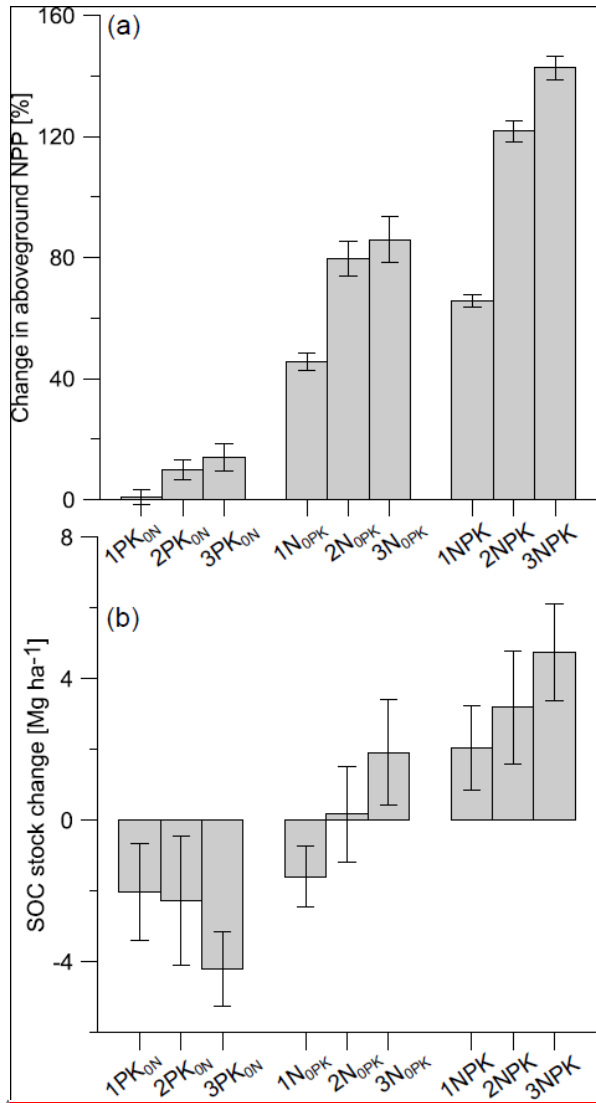
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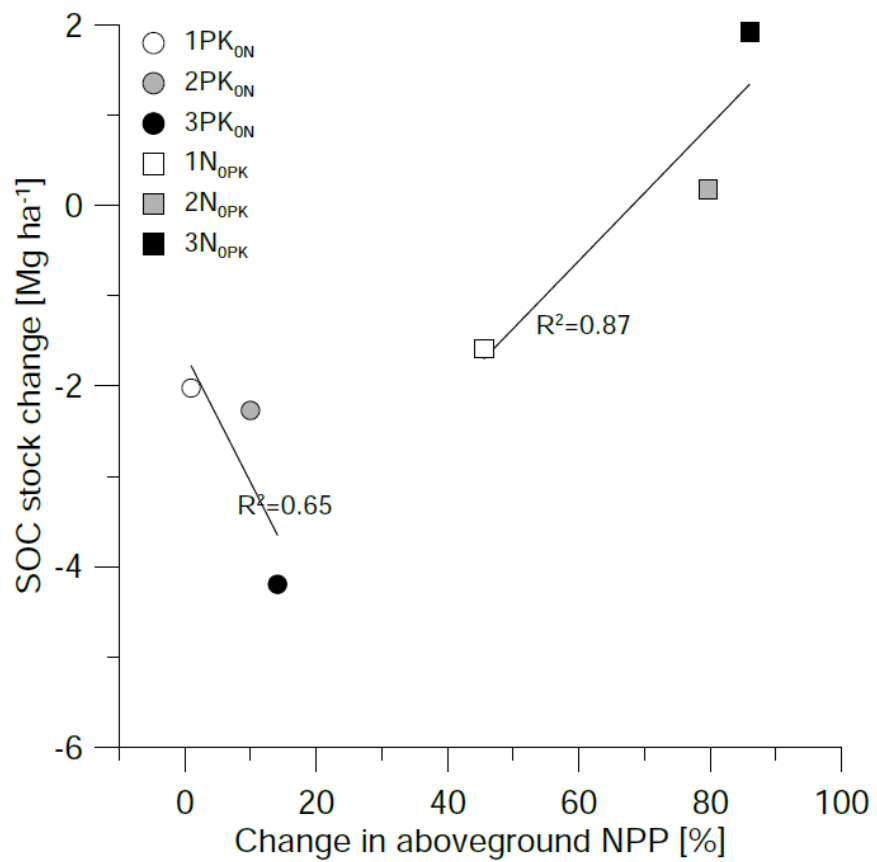
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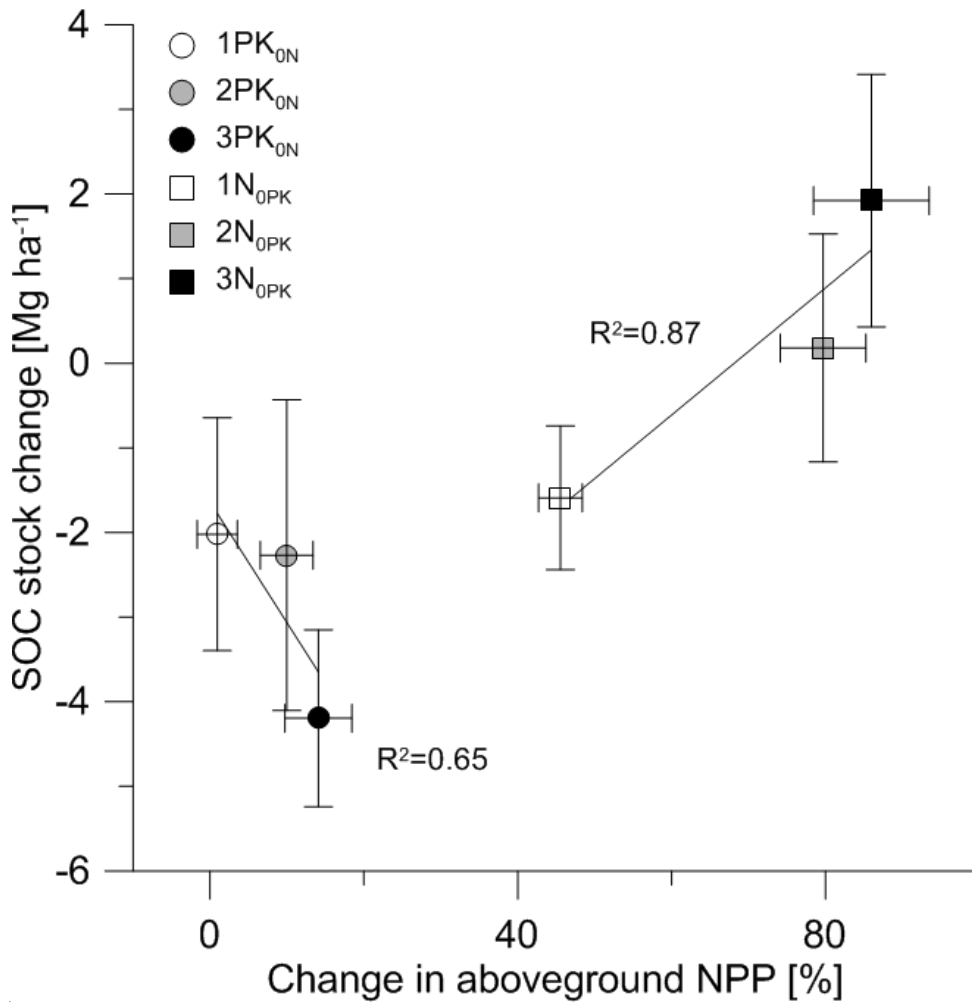
Figure 1: **Aa)** Average relative yield differences and **Bb)** Average soil organic carbon (SOC) stock differences between fertilised and unfertilised soils for all investigated levels of fertilisation (PK_{0N}, N_{0PK}, NPK) with standard deviations errors.

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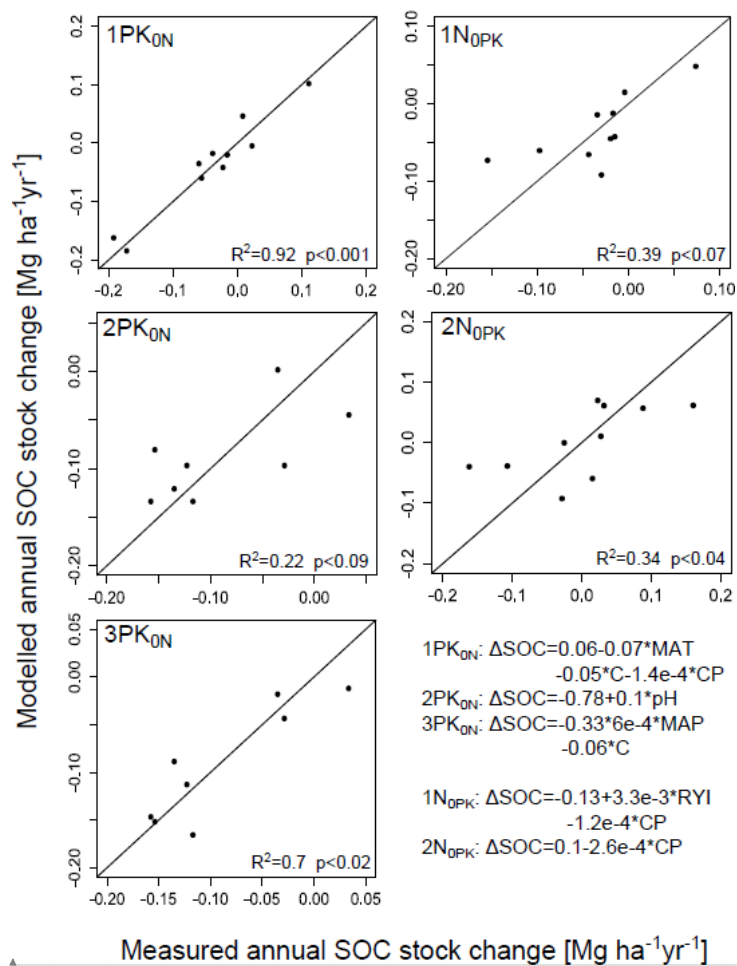
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574 Figure 2: Correlation of average SOC stock differences between fertilised and unfertilised soils
 575 and average yield differences in fertilised and unfertilised soils with regression lines for the P
 576 and N fertilised regimes separately. Error bars indicate standard errors.

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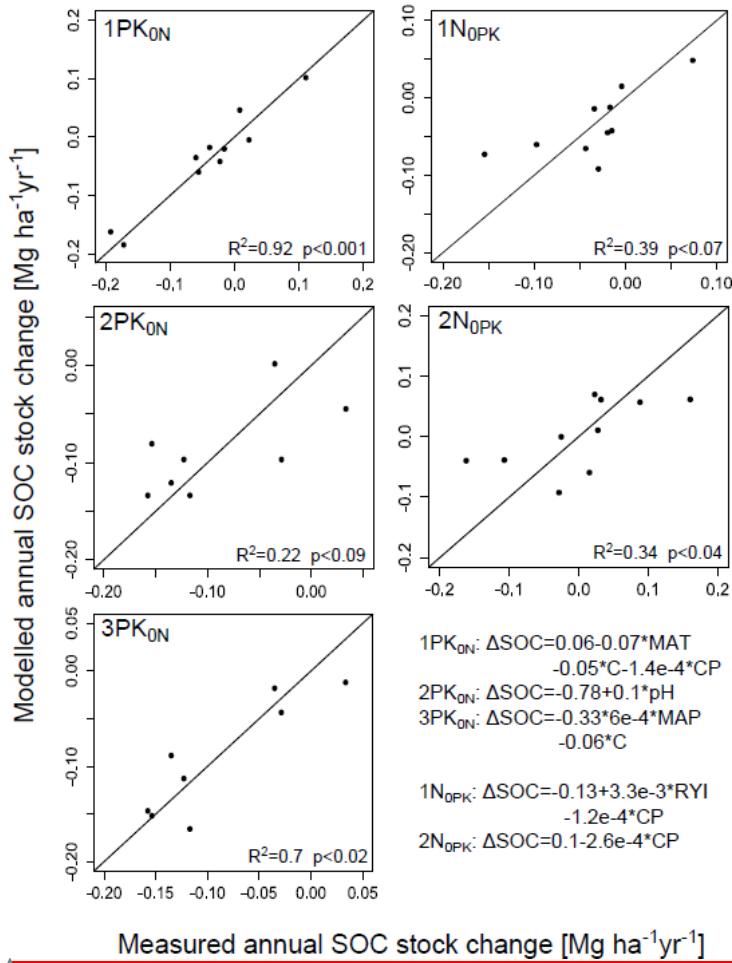
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582 Figure 3: Comparison of measured and modelled annual soil organic carbon (SOC) stock change
 583 [Mg ha⁻¹yr⁻¹] for all phosphorus-potassium (PK) levels and two N levels, with model equations
 584 in which MAT is mean annual temperature [°C], MAP mean annual precipitation [mm], pH
 585 average initial soil pH, C average initial SOC concentration [%], CP the average initial C:P_{available}
 586 ratio and RYI the relative yield increase [%].