

1 **Title:** Temperature-mediated changes in microbial carbon use efficiency and  $^{13}\text{C}$  discrimination

2

3 **Author list and affiliations:**

4 C. A. Lehmeier<sup>1</sup>, F. Ballantyne IV<sup>1,2</sup>, K. Min<sup>1</sup>, S. A. Billings<sup>1\*</sup>

5 <sup>1</sup>Department of Ecology and Evolutionary Biology, Kansas Biological Survey, University of  
6 Kansas, 2101 Constant Ave., Lawrence, KS 66047, USA.

7 <sup>2</sup>now: Odum School of Ecology, University of Georgia, 140 E. Green St., Athens, GA 30602,  
8 USA.

9 **\*Corresponding author:**

10 Sharon A. Billings, Kansas Biological Survey, University of Kansas, 2101 Constant Ave.,  
11 Lawrence, KS 66047, USA. Tel.: 001-785-864-1560, Fax: 001-785-864-1534, email:

12 [Sharon.Billings@ku.edu](mailto:Sharon.Billings@ku.edu)

13

14

15

16

17

18

19 **Abstract**

20 Understanding how carbon dioxide (CO<sub>2</sub>) flux from ecosystems feeds back to climate warming  
21 depends in part on our ability to quantify the efficiency with which microorganisms convert  
22 organic carbon (C) into either biomass or CO<sub>2</sub>. Quantifying ecosystem-level respiratory CO<sub>2</sub>  
23 losses often also requires assumptions about stable C isotope fractionations associated with the  
24 microbial transformation of organic substrates. However, the diversity of organic substrates'  
25  $\delta^{13}\text{C}$  and the challenges of measuring microbial C use efficiency (CUE) in their natural  
26 environment fundamentally limit our ability to project ecosystem C budgets in a warming  
27 climate. Here, we quantify the effect of temperature on C fluxes during metabolic  
28 transformations of cellobiose, a common microbial substrate, by a cosmopolitan microorganism  
29 growing at a constant rate. Specific respiration rate increased by 250% between 13 °C and 26.5  
30 °C, decreasing CUE from 77% to 56%. Specific respiration rate was positively correlated with  
31 an increase in respiratory <sup>13</sup>C discrimination from 4.4‰ to 6.7‰ across the same temperature  
32 range. This first demonstration of a direct link between temperature, microbial CUE and  
33 associated isotope fluxes provides a critical step towards understanding  $\delta^{13}\text{C}$  of respired CO<sub>2</sub> at  
34 multiple scales, and towards a framework for predicting future ecosystem C fluxes.

35

36 **1 Introduction**

37 Because Earth's C cycle is a key regulator of climate, a central goal of biogeochemistry is to  
38 understand biosphere-atmosphere C exchange. Globally, almost all C initially assimilated via  
39 photosynthesis is respired back to the atmosphere as CO<sub>2</sub> by auto- and heterotrophic organisms

40 (Schimel, 1995; Trumbore, 2006). Though we have a reasonably comprehensive understanding  
41 of how environmental conditions influence CO<sub>2</sub> uptake by photosynthetic organisms, our  
42 understanding of how respiratory CO<sub>2</sub> fluxes respond to environmental conditions significantly  
43 lags behind. This is especially true for respiratory CO<sub>2</sub> derived from heterotrophs, which may  
44 account for more than half of respiratory C losses from soils and aquatic systems (Kucera and  
45 Kirkham, 1971; Hanson et al., 2000; Cotner and Biddanda, 2002; Subke et al., 2006). Metabolic  
46 rates of heterotrophs are expected to increase with rising temperatures (Gillooly et al., 2001;  
47 Pomeroy and Wiebe, 2001; Hall et al., 2008), which is of great concern given Earth's large  
48 reservoir of reduced organic matter (OM) that may be mineralized to CO<sub>2</sub> via metabolism  
49 (Hedges et al., 2000; Kirschbaum, 2006). The influence of temperature on the physiology of  
50 heterotrophic microbes must therefore be well understood to project shifts in the global C  
51 balance in a warmer climate.

52 Existing knowledge of Earth's terrestrial C balance has been bolstered by the use of stable  
53 isotopes. A milestone for progress was when photosynthetic responses to environmental  
54 conditions were linked to differences between the stable C isotopic composition ( $\delta^{13}\text{C}$ ) of  
55 atmospheric CO<sub>2</sub> and that of plant products (Farquhar et al., 1982). These differences, caused by  
56 C isotope fractionation during CO<sub>2</sub> diffusion into leaves and subsequent carboxylation (Park and  
57 Epstein, 1961; O'Leary, 1981), impart an isotopic fingerprint on ecosystem C pools and permit  
58 inference about C fluxes from  $\delta^{13}\text{C}$  of ecosystem C pools at multiple spatio-temporal scales  
59 (Farquhar and Richards, 1984; Pataki et al., 2003; Dijkstra et al., 2004; Barbosa et al., 2010).  
60 Recent studies remind us that respiratory C losses also leave an isotopic fingerprint on  $\delta^{13}\text{C}$   
61 values of plant tissues via respiration of substrates with distinct  $\delta^{13}\text{C}$  (Bathellier et al., 2009;  
62 Brüggemann et al. 2011; Ghashghaie and Badeck, 2014), and via C isotope fractionation during

63 decarboxylation in respiratory pathways (Werner and Gessler, 2011; Werner et al., 2011;  
64 Tcherkez et al., 2012). Though not all C isotope fractionations during metabolism are well-  
65 characterized,  $\delta^{13}\text{C}$  of metabolic reaction substrates and products can vary predictably, caused by  
66 kinetic or thermodynamic isotope effects (Rossmann et al., 1991; Gleixner and Schmidt, 1997;  
67 Cleland, 2005; Tcherkez et al., 2012). Accounting for isotope effects in plant respiratory C  
68 losses improves our ability to quantify the contributions of different pools to  $\text{CO}_2$  fluxes and thus  
69 our predictions of terrestrial ecosystem C budgets under changing environmental conditions.  
70 Using  $\delta^{13}\text{C}$  of heterotrophically respired  $\text{CO}_2$  holds similar promise, but if and how changing  
71 environmental conditions influence any fractionation factors for the fluxes associated with the  
72 liberation of C from OM is unknown.

73 Significant uncertainty about the direction and magnitude of C isotope fractionation during  
74 microbial C transformations (Bowling et al., 2008; Werth and Kuzyakov, 2010) renders  
75 quantifying microbial  $\text{CO}_2$  fluxes in ecosystems difficult. Difficulties arise because microbes in  
76 natural systems can access a diverse array of organic substrates with distinct  $\delta^{13}\text{C}$  signatures  
77 (Park and Epstein, 1961; Billings, 2006), the respiration of which influences  $\delta^{13}\text{C}$  of respired  
78  $\text{CO}_2$ . Though we know the growth rate of microbial populations influences C flux into and  
79 through biomass (Kayser et al., 2005), it is impossible to directly quantify microbial growth *in*  
80 *situ*. Furthermore, absence of steady state conditions over a course of  $\text{CO}_2$  flux measurements  
81 makes the interpretation of temperature effects on the magnitude and the  $\delta^{13}\text{C}$  of ecosystem  
82 respiration an even greater challenge (Gamnitzer et al., 2011; Nickerson et al., 2013). Thus,  
83 establishing a mechanistic understanding of the links between temperature, microbial respiration  
84 rates and C isotope fractionation during substrate transformations at a fundamental level requires

85 that we characterize these processes as temperature changes in isolation from other factors that  
86 influence microbial C transformations.

87 To assess the influence of temperature on microbial growth and respiration rates, we grew a  
88 widely distributed Gram-negative, heterotrophic bacterium (*Pseudomonas fluorescens*) in  
89 continuous culture bioreactors (chemostats) at seven temperatures ranging from 13 °C to 26 °C  
90 (Fig. 1) at reactor dilution rates of approximately 0.14 h<sup>-1</sup>, which is equivalent to the relative  
91 growth rates of the microbial populations (Dawson, 1974; Smith and Waltman, 1995; Goldman  
92 and Dennett, 2000; Chrzanowski and Grover, 2008; Ferenci, 2008; Bull, 2010; Egli, 2015). We  
93 measured microbial respiration rates and δ<sup>13</sup>C of respired CO<sub>2</sub> in this open, flow-through system  
94 at steady-state (Craig and Gordon, 1965; Fry, 2006; see Supplementary Material for a detailed  
95 elaboration of this approach). We computed the temperature dependence of a widely applied  
96 metric of microbial C use efficiency (CUE), defined as SGR / (SGR+SRR), where SGR and SRR  
97 are specific growth and specific respiration rates respectively, with units of C per microbial  
98 biomass-C and time. Our simplified system eliminates factors present in natural environments  
99 that preclude accurate assessment of specific growth and respiration rates, and thus accurate  
100 estimates of CUE as defined above. Obtaining accurate estimates of microbial CUE is critical  
101 for projecting C fluxes into the future because the particular value of CUE significantly  
102 influences CO<sub>2</sub> loss rates from ecosystems in models of OM decomposition (Allison et al., 2010;  
103 Wieder et al., 2013). Finally, simultaneously quantifying differences in δ<sup>13</sup>C of organic substrate,  
104 microbial biomass and respired CO<sub>2</sub> along a temperature gradient is critical for partitioning  
105 synoptic CO<sub>2</sub> measurements into component fluxes.

106

## 107 **2 Materials and Methods**

### 108 **2.1 Pre-cultivation of microorganisms for chemostat inoculation**

109 We pre-cultivated *Pseudomonas fluorescens* (Carolina Biological Supply, USA) in nutrient  
110 solution containing 10 mM NH<sub>4</sub>Cl, 1.6 mM KNO<sub>3</sub>, 2.6 mM K<sub>2</sub>HPO<sub>4</sub>, 1.0 mM KH<sub>2</sub>PO<sub>4</sub>, 0.8 mM  
111 MgSO<sub>4</sub>, 0.2 mM CaCl<sub>2</sub>, 0.1 mM CuCl<sub>2</sub>, 0.04 mM FeSO<sub>4</sub>, 0.03 mM MnCl<sub>2</sub> and 0.02 mM ZnSO<sub>4</sub>,  
112 modified from Abraham et al. (1998). The sole C source in the nutrient medium was 10 mM  
113 cellobiose (C<sub>12</sub>H<sub>22</sub>O<sub>11</sub>; with a δ<sup>13</sup>C of -24.2‰); cellobiose is a disaccharide consisting of two  
114 glucose molecules and a basic module of cellulose. Thus, the C to N to P atomic ratio of the  
115 autoclaved, sterile nutrient solution was 100 to 10 to 3.3; its pH was adjusted to 6.5. The  
116 bacteria grew for a few days in batch culture in a flask fitted with a vent for air exchange covered  
117 by a 0.22 μm filter (Fisher Scientific, USA) to avoid contamination. Vessel contents were stirred  
118 continuously in an incubator maintained at 10 °C.

### 119 **2.2 The laboratory mesocosm – the chemostat**

120 The chemostat system was composed of two 1.9 L vessels, a medium reservoir tank and a  
121 reactor, each maintained on separate heating/stirring plates (Fig. 1) in separate incubators. The  
122 reactor volume was on average 870 mL (Supplementary Table 1). The reservoir tank was  
123 connected via a flexible tube to the reactor (Tygon E-LFL pump tubing, Masterflex, USA),  
124 which in turn had an outlet tube (Fig. 1; both tubes had a 1.6 mm inner diameter). When the  
125 chemostat was operated in “continuous culture mode” a peristaltic pump transported fresh  
126 medium to the reactor and simultaneously removed medium from the reactor at the same rate.  
127 Thus, reactor volume remained constant during all chemostat runs. The 0.22 μm filter in the

128 reservoir tank lid allowed for pressure compensation during withdrawal of nutrient solution in  
129 the continuous flow mode. Experimental temperatures were continuously measured with a  
130 thermometer (Oakton, USA) placed in the reactor medium (Fig. 1). This thermometer was  
131 routinely compared against an internal laboratory standard mercury thermometer, before and at  
132 the end of each experiment. The reactor temperatures were adjusted with heating/stirring plate  
133 and incubator settings, and kept constant during all experimental runs.

134 The reactor lid had two ports for gas lines. The outlet port tube was connected to a  $^{13}\text{CO}_2/^{12}\text{CO}_2$   
135 analyzer (G2101-i, Picarro, USA) containing a pump that continuously removed air from the  
136 reactor headspace at an average rate of  $0.025 \text{ L min}^{-1}$ . A water trap (magnesium perchlorate,  
137 Costech, USA) was installed between outlet port of the reactor and the gas analyzer. The  $\text{CO}_2$   
138 analyzer recorded the concentration and the  $\delta^{13}\text{C}$  of the reactor headspace  $\text{CO}_2$  at 0.5 Hz. The  
139 reactor's inlet tube was connected to a mass flow controller (MC-50SCCM, Alicat Scientific,  
140 USA), which in turn, was connected to a gas cylinder containing  $\text{CO}_2$ -free air (Fig. 1). The mass  
141 flow controller was programmed to maintain the reactor headspace at constant atmospheric  
142 pressure; thus, the  $0.025 \text{ L min}^{-1}$  headspace air removed by the  $^{13}\text{CO}_2/^{12}\text{CO}_2$  analyzer pump was  
143 instantaneously replaced with  $\text{CO}_2$ -free air flowing from the gas cylinder into the reactor  
144 medium. Assuming (1) that 1 mol of  $\text{O}_2$  is consumed per 1 mol of  $\text{CO}_2$  produced in aerobic  
145 respiration, (2) a typical reactor headspace  $\text{CO}_2$  concentration of around 2000 ppm at steady state  
146 (see Fig. 2A and below), and (3) an  $\text{O}_2$  concentration of 21% in the air supply to the reactor, the  
147 air supply permitted continuous aerobic metabolism. Routine tests with  $\text{CO}_2$ -free air in sterile  
148 chemostats were performed to ensure there were no leaks in the system.

### 149 **2.3. The chemostat run – standardized protocol and description of events**

150 We conducted seven independent chemostat runs, at temperatures of 13, 14.5, 16, 18, 21, 23.5  
151 and 26.5 °C, in random temporal order. For each of the chemostat runs, we inoculated the  
152 reactor with a 10 mL aliquot of the *P. fluorescens* pre-culture and activated the flow of CO<sub>2</sub>-free  
153 air through the reactor; this was considered time 0. At the initial stage of a chemostat run, the  
154 bacteria grew in batch culture, that is, there was no flow of fresh nutrient medium from the  
155 reservoir tank to the reactor, and no removal of medium from the reactor (Fig. 1).

### 156 **2.3.1 Respiration measurements at chemical and isotopic equilibrium in the continuous** 157 **flow chemostat at steady-state**

158 At the initial pH of 6.5, inorganic C in the fresh reactor medium was mainly in the form of  
159 H<sub>2</sub>CO<sub>3</sub> (aq) and HCO<sub>3</sub><sup>-</sup> (Stumm and Morgan, 1981). By continuously bubbling CO<sub>2</sub>-free air into  
160 the reactor, we expelled these initial inorganic C pools from the reactor medium. This was  
161 evident by concentrations of reactor headspace CO<sub>2</sub> of virtually zero in the early stages of batch  
162 culture after each run's inoculation (Fig. 2A). During the phase of rising reactor headspace CO<sub>2</sub>  
163 via respiratory activity of the exponentially growing population (Fig. 2A), inorganic C in the  
164 reactor medium accrued with the increasing addition of CO<sub>2</sub> from microbial respiration. That is,  
165 at any point in time during the phase of increasing reactor headspace CO<sub>2</sub> concentration, the  
166 nutrient medium acted as a sink for respired CO<sub>2</sub> (see also Supplementary Material).

167 Once the respiratory activity of the growing microbial population pushed the reactor headspace  
168 CO<sub>2</sub> concentration above 500 ppm, we transferred the chemostat into the “continuous culture,  
169 open system” mode (Figs. 1, 2; Ferenci, 2008; Bull, 2010). The peristaltic pump henceforth  
170 transported fresh nutrient medium from the reservoir tank to the reactor at a constant rate of, on  
171 average, 118 mL h<sup>-1</sup> (Supplementary Table 1), and simultaneously removed medium from the



172 reactor at the same rate so that the reactor volume remained constant. Initial chemostat  
173 experiments indicated that when headspace CO<sub>2</sub> concentrations reached 500 ppm, the bacterial  
174 population was sufficiently dense to maintain itself without being washed out via dilution.  
175 Depending on the reactor temperature, the onset of the continuous culture mode occurred  
176 between 40 h (at 26.5 °C) and 72 h (at 13 °C) after inoculation.

177 After the switch from batch to continuous culture, the rate of increase in reactor headspace CO<sub>2</sub>  
178 concentration gradually slowed because cells were continuously diluted into the waste stream  
179 (Fig. 1), and approached a phase where the CO<sub>2</sub> concentration became stable (Fig. 2A). At this  
180 point, bacteria grown in continuous culture had reached the phase of steady-state growth and  
181 physiology (see Ferenci, 2008; Bull, 2010). A key feature of the continuous culture chemostat  
182 relevant to our study is that at this steady-state, the constant dilution rate of the reactor (the  
183 medium flow rate divided by the reactor volume) is equivalent to the specific growth rate of the  
184 microbial culture (Bull, 2010). That is, washout of cells with the nutrient medium flow is  
185 balanced by cell division so that the size of the population in the reactor can be expected to be  
186 reasonably constant in the time frames employed here (see discussion in Ferenci, 2008; Bull,  
187 2010).

188 Critically, when reactor headspace CO<sub>2</sub> concentrations approached the steady-state, inorganic C  
189 pools came to their respective equilibria as well (Stumm and Morgan, 1981). At this point, pools  
190 of H<sub>2</sub>CO<sub>3</sub> (aq) and HCO<sub>3</sub><sup>-</sup> were no longer a *net* sink for respired CO<sub>2</sub>. As reactor headspace CO<sub>2</sub>  
191 concentrations reached steady state, the system supported constant microbial CO<sub>2</sub> production  
192 reflective of steady-state growth under constant environmental conditions, and reflected  
193 chemical equilibrium (i.e., constant size) of the dissolved inorganic C pools. Thus, the rate of

194 CO<sub>2</sub> addition to the reactor headspace volume at steady-state accurately represented the CO<sub>2</sub>  
195 released during microbial respiration (see also Supplementary Material).

196 We calculated the molar CO<sub>2</sub> production rate of the microbial population as the product of the  
197 average molar CO<sub>2</sub> concentration measured by the <sup>13</sup>CO<sub>2</sub>/<sup>12</sup>CO<sub>2</sub> analyzer for 5 hours at steady  
198 state (Fig. 2A) multiplied by the molar air flow rate through the reactor, which was calculated as

199 
$$\text{Air flow (mol min}^{-1}\text{)} = 0.96 \text{ atm} * 0.025 \text{ L min}^{-1} / (0.082 \text{ atm L mol}^{-1} \text{ K}^{-1} * 296 \text{ K}),$$

200 with 0.96 atm and 296 K being the barometric pressure and the temperature in the lab where the  
201 experiments were performed, 0.025 L min<sup>-1</sup> the average volumetric headspace flow rate through  
202 the reactor and 0.082 atm L mol<sup>-1</sup> K<sup>-1</sup> the gas constant.

203 The δ<sup>13</sup>C of the reactor headspace CO<sub>2</sub> during the earliest batch culture phase was generally very  
204 negative; the <sup>13</sup>CO<sub>2</sub>/<sup>12</sup>CO<sub>2</sub> analyzer cannot accurately measure <sup>13</sup>C and <sup>12</sup>C in very low CO<sub>2</sub>  
205 concentrations (Fig. 2B). The δ<sup>13</sup>C of reactor headspace CO<sub>2</sub> became less negative as the CO<sub>2</sub>  
206 concentration increased (Fig. 2B). During the “climbing” phase of the reactor headspace CO<sub>2</sub>,  
207 the δ<sup>13</sup>C of the CO<sub>2</sub> pool was influenced by isotopic fractionation among gaseous CO<sub>2</sub>, H<sub>2</sub>CO<sub>3</sub>  
208 (aq) and HCO<sub>3</sub><sup>-</sup> (Vogel et al., 1970; Mook et al., 1974; Stumm and Morgan, 1981; Szaran, 1997),  
209 because the dissolved inorganic C pools functioned as a net sink for respired CO<sub>2</sub>. At steady-  
210 state, with constant headspace CO<sub>2</sub> concentrations and constant size of the dissolved inorganic C  
211 pools (see above), isotopic equilibrium was achieved, evidenced by constant δ<sup>13</sup>C readings of  
212 reactor headspace CO<sub>2</sub> (Fig. 2B). As such, in this open system at steady-state, the δ<sup>13</sup>C of the  
213 CO<sub>2</sub> leaving the reactor (the CO<sub>2</sub> measured by the analyzer) is identical to the δ<sup>13</sup>C of microbial  
214 respiration (Craig and Gordon, 1965; Fry, 2006). Importantly, this principle is valid irrespective

215 of temperature, microbial growth rate or microbial biomass in the reactor. (See Supplementary  
216 Material for an elaboration of the principle of chemical and isotopic equilibrium.)

217 We used the average  $\delta^{13}\text{C}$  measurement of reactor headspace  $\text{CO}_2$  over the same five hours in the  
218 stable phase employed for calculations of microbial respiration rates (see above) as the isotopic  
219 signature of  $\text{CO}_2$  respired by the microbial culture at each temperature. Any measurements of  
220 headspace  $\text{CO}_2$  and  $\delta^{13}\text{C}$  during the climbing phase before steady-state (Fig. 2) were not used in  
221 these calculations.

222 For the example chemostat at 23.5 °C, the half-life of the reactor ( $t_{1/2}$ ), i.e., the time it took until  
223 50% of the reactor medium was exchanged with fresh tank medium, was 5.2 h (with  $t_{1/2} = \ln(2) /$   
224 (medium flow rate / reactor volume); Supplementary Table 1). In a homogeneous, well-mixed  
225 system such as that employed here, 95% of the pool (i.e., the reactor) is exchanged with new  
226 medium within approximately five times the half-life. Thus, during the respiration  
227 measurements between time 70 h and 74 h (in the example time course in Fig. 2), any “leftovers”  
228 from the batch culture mode were insignificant, and the microbial culture could be considered  
229 homogeneous. This principle was applicable to all chemostat runs we performed.

230 After the 5-hour respiration measurements were completed, we disconnected the gas lines from  
231 the reactor, connected the mass flow controller directly to the  $^{13}\text{CO}_2/^{12}\text{CO}_2$  analyzer, and replaced  
232 the  $\text{CO}_2$ -free air cylinder with a reference gas cylinder containing 1015 ppm  $\text{CO}_2$  at a  $\delta^{13}\text{C}$  of -  
233 48.9‰ (Matheson, USA). This laboratory standard gas was previously calibrated against  
234 secondary  $\text{CO}_2$  standards (Oztech, USA) and served for any necessary corrections of the  $\delta^{13}\text{C}$  of  
235 the reactor headspace  $\text{CO}_2$  measurements. Across the seven standard measurement procedures  
236 after each individual chemostat run, the  $\delta^{13}\text{C}$  measured for the laboratory standard gas showed

237 only slight variation (1 SD = 0.16‰). CO<sub>2</sub> concentration measurements needed no correction;  
238 measurements of lab-internal gases with previously determined CO<sub>2</sub> concentrations between  
239 chemostat runs showed very stable and accurate analyzer performance.

### 240 **2.3.2 Measurements of extracellular enzyme activities at steady-state**

241 Using principles detailed by Lehmeier et al. (2013) and Min et al. (2014), we tested reactor  
242 medium for activity of the extracellular enzymes  $\beta$ -glucosidase and  $\beta$ -N-acetyl glucosaminidase  
243 across all chemostat temperatures; we never detected extracellular activity of either enzyme. The  
244 lack of extracellular  $\beta$ -glucosidase activity indicates that the sole C source, cellobiose, was  
245 directly taken up by microbes and cleaved intracellularly into glucose monomers for further  
246 metabolism. The lack of extracellular  $\beta$ -N-acetyl glucosaminidase activity suggests that the  
247 inorganic N provided in nutrient medium was the sole source of N taken up by *P. fluorescens*.  
248 These findings do not rule out the possibility that *P. fluorescens* may have taken up (i.e. recycled)  
249 any exuded C-based metabolic compounds, although such a scenario in continuous culture  
250 conditions may not appear to be energetically favorable. Thus, the assumption that the sole  
251 resources used by *P. fluorescens* were the cellobiose and the nutrient medium appears reasonable.

### 252 **2.3.3 Harvest of microbial biomass at steady-state**

253 Immediately after completing the 5 hour respiration measurements, we filtered approximately  
254 300 mL of reactor medium for steady-state microbial biomass using 0.2  $\mu$ m filters made of  
255 polyethersulfone (Pall, USA) and a vacuum pump. The filters had previously been oven-dried  
256 for 48 h at 75 °C and their dry weight determined. We then oven-dried the filters post filtration  
257 for 48 h at 75 °C, removed some of the dry biomass and weighed 1.2 mg of the material into tin

258 cups for subsequent combustion in an elemental analyzer (1110 CHN Combustion Analyzer,  
259 Carlo Erba Strumentazione, Italy) coupled to a ThermoFinnigan DeltaPlus mass spectrometer  
260 (Finnigan MAT, Germany) at the Keck Paleoenvironmental and Environmental Stable Isotope  
261 Laboratory (The University of Kansas, USA). We thus determined biomass C (and N) elemental  
262 content, as well as the  $\delta^{13}\text{C}$  of the biomass. In this analysis, the samples were compared against  
263 a laboratory standard  $\text{CO}_2$  previously calibrated against the same secondary  $\text{CO}_2$  standards as  
264 used in calibration of the  $\text{CO}_2$  standard used for respiration measurements (see above). The  $\delta^{13}\text{C}$   
265 of the substrate cellobiose was measured likewise. Dry weight of the sampled reactor aliquots  
266 and the C content obtained from elemental analysis served to calculate total microbial C content  
267 in the steady-state reactor and to calculate specific respiration rates.

268 At all temperatures studied, biomass C and N contents were virtually the same, on average 27%  
269 and 8% of microbial dry mass, respectively (Supplementary Table 1). From mass balance  
270 calculations, we determined that only a small percentage of the C and N supplied via fresh  
271 medium from the tank was taken up to fuel microbial growth (2.8 % and 4.3 % on average for C  
272 and N, respectively). This suggests that the observed temperature effects on specific respiration  
273 rates and CUE were not confounded by any differences in C and N limitations at the different  
274 temperatures (Goldmann and Dennet, 2000; Cotner et al., 2006; Chrzanowski and Grover, 2008).

275

### 276 **3 Results and Discussion**

277 For *P. fluorescens* grown in continuous culture, CUE, defined as  $\text{SGR} / (\text{SGR} + \text{SRR})$ , declined  
278 with increasing temperature, from 77% at 13 °C to 56% at 26.5 °C (Fig. 3A). Because specific

279 growth rates were similar across the experimental temperatures ( $137 \text{ mg g}^{-1} \text{ h}^{-1}$ ,  $\pm 8$  (1 SD); or  
280  $13.7 \% \text{ h}^{-1}$  in relative terms; Fig. 3B), the decline in CUE was due to the 2.5 fold increase of SRR  
281 with temperature, which rose from  $45 \text{ mg g}^{-1} \text{ h}^{-1}$  at  $13 \text{ }^\circ\text{C}$  to  $113 \text{ mg g}^{-1} \text{ h}^{-1}$  at  $26.5 \text{ }^\circ\text{C}$  (Fig. 3B).  
282 The decline in CUE is also evident in the more than 50% reduction in steady-state dry microbial  
283 biomass with increasing temperature (Fig. 3A). For example, though SGR was approximately  
284 the same and thus the fraction of biomass replaced per time similar across all experimental  
285 temperatures (0.147 and 0.141 at  $13 \text{ }^\circ\text{C}$  and  $26.5 \text{ }^\circ\text{C}$ , respectively; Fig. 3B) microbes at  $13 \text{ }^\circ\text{C}$   
286 generated  $20.9 \text{ mg biomass h}^{-1}$  while those at  $26.5 \text{ }^\circ\text{C}$  generated less than half the amount ( $9.5$   
287  $\text{mg biomass h}^{-1}$ ).

288 Because we did not quantify possible C losses from the population at steady-state such as  
289 secretion of organic acids or other compounds (El-Mansi and Holms, 1989; Nanchen et al.,  
290 2006), gross rates of steady-state cellobiose C uptake may have been slightly higher than what  
291 was calculated from the sum of SGR and SRR. However, the direct observation of *P.*  
292 *fluorescens*' CUE is consistent with the negative effect of increasing temperature on microbial  
293 CUE widely reported in literature for soils and aquatic ecosystems (del Giorgio and Cole, 1998;  
294 Gillooly et al., 2001; Rivkin and Legendre, 2001; Apple et al., 2006; Manzoni et al., 2012; Frey  
295 et al., 2013; Tucker et al., 2013).

296 Across the chemostat runs, we observed strong C isotope fractionations, which created  
297 pronounced differences in  $\delta^{13}\text{C}$  between microbial biomass and the sole C substrate, cellobiose,  
298 and between microbial biomass and respired  $\text{CO}_2$  (Fig. 4). Microbial biomass exhibited 5.5 to  
299  $10.5\%$  more negative  $\delta^{13}\text{C}$  values than the cellobiose and respired  $\text{CO}_2$  was even more  $^{13}\text{C}$   
300 depleted, at least  $4.4\%$  more negative than the biomass (Fig. 4A). Because each chemostat was

301 at steady-state, isotopic mass balance dictates that  $^{13}\text{C}$  “missing” from cellobiose had to  
302 accumulate in another pool in the reactor. The only pool that could have been enriched with the  
303 “missing”  $^{13}\text{C}$  was reactor DOC, which we analyzed for  $\delta^{13}\text{C}$  in four out of the seven chemostat  
304 runs (Fig. 4A). Reactor DOC consisted of a large pool of cellobiose (because the rate of C  
305 consumption by the chemostat cultures was, on average, only 2.8% of the rate of C supply) and  
306 presumably a pool of additional organic compounds (e.g. acetate). Such compounds appear to be  
307 typically secreted from microbial cells at low rates in aerobic chemostats operated at dilution  
308 rates similar to those of our runs (El-Mansi and Holms, 1989; Nanchen et al., 2006), and have  
309 been shown to be enriched in  $^{13}\text{C}$  compared to cellular biomass (Blair et al., 1985). However,  
310 because such a small fraction of the available cellobiose was taken up by *P. fluorescens*, the  
311 fraction of total DOC comprised of secreted organic compounds was small. As a result,  $^{13}\text{C}$   
312 enrichment of any microbial exudates was insufficient to enrich bulk DOC to an extent  
313 detectable by the isotope-ratio mass spectrometer (Fig. 4A).

314 The majority of the fractionation between *P. fluorescens* biomass and the substrate was most  
315 likely due to discrimination against  $^{13}\text{C}$  during cellobiose uptake. If we assume that *P.*  
316 *fluorescens* secreted organic compounds at a rate of 10% of the sum of SGR and SRR (El-Mansi  
317 and Holms, 1989) and that the bacteria did not discriminate against  $^{13}\text{C}$ -containing cellobiose  
318 during uptake (and thus assimilated cellobiose possessed a  $\delta^{13}\text{C}$  of -24.2‰), isotopic mass  
319 balance dictates that the  $\delta^{13}\text{C}$  of the C secretion flux (Fig. 5) would have to be +70‰, at  
320 minimum, across all temperatures. To our knowledge, such high metabolic discrimination  
321 against  $^{13}\text{C}$  would be very unusual for biological systems (O’Leary, 1981). An alternative and  
322 more likely scenario is therefore that *P. fluorescens* took up less  $^{13}\text{C}$ -containing cellobiose than  
323 was supplied as substrate, and that discrimination during uptake contributed substantially to *P.*

324 *fluorescens* biomass and respired CO<sub>2</sub> being more <sup>13</sup>C depleted than the cellobiose supplied.  
325 This conclusion holds for all temperatures studied. If we assume for the example 23.5 °C  
326 chemostat run at steady-state (Figs. 2, 5) that *P. fluorescens* secreted organic compounds at a rate  
327 of 10% of the sum of SGR and SRR, and we further assume that the δ<sup>13</sup>C of secreted compounds  
328 was 11.7‰ less negative than that of the biomass (Blair et al., 1985), that is -18.1‰, then the  
329 δ<sup>13</sup>C of the cellobiose taken up would have been -31.1‰, which is only a 10‰ difference from  
330 the cellobiose provided, and therefore probably a more likely scenario (Fig. 5).

331 Substantial <sup>13</sup>C depletion of respired CO<sub>2</sub> relative to microbial biomass has not, to our  
332 knowledge, been reported in other studies. With the microbial C consumption rate amounting to  
333 only 2.8% of the rate of C supply, C availability was high compared to what microbes in their  
334 natural environments typically experience (Tempest and Neijssel, 1978; Cole et al., 1988;  
335 Hobbie and Hobbie, 2013), potentially promoting enzymatic discrimination. The cellobiose δ<sup>13</sup>C  
336 of -24.2 ‰ implies a <sup>13</sup>C/<sup>12</sup>C ratio of ~1/91. Considering the molecular formula of cellobiose  
337 C<sub>12</sub>H<sub>22</sub>O<sub>11</sub>, this means that not more than about one out of eight cellobiose molecules in the  
338 supplied substrate included a <sup>13</sup>C atom. Faster diffusion of the isotopically lighter cellobiose  
339 molecules may have contributed to a lower probability of <sup>13</sup>C-containing cellobiose approaching  
340 bacterial membrane uptake sites, and hence, to the differences between δ<sup>13</sup>C of substrate and  
341 biomass (Fig. 4A). However, isotope fractionation during diffusion – a physical process  
342 dependent on compound mass – would likely exhibit a continuous temperature response. Thus,  
343 it seems unlikely that fractionation during diffusion was the primary driver of the pronounced,  
344 discontinuous changes in the difference between substrate and biomass δ<sup>13</sup>C, which ranged from  
345 5.5 to 10.5‰ (Fig. 4A). Rather, this variation, with one apparently linear part between 13 °C and  
346 16 °C and another between 18 °C and 23.5 °C, may be explained parsimoniously by a



347 significant, discontinuous reorganization of enzyme-mediated C fluxes into and out of bacterial  
348 cells (see Nanchen et al., 2006), induced by differences in temperature at which *P. fluorescens*  
349 was growing and the related differences in substrate uptake rates.

350 Evidence from work on C isotope distribution within carbohydrate molecules (e.g., Rossmann et  
351 al., 1991; Gleixner and Schmidt, 1997) suggests non-random distribution of  $^{13}\text{C}$  in biological  
352 molecules (such as cellobiose). Based on such phenomena, it is probable that the  $^{13}\text{C}$  atom in  
353  $^{13}\text{C}$ -containing cellobiose was consistently at the same position within the molecule, which rules  
354 out the possibility that any changes in the intramolecular  $^{13}\text{C}$  distribution of cellobiose were  
355 responsible for the observed  $\delta^{13}\text{C}$  patterns in biomass and respired  $\text{CO}_2$  (Fig. 4). Hence, the  
356 discontinuous pattern of  $\delta^{13}\text{C}$  of respired  $\text{CO}_2$  with temperature, similar to the pattern for  $\delta^{13}\text{C}$  of  
357 the biomass (Fig. 4A), presumably reflects the downstream consequence of an upstream change  
358 in  $\delta^{13}\text{C}$  of the metabolic substrate taken up and ultimately respired. However, the more negative  
359  $\delta^{13}\text{C}$  of respired  $\text{CO}_2$  compared to that of biomass is, to our knowledge, the most direct evidence  
360 to date for  $^{13}\text{C}$  discrimination during respiration of a heterotrophic microbe. The observation of a  
361 substantial respiratory  $^{13}\text{C}$  discrimination corroborates inferences drawn in earlier studies  
362 (Šantrůčková et al., 2000; Fernandez and Cadisch, 2003) and is also consistent with plant studies  
363 reporting C isotope discrimination during dark respiration in roots (Klumpp et al., 2005;  
364 Bathellier et al., 2009; Ghashghaie and Badeck, 2014). Our observations of respiratory  
365 discrimination against  $^{13}\text{C}$  highlight the similarity of heterotrophic, aerobic respiratory pathways,  
366 and isotope effects within them, across life's domains.

367 In contrast to the discontinuous relationship between biomass  $\delta^{13}\text{C}$  and temperature, we observed  
368 a comparably continuous and linear increase in respiratory discrimination against  $^{13}\text{C}$  with

369 temperature (Fig. 4B). This increase generated a marginally positive significant ( $P=0.08$ )  
370 correlation with SRR (Fig. 6), and hence a marginally significant ( $P=0.07$ ) correlation with CUE.  
371 A physiological interpretation of this finding is not straightforward, as multiple, possibly  
372 simultaneous enzymatic fractionations may have contributed to the observed  $\delta^{13}\text{C}$  of respired  
373  $\text{CO}_2$  (Dijkstra et al., 2011; Tcherkez et al., 2012). It could simply result from a proportionally  
374 increasing flux through respiratory pathways, with associated stronger expression of  $^{13}\text{C}$   
375 discrimination by the enzymes involved (Tcherkez et al., 2012), or could result from increasing  
376 temperatures altering the relative fluxes through respiratory pathways (Chung et al., 1976;  
377 Wittmann et al., 2007; Dijkstra et al., 2011) such that the overall observed respiratory  $^{13}\text{C}$   
378 discrimination increased with temperature. This may be possible given that respiratory pathways  
379 can exhibit distinct fractionation factors (Bathellier et al., 2009 and references therein) and  
380 prompt different, specific C atoms to undergo decarboxylation from the two glucose units of the  
381 substrate cellobiose, which contain non-randomly distributed  $^{13}\text{C}$  atoms (Rossmann et al., 1991;  
382 Gleixner and Schmidt, 1997). If relative fluxes through different respiratory pathways changed  
383 with temperature, the continuous nature of the relationship between temperature and respiratory  
384  $^{13}\text{C}$  discrimination suggests a smooth transition compared to the abrupt and discontinuous shifts  
385 in apparent uptake and/or secretion discrimination described above. Future metabolic flux  
386 analyses linked to isotopic approaches sensitive enough to quantify C isotopes in microbial  
387 exudation will be well-suited to explore how C allocation to distinct, aerobic respiratory  
388 pathways may vary with temperature and result in varying  $\delta^{13}\text{C}$  of respired  $\text{CO}_2$ .

389

#### 390 **4. Conclusions**

391 Our observations clearly show a decline in microbial CUE with increasing temperature when C  
392 substrate is plentiful and demonstrate the mechanism driving it – an increase in SRR. The  
393 relationship between CUE and temperature underscores the importance of incorporating variable,  
394 temperature dependent SRR, which influences CUE and growth efficiency, in ecosystem process  
395 models. The temperature-driven changes in SRR and respiratory discrimination against  $^{13}\text{C}$  were  
396 not independent of each other, suggesting that increasing SRR, to some degree, drives enhanced  
397 C isotopic discrimination. We demonstrate that C isotope discrimination associated with  
398 microbial decomposition of OM can impart large and variable isotopic signatures on C pools  
399 typically characterized and interpreted in biogeochemical studies at any scale. To date, efforts to  
400 partition flux components of net ecosystem exchange have assumed little to no fractionation  
401 between respired substrates and the resultant  $\text{CO}_2$ . Our results suggest that this assumption must  
402 be reevaluated, and represent a first step towards an isotopically explicit, mechanistic framework  
403 for microbial C isotope fluxes in Earth system models.

404

405 **Author contribution:**

406 C.A.L. and K.M. performed the experiments; all authors contributed to all other parts and stages  
407 of the manuscript.

408

409 **Data availability:** The data presented in this study are available for collaborative use by anyone  
410 interested; contact the corresponding author for access to the data.

411

412 **Acknowledgements:**

413 We thank Dr. Susan Ziegler, Dr. Jarad Mellard and Chao Song for helpful discussions during the  
414 design of the experiments, Greg Caine for expert stable isotope analysis and the provision of  
415 isotopic standards, Drs. Susan Ziegler, Mike Burrell, Karl Auerswald, Hanns-Ludwig Schmidt  
416 and John Kelly for comments on the manuscript, Gil Ortiz with assistance generating Figure 5,  
417 anonymous referees for their time and constructive comments, and the National Science  
418 Foundation of the USA for funding (grant no. DEB-0950095 and EAR-1331846).

419

420 **References:**

- 421 Abraham, W.-R., Hesse, C., and Pelz, O.: Ratios of carbon isotopes in microbial lipids as an  
422 indicator of substrate usage, *App. Environ. Microbiol.*, 64, 4202-4209, 1998.
- 423 Allison, S.D., Wallenstein, M.D., and Bradford, M.A.: Soil-carbon response to warming  
424 dependent on microbial physiology, *Nat. Geosci.*, 3, 336-340, 2010.
- 425 Apple, J.K., del Giorgio, P.A., and Kemp, W.M.: Temperature regulation of bacterial production,  
426 respiration, and growth efficiency in a temperate salt-marsh estuary, *Aquat. Microb. Ecol.*, 43,  
427 243-254, 2006.
- 428 Barbosa, Í.C.R., Köhler, I.H., Auerswald, K., Lüps, P., and Schnyder, H.: Last-century changes of  
429 alpine grassland water-use efficiency: a reconstruction through carbon isotope analysis of a time-  
430 series of *Capra ibex* horns, *Glob. Change Biol.*, 16, 1171-1180, 2010.

431 Bathellier, C., Tcherkez, G., Bligny, R., Gout, E., Cornic, G., and Ghashghaie, J.: Metabolic  
432 origin of the  $\delta^{13}\text{C}$  of respired  $\text{CO}_2$  in roots of *Phaseolus vulgaris*, *New Phytol.*, 181, 387-399,  
433 2009.

434 Billings, S.: Soil organic matter dynamics and land use change at a grassland/forest ecotone, *Soil*  
435 *Biol. Biochem.*, 38, 2934-2943, 2006.

436 Blair, N., Leu, A., Muñoz, E., Olsen, J., Kwong, E., and Des Marais, D.: Carbon isotopic  
437 fractionation in heterotrophic microbial metabolism, *Appl. Environ. Microbiol.*, 50, 996-1001,  
438 1985.

439 Bowling, D.R., Pataki, D.E., and Randerson, J.T.: Carbon isotopes in terrestrial ecosystem pools  
440 and  $\text{CO}_2$  fluxes, *New Phytol.*, 178, 24-40, 2008.

441 Brüggemann, N., Gessler, A., Kayler, Z., Keel, S.G., Badeck, F., Barthel, M., Boeckx, P.,  
442 Buchmann, N., Bruognoli, E., Esperschütz, J., Gavrichkova, O., Ghashghaie, J., Gomez-  
443 Casanovas, N., Keitel, C., Knohl, A., Kuptz, D., Palacio, S., Salmon, Y., Uchida, Y., and Bahn,  
444 M.: Carbon allocation and carbon isotope fluxes in the plant-soil-atmosphere continuum: a  
445 review, *Biogeosci.*, 8, 3457-3489, 2011.

446 Bull, A.T.: The renaissance of continuous culture in the post-genomics age, *J. Ind. Microbiol.*  
447 *Biotechnol.*, 37, 993-1021, 2010.

448 Chrzanowski, T.H., and Grover J.P.: Element content of *Pseudomonas fluorescens* varies with  
449 growth rate and temperature: A replicated chemostat study addressing ecological stoichiometry,  
450 *Limnol. Oceanogr.*, 53, 1242-1251, 2008.

451 Chung, B.H., Cannon, R.Y., and Smith, R.C.: Influence of growth temperature on glucose  
452 metabolism of a psychrotrophic strain of *Bacillus cereus*, *Appl. Environ. Microb.*, 31, 39-45,  
453 1976.

454 Cleland, W.W.: The use of isotope effects to determine enzyme mechanisms, *Arch. Biochem.*  
455 *Biophys.*, 433, 2-12, 2005.

456 Cole, J.J., Findlay, S., and Pace M.L.: Bacterial production in fresh and saltwater ecosystems: a  
457 cross-system overview, *Mar. Ecol. Prog. Ser.*, 43, 1-10, 1988.

458 Cotner, J.B., and Biddanda, B.A.: Small players, large role: microbial influence on  
459 biogeochemical processes in pelagic aquatic ecosystems, *Ecosystems*, 5, 105-121, 2002.

460 Cotner, J.B., Makino, W., and Biddanda, B.A.: Temperature affects stoichiometry and  
461 biochemical composition of *Escherichia coli*, *Microb. Ecol.*, 52, 26-33, 2006.

462 Craig, H., and Gordon, L.I.: Deuterium and 18 oxygen variations in the ocean and the marine  
463 atmosphere. In: *Proceedings of a conference on stable isotopes in oceanographic studies and*  
464 *paleotemperatures*. Ed: E. Tongioli, Lischi & Figli, Pisa, 1965.

465 Dawson, P.S.S.: *Microbial growth*, Dowden, Hutchinson & Ross, Stroudsburg, 1974.

466 del Giorgio, P.A., and Cole, J.J.: Bacterial growth efficiency in natural aquatic systems, *Annu.*  
467 *Rev. Ecol. Syst.*, 29, 503-541, 1998.

468 Dijkstra, F.A., Hobbie, S.E., Knops, J.M.H., and Reich, P.B.: Nitrogen stabilization and plant  
469 species interact to influence soil carbon stabilization, *Ecol. Lett.*, 7, 1192-1198, 2004.

470 Dijkstra, P., Thomas, S.C., Heinrich, P.L., Koch, G.W., Schwartz, E., and Hungate, B.A.: Effect  
471 of temperature on metabolic activity of intact microbial communities: evidence for altered  
472 metabolic pathway activity but not for increased maintenance respiration and reduced carbon use  
473 efficiency, *Soil Biol. Biochem.*, 43, 2023-2031, 2011.

474 Egli, T.: Microbial growth and physiology: a call for better craftsmanship, *Front. Microbiol.*,  
475 6:287, DOI: 10.3389/fmicb.2015.00287, 2015.

476 El-Mansi, E.M.T., Holms, W.H.: Control of carbon flux to acetate excretion during growth of  
477 *Escherichia coli* in batch and continuous cultures, *J. General Microbiol.*, 135, 2875-883, 1989.

478 Farquhar, G.D., O'Leary, M.H., and Berry, J.A.: On the relationship between carbon isotope  
479 discrimination and the intercellular carbon dioxide concentration in leaves, *Aust. J. Plant*  
480 *Physiol.*, 9, 121-137, 1982.

481 Farquhar, G.D., and Richards, R.A.: Isotopic composition of plant carbon correlates with water-  
482 use efficiency of wheat genotypes, *Aust. J. Plant Physiol.*, 11, 539-552, 1984.

483 Ferenci, T.: Bacterial physiology, regulation and mutational adaptation in a chemostat  
484 environment, *Adv. Microb. Physiol.*, 53, 169-230, 2008.

485 Fernandez, I., and Cadisch, G.: Discrimination against  $^{13}\text{C}$  during degradation of simple and  
486 complex substrates by two white rot fungi, *Rapid Commun. Mass Spectrom.*, 17, 2614-2620,  
487 2003.

488 Frey, S., Lee, J., Melillo, J.M., and Six, J.: The temperature response of soil microbial efficiency  
489 and its feedback to climate, *Nat. Clim. Change*, 3, 395-398, 2013.

490 Fry, B.: *Stable Isotope Ecology*, Springer, New York, 2006.

491 Gamnitzer, U., Moyes, A.B., Bowling, D.R., and Schnyder, H.: Measuring and modeling the  
492 isotopic composition of soil respiration: insights from a grassland tracer experiment, *Biogeosci.*,  
493 8, 333-1350, 2011.

494 Ghashghaie, J., and Badeck, F.W.: Opposite carbon isotope discrimination during dark  
495 respiration in leaves versus roots – a review, *New Phytol.*, 201, 751-769.

496 Gillooly, J.F., Brown, J.H., West, G.B., Savage, V.M., and Charnov, E.L.: Effects of size and  
497 temperature on metabolic rate, *Science*, 293, 2248-2251, 2001.

498 Gleixner, G., and Schmidt H.-L.: Carbon isotope effects on the Fructose-1,6-bisphosphate  
499 aldolase reaction, origin for non-statistical  $^{13}\text{C}$  distributions in carbohydrates, *J. Biol. Chem.*,  
500 272, 5382-5387, 1997.

501 Goldman, J.C., and Denner, M.R.: Growth of marine bacteria in batch and continuous culture  
502 under carbon and nitrogen limitation, *Limnol. Oceanogr.*, 45, 789-800, 2000.

503 Hall, E.K., Neuhauser, C., and Cotner, J.B.: Toward a mechanistic understanding of how natural  
504 bacterial communities respond to changes in temperature in aquatic ecosystems. *ISME J*, 2, 471-  
505 481, 2008.

506 Hanson, P.J., Edwards, N.T., Garten, C.T., and Andrews, J.A.: Separating root and microbial  
507 contributions to soil respiration: a review of methods and observations. *Biogeochem.*, 48, 115-  
508 146, 2000.



509 Hedges, J.I., Eglinton, G., Hatcher, P.G., Kirchman, D.L., Arnosti, C., Derenne, S., Evershed,  
510 R.P., Kögel-Knabner, I., de Leeuw, J.W., Littke, R., Michaelis, W., Rullkötter, J.: The  
511 molecularly-uncharacterized component of nonliving organic matter in natural environments,  
512 *Org. Geochem.*, 31, 945-958, 2000.

513 Hobbie, J.E., and Hobbie, E.A.: Microbes in nature are limited by carbon and energy: the  
514 starving-survival lifestyle in soil and consequences for estimating microbial rates, *Front.*  
515 *Microbiol.*, 4, doi: 10.3389/fmicb.2013.00324, 2013.

516 Kayser, A., Weber, J., Hecht, V., and Rinas, U.: Metabolic flux analysis of *Escherichia coli* in  
517 glucose-limited continuous culture. I. Growth-rate dependent metabolic efficiency at steady-  
518 state, *Microbiol.*, 151, 693-706, 2005.

519 Kirschbaum, M.U.F.: The temperature-dependence of organic-matter decomposition – still a  
520 topic of debate, *Soil Biol. Biochem.*, 38, 2510-2518, 2006.

521 Klumpp, K., Schäufele, R., Lötscher, M., Lattanzi, F.A., Feneis, W., and Schnyder, H.: C-isotope  
522 composition of CO<sub>2</sub> respired by shoots and roots: fractionation during dark respiration?, *Plant*  
523 *Cell Environ.*, 28, 241-250, 2005.

524 Kucera, C.L., and Kirkham, D.R.: Soil respiration studies in tallgrass prairie in Missouri,  
525 *Ecology*, 52, 912-915, 1971.

526 Lehmeier, C.A., Min, K., Niehues, N.D., Ballantyne, F.IV., and Billings, S.A.: Temperature-  
527 mediated changes of exoenzyme-substrate reaction rates and their consequences for the carbon to  
528 nitrogen flow ration of liberated resources, *Soil Biol. Biochem.*, 57, 374-382, 2013.

529 Manzoni, S., Taylor, P., Richter, A., Porporato, A., and Ågren, G.I.: Environmental and  
530 stoichiometric controls on microbial carbon-use efficiency in soils, *New Phytol.*, 196, 79-91,  
531 2012.

532 Min, K., Lehmeier, C.A., Ballantyne, F., Tatarko, A., and Billings, S.A.: Differential effects of  
533 pH on temperature sensitivity of organic carbon and nitrogen decay, *Soil Biol. Biochem.*, 76,  
534 193-200, 2014.

535 Mook, W.G., Bommerson, J.C., and Staverman, W.H.: Carbon isotope fractionation between  
536 dissolved bicarbonate and gaseous carbon dioxide, *Earth Planet. Sci. Lett.*, 22, 169-176, 1974.

537 Nanchen, A., Schicker A., and Sauer U.: Nonlinear dependency of intracellular fluxes on growth  
538 rate in miniaturized continuous cultures of *Escherichia coli*, *Appl. Environ. Microb.*, 73, 1164-  
539 1172, 2006.

540 Nickerson, N., Egan, J., and Risk, D.: Iso-FD: A novel method for measuring the isotopic  
541 signature of surface flux, *Soil Biol. Biochem.*, 62, 99-106, 2013.

542 O'Leary, M.H.: Carbon isotope fractionation in plants, *Phytochem.*, 20, 553-567, 1981.

543 Park, R., and Epstein, S.: Metabolic fractionation of  $C^{13}$  &  $C^{12}$  in plants, *Plant Physiol.*, 36, 133-  
544 138, 1961.

545 Pataki, D.E., Ehleringer, J.E., Flanagan, L.B., Yakir, D., Bowling, D.R., Still, C.J., Buchmann,  
546 N., Kaplan, J.O., and Berry, J.A.: The application and interpretation of Keeling plots in terrestrial  
547 carbon cycle research, *Glob. Biogeochem. Cycles*, 17, doi:10.1029/2001GB001850, 2003.

548 Pomeroy, L.R., and Wiebe, W.J.: Temperature and substrate as interactive limiting factors for  
549 marine heterotrophic bacteria, *Aquat. Microb. Ecol.*, 23, 187-204, 2001.

550 Rivkin, R.B., and Legendre, L.: Biogenic carbon cycling in the upper ocean: effects of microbial  
551 respiration, *Science*, 291, 2398-2400, 2001.

552 Rossmann, A., Butzenlechner, M., and Schmidt, H.L.: Evidence for a nonstatistical carbon  
553 isotope distribution in natural glucose, *Plant Physiol.*, 96, 609-614, 1991.

554 Šantrůčková, H., Bird, M.I., and Lloyd, J.: Microbial processes and carbon-isotope fractionation  
555 in tropical and temperate grassland soils, *Funct. Ecol.*, 14, 108-114, 2000.

556 Schimel, D.S.: Terrestrial ecosystems and the carbon-cycle, *Glob. Change Biol.*, 1, 77-91, 1995.

557 Smith, H.L., and Waltman, P.E.: The theory of the chemostat: dynamics of microbial  
558 competition, Cambridge University Press, New York, 1995.

559 Stumm, W., and Morgan, J.J.: *Aquatic Chemistry: An introduction emphasizing chemical*  
560 *equilibria in natural waters*, John Wiley & Sons, New York, 1981.

561 Subke, J.-A., Inglima, I., and Cotrufo, M.F.: Trends and methodological impacts in soil CO<sub>2</sub> flux  
562 partitioning: A metaanalytical review, *Glob. Change Biol.*, 12, 921-943, 2006.

563 Szaran, J.: Achievement of carbon isotope equilibrium in the system HCO<sub>3</sub><sup>-</sup> (solution) – CO<sub>2</sub>  
564 (gas). *Chem. Geol.*, 142, 79-86, 1997.

565 Tcherkez, G., Mahé, A., and Hodges, M.: <sup>12</sup>C/<sup>13</sup>C fractionations in plant primary metabolism,  
566 *Trends Plant Sci.*, 16, 499-506, 2012.

567 Tempest, D.W., and Neijssel, O.M.: Eco-physiological aspects of microbial growth in aerobic  
568 nutrient-limited environments, *Adv. Microb. Ecol.*, 2, 105-153, 1978.

569 Trumbore, S.: Carbon respired by terrestrial ecosystems – recent progress and challenges, *Glob.*  
570 *Change Biol.*, 12, 141-153, 2006.

571 Tucker, C.L., Bell, J., Pendall, E., and Ogle, K.: Does declining carbon-use efficiency explain  
572 thermal acclimation of soil respiration with warming?, *Glob. Change Biol.*, 19, 252-263, 2013.

573 Vogel, J.C., Grootes, P.M., and Mook, W.G.: Isotopic fractionation between gaseous and  
574 dissolved carbon dioxide, *Z. Physik*, 230, 225-238, 1970.

575 Werner, C., and Gessler, A.: Diel variations in the carbon isotope composition of respired CO<sub>2</sub>  
576 and associated carbon sources: a review of dynamics and mechanisms, *Biogeosci.*, 8, 2437-2459,  
577 2011.

578 Werner, R.A., Buchmann, N., Siegwolf, R.T.W., Kornexl, B.E., and Gessler, A.: Metabolic  
579 fluxes, carbon isotope fractionation and respiration – lessons to be learned from plant  
580 biochemistry, *New Phytol.*, 191, 10-15, 2011.

581 Werth, M., and Kuzyakov, Y.: <sup>13</sup>C fractionation at the root-microorganisms-soil interface: a  
582 review and outlook for partitioning studies, *Soil Biol. Biochem.*, 42, 1372-1384, 2010.

583 Wieder, W.R., Bonan, G.B., and Allison, S.D.: Global soil carbon projections are improved by  
584 modelling microbial processes, *Nat. Clim. Change*, 3, 909-912, 2013.

585 Wittmann, C., Weber, J., Betiku, E., Krömer, J., Böhm, D., and Rinas, U.: Response of fluxome  
586 and metabolome to temperature-induced recombinant protein synthesis in *Escherichia coli*, *J.*  
587 *Biotechnol.*, 132, 375-384, 2007.

588

589

590

591

592

593

594

595

596

597

598

599

600

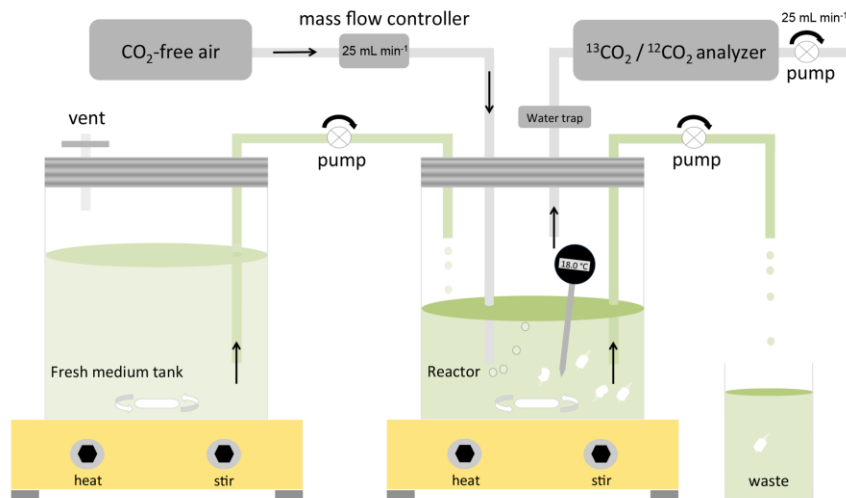
601

602

603 **Figures:**

604 **Fig. 1.** Chemostat system comprised of *P. fluorescens* growing on cellobiose. Seven independent  
605 experiments were conducted, with reactor temperatures of 13, 14.5, 16, 18, 21, 23.5 and 26.5 °C;  
606 all other conditions were identical. During continuous flow, dilution rate of the reactor  
607 (mean=0.137±0.01 h<sup>-1</sup> across all experiments) equals microbial growth rate. A peristaltic pump  
608 supplied fresh nutrient medium from a reservoir tank to the reactor and removed reactor medium  
609 (including biomass) at a constant rate. Headspace volume was flushed with CO<sub>2</sub>-free air,  
610 bubbling through reactor medium and supplying microorganisms with O<sub>2</sub>. A <sup>13</sup>CO<sub>2</sub>/<sup>12</sup>CO<sub>2</sub>  
611 analyzer continuously sampled reactor headspace and measured the concentration and δ<sup>13</sup>C of  
612 respired CO<sub>2</sub>.

613

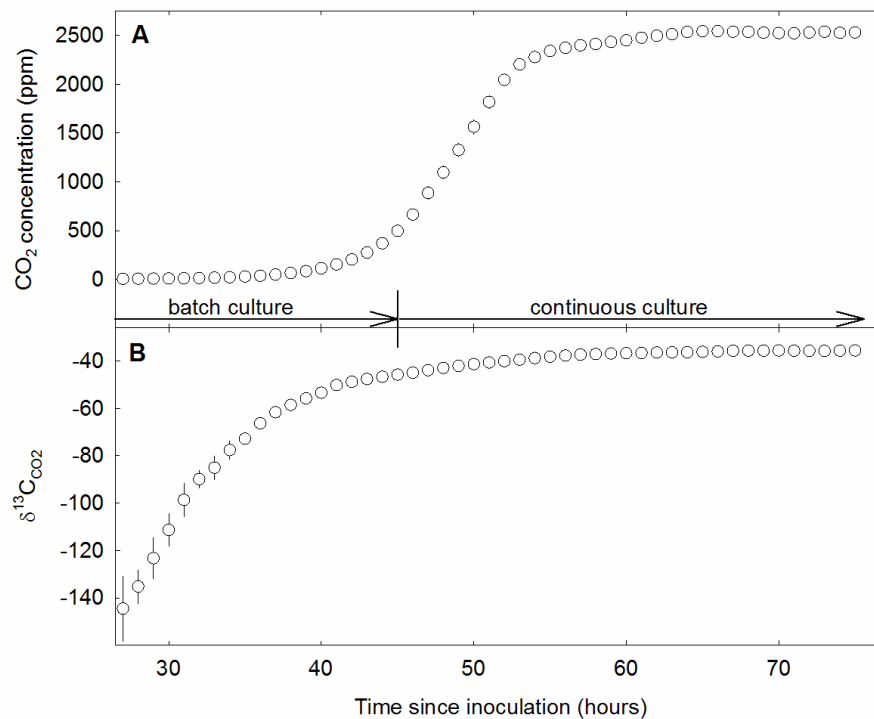


614

615

616 **Fig. 2.** Example time course of the evolution of reactor headspace CO<sub>2</sub> concentration (A) and  
617  $\delta^{13}\text{C}$  of the CO<sub>2</sub> (B) of the chemostat run at 23.5 °C in hours since inoculation of the reactor with  
618 pre-cultured *P. fluorescens*. Data points are hourly means. Error bars (where visible) denote  $\pm 1$   
619 SD. The reactor was shifted from batch to continuous culture mode 45 h after inoculation.  
620 Microbial respiration rate and the  $\delta^{13}\text{C}$  of respired CO<sub>2</sub> were measured between 70 and 74 h after  
621 inoculation when the culture reached steady-state.

622

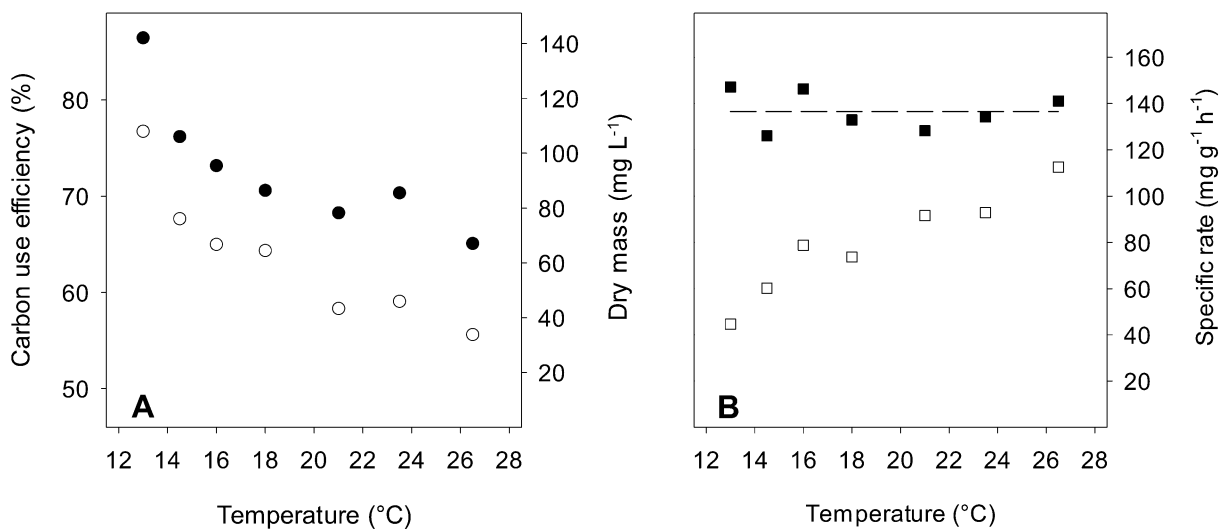


623

624

625 **Fig. 3.** Steady-state process variables of *P. fluorescens* growing in chemostats at specified  
 626 temperatures. Microbial carbon use efficiency (○; A), dry microbial biomass (●; A), specific C  
 627 growth rate (■; B), and specific C respiration rate (□; B), expressed per unit of microbial  
 628 biomass-C. The dashed line denotes the average of the seven specific growth rates ( $137 \text{ mg g}^{-1}$   
 629  $\text{h}^{-1}$ ).

630



631

632

633

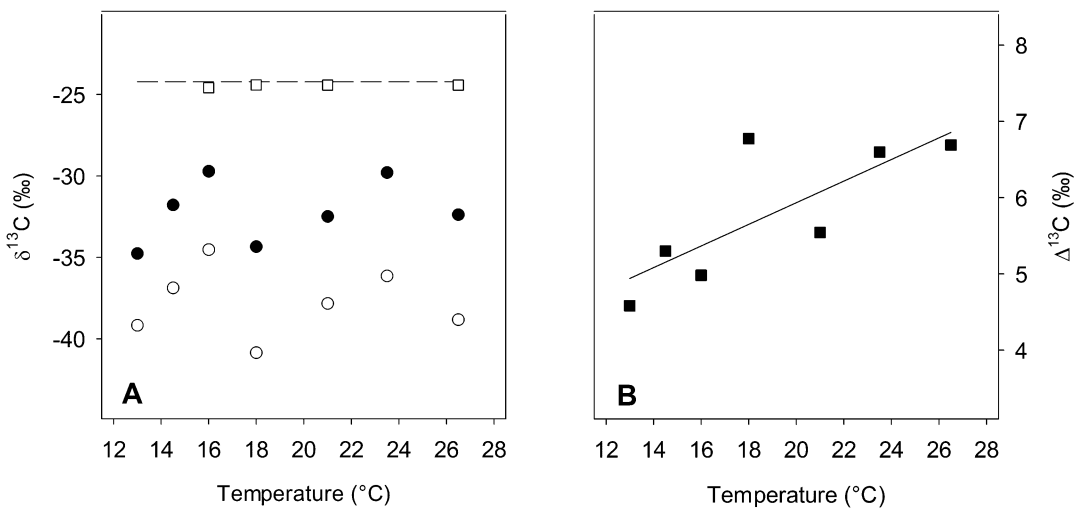
634

635



636 **Fig. 4.** Steady-state  $\delta^{13}\text{C}$  of microbial biomass ( $\bullet$ ; A) and of respired  $\text{CO}_2$  ( $\circ$ ; A), and C isotope  
637 discrimination during respiration ( $\Delta^{13}\text{C}$ ; B) of *P. fluorescens* growing in chemostats at specified  
638 temperatures. In panel A, the dashed line denotes the  $\delta^{13}\text{C}$  of the substrate cellobiose (-24.2‰),  
639 and  $\delta^{13}\text{C}$  of reactor filtrate is shown as open squares; standard errors, derived from multiple  
640 measurements across time during steady-state, are smaller than the size of the symbols.  $\Delta^{13}\text{C}$  is  
641 calculated as  $\Delta^{13}\text{C} = (\delta^{13}\text{C}_{\text{biomass}} - \delta^{13}\text{C}_{\text{respired CO}_2}) / (1 + \delta^{13}\text{C}_{\text{respired CO}_2})$ . The solid line denotes  
642 linear regression of  $\Delta^{13}\text{C}$  vs. temperature ( $y = 0.14x + 3.1$ ;  $R^2 = 0.61$ ;  $P=0.04$ ).

643



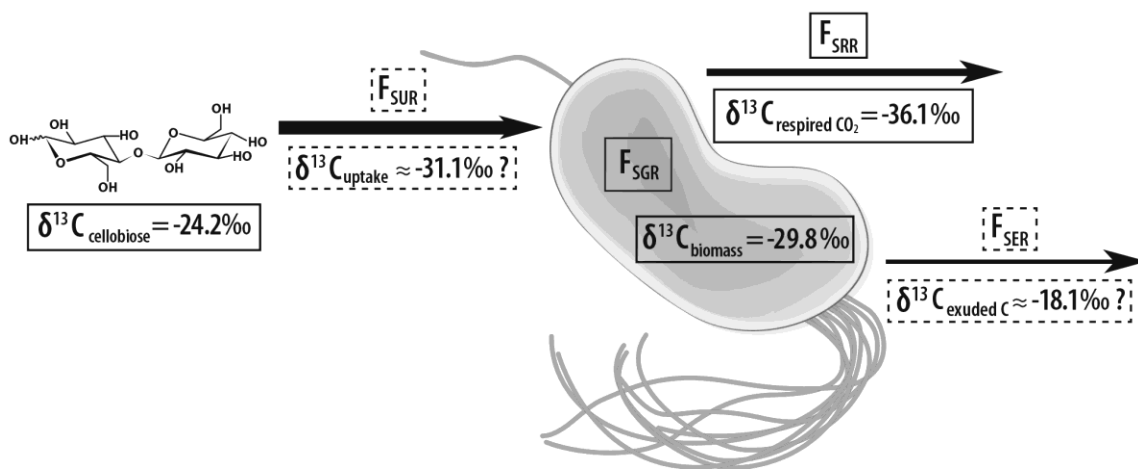
644

645

646

647

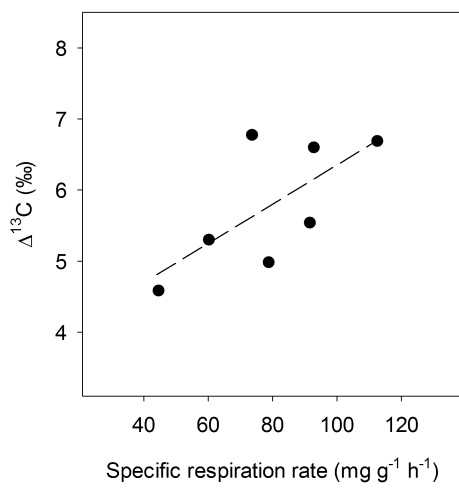
648 **Fig. 5.** Schematic of an individual *P. fluorescens* cell, representing the sample population  
 649 growing at 23.5 °C with one available substrate at a constant relative growth rate of 0.13 h<sup>-1</sup> at  
 650 steady-state (Fig. 2), with measured (solid boxed) and estimated (dashed boxed) magnitudes of C  
 651 and <sup>13</sup>C fluxes into and out of the population. Designated fluxes include specific uptake rate of  
 652 cellobiose (F<sub>SUR</sub>), specific growth rate (F<sub>SGR</sub>), specific respiration rate (F<sub>SRR</sub>) and specific  
 653 excretion rate (F<sub>SER</sub>), in relation to steady-state biomass-C in the chemostat, where F<sub>SUR</sub> = F<sub>SGR</sub> +  
 654 F<sub>SRR</sub> + F<sub>SER</sub>. The estimate of δ<sup>13</sup>C<sub>uptake</sub> is based on the assumption that δ<sup>13</sup>C<sub>exuded C</sub> is 11.7‰ less  
 655 negative than δ<sup>13</sup>C<sub>biomass</sub> (Blair et al., 1985) and that F<sub>SER</sub> is 10% of the sum of F<sub>SGR</sub> and F<sub>SRR</sub> (El-  
 656 Mansi and Holms, 1989). In contrast to this experimental system, in natural environments  
 657 measurements of boxed pools and fluxes can be confounded by the presence of dormant  
 658 microorganisms, unknown microbial growth rates, diverse available substrates, and a lack of  
 659 steady-state CO<sub>2</sub> fluxes.



662

663 **Fig. 6.** Correlation between the specific respiration rate of *P. fluorescens* growing in continuous  
664 chemostat culture at temperatures ranging from 13 °C to 26.5 °C and the carbon isotope  
665 discrimination during respiration. The dashed line denotes a linear regression of the form  $y =$   
666  $0.03x + 3.6$ ;  $R^2 = 0.48$ ;  $P=0.08$ .

667



668