Author's response to Dr. G. Darnis

We are grateful to your comments and useful suggestions that improved our manuscript greatly. As described below, we have revised our manuscript. Please note that the expression in black colored letters are the ones provided by you whereas those in red are our replies.

Specific comments

Title

Could it better reflect the changes related to the strong wind event that occurred during the field work? On the other hand, I understand that the chosen title fits with that of a companion paper addressing the case of microplankton community during the same field campaign.

Answer: We changed the title to "Short-term changes of the mesozooplankton community and copepod gut pigment in the Chukchi Sea in autumn: Reflection of strong wind event" (marked manuscript p1 L1-3).

Abstract

The abstract is too wordy. The content could be synthesized, and multiple uses of words limited. A lot of results reported but few conclusions in comparison. Answer: We shortened the abstract (marked manuscript p1 L16-p2 L14).

Line 6: Change "high-frequent" to "high-frequency", here and elsewhere in the text. Answer: We changed to "high-frequency" through the manuscript (marked manuscript p1 L20, p3 L4, L12, L19).

Line 12: "dominant" and not "most dominated" Answer: We changed to "dominant" (marked manuscript p1 L26).

Line 20: If you decide to keep this part add the value of the C:Chl a ratio, although I don't think that this is necessary to mention in the abstract the assumption of this ratio in the calculation of the estimate.

Answer: We deleted "Assuming C:Chl. a ratio" (marked manuscript p2 L4).

Introduction

More emphasis should be put on Calanus glacialis in the Introduction since it is clearly

the dominant species in terms of biomass and one focus of the study with the gut content analysis. My suggestion is to move one paragraph on this species from the Discussion to the Introduction.

Answer: We moved one paragraph of *C. glacialis* from Discussion to Introduction (marked manuscript p2 L20-26).

Line 6. Mesozooplankton are secondary producers everywhere. This trivial statement can be removed.

Answer: We deleted "is a secondary producer and" (marked manuscript p2 L17).

Line 11: "dominate" and not "dominated" Answer: We deleted this part owing above change (marked manuscript p2 L30).

Line 13: Maybe "origin" would be better here than "formation" Answer: We deleted this part owing above change (marked manuscript p2 L31).

Materials and methods Page 6, line 6: replace "with" by "of" Answer: We changed to "by" (marked manuscript p5 L7).

Page 6, line 8: replace "with" by "to" Answer: We changed to "to" (marked manuscript p5 L9).

Results

For section 3.1, I am not sure that "Hydrography" is the right title since chl a, a biological feature, is described there also. Description of vertical distribution of chl a (Figure 2c) could be relevant for the study.

Answer: We changed the title as "Hydrography and chlorophyll *a*" (marked manuscript p6 L11).

Section 3.3 on *Calanus glacialis* I would start by describing population structure instead of DVM right away. Swap the paragraphs.

Answer: We swapped the paragraphs (marked manuscript p7 L7-18). Owing this change in text, Figs. 3 and 4 are also swapped (Figs. 3 and 4).

Detail the composition of the population by giving the percentage of other stages than

CV. Mention that the young copepodites abundances are underestimated due to the coarse mesh net used.

Answer: Concerning underestimation of small zooplankton, we added short note in Materials and methods (marked manuscript p4 L1-2).

Page 8, line 16: "their population" by "the population" Answer: We changed to "the population" (marked manuscript p7 L10).

Discussion

Page 9, line 8: Greater spatial and temporal change than what? I did not get the meaning of the sentence. It should be clarified. In fact, the whole paragraph needs to be rewritten as it's difficult to differentiate what's comes from this study and what are Matsuno et al (2012) findings.

Answer: We changed the word from "greater" to "large" (marked manuscript p8 L11). Concerning comparison with Matsuno et al. (2012), we rewrote this part (marked manuscript p8 L12-28).

The structure of the sentences should be simpler and more direct. Most sentences should start simply with the subject for sake of clarity. As an example: "Comparing the above characteristics by Matsuno et al. (2012), the zooplankton abundance of this study was nearly half (mean: 34 059 ind.m⁻²), there was a low abundance of small copepod *Pseudocalanus* spp. and cyclopoid copepods, and no occurrence of Arctic copepod *Metridia longa* was remarkable (Table 1)." could well be: "Total zooplankton abundance in this study was approximately half (mean: 34 059 ind.m⁻²) the abundance reported by Matsuno et al. (2012) on the Chukchi shelf (mean: 75 683 ind.m-2), with low abundance of small copepods (*Pseudocalanus* spp. and cyclopoids) and the remarkable absence of the Arctic copepod *Metridia longa*."

Answer: We changed the sentence following referee's comment (marked manuscript p8 L12-16).

Page 10, line 5: Again, this sentence is not clear. What does the addition of holoplankton mean?

Answer: We deleted the sentence "the addition of holoplankton and" (marked manuscript p9 L10).

Page 10, line 8-11: "Benthic barnacle adults released their larvae when they met

phytoplankton blooms (Crisp, 1962; Clare and Walker, 1986), and their larvae spent two to three weeks at water columns and then settled (Herz, 1933)." If this is usual behavior on the part of adult barnacles, present tense should be used. Furthermore, replace "at water columns" by "in the water column".

Answer: We changed to "Benthic barnacle adults release their larvae when they meet phytoplankton blooms (Crisp, 1962; Clare and Walker, 1986), and their larvae spend two to three weeks in the water column and then settle (Herz, 1933)." (marked manuscript p9 L13-15).

Page 10, line 17: replace "several limited" by "a few" Answer: We changed to "a few" (marked manuscript p9 L21).

Page 10, line 18: this issue has been already addressed in the previous paragraph. Answer: We deleted the sentence on barnacle larvae (marked manuscript p9 L22-23).

Section 4.2 Population structure of Calanus glacialis

The first paragraph belongs to the Introduction section. It's a description of the status and life cycle of this important arctic shelf copepod, which is a focus of the work. Answer: We moved this paragraph into Introduction (marked manuscript p10 L2-11).

Page 13, line 9-11. What is the value measured by Tande and Bamstedt (1985)? Why make this comparison with the situation in the Barents Sea in spring-summer if it's not interpreted further. Wouldn't it be more relevant to try a comparison with grazing impacts estimated by Campbell et al (2009) for roughly the same region? At least conclusions of this study should be better addressed in the present work.

Answer: We deleted sentence comparing with Tande and Bamstedt (1985) (marked manuscript p11 L31-32).

Page 13, line 18. Change "proportion" for "potential contribution" Answer: We changed to "potential contribution" (marked manuscript p12 L6-7).

Conclusion

Page 14, line 3. In fact, grazing impacts was only estimated for the dominant stage of *C*. *glacialis*, excluding other dominant copepods such as *Pseudocalanus*. The sentence should be changed accordingly.

Answer: We changed from "dominant copepods" to "C. glacialis" (marked manuscript

p12 L20).

Page 14, line 12. This conclusion is rather trivial. We certainly could not expect metazoan plankton demography to respond so fast to a short surge in phytoplankton stock. However, it would be particularly interesting to speculate on the cumulative effects of late summer-early autumn strong winds in a region more and more impacted by the reduction in sea ice cover. Could some zooplankton benefit from an extension of the primary production season with more turbulence and later freeze-up of the Chukchi Sea?

Answer: Yes, the case is expected. Owing to comment, we added short note on this issue at end of the conclusion (marked manuscript p12 L29-31).

Tables

Table 2. Some information given in the legend for the calculations should be moved to the Materials and methods section.

Answer: Since this Table 2 is only cited in Discussion (note that not in Results), we remained that the legend of this Table as it was.

Figure 1. No need to mention in the legend that the depth contours are superimposed since it's obvious.

Answer: We deleted the sentence (marked manuscript p19 L3).

Author's response to Dr. V. Dvoretsky

We are grateful to your comments and useful suggestions that improved our manuscript greatly. As described below, we have revised our manuscript. Please note that the expression in black colored letters are the ones provided by you whereas those in red are our replies.

Specific comments.

Abstract L 11. 'ranged 23 610–56 809 ind.m⁻²' replace by 'ranged from 23 610 to 56 809 ind.m⁻²'

Answer: We changed to "ranged from 23 610 to 56 809 ind. m^{-2} " (marked manuscript p1 L24-25).

Results L 21 'ranged 23 610–56 809 ind.m⁻²' replace by 'ranged from 23 610 to 56 809 ind.m⁻²'

Answer: We changed to "ranged from 23 610 to 56 809 ind. m^{-2} " (marked manuscript p6 L23).

Methods

In general, methods are good described. Therefore I suggest to mention that NORPAC nets (mouth diameter 45 cm, mesh size 335 μ m) are rather coarse to quantitatively catch smaller zooplankton forms (e.g. Pseudocalanus spp. C1-4, Cyclopoida, the youngest stages of other copepods) and there may be some underestimations of these zooplankton taxa.

Amswer: We added short note on the underestimation of small copepods in the Materials and Methods (marked manuscript p4 L1-2).

Results

I recommend calculating the Shannon diversity index and the Pielou evenness of the zooplankton communities. These indices may give interesting information on the shortterm dynamics of zooplankton structure in the Chukchi Sea. Estimated zooplankton biodiversity should be described in the Results in detail, and therefore compared (if possible) with previous studies in the Discussion section.

Answer: We calculated both Shannon species diversity index and Pielou evenness and compared with the reported values in this region (marked manuscript p5 L22-28, p6 L27-28, p7 L1-2, p9 L5-8).

Short-term changes of the mesozooplankton community and copepod gut pigment in the Chukchi Sea in autumn: <u>Reflection of strong wind event</u>

4

5 K. Matsuno^{1*}, A. Yamaguchi², S. Nishino³, J. Inoue¹ and T. Kikuchi³

6 [1]{Arctic Environment Research Center, National Institute of Polar Research, 10-3 Midori-

7 cho, Tachikawa, Tokyo 190-8518, Japan}

8 [2]{Laboratory of Marine Biology, Graduate School of Fisheries Sciences, Hokkaido
9 University, 3-1-1 Minato-cho, Hakodate, Hokkaido 041-8611, Japan}

- 10 [3] [Japan Agency for Marine-Earth Science and Technology, 2-15 Natsushima-cho,
- 11 Yokosuka, Kanagawa 237-0061, Japan}
- 12

13 * Correspondence to: K. Matsuno (matsuno.kohei@nipr.ac.jp, k.matsuno@fish.hokudai.ac.jp)

14

15 Abstract

In the Chukehi Sea, due to the recent drastic reduction of sea-ice during the summer, an 16 increasing formation of atmospheric turbulence has been reported. However, the importance 17 and effects of atmospheric turbulence on the marine ecosystem are not fully understood in this 18 19 region. To evaluate the effect of atmospheric turbulence on the marine ecosystem, high-20 frequencyt sampling (two to four times per day) on the mesozooplankton community and the 21 gut pigment of dominant copepods were made at a fixed station in the Chukchi Sea from 10 to 22 25 September 2013. During the study period, a strong wind event (SWE) was observed on 18 September. After the SWE, the standing stock of chlorophyll a (chl. a) was increased, 23 especially for micro-size (> 10 µm) fractions. Zooplankton abundance ranged from 23610 to 24 =56809 ind. m⁻² and exhibited no clear changes with SWE. In terms of abundance, calanoid 25 copepods constituted the most-dominanted taxa (mean: 57%), followed by barnacle larvae 26 27 (31%). Within the calanoid copepods, small-sized Pseudocalanus spp. (65%) and large-sized 28 Calanus glacialis (30%) dominated. In the population structure of C. glacialis, copepodid 29 stage 5 (C5) dominated, and the mean copepodid stage did not vary with SWE. The

dominance of accumulated lipids in C5 and C6 females with immature gonads indicated that 1 2 they were preparing for seasonal diapause. The gut pigment of C. glacialis C5 was higher at 3 night and was correlated with ambient chl. a, and a significant increase was observed after SWE (2.6 vs 4.5 ng pigment ind.⁻¹). Assuming C:Chl. *a* ratio, t_{T} he grazing impact by C. 4 glacialis C5 was estimated to be 4.14 mg C m^{-2} day⁻¹, which corresponded to 0.5–4.6% of 5 the standing stock of micro-size phytoplankton. Compared with the metabolic food 6 7 requirement, their feeding on phytoplankton accounted for 12.6% of their total food 8 requirement. These facts suggest that C. glacialis could not maintain their population on 9 solely phytoplankton food, and other food sources (i.e., microzooplankton) are important in 10 autumn. As observed for the increase in gut pigment, temporal phytoplankton bloom, which 11 is enhanced by the atmospheric turbulence (SWE) in autumn, may have a positive effect on 12 copepod nutrition. However, because of the relatively long generation length of copepods, a smaller effect was detected for their abundance, population structure, lipid accumulation and 13 14 gonad maturation within the short term period (16 days).

15

16 **1 Introduction**

17 In marine ecosystems of the western Arctic Ocean, mesozooplankton is a secondary producer 18 and an important food resource for pelagic fishes and whales (Lowry et al., 2004; Ashjian et 19 al., 2010). In terms of biomass, mesozooplankton in the western Arctic Ocean is dominated 20 by Arctic copepods. Within the Arctic copepods, Calanus glacialis is a key species that 21 dominates in the zooplankton biomass and commonly occurs in the continental shelf 22 throughout the Arctic Ocean (Conover and Huntley, 1991; Lane et al., 2008). For the life 23 cycle of C. glacialis, they grow to C3-C4 at the epipelagic layer during the first summer, then 24 descend to the deep layer and enter diapause, developing to C5 at the epipelagic layer in the 25 second summer, then down to the deep layer and molting to the adult stage (C6) using stored lipid as energy (Longhurst et al., 1984; Ashjian et al., 2003; Falk-Petersen et al., 2009).: i.e., 26 27 Calanus elacialis, Calanus hyperboreus, Metridia longa and Pseudocalanus spp. In the 28 southern Chukehi Sea, large Pacific copepods (e.g., Calanus marshallae, Neocalanus 29 cristatus, Neocalanus flemingeri, Neocalanus plumchrus, Eucalanus bungii and Metridia 30 pacifica) dominated in biomass because of the inflow of Pacific water. Thus, the zooplankton 31 community in this region is greatly governed by water mass formation (Hoperoft et al., 2010; 32 Matsuno et al., 2011, 2012). Concerning seasonal change, a comparison of the zooplankton

community between summer and autumn (Llinás et al., 2009) and year-round changes based
 on the two-week sampling interval at a drifting ice station in the Arctic basin were reported
 (Ashjian et al., 2003). However, short-term changes in the zooplankton community based on

4 high-frequencyt sampling (two to four times per day) have not been reported.

5 Recently, a drastic reduction in the area of sea ice has been observed in the Arctic 6 Ocean during summer, and the sea ice reduction was greatest in the western Arctic Ocean 7 (Stroeve et al., 2007; Comiso et al., 2008; Markus et al., 2009). Furthermore, increases in 8 frequency and magnitude of cyclones and a northward shift of their track during summer have 9 been reported in recent years (Serreze et al., 2000; Cabe et al., 2001; Zhang et al., 2004; Sepp 10 and Jaagus, 2011). While the importance of such changes is clear, little information is 11 available regarding their effect on atmospheric turbulence in marine ecosystems in the western Arctic Ocean. From 10 to 25 September 2013, high-frequencyt sampling/observation 12 13 was made at a fixed station in the Chukchi Sea, and the occurrence of strong wind events 14 (SWE), vertical flux of nutrients and changes in primary production and microplankton 15 communities were reported (Nishino et al., 2015 in press; Yokoi et al., submitted in preparation). However, no information is available regarding how the mesozooplankton community 16 17 responds to such atmospheric turbulence and oceanic environmental changes.

In the present study, we evaluated short-term changes of the mesozooplankton community in the Chukchi Sea during autumn based on the high-frequencyt samplings made simultaneously by Nishino et al. (<u>2015in press</u>) and Yokoi et al. (<u>submittedin preparation</u>). We estimated the grazing impact of dominant copepods based on their gut pigments and evaluated the effect of SWE (Nishino et al., <u>2015in press</u>) and short-term changes of microplankton assemblages (Yokoi et al., <u>submittedin preparation</u>) on the mesozooplankton community in the Chukchi Sea in autumn.

25

26 2 Materials and Methods

27 2.1 Field sampling

28 Zooplankton samplings were conducted at a fixed station in the Chukchi Sea (72°45′N, 29 168°15′W, depth 56 m) from 10 to 25 September 2013 (Fig. 1) (Nishino et al., 2015 in press). 30 Zooplankton samples were collected by vertical hauls with a NORPAC net (mouth diameter 31 45 cm, mesh size 335 μ m; Motoda, 1957) and ring net (mouth diameter 80 cm, mesh size 335 32 μ m) from 49 m depth to the sea surface two to four times per day (total of 47 times) including **書式変更:**フォント : 12 pt

書式変更:フォント: 12 pt

both day and night. The large mesh size of NORPAC net (335 µm) may induce somewhat 1 2 underestimation of smaller zooplankton species and early stages of larger zooplankton species. Zooplankton samples collected by NORPAC nets were preserved with 5% buffered formalin 3 4 immediately on board. The ring net samples were used for copepod gut pigment 5 measurement. For evaluation of the diel vertical migration (DVM) of copepods, day and night vertical stratified hauls with closing PCP nets (mouth diameter 45 cm, mesh size 62 µm; 6 7 Kawamura, 1989) from two layers (0-20 m and 20-49 m) were made on 14 and 22 8 September. The samples from the PCP net were preserved with 5% buffered formalin.

At the fixed station, temperature, salinity and chlorophyll *a* (chl. *a*) fluorescence were measured by CTD (Sea-Bird Electronics Inc., SBE911Plus) casts at a frequency of two to four times per day. To evaluate size fractionated chl. *a*, water samples from the sea surface and the maximum fluorescence layer (16.8-27.7 m depth) were collected by a bucket and rosette multi-sampler mounted on CTD, respectively. Water samples were filtered through 20, 10 and 2-µm pore-size membranes and GF/F filters, and chl. *a* was extracted with N,Ndimethylformamide and measured with a fluorometer (Turner Designs, Inc., 10-AU-005).

16 **2.2 Gut pigment**

17 For fresh samples collected by ring nets, 10% v/v soda (saturated CO₂ in water) was added to 18 avoid copepod grazing, gut evacuation and decomposition of gut pigment. Fresh specimens 19 of C. glacialis copepodid stage 5 (C5) were sorted under a stereomicroscope. Batches of 20 fifteen specimens were immersed in 6 ml N,N-dimethylformamide and stored in dark, cold 21 conditions overnight to extract chlorophyll and phaeopigments. After the extraction of 22 pigment, the chlorophyll and phaeopigments were measured with a fluorometer (Turner 23 Designs, Inc., 10-AU-005). Chlorophyll and phaeopigments were summed and expressed as 24 gut pigments (ng pigment ind.⁻¹) (cf. Mackas and Bohrer, 1976).

The gut pigment of *C. glacialis* C5 was higher at nighttime than daytime. Assuming their grazing at night, the grazing rate ($GR_{ind.}$, mg pigment ind.⁻¹ day⁻¹) of *C. glacialis* C5 was calculated from following equation:

$$28 \qquad GR_{ind.} = GP \times k \times T / 10^6, \tag{1}$$

where *GP* is the individual gut pigment at night (ng pigment ind.⁻¹), k is the gut evacuation rate (0.017 min⁻¹, Tande and Båmstedt, 1985), and *T* is the length of the nighttime (mean 13

- 1 hours = 780 min during study period). The grazing impact of *C. glacialis* C5 on micro-size (>
- 2 10 μ m) chl. *a* (*GI*, % on chl. *a* standing stock day⁻¹) was calculated by the following equation:
- 3 $GI = GR_{ind.} \times N / Int. chl. a \times 100,$

4 where *N* is the abundance of *C. glacialis* C5 (ind. m^{-2}), and *Int. chl. a* is standing stock of 5 large-sized (> 10 µm) chl. *a* (mg m⁻²).

6 **2.3 Zooplankton community**

7 In the land laboratory, identification and enumeration bywith taxa were performed on 8 zooplankton samples collected by NORPAC nets under a stereomicroscope. For dominant 9 taxa (calanoid copepods), identification was made towith species and copepodid stage levels. For species identification of calanoid copepods, we referred mainly to Brodsky (1967) and 10 Frost (1974) for Calanus spp., Miller (1988) for Neocalanus spp. and Frost (1989) for 11 12 Pseudocalanus spp. For Pseudocalanus spp., species identification was made only for late copepodid stage C5 females/males (C5F/M) and C6F/M, and their early copepodid stages 13 14 (C1-C4) were treated as *Pseudocalanus* spp.

For the evaluation of the DVM of large dominant copepods, we enumerated *C*. *glacialis* from PCP net samples. For *C. glacialis*, the lipid accumulation of C5 was scored into three categories: I (the oil droplet length (ODL) was 0–4% of the prosome length (PL)), II (ODL was 4–40% of PL) and III (ODL was > 40% of PL). Gonad maturation of *C. glacialis* C6F was also scored into three categories: I (immature), II (small oocytes in the ovary or oviduct) and III (large eggs or distended opaque filled-in oviduct). For this gonad maturation index, we cited those for *C. hyperboreus* (Hirche and Niehoff, 1996).

22

23

A species diversity index (\underline{H}) in each sample was calculated using the equation:

書式変更:フォント : 斜体

 $H' = -\sum n / Ni \times \ln n / Ni$

24 where *p* is the abundance (ind. m_{r}^{-2}) of *jth* species and *Ni* is the abundance (ind. m_{r}^{-2}) of total 25 calanoid copepods in the sample (Shannon and Weaver, 1949). Pielou evenness (*J*) was also 26 calculated using the equation:

- 27 <u>J'=H'/ln (s)</u>
- 28 where *s* is total number of observed species in the community (Pielou, 1966).

Λ	書式変更:	フォント : 斜体
(書式変更:	上付き
\square	書式変更:	フォント : 斜体
$\backslash \gamma$	書式変更:	フォント : 斜体
$\backslash \rangle$	書式変更:	上付き
Y	書式変更:	フォント : 斜体
\neg	書式変更:	フォント : 斜体
\bigvee	書式変更:	フォント : 斜体
<u> </u>	書式変更:	フォント : 斜体
Y	書式変更:	フォント : 斜体

(2)

1 From NORPAC net samples, the mean copepodid stage (*MCS*) of *C. glacialis* was 2 calculated from the following equation:

3
$$MCS = \frac{\sum_{i=1}^{6} i \times Ai}{\sum_{i=1}^{6} Ai},$$

4 where *i* is the number of the copepodid stage (1–6 indicates C1–C6), and Ai (ind. m⁻²) is the 5 abundance of the *i*th copepodid stage (cf. Marin, 1987).

During the study period, SWE was observed from 18 to 19 September (Nishino et al.,
<u>2015in press</u>). To evaluate the effect of SWE, the abundances of each zooplankton taxon and
species were compared between "before SWE (10–18 September)" and "after SWE (19–25
September)" using the *U*-test. This statistical analysis was performed with Stat View.

10 3 Results

11 3.1 Hydrography and chlorophyll a

During the sampling period, the temperature ranged from -1.5°C to 3.3°C, and thermocline 12 was observed at a depth of approximately 25 m (Fig. 2a). Cold water below 0°C continuously 13 14 occurred below thermocline, whereas the temperature above thermocline decreased from 3.3 15 to 1.5°C during the study period (Fig. 2a). Salinity ranged from 31.1 to 32.8, and halocline 16 was observed at approximately 25 m, which was parallel to thermocline (Fig. 2b). Salinity below halocline was constant ca. > 32, whereas salinity in the upper layer increased from 31.1 17 to 31.6 throughout the study period. Chl. a ranged from 0.08 to 3.25 mg m⁻³ and increased 18 19 after the SWE (Fig. 2c). The T-S diagram showed that the hydrographic conditions in the 20 upper layer changed temporally; thus, temperature decreased while salinity increased during 21 the study period (Fig. 2d).

22 3.2 Zooplankton community

Zooplankton abundance ranged from 23610=<u>to</u>56809 ind. m⁻², and calanoid copepods and
 barnacle larvae composed 57% and 31% of the community, respectively (Table 1). For
 calanoid copepods, 15 species belonging to nine genera were identified. Within them,
 Pseudocalanus spp. and *C. glacialis* dominated and composed 65% and 30% of the total
 copepods, respectively. Shannon species diversity and Pielou evenness for the copepods
 commuity were 1.87±0.12 and 0.81±0.06, respectively. According to a comparison of time

書式変更:フォント : 斜体

(3)

periods before and after the SWE, the total zooplankton abundance, the Shannon species
 diversity and the Pielou evenness did not change, whereas one calanoid copepod *Centropages abdominalis* and cyclopoid copepods increased after the SWE (Table 1). On the other hand,
 one calanoid copepod *Metridia pacifica*, appendicularians, barnacle larvae, euphausiids and
 gymnosomes (*Clione limacina*) decreased after the SWE.

6 3.3 Calanus glacialis

7 Throughout the study period, the standing stock of *C. glacialis* ranged from 1990 to 14554 8 ind. m⁻², and no significant changes were detected with the SWE (Fig. 3a, Table 1). For the 9 population structure, all copepodid stages (C1 to C6F/M) occurred, and C5 was the most 10 dominant stage (36%) of the population (Fig. 3). The *MCS* did not vary with SWE (Fig. 3a). 11 Throughout the study period, the lipid accumulation of C5 was high (Fig. 3b), and the gonad 12 maturation of C6F was dominated by immature specimens (Fig. 3c). These parameters 13 exhibited no significant change with the SWE (*U*-test, p = 0.285-0.303).

On both 14 and 22 September, the C1–C4 and C5 populations of *C. glacialis* occurred mainly at lower layers (20–49 m) during the daytime, and they migrated to upper layers (0–20
m) at night (Fig. <u>43</u>). It should be noted that nearly half of C5 remained in the lower layer both day and night. Lipid accumulation was higher for the C5 population residing in the lower layer. C6F occurred at a lower layer throughout the day.

19 Throughout the study period, the standing stock of *C. glacialis* ranged from 1990 to 20 14554 ind. m⁻², and no significant changes were detected with the SWE (Fig. 4a, Table 1). 21 For the population structure, all copepodid stages (C1 to C6F/M) occurred, and C5 was the 22 most dominant stage (36%) of their population (Fig. 4). The *MCS* did not vary with SWE 23 (Fig. 4a). Throughout the study period, the lipid accumulation of C5 was high (Fig. 4b), and 24 the gonad maturation of C6F was dominated by immature specimens (Fig. 4c). These 25 parameters exhibited no significant change with the SWE (*U* test, p = 0.285 = 0.303).

The gut pigment of *C. glacialis* C5 ranged from 0.6 to 12.3 ng pigment ind.⁻¹ and showed a significant increase after the SWE (mean values: 2.6 vs 4.5 ng pigment ind.⁻¹, *U*test, p < 0.01) (Fig. 5a). In a comparison between day and night, most dates except 18 and 23 September showed high gut pigment levels at night by a factor of two to five times those of the daytime. In both day and night, the gut pigments were correlated with a standing stock of chl. a (p < 0.05) (Fig. 5b). 1 The population grazing rate of *C. glacialis* C5 ranged from 0.04–0.28 mg pigment m⁻² 2 day⁻¹, peaked on 20 September (Fig. 6a) and increased significantly after the SWE (0.11 vs 3 0.18 mg pigment m⁻² day⁻¹, *U*-test, p < 0.05). During the study period, chl. *a* peaked at 18 4 September, and micro-size (> 10 µm) dominated (54% of total chl. *a*), especially after the 5 SWE (66%) (Fig. 6b). The grazing impact of *C. glacialis* C5 on micro-size chl. *a* ranged 6 from 0.5–4.6% on chl. *a* standing stock day⁻¹ and was high before the SWE from 10 to 15 7 September (Fig. 6c).

8

9 4 Discussion

10 4.1 Zooplankton community

11 The zooplankton community in the Chukchi Sea is known to have greater large spatial and 12 temporal changes (Springer et al., 1989; Llinás et al., 2009; Matsuno et al., 2011). Total zooplankton abundance in this study was approximately half (mean: 34059 ind. m⁻²) the 13 14 abundance reported by Matsuno et al. (2012) on the Chukchi shelf (mean: 75683 ind. m⁻²), with low abundance of small copepods (Pseudocalanus spp. and cyclopoids) and the 15 remarkable absence of the Arctic copepod Metridia longa. According to Matsuno et al. 16 (2012), the calanoid copepod community in the western Arctic Ocean during the autumn 17 18 could be divided into three groups (Shelf, Slope and Basin) that vary with depth. According 19 to their criteria, a shallower bottom depth and higher temperature and salinity indicate that the zooplankton community in this study corresponded to the Shelf group defined by Matsuno et 20 al. (2012). As characteristics of the Shelf group, high zooplankton abundance (mean: 75683 21 22 ind. m⁻²) and dominance of neritic (Pseudocalanus spp. and Centropages abdominalis) and 23 Pacific copepods (Eucalanus bungii, Neocalanus cristatus and Metridia pacifica) were 24 reported (Matsuno et al., 2012). 25 Comparing the above characteristics by Matsuno et al. (2012), the zooplankton abundance of this study was nearly half (mean: 34059 ind. m^{-2}), there was a low abundance 26 27 of small copepod Pseudocalanus spp. and cyclopoid copepods, and no occurrence of Arctic copepod *Metridia longa* was remarkable (Table 1). For the hydrography of this station, 28 29 Nishino et al. (2015in press) noted that the upper warm and less saline water was the Pacific

30 Summer water, which was transported to the Arctic Ocean during summer, and the cold saline

31 water of the lower layer was the remnant brine of the Pacific Winter water, which was

書式変更:上付き

transported to the Arctic Ocean during winter. Geographically, the present station was located 1 2 at a main stream of the Pacific water (Weingartner et al., 2005). The high abundance of the 3 Pacific copepod M. pacifica and Neocalanus spp. and the lack of the occurrence of the Arctic 4 *M. longa* in this study was thought to be a reflection of the water mass covering the station. 5 Because of these reasons (less small copepods and high abundance of Pacific copepods), the 6 Shannon species diversity and Pielou evenness in this study (1.87 and 0.81) are higher than 7 the reported values for the whole Chukchi Sea (1.79 and 0.62 calculated from Matsuno et al., 8 2012).

9

Seasonal characteristics during summer included the addition of holoplankton and the 10 11 dominance of the meroplankton (barnacle and bivalve larvae), which composed 39% of zooplankton abundance (Hopcroft et al., 2010). Dominance of the barnacle larvae also 12 13 occurred in this study (Table 1). Benthic barnacle adults released their larvae when they meet 14 phytoplankton blooms (Crisp, 1962; Clare and Walker, 1986), and their larvae spendt two to 15 three weeks atin the water columns and then settled (Herz, 1933). The abundance of barnacle larvae in this study (mean 10430 ind. m⁻²) was 13-55% lower than that in summer 16 (19114-79899 ind. m⁻², Matsuno et al., 2011). It also should be noted that the abundance of 17 18 barnacle larvae decreased significantly during the study period (Table 1). These facts suggest 19 that most of the barnacle larvae may have ended the planktonic phase and settled to the sea 20 bottom during the study period (autumn).

21 Concerning the effect of the SWE, several limited a few taxa and species showed 22 significant changes in abundance (Table 1). Within them, the decrease of barnacle larvae may have been caused by their seasonal settlement and was not related to the SWE. For dominant 23 24 species, cyclopoid copepods were increased after the SWE (Table 1). The generation length 25 of cyclopoid copepods was reported to be two to three months in the Arctic Ocean (Dvoretsky 26 and Dvoretsky, 2009). In ambient temperature $(-1.5 \text{ to } 3.3^{\circ}\text{C})$, the egg hatching of this taxon 27 is estimated to be 11-41 days (Nielsen et al., 2002). These facts suggest that the increase of 28 cyclopoid copepods would not be caused by their reproduction within the study period (16 29 days). As an alternative cause, the horizontal advection of the water mass during the study 30 period, which was reported by Nishino et al. (2015in press), should be considered. These 31 results suggest that the effect of SWE on zooplankton abundance was relatively small because 32 of the longer generation length of mesozooplankton in this region.

1 4.2 Population structure of *C. glacialis*

2 Calanus glacialis is a key species that dominates in the zooplankton biomass and commonly 3 occurs in the continental shelf throughout the Arctic Ocean (Conover and Huntley, 1991; Lane et al., 2008). The generation length of this species is known to vary from one to three years 4 5 (mostly two years) depending on annual primary productivity in each region (Falk Petersen et 6 al., 2009). For the life cycle of C. glacialis, they grow to C3-C4 at the epipelagic layer 7 during the first summer, then descend to the deep layer and enter diapause, developing to C5 8 at the epipelagic layer in the second summer, then down to the deep layer and molting to the 9 adult stage (C6) using stored lipid as energy. They reproduce at the epipelagic layer, using the 10 grazed phytoplankton bloom as energy (Longhurst et al., 1984; Ashjian et al., 2003; Falk 11 Petersen et al., 2009).

12 -Concerning the population structure, Ashjian et al. (2003) reported that C. glacialis 13 around the Northwind Abyssal Plain was dominated by C5 and C6F in September. In the 14 present study, the population structure of C. glacialis was dominated by C5 (Fig. 34a), and 15 their MCS (mean \pm sd: 3.77 \pm 0.20) was similar to the reported value for autumn in this region 16 (3.58, Matsuno et al., 2012). Most of the C6F had immature gonads and no ovigerous C6Fs 17 were observed (Fig. 34c). These results corresponded with the year-round observation around 18 the Northwind Abyssal Plain (Ashjian et al., 2003). Calanus glacialis C6F is known to occur 19 at the epipelagic layer in April, just before sea-ice melting (Kosobokova, 1999), and to 20 reproduce with grazing ice algae and ice-edge bloom (Campbell et al., 2009). Thus, because 21 this study period (September) greatly varied with regards to their reproduction period (April), 22 most C6Fs were considered to have immature gonads when residing in the lower layer 23 (diapause).

24 Nocturnal ascent DVM, which is related to nighttime grazing on phytoplankton, was 25 reported for C. glacialis in the Arctic Ocean during spring and autumn (Runge and Ingram, 26 1988; Conover and Huntley, 1991). In this study, DVM was observed for C3, C4 and C5 (Fig. 43). At high-latitude seas, the magnitude of Calanus spp. DVM is known to vary with the 27 28 season and copepodid stage, and their DVM intensity is greater during spring and autumn 29 when the diel changes in light penetration are large (Falkenhaug et al., 1997). No DVM of 30 Calanus spp. was reported for lipid accumulated C5 (Falk-Petersen et al., 2008). In the 31 present study, nearly half of the C5 population, which was characterized as having a large 32 lipid accumulation, remained in the lower layer throughout the day (Fig. 43). The deep C5 33 population may have already completed lipid accumulation and ceased DVM in the study

1 period (September), while the remaining C5 population with active DVM may have grazed on

2 phytoplankton in the upper layer during the night and stored lipid in preparation for diapause.

3 These results suggest that the C. glacialis population in this study was at the seasonal phase

4 just before entering diapause, and this interpretation corresponded well with their life cycle in

5 this region (Ashjian et al., 2003).

6 4.3 Grazing of C. glacialis

7 The gut pigment of C. glacialis in the Arctic Ocean is known to exhibit diel changes and 8 higher levels at night than in the daytime (Conover and Huntley, 1991). Higher gut pigment 9 levels at night were also observed in this study (U-test, p < 0.001) (Fig. 5a). Their gut pigments were correlated with the standing stock of chl. a (Fig. 5b) and increased during the 10 high chl. a period after the SWE (Fig. 5a). These facts suggest that C. glacialis feeding 11 12 responded to the small phytoplankton bloom, which was enhanced by the nutrient supply and vertical mixing caused by the SWE (Nishino et al., 2015in press; Yokoi et al., submittedin 13 preparation). 14

15 Concerning gut pigment measurement, underestimation by decomposition of phytoplankton pigment through gut passage has been reported (Conover et al., 1986; Head, 16 1992). This underestimation is reported to be ca. 0.1-10% of grazing (Conover et al., 1986) 17 and varies with light conditions, grazing behavior and phytoplankton species (Head, 1992). 18 19 To estimate the grazing impact, data on the gut evacuation rate (k, \min^{-1}) is needed 20 (Mauchline, 1998). The gut evacuation rate is known to have a positive correlation with 21 temperature (Dam and Petersen, 1988). From the equation between k and temperature (T: 22 °C): k = 0.00941 + 0.002575T (Mauchline, 1998) and ambient temperature in this study (T: -1.5 to 3.3°C), k would be estimated to be 0.0055–0.0179 min⁻¹. This range covers the value 23 applied in the present study (k = 0.017, Tande and Båmstedt, 1985). The value was also at the 24 25 range observed by our independent laboratory experiments in September 2010 (0.006-0.041, Matsuno et al., unpublished data). These facts suggest that the value applied in this study (k =26 27 0.017) was reasonable for C. glacialis in this region.

Assuming that half of the C5 population performed nocturnal ascent and grazing on phytoplankton at night, using C:Chl. *a* ratio (29.9, Sherr et al., 2003), the grazing impact (mg $C m^{-2} day^{-1}$) of *C. glacialis* C5 was calculated (Table 2). The grazing impact of this study was estimated to be 4.14 mg C m⁻² day⁻¹, which was lower than the reported value (Tande and Båmstedt, 1985). In the Arctic Ocean, the grazing impact of copepods is greatly affected **書式変更:**フォント: 12 pt

1 by the extent of sea-ice, environmental conditions and food concentration (Hirche et al., 1991;

Campbell et al., 2009). The low grazing impact of this study may be caused by the low
 feeding activity of C5 just before their diapause and may also be related to the low ambient

4 chl. a (mean: 14.3 mg m⁻²) during the study period.

5 We also estimated the food requirement of C. glacialis C5 to support their metabolism 6 under ambient temperature (Ikeda and Motoda, 1978; Ikeda et al., 2001). The potential 7 contribution proportion of phytoplankton food to food requirements was 12.6% for C. glacialis C5 (Table 2). This result indicates that C. glacialis C5 could not maintain its 8 9 population solely on phytoplankton food; other food sources are important. Regarding food 10 for C. glacialis, Campbell et al. (2009) reported that this species prefers microzooplankton 11 rather than phytoplankton in the Chukchi Sea, and Levinsen et al. (2000) noted that the micro-12 size (> 10 µm) ciliates and dinoflagellates are important food sources during post-bloom. For 13 the microplankton community during the study period, Yokoi et al. (submittedin preparation) 14 noted that not only diatoms (1.64-14.11 cells ml⁻¹) but also dinoflagellates (0.54-2.42 cells 15 ml^{-1}) and ciliates (0.14–2.76 cells ml^{-1}) were abundant. Thus, as food sources for C. glacialis, microzooplankton (dinoflagellates and ciliates) are considered to be important in the Chukchi 16 17 Sea during autumn.

18 5 Conclusions

19 Throughout this study, short-term changes in the mesozooplankton community and the grazing impact of <u>C. glacialis</u> dominant copepods were evaluated in the Chukchi Sea during 20 21 autumn. During the 16-day sampling/observation period, the zooplankton community showed 22 no clear changes related to the SWE and dominant copepods prepared for diapause (i.e., 23 stored lipids in the pre-adult stage or adults with immature gonads). On the other hand, the 24 feeding intensity of dominant copepods increased with the reflection of the temporal phytoplankton bloom, which was enhanced by the SWE (Nishino et al., 2015in press). Thus, 25 26 the temporal phytoplankton bloom caused by the atmospheric turbulence (SWE) during 27 autumn may have had a positive indirect effect on mesozooplankton (SWE \rightarrow nutrient supply 28 from the deep layer \rightarrow small phytoplankton bloom \rightarrow copepod feeding) within a short period. 29 These facts suggest that some zooplankton may obtain benefit from an extension of the 30 primary production season with more turbulence and later freeze-up of the Chukchi Sea.However, because of the relatively long generation length of copepods, a smaller effect 31 was detected for their abundance, population structure, lipid accumulation and gonad 32

書式変更:フォント : 斜体

maturation within the short term period. 1

Author contributions 2

3 S.N., J.I. and T.K. designed and coordinated this research project. S.N. and J.I. were chief scientists during the MR13-06 cruise of R.V. Mirai. K.M. collected the zooplankton samples, 4 measured copepod gut pigments during the cruise and performed species identification and 5 enumeration of the zooplankton samples in the land laboratory. K.M. and A.Y. wrote the 6 7 manuscript with contributions from all co-authors.

8 Acknowledgements

9 We are grateful to the captain, officers and crew of the R.V. Mirai (JAMSTEC), operated by 10 GODI, for their help in the sample collection. This study was supported by the Green Network of Excellence Program's (GRENE Program) Arctic Climate Change Research 11 12 Project: 'Rapid Change of the Arctic Climate System and its Global Influences.' This study 13 was partially supported by a Grant-in-Aid for Scientific Research (A) (24248032) and a 14 Grant-in-Aid for Scientific Research on Innovative Areas (24110005) from the JSPS. 15

1 References

2	Ashjian, C. J., Campbell, R. G., Welch, H. E., Butler, M., and Keuren, D. V.: Annual cycle in
3	abundance, distribution, and size in relation to hydrography of important copepod
4	species in the western Arctic Ocean, Deep-Sea Res. I, 50, 1235–1261, 2003.
5	Ashjian, C. J., Braund, S. R., Campbell, R. G., George, J. C. C., Kruse, J., Maslowski, W.,
6	Moore, S. E., Nicolson, C. R., Okkonen, S. R., Sherr, B. F., Sherr, E. B., and Spitz, Y.
7	H.: Climate variability, oceanography, bowhead whale distribution, and Iñupiat
8	subsistence whaling near Barrow, Alaska, Arctic, 63, 179-194, 2010.
9	Brodsky, K. A.: Calanoida of the far-eastern seas and polar basin of the USSR. Israel Program
10	Scientific Translation, Jerusalem, 1967.
11	Cabe, G. J., Clark, M. P., and Serreze, M. C.: Trends in northern hemisphere surface cyclone
12	frequency and intensity, J. Climate, 14, 2763-2768, 2001.
13	Campbell, R. G., Sherr, E. B., Ashjian, C. J., Plourde, S., Sherr, B. F., Hill, V., and Stockwell,
14	D. A.: Mesozooplankton prey preference and grazing impact in the western Arctic
15	Ocean, Deep-Sea Res. II, 56, 1274–1289, 2009.
16	Clare, A. S. and Walker, G.: Further studies on the control of the hatching process in Balanus
17	balanoides (L), J. Exp. Mar. Biol. Ecol., 97, 295-304, 1986.
18	Comiso, J. C., Parkinson, C. L., Gersten, R., and Stock, L.: Accelerated decline in the Arctic
19	sea ice cover, Geophys. Res. Lett., 35, L01703. doi: 10.1029/2007GL031972, 2008.
20	Conover, R. J. and Huntley, M.: Copepods in ice-covered seas-distribution, adaptations to
21	seasonally limited food, metabolism, growth patterns and life cycle strategies in polar
22	seas, J. Mar. Syst., 2, 1–41, 1991.
23	Conover, R. J., Durvasula, R., Roy, S., and Wang, R.: Probable loss of chlorophyll-derived
24	pigments during passage through the gut of zooplankton, and some of the
25	consequences, Limnol. Oceanogr., 31, 878-887, 1986.
26	Crisp, D. J.: Release of larvae by barnacles in response to the available food supply, Anim.
27	Behav., 10, 382–383, 1962.
28	Dam, H. G. and Peterson, W. T.: The effect of temperature on the gut clearance rate constant
29	of planktonic copepods, J. Exp. Mar. Biol. Ecol., 123, 1-14, 1988.

2	Cyclopoida) in Kola Bay (Barents Sea), Mar. Biol., 156, 1433-1446, 2009.	
3	Falkenhaug, T., Tande, K. S., and Semenova, T.: Diel, seasonal and ontogenetic variations in	
4 5	the vertical distributions of four marine copepods, Mar. Ecol. Prog. Ser., 149, 105-119, 1997.	
6	Falk-Petersen, S., Leu, E., Berge, J., Kwasniewski, S., Nygård, H., Røstad, A., Keskinen, E.,	
7	Thormar, J., von Quillfeldt, C., Wold, A., and Gulliksen, B.: Vertical migration in	
8	high Arctic waters during autumn 2004, Deep-Sea Res. II, 55, 2275–2284, 2008.	
9 10	Falk-Petersen, S., Mayzaud, P., Kattner, G., and Sargent, J. R.: Lipids and life strategy of Arctic <i>Calanus</i> , Mar. Biol. Res., 5, 18–39, 2009.	
11	Frost, B. W.: Calanus marshallae, a new species of calanoid copepod closely allied to the	
12	sibling species C. finmarchicus and C. glacialis, Mar. Biol., 26, 77–99, 1974.	
13 14	Frost, B. W.: A taxonomy of the marine calanoid copepod genus <i>Pseudocalanus</i> , Can. J. Zool., 67, 525–551, 1989.	
15	Gnaiger, E.: Calculation of energetic and biochemical equivalents of respiratory oxygen	
16 17	consumption, in: Polarographic oxygen sensors, edited by: Gnaiger, E., and Horstner, H., Springer-Verlag, Berlin, 337–345, 1983.	
18 19	Grebmeier, J. M.: Shifting patterns of life in the pacific Arctic and sub-Arctic Seas, Annu. Rev. Mar. Sci., 4, 63–78, 2012.	
20 21	Head, E. J. H.: Gut pigment accumulation and destruction by arctic copepods <i>in vitro</i> and <i>in situ</i> , Mar. Biol., 112, 583–592, 1992.	
22 23	Herz, L. E.: The morphology of the later stages of <i>Balanus crenatus</i> Bruguiere, Biol. Bull., 64, 432–442, 1933.	
24 25	Hirche, HJ. and Niehoff, B.: Reproduction of the Arctic copepods <i>Calanus hyperboreus</i> in the Greenland sea-field and laboratory observations, Polar Biol., 16, 209–219, 1996.	
26 27 28	Hirche, HJ., Baumann, M. E. M., Kattner, G., and Gradinger, R.: Plankton distribution and the impact of copepod grazing on primary production in Fram Strait, Greenland Sea, J. Mar. Syst., 2, 477–494, 1991.	

Dvoretsky, V. G. and Dvoretsky, A. G.: Life cycle of Oithona similis (Copepoda:

2	the Chukchi Sea during summer 2004, Deep-Sea Res. II, 57, 27–39, 2010.
3 4	Ikeda, T. and Motoda, S.: Estimated zooplankton production and their ammonia excretion in the Kuroshio and adjacent seas, Fish. Bull., 76, 357–367, 1978.
5 6 7	Ikeda, T., Kanno, Y., Ozaki, K., and Shinada, A.: Metabolic rates of epipelagic marine copepods as a function of body mass and temperature, Mar. Biol., 139, 587–596, 2001.
8 9	Kawamura, A.: Fast sinking mouth ring for closing Norpac net, Bull. Japan. Soc. Sci. Fish., 55, 1121, 1989.
10 11	Kosobokova, K. N.: The reproduction cycle and life history of the Arctic copepod <i>Calanus</i> glacialis in the White Sea, Polar Biol., 22, 254–263, 1999.
12 13 14	Lane, P. V. Z., Llinás, L., Smith, S. L., and Pilz, D.: Zooplankton distribution in the western Arctic during summer 2002: Hydrographic habitats and implications for food chain dynamics, J. Mar. Syst., 70, 97–133, 2008.
15 16 17	Levinsen, H., Turner, J. T., Nielsen, T. G., and Hansen, B. W.: On the trophic coupling between protists and copepods in arctic marine ecosystem, Mar. Ecol. Prog. Ser., 204, 65–77, 2000.
18 19	Llinás, L., Pickart, R. S., Mathis, J. T., and Smith, S. L.: Zooplankton inside an Arctic Ocean cold-core eddy: probable origin and fate, Deep-Sea Res. II, 56, 1290–1304, 2009.
20 21	Longhurst, A., Sameoto, D., and Herman, A.: Vertical distribution of Arctic zooplankton in summer: eastern Canadian archipelago, J. Plankton Res., 6, 137–168, 1984.
22 23 24	Lowry, L. F., Sheffield, G., and George, C.: Bowhead whale feeding in the Alaskan Beaufort Sea, based on stomach contents analyses. J. Cetacean Res. Manage., 6, 215–223, 2004.
25 26	Mackas, D. and Bohrer, R.: Fluorescence analysis of zooplankton gut contents and an investigation of diel feeding patterns. J. Exp. Mar. Biol. Ecol., 25, 77–85, 1976.
27 28	Marin, V.: The oceanographic structure of eastern Scotia Sea-IV. Distribution of copepod species in relation to hydrography in 1981, Deep-Sea Res. A, 34, 105–121, 1987.

Hopcroft, R. T., Kosobokova, K. N., and Pinchuk, A. I.: Zooplankton community patterns in
 the Chukchi Sea during summer 2004, Deep-Sea Res. II, 57, 27–39, 2010.

- Markus, T., Stroeve, J. C., and Miller, J.: Recent changes in Arctic sea ice melt onset,
 freezeup, and melt season length. J. Geophys. Res., 114, C12024. doi:
 10.1029/2009JC005436, 2009.
- Matsuno, K., Yamaguchi, A., Hirawake, T., and Imai, I.: Year-toyear changes of the
 mesozooplankton community in the Chukchi Sea during summers of 1991, 1992 and
 2007, 2008. Polar Biol., 34, 1349–1360, 2011.
- Matsuno, K., Yamaguchi, A., Shimada, K., and Imai, I.: Horizontal distribution of calanoid
 copepods in the western Arctic Ocean during the summer of 2008. Polar Sci., 6,
 105–119, 2012.
- 10 Mauchline, J.: The biology of calanoid copepods, Adv. Mar. Biol., 33, 1–660, 1998.
- Miller, C. B.: *Neocalanus flemingeri*, a new species of calanidae (copepoda: calanoida) from
 the subarctic Pacific Ocean, with a comparative redescription of *Neocalanus plumchrus* (Marukawa) 1921, Prog. Oceanog., 20, 223–273, 1988.
- 14 Motoda, S.: North Pacific standard plankton net, Inf. Bull. Planktol. Jpn., 4, 13–15, 1957.
- Nielsen, T. G., Møller, E. F., Satapoomin, S., Ringuette, M., and Hopcroft, R. R.: Egg
 hatching rate of the cyclopoid copepod *Oithona similis* in arctic and temperate waters,
 Mar. Ecol. Prog. Ser., 236, 301–306, 2002.
- Nishino, S., Kawaguchi, Y., Inoue, J., Hirawake, T., Fujiwara, A., Futsuki, R., Onodera, J.,
 and Aoyama, M.: Nutrient supply and biological response to wind-induced mixing,
 inertial motion, <u>internal waves</u>, and currents in the northern Chukchi Sea, J. Geophys.
 Res. C, 120, 1975–1992, 2015. in press.
- Pielou, E. C.: The measurement of diversity in different types of biological collections. J.
 Theor. Biol., 13, 131–144, 1966.
- Runge, J. A. and Ingram, R. G.: Underice grazing by planktonic, calanoid copepods in
 relation to a bloom of ice microalgae in southeastern Hudson Bay, Limnol.
 Oceanogr., 33, 280–286, 1988.
- Sepp, M. and Jaagus, J. Changes in the activity and tracks of Arctic cyclones, Climate Change,
 105, 577–595, 2011.
- Serreze, M. C., Walsh, J. E., Chapin III, F. S., Osterkamp, T., Dyurgerov, M., Romanovsky,
 V., Oechel, W. C., Morison, J., Zhang, T., and Barry, R. G. Observational evidence

1	of recent change in the northern high-latitude environment, Climate Change, 46,
2	159–207, 2000.
3	Shannon, C. E., and Weaver, W.: The Mathematical Theory of Communication. The
4	University of Illinois Press, Urbana, 1949.
5	Sherr, E. B., Sherr, B. F., Wheeler, P. A., and Thompson, K.: Temporal and spatial variation
6	in stocks of autotrophic and heterotrophic microbes in the upper water column of the
7	central Arctic Ocean, Deep-Sea Res. I, 50, 557-571, 2003.
8	Springer, A. M., McRoy, C. P., and Turco, K. R.: The paradox of pelagic food web in the
9	northern Bering Sea-II. Zooplankton communities. Cont. Shelf Res., 9, 359-386,
10	1989.
11	Stroeve, J., Holland, M. M., Meier, W., Scambos, T., and Serreze, M.: Arctic sea ice decline:
12	Faster than forecast. Geophys. Res. Lett., 34, L09501. doi: 10.1029/2007GL029703,
13	2007.
14	Tande, K. S. and Bámstedt, U.: Grazing rates of the copepods Calanus glacialis and C.
15	finmarchicus in arctic waters of the Barents Sea, Mar. Biol., 87, 251-258, 1985.
16	Weingartner, T., Aagaard, K., Woodgate, R., Danielson, S., Sasaki, Y., and Cavalieri, D.:
17	Circulation on the north central Chukchi Sea shelf, Deep-Sea Res. II, 52, 3150-3174,
18	2005.
19	Yokoi, N., Matsuno, K., Ichinomiya, M., Yamaguchi, A., Nishino, S., Onodera, J., Inoue, J.,
20	and Kikuchi, T.: Short-term changes of microplankton community in the Chukchi
21	Sea during autumn: Consequence of strong wind event, Biogeosciences, submitted.in
22	preparation.
23	Zhang, X., Walsh, J. E., Zhang, J., Bhatt, U. S., and Ikeda, M.: Climatology and interannual
24	variability of Arctic cyclone activity: 1948-2002, J. Climate, 17, 2300-2317, 2004.

1 Figure captions

2	Figure 1. Location of sampling station in the Chukchi Sea from 10 to 25 September 2013.					
3	Depth contours (50, 100, 1000, 2000 and 3000 m) are superimposed. Arrows					
4	indicate major current flows in this region (cf. Grebmeier, 2012).					
5	Figure 2. Temporal changes in temperature (a), salinity (b), chlorophyll a (c) and T-S					
6	diagram (d) at a fixed-station in the Chukchi Sea from 10 to 25 September 2013.					
7	Figure 3. Calanus glacialis: temporal changes in the population structure (a), lipid					
8	accumulation of C5 (b) and gonad maturation of C6F (c) at a fixed station in the					
9	Chukchi Sea from 10 to 25 September 2013. Diel vertical distribution of Calanus					
10	glacialis at a fixed station in the Chukchi Sea, 14 and 22 September 2013. For C5,					
11	the mean lipid accumulation score (1-3) is also shown. Note that no C6Ms were					
12	collected.					
13	Figure 4. Diel vertical distribution of Calanus glacialis at a fixed-station in the Chukchi Sea,					
14	14 and 22 September 2013. For C5, the mean lipid accumulation score (1-3) is also					
15	shown. Note that no C6Ms were collected. Calanus glacialis: temporal changes in					
16	the population structure (a), lipid accumulation of C5 (b) and gonad maturation of					
17	C6F (c) at a fixed station in the Chukchi Sea from 10 to 25 September 2013.					
18	Figure 5. Calanus glacialis: temporal and diel changes in the gut pigment of C5 (a) at a fixed					
19	station in the Chukchi Sea from 10 to 25 September 2013. The relationship between					
20	the gut pigment of C5 and standing stock chlorophyll a (b). Dotted and dashed lines					
21	indicate regressions for day and night, respectively. For (b), "whole" indicates					
22	including all data, both day and night. **: $p < 0.01$, *: $p < 0.05$.					
23	Figure 6. Temporal changes in the grazing rate of Calanus glacialis C5 (a), integrated size-					
24	fractionated chlorophyll a (b) and the grazing impact of C. glacialis C5 on micro-					
25	phytoplankton (> 10 μ m) (c) at a fixed station in the Chukchi Sea from 10 to 25					
26	September 2013.					
07						

1	Table 1. List of mesozooplankton taxa and calanoid copepod species and their mean
2	abundances (ind. m^{-2}) at a fixed station in the Chukchi Sea from 10 to 18 September,
3	before the strong wind event (SWE), and 19-25 September (after the SWE), 2013.
4	For calanoid copepods, Shannon species diversity and Pielou evenness were
5	calculated. Differences between the two periods (before vs. after the SWE) were
6	tested with the U-test. *: $p < 0.05$, **: $p < 0.01$, ***: $p < 0.0001$, NS: not significant.

Species <u>/taxa</u> <u>Community parameters</u>	Before SWE (10–18 Sep.)	After SWE (19–25 Sep.)	U-test
Calanoid copepods			
Acartia longiremis	604	542	NS
Calanus glacialis	6714	5658	NS
Calanus hyperboreus	0	5	NS
Centropages abdominalis	9	29	*
Eucalanus bungii	6	6	NS
Eurytemora herdmani	0	2	NS
Metridia pacifica	251	154	*
Microcalanus pygmaeus	6	3	NS
Neocalanus cristatus	6	5	NS
Neocalanus flemingeri	46	65	NS
Neocalanus plumchrus	12	15	NS
Pseudocalanus acuspes	3393	3254	NS
Pseudocalanus mimus	1194	1296	NS
Pseudocalanus minutus	2178	2387	NS
Pseudocalanus newmani	2805	2774	NS
Pseudocalanus spp. (C1–C4)	2758	2980	NS
Cyclopoid copepods	511	1153	**
Poecilostomatoid copepods	0	3	NS
Amphipoda	9	5	NS
Appendicularia	707	442	*

Barnacle larva	12118	8945	***
Chaetognatha	1281	1039	NS
Echinodermata larva	31	61	NS
Eubrachyura zoea	41	26	NS
Euphausiacea	18	3	*
Gymnosomata	172	84	**
Hydrozoa	209	205	NS
Isopoda	3	3	NS
Polychaeta	1124	1005	NS
Thecosomata	16	8	NS
Total zooplankton	36223	32154	NS
Shannon species diversity	<u>1.85</u>	<u>1.90</u>	<u>NS</u>
Pielou evenness	<u>0.80</u>	0.82	<u>NS</u>

表の書式変更

1	Table 2. Calanus glacialis: Comparison of food requirements (ingestion) and grazing rate on				
2	phytoplankton and the proportion of phytoplankton food in the Chukchi Sea from 10				
3	to 25 September 2013. Food requirements were calculated from the metabolism,				
4	which was estimated by the body mass, temperature (Ikeda et al., 2001), respiratory				
5	quotient (Gnaiger, 1983) and individual carbon budget (Ikeda and Motoda, 1978).				
6	For details on the values used in this calculation, see the footnotes.				

Day/night		Flux (mg C m^{-2} da	Proportion of phytoplankton	
(depth, temp.: <i>T</i> , period)	Metabolism	Food requirement (ingestion) (A)	Grazing rate on phytoplankton (B)	food (%: B/A)
Day-lower layer	5.64	14.11	_	_
(20–49 m, –0.74°C, 11 h)				
Night-upper layer	4.12	10.29	_	-
(0–20 m, 2.31°C, 13 h)				
Night-lower layer	3.34	8.34	-	_
(20–49 m, –0.74°C, 13 h)				
Daily-water column	13.10	32.74	4.14	12.6

7 Dry mass (*DM*: mg ind.⁻¹) of *C. glacialis* C5 was 0.654 (Matsuno et al., unpublished data).

8 Oxygen consumption (μ L O₂ ind.⁻¹ h⁻¹) = exp(-0.399+0.801×Ln(*DM*)+0.069×*T*)) (Ikeda et al.,

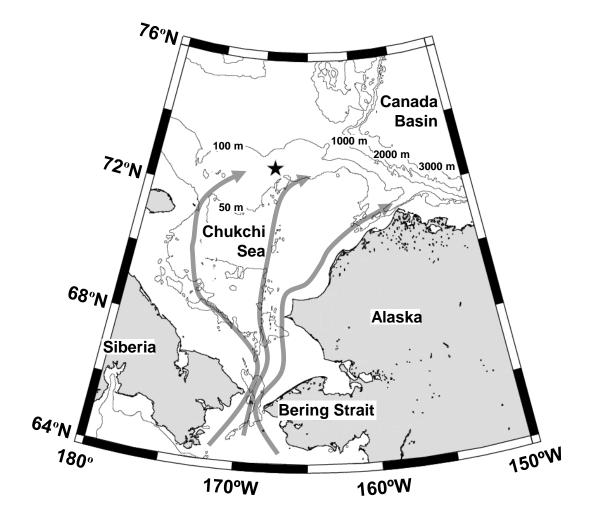
9 2001).

10 Respiratory quotient ([CO₂/O₂]) was assumed to be 0.97 (Gnaiger, 1983).

11 Metabolism was assumed to be 0.4 (40%) of ingestion (Ikeda and Motoda, 1978).

12 Mean abundance of *C. glacialis* C5 was 2176 ind. m^{-2} during this study period.

13 C:Chl ratio was 29.9 for September in this region (Sherr et al., 2003).



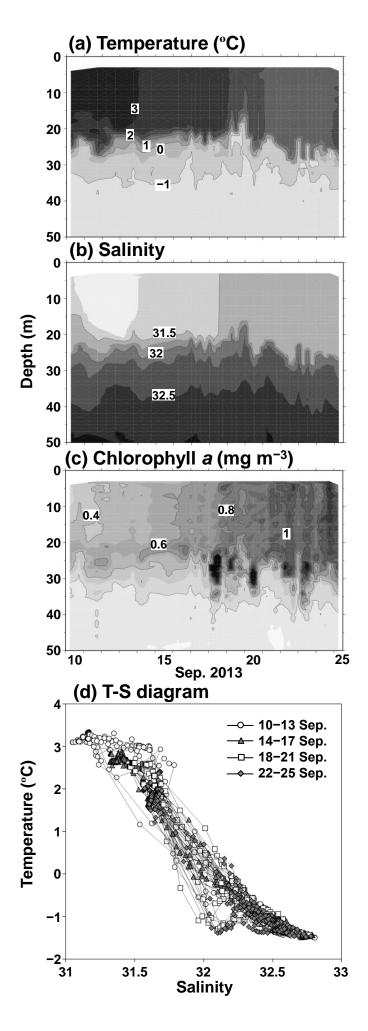
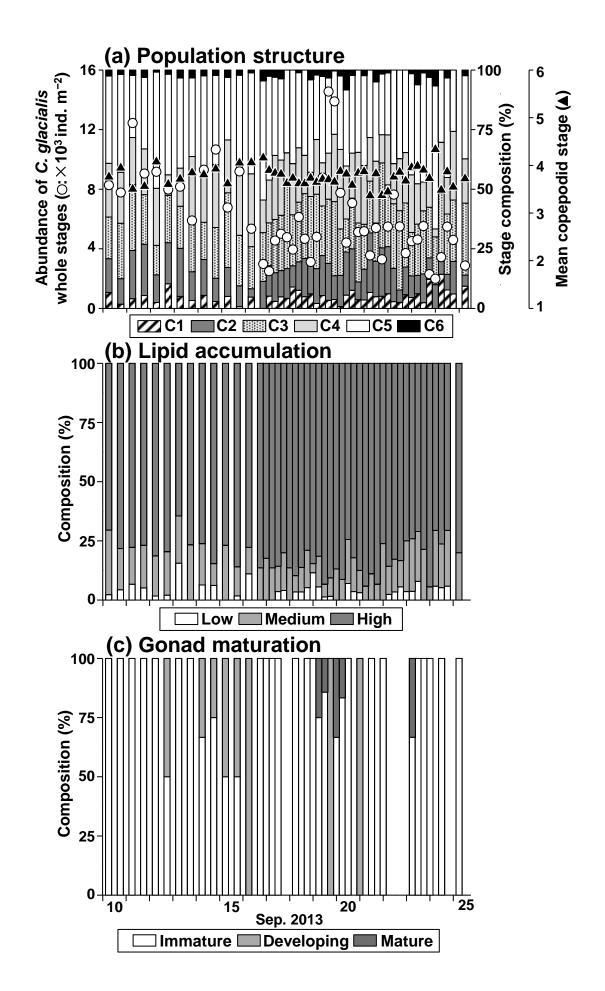
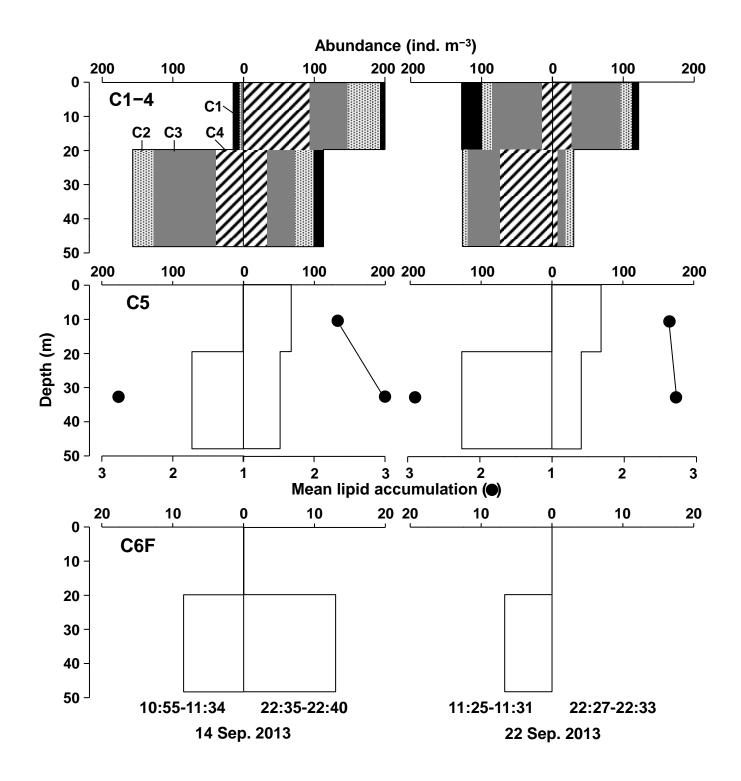
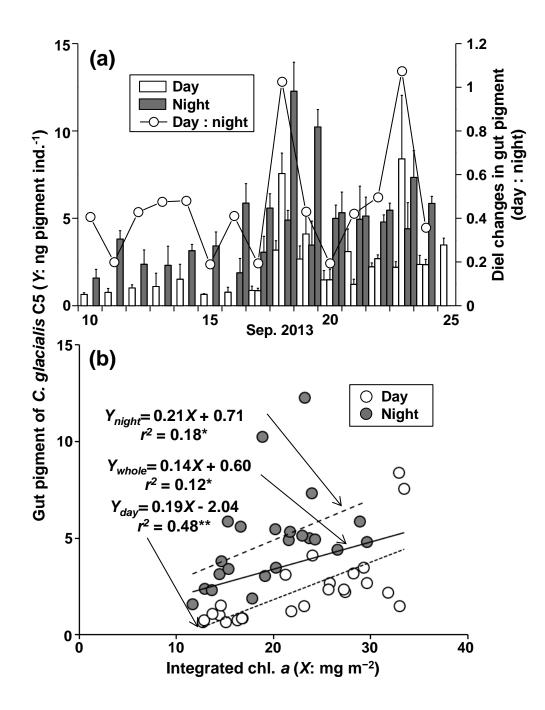


Fig. 2. (Matsuno et al.)







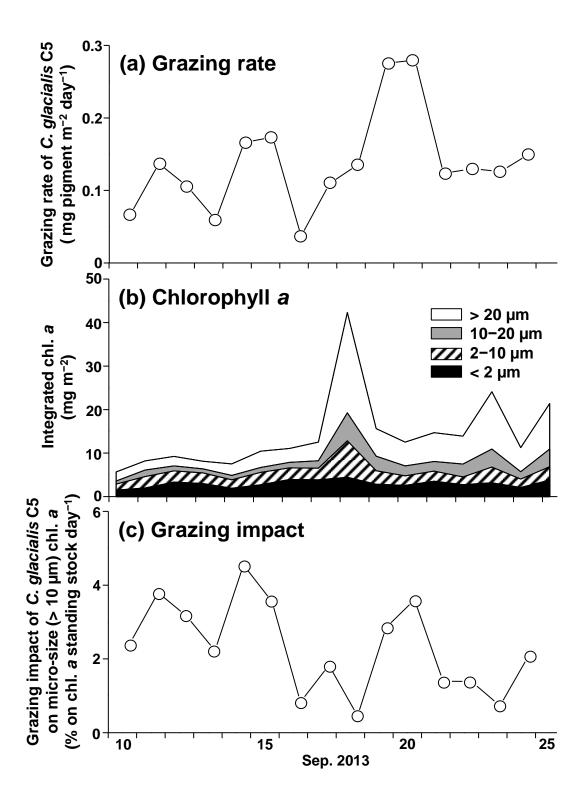


Fig. 6. (Matsuno et al.)