Author's response

We thank the two referees for their useful comments and address them below. The discussion has been modified according to their advices and corrected by an English native speaker. The full modified discussion is available on page 9. The manuscript with tracked changes is available at the end of the document.

In addition to the modifications proposed by the referees we have replaced the reference "Falkowski et al. 1997" by "Moore et al. 2013" as requested by the editor (*Moore, J. K., Geider, R. J., Guieu, C., Jaccard, S. L., Jickells, T. D., LaRoche, J., Lenton, T. M., Mahowald, N. M., Marañón, E., Marinov, I., Nakatsuka, T., Oschlies, A., Saito, M. A., Thingstad, T. F., Tsuda, A., Ulloa, O.: Processes and patterns of nutrient limitation, Nat. Geosci., 6, 701–710, 2013*).

Anonymous Referee #1:

Referee: If diazotrophs release almost nitrogen they fixed and primary production increased using the released nitrogen, f-ratio could decrease since regenerated production is likely to be enhanced. The decrease of f-ratio was observed in a time-series experiment in western Pacific warm pool when nitrogen fixation and primary production showed an increasing trend (Shiozaki et al., 2013, L&O). Please consider the difference between their results and your ones.

Response: We did not measure the f-ratio and thus we cannot compare directly our results with those presented in the Shiozaki et al. paper. However, an enhanced N remineralization under high PP and N₂ fixation rates may explain the higher efficiency of UCYN-C to promote C export. Therefore, we have added the reference at the end of section 4.2 to support the remineralization hypothesis: *"Additionally, when UCYN-C dominated, an enhanced N remineralization may have enabled more C to be fixed per unit of fixed N₂ leading to a higher <i>e-ratio. A proportionally higher N remineralization following high PP and N₂ fixation rates is supported by similar findings in the western North Pacific warm pool (Shiozaki et al., 2013)."*

Referee: Recent study showed that DDA contributed to export production efficiently (Karl et al., 2012, PNAS). Meanwhile, the present study indicated that the production driven by UCYN-C was more efficient to promote export production than by DDA (P4293, L4-5). What environment makes UCYN-C flourish? Is this result consistent with the result of Karl et al. (2012)?

Response: Here we show a tight coupling between N_2 fixation and N export when DDAs dominated the diazotrophic community. This suggests a direct sink of the DDAs which is in good agreement with Karl et al. (2012) findings. Alternatively, when UCYN-C dominated, the use of DON and an enhanced remineralization of N led to a higher export production efficiency. We cannot conclude that UCYN-C are intrinsically more efficient to promote

export production than DDAs. However, we propose that the flourishing of UCYN-C have triggered directly or indirectly the process that has enhanced the export. The ecology of UCYN-C is not well constrained and the factor controlling their development are poorly understood. This is discussed in a companion paper under submission in the VAHINE special issue (Turk et al, 2015).

Specific comment

Referee: P4275 L15 Correct to N2 fixation.

Response: It has been corrected in the new version.

Referee: P4290 L8 The unit of nitrate concentration is nM here. But that is μ M in P4285 L2,3. The authors should use unit consistently.

Response: All the units of concentration are now expressed in μ mol L⁻¹

Referee: P4290 L10 Jickells et al (2005) is not appropriate reference here.

Response: The reference has been deleted

Referee: P4292 L1-5 This discussion seems not to fit the context.

Response: The discussion has been modified as follows: "Assuming that DON and N_2 fixation are the only possible sources of N in the mesocosms, we calculated that a DON use of 0.9 μ mol L⁻¹ would have supported up to 78 % of the PON production during P2, and potentially fueled the PON export to the same extent. This is in agreement with Torres-Valdes et al. (2009) and Letscher et al. (2013) who showed that DON pool is a dynamic contributor of the N cycle able to supports up to 40 % of the vertical PON export in the oligotrophic gyres of the Pacific and Atlantic Oceans."

Referee: P4294 "the diazotrophs are known to over-fix C relative to N" Need citation.

Response: The reference Mulholland et al. (2007) has been added (Mulholland, M.: The Fate of nitrogen fixed by diazotrophs in the ocean, Biogeosciences, 4, 37–51,doi:10.5194/bg-4-37-2007, 2007)

Anonymous Referee #2:

Specific comments

Referee: P4278, L13: do the authors mean phosphorous instead of phosphate? DIP is a commonly used abbreviation for dissolved inorganic phosphorous. Phosphate is in fact a form of this DIP, so this abbreviation seems redundant. If authors mean DIP in its original

meaning please change phosphate by phosphorous. Otherwise, consider using PO4 instead. And apply these changes to the whole text.

Referee: P4278, L15: again, it is confusing that P is used for phosphate, as it is the name of the element. Please use PO4 or something similar instead.

Response: It is clear that phosphate (oxidation degree V) is by far the most important component of the "Phosphorous cycle" and that most of natural inorganic and organic compounds, in dissolved and particulate forms, contain phosphate group (H₃PO₄, ATP, ADP, Phospholipids, Glucose-6 phosphate...). It would be of great interest to speak about the Phosphate cycle and not the Phosphorous cycle. The Phosphorous in an element spontaneously oxidized which cannot exist in this elemental form in water. Nevertheless, we understand that it is confusing that P is used for phosphate, as it is the name of the element. We have replaced in the text DIP by phosphate* (PO₄³⁻) and define DOP by dissolved organic phosphorous.

*Whatever the method currently used, the basis for phosphate measurements in seawater is the Murphy and Riley method (A modified single solution for the determination of phosphate in natural waters, Analytica Chimica Acta, Volume 27, Pages 31-36, 1962).

Referee: P4280, L13: please do not use unexplained abbreviations in section headers, write the complete term and add the acronym in parentheses, especially when it is not the commonly used one as DIP for dissolved inorganic phosphate. This helps the reader to skip through the sections without coming back to look for the meaning of the acronym. Do the same in the next headers whenever necessary.

Response: We deleted all the abbreviations in the headers

Referee: P4280, L17: Please could the authors explain why adding an excess of KH2PO4 stops PO4 assimilation? Is it an effect of dilution of the tracer?

Response: The addition of KH_2PO_4 is performed in order to dilute the tracer. The explanation has been added in the text as follow: "Incubations were stopped by adding 50 μ L of KH_2PO_4 solution (10 mmol L⁻¹) in order to reduce to a minimum the ³³P assimilation by dilution effect [...]".

Referee: P4286, L9: please explain what the authors mean with "increased faster". It is slightly ambiguous, it suggests a sharp change that it not so obvious looking at the graph.

Response: A sharp change occurred in M3 but is not obvious in M1, M2. The sentence has been modified by "*PP continued to increase*" in order to include the PP evolution in all mesocoms.

Referee: P4286, L19: the reviewer wonders if the authors consider that this figure is essential? Does it add to the information in the text? As most of the information of this figure is given or could be given in the text, is it essential to keep this fig 2 and the next fig 7?

Referee: P4288, L25-23: this sentence is unnecessary, the information is in the figure caption and the e ratio was defined in the previous sentence, simply add (Fig. 7) to the previous sentence, please. And again, is it essential to use this graph, given that more of its information is in the text?

Response: We agree that figures 2 and 7 do not add supplementary information and have thus been deleted. In the text, we added the size of the population after the p value when the statistical difference between P1 and P2 is tested "(p<0.05, n=57)".

Referee: P4290, L8: units of nitrate in nM, however, measurements were made in μ M, as explained in the methods section (P4285) and in the results section (P4285). The authors should try to be consistent with units through the text.

Response: All the unit of concentrations are now expressed in μ mol L⁻¹.

Referee: P4290, L10: Jickells et al (2005) does not apply here. Pathways of atmospheric nitrogen deposition are not only limited to dust, like Fe, which is exactly the scope of this study.

Response: The reference has been deleted.

Referee: P4291, L21 to P4292, L10: the argumentation of this part is difficult to follow. It is very difficult to grasp the different arguments supporting that the DON pool produced or maintained in the previous part of the experiment (P1) is now supporting part of the production of PON in addition to N2 fixation, and thus supporting the export of organic matter. Relating the effect of a lateral transport of DON to a close system seems difficult to fit in order to explain this argument.

Response: The conditions for the DON mobilization and the potential DON users are discussed in section 4.3. From P4291, L21 to P4292, L10 we showed the evidences of the DON consumption. As it was difficult to follow, we modified the paragraph as follows:" *During P2, the increase in PON concentrations (Fig. 3B) suggests that part of the freshly produced biomass remained in the water column. The accumulation of PON probably favored remineralization processes, explaining the increase in NH*⁴⁺ concentrations. This may have enhanced the transfer of the recently fixed N₂ to the non diazotrophic plankton as demonstrated by Bonnet et al. (2015a) and explains the development of picocyanobacteria during P2 (Leblanc et al., 2015). Additionally, the total amount of N provided by N₂ fixation did not account for all the exported PON during P2 (Fig. 6), implying that an additional N source played a significant role in promoting the export. The only alternative N source is DON

which, indeed, exhibited a significant decrease in concentration of $0.9 \pm 0.7 \mu$ mol L-1 in the mesocosms during P2 (see section 4.3 for further discussion on DON consumption). Assuming that DON and N₂ fixation are the only possible sources of N in the mesocosms, we calculated that a DON use of 0.9μ mol L⁻¹ would have supported ~78 % of the PON production during P2, and potentially fueled the PON export to the same extent. This is in agreement with Torres-Valdes et al. (2009) and Letscher et al. (2013) who showed that DON pool is a dynamic contributor of the N cycle able to supports up to 40 % of the vertical PON export in the oligotrophic gyres of the Pacific and Atlantic Oceans. A quantification of the diazotrophs in the sediment traps, performed on day 19, shows that ~10 % of the UCYN-C biomass in the mesocosms was exported this day, explaining ~ 7 % of the PON export (Bonnet et al., 2015a). Thus, the recently fixed N₂ by UCYN-C can directly be exported but is probably more efficiently transferred to non diazotrophic plankton through remineralization processes.

Referee: P4292, L24-27: This sentence is incomplete or has some extra particles, probably it should say "at LEAST 20

Response: The sentence has been replaced by: "Bonnet et al. (2015a) demonstrated that ~ 10% of the recently fixed N_2 during P2 was transferred toward non diazotrophic plankton i.e. in picoplankton and bacteria".

Referee: P4294, L15: why do the authors relate a canonical Redfield ratio (6.6) with a cite of Fukuda et al (1998)?

Response: The Fukuda et al paper's show that the averaged C/N ratio was equal to the Redfield ratio in bacteria specifically. We modified the text as follow: "Based on BP data and assuming a bacterial growth efficiency between 10 and 30% (del Giorgio and Cole, 1998) and a C/N ratio of 6.6 in bacteria cells (Fukuda et al. 1998), we calculated that bacterial respiration would have led to a DON consumption of 0.2 to 0.7 μ mol L-1 during P2."

Technical corrections

Referee: P4275, L15: please correct "v fixation".

Response: It has been replaced by " N_2 fixation".

Referee: P4276, L12: the term "alighted" is referred to fire, may the authours mean "lit up"

instead?

Response: "Alighted" has been replaced by "photic"

Referee: P4277, L23: please change "If heterotrophic although..." by "Although heterotrophic bacteria...."

Response: The correction has been applied

Referee: P4278, L6: please add "export, AND (2) to trace...."

Response: The correction has been applied

Referee: P4280, L7: please add "organic" to "dissolved matter" to be consistent with other sections.

Response: The correction has been applied

Referee: P4281, L11-14: please re write the sentence "Incubation bottles...for 24h". It is a bit too long and wordy sentence, difficult to understand.

Response: The sentence has been replaced by: "Incubation bottles were then amended with 5 % (vol:vol) of ${}^{15}N_2$ enriched seawater and closed without headspace with silicone septum caps. The latter were incubated on an in situ mooring line close to the mesocosms at the appropriate sampling depths for 24 h."

Referee: P4281, L14: delete of in "After of incubation". It should say, "after incubation"

Response: The correction has been applied

Referee: P4283, L14: correct "...DON and DOP, respectively".

Response: The correction has been applied

Referee: P4283, L17: add "were found to be A negligible source OF particulate...."

Response: The correction has been applied

Referee: P4284, L10-13: please re write this sentence, it is slightly wordy and takes some readings to get through it.

Response: The sentence has been replaced by:" The values of fluxes and concentrations presented in the text are averaged over the three depths (no significant differences between depths have been evidenced, paired Friedman test, α =0.05)."

Referee: P4284, L18: correct, "according TO the propagation of errors".

Response: The correction has been applied

Referee: P4284, L24: please consider changing "similarly inside the mesocosms and in surrounding waters". The sentence looks incomplete, awkward, as it needs a verb.

Response: The sentence has been replaced by:"*Briefly, seawater temperature increased inside the mesocosms and in surrounding waters from 25.5 to 26.7*°C over the course of the experiment."

Referee: P4285, L8-10: another wordy sentence, too long, please try to re write.

Response: The sentence has been separated in two sentences: "The day after the fertilization, PO_4^{3-} concentrations reached 0.08 µmol L^{-1} in all mesocosms. Then, the concentrations decreased steadily, tending to initial concentrations (0.02-0.08 µmol L^{-1}) at the end of the experiment."

Referee: P4285, L15: try to keep consistency between figures and the mention in the text, change Fig. 1a. Use upper case for A in the text or change the A in the figure to a lower case. The same for the rest of figure mentions through the text.

Response: The upper case (Fig. 1A) is now used through the whole text

Referee: P4285, L23: add (PP) after primary production.

Response: The correction has been applied

Referee: P4285, L25: 7.3 nM d-1. This is an awkward way of expressing a rate, it might be confusing for many readers. It is commonly accepted to write all the units in rates, please change rates in the text to a format like nmol L-1 d-1 (as it was used in P4282,L5).

Response: The rates are now expressed in nmol $L^{-1} d^{-1}$.

Referee: P4288, L8: correct, "DON dynamicS".

Response: The correction has been applied

Referee: P4288, L10-12: wordy, please re write the sentence, "decreased of" in this context is confusing.

Response: The sentence has been separated in two sentences: "After these days, DOP concentrations significantly decreased (p< 0.05) and reached 0.09 ± 0.01 µmol L⁻¹ on average. DOP concentrations also decreased in surrounding waters from day 18 but to a lesser extent than in the mesocosms."

Referee: P4288, L17: delete at, "during P1 averaging at..."

Response: The correction has been applied

Referee: P4288, L18: comma before respectively. This happens through the text in several occasions, please, correct all of them.

Response: We added a comma before each "respectively".

Referee: P4288, L19: change to "M1 were higher than THOSE in..."

Response: The correction has been applied

Referee: P4288, L23: re write "much more stochastic", too informal and vague.

Response: The sentence has been replaced by:" The daily POP_{export} ranged from 1.0 to 18.4 nmol $L^{-1} d^{-1}$."

Referee: P4289, L1: "... integrated over P1 WERE..." (rates...were).

Response: The correction has been applied

Referee: P4289, L1: "... all the mesocosms, AND DID not significantly differ from"

Response: The correction has been applied

Referee: P4289, L3: "the resulting change ON the..."

Response: The correction has been applied

Referee: P4289, L7-8: "...but deviated negatively from 0... FROM day 19 TO day 23"

Response: The correction has been applied

Referee: P4289, L8: "Thus, even though..." Please re write this, too wordy.

Response: The sentence has been modified: "At the end of P2, the decrease of the TN_{calc} pool was 0.20 ± 0.04 µmol L⁻¹."

Referee: P4289, L9: delete of in "the TNcalc pool decreased of"

Response: The sentence has been modified has suggested in the previous comment.

Referee: P4291, L5: change "trough" by "through". This is repeated through the discussion, please check all the misspellings and correct them.

Response: This change has been applied through the whole text.

Referee: P4294, L19-22: please, re write this sentence, wordy, difficult to understand. Add a cite to the first argument (diazotrophs over-fix C).

Response: The sentence has been modified as follow: "Diazotrophs are known to over-fix C relative to N (Mulholland et al., 2007), which may explain why the POC:PON ratio was above the Redfield ratio during the experiment. The resulting N deficit for bacterial mineralization may have been found in the labile or semi labile DON pool. This hypothesis is supported by Van Wambeke et al. (2015) who showed that BP was limited by N availability."

Referee: Figures. Please do not use unexplained acronyms and abbreviations in figures captions, do write the whole term before adding the acronym, e.g. DDAs, PP... Figures should be self-contained, so anyone looking at them without reading the text can understand them. Consider joining both boxplots in one figure.

Response: We remove (or define) all the acronyms and abbreviations in the figure captions. The two box plots have been deleted.

The discussion and the conclusions have been modified has follows:

4. Discussion

4.1. Phosphate availability sustained N₂ fixation and in turn new primary production

 N_2 fixation is limited by micronutrients availability such as iron (Fe) (Dekaezemacker et al., 2013; Monteiro et al., 2011; Moore et al., 2009; Shiozaki et al., 2014), temperature (Breitbarth et al., 2007; Fu et al., 2014; Mulholland and Bernhardt, 2005), light (Garcia et al., 2013; Kranz et al., 2010) and ultimately by PO_4^{3-} availability (Dore et al., 2008; Karl et al., 1997; Moutin et al., 2008). During this experiment, PO_4^{3-} fertilization was carried out as P limitation would have prevented diazotrophs growth. This fertilization associated with the relatively high micronutrient concentrations (e.g. Fe) (Ambatsian et al., 1997), high seawater temperature (>25 °C) and the maintenance of the water mass in the photic layer, provided optimal conditions for diazotrophs growth. As expected, diverse diazotroph phylotypes developed extensively in the mesocosms with abundances of 1.10^5 to 5.10^5 nifH copies L⁻¹ (Turk et al., 2015). Furthermore, N₂ fixation rates in the mesocosms were high (18.5 ± 13.1 nmol N L⁻¹ d⁻¹) compared to those in surrounding waters (9.2 ± 4.8 nmol N L⁻¹ d⁻¹) (Fig. 1B) and are among the highest reported in the literature (Luo et al., 2012).

The contribution of N₂ fixation to PP (10.8 ± 5.0 %) in the mesocosms and in surrounding waters (5.7 ± 2.0 %) was in the upper range of previous studies in the Pacific Ocean (Raimbault and Garcia, 2008; Shiozaki et al., 2013) and the Mediterranean Sea (Bonnet et al., 2011; Ridame et al., 2013). Prior to PO_4^{3-} fertilization, NO_3^- concentrations were <0.04 µmol L⁻¹. As there was no external supply of NO_3^- , the potential consumption of the initial NO_3^- in the mesocosms represented <11.5 % of the integrated N₂ fixation rates over P1 and P2 (0.35 ± 0.08 µmol L⁻¹) (Fig. 6A). Thus, N₂ fixation supplied nearly all the new production during the experiment. These results indicate that in a N depleted system, diazotrophs can provide enough new N to sustain high PP (exceeding 2 µmol C L⁻¹ d⁻¹) and biomass (up to 1.42 µg L⁻¹ of Chl a), as long as PO_4^{3-} does not limit N₂ fixation.

4.2. The relative efficiency of different diazotrophs to export particulate matter

Only few studies have focused on the direct coupling between N_2 fixation and particulate export (Dore et al., 2008; Karl et al., 2012; White et al., 2012). To our knowledge, the only study comparing the export efficiency of different diazotrophs reports that DDAs blooms could contribute up to 44 % of the direct export in the North East Pacific, while UCYN (Group A) and Trichodesmium spp. could account for only 0 to 10 % of the export (White et al., 2012). The scarcity of data is due to methodological issues associated with the use of sediment traps in the open ocean due to (1) the patchy distribution of N_2 fixers that are not necessarily collected by the sediment traps, and (2) the temporal lag between the production and the export which is difficult to assess (Nodder and Waite, 2001). The mesocosm approach was designed to overcome these experimental limitations. The shallow depth of the traps (~15 m) and the absence of NO_3^- normally supplied via the nitracline prevent any comparison with open ocean studies. Nevertheless, the mesocosm approach enables a comparison of the export efficiency under contrasting ecological situations. In this case, the period dominated by DDAs (P1) is compared with the period dominated by UCYN-C (P2).

During P1, the biomass was stable in the mesocosms and the amount of recently fixed N₂ was equal to the amount of exported PON, suggesting a tight coupling between the two processes (Fig. 6). It has been shown that large aggregates of the diatom Rhizosolenia spp., representing the majority of DDAs during P1 (Turk et al., 2015), can sink at high rates (Villareal et al., 1996). This suggests that during this experiment, the recently fixed N₂ by DDAs remained within the symbiotic association and was quickly exported in the settling particles. This agrees with Karl et al. (2012) who showed that DDAs support the export pulses regularly observed in late summer in the tropical North Pacific Ocean.

During P2, the increase in PON concentrations (Fig. 3B) suggests that part of the freshly produced biomass remained in the water column. The accumulation of PON probably favored remineralization processes, explaining the increase in NH_4^+ concentrations. This may have enhanced the transfer of the recently fixed N_2 to the non diazotrophic plankton as demonstrated by Bonnet et al. (2015a) and explains the development of picocyanobacteria (Leblanc et al., 2015). Additionally, the total amount of N provided by N₂ fixation did not account for all the exported PON during P2 (Fig. 6), implying that an additional N source played a significant role in promoting the export. The only alternative N source is DON which, indeed, exhibited a significant decrease in concentration of 0.9 \pm 0.7 μ mol L⁻¹ in the mesocosms during P2 (see section 4.3 for further discussion on DON consumption). Assuming that DON and N_2 fixation are the only possible sources of N in the mesocosms, we calculated that a DON use of 0.9 μ mol L⁻¹ would have supported ~78 % of the PON production during P2, and potentially fueled the PON export to the same extent. This is in agreement with Torres-Valdes et al. (2009) and Letscher et al. (2013) who showed that DON pool is a dynamic contributor of the N cycle able to supports up to 40 % of the vertical PON export in the oligotrophic gyres of the Pacific and Atlantic Oceans. A quantification of the diazotrophs in the sediment traps, performed on day 19, shows that ~10 % of the UCYN-C biomass in the mesocosms was exported this day, explaining ~ 7 % of the PON export (Bonnet et al., 2015a). Thus, the recently fixed N_2 by UCYN-C can directly be exported but is probably more efficiently transferred to non diazotrophic plankton through remineralization processes.

The contrast between P1 and P2 is also observed using the e-ratio. The production driven by UCYN-C was more efficient in promoting POC export than the production driven by DDAs. During P1, it is probable that the recently fixed C by DDAs remained within the symbiotic association and sunk with the recently fixed N₂ constituting a direct and net C export. During P2, the higher efficiency of C export strongly suggests that the DON ultimately fueled PP which, in turn, increased POC export. Additionally, when UCYN-C dominated, an enhanced N

remineralization may have enabled more C to be fixed per unit of fixed N_2 leading to a higher e-ratio. A proportionally higher N remineralization following high PP and N_2 fixation rates is supported by similar findings in the western North Pacific warm pool (Shiozaki et al., 2013).

4.3. The unexpected high dissolved organic matter consumption

The use of dissolved organic compounds and their implications on PP in the open ocean has long been demonstrated (Antia et al., 1991). The use of DOP by plankton communities in the oligotrophic ocean has been observed in the North Pacific Ocean (Bjorkman and Karl, 2003) and in the Atlantic Ocean (Lomas et al., 2010; Mather et al., 2008) and generally occurs under PO_4^{3-} limitation. In this study, the decrease in DOP concentrations during P2 occurred when T_{PO4} reached the lowest levels, confirming the ability of the planktonic community to use DOP under low PO_4^{3-} availability. More surprisingly, the significant and rapid decrease in DON concentrations (Fig. 4) observed during the development of UCYN-C (P2) in the mesocosms was associated with a rapid increase in PP (Fig. 1C), biomass (Figs. 2 and 3) and bacterial production (BP) (Van Wambeke et al., 2015), suggests high consumption of DON directly or indirectly by primary producers. In the open ocean, DON is mainly refractory. Nevertheless, it is now recognized that a fraction of the DON is labile and can directly support phytoplankton growth, while a semi labile fraction can be mineralized by bacterioplankton (Antia et al., 1991; Bronk, 2002; Bronk et al., 2007). In this study, we propose three hypotheses that could explain the observed decrease in DON concentrations during P2: (i) bacterial mineralization of DON triggered by high PP, (ii) direct uptake of DON by primary producers including UCYN-C and (iii) abiotic photo-degradation of DON into NH₄⁺.

(i) The increase in PP driven by high N₂ fixation rates during P2 led to an increase in bacterial production (Van Wambeke et al., 2015). The significant negative correlation between BP and DON concentrations (Spearman rank correlation, r=-0.35; p< 0.001) indicate significant consumption of DON by bacterial mineralization. Diazotrophs are known to over-fix C relative to N (Mulholland et al., 2007), which may explain why the POC:PON ratio was above the Redfield ratio during the experiment. The resulting N deficit for bacterial mineralization may have been found in the labile or semi labile DON pool. This hypothesis is supported by Van Wambeke et al. (2015) who showed that BP was limited by N availability in the mesocosms during the experiment. Based on BP data and assuming a bacterial growth efficiency of 10 to 30 % (del Giorgio and Cole, 1998) and a C:N ratio of 6.6 in bacteria cells (Fukuda et al., 1998), bacterial respiration would have led to a DON consumption of 0.2 to 0.7 µmol L⁻¹ during P2, supporting at least part of the DON removal of ~0.9 µmol L⁻¹ reported here.

(ii) An alternative explanation for the decrease in DON concentrations is direct consumption by primary producers. Cyanobacteria are known to use DON compounds such as urea (Collier et al., 2009; Painter et al., 2008) to such an extent that DON has been reported to be the main source of N fueling cyanobacterial blooms in coastal waters (Glibert et al., 2004). The DON decrease occurring during the development of UCYN-C, whose abundances reached 5.10^5 nifH copies L⁻¹ (Turk et al., 2015), questions their ability to use DON to meet their N requirements. Direct uptake of glutamate and amino acids (constitutive components of the DON pool) has been reported in natural and laboratory populations of Trichodesmium spp. (Mulholland and Capone, 1999; Mulholland et al., 1999). Furthermore, large decrease in DON concentrations were observed after blooms of the diazotroph Aphanizomenon ovalisporum in Lake Kinneret (Berman, 1997). The hypothesis of a direct use of DON by A. ovalisporum was confirmed in culture experiment where the development of this diazotroph was stimulated by DON additions (Berman, 1997, 1999). To our knowledge, no direct uptake measurements of DON compounds have been performed on UCYN. However, the ureA gene implicated in the urea assimilation has been identified in the cyanobacterial diazotrophic strain Cyanothece PCC 7822 (Bandyopadhyay et al., 2011) closely related to the UCYN-C cluster. These pieces of evidences suggest that in addition to N₂ fixation, UCYN-C might be able to use the DON pool to meet their N requirements.

(iii) Finally, photo-degradation could be a possible sink of DON in surface waters (Bronk, 2002). A field study performed in the ultraoligotrophic eastern basin of the Mediterranean Sea indicates a production of NH_4^+ from DON of 0.2-2.9 nmol N L⁻¹ d⁻¹ in surface waters (Kitidis et al., 2006). Taking into account the highest rates reported above, this process cannot explain more than 10 % of the observed DON removal. Moreover, the DON decrease occurred only during P2 whereas photo-degradation would be occurring continuously over the experiment.

The first two hypotheses (i and ii) are more likely to explain the DON decrease during P2. None of these two hypotheses can be excluded even though direct proof of large uptake of DON by UCYN-C is lacking. Thus, in this study the DON use was directly or indirectly triggered by the UCYN-C activity.

Conclusions

This study confirms that in the South West Pacific, N₂ fixation is a biogeochemically relevant process able to provide sufficient new N to drive new PP, biomass accumulation and organic matter export as long as P is not limiting. The fate of the recently fixed N appears to be closely related to the diazotrophic community involved in N₂ fixation. A strong coupling of N₂ fixation and PON export occurred when DDAs dominated the diazotrophic community suggesting their direct export. When the community was dominated by UCYN-C, biomass accumulation was observed together with an efficient particulate export. A significant decrease in DON concentrations was observed during the same period indicating a direct or indirect use of DON by UCYN-C. Thus, in addition to fueling primary production, UCYN-C appear to be able to enhance regenerated production based both on the transfer of recently fixed N₂ toward non fixing planktonic groups and on the use of the DON pool. This use of DON exceeded the new N provided by N₂ fixation even though the N₂ fixation rates were among the highest reported in literature for the global ocean. These results suggest that DON has to be considered as a dynamic pool in LNLC areas as it may provide significant amounts of N and contribute significantly to particulate export.

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Dinitrogen fixation and dissolved organic nitrogen fueled primary production and particulate export during the VAHINE mesocosms experiment (New Caledonia lagoon)

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Abstract

In the oligotrophic ocean characterized by nitrate (NO_3) depletion in surface waters, dinitrogen (N_2) fixation and dissolved organic nitrogen (DON) can represent significant nitrogen (N) sources for the ecosystem. Here we deployed in New Caledonia large in situ mesocosms in order to investigate (1) the contribution of N₂ fixation and DON use to primary production (PP) and particle export and (2) the fate of the freshly produced particulate organic N (PON) i.e. whether it is preferentially accumulated and recycled in the water column or exported out of the system. The mesocosms were fertilized with phosphate ($\frac{PPO_4^{3-}}{1}$) in order to prevent P-limitation phosphorus (P)-limitation and promote N_2 fixation. The diazotrophic community was dominated by diatoms-diazotrophs associations (DDAs) during the first 10 part of the experiment for 10 days (P1) followed by the unicellular N2-fixing cyanobacteria UCYN-C the 9 last last 9 days (P2) of the experiment. N₂ fixation rates averaged 9.8 ± 4.0 and $27.7 \pm 8.6 \text{ nM d}^{-1}$ nmol L⁻¹ d⁻¹ during P1 and P2, respectively. NO₃⁻ concentrations $(< 0.04 \,\mu\text{mol}\,\text{L}^{-1})$ in the mesocosms were a negligible source of N indicating that N₂ fixation was the main driver of new production all along the experiment. The contribution of 15 N₂ fixation to PP was not significantly different (p > 0.05) during P1 (9.0 ± 3.3 %) and P2 $(12.6 \pm 6.1 \%)$. However, the eratio ratio that quantifies the efficiency of a system to export particulate organic carbon (POC_{export}) compared to PP ($e_{ratio-ratio} = POC_{export}/PP$) was significantly higher (p < 0.05) during P2 (39.7 ± 24.9 %) than during P1 (23.9 ± 20.2 %), indicating that the production sustained by UCYN-C was more efficient at promoting C export 20 than the production sustained by DDAs. During P1, PON was stable and the total amount of N provided by N₂ fixation $(0.10 \pm 0.02 \text{ M} \mu \text{mol L}^{-1})$ was not significantly different (p > 0.05)from the total amount of PON exported ($0.10 \pm 0.04 \frac{M}{\mu}$ mol L⁻¹), suggesting a rapid and probably direct export of the recently fixed N_2 by the DDAs. During P2, both PON concentrations and PON export increased in the mesocosms by a factor 1.5-2. Unlike in P1, this 25 PON production was not totally explained by the new N provided by N_2 fixation. The use of DON, whose concentrations decreased significantly (p < 0.05) from 5.3 ± 0.5 M-µmol L⁻¹ to 4.4 ± 0.5 M µmol L⁻¹, appeared to be the missing N source. The DON consumption of about $0.9(\sim 0.9 \text{ M} \text{ }\mu\text{mol }L^{-1})$ during P2 is even higher than the total amount of new N brought by N₂ fixation (about $0.25 \sim 0.25 \text{ M}\mu\text{mol }L^{-1}$) during the same period. These results suggest that while DDAs mainly rely on N₂ fixation for their N requirement/requirements, both N₂ fixation and DON can be significant N-sources for primary production and particulate export following UCYN-C blooms in the New Caledonia lagoon and by extension in the N-limited Ocean oceans where similar events are likely to occur.

1 Introduction

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Nitrogen (N) availability constitutes one of the most limiting factors for marine primary production (PP) (Falkowski, 1997Moore et al., 2013). About 80 % of the global ocean surface is depleted in dissolved inorganic N (nitrate (NO₃⁻) and ammonium (NH₄⁺) < 1 µmol L⁻¹) and characterized by low PP, low biomass and low particulate matter export fluxes (Longhurst, 2007). In these Low Nitrate, Low Chlorophyll (LNLC) ecosystems, the strong stratification of the photic surface layer prevents the mixing with NO₃⁻-replete deep waters and imposes on phytoplanktonic communities to rely on alternative N sources for growth. These sources the virtually inexhaustible dissolved dinitrogen (N₂) pool (~400~400 µmol L⁻¹), and the large (~5~5 µmol L⁻¹) but mainly refractory dissolved organic N (DON) pool.

The first pool (N₂) is only accessible to prokaryotic organisms possessing the nifH gene and able to reduce the N₂ gas molecule into bioavailable NH₄⁺. This process called N₂ fixation (or diazotrophy) is responsible for the main external source of N for the upper ocean (Gruber and Galloway, 2008; Mahaffey, 2005) and fuels PP in LNLC ecosystems (e.g. Dugdale and Goering, 1967; Karl et al., 1997). However, the fate of recently fixed N₂ in the planktonic food web and its potential impact on carbon export are poorly understood. Moreover, this fate may differ according to the diazotrophic species involved in N₂ fixation. The widely distributed filamentous cyanobacterium *Trichodesmium* spp., one of the main contributors to global N₂ fixation (Capone et al., 1997), is rarely found in sediment traps (Chen

et al., 2003; Walsby, 1992) indicating that *Trichodesmium* <u>spp</u> has a low direct export efficiency. However, a recent study performed in the South West Pacific indicates that N fixed

by *Trichodesmium* <u>spp</u>. is preferentially and rapidly (within few days) transferred to diatoms and bacteria (Bonnet et al., 2015b), potentially resulting in indirect carbon export. Conversely, diatoms-diazotrophs associations (DDAs) drive an efficient biological carbon pump in the Amazon River plume (Subramaniam et al., 2008) and in the North Pacific Gyre (Dore et al., 2008; Karl et al., 2012), indicating an efficient export of the N fixed by these organisms. Finally, unicellular N₂-fixing cyanobacteria (UCYN), are presumably the most abun-

dant in the global ocean and are also major contributor contributors to global N₂ fixation (e.
 g. Moisander et al., 2010; Montoya et al., 2004). However, little is known regarding the fate of the N fixed by UCYN, whether it is directly or indirectly exported out of the euphotic zone
 or recycled in surface waters (Thompson and Zehr, 2013).

The second N pool (DON) may constitute a significant N source for planktonic communities but remains poorly constrained (Bronk, 2002). The DON pool is a "Black box" composed of various chemical products more or less refractory with their own specific turnover time (Bronk et al., 2007). The persistence of high DON concentrations in surface oceanic waters has formerly led to consider that it is unavailable for the marine biota. However, the determination of DON concentrations is submitted to high analytical uncertainties (Czerny et al., 2013) that may hide low but ecologically relevant changes in concentration concentrations resulting from consumption of the labile or semi labile fraction of the DON pool. Furthermore, due to the heterogeneous composition of DON, isotopic labeling experiments using tracers are difficult to conduct, which explains the lack of information on the fluxes transiting in and out of the DON pool.

- in and out of the DON pool (Bronk, 2002; Bronk et al., 2007). Although heterotrophic bacteria are presumably the main users of this organic pool, it has been shown that primary producers can also use it to meet their N requirements (Antia et al., 1991; Berman and Bronk, 2003). Similarly to fixed fixed N₂, the fate of DON assimilated by marine plankton ²⁵ will depend on the consumers, whether they would remineralize or export the particulate
 - organic matter produced.

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Studying the fate of N in the ocean is complex as it requires to follow the biogeochemical characteristics, the succession of planktonic species and the potential export from the same water mass for several weeks. In the open ocean, such studies are further complicated by

physical processes (e.g. lateral advection) that spread the water masses. Here, we isolated a part of the water column from physical dispersion using in situ large mesocosms (52 m^3) equipped with sediment traps in order to overcome this issue. The objectives of this study were (1) to investigate the contribution of N₂ fixation and DON use on PP and particle export, and (2) to trace the fate of these N sources in the ecosystem, i.e. whether the freshly produced particulate organic N (PON) is accumulated or exported out of the system.

The mesocosms were deployed in the subtropical New Caledonian lagoon (South West Pacific), characterized by LNLC conditions (Fichez et al., 2010), where high N_2 fixation rates and abundances of *Trichodesmium* spp. and UCYN communities have been reported

- (Biegala and Raimbault, 2008; Garcia et al., 2007; Rodier and Le Borgne, 2008, 2010).
 Dissolved inorganic phosphorus (DIPPhosphate (PO₄³⁻) availability has previously been reported to control N₂ fixation in the South West Pacific (Moutin et al., 2005, 2008). In order to avoid phosphorus (P) PO₄³⁻ limitation and to create favorable conditions for diazotrophs growth, the mesocosms were fertilized with DIP PO₄³⁻ at the beginning of the experiment.
- ¹⁵ Diazotrophs developed extensively in the mesocosms during the 23-days experiment (Turk et al., 2015). The diazotroph community was dominated by DDAs during the first half of the experiment (from day 5 to day 14, hereafter called P1), then shifted towards a large dominance of UCYN from group C (closely related to *Cyanothece* sp.) which developed extensively during the second half of the experiment (from day 15 to day 23, hereafter called P2). PP and N₂ fixation rates were monitored for the 23 days together with C, N and P pools dynamics in the water column and in export material. The results are discussed under the light of the shift of dominance between the two diazotrophic communities.

2 Material and method

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2.1 Study site and mesocosm description

²⁵ Three mesocosms were deployed at the exit of the oligotrophic New Caledonian lagoon (22°29.1' S–166°26.9' E), 28 km off the coast of New Caledonia from 13 January 2013

(day 1) to 4 February 2013 (day 23). The site was 25 m depth on sandy bottom, protected from the dominant trade winds (southeast sector) and characterized by a high influence of oceanic oligotrophic waters coming from outside the lagoon through the Boulari passage (Ouillon et al., 2010). The complete description of the mesocosms design is detailed in Bonnet et al. (2015c). Briefly, the enclosures were cylindrical bags 2.3 m in diameter and reaching about 15 m deep into the water. The bags were maintained 1 m above the surface to prevent external water inclusions. They were supported by a polyethylene frame and kept

at the surface with floats. The bags were straightened by weights at the bottom of the mesocosms. After deployment, the mesocosms were left opened at the bottom for 24 h to insure a total homogeneity in the water column. On day 1, the bottom was closed with a sediment trap consisting in a funnel shape end fitted with a 3" adapter for fastening a collection bottle for the daily sinking material.

The DIP PO_4^{3-} fertilization was conducted in the evening of day 4. The fertilization consisted in an addition in each mesocosm of 20 L of a filtered seawater solution enriched with KH₂PO₄ (41.6 mM) leading to a final concentration of $\sim 800 \sim 0.8$ nM µmol L⁻¹ in the mesocosms. To insure homogenization, the solution was added to each mesocosm using a polyethylene tubing connected to a Teflon pump lifted regularly from the bottom to the surface of the mesocosms.

2.2 In situ monitoring and water sampling

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CTD casts and water collection were conducted daily in each of the three replicate mesocosms (hereafter called M1, M2, and M3) and in surrounding waters. Seawater sampling was performed from a floating platform moved around the mesocosms. CTD casts were performed at 10 a.m. (local time) in each mesocosm and in surrounding waters using a memory probe SBE 911 plus (Sea-Bird Electronics, Inc.) equipped with conductivity, turbidity, fluorimetry, temperature and dissolved oxygen sensors. The CTD was handled with a speed of 0.2–0.3 m s⁻¹. The water was collected just after before the CTD casts at three depths in each mesocosm (1, 6 and 12 m) using an air-compressed Teflon pump (AstiPure[™]) connected to a polyethylene tubing. Samples for particulate and dissolved organic and inorganic matter (C, N and P) and PP determination were first collected in 50 L polypropylene carboys at the three depths and sub-sampled back on the R/V *Alis* moored 1 nautical mile away from the mesocosms site. Samples for N₂ fixation rate and PP determination were directly collected from the pump in polycarbonate bottles (4.5 L) for each depth in each mesocosm and in surrounding waters. Sinking material was collected every day by divers from sediment traps –as decribed in Bonnet et al. (2015c).

2.3 Primary production rates and dissolved inorganic phosphorus phosphate turnover time

PP rates and DIP PO₄³⁻ turnover time (*T*_{DIP}-*T*_{PO4} i.e. the ratio of DIP PO₄³⁻ concentration and uptake), were measured using the ¹⁴C/³³P dual labeling method (Duhamel et al., 2006). 60 mL bottles were amended with ³³P and ¹⁴C and incubated for 3 to 4 h on a mooring line close to the mesocosms site at the sampling depths. After incubation, Incubations were stopped by adding 50 μL of KH₂PO₄ solution (10) was added mmol L⁻¹) in order to reduce to a minimum the ³³P assimilation by dilution and kept effect and by keeping them in the dark to stop DIC uptake. Samples were then filtered on 0.2 μmol L⁻¹ polycarbonate membrane filters, and placed into scintillation vials with 250 μL of HCI 0.5 M. After 12 h, 5 mL of scintillation liquid (ULTIMA Gold MV, PerkinElmer Inc.) was were added to each vial before the first count on a Packard Tri-Carb 2100TR scintillation counter. The activity of ³³P and ¹⁴C was separated using a second count made 5 months later taking into account the

²⁰ half-life of the two isotopes.PP and T_{DIP} ³³P: 25.38 days (Duhamel et al., 2006). PP and T_{PO4} were calculated according to Moutin et al. (2002).

2.4 N₂ fixation rates

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Samples for N₂ fixation incubations were collected in 4.5 L polycarbonate bottles. The latter were amended with ¹⁵N₂ enriched seawater according to the protocol developed by Mohr et al. (2010). Briefly, the ¹⁵N₂ enriched seawater was prepared daily from 0.2 μ mol L⁻¹ filtered seawater collected from the same site in a 4.5 L polycarbonate bottle. Seawater was

first degassed through a degassing membrane (Membrana, Minimodule, flow rate fixed at 450 mL min^{-1}) connected to a vacuum pump (< 200 mbar absolute pressure) during for at least 1 h. It was then tightly closed with no head space with a silicone septum cap and amended with 1 mL of ¹⁵N₂ (98.9% Cambridge isotope) per 100 mL. The bottle was then shaken vigorously and incubated overnight at 3 bars (20 m depth) to promote ¹⁵N₂ disso-5 lution. Incubation bottles were then amended with 5% (vol:vol) of $^{15}N_2$ enriched seawater and closed without headspace with silicone septum capsand. The latter were incubated on an in situ mooring line close to the mesocosms at the appropriate sampling depths for 24 h. After of incubation, 12 mL of incubated water was sampled in Exetainers on 10 replicate samples and analyzed using a Membrane Inlet Mass Spectrometer (Kana et al., 1994) to 10 estimate the final enrichment of the ${}^{15}N_2$ pool during the incubation. The measured final 15 N/ 14 N ratio of the N₂ in the incubation bottles was found to be 2.4 ± 0.2 atom% (n = 10). Samples were then filtered on combusted (450 °C, 4 h) GF/F filters and stored at -20 °C for the duration of the cruise. Every day, T0 samples were spiked with $^{15}N_2$ and immediately filtered in order to determine the initial background of ¹⁵N/¹⁴N ratio of PON for calculation of 15 N₂ fixation rates. Filters were then dried at 60 °C during 24 h prior to analysis using a mass spectrometer (Delta plus, Thermo Fisher Scientific) coupled with an elemental analyzer (Flash EA, Thermo Fisher Scientific) for the concentration of PON concentrations and PON ¹⁵N enrichment determination. Standard deviation was 0.004 µmol for PON and 0.0001 atom% for the ¹⁵N/¹⁴N isotopic ratio. The fluxes were defined as significant when ¹⁵N en-20 richment was higher than three times the standard deviation obtained on initial T0 samples. The fluxes were calculated according to the equation detailed in Montoya et al. (1996). A recent study (Dabundo et al., 2014) reports potential contamination of commercial $^{15}N_2$ gas stocks with ¹⁵N-enriched NH_4^+ , NO_3^- and/or nitrite (NO_2^-), and nitrous oxide. The ¹⁵N₂ Cambridge Isotopes stocks analyzed contained low concentrations of ¹⁵N contaminants, 25 and the potential overestimated N₂ fixation rates modeled using this contamination level would range from undetectable to $0.02 \text{ nmol N L}^{-1} \text{ d}^{-1}$. These rates are in the lower end of the range of rates measured in this study and we thus considered that this issue did not affect the results reported here.

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2.5 Chlorophyll *a*, inorganic and organic matter analyses

Samples for chlorophyll *a* (Chl *a*) concentrations determination were collected by filtering 550 mL of seawater on GF/F filters. Filters were directly stored in liquid nitrogen. Chl *a* was extracted in methanol and measured by fluorometry (Herbland et al., 1985).

- Samples for total organic carbon (TOC) concentrations were collected in duplicate at only one depth (6 m) in each mesocosm and in surrounding waters in precombusted (450 °C, 4 h) 12 mL sealed glassware flasks, acidified with H₃PO₄ and stored in the dark at 4 °C until analysis. Samples were analyzed on a Shimadzu TOCV analyzer with a typical precision of 2 μmol L⁻¹. Samples for particulate organic carbon (POC) concentrations were collected
- by filtering 2.3 L of seawater through precombusted (450 °C, 4 h) GF/F filter and determined using the combustion method (Strickland and Parsons, 1972) on an EA 2400 CHN analyzer. Filters were not acidified to remove inorganic carbon as it is assumed to be < 10 % of the total particulate C (Wangersky, 1994). Dissolved organic carbon (DOC) concentrations were calculated as the difference between TOC and POC concentrations. The DOC precision
 calculated from the analytical precision of each term according to the errors propagation
 - law was $5 \mu mol L^{-1}$.

Samples for NH_4^+ were collected in 40 mL glass vials and analyzed by the fluorescence method according to Holmes et al. (1999) on a trilogy fluorometer (Turner Design). The detection limit was 0.01 µmol L⁻¹. Samples for NO_3^- , NO_2^- , $\frac{\text{DIP}PO_4^{3-}}{100}$, total N (TN) and to-

- tal P (TP) concentrations determination were collected in 40 mL glass bottles and stored at -20 °C until analysis. NO₃⁻, NO₂⁻ and DIP PO₄³⁻ concentrations were determined on a segmented flow analyzer according to Aminot and Kérouel (2007). The detection limit was 0.01 and 0.005 µmol L⁻¹ for NO₃⁻ + NO₂⁻ and DIP PO₄³⁻, respectively. TN and TP concentrations were determined according to the wet oxidation procedure described in Pujo-Pay and Raim bault (1994). The precision was 0.5 µmol L⁻¹ and 0.02 µmol L⁻¹ for TN and TP, respectively. Samples for PON and particulate organic P (POP) concentrations were collected by filter
 - ing 1.2 L of water on precombusted (450 °C, 4 h) and acid washed (HCl, 10%) GF/F filters and analyzed according to the wet oxidation protocol described in Pujo-Pay and Raimbault

(1)

(1994) with a precision of 0.06 and 0.007 μ mol L⁻¹ for PON and POP, respectively. DON concentrations were calculated from TN concentrations subtracted by PON, NO₃⁻, NO₂⁻ and NH⁺₄ concentrations. Dissolved organic P (DOP) concentrations were calculated from TP concentrations subtracted by POP concentrations and DIP and PO_4^{3-} concentrations. The precision calculated according to the propagation law of analytical precision associated with each parameter was 0.5 and 0.03 μ mol L⁻¹ for DON and DOP, respectively.

Samples from sediment traps were collected daily and preserved in a 5% buffered solution of formaldehyde and stored at 4 °C until analysis. All the swimmers were handpicked from each sample, and were found to be negligible source a negligible source of particu-

late matter compared to the total particulate matter exported (< 5%). Samples were then 10 desalted using ultrapure water (Milli-Q grade) and frozen dried. The daily amount of POC exported (POC_{export}), and PON exported (PON_{export}) were measured using a CHN analyzer (Perkin Elmer 2400). The POP exported (POP_{export}) was measured after mineralization using nitric acid and further determination of mineralized P according to Pujo-Pay and Raimbault (1994). 15

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2.6 Data presentation and statistical analyses

The build-up of an elemental mass balanced budget is theoretically accessible as all the stocks and fluxes were sampled in the mesocosms. However, the attempts of closing a elemental mass budget in similar mesocosms studies were limited by the large analytical uncertainties on the suspended organic pool determination (Czerny et al., 2013; Guieu et al., 2014). Nevertheless, as the only incoming (N₂ fixation) and export fluxes (PON_{export}) of N were characterized accurately, we were able to calculate the change in TN content (ΔTN_{calc}) of a mesocosm. It was defined as:

 $\Delta TN_{calc} = \Sigma N_{2 \text{ fix}} + \Sigma PON_{export}$

where $\Sigma N_{2,fix}$ and ΣPON_{export} are the cumulated average over depth N₂ fixation rates and 25 PON_{export}. This approach does not discriminate the different N pools in the water column but allows a precise evaluation of the TN variation in the mesocosm and a direct comparison of the N_2 fixation and the PON_{export}.

Considering the absence of significant differences between the three depths sampled of most of the concentrations and fluxes measured (paired non parametric Friedman test,

- $\alpha = 0.05$), the values The values of fluxes and concentrations presented in the text are averaged over the three depths (no significant differences have been evidenced between depths, paired Friedman test, $\alpha = 0.05$). Statistical differences between each mesocosm or between the mesocosms and surrounding waters were tested using the paired non parametric Wilcoxon signed-rank test ($\alpha = 0.05$) for each parameter presented. If no significant
- ¹⁰ differences between the mesocoms were detected, the values were averaged between the mesocosms. The associated uncertainties were calculated as the analytical precision cumulated to the standard deviation of each term according to the propagation of errors law. Differences between P1 and P2 were tested using the non parametric Kruskal–Wallis test ($\alpha = 0.05$).

15 3 Results

3.1 Hydrological background

The detailed description of hydrological and inorganic nutrients conditions during the experiment is extensively presented in Bonnet et al. (2015c). Briefly, seawater temperature increased inside the mesocosms and in surrounding waters from 25.5 to 26.7 °C over the course of the experiment, similarly inside the mesocosms and in surrounding waters. The water column was well mixed in the mesocosms as temperature and salinity were homogeneous with depth over the course of the experiment. The sum of NO₃⁻ and NO₂⁻ concentrations averaged over depth in the mesocosms were below 0.04 µmol L⁻¹ the day before the DIP PO₄³⁻ fertilization (day 4) and decreased to 0.01 toward µmol L⁻¹ towards the end of the experiment. NH₄⁺ concentrations were close to the detection limit of 0.01 µmol L⁻¹ from day 1 to day 18 and increased in all the mesocosms up to 0.06 µmol L⁻¹ toward the end of

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the experiment. Prior to the DIP fertilization, DIP PO_4^{3-} fertilization, PO_4^{3-} concentrations in the mesocosms ranged from 0.02 to 0.05 µmol L⁻¹. The day after the fertilization, DIP PO_4^{3-} concentrations reached ~ 0.8 µmol L⁻¹ in all mesocosmsand decreased steadilyduring the course of the experiment and tended. Then, the concentrations decreased steadily, tending to initial concentrations (0.02–0.08 µmol L⁻¹) at the end of the experiment. In surrounding waters, NO_3^- remained below 0.20 and DIP µmol L⁻¹ and PO_4^{3-} averaged 0.05 µmol L⁻¹ all along the experiment.

3.2 DIP Phosphate turn over time

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The evolution of $T_{\text{DIP}} T_{\text{PO4}}$ was closely related to the dynamics of $\overline{\text{DIP}} \text{-} \text{PO}_4^{3-}$ concentrations. Before the $\overline{\text{DIP}}$ fertilization, $T_{\text{DIP}} \text{-} \text{PO}_4^{3-}$ fertilization, T_{PO4} decreased from 1.0 ± 0.1 d on day 3 to 0.4 ± 0.1 d on day 4 in all the mesocosms (Fig. 1a1A). At the start of P1, $T_{\text{DIP}} - T_{\text{PO4}}$ dramatically increased in all mesocosms reaching 35.7 ± 15.7 , 30.1 ± 8.6 and 35.8 ± 10.5 d in M1, M2 and M3respectively. T_{DIP} , respectively. T_{PO4} then decreased steadily in all the mesocosms reaching 1 d on day 14, day 19 and day 21 for in M1, M2 and M3, respectively.

At the end of the experiment (day 23), T_{DIP} - T_{PO4} values were the lowest reached over the experiment in all the mesocosms with values below 0.2 d. In surrounding waters, T_{DIP} - T_{PO4} was stable around 1.8 ± 0.7 d from the start of the experiment to day 15 and then decreased to reach 0.5 ± 0.1 d on day 23 (Fig. 1a1A).

3.3 N₂ fixation rates and primary production rates

Before the DIP-PO₄³⁻ fertilization, N₂ fixation rates inside the mesocosms were 17.4±
7.3 nmol N L⁻¹ d⁻¹ and decreased the days after following the fertilization (Fig. 1b1B). During P1, the average N₂ fixation rates in the mesocosms were 9.8±4.0 nmol N L⁻¹ d⁻¹. During P2, N₂ fixation rates in the mesocosms were significantly (*p* < 0.05) higher than during P1 averaging 27.7±8.6 nmol N L⁻¹ d⁻¹. N₂ fixation rates were not significantly different (*p* > 0.05) between the three mesocosms all along the experiment. In surrounding

waters, N₂ fixation rates did not show any clear pattern along the experiment and averaged 9.2 nmol N L⁻¹ d⁻¹, ranging from 1.9 to 29.3 nmol N L⁻¹ d⁻¹ (Fig. 1b1B).

The day before the DIP PO₄³⁻ fertilization, PP was not significantly different (p > 0.05) between the three mesocosms and averaged $0.4 \pm 0.1 \,\mu$ mol C L⁻¹ d⁻¹ (Fig. 1e1C). Dur-

- ing P1, PP increased steadily in the mesocosms to reach 0.9±0.1 C d⁻¹ μmol C L⁻¹ d⁻¹ at the end of P1. During P2, while T_{DIP} T_{PO4} was decreasing, PP increased faster than during P1 continued to increase in the mesocosms with values generally exceeding 1.5 μmol C L⁻¹ d⁻¹. During P2, PP was significantly higher in M3 than in M1 and M2 (p < 0.05). In surrounding waters, PP was stable before and during P1 at 0.9±0.3 μmol C L⁻¹ d⁻¹.
 and increased during P2 reaching 1.5±0.2 μmol C L⁻¹ d⁻¹ on day 23. Over the whole ex-
- periment, PP in the mesocosms was not significantly different from surrounding waters (p > 0.05) except for in M3 during P2 (p < 0.05).

Assuming that all the diazotrophs are primary producers and have a C:N fixation ratio of 6.6 (Redfield, 1934), we calculated that N₂ fixation sustained 10.8 ± 5.0 % (range 3.7– 32.2%) of PP in the mesocosms after the DIP PO₄³⁻ fertilization and 5.7 ± 2.0% (range 2.2–9.1%) in surrounding waters. The contribution of N₂ fixation to PP was not significantly different (p > 0.05, n = 57) during P1 (9.0 ± 3.3%) and P2 (12.6 ± 6.1%)(Fig. 2).

3.4 Chl a and particulate organic matter dynamics

The day before the fertilization, Chl *a* concentrations in the mesocosms were $0.21 \pm 0.05 \,\mu\text{g} \,\text{L}^{-1}$ (Fig. 32). During P1, Chl *a* did not show any clear pattern; concentrations were in the 0.12 to $0.28 \,\mu\text{g} \,\text{L}^{-1}$ range. During P2, Chl *a* increased in all the mesocosms but to a greater extend in M3 compared to M1 and M2 and reached maximal depth-averaged concentrations of 0.55 ± 0.01 , 0.47 ± 0.08 and $1.29 \pm 0.22 \,\mu\text{g} \,\text{L}^{-1}$ for in M1, M2 and M3, respectively. Before and during P1, Chl *a* concentrations in surrounding waters were close to the concentrations in the mesocosms and ranged between 0.09 and $0.28 \,\mu\text{g} \,\text{L}^{-1}$. During P2, they increased but to a lower extent than in the mesocosms with daily averaged concentrations of $0.42 \pm 0.03 \,\mu\text{g} \,\text{L}^{-1}$ on day 23.

The day before the $\frac{\text{DIP}}{\text{PO}_{4}^{3-}}$ fertilization, POC concentrations ranged between from 9 and to 15 µmol L⁻¹ (Fig. 4a3A). During P1, POC concentrations did not show any clear pattern in the mesocosms and concentrations ranged between from 6 and to 13 μ mol L⁻¹. During P2, POC concentrations increased in M3 reaching $18 \mu mol L^{-1}$ on day 21, whereas they remained stable in M1 and M2. POC concentrations in surrounding waters were stable all along the experiment and were significantly lower (p < 0.05) than in the mesocosms. Initial PON concentrations were about $0.9 \,\mu$ mol L⁻¹ and remained relatively stable during P1 (Fig. 4b3B). During P2, PON concentrations increased in all the mesocosms by a factor 1.5 in M1 and M2, and by a factor 2 for in M3 at day 23 reaching 2 μ mol L⁻¹. PON concentrations also increased outside but to a lesser extent with values remaining below 1 .- POP 10 μ mol L⁻¹. POP concentrations showed the same pattern than PON concentrations. The day before the $\frac{\text{DIP}}{\text{PO}_{4}}$ PO³⁻ fertilization, POP concentrations were not significantly different (p > 0.05) between the mesocosms and averaged 0.05 µmol L⁻¹ (Fig. 4c3C). During P1, the concentrations in the mesocosms remained relatively stable. During P2, POP concentrations increased in all the mesocosms but to a higher extent in M3, reaching 0.07, 0.07 15 and 0.12 μ mol L⁻¹ in M1, M2 and M3, respectively. Particulate POC/PON ratio decreased during the experiment from the initial averaged ratio of 12 to 8 at the end of the experiment, but remained higher than the Redfield ratio (6.6).

3.5 Dissolved organic matter dynamics

DOC concentrations ranged from 50 to 74 µmol L⁻¹ (average value of 60 ± 4 µmol L⁻¹) over the course of the experiment without any clear trend over the 23 days (Fig. 5a4A). Furthermore no significant differences were measured between the mesocosms and surrounding waters (*p* > 0.05). Before the DIP PO₄³⁻ fertilization, DON concentrations averaged 5.2±0.5 µmol L⁻¹ on day 4 (Fig. 5b4B). Concentrations remained stable during P1
in and out the mesocosms. In contrast during During P2, DON concentrations decreased significantly (*p* < 0.05) in the three mesocosms. The decrease was of 0.7±0.5, 0.8±0.5 and 1.0±0.5 µmol L⁻¹ between day 17 and day 23 in M1, M2 and M3, respectively, and was not significantly different between the mesocosms (*p* > 0.05). DON concentrations were not

significantly different in surrounding waters and in the mesocoms up to day 17 (p > 0.05). From this day, even though a significant decrease (p < 0.05) in DON concentrations was also observed in surrounding waters, the resulting concentrations were significantly higher outside than in the mesocosms (p < 0.05).

- ⁵ The DOP dynamics was similar to the DON dynamicdynamics: during P1, DOP concentrations were on average $0.14\pm0.03\,\mu\text{mol}\,\text{L}^{-1}$ and remained stable up to day 14, day 16 and day 17 for M1, M2 and M3, respectively (Fig. 5c4C). After these days, DOP concentrations significantly decreased (p < 0.05) of 0.10, 0.07 and 0.06 in M1, M2 and M3, respectively, which also occurred in a lower extent in reached $0.09\pm0.01\,\mu\text{mol}\,\text{L}^{-1}$ on average. DOP
- ¹⁰ concentrations also decreased in surrounding waters from day 18. <u>18</u>, but to a lesser extent than in the mesocosms.

3.6 Export fluxes and their coupling with PP primary production and N₂ fixation

Before the $\frac{\text{DIP}}{\text{PO}_{4}^{3-}}$ fertilization (day 4), the exported fluxes were not significantly different (p > 0.05) between the three mesocosms (104±35, 6.5±1.7 and 0.35±0.09 in nmol L⁻¹ d⁻¹ on average for POC_{export}, PON_{export} and POP_{export}, respectively) (Fig. 65). The daily ex-15 ported particulate matter remained relatively stable during P1 averaging at 164 ± 141 , 10.2 ± 7.1 and 0.9 ± 1.3 nmol L⁻¹ d⁻¹ for POC_{export}, PON_{export} and POP_{export}respectively. Fluxes in M1 were higher than in M2 and M3 during this period, which explains the large standard deviation of the averaged fluxes., respectively. During P2, the daily export fluxes increased continuously in all the mesocosms to a higher extent than during P1 for C and 20 N, peaking at 1197 ± 257 and 106.1 ± 20.1 nmol L⁻¹ d⁻¹ for C and N, respectively. The daily POP_{export} was much more stochastic ranging ranged from 1.0 to 18.4 nmol L⁻¹ d⁻¹. The e-ratio is -ratio, defined as the amount of exported carbon (POC export) relative to the fixed carbon (PP). The box-plot of the daily e ratio data collected in the three mesocosms during P1 and P2 is shown in Fig. 7. The e ratio, was significantly higher (p < 0.05, 25 n = 57) during P2 (39.7 ± 4.9%) than during P1 (23.9 ± 20.2%). Integrated N₂ fixation rates integrated rate over P1 was $0.10 \pm 0.02 \,\mu$ mol L⁻¹ on average for all the mesocosms and $0.10 \pm 0.04 \,\mu$ mol L⁻¹ (Fig. 86). The resulting change of on the TN pool in the mesocosms remained between -0.01 and $0.01 \,\mu$ mol L⁻¹ and was not significantly different from 0 (sign test, p > 0.05). During P2, Time integrated N₂ fixation rate over P2 was $0.25 \pm 0.06 \,\mu$ mol L⁻¹ and the PON_{export} was $0.45 \pm 0.04 \,\mu$ mol L⁻¹ (Fig. 86). The resulting change of the TN_{calc} pool in the mesocosms remained not significantly different from 0 (sign test, p > 0.05) up to day 18 but deviated negatively from 0 (sign test, p < 0.05) on from day 19 up to day 23 (Fig. 8).

18 but deviated negatively from 0 (sign test, p < 0.05) on from day 19 up to day 23 (Fig. 8). Thus, even though PON concentrations increased during 6). At the end of P2, the decrease of the TN_{calc} pool decreased of was $0.20 \pm 0.04 \,\mu$ mol L⁻¹.

4 Discussion

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¹⁰ 4.1 DIP Phosphate availability sustained N₂ fixation and in turn new PPprimary production

A widely accepted concept in marine biogeochemistry states that N₂ fixation is locally limited by micronutrient limited by micronutrients availability such as iron (Fe) (Dekaezemacker et al., 2013; Monteiro et al., 2011; Moore et al., 2009; Shiozaki et al., 2014), temperature (Breitbarth et al., 2007; Fu et al., 2014; Mulholland and Bernhardt, 2005)or light availability, light (Garcia et al., 2013; Kranz et al., 2010). However, fixation is ultimately driven by the DIP excess and ultimately by PO_4^{3-} availability (Dore et al., 2008; Karl et al., 1997; Moutin et al., 2008), resulting from N loss through denitrification or anammox (Weber and Deutsch, 2014). During our mesocosms experimentwhere no

P supply could be provided, the artificial DIP fertilization was performed in order to alleviate a limitation that. During this experiment, PO₄³⁻ fertilization was carried out as P limitation would have prevented diazotrophs growth. This fertilization associated with the relatively high micronutrients micronutrient concentrations (e.g. Fe) concentrations previously reported in the lagoon (Ambatsian et al., 1997), the seawater temperature above high

²⁵ seawater temperature (>25 °C) and the maintenance of the water mass in the lighted upper layerare considered as ideal photic layer, provided optimal conditions for diazotrophs

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growth. As a result expected, diverse diazotroph phylotypes developed extensively in the mesocosms at abundances comprised between with abundances of 1.10^5 and to 5.10^5 nifH copies L⁻¹ (Turk et al., 2015). Furthermore, N₂ fixation rates measured in the mesocosms were high (18.5 ± 13.1 in averagenmol N L⁻¹ d⁻¹) compared to those measured in surrounding waters (9.2 ± 4.8 in averagenmol N L⁻¹ d⁻¹) (Fig. 1b1B) and are among the highest previously reported in the literature (Luo et al., 2012).

The contribution of N₂ fixation to PP (10.8 ± 5.0 %on average) in the mesocosms and in surrounding waters (5.7 ± 2.0 %on average) was in the upper range reported in of previous studies in the Pacific Ocean (Raimbault and Garcia, 2008; Shiozaki et al., 2013) and in the Mediterranean Sea (Bonnet et al., 2011; Ridame et al., 2013). Before the DIP Prior to PO₄³⁻ fertilization, NO₃⁻ concentrations were below 40 nM. As <0.04 µmol L⁻¹. As there was no external supply of NO₃⁻ can be provided to the system and considering N inputs from atmospheric deposition as negligible in the studied region (Jickells et al., 2005), we estimated that, the potential consumption of the initial NO₃⁻ initially present in the mesoto cosms would have represented represented <11.5 % of the integrated N₂ fixation rates

- ¹⁵ cosms would have represented represented <11.5% of the integrated N₂ fixation rates over P1 and P2 (0.35±0.08 µmol L⁻¹) (Fig. 76). Thus, N₂ fixation supplied nearly all the new production (sensu Dugdale and Goering, 1967) during P1 and P2. The mesocosms isolates only a part of the water column, which hinders any comparison with open ocean field studies. However, this experimentshows during the experiment. These results indicate
 ²⁰ that in a N depleted system, diazotrophs can provide enough new N to sustain high PP
- (exceeding at times 2 μ mol C L⁻¹ d⁻¹) and biomass (up to 1.42 μ g L⁻¹ of Chl *a*), as long as P-PO₄³⁻ does not limit N₂ fixation.

4.2 The relative efficiency of different diazotrophs to export particulate matter

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Only few studies have focused on the direct coupling between N₂ fixation and particulate export (Dore et al., 2008; Karl et al., 2012; White et al., 2012). To our knowledge, the only study that has compared comparing the export efficiency of different diazotrophs reports that in the North East Pacific DDAs blooms could contribute up to 44 % of the direct export in the North East Pacific, while UCYN (Group A) and *Trichodesmium* spp. could account for only 0

to 10% of the particulate export at the base of the euphotic layer export (White et al., 2012). The scarcity of data is explained by methodological issues that are often encountered when using sediment traps to collect the sinking material due to methodological issues associated with the use of sediment traps in the open ocean. They include due to (1) the patchy geographical distribution of N₂ fixers that is are not necessarily collected in by the sediment traps, and (2) the temporal lag between production of fixers and the production and the

export which is difficult to assess (Nodder and Waite, 2001). The mesocosms approach used here mesocosm approach was designed to overcome these limitations. Thus, even if the experimental limitations. The shallow depth of the traps (\sim 15 m) and the absence of

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- NO₃⁻ supply trough normally supplied via the nitracline prevent any comparison with open ocean studies, it allows. Nevertheless, the mesocosm approach enables a comparison of the export efficiency under contrasted ecological situations, here between P1, contrasting ecological situations. In this case, the period dominated by DDAs, and P2, (P1) is compared with the period dominated by UCYN-C (P2).
- ¹⁵ During P1, the biomass was stable in the mesocosms and the amount of recently fixed recently fixed N₂ was equal to the amount of N exported exported PON, suggesting a tight coupling between the two processes (Fig. 86). It has been shown that large aggregates of the diatom *Rhizosolenia* spp., that represented more than 80% of the diatoms involved in representing the majority of DDAs during P1 (Turk et al., 2015), may directly can sink
- ²⁰ at high rates (Villareal et al., 1996). *Rhizosolenia* spp. may also have been ingested by zooplankton producing fast sinking fecal pellets. However, the large size of this diatom limits its grazing by mesozooplankton (Perissinotto, 1992) and this species is generally poorly represented in copepod guts (Haq, 1967; Marshall and Orr, 1966). This suggests that during this experiment, the recently fixed N₂ by DDAs remained within the symbiotic association
- and was quickly exported in the settling particles. This is in good agreement agrees with Karl et al. (2012) findings who showed that DDAs support the pulses of particle export export pulses regularly observed in late summer in the tropical North Pacific Ocean.

The situation was contrasted During P2, the increase in PON concentrations during P2 : the total amount of exported PON exceeded the (Fig. 3B) suggests that part of the freshly

produced biomass remained in the water column. The accumulation of PON probably favored remineralization processes, explaining the increase in NH_{4}^{+} concentrations. This may have enhanced the transfer of the recently fixed N_2 to the non diazotrophic plankton as demonstrated by Bonnet et al. (2015a) and explains the development of picocyanobacteria (Leblanc et al., 2015). Additionally, the total amount of N provided by N₂ fixation did not 5 account for all the exported PON during P2 (Fig. -8)suggesting 6), implying that an additional N source . Furthermore, a part of the PON produced was accumulated in the water column increasing the N source deficit. The equilibrium can only be reached considering the DON pool as a source for the PON production (including PON_{export}), which is consistent with the significant DON concentrations decrease played a significant role in promoting the 10 export. The only alternative N source is DON which, indeed, exhibited a significant decrease in concentration of 0.9 ± 0.7 in the mesocoms μ mol L⁻¹ in the mesocosms during P2 (see next section section 4.3 for further discussion on the DON consumption). Assuming a DON use of 0.9 and neglecting the that DON and sources N₂ fixation are the only possible sources of N in the mesocosms, we calculated that the DON pool supported 15 78a DON use of 0.9 μ mol L⁻¹ would have supported ~ 78% of the PON production during P2and thus, and potentially fueled the PON export in to the same extent. This is in agreement with Torres-Valdes et al. (2009) support the idea of a lateral transport of DON and DOP from the productive area toward the equatorial gyres of the Atlantic ocean were it can be assimilated and support up to 40% of the vertical PON and POP export. 20 Recently Letscher et al. and Letscher et al. (2013) confirmed the importance of the lateral advection of DON toward the Atlantic and Pacific gyres. The DON concentrations measured during this experiment (4.5-5.5) are in the lower range of the reported values for the surface Pacific Ocean (Bronk, 2002). Together with low inorganic nutrient availability, low DON concentrations indicate that the studied water mass was characteristic of the open 25 ocean. Thus, even in oligotrophic areas, DON pool appears as a dynamic contributor to who showed that DON pool is a dynamic contributor of the N cycle able to fuel particulate export.

Using our data only, it is difficult to assess the contribution of fixation and DON use to PON_{export} during P2supports up to 40% of the vertical PON export in the oligotrophic gyres of the Pacific and Atlantic Oceans. A quantification of the diazotrophs in the sediment traps, performed on day 19, shows that $\sim 10\%$ of the UCYN-C biomass in the mesocosms was directly exported daily in the sediment traps and may explain exported 5 this day, explaining \sim 7% of the total PON export (Bonnet et al., 2015a). The latter paper suggests that this direct export may have been promoted by the formation of high sinking aggregates (~1size) of UCYN-C observed in the mesocosms. The formation of these aggregates may have been enhanced by the production of extracellular polysaccharides (EPS) by UCYN, as described in culture by Sohm et al. (2011). However, the direct export 10 of UCYN-C biomass does not completely explain the biomass exported. This suggests that a part of the Thus, the recently fixed N₂ was accumulated in the water column and may have been recycled through the bacterial loop or directly released under dissolved inorganic or organic N compounds and potentially transferred to non-diazotrophic plankton cells during the experiment that in turn, were exported in the sediment traps. On day 17, 15 Bonnet et al. (2015a) demonstrated that at $\sim 20\%$ of the recently fixed by the diazotrophs dominated by by UCYN-C was transferred toward can directly be exported, but is probably more efficiently transferred to non diazotrophic plankton i. e. in picoplankton and bacteria. This may explain the accumulation of picocyanobacteria observed during P2 in all the mesocosms (Leblanc et al., 2015). Zooplankton may have also played a role in the transfer 20 of the recently fixed N by grazing on UCYN and may have potentially been guickly exported trough the production of rapid sinking fecal pellet (Hunt et al., 2015). through minerealization processes.

The contrasted pattern contrast between P1 and P2 also appeared regarding is also observed using the *e*-ratio. Indeed the -ratio. The production driven by UCYN-C was more efficient to promote POC_{export} in promoting POC export than the production driven by DDAs(Fig. 7)... During P1, it is probable that the recently fixed C by DDAs remained within the symbiotic association and sunk together with the recently fixed N₂ constituting a direct and net C export. During P2, the higher efficiency of C export indicates that the use of

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the DON pool strongly suggests that the DON ultimately fueled PP whichin turn enhanced the POC_{export}, in turn, increased POC export. Additionally, a preferential recycling of N compared to C when UCYN-C dominated, may have allowed more carbon an enhanced N remineralization may have enabled more C to be fixed per unit of fixed N₂ -leading to a higher *e*-ratio. A proportionally higher N remineralization following high PP and N₂ fixation rates is supported by similar findings in the western North Pacific warm pool (Shiozaki et al., 2013).

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4.3 Who were the DON consumers? The unexpected high dissolved organic matter consumption

- ¹⁰ The use of dissolved organic compounds and their implications on PP in the <u>open</u> ocean has long been <u>suggested demonstrated</u> (Antia et al., 1991). The use of DOP by plankton communities in the oligotrophic ocean has been observed in the North Pacific Ocean (Bjorkman and Karl, 2003) or and in the Atlantic Ocean (Lomas et al., 2010; Mather et al., 2008) and generally occurs under <u>DIP scarcity. In the present PO₄³⁻ limitation. In this study</u>,
- ¹⁵ the decrease of in DOP concentrations during P2 occurred when T_{DIP} - T_{PO4} reached the lowest values levels, confirming the ability of the planktonic community to significantly use the use DOP under low $\frac{\text{DIP}}{\text{PO4}}$ availability. More surprisingly, the significant and rapid decrease of in DON concentrations (0.9 ± 0.7 in average in the mesocomsFig. 4) observed during the development of UCYN-C (P2) in the mesocosms was associated with a rapid
- increase in PP (Fig. 1e1C), biomass (Figs. 3 and 42 and 3) and bacterial production (BP) (Van Wambeke et al., 2015), suggests a considerable consumption of this organic N pool high consumption of DON directly or indirectly by primary producers. In the open ocean, DON is considered as mainly refractory. Nevertheless, it is now recognized that a fraction of the DON is labile and can directly support phytoplankton growth, while a semi labile frac-
- tion can be mineralized by bacterioplankton (Antia et al., 1991; Bronk, 2002; Bronk et al., 2007). The large analytical uncertainties on DON determination (~0.5) imply that small but relevant changes in concentrations in the open ocean are difficult to trace. Furthermore, DON is composed of a heterogenic assemblage of organic compounds such as urea, amino

acid, nucleic acids, humic and fulvic substances. Due to this heterogeneous composition, experiments using isotopically labeled compounds are difficult to conduct to measure the DON uptake. As a result, it is challenging to identify the producers and consumers of the DON pool. In the present In this study, we propose three hypotheses that could explain in

- part or totally the observed DON concentrations decrease the observed decrease in DON concentrations during P2: (1) a i) bacterial mineralization of DON triggered by high PP, (2) a ii) direct uptake of DON by primary producers including UCYN-C and (3) an iii) abiotic photo-degradation of DON into NH₄⁺.
- (i) The increase in PP driven by high N_2 fixation rates during P2 has led to an increase ef-in bacterial production (Van Wambeke et al., 2015). The significant negative correlation 10 between BP and DON concentrations (Spearman rank correlation, r = -0.35; p < 0.001) would argue for a indicate significant consumption of the DON by bacterial mineralization. Based on BP data and assuming a bacterial growth efficiency between 10 and 30% (del Giorgio and Cole, 1998) and a C : N ratio of 6.6 (Fukuda Diazotrophs are known to over-fix C relative to N (Mulholland et al., 1998), we calculated that bacterial respiration 15 could have led to a DON consumption of 0.2 to 0.7 during P2, which could explain the DON removal reported here. Therefore, during P2, the increasing PP stimulated by the new N provided by UCYN-C may have stimulated in turn bacterial activity. Figuring that (1) the diazotrophs are known to over-fix C relative to N and (2) the 2007), which may explain why the POC: PON ratio was well above the Redfield one ratio during the ex-20 periment, the N demand. The resulting N deficit for bacterial mineralization may have been ultimately found in found in the labile or semi labile DON pool. This hypothesis is supported by Van Wambeke et al. (2015) who showed that BP was limited by N availability . Furthermore, the increasing concentrations suggest an increasing regenerated production potentially mediated by bacterial ammonification that support the hypothesis 25 of in the mesocosms during the experiment. Based on BP data and assuming a bacterial
- mediated DON consumption growth efficiency of 10 to 30% (del Giorgio and Cole, 1998) and a C:N ratio of 6.6 in bacteria cells (Fukuda et al., 1998), bacterial respiration would

have led to a DON consumption of 0.2 to $0.7 \,\mu$ mol L⁻¹ during P2, supporting at least part of the DON removal of ~0.9 μ mol L⁻¹ reported here.

(ii) An alternative explanation for the DON decrease in concentration is a decrease in DON concentrations is direct consumption by primary producers. Cyanobacteria are known

- to be able to use DON compounds such as urea (Collier et al., 2009; Painter et al., 2008) in such to such an extent that DON has been reported to be the main source of N fueling cyanobacterial blooms in coastal waters (Glibert et al., 2004). In our system, the DON decrease that occurred during the large development of the diazotrophic cyanobacteria The DON decrease occurring during the development of UCYN-C, whose abundances
- reached 5.10⁵ nifH copies L⁻¹ (Turk et al., 2015), questions the ability of the latter their ability to use DON to meet their N requirements. Direct uptake of glutamate and amino acid acids (constitutive components of the DON pool) has been reported in natural and laboratory populations of *Trichodesmium* spp. (Mulholland and Capone, 1999; Mulholland et al., 1999). Furthermore, similarly to our results, Berman (1997) observed that large de-
- ¹⁵ crease in DON concentration were followed by concentrations were observed after blooms of the diazotroph *Aphanizomenon ovalisporum* in Lake Kinneret . Their (Berman, 1997). The hypothesis of a direct use of DON by the diazotroph *A. ovalisporum* was confirmed in culture experiment where *A. ovalisporum* development the development of this diazotroph was stimulated by DON additions (Berman, 1997, 1999). To our knowledge, no direct up-
- take measurement measurements of DON compounds has have been performed on UCYN. However, the ureA gene implicated in the urea assimilation has recently been identified in the cyanobacterial diazotrophic strain *Cyanothece* PCC 7822 (Bandyopadhyay et al., 2011) that is closely related to the UCYN-C cluster. This would These pieces of evidences suggest that in addition to provide new N trough N₂ fixation, the UCYN-C present during P2 would
- have been might be able to use the DON pool to meet a large part of their N requirements, which may explain the observed significant decrease in DON concentrations.

(iii) Finally, photo-degradation has been pointed out as could be a possible sink of DON in surface waters (Bronk, 2002). Only few measurements have been performed in open ocean systems, with only one A field study performed in the ultraoligotrophic eastern basin

of the Mediterranean Sea indicating indicates a production of NH_4^+ from DON of 0.2–2.9 in surface water nmol N L⁻¹ d⁻¹ in surface waters (Kitidis et al., 2006). However, even taking Taking into account the highest rates reported above, this process can only explain a small part of the DON removal observed herecannot explain more than 10% of the observed DON removal. Moreover, the DON decrease occurred only during P2 whereas photo-degradation was likely to occur would be occurring continuously over the experiment. Consequently, the

two first hypotheses are preferred

The first two hypotheses (i and ii) are more likely to explain the DON decrease and the concomitant PON increase during P2. None of these two hypotheses can be excluded even

though direct evidence proof of large uptake of DON by UCYN-C is lacking. AnyhowThus, in this study the DON use appears was directly or indirectly triggered by the UCYN-C activity.

5 Conclusions

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This study confirms that in the South West Pacific, N₂ fixation is a biogeochemically relevant process able to provide sufficient new N to drive new PP, biomass accumulation and organic matter export as long as P is not limiting. It appears that the The fate of the recently fixed N 15 is appears to be closely related to the diazotrophic community involved in N₂ fixation. Thus, aA strong coupling of N₂ fixation and PON_{export} PON export occurred when DDAs dominated the diazotrophic community suggesting their direct export. Conversely, when When the community was dominated by UCYN-C, biomass accumulation was observed together with an efficient particulate export. Moreover, aA significant decrease of in DON concen-20 trations was observed during the same period indicating a direct or indirect use of DON by UCYN-C. Thus, in addition to fuel fueling primary production, UCYN-C appears appear to be able to enhance regenerated production based both on the transfer of recently fixed N₂ toward non fixing planktonic groups and on the use of the DON pool. This use of DON exceeded the new N provided by N_2 fixation even though the N_2 fixation rates were among 25 the highest reported in literature for the global ocean. These results suggest suggest that

DON has to be considered as a dynamic pool even in LNLC area in LNLC areas as it may provide significant amounts of N and contribute significantly to the particulate export.

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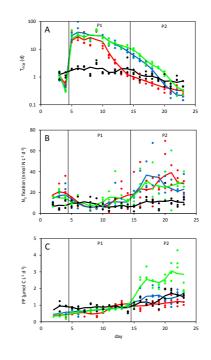


Figure 1. Temporal evolution of (a) dissolved inorganic (A) phosphate turn over time ($T_{DIP}T_{PO4}$, d), (b) (B) dinitrogen fixation (N₂ fixation) rates (nmol N L⁻¹ d⁻¹), (c) (C) primary production (PP) rates (µmol C L⁻¹ d⁻¹) in the mesocosms M1 (red), M2 (blue) and M3 (green) and in surrounding waters (black). The three dots of each color represent the measured values on the three sampled depths. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by DDAs diatoms diazotrophic associations and UCYN-Gunicellular N₂-fixing cyanobacteria (group C), respectively.

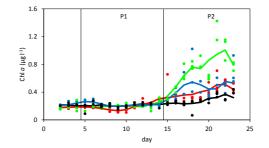


Figure 2. Boxplot Temporal evolution of the contribution of fixation to PP during P1 Chlorophyll a (dominance of DDAsChl a) and P2-concentrations (dominance of UCYN-Cµg L⁻¹) in the mesocoms (see text for details) mesocosms and surrounding waters. No significant differences were observed between The color code is identical to Fig. 1. The three dots of each color represent the measured values on the three sampled depths. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by diatoms-diazotrophic associations and unicellular N2-fixing cyanobacteria (Krustal-Wallis test, $\alpha = 0.05$ group C), respectively.

Temporal evolution of Chlorophyll a (Chl a) concentrations () in the mesocosms and surrounding waters. The color code is identical to Fig. 1. The three dots of each color represent the measured values on the three sampled depths. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by **DDAs and UCYN-C, respectively.**

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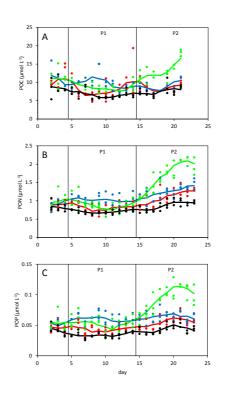


Figure 3. Temporal evolution of (a) (A) particulate organic carbon (POC), (b) (B) particulate organic nitrogen (PON), (c) (C) particulate organic phosphorus (POP) concentrations (μ mol L⁻¹) in the mesocosms and surrounding waters. The color code as is identical to Fig. 1. The three dots of each color represent the measured values on the three sampled depths. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by DDAs diatoms diazotrophic associations and UCYN-Cunicellular N₂-fixing cyanobacteria (group C), respectively.

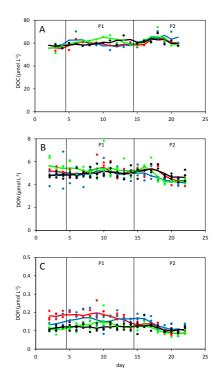


Figure 4. Temporal evolution of (a) (A) dissolved organic carbon (DOC), (b) (B) dissolved organic nitrogen (DON), (c) (C) dissolved organic phosphorus (DOP) concentrations (μ mol L⁻¹) in the meso-cosms and surrounding waters. The color code is identical to Fig. 1. The three dots of each color represent the measured values on the three sampled depths. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by DDAs diatoms-diazotrophic associations and UCYN-CuniceIlular N₂-fixing cyanobacteria (group C), respectively.

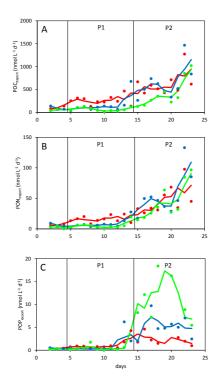
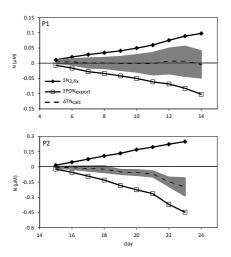


Figure 5. Temporal evolution of (a) (A) particulate organic carbon exported (POC_{export}), (b) (B) particulate organic nirtogen exported (PON_{export}) and (c) (C) particulate organic phosphorus exported (POP_{export}) fluxes (nmol L⁻¹ d⁻¹) in the mesocosms expressed in equivalent water volume. The color code is identical to Fig. 1. The solid lines are the three days running mean value. P1 and P2 denote the two phases of the experiment when the diazotrophic community was dominated by DDAs diatoms diazotrophic associations and UCYN-Cunicellular N₂-fixing cyanobacteria (group C), respectively.



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Figure 6. Boxplot Integration of the *c* ratio dinitrogen fixation rates (defined as $POC_{export}/PP\sumN_{2,fix}$) and particulate organic nitrogen exported (ΣPON_{export}) during P1 (dominated by DDAsdominance of diatoms-diazotrophic associations) and P2 (dominated by UCYN-Cdominance of unicellular N₂-fixing cyanobacteria group C) in the mesocoms mesocosms together with the calculated change in Total N content (see text for details ΔTN_{calc}) - A significant defined as the difference were observed between $\Sigma N_{2,fix}$ and ΣPON_{export} . The shaded areas represent the two phases (Krustal–Wallis test, $\alpha = 0.05$) uncertainty associated with the ΔTN_{calc} change.

Integration of fixation rates $(\Sigma N_{2,fix})$ and PON exported (ΣPON_{export}) during P1 (dominance of DDAs) and P2 (dominance of UCYN-C) in the mesocosms together with the calculated change in Total N content (ΔTN_{calc}) defined as the difference between integrated fixation and PON exported. The shaded areas represent the uncertainty associated with the ΔTN change.