Co-variation of metabolic rates and cell-size in coccolithophores

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Abstract

Coccolithophores are sensitive recorders of environmental change. The size of their coccosphere varies in the ocean along gradients of environmental conditions and provides a key for understanding the fate of this important phytoplankton group in the future ocean. But interpreting field changes in coccosphere size in terms of laboratory observations is hard, mainly because the marine signal reflects the response of multiple morphotypes to changes in a combination of environmental variables. In this paper I examine the large corpus of published laboratory experiments with coccolithophores looking for relations between environmental conditions, metabolic rates and cell size (a proxy for coccosphere size). I show that growth, photosynthesis, and to a lesser extent calcification, co-vary with cell size when pCO_2, irradiance, temperature, nitrate, phosphate and iron conditions change. With the exception of phosphate and temperature, a change from limiting to non-limiting conditions always results in an increase in cell size. An increase in phosphate or temperature (below the optimum temperature for growth) produces the opposite effect. The magnitude of the coccosphere size changes observed in the laboratory is comparable to that observed in the ocean. If the biological reasons behind the environment-metabolism-size link are understood, it will be possible to use coccosphere size changes in the modern ocean and in marine sediments to investigate the fate of coccolithophores in the future ocean. This reasoning can be extended to the size of coccoliths if, as recent experiments are starting to show, coccolith size reacts to environmental change proportionally to coccosphere size. The coccolithophore

database is strongly biased in favor of experiments with the coccolithophore *Emiliania huxleyi* (*E. huxleyi*) (82% of database entries), and more experiments with other species are needed to understand if these observations can be extended to coccolithophores in general. I introduce a simple model that simulates the growth rate and the size of cells forced by nitrate and phosphate concentrations. By considering a simple rule that allocates the energy flow from nutrient acquisition to cell structure (biomass) and cell maturity (biological complexity, eventually leading to cell division), the model is able to reproduce the co-variation of growth rate and cell size observed in laboratory experiments with *E. huxleyi* when these nutrients become limiting. These results support ongoing efforts to interpret coccosphere and coccolith size measurements in the context of climate change.

1 Introduction

Coccolithophores, the main calcifying phytoplankton group, are an important component of the oceanic carbon cycle (Broecker and Clark, 2009; Poulton et al., 2007). Through their cellular processes of photosynthesis (a CO_2 sink) and calcification (a source of CO_2), they contribute in defining the magnitude of the ocean-atmosphere CO_2 flux (Shutler et al., 2013). The calcium carbonate platelets (coccoliths) that make up their exoskeleton (coccosphere) provide ballast for dead organic matter in the photic zone, accelerating the export of carbon from the upper ocean to the sediments (Honjo et al., 2008). There is laboratory and field evidence that climate change is affecting the cellular processes and global distribution of coccolithophores, with potential consequences on the magnitude of the carbon fluxes introduced above (Gehlen, 2007; Wilson et al., 2012). For example, in laboratory cultures, the coccolithophore *E. huxleyi* shows reduced calcification-to-photosynthesis ratios when CO_2 is changed from pre-industrial levels to those predicted for the future, acidic ocean (Hoppe et al., 2011; Langer et al., 2009; Riebesell et al., 2000; Zondervan et al., 2002). In the ocean, the coccolithophore *E. huxleyi* has been expanding polewards in the past sixty years, most likely driven by rising sea surface temperatures and the fertilizing effect of increased CO_2 levels (Winter et al., 2013). Despite the great number of laboratory experiments testing the effect of multiple environmental conditions on coccolithophore physiology (Iglesias-Rodriguez et al., 2008; Langer et al., 2012; Paasche et al., 1996; Riebesell et al., 2000; Riegman et al., 2000; Rouco et al., 2013; Sett et al., 2014; Zondervan, 2007; Zondervan et al.,

2002), it is hard to link laboratory results with field observations to obtain a unified picture of how coccolithophores respond to changing environmental conditions (Poulton et al., 2014).

E. huxleyi is the most abundant, geographically distributed and studied coccolithophore (Iglesias-Rodríguez, 2002; Paasche, 2001; Winter et al., 2013). It exhibits a strong genetic diversity, with the different genotypes adapted to distinct environmental conditions (Cook et al., 2011; Iglesias-Rodríguez et al., 2006; Medlin et al., 1996) – a characteristic that explains its global distribution and ecological success in the modern ocean (Read et al., 2013). *E. huxleyi* morphotypes, which differ for their coccosphere size, as well as shape, size and degree of calcification of coccoliths (Young and Henriksen, 2003), correspond to at least three genetically distinct genotypes (Cook et al., 2011; Schroeder et al., 2005). The geographical distribution of *E. huxleyi* morphotypes in the ocean is controlled by environmental conditions (Beaufort et al., 2008; Beaufort et al., 2011; Cubillos et al., 2007; Henderiks et al., 2012; Poulton et al., 2011; Schiebel et al., 2011; Smith et al., 2012; Young et al., 2014). But the physiological role of key factors such as pCO_2 is controversial, with a study showing that high pCO_2 favors morphotypes with smaller and lighter coccoliths, (Beaufort et al., 2011), and other studies showing the opposite (Grelaud et al., 2009; Iglesias-Rodriguez et al., 2008; Smith et al., 2012). Next to pCO_2, there is growing evidence that irradiance, nutrients and temperature also play a role in controlling morphotype biogeography (Berger et al., 2014; Henderiks et al., 2012; Smith et al., 2012). Despite the need for a better understanding, it is clear that the geographical distribution of *E. huxleyi* morphotypes carries precious information on how this key coccolithophore species will react to climate change.

But there is another, more subtle effect of climate change on coccolithophores: as living conditions evolve, cell-size and coccosphere-size adapt, due uniquely to a physiological response to environmental change. At the cellular scale, laboratory experiments with *E. huxleyi* show that pCO_2, irradiance, temperature and nutrient concentrations affect not only rates of photosynthesis and calcification, but also cell and coccosphere size, without inducing a change in morphotype (Bach et al., 2011; De Bodt et al., 2010; Iglesias-Rodriguez et al., 2008; Muller et al., 2008; Müller et al., 2012; Oviedo et al., 2014; Rouco et al., 2013). Culture conditions also affect the size and mass of coccoliths (Bach et al., 2012; Bollmann and Herrle, 2007; Müller et al., 2012; Paasche et al., 1996; Satoh et al., 2008; Young and Westbroek, 1991). Coccolith size (length, volume) and weight are used as proxies for coccolithophore calcification because they are related to the total mass of calcite in the cell (Beaufort et al., 2011) (although multiple layers of coccoliths around cells may complicate

this simple picture). The size of coccoliths is positively related to that of coccospheres in laboratory experiments (Müller et al., 2012), in the ocean (Beaufort et al., 2008) and in marine sediments (Henderiks, 2008), and the mass of coccoliths is positively related to that of coccospheres in the ocean (Beaufort et al., 2011). These observations suggest that the physiological sensitivity of coccosphere and coccolith size to environmental conditions carries supplementary information on the reaction of *E. huxleyi* to climate change.

In the ocean, attempts are made to disentangle the effect of multiple environmental parameters on the size and mass of *E. huxleyi* coccospheres and coccoliths (Beaufort et al., 2008; Beaufort et al., 2011; Cubillos et al., 2007; Hagino et al., 2005; Henderiks et al., 2012; Meier et al., 2014; Poulton et al., 2011; Young et al., 2014). This is a complicated task. Primarily, as explained above, because changes in cell size are partly ecological in origin and some automatic measuring procedures do not distinguish between the different morphotypes (Beaufort et al., 2008; Beaufort et al., 2011; Meier et al., 2014). Second, because environmental parameters co-vary in the field, making it hard to interpret size changes observed in the ocean in terms of those recorded in the laboratory. Nevertheless, a recent study based on scanning electron microscope observations suggests that the coccosphere size of *E. huxleyi* within a population of a given morphotype varies considerably and is likely under physiological control (Henderiks et al., 2012). Also the size of coccoliths of a given morphotype varies in the modern ocean (Hagino et al., 2005; Henderiks et al., 2012; Poulton et al., 2011) as well as the recent geological past (Berger et al., 2014; Horigome et al., 2014), and is likely to be under the control of parameters other than pCO_2 (Horigome et al., 2014; Young et al., 2014). To take advantage of the physiological and environmental information carried by coccosphere and coccolith size, two steps need to be taken: first, the effect of single environmental parameters on coccosphere and coccolith size has to be systematically observed in the laboratory and, second, an understanding of the biological reasons behind cell-size changes needs to be developed.

In this paper I explore the available laboratory data of coccolithophore metabolic rates and cell-size. The metabolic rates considered are the growth rate (in units of day^{-1}), the rate of photosynthesis (in units of pg_C $cell^{-1}$ day^{-1}) and the rate of calcification (in units of pg_C $cell^{-1}$ day^{-1}). First, I investigate how coccolithophore metabolic rates scale with cell-size in five species of coccolithophores, and how this scaling compares to that of other phytoplankton groups. Second, I discuss how metabolic rates and coccosphere size of a given coccolithophore species are affected by changes in environmental culture conditions. The

laboratory changes in *E. huxleyi* coccosphere-size are compared to coccosphere size changes observed in the modern ocean across gradients of environmental change. Finally, I propose a simple model that explains why metabolic rates and cell-size co-vary, with the hope that a few basic principles may be used in the future to extract environmental and metabolic information from coccosphere and coccolith measurements obtained in the field. This paper is based on a database of published results of culture experiments with coccolithophores - the next section introduces this database.

2 A database of coccolithophore metabolism and cell size

The database (Table 1, appendix A1) is composed of data collected in 369 separate culture experiments with 28 strains belonging to five species of coccolithophores (*E. huxleyi*, *Gephyrocapsa oceanica*, *Calcidiscus leptoporus*, *Syracosphaera pulchra* and *Coccolithus braarudii* (formerly known as *Coccolithus pelagicus*)). These studies were carried out in batch reactors or chemostats, in a wide range of culture conditions, including variable irradiance, light cycle, temperature, nutrient concentration (NO_3, PO_4 and Fe) calcium and inorganic carbon concentrations (pCO_2, DIC, total alkalinity). The salinity and the concentration of magnesium are similar to that of seawater. The database reports measured values of growth rate μ, in units of day^{-1}, the organic (POC) and inorganic (PIC) carbon quota, in units of pgC $cell^{-1}$, and the cell-specific rates of photosynthesis (RPh) and calcification (RCa), in units of pgC $cell^{-1}$ day^{-1}. These quantities are interrelated according to the following expressions:

$$RPh = \mu \times POC \quad (1)$$

and

$$RCa = \mu \times PIC \quad (2)$$

Equations 1 and 2 were used to complete the database when only two out of three of growth rate, carbon content and cell-specific metabolic rates are presented in a given literature source. When possible, the DIC system data has been converted to the total pH scale so that pCO_2 can be compared across the dataset. The database includes 120 measurements of coccosphere size carried out with coulter counters, flow cytometers and optical and scanning electron (SEM) microscopes.

158 Some consideration of growth rate measurements in conditions of nutrient limitation is
159 necessary. In nutrient-limited batch cultures, the growth rate decreases in time as nutrients are
160 depleted, so that determining growth rates via cell counts yields erroneous results (Langer et
161 al., 2013). Reliable growth rates in conditions of nutrient limitation can be obtained in
162 chemostats, where the growth rate is controlled by setting the dilution rate of the medium and
163 the cell population is continuously renovated (Langer et al., 2013). An alternative are semi-
164 continuous cultures where cells are periodically harvested and inoculated into new medium,
165 allowing relatively constant growth conditions (LaRoche et al., 2010). When considering
166 nutrient limitation, I thus chosen to use only data produced in chemostat and semi-continuous
167 culture experiments.

168 ## 2.1 Normalized growth rates

169 The light cycle varies from experiment to experiment, ranging from continuous light to a 12-
170 12h light-dark cycle. In order to compare the growth rates from experiments with different
171 light/dark cycles, the data needs to be normalized with respect to the duration of the light
172 period. Since photosynthesis is restricted to the light period, growth rates (μ, in day^{-1}) have
173 been normalized to the length of the light period. This is done applying the following
174 relationship (Rost et al., 2002):

175 $$\mu_i = \frac{\mu \times (L + D)}{L - D \times r} \quad (3)$$

176 where μ_i (in day^{-1}) is the normalized, instantaneous growth rate, μ (in day^{-1}) is the growth rate
177 measured via cell counts, L and D are the length (in hours) of the light and dark periods and r,
178 the factor which accounts for the respiratory loss of carbon during the dark period, is set to
179 0.15 (Laws and Bannister, 1980). Thus, the instantaneous growth rate μ_i, in units of day^{-1}, is
180 the growth rate normalized to a light period of 24 hrs.

181 ## 2.2 Normalized cell carbon quotas

182 The organic carbon quota (POC) is positively related to cell volume. To compare POC across
183 the database, a large bias introduced by the sampling strategy needs to be considered.
184 Specifically, in experiments with a light/dark cycle, POC increases during the day as small
185 cells formed during nighttime division assimilate carbon and increase in size (Linschooten et
186 al., 1991; Muller et al., 2008; Vanbleijswijk et al., 1994; Zondervan et al., 2002). Typically,
187 sampling for POC measurements is carried out at different times during the light period in

different experiments. This introduces variability in the POC data that is not related to the experimental growth conditions. When the time of sampling in the light cycle is reported, POC data have been normalized with respect to the time of sampling using the following equation (the derivation of this equation is given in appendix A1):

$$POC(t) = \frac{L \cdot POC(S_T)}{L + S_T} \cdot \left(1 + \frac{t}{L}\right) \quad (4)$$

where L is the length (in hours) of the light period, S_T is the sampling time in hours after the beginning of the light period, $POC(S_T)$ is the POC measured in the experiment at time S_T and t is the time at which the corrected POC value is calculated.

For experiments with a light/dark cycle where the sampling time is reported, I imposed t = L/2 in equation 4 to estimate the POC in the middle of the light phase. When the time of sampling is not reported, equation 4 was used to estimate a minimum and a maximum POC in the middle of the light phase assuming that the reported POC value was measured at the end and at the beginning of the light phase, respectively. This procedure was applied also to PIC values because inorganic carbon ($CaCO_3$) production takes place nearly exclusively during the light phase in coccolithophores (Muller et al., 2008) and PIC shows an evolution similar to POC during the light period (Zondervan et al., 2002). In experiments with continuous light the cell-cycle is desynchronized such that the average cell diameter remains constant if environmental conditions do not change (Muller et al., 2008; Müller et al., 2012). Thus, the POC measurements were not corrected in these experiments. Interestingly, fossil coccolithophores represent an integrated sample over the whole light:dark cycle and thus should be more comparable to laboratory samples from desynchronizes cultures – something to keep in mind as the amount of morphological data of coccolithophores from marine sediments is growing (Beaufort et al., 2011; Grelaud et al., 2009).

2.3 Normalized cell-specific rates of photosynthesis and calcification

The normalized growth rates and normalized cell carbon quota are used to calculate normalized, cell-specific rates of photosynthesis (RPh_i, in pgC $cell^{-1}$ day^{-1}) and calcification (RCa_i, in pgC $cell^{-1}$ day^{-1}):

$$RPh_i = \mu_i \times POC_C \quad (5)$$

$$RCa_i = \mu_I \times PIC_C \quad (6)$$

where the subscript C indicates that the carbon quota refers to the value in the middle of the light phase (calculated imposing t = L/2 in equation 4) and the subscript *i* indicates that the metabolic rates are normalized with respect to the light period (equation 3). Thus, RPh_i and RCa_i are the metabolic rates normalized to a light period of 24 hrs. When the time at which sampling occurred during the light period is not known, minimum and a maximum cell-specific rates of photosynthesis and calcification are calculated assuming that the reported POC and PIC values were measured at the end and at the beginning of the light phase, respectively.

2.4 Estimating cell and coccosphere size from carbon quota

Coccosphere size data is reported only in a third of the experiments included in the dataset (of which more than 80% of measurements are for *E. huxleyi*), while no cell-size measurements are included in the database. To take advantage of the full set of metabolic measurements available, cell-size and coccosphere size were estimated from the particulate organic (POC) and inorganic (PIC) carbon content per cell with the following expression (the full derivation is given in appendix A2):

$$V_{Sphere} = \frac{1.8 \times POC}{d_{POM}} \cdot \left(1 + \frac{f_{CY}}{1 - f_{CY}}\right) + \frac{100}{12} \cdot \frac{PIC}{d_{CaCO3}} \cdot \left(1 + \frac{f_{Sh}}{1 - f_{Sh}}\right) \quad (7)$$

where V_{Sphere} is the volume of the coccosphere (Fig. 1), the volume of the cell and shield are equal to the first and second term on the right in equation 7, respectively, d_{POM} (in g cm^{-3}) is the density of organic matter, d_{CaCO3} (equal to 2.7 g cm^{-3}) is the density of $CaCO_3$ and f_{CY} and f_{Sh} are the volume fraction occupied by water in the cell and shield, respectively. Equation 7 assumes that cell volume scales linearly with cellular carbon content. This assumption is reasonable for coccolithophores due to the absence of large vacuoles (Paasche, 1967).

I used equation 7 to calculate the diameter of the cell and the coccosphere for all the experiments in the database for which POC and PIC data are available (Fig. 2). The unknowns in this equation are d_{POM}, f_{CY} and f_{SH}. First, d_{POM} was set to 1.5 g cm^{-3}, which lies at the center of the range of values proposed by Walsby and Raynolds (1980) (1.3 – 1.7 g cm^{-3}). Then f_{CY} and f_{SH} were varied so that the resulting diameter of the great majority of *E. huxleyi* spheres fell in the range 3 – 7.5 μm, which corresponds approximately to the range reported in culture experiments (Fig. 2) and to that measured microscopically in surface waters off the coast of the Benguela upwelling system (Henderiks et al., 2012). The chosen

values of f_{CY} (0.79) and f_{SH} (0.66) results in a difference between the diameter of the coccosphere and that of the cell of about 1.5 μm for most of *E. huxleyi* the cells (values significantly smaller or larger than 1.5 μm are interpreted in appendix A2). This value, observed in cultures of *E. huxleyi* (Henderiks, pers. comm.), corresponds roughly to twice the thickness of one layer of coccoliths (and thus to one layer of coccoliths in the shield around one cell). This is consistent with the laboratory observation that in most calcifying *E. huxleyi* cells regulate their calcification rates/division rates in order to maintain at least a complete layer of coccoliths, even in growth-limited conditions (Paasche, 1999). With these parameter settings, the resulting density of the naked *E. huxleyi* cell is 0.18 pgC μm^{-3}, which is comparable to that of carbon in protist plankton of similar size determined by Menden-Deuer and Lessard (2000). The cell diameter obtained with this procedure is compared with that obtained applying an existing relation between POC and cell volume (Montagnes et al., 1994) in appendix A1.

The calculated coccosphere diameter of *E. huxleyi* is compared to the measured coccosphere diameter for the experiments in the database where POC, PIC and cell size data are reported (Fig. 2b). Although a clear positive relation between measured and calculated coccosphere size exists, the calculated diameters are always larger than the measured diameters (except for two experiments in Kaffes et al. (2010)). The large majority of coccosphere size measurements in the database were carried out with Coulter counters, which often do not include the coccolith shield in the size measurement (Iglesias-Rodriguez et al., 2008; Oviedo et al., 2014; van Rijssel and Gieskes, 2002). Consistently, the Coulter counter diameter for *E. huxleyi* corresponds to the cell diameter calculated with equation 7 (Fig. 2b). Another source for the observed discrepancy is the fact that in some experiments cells are fixed chemically prior to size measurements, a treatment that induces cell shrinkage. Appendix A1 discusses the discrepancy between measured and calculated coccosphere size more in detail. With these consideration in mind, the choice made above of constraining equation 7 with the range of *E. huxleyi* coccosphere diameters measured with the microscope (Henderiks, 2008) appears to be the safest.

In figure 2c, the same parameterization of equation 7 is applied to the POC and PIC data available for the other coccolithophore species. A comparison with published coccosphere size data for some of these species suggests that approach is reasonable. Most of the calculated coccosphere diameters for *Coccolithus braarudii*, for example, fall in the range

17-24 μm, which is slightly more extended that that reported by Henderiks (unpublished data reported graphically in figure 7 of Henderiks (2008)) (18 – 22 μm). The corresponding shield thickness for *Coccolithus braarudii* falls in two groups (4.5 μm and 7.5 μm) suggesting the presence of more than one layer of coccoliths per cell in some cases. Similar to *E. huxleyi*, the coccosphere diameter measured with Coulter counters is always smaller than the calculated diameter (Fig. 2d). However, the discrepancy is small for these larger-sized species. Significantly, the coccosphere diameter of *Calcidiscus leptoporus* measured with SEM without prior fixing of cells by Langer et al. (2006) coincides with the calculated coccosphere diameter using equation 7 (Fig. 2d). When discussing cell and coccosphere size from experiments in the database I use equation 7 throughout the rest of this manuscript, regardless if size measurements are reported in the literature sources or not.

3 The allometric scaling of coccolithophore metabolism

In this section the coccolithophore database is used to investigate relationships between cell volume and metabolic rates across different taxa under comparable growth conditions (allometric relations). The differences in metabolic rates we will deal with are largely due to differences in characteristic cell size across different taxa. Allometric relationships for coccolithophores will be compared with similar relations for other phytoplankton groups compiled by Marañón (2008). The Marañón (2008) dataset includes cell volume and metabolic rate data measured in the field for a vast array of unicellular photosynthetic organisms spanning 9 orders of magnitude in size, from photosynthetic cyanobacteria (volume = 0.1 μm^3) to large diatoms (volume = 10^8 μm^3) and including dinoflagellates and haptophytes. The Marañón (2008) dataset reports rate measurements that mostly reflect *in situ* optimum growth conditions; thus, in this section, I focus on experiments in the coccolithophore database that were carried out in optimum conditions (Table 2). The assumptions made in comparing metabolic rates from the coccolithophore database with those measured in the field by Marañón (2008) are detailed in appendix A2.

Figures 3a and 3b compare the allometric relations of photosynthesis and growth for coccolithophores with those established by Marañón (2008) for phytoplankton. Figures 3c and 3d show the allometric relations for photosynthesis and calcification in coccolithophores,

highlighting the position of the five different coccolithophore species considered. Linear regressions through the optimum coccolithophore dataset yield the following equations:

$$Log_{10}(RPh_i) = 0.89 \times Log_{10}(Volume) - 0.66 \quad (8)$$

$$Log_{10}(\mu_i) = -0.11 \times Log_{10}(Volume) + 0.1 \quad (9)$$

$$Log_{10}(RCa_i) = 1.02 \times Log_{10}(Volume) - 1.02 \quad (10)$$

The slope of the photosynthesis (0.89) and growth rate (-0.11) regressions for coccolithophores is very similar to that of the Marañón (2008) dataset (0.91 and -0.09, respectively) and comparable to the slope of the regression through the calcification rate data (1.02). Furthermore, the different coccolithophore species occupy a position on the volume-photosynthesis diagram that is dictated by their cell size (figure 3c). These plots show that, for coccolithophores grown in optimum conditions, 1) photosynthesis in coccolithophores – including five different species spanning nearly three orders of magnitude in cell size - scales to cell volume in a comparable way as it does in other phytoplankton, 2) the size dependence of growth rates is very small for coccolithophores, 3) calcification in optimum growth conditions scales isometrically with cell volume.

The finding of a near-isometric scaling of coccolithophore growth in laboratory experiments has implications for the scaling of phytoplankton population abundance with body size in the ocean. In the ocean, including a variety of contrasting marine environments, phytoplankton population abundance scales with body size with an exponent equal to -¾: in other words, small cells are more abundant than large cells (Cermeño et al., 2006). Reviews of laboratory culture experiments with phytoplankton growth under optimal growth conditions suggest that cell-specific photosynthesis rates scale with cell volume with an exponent of ¾ (Lopez-Urrutia et al., 2006; Niklas and Enquist, 2001), possibly a consequence of the generic properties of transportation networks inside the organisms (Banavar et al., 2002; West et al., 1997). According to this scaling rule, growth rates scale with cell-size with an exponent of -¼, implying that large cells grow more slowly than small cells and offering an explanation for the size scaling of population abundance with cell size observed in the field (Cermeño et al., 2006).

However, the laboratory -¼ scaling of growth rate to cell size has been challenged by the observation that the same scaling in natural communities of phytoplankton is nearly isometric

(Huete-Ortega et al., 2012; Marañón, 2008; Marañón et al., 2007) (i.e. a slope in eq. 9 nearly equal to 0 and no effect of cell size on growth rate). The size exponent for different phytoplankton groups varies, with diatoms having a higher exponent (0.01) that of dinoflagellates (-0.11) (Marañón, 2008) and whole community exponents varying from -0.01 (Marañón, 2008) to 0.16 (Huete-Ortega et al., 2012). An isometric scaling of growth rates to cell volume has recently been also observed in laboratory experiments with 22 species of phytoplankton ranging from 0.1 to 10^6 μm^3 in volume (López-Sandoval et al., 2014; Marañón et al., 2013). In this context the coccolithophore dataset is particularly relevant because it fills in the gap of sizes between 10^0 and 10^3 μm^3 that is underrepresented in Marañóns' (2008) dataset. Furthermore, it confirms that a scaling exponent significantly smaller than -¼ occurs in laboratory conditions, in addition to field situations, suggesting that cell-size is not an important factor in determining the size distribution of coccolithophore populations. Taken together, the near-isometric scaling of growth rate with cell size observed in the ocean by Marañón (2008) and in the laboratory (López-Sandoval et al., 2014; Marañón et al., 2013) suggest that the -¾ scaling of phytoplankton population abundance with cell size is not due uniquely to an effect of cell size on growth rates.

We are left with a contradiction that needs to be explained: whereas in some cases growth rates in the laboratory scale with cell-size with an exponent of -¼ (Lopez-Urrutia, 2006; Niklas and Enquist, 2001), this is not the case in the ocean (Huete-Ortega et al., 2012; Marañón, 2008; Marañón et al., 2007) and in some laboratory experiments (López-Sandoval et al., 2014; Marañón et al., 2013). With regard to laboratory experiments, López-Sandoval et al. (2014) point out that this difference could be in part due to the fact that older compilations of experimental data do not include cells smaller than 100 μm^3. In the ocean, the larger phytoplankton (e.g. diatoms) have the ability to move vertically in the water column and adapt to variable nutrient and light conditions (Mitrovic et al., 2005; Stolte et al., 1994). This confers an advantage over small phytoplankton cells and provides a possible explanation for the near-isometric scaling of natural phytoplankton communities (Marañón, 2008). In laboratory experiments, where environmental parameters are typically constant, such extrinsic factors cannot be at play and some intrinsic, cellular-level, property of coccolithophore cells must exist that allows larger coccolithophores to overcome the geometrical constraints imposed by cell size on resource acquisition (Raven, 1998). Some coccolithophores posses carbon concentrating mechanisms (CCMs) that enable cells to take up HCO_3^-, as well as CO_2, for photosynthesis, and interconvert HCO_3^- to CO_2 internally via

the carbonic anhydrate enzyme (Reinfelder, 2011; Rost et al., 2003). There is evidence from the carbon stable isotope composition of coccolithophore calcite that large coccolithophore species employ CCMs more efficiently than small species when CO_2 is scarce (Bolton and Stoll, 2013). This differential use of CCMs in large and small coccolithophore species offers a plausible (even if not excusive) explanation of why coccolithophore growth rate scales nearly isometrically with cell size in laboratory experiments.

4 Environmental controls on cell size and metabolic rates in coccolithophores

In this section I investigate how changes in environmental conditions affect cell size and metabolic rates in coccolithophores. The changes we will deal with are produced by the physiological response of a given taxon to environmental change; I will discuss the effects of six environmental variables: pCO_2, irradiance, temperature, nitrate, phosphate and iron. Next to the optimum group of experiments introduced in section 3, I highlight light-limited, nitrate-limited, phosphate-limited and iron-limited experiments. The set of conditions defining these groups is detailed in table 2. Most of the data (82 % of database entries) comes from cultures of *E. huxleyi*, the more thoroughly studied coccolithophore; experiments with the other four coccolithophores in the database have essentially tested the effect of pCO_2 conditions on growth, photosynthesis and calcification.

Within the optimum group of experiments, the position of the high-CO_2 subgroup largely corresponds to that of the low pCO_2 group (Fig. 4). A considerable number of data points collected in sub-optimal growth conditions, however, fall below the regression line through the optimal data. The scatter is greater for *E. huxleyi* reflecting the fact that a much smaller number of environmental conditions have been tried out for the other species. For all rates of growth, photosynthesis and calcification, the light-limited experiments consistently plot below the optimum experiments (Fig. 4). The position of the nutrient-limited experiments below the optimum experiments is even more evident (Fig. 4): light-limited and nutrient limited cells have smaller metabolic rates than cells of comparable size grown in optimum conditions. For experiments where the sampling time during the light period is unknown, the range of values for the photosynthetic rate (error bars) is large and an overlap with optimum group of experiment exists. However, only 5 out of 30 experiments in the light-limited group and 9 out of 31 nutrient-limited experiments have unknown sampling times, such that the

position of the experiments run in limiting conditions under the optimum group of experiments is significant.

The plots of volume against metabolic rates introduced above do not take advantage of the whole potential of the experimental dataset. This is because part of the variability in metabolic rates observed is due to differences in the pre-culture conditions and, very likely, to biological variability, rather than to the experimental conditions that the experiments are designed to test. A better picture is obtained if *changes* in cell volume are plotted against *changes* in metabolic rates. I have explored the database for sets of experiments were only one experimental condition is changed at a time, so that the change in volume and metabolic rates can be calculated by subtraction and plotted. In this way different sets of experiments can be compared on the same plot (this procedure is explained in detail in appendix A3). The plots show the changes in metabolic rates and cell size induced by an increase in pCO_2, an increase in irradiance starting from light-limited conditions, an increase in temperature and a decrease in nitrate, phosphate or iron starting from nutrient-replete conditions (figures 5 and 6). These changes correspond to the evolution of the living conditions that phytoplankton are experiencing (warming, acidification) or are planned to experience (ocean stratification leading to increased irradiance and oligotrophy) in the coming centuries (Behrenfeld et al., 2006; Bopp, 2005; Bopp et al., 2001). Tables 3 and 4 summarize the changes in cell and coccosphere diameter and volume induced by changes in experimental culture conditions. They highlight an important fact: changes in pCO_2 produce only limited variations in coccosphere size compared to variations in other parameters such as irradiance, temperature and nutrients.

4.1 pCO_2 increase

For the low-pCO_2 group of experiments run in optimum conditions (Fig. 5), an increase in pCO_2 leads to an increase in cell size and little change in the growth rate. The rate of photosynthesis increases with pCO_2, indicating that *E. huxleyi* is carbon-limited in this range of pCO_2. The biomass-specific calcification rate decreases in the great majority of the experiments, while the change in the rate of calcification can be positive or negative. Interestingly, the response of photosynthesis and calcification differ not only in sign, but also in homogeneity: while the change in photosynthetic rate defines a clear trend in the volume-metabolism space, the change in calcification rate is poorly correlated with the change in cell volume. This is not surprising given that the rate of photosynthesis increases both due to the

fertilizing effect of CO_2 (physiological effect) and due to the increase in cell size (geometric effect), while the rate of calcification is positively affected by the increase in cell-size (geometric effect) but inhibited physiologically by acidification (Raven and Crawfurd, 2012). This complex reaction of calcification to changes in the DIC system has been elegantly captured in a recent model equation developed by (Bach et al., 2015). Furthermore, the response of calcification to a rise in pCO_2 is modulated by the growth temperature (which varies between experiments) and can be negative or positive (Sett et al., 2014). Finally, the response of calcification in *E. huxleyi* to an increase in pCO_2 is known to be strain-specific, with a large span of responses possible (Langer et al., 2006). In all experiments but 3, the ratio of calcification to photosynthesis decreases following the pCO_2 increase. Overall, the changes observed for the low pCO_2 group of optimum experiments occur also in the high-pCO_2 group of experiments (albeit with a larger scatter) and in the experiments run in conditions of light limitation (Fig. 5). The few experiments available where pCO_2 is varied in conditions of nitrate limitation seem to point to a similar behavior (see appendix A3), as do the data available for other coccolithophore species (Fig. 5).

4.2 Irradiance increase in light-limited conditions

Increasing irradiance from irradiance-limited conditions leads to a large increase in cell-size, growth rate and rate of photosynthesis (Fig. 6). In the majority of experiments also the biomass-specific and cell-specific rate of calcification increase with irradiance. The effects on the calcification-to-photosynthesis ratio are large, with most experiments showing an increase in calcification compared to photosynthesis. These effects are observed both in low pCO_2 and in high-pCO_2 conditions; they can be understood considering that both photosynthesis and calcification are light-dependent, energy-requiring processes (Brownlee et al., 1995; Raven and Crawfurd, 2012). Interestingly, there is a smaller dispersion in the calcification rate data compared to the set of experiments where pCO_2 is increased (figure 5). This is because both the geometric and physiological consequences of an irradiance increase concur in increasing the rate of calcification (geometric and physiological effects have contrasting influence on calcification rate for a pCO_2 rise). The experiments showing a negative response of the PIC/POC ratio with increased irradiance are from Rokitta and Rost (2012) and Feng et al. (2008) where high light intensities where used (300 and 400 $\mu mol\ m^{-2}\ s^{-1}$, respectively), possibly inducing photoinhibition of calcification (Feng et al., 2008).

4.3 Temperature

Both in optimum and in light-limited conditions, an increase in temperature leads to an increase in the growth, photosynthesis and calcification rate and a decrease in cell size in the majority of the experiments considered (the scatter is considerable). This is consistent with the observation that *E. huxleyi* has highest growth rate at temperatures 5-10°C higher than the maxima observed at the isolation sites (Sett et al., 2014) – a pattern that seems to apply in general to phytoplankton from polar and temperate regions (Atkinson et al., 2003; Thomas et al., 2012). This trend has also been described in an long-term experiment during which *E. huxleyi* was allowed to adapt for 1 year (roughly 460 asexual generations) to high temperatures (Schlüter et al., 2014).

4.4 NO_3, PO_4 and Fe limitation

Under nitrogen limitation all cell-specific and biomass-specific metabolic rates decrease and cells become smaller (Fig. 6). The same effect on metabolic rates is observed under phosphorous limitation, but the effect on cell size is opposite (Fig. 6). The contrasting effect of nitrogen and phosphorous limitation on cell size depends on the different role of these nutrient in the cell cycle (Muller et al., 2008). In the G1 (assimilation) phase of the cell-cycle, nitrogen consumption by *E. huxleyi* cells is high because cells are synthesizing and accumulating biomass (Muller et al., 2008). Therefore, nitrogen depletion decreases assimilation rates and leads to smaller cells. The result is not dissimilar from what happens during light limitation. Phosphorous consumption, instead, is highest during the S and G2 + M phases, due to synthesis of nucleic acids and membrane phospholipids immediately before cell division (Geider and La Roche, 2002; Muller et al., 2008). Thus, phosphorous limitation is though to arrest the cells in the G1 (assimilation) phase of the cell cycle, increasing the length of this phase and leading to an increase in the cell-size. Thus, in phosphorous limited cells, cell-size does not increase because the assimilation rate increases but because the assimilation period is longer. The change in the ratio of photosynthesis to calcification is generally positive. In the only set of experiments considering iron limitation (Schulz et al., 2007), cell-size co-varies with growth and photosynthesis rates in a similar way as in nitrate-limited experiments (Fig. 6). Iron is a key component of carbon concentrating mechanisms (CCMs) that increase the rate of import of inorganic carbon (CO_2 and HCO_3^-) for photosynthesis, and of chlorophyll; thus, under iron-limiting conditions, the decrease in metabolic rates is produced by carbon-limitation (Schulz et al., 2007). The concomitant decrease in cell-size is consistent with the size shifts observed in the experiments where pCO_2 is varied (Fig. 5).

It should be noted that the coccolithophore database is strongly biased in favor of experiments with the coccolithophore *E. huxleyi* (82% of database entries), and more experiments with other species are needed to understand if the above relations between environment, cell size and metabolic rates can be extended to coccolithophores in general. Furthermore, the experiments included in the coccolithophore dataset are designed to quantify the instantaneous (meaning a few generations) response of coccolithophores to changing growth conditions. In longer-term experiments, lasting several hundred generations, (Lohbeck et al., 2012; Schlüter et al., 2014) *E. huxleyi* has been observed to adapt to elevated temperatures and pCO_2 conditions simulating future ocean conditions. This implies that the trends of metabolic rates and cell-size with changing environmental conditions that are described in this section will be modulated by evolutionary adaptation, adding further complexity to the interpretation of past and future response of coccolithophores to climate change. The results of these experiments show, however, that the long-term response of growth rate and cell size to increased temperature and increased pCO_2 are qualitatively comparable: cells adapted to high temperature decrease their cell-size while cells adapted to high pCO_2 increase their cell size (Schlüter et al., 2014).

5 The size of *E. huxleyi* in the ocean: is there hope of detecting a physiological signal ?

In the previous section we saw that a change in laboratory culture conditions nearly always results in a change of cell and coccosphere-size of coccolithophores. In this section the changes in coccosphere size observed in laboratory experiments are compared to those observed in the ocean. I will consider in some detail the BIOSOPE transect that crosses the south pacific gyre from the Marquises islands to the Peru upwelling zone (Beaufort et al., 2008). Figure 7a shows the BIOSOPE transect superimposed on a surface ocean chrolophyll concentraion map obtained from satellite observations. Figure 7b is a vertical transect in the upper 300 m of the ocean showing the variability of the diameter of coccospheres belonging to the order Isochrysidales. The order Isochrysidales is composed of the genera *Emiliania*, *Geophyrocapsa* and *Crenalithus*. These genera cannot be distinguished from one another by the automated SYRACO system used to measure coccosphere diameter and generate figure 7b. In addition to SYRACO, the BIOSOPE samples were examined with a Scanning Electron

Microscope and a light microscope which process less samples than SYRACO but are able to distinguish the different Isochrysidales genera.

Along the BIOSOPE transect the diameter and volume of Isochrysidales coccospheres measured with SYRACO varies considerably (from 4.5 to 8 μm figure 7b). Scanning Electron Microscope and light microscope observations show that between 140°W and 130°W, where coccospheres are largest (mostly > 6 μm in diameter), *Gephyrocapsa oceanica* dominates the Isochrysidales assemblage (Beaufort et al., 2008). *Gephyrocapsa oceanica* has a characteristic cell size which is slightly larger than *E. huxleyi* (figure 3). In the Peru upwelling zone (75°W) where SYRACO detects large coccospheres (mostly > 6 μm in diameter), microscope observations show that *E. huxleyi* morphotype R, which is characteristically large ("over-calcified"), is abundant. Clearly, changes in coccosphere size along the BIOSOPE transect are partly ecological in origin – an observation that can be exported to the global ocean (Beaufort et al., 2011).

But how do the cell-size changes observed along the BIOSOPE transect compare with those observed in laboratory experiments? Whereas in the ocean changes in cell size can be due to both ecological and physiological effects, in the laboratory only physiological effects are expected. The histograms of figure 8a and 8b show the coccosphere diameter and volume of cultured *E. huxleyi* cells and of the *Isochrysidales* coccolithophores in the BIOSOPE transect. Laboratory and field measurements compare well. The red horizontal bar graphs of figures 8a and 8b are the changes in coccosphere diameter and coccosphere volume observed in laboratory experiments for given variations in culture conditions (see also Tables 3 and 4). The comparison of histograms and bar charts shows that the variability of cell-size in laboratory cultures is similar to that observed in the BIOSOPE transect. In figure 8c, the range of environmental conditions imposed in laboratory cultures are compared with the range of environmental conditions along the BIOSOPE transect. Large differences in the total range exist only for phosphate and iron, with concentrations in limited experiments being much lower than those measured in the BIOSOPE transect. Even discarding the phosphate and iron limitation experiments, it is clear that changes in environmental conditions along the BIOSOPE transect are very likely to be an important driver of coccosphere size variability: *physiological* effects concur with *ecological* effects in determining coccolithophore cell-size variability.

Further evidence for a physiological control on coccosphere size in the ocean comes from the Benguela coastal upwelling system, where the size of the well-calcified *E. huxleyi* morphotype A* (determined by SEM observations) changes considerably with environmental conditions (Henderiks et al., 2012). The largest coccospheres occurred at the depth of the deep chlorophyll maximum (DCM) – where growth conditions can be assumed to have been more favorable than in the overlying and underlying water masses - whereas coccospheres above and below the DCM were significantly smaller. This is consistent with the laboratory observations (section 4) that environmental conditions which result in large growth rates (and thus lead to large populations in the field) are also those that give rise to large cells (phosphate concentrations in the Benguela upwelling system were much larger than those which induce an increase in cell size in culture experiments).

Another, even less explored (but equally promising), avenue of research is that of the physiological control of environmental conditions on the size of coccoliths. Field measurements of coccolith size are more abundant that measurements of coccosphere size. However, as for coccospheres, it is difficult to disentangle physiological from ecological effects. Clearly, different morphotypes occupy distinct ecological niches characterized by different environmental conditions. For example, Cubillos et al. (2007) show that Type A ("overcalcified") and Type B/C morphotypes occupy distinct latitudinal zones in the southern ocean. Environmental conditions likely control the geographical distribution of different morphotypes on the east coast of Japan (Hagino et al., 2005), the Bay of Biscay (Smith et al., 2012), the Patagonian shelf (Poulton et al., 2011) and the South East Pacific (Beaufort et al., 2008). Clearly, part of the variability in coccolith size distribution in the global ocean is due to ecological effects (Beaufort et al., 2011).

There is laboratory and field evidence, however, that coccolith size is affected by environmental conditions also via physiological effects. Coccosphere and coccolith size are related (Henderiks, 2008). In laboratory cultures subject to varying pCO_2 and nitrate levels, coccolith volume (which is related to coccolith length) is positively correlated to both cell and coccosphere size (Müller et al., 2012), leading to the counterintuitive co-existence of large coccoliths and acidic conditions. An increase in the size of coccoliths with increasing pCO_2 has also been observed in nutrient replete, nitrogen-limited and phosphate-limited experiments (Rouco et al., 2013). In the Benguela coastal upwelling system a significant positive correlation has been found between the coccosphere diameter and coccolith length of *E. huxleyi* morphotype A* (Henderiks et al., 2012). Since the Benguela correlation is based

on SEM observaions, it is likely that ecological effects can be excluded and that the physiological effects that produce larger coccospheres also result in the production of larger coccoliths. More in general, when the coccolith size from individual morphotypes is measured along gradients of environmental conditions, it results that coccolith size varies significantly; for example off the eastern coast of Japan (Hagino et al., 2005) and along the Patagonian shelf (Poulton et al., 2011). More experiments and field observations are needed to understand how other environmental parameters (e.g. temperature, irradiance and nutrient availability) affect coccolith size, and to what extent laboratory observations can be exported to the ocean. The available information suggests, however, that the environment controls coccolith size via a physiological effect and that there could be as much hidden information in the size of coccoliths as there is in the size of coccospheres - in the next section I propose a way to extract this information from the modern ocean and sedimentary record.

6 A theoretical basis for interpreting the co-variation of metabolic rates and cell size

We saw that metabolic rates and cell-size co-vary in coccolithophores subject to changes in laboratory environmental conditions (section 4) and that the changes in coccosphere size observed in the laboratory are comparable in magnitude to those observed in the field along gradients of environmental change (section 5). If the cellular processes that give rise to this co-variation are understood, there is hope that coccosphere-size measurements from the field will yield information on the metabolic status of cells in the modern ocean and, possibly, on past environmental conditions. In this section I introduce a simple model that provides a theoretical basis for understanding how cellular metabolism - forced by environmental conditions - controls cell-size, giving rise to the correlations described in section 4.

The mean size of dividing cells is the result of two factors: the rate of nutrient assimilation into biomass and the length of the generation time (the time between two successive cell divisions) - long generation times and large rates of nutrient assimilation give rise to large cells, and vice versa. The changes in cell size observed in the previous section can be interpreted within this simple scheme. The central concept I use – that of separation of *structure* (biomass) from *maturity* (biological complexity, eventually leading to cell division) - is taken from the Dynamic Energy Budget (DEB) Theory (Kooijman, 2010). The model presented here is much simplified compared to existing DEB models of phytoplankton cells

(Lorena et al., 2010; Muller et al., 2011; Muller and Nisbet, 2014). However, it considers the minimum number of concepts that are necessary to explain the co-variance of metabolic rates and cell-size we are dealing with. The most important simplifications I introduce are discussed in appendix A4; the mathematical notation in this section follows that of (Kooijman, 2010).

Consider a spherical growing cell assimilating NO_3 and PO_4 (CO_2 is considered to be non-limiting). The assimilation rate of nutrients, $\dot{J}_i$ (in μmol cell^{-1} day^{-1}), is proportional to the surface of the cell (Figure 9):

$$\dot{J}_i = S \cdot j_{i\max} \cdot \frac{[i]}{[i] + K_i} \tag{11}$$

where the subscript i represents either NO_3 or PO_4, $j_{i\max}$ (in μmol μm^{-2} day^{-1}) is the surface-specific maximum nutrient uptake rate, S (in μmol μm^{-2}) is cell surface, K_i (mol litre^{-1}) is a Monod constant for nutrient uptake and $[i]$ (in mol litre^{-1}) is the nutrient concentration. Both the cell surface and the rate of nutrient assimilation are time dependent because the model simulates a growing cell. Values of $j_{i\max}$ were set equal to 4×10^{-9} and values of K_i were set equal to 0.2 μmol litre^{-1} and 2 nmol litre^{-1} for NO_3 and PO_4, respectively, which is in the range of values determined for *E. huxleyi* (Riegman et al., 2000).

Assimilated nutrients are used to undertake two fundamental tasks (figure 9): 1) increase the cellular biomass via production of *structure* and 2) increase the *maturity* of the organism. In DEB theory the *structure* (quantified in moles of carbon per cell) contributes to the biomass of the organism (and thus cell volume) and is composed of organic compounds that have a long residence time in the cell. *Maturity* (quantified in Joules per cell) has the formal status of information and is a measure of the complexity of the organism (Kooijman, 2010). Fundamental biological events in the lifespan of an organism, such as cell division, take place at a threshold level of *maturity*. Assimilated N and P both contribute to structure and maturity via the fluxes $\dot{J}_{G_i}$ and $\dot{J}_{MAT_i}$ such that mass is conserved:

$$\dot{J}_{G_i} = \kappa \cdot \dot{J}_i \tag{12}$$

and

$$\dot{J}_{MAT_i} = (1-\kappa)\cdot \dot{J}_i \quad (13)$$

where κ, which takes a value from 0 to 1, is the portion of the nutrient uptake flux which is dedicated to growth, and $\dot{J}_{G_i}$ and $\dot{J}_{MAT_i}$ (in μmol_i cell^{-1} day^{-1}) are the fluxes dedicated to growth and maturity, respectively. Dimensionless parameter κ was set equal to 0.5 both for NO_3 and for PO_4.

The growth fluxes generated from nutrient uptake, $\dot{J}_{G_i}$, are sent to a synthesizing unit (SU) for growth where biomass is synthesized at a rate $\dot{J}_G$ (in mol_C cell^{-1} day^{-1}) :

$$\dot{J}_G = 10^{-6}\cdot CN_{BIO}\cdot \left[\sum_{i=N,P}\left(\frac{\dot{J}_{G_i}}{y_{G_i}}\right)^{-1} - \left(\sum_{i=N,P}\frac{\dot{J}_{G_i}}{y_{G_i}}\right)^{-1}\right]^{-1} \quad (14)$$

where CN_{BIO} is the Redfield C/N ratio (equal to 106/16), necessary to transform the growth rate from units of mol_N cell^{-1} day^{-1} to mol_C cell^{-1} day^{-1}, and parameters y_{G_i} are the yield of nutrient flux *i* to the structure. The maturation fluxes generated from nutrient uptake, $\dot{J}_{MAT_i}$, are sent to another SU which tracks the build up of maturity in the cell with a rate $\dot{p}_R$ (in Joules cell^{-1} day^{-1}):

$$\dot{p}_R = 10^{-6}\cdot CN_{BIO}\cdot \mu_{MAT}\cdot \left[\sum_{i=N,P}\left(\frac{\dot{J}_{MAT_i}}{y_{MAT_i}}\right)^{-1} - \left(\sum_{i=N,P}\frac{\dot{J}_{MAT_i}}{y_{MAT_i}}\right)^{-1}\right]^{-1} \quad (15)$$

where μ_{MAT} (in Joules mol_C^{-1}) is the chemical potential of maturity (set equal to 10^5 joules mol_C^{-1}) and the parameters y_{MAT_i} are the yield of nutrient flux *i* to maturity. In this simple model, I set the yield parameters in equations x and y such that NO_3 contributes primarily to the structure ($y_{G_{NO3}} = 1$; $y_{G_{PO4}} = 0.6$) and PO_4 to maturity ($y_{MAT_{NO3}} = 0.6$; $y_{MAT_{PO4}} = 1$).

The build-up of structure M_V (in $mol_C\ cell^{-1}$) and maturity E_H (in Joules $cell^{-1}$) is tracked by the following differential equations:

$$\frac{dM_V}{dt} = \dot{J}_G \tag{16}$$

$$\frac{dE_H}{dt} = \dot{p}_R \tag{17}$$

in DEB theory volume, V (in μm^3) , is obtained from the structural mass (the maturation flux is considered to dissipate in the environment and thus does not contribute to cell volume):

$$V = \frac{M_V \cdot \mu_V}{[E_G]} \tag{18}$$

where μ_V (Joules mol_C^{-1}) is the chemical potential of the structure and $[E_G]$ (in Joules μm^{-3}) represents the volume-specific growth costs. In equation 18, the ratio of the chemical potential of the structure to the volume-specific growth costs can be obtained from the density of carbon in biomass, C_{BIO}, which is equal to 0.18 $pg_C\ \mu m^{-3}$ for *E. huxleyi* (section 3):

$$\frac{\mu_V}{[E_G]} = 10^{12} \cdot \frac{m_C}{C_{BIO}} \tag{19}$$

where m_C (=12) is the molecular weight of carbon and the factor 10^{12} is needed to convert pg_C to g_C. Thus, substituting the right hand side of equation 19 in equation 18, the model calculates cell volume as follows:

$$V = \frac{M_V \cdot m_C \cdot 10^{12}}{C_{BIO}} \tag{20}$$

At any time, the instantaneous growth rate μ_{INST} (in day^{-1}) can be calculated as the ratio of the carbon uptake rate and the cellular carbon quota:

$$\mu_{INST} = \frac{\dot{J}_G}{M_V} \tag{21}$$

Figure 10 shows how maturity, cell volume and the instantaneous growth rate (calculated with equation 21) evolve during a typical model run in non-limiting conditions. The model is run starting with initial cell size equal to 10 μm^3. As nutrients are taken up, they contribute to the structure. Biomass and cell size increase. As the cell grows maturity accumulates, until the threshold maturity for cell division is attained (dashed red line in figure 10a). The cell divides and a new cell cycle starts. After cell division the cell volume of the daughter cell is equal to half the volume of the parent cell, while the maturity buffer is emptied and the maturity of the daughter cell is set to zero. The instantaneous growth rate (equation 21) decreases during growth within a given cell cycle consistent with the fact that the growth rate is proportional to the surface/volume ratio of cells. After a few cell cycles model variables (structure, maturity, volume etc.) repeat themselves from one cycle to another: the model has reached steady state. A full model run which brings the system into steady state lasts about ten cell cycles. The final steady state condition is independent of the initial cell size and depends only on nutrient concentrations and biological model parameters. The generation time is graphically visible as the horizontal distance between two successive division events. At steady state, the growth rate μ (in day^{-1}) can be approximated from the generation time G_T (in days) (Powell, 1956):

$$\mu = \frac{\log 2}{G_T} \qquad (22)$$

The growth rate calculated from the generation time (equation 22) is numerically equivalent to the average value of the instantaneous growth rate calculated with equation 21 (red dashed line in figure 10c). In the following, I will discuss average cell volumes and growth rates at steady state (dashed red lines in figures 10b and 10c).

Next, the model is used to investigate how cell-size and growth rate vary in conditions of nutrient limitation. The model is run changing NO_3 and PO_4 concentrations while keeping all the other model parameters unchanged. As explained above, the SUs were parameterized such that NO_3 contributes primarily to the structure (and to a lesser extent to maturity) and PO_4 contributes primarily to maturity (and to a lesser extent to the structure). The model was run ten thousand times with combinations of NO_3 and PO_4 concentrations included between 10^{-2} to 1 mmol $litre^{-1}$ (NO_3) and 10^{-4} and 10^{-2} mmol $litre^{-1}$ (PO_4) (Figure 11). Figure 11a shows how cell volume (blue contour lines) and growth rate (red dashed lines) depend on NO_3 and PO_4 concentrations: while NO_3 and PO_4 limitation both result in a decrease of the

731 growth rate, they have contrasting effects on cell size, with NO_3 limitation resulting in a
732 decrease size and PO_4 limitation in an increase of cell size. These trends are further displayed
733 in figures 11b to 11e: figures 11b and 11c are plots of how growth rate and cell-size vary
734 when PO_4 is kept at non-limiting levels (10^{-2} mmol litre^{-1}) and NO_3 varies. Figures 11d and
735 11e are plots of how growth rate and cell-size vary when NO_3 is kept at non-limiting levels (1
736 mmol litre^{-1}) and PO_4 varies. Figures 11c and 11e are of the same sort of those presented in
737 section 5 where changes in growth rate and cell volume induced by NO_3 and PO_4 limitation
738 are represented on log scales. The experimental data from Riegman et al. (2000) (orange
739 points: NO_3 limitation, brown points: PO_4 limitation) are included in figures 11c and 11e.

740 These simulations show that the model can reproduce trends in growth rate and cell size
741 observed in laboratory experiments when NO_3 and PO_4 become limiting (section 5). In the
742 following I discuss the features of the model that produce these trends. The growth rate is
743 directly related to the generation time (equation 22). The generation time depends on the rate
744 at which the maturity buffer is filled. Since both NO_3 and PO_4 contribute to the maturation
745 flux, limitation in NO_3 and PO_4 both result in an increase in the generation time and a
746 decrease in the growth rate. The link between growth rate and maturation flux is obvious if
747 the maturation power is plotted as a function of NO_3 and PO_4 concentrations: the isolines of
748 growth rate (figure 11a) follow those of the maturation power (figure 11b). Controls on cell
749 size are slightly more complicated. Cell size is affected both by the rate of biomass increase
750 and by the generation time. Specifically, cell size is proportional both to the rate of biomass
751 increase and to the generation time (and thus inversely proportional to the growth rate). The
752 key model quantity determining how the average cell size changes following a change in
753 nutrient concentrations is the ratio of the energy fluxes dedicated to growth and maturation:
754

755 $$\frac{\dot{p}_G}{\dot{p}_R} \qquad (23)$$

756 Figure 11b shows the value of this ratio as a function of NO_3 and PO_4 concentrations. On a
757 diagonal line along which NO_3 and PO_4 decrease by proportionally the same amount, the
758 growth/maturity ratio is constant and equal to 1 and cell volume does not change (figure 11a).
759 If NO_3 decreases more than PO_4, then growth is more affected than maturity, leading to a
760 decrease in cell size, and vice versa.

We conclude that changes in simple model quantities, which have a sound basis in biological metabolic theory, can explain the co-variance of metabolic rates and cell-size observed in laboratory experiments where nitrate and phosphate are limiting. Although the model was run with the uptake parameters of NO_3 and PO_4, the same trend of growth rate and cell size decrease with decreasing NO_3 concentrations is obtained if NO_3 is replaced by CO_2, or the Monod term for NO_3 is replaced by a Monod term for irradiance, suggesting that the simple set of rules discussed here can potentially explain the majority of the trends in metabolic rates and cell-size described in section 4. More work is needed to expand this simple physiological model to include other important features of full DEB models such as the distinction between reservoirs and structure, and to consider the interacting effect of multiple environmental changes. There is hope, however, that this effort will be rewarded by a better understanding of how environment affects the metabolic performance of coccolithophores in the modern ocean - a fundamental step in predicting how this important group of phytoplankton will be affected by climate change.

7 Conclusions

The examination of published results of coccolithophore culture experiments allows the following conclusions. The scaling of coccolithophore metabolism to cell size in optimal growth conditions is comparable to that observed in other phytoplankton groups by Marañón (2008). Larger taxa experience greater photosynthesis and calcification rates, while the growth rate is weakly dependent on cell-size. In addition cell size in *E. huxleyi* depends on environmental conditions. When only one of pCO_2, irradiance, temperature, NO_3, PO_4 or Fe is varied, cell-size and metabolic rates co-vary, defining clear trends in the 2D metabolism-cell size space. An exception is calcification under variable pCO_2 that does not show clear trends. The magnitude of coccosphere size changes observed by varying environmental culture conditions in the laboratory is comparable to the variability of *E. huxleyi* coccosphere size in the ocean. This suggests the existence of at least two controls on *E. huxleyi* cell size in the ocean: 1) the change in the relative abundance of *E. huxleyi* morphotypes with different characteristic cell sizes (*ecological* control) and 2) the change in coccosphere size induced by fluctuating environmental conditions (*physiological* control). Simple rules that regulate the partitioning of energy amongst growth and maturity explain the co-variance of cell-size and metabolic rates observed in laboratory experiments. There is hope that the Dynamic Energy

Budget Theory - which formalizes this fundamental energy partitioning - can be used to interpret coccosphere and coccolith cell-size in the past and modern ocean in terms of environmental change, providing a key for predicting the fate of coccolithophores in the future. In an evolutionary perspective, we can expect that adaptation to changing environmental conditions will modulate the observed metabolism-cell size trends, adding further complexity in the study of past and future response of coccolithophores to climate change.

Appendix

A1. The coccolithophore database

The full coccolithophore database is presented in Table A1.

Normalized cell carbon quota

Due to cell division during the dark phase, POC at the end of the light phase (P_{END}) is double the POC at the beginning of the light phase (P_0) :

$$POC_{END} = 2 \cdot POC_0 \quad \text{(A1)}$$

Thus, if POC increases linearly during the day, its evolution in time during the light phase can be expressed as follows:

$$POC(t) = POC_0 + \frac{t}{L} \cdot POC_0 \quad \text{(A2)}$$

where t is time in hours and L is the length of the light period in hours. To obtain an expression that calculates the carbon quota at any given time in the light phase, let S_T and $POC(S_T)$ be the sampling time and the corresponding POC value measured in an experiment. By substituting these values for POC(t) and t in equation A2 and rearranging we can calculate POC_0:

$$POC_0 = \frac{L \cdot POC(S_T)}{L + S_T} \quad \text{(A3)}$$

We can then substitute this expression for POC_0 in equation A1 to obtain an expression calculating the POC at any time during the light period:

$$POC(t) = \frac{L \cdot POC(S_T)}{L + S_T} \cdot \left(1 + \frac{t}{L}\right) \quad \text{(A4)}$$

Estimating cell and coccosphere size from carbon quota

The volume of the coccosphere can be thought of as the volume of the cell (V_{Cell}) plus that of the coccolith shield (V_{Shield}) (see figure 1):

$$V_{Sphere} = V_{Cell} + V_{Shield} \quad \text{(A5)}$$

Both the cell and the shield contain water. Therefore, the volume of the cell can be expressed as:

$$V_{Cell} = V_{POM} + V_{H_2OCell} \quad \text{(A6)}$$

where V_{POM} is the volume occupied by organic matter and $V_{H2OCell}$ is the volume occupied by water in the cell. Similarly, the volume of the shield can be expressed as:

$$V_{Shield} = V_{CaCO3} + V_{H_2OShield} \quad \text{(A7)}$$

where V_{CaCO3} is the volume of the $CaCO_3$ in all the coccoliths of the shield and $V_{H2OShield}$ is the volume of water contained in the shield. Defining f_{CY} and f_{SH} as the volume fractions of water in the cell and shield, respectively, the volume of the coccosphere can be expressed as:

$$V_{Sphere} = V_{POM} + \frac{f_{CY}}{1 - f_{CY}} \cdot V_{POM} + V_{CaCO3} + \frac{f_{Sh}}{1 - f_{Sh}} \cdot V_{CaCO3} \quad \text{(A8)}$$

Expressing volumes in terms of mass divided by density, the above equation becomes:

$$V_{Sphere} = \frac{M_{POM}}{d_{POM}} \cdot \left(1 + \frac{f_{CY}}{1 - f_{CY}}\right) + \frac{M_{CaCO3}}{d_{CaCO3}} \cdot \left(1 + \frac{f_{Sh}}{1 - f_{Sh}}\right) \quad \text{(A9)}$$

where M_{POM} and M_{CaCO3} are the mass of organic matter and $CaCO_3$ in the coccosphere, respectively, and d_{POM} (1.3 – 1.7 g cm^{-3}; (Walsby and Raynolds, 1980)) and d_{CaCO3} (2.7 g cm^{-3}) are the density of organic matter and $CaCO_3$, respectively. M_{POM} is related to the organic carbon per cell (POC) (Muller et al., 1986):

$$M_{POM} = 1.8 \times POC \quad \text{(A10)}$$

while the total mass of the coccoliths is related to the inorganic carbon content (PIC) per cell by:

$$M_{CaCO3} = \frac{MW_{CaCO3}}{MW_C} \cdot PIC \quad \text{(A11)}$$

where MW_C is the molecular weight of carbon (12) and MW_{CaCO3} is the molecular weight of $CaCO_3$ (100).

Substituting equations A10 and A11 in equation A9, the volume of the coccosphere can be expressed as:

$$V_{Sphere} = \frac{1.8 \times POC}{d_{POM}} \cdot \left(1 + \frac{f_{CY}}{1 - f_{CY}}\right) + \frac{100}{12} \cdot \frac{PIC}{d_{CaCO3}} \cdot \left(1 + \frac{f_{Sh}}{1 - f_{Sh}}\right) \quad \text{(A12)}$$

As explained in section 2.4, the values chosen for f_{CY} (0.79) and f_{SH} (0.66) results in a difference between the diameter of the coccosphere and that of the cell of about 1.5 μm for most of *E. huxleyi* the cells. Values significantly smaller than 1.5 are observed when cells are cultured in Ca^{2+}-poor fluids (Riegman et al., 2000; Trimborn et al., 2007), low saturation states or undersaturation wth respect to $CaCO_3$ (Bach et al., 2011; Borchard et al., 2011) or at very low light irradiances of 15 and 30 μmol photons m^2 s^{-1} in (Zondervan et al., 2002). In one case (Feng et al., 2008) small values of the coccosphere-cell diameter difference occur at high irradiances (400 μmol photons m^2 s^{-1}) and are interpreted by these authors as reflecting inhibition of calcification at high irradiance. In three of the experiments carried out by (De Bodt et al., 2010), the coccosphere-cell diameter difference is roughly double (~ 3 μm), suggesting the presence of two layers of coccoliths making up the shield that surrounds the cell.

The reconstruction of cell geometry obtained by applying equation 7 is compared to that obtained applying the equation of Montagnes et al. (1994) which relates cell carbon content (C, in pg $cell^{-1}$) to cell volume (V, in μm^3):

$$C = 0.109 \times V^{0.991} \quad \text{(A13)}$$

The diameter of *E. huxleyi* cells calculated with this formula is shown in figure 2a. The resulting cell diameter is up to 1.5 μm larger than that obtained with equation 7. I decided to use equation 7, rather than use the equation of Montagnes et al. (1994), because the equation of Montagnes et al. (1994) implies a much lower density of carbon per cell (0.1 pgC μm^{-3}) and would result in *E. huxleyi* spheres larger (up to 12 μm diameter) than those observed in culture and in the field. Similar to *E. huxleyi*, if the relation between cell volume and carbon quota per cell of Montagnes et al. (1994) (equation A13) is applied to the *Coccolithus braarudii* POC data, then the resulting coccosphere diameters for most of the coccospheres in the database (20-25 μm) are higher than those reported in Henderiks (Henderiks, 2008) (18 – 22 μm) (Fig. 2c).

Figure 2 shows that the measured coccosphere diameter is always smaller than the coccosphere diameter calculated with the geometric model (equation 7). The large majority of coccosphere size measurements in the database were carried out with Coulter counters (Table A2). It is known that cell-size measurements obtained with the coulter counter underestimates the real coccosphere size as measured by scanning electron microscope (SEM), possibly because the coulter counter does not see the coccolith shield (Oviedo et al., 2014). Iglesias-Rodriguez et al. (2008) also report coccosphere size measurements obtained with coulter counters that are significantly smaller those obtained with flow cytometry. In fact, their coulter counter measurements are very similar to the flow cytometer measurements after acidification of the sample, consistent with the idea that the coulter counter does not see the coccolith shield (Oviedo et al., 2014). Similarly, by comparing light microscope measurements with Coulter counter measurements van Rijssel and Gieskes (2002) report that coulter counter does not see the coccosphere. These considerations seem to be confirmed by the experiments of Langer et al. (2006) with *Calcidiscus leptoporus* for which the coccosphere volume determined with equation 7 coincides with the SEM-derived volume (without prior fixing of the cells).

Table A2 – Summary of methods used to determine the size of coccospheres in experiments included in the coccolithophore database

	Measurement type	Fixation reported	Notes
Müller et al. (2012)	CC[a]	no	Reports difference between non-acidified and acidified samples
Lefebvre et al. (2011)	FC[b]	no	
Borchard et al. (2011)	CC	no	
Bach et al. (2011)	CC	no	
Krug et al. (2011)	?	no	
Kaffes et al. (2010)	CC	no	
Fiorini et al. (2011)	CC	no	
De Bodt et al. (2010)	CC	yes	
Iglesias-Rodriguez et al. (2008)	CC and FC	yes, both	Coulter size << Cytometer size. Coulter = Cytometer after acidification
Langer et al. (2006)	SEM[c]	no	SEM-measured size coincides with size calculated with equation 7
Sciandra et al. (2003)	HOPC[d] and CC	no	HOPC results similar to CC results
Riegman et al. (2000)	CC	no	
van Rijssel and Gieskes (2002)	CC and LM[e]	no	LM measurement shows that coccosphere is not included in CC measurement
Arnold et al. (2013)	CC	no	

[a]Coulter counter, [b]Flow cytometer, [c]SEM, [d]Hiac optical particle counter, [e]Light microscope.

A2. Comparing the coccolithophore database with the Marañón (2008) phytoplankton database

Marañón (2008) reports metabolic rate measurements carried out in the field (via cell counts and ^{14}C-radiolabelling during incubation experiments lasting a maximum of approximately 1 day) that are as far as possible representative of *in situ* rates. Further, he chose to plot data for organisms growing in conditions of irradiance and nutrient availability that were more favorable for growth, and ran incubations at *in situ* temperature. However, nutrient limitation and sub-optimal irradiance conditions cannot be excluded for some of the measurements included in his review (Marañón, personal communication). In his compilation, the photosynthetic rates reported in units of pgC $cell^{-1}$ h^{-1} are converted in pgC $cell^{-1}$ day^{-1} by multiplying by the length of the photoperiod that may be different for different locations. When the length of the photoperiod was not available, Marañón (2008) used a photoperiod of 12h (Marañón, personal communication). In comparing the data of my dataset with the data of Marañón (2008), I divided the instantaneous growth rate (μ_i) and cell-specific metabolic rates (RPh_i and RCa_i) obtained with equations 5 and 6 by 2, obtaining rates that refer to a photoperiod of 12h. Furthermore, I concentrate on the experiments from the coccolithophore database that were carried out in culture conditions that presumably do not depart too much from those of Marañón (2008). I thus selected 172 "optimum experiments" (Table 2) carried out in conditions of high irradiance (≥ than 80 μmol photons m^2 s^{-1}), nutrient replete conditions (dissolved PO_4 and $NO_3 \geq 4$ and 64 μM, respectively) and dissolved Ca between 9 and 11.3 mM. I further subdivided these optimum experiments in a "low pCO_2" sub-group, with pCO_2 included between 150 and 550 μatm and total alkalinity between 2.1 and 2.45 mol kg^{-1}, and a "high pCO_2" sub-group, with pCO_2 included between 551 and 1311 μatm and total alkalinity between 1.9 and 2.6 mol kg^{-1}. The low pCO_2 sub-group is representative of the ranges of the monthly means values of pCO_2 and total alkalinity in the surface ocean (Lee et al., 2006; Takahashi, 2009). No distinction between low-pCO_2 and high-pCO_2 sub-groups is made in section 3 where both groups are collectively referred to as the "optimum" group. Instead, the low-pCO_2 and high pCO_2 subgroups are discussed separately and have distinct symbols in the plots of section 4 and 5.

A3. Comparison of changes in cell size with changes in metabolic rates

Method

In section 4 the changes in cell-size and metabolic rates induced by a shift of a given environmental parameter are discussed. For example, with regards to variations in pCO_2, I singled out groups of culture experiments where pCO_2 was the only environmental parameter that varied while all other culture and pre-culture conditions were reported to be constant. For every such group of experiments I recorded the difference in cell volume and metabolic rates between cells grown at a given pCO_2 and those of the experiment carried out at the lowest pCO_2 level. For example, Langer et al. (2009) carried out four experiments with *E. huxleyi* clone RCC 1238 at pCO_2 levels of 218, 412, 697 and 943 μatm. Except for the DIC parameters that co-vary with pCO_2, all other pre-culture and experimental conditions were the same. For this group of four experiments I calculated the difference in cell volume and metabolic rates between the experiments at 412, 697 and 943 μatm and the experiment at 218 μatm, obtaining the displacement in the 2D volume-metabolism space for the three experiments carried out at 412, 697 and 943 μatm.

Irradiance and temperature changes

Ideally, when comparing experiments at different irradiance and temperature levels, all other experimental parameters should be constant. In the Zondervan et al. (2002) experiments I selected couples of experiments with different irradiance and similar DIC system parameters. Similarly, I compared experiments at different temperature but similar pCO_2 conditions in the set of experiments by Sett et al. (2014). The difference in pCO_2 between different irradiance or temperature conditions was never greater than 150 μatm. Given the effect of pCO_2 on cell-size and metabolic rates (Fig. 5), some of the variability shown in the plots that show how metabolic rates co-vary with cell-size when irradiance or temperature increases (Fig. 6) will be due to variations in pCO_2.

Nutrient limitation

In Müller et al. (2012) the evolution in the 2D volume-metabolism space is obtained by comparing nitrate-replete, batch and nitrate-limited chemostat experiments with comparable DIC systems. In this way the only aquatic chemistry difference is in the dissolved nitrate concentration. In the N-limited chemostat experiments of Riegman et al. (2000), the

displacement in the 2D size-metabolism space is obtained by the difference between the highest growth rate (0.61 day^{-1}) and the nitrate-limited experiments that have lower growth rates (0.15 to 0.45 day^{-1}). In the semi-continuous cultures of Kaffes et al. (2010) the data obtained in NO_3-replete conditions (~ 280μM) was compared with that obtained at "ambient" (N. Atlantic) NO_3 concentrations (~ 10μM). Similar to the nitrate-limited experiments of Riegman et al. (2000), in the P-limited experiment of Borchard et al. (2011) and Riegman et al. (2000), the displacement in the size-metabolism space is obtained by the difference of size and metabolism at the different dilution rates (which have different dissolved P concentrations).

The shift in cell-size, growth and photosynthesis rate produced by iron limitation is deduced from the experiments of Schultz et al. (2007). These are batch experiments, so the growth rates estimated from cell counts are not reliable (Langer et al., 2013). Nevertheless, the iron-limited experiment was included because the batch experiments inform on the direction of change (positive or negative) of cell-size and metabolic rates. The net fixation rates in pmol $cell^{-1}$ hr^{-1} measured by membrane-inlet mass spectrometry by Schultz et al. (2007) (their figure 3) were converted in pgC $cell^{-1}$ day^{-1} considering 12 hours of light. The organic carbon quota per cell was then calculated from the carbon uptake rate and the growth rate (their table 1) using equation 5. The shift in metabolic rates and cell-size for iron limitation was obtained from the difference between the iron-replete and iron-limited experiments.

Increase in pCO_2 in nitrate-limited conditions

The evolution in the metabolism-volume space following an increase in pCO_2 in nutrient-limited conditions is hard to assess. Ideally, when pCO_2 is changed in the chemostat, the dilution rate should be adjusted so that the nutrient concentration remains unaltered. In this way, two nutrient-limited chemostat experiments with different pCO_2 levels could be compared. To the best of my knowledge this has not been done. However, the results of Müller et al. (2012) suggest that the growth rate changes little with pCO_2 in conditions of nitrate limitation. In these experiments, the cell-size and cell-specific photosynthesis rate of nitrate-limited cells increases with pCO_2. Nitrate is below the detection limit in all of these chemostat experiments. However, the extent to which the N/C ratio is lower in nitrate-depleted cells compared to nitrate-replete cells does not vary with pCO_2. Since decreased biomass N/C ratios are an indication of the extent of nitrate-limitation, we can conclude that the level of limitation is similar in the nitrate-limited experiments. With this in mind, the

behavior of the cells in the Müller et al. (2012) experiment is comparable to that of the cell which experience a pCO_2 increase in optimum conditions: little or no change in the growth rate, an increase in rate of photosynthesis and a decreased in calcification.

A4. Limitations of the simple DEB approach

Proper DEB models of dividing unicellular organisms are more complex than the simple version introduced in section 6. Specifically, 1) full DEB models include reserves, as well as structure and maturity, so that uptake and assimilation are decoupled and biomass stoichiometry varies with changes in nutrient availability (stoichiometry is fixed in the model used in this manuscript), 2) full DEB models consider the energy flow devoted to somatic maintenance and maturity maintenance, 3) part of the energy rejected by the growth SU is re-absorbed into the reserves in full DEB models. Notwithstanding these limitations, the simple model presented in this manuscript has the minimum characteristics of DEB models that are necessary to reproduce typical co-variations of metabolic rates and cell size.

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REFERENCES

Arnold, H. E., Kerrison, P., and Steinke, M.: Interacting effects of ocean acidification and warming on growth and DMS-production in the haptophyte coccolithophoreEmiliania huxleyi, Global Change Biology, 19, 1007-1016, 2013.

Atkinson, D., Ciotti, B. J., and Montagnes, D. J. S.: Protists decrease in size linearly with temperature: ca. 2.5% C-1, Proceedings of the Royal Society B: Biological Sciences, 270, 2605-2611, 2003.

Bach, L. T., Bauke, C., Meier, K. J. S., Riebesell, U., and Schulz, K. G.: Influence of changing carbonate chemistry on morphology and weight of coccoliths formed by Emiliania huxleyi, Biogeosciences, 9, 3449-3463, 2012.

Bach, L. T., Riebesell, U., and Georg Schulz, K.: Distinguishing between the effects of ocean acidification and ocean carbonation in the coccolithophore Emiliania huxleyi, Limnol. Oceanogr., 56, 2040-2050, 2011.

Bach, L. T., Riebesell, U., Gutowska, M. A., Federwisch, L., and Schulz, K. G.: A unifying concept of coccolithophore sensitivity to changing carbonate chemistry embedded in an ecological framework, Prog Oceanogr, 135, 125-138, 2015.

Banavar, J. R., Damuth, J., Maritan, A., and Rinaldo, A.: Supply-demand balance and metabolic scaling, Proceedings of the National Academy of Sciences, 99, 10506-10509, 2002.

Beaufort, L., Couapel, M., Buchet, N., Claustre, H., and Goyet, C.: Calcite production by coccolithophores in the south east Pacific Ocean, Biogeosciences, 5, 1101-1117, 2008.

Beaufort, L., Probert, I., de Garidel-Thoron, T., Bendif, E. M., Ruiz-Pino, D., Metzl, N., Goyet, C., Buchet, N., Coupel, P., Grelaud, M., Rost, B., Rickaby, R. E. M., and de Vargas, C.: Sensitivity of coccolithophores to carbonate chemistry and ocean acidification, Nature, 476, 80-83, 2011.

Behrenfeld, M. J., O'Malley, R. T., Siegel, D. A., McClain, C. R., Sarmiento, J. L., Feldman, G. C., Milligan, A. J., Falkowski, P. G., Letelier, R. M., and Boss, E. S.: Climate-driven trends in contemporary ocean productivity, Nature, 444, 752-755, 2006.

Berger, C., Meier, K. J. S., Kinkel, H., and Baumann, K. H.: Changes in calcification of coccoliths under stable atmospheric CO₂, Biogeosciences, 11, 929-944, 2014.

Bollmann, J. and Herrle, J. O.: Morphological variation of Emiliania huxleyi and sea surface salinity, Earth and Planetary Science Letters, 255, 273-288, 2007.

Bolton, C. T. and Stoll, H. M.: Late Miocene threshold response of marine algae to carbon dioxide limitation, Nature, 500, 558-562, 2013.

Bopp, L.: Response of diatoms distribution to global warming and potential implications: A global model study, Geophysical Research Letters, 32, 2005.

Bopp, L., Monfray, P., Aumont, O., Dufresne, J.-L., Le Treut, H., Madec, G., Terray, L., and Orr, J. C.: Potential impact of climate change on marine export production, Glob. Biogeochem. Cycle, 15, 81-99, 2001.

Borchard, C., Borges, A. V., Händel, N., and Engel, A.: Biogeochemical response of Emiliania huxleyi (PML B92/11) to elevated CO2 and temperature under phosphorous limitation: A chemostat study, J. Exp. Mar. Biol. Ecol., 2011.

Broecker, W. and Clark, E.: Ratio of coccolith CaCO(3) to foraminifera CaCO(3) in late Holocene deep sea sediments, Paleoceanography, 24, 2009.

Brownlee, C., Davis, M., Nimer, N., Dong, L. F., and Merret, M. J.: Calcification, photosynthesis and intracellular regulation of Emiliania huxleyi, Bulletin de l'Institut Océanographique de Monaco, 14, 19-35, 1995.

Cermeño, P., Marañón, E., Harbour, D., and Harris, R. P.: Invariant scaling of phytoplankton abundance and cell size in contrasting marine environments, Ecol Lett, 9, 1210-1215, 2006.

Cook, S. S., Whittock, L., Wright, S. W., and Hallegraeff, G. M.: Photosynthetic pigment and genetic differences between two southern ocean morphotypes of emiliania huxleyi (haptophyta), Journal of Phycology, 47, 615-626, 2011.

Cubillos, J. C., Wright, S. W., Nash, G., de Salas, M. F., Griffiths, B., Tilbrook, B., Poisson, A., and Hallegraeff, G. M.: Calcification morphotypes of the coccolithophorid Emiliania huxleyi in the Southern Ocean: changes in 2001 to 2006 compared to historical data, Marine Ecology Progress Series, 348, 47-54, 2007.

De Bodt, C., Van Oostende, N., Harlay, J., Sabbe, K., and Chou, L.: Individual and interacting effects of pCO2 and temperature on Emiliania huxleyi calcification: study of the calcite production, the coccolith morphology and the coccosphere size, Biogeosciences, 7, 1401-1412, 2010.

Feng, Y., Warner, M. E., Zhang, Y., Sun, J., Fu, F.-X., Rose, J. M., and Hutchins, D. A.: Interactive effects of increased pCO₂, temperature and irradiance on the marine coccolithophore Emiliania huxleyi (Prymnesiophyceae), European Journal of Phycology, 43, 87 - 98, 2008.

Fiorini, S., Middelburg, J. J., and Gattuso, J.-P.: Testing the Effects of Elevated Pco2 on Coccolithophores (Prymnesiophyceae): Comparison between Haploid and Diploid Life Stages1, Journal of Phycology, no-no, 2011.

Gehlen, M., Gangsto, R., Schneider, B., Bopp, L., Aumont, O. and Ethe, C. : The fate of pelagic CaCO3 production in a high CO2 ocean: a model study, Biogeosciences, 4, 505-519, 2007.

Geider, R. and La Roche, J.: Redfield revisited: variability of C:N:P in marine microalgae and its biochemical basis, European Journal of Phycology, 37, 1-17, 2002.

Grelaud, M., Schimmelmann, A., and Beaufort, L.: Coccolithophore response to climate and surface hydrography in Santa Barbara Basin, California, AD 1917-2004, Biogeosciences, 6, 2025-2039, 2009.

Hagino, K., Okada, H., and Matsuoka, H.: Coccolithophore assemblages and morphotypes of Emiliania huxleyi in the boundary zone between the cold Oyashio and warm Kuroshio currents off the coast of Japan, Marine Micropaleontology, 55, 19-47, 2005.

Henderiks, J.: Coccolithophore size rules - Reconstructing ancient cell geometry and cellular calcite quota from fossil coccoliths, Marine Micropaleontology, 67, 143-154, 2008.

Henderiks, J., Winter, A., Elbrächter, M., Feistel, R., der Plas, A., Nausch, G., and Barlow, R.: Environmental controls on Emiliania huxleyi morphotypes in the Benguela coastal upwelling system (SE Atlantic), Marine Ecology Progress Series, 448, 51-66, 2012.

Honjo, S., Manganini, S. J., Krishfield, R. A., and Francois, R.: Particulate organic carbon fluxes to the ocean interior and factors controlling the biological pump: A synthesis of global sediment trap programs since 1983, Prog Oceanogr, 76, 217-285, 2008.

Hoppe, C. J. M., Langer, G., and Rost, B.: Emiliania huxleyi shows identical responses to elevated pCO2 in TA and DIC manipulations, J. Exp. Mar. Biol. Ecol., 406, 54-62, 2011.

Horigome, M. T., Ziveri, P., Grelaud, M., Baumann, K. H., Marino, G., and Mortyn, P. G.: Environmental controls on the Emiliania huxleyi calcite mass, Biogeosciences, 11, 2295-2308, 2014.

Huete-Ortega, M., Cermeno, P., Calvo-Diaz, A., and Marañón, E.: Isometric size-scaling of metabolic rate and the size abundance distribution of phytoplankton, Proceedings of the Royal Society B: Biological Sciences, 279, 1815-1823, 2012.

Iglesias-Rodríguez, D. M., Schofield, O. M., Batley, J., Medlin, L. K., and Hayes, P. K.: Intraspecific genetic diversity in the marine coccolithophore emiliania huxleyi (prymnesiophyceae): the use of microsatellite analysis in marine phytoplankton population studies1, Journal of Phycology, 42, 526-536, 2006.

Iglesias-Rodríguez, M. D.: Representing key phytoplankton functional groups in ocean carbon cycle models: Coccolithophorids, Glob. Biogeochem. Cycle, 16, 2002.

Iglesias-Rodriguez, M. D., Halloran, P. R., Rickaby, R. E. M., Hall, I. R., Colmenero-Hidalgo, E., Gittins, J. R., Green, D. R. H., Tyrrell, T., Gibbs, S. J., von Dassow, P., Rehm, E., Armbrust, E. V., and Boessenkool, K. P.: Phytoplankton calcification in a high-CO2 world, Science, 320, 336-340, 2008.

Kaffes, A., Thoms, S., Trimborn, S., Rost, B., Langer, G., Richter, K.-U., Köhler, A., Norici, A., and Giordano, M.: Carbon and nitrogen fluxes in the marine coccolithophore Emiliania huxleyi grown under different nitrate concentrations, J. Exp. Mar. Biol. Ecol., 393, 1-8, 2010.
Kooijman, S. A. L. M.: Dynamic Energy Budget Theory for Metabolic Organisation, Cambridge University Press, Cambridge, 2010.
Krug, S. A., Schulz, K. G., and Riebesell, U.: Effects of changes in carbonate chemistry speciation on Coccolithus braarudii: a discussion of coccolithophorid sensitivities, Biogeosciences, 8, 771-777, 2011.
Langer, G., Geisen, M., Baumann, K. H., Klas, J., Riebesell, U., Thoms, S., and Young, J. R.: Species-specific responses of calcifying algae to changing seawater carbonate chemistry, Geochem. Geophys. Geosyst., 7, 2006.
Langer, G., Nehrke, G., Probert, I., Ly, J., and Ziveri, P.: Strain-specific responses of Emiliania huxleyi to changing seawater carbonate chemistry, Biogeosciences 6, 2637-2646, 2009.
Langer, G., Oetjen, K., and Brenneis, T.: Calcification of Calcidiscus leptoporus under nitrogen and phosphorus limitation, J. Exp. Mar. Biol. Ecol., 413, 131-137, 2012.
Langer, G., Oetjen, K., and Brenneis, T.: Coccolithophores do not increase particulate carbon production under nutrient limitation: A case study using Emiliania huxleyi (PML B92/11), J. Exp. Mar. Biol. Ecol., 443, 155-161, 2013.
LaRoche, J., Rost, B., and Engel, A.: Bioassays, batch culture and chemostat experimentation. In: Guide of best practices for ocean acidification research and data processing, Riebesell, U., Fabry, V. J., Hansson, L., and Gattuso, J.-P. (Eds.), Publications office of the European Union, Luxembourg, 2010.
Laws, E. A. and Bannister, T. T.: Nutrient-Limited and Light-Limited Growth of Thalassiosira-Fluviatilis in Continuous Culture, with Implications for Phytoplankton Growth in the Ocean, Limnol. Oceanogr., 25, 457-473, 1980.
Lee, K., Tong, L. T., Millero, F. J., Sabine, C. L., Dickson, A. G., Goyet, C., Park, G.-H., Wanninkhof, R., Feely, R. A., and Key, R. M.: Global relationships of total alkalinity with salinity and temperature in surface waters of the world's oceans, Geophysical Research Letters, 33, 2006.
Lefebvre, S. C., Benner, I., Stillman, J. H., Parker, A. E., Drake, M. K., Rossignol, P. E., Okimura, K. M., Komada, T., and Carpenter, E. J.: Nitrogen source and pCO2 synergistically affect carbon allocation, growth and morphology of the coccolithophore Emiliania huxleyi: potential implications of ocean acidification for the carbon cycle, Global Change Biology, n/a-n/a, 2011.
Linschooten, C., Vanbleijswijk, J. D. L., Vanemburg, P. R., Devrind, J. P. M., Kempers, E. S., Westbroek, P., and Devrinddejong, E. W.: Role of the Light-Dark Cycle and Medium Composition on the Production of Coccoliths by Emiliania-Huxleyi (Haptophyceae), Journal of Phycology, 27, 82-86, 1991.
Lohbeck, K. T., Riebesell, U., and Reusch, T. B. H.: Adaptive evolution of a key phytoplankton species to ocean acidification, Nature Geoscience, 2012.
López-Sandoval, D. C., Rodríguez-Ramos, T., Cermeño, P., Sobrino, C., and Marañón, E.: Photosynthesis and respiration in marine phytoplankton: Relationship with cell size, taxonomic affiliation, and growth phase, J. Exp. Mar. Biol. Ecol., 457, 151-159, 2014.
Lopez-Urrutia, A., San Martin, E., Harris, R.P. and Irigoien, X.: Scaling the metabolic balance of the oceans, Proceedings of the National Academy of Sciences, 103, 8739-8744, 2006.
Lorena, A., Marques, G. M., Kooijman, S. A. L. M., and Sousa, T.: Stylized facts in microalgal growth: interpretation in a dynamic energy budget context, Philosophical Transactions of the Royal Society B: Biological Sciences, 365, 3509-3521, 2010.
Marañón, E.: Inter-specific scaling of phytoplankton production and cell size in the field, Journal of Plankton Research, 30, 157-163, 2008.
Marañón, E., Cermeño, P., López-Sandoval, D. C., Rodríguez-Ramos, T., Sobrino, C., Huete-Ortega, M., Blanco, J. M., Rodríguez, J., and Fussmann, G.: Unimodal size scaling of phytoplankton growth and the size dependence of nutrient uptake and use, Ecol Lett, 16, 371-379, 2013.

1211 Marañón, E., Cermeño, P., Rodriguez, J., Zubkov, M. V., and Harris, R. P.: Scaling of phytoplankton
1212 photosynthesis and cell size in the ocean, Limnol. Oceanogr., 52, 2190-2198, 2007.
1213 Medlin, L. K., Barker, G. L. A., Campbell, L., Green, J. C., Hayes, P. K., Marie, D., Wrieden, S., and
1214 Vaulot, D.: Genetic characterisation of Emiliania huxleyi (Haptophyta), Journal of Marine Systems,
1215 9, 13-31, 1996.
1216 Meier, K. J. S., Beaufort, L., Heussner, S., and Ziveri, P.: The role of ocean acidification in Emiliania
1217 huxleyi coccolith thinning in the Mediterranean Sea, Biogeosciences, 11, 2857-2869, 2014.
1218 Menden-Deuer, S. and Lessard, E. J.: Carbon to volume relationships for Dinoflagellates, Diatoms,
1219 and other protist plankton, Limnol. Oceanogr., 45, 569-579, 2000.
1220 Mitrovic, S. M., Howden, C. G., and Bowling, L. C.: Unusual allometry between in situ growth of
1221 freshwater phytoplancton under static and fluctuating light environments: possible implications
1222 for dominance, Journal of Plankton Research, 25, 517-526, 2005.
1223 Montagnes, D. J. S., Berges, J. A., Harrison, P. J., and Taylor, F. J. R.: Estimating Carbon, Nitrogen,
1224 Protein, and Chlorophyll-a from Volume in Marine-Phytoplankton, Limnol. Oceanogr., 39, 1044-
1225 1060, 1994.
1226 Muller, E. B., Ananthasubramaniam, B., Klanjšček, T., and Nisbet, R. M.: Entrainment of cell division
1227 in phytoplankton with dynamic energy budgets, Journal of Sea Research, 66, 447-455, 2011.
1228 Muller, E. B. and Nisbet, R. M.: Dynamic energy budget modeling reveals the potential of future
1229 growth and calcification for the coccolithophoreEmiliania huxleyiin an acidified ocean, Global
1230 Change Biology, 20, 2031-2038, 2014.
1231 Muller, M. N., Antia, A. N., and LaRoche, J.: Influence of cell cycle phase on calcification in the
1232 coccolithophore Emiliania huxleyi, Limnol. Oceanogr., 53, 506-512, 2008.
1233 Müller, M. N., Beaufort, L., Bernard, O., Pedrotti, M. L., Talec, A., and Sciandra, A.: Influence of
1234 CO₂ and nitrogen limitation on the coccolith volume of Emiliania huxleyi
1235 (Haptophyta), Biogeosciences Discussions, 9, 4979-5010, 2012.
1236 Muller, P. J., Suess, E., and Ungerer, C. A.: Amino-Acids and Amino-Sugars of Surface Particulate and
1237 Sediment Trap Material from Waters of the Scotia Sea, Deep-Sea Res, 33, 819-838, 1986.
1238 Niklas, K. J. and Enquist, B. J.: Invariant scaling relationships for interspecific plant biomass
1239 production rates and body size, Proceedings of the National Academy of Sciences, 98, 2922-2927,
1240 2001.
1241 Oviedo, A. M., Langer, G., and Ziveri, P.: Effect of phosphorus limitation on coccolith morphology and
1242 element ratios in Mediterranean strains of the coccolithophore Emiliania huxleyi, J. Exp. Mar.
1243 Biol. Ecol., 459, 105-113, 2014.
1244 Paasche, E.: Marine plankton algae grown with light-dark cycles. I. Coccolithus huxlei, Physiologia
1245 Plantarum, 20, 946-956, 1967.
1246 Paasche, E.: Reduced coccolith calcite production under light-limited growth: a comparative study of
1247 three clones of Emiliania huxleyi (Prymnesiophyceae), Phycologia, 38, 508-516, 1999.
1248 Paasche, E.: A review of the coccolithophorid Emiliania huxleyi (Prymnesiophyceae), with particular
1249 reference to growth, coccolith formation, and calcification-photosynthesis interactions,
1250 Phycologia, 40, 503-529, 2001.
1251 Paasche, E., Brubak, S., Skattebol, S., Young, J. R., and Green, J. C.: Growth and calcification in the
1252 coccolithophorid Emiliania huxleyi (Haptophyceae) at low salinities, Phycologia, 35, 394-403,
1253 1996.
1254 Poulton, A. J., Adey, T. R., Balch, W. M., and Holligan, P. M.: Relating coccolithophore calcification
1255 rates to phytoplankton community dynamics: Regional differences and implications for carbon
1256 export, Deep Sea Research Part II: Topical Studies in Oceanography, 54, 538-557, 2007.
1257 Poulton, A. J., Stinchcombe, M. C., Achterberg, E. P., Bakker, D. C. E., Dumousseaud, C., Lawson, H.
1258 E., Lee, G. A., Richier, S., Suggett, D. J., and Young, J. R.: Coccolithophores on the north-west
1259 European shelf: calcification rates and environmental controls, Biogeosciences Discussions, 11,
1260 2685-2733, 2014.

Poulton, A. J., Young, J. R., Bates, N. R., and Balch, W. M.: Biometry of detached Emiliania huxleyi coccoliths along the Patagonian Shelf, Marine Ecology Progress Series, 443, 1-17, 2011.
Powell, E. O.: Growth rate and generation time of bacteria, with special reference to continuous culture, Journal of General Microbiology, 15, 492-511, 1956.
Raven, J. A.: The twelfth Tansley Lecture. Small is beautiful: the picophytoplankton, Functional Ecology, 12, 503-513, 1998.
Raven, J. A. and Crawfurd, K.: Environmental controls on coccolithophore calcification, Marine Ecology Progress Series, 470, 137-166, 2012.
Read, B. A., Kegel, J., Klute, M. J., Kuo, A., Lefebvre, S. C., Maumus, F., Mayer, C., Miller, J., Monier, A., Salamov, A., Young, J., Aguilar, M., Claverie, J.-M., Frickenhaus, S., Gonzalez, K., Herman, E. K., Lin, Y.-C., Napier, J., Ogata, H., Sarno, A. F., Shmutz, J., Schroeder, D., de Vargas, C., Verret, F., von Dassow, P., Valentin, K., Van de Peer, Y., Wheeler, G., Dacks, J. B., Delwiche, C. F., Dyhrman, S. T., Gloeckner, G., John, U., Richards, T., Worden, A. Z., Zhang, X., Grigoriev, I. V., Allen, A. E., Bidle, K., Borodovsky, M., Bowler, C., Brownlee, C., Cock, J. M., Elias, M., Gladyshev, V. N., Groth, M., Guda, C., Hadaegh, A., Iglesias-Rodriguez, M. D., Jenkins, J., Jones, B. M., Lawson, T., Leese, F., Lindquist, E., Lobanov, A., Lomsadze, A., Malik, S.-B., Marsh, M. E., Mackinder, L., Mock, T., Mueller-Roeber, B., Pagarete, A., Parker, M., Probert, I., Quesneville, H., Raines, C., Rensing, S. A., Riano-Pachon, D. M., Richier, S., Rokitta, S., Shiraiwa, Y., Soanes, D. M., van der Giezen, M., Wahlund, T. M., Williams, B., Wilson, W., Wolfe, G., Wurch, L. L., and Emiliania Huxleyi, A.: Pan genome of the phytoplankton Emiliania underpins its global distribution, Nature, 499, 209-213, 2013.
Reinfelder, J. R.: Carbon Concentrating Mechanisms in Eukaryotic Marine Phytoplankton, Annual Review of Marine Science, 3, 291-315, 2011.
Riebesell, U., Zondervan, I., Rost, B., Tortell, P. D., Zeebe, R. E., and Morel, F. M. M.: Reduced calcification of marine plankton in response to increased atmospheric CO2, Nature, 407, 364-367, 2000.
Riegman, R., Stolte, W., Noordeloos, A. A. M., and Slezak, D.: Nutrient uptake, and alkaline phosphate (EC 3 : 1 : 3 : 1) activity of Emiliania huxleyi (Prymnesiophyceae) during growth under N and P limitation in continuous cultures, Journal of Phycology, 36, 87-96, 2000.
Rokitta, S. D. and Rost, B.: Effects of CO2 and their modulation by light in the life-cycle stages of the coccolithophore Emiliania huxleyi, Limnol. Oceanogr., 57, 607-618, 2012.
Rost, B., Riebesell, U., Burkhardt, S., and Sultemeyer, D.: Carbon acquisition of bloom-forming marine phytoplankton, Limnol. Oceanogr., 48, 55-67, 2003.
Rost, B., Zondervan, I., and Riebesell, U.: Light-dependent carbon isotope fractionation in the coccolithophorid Emiliania huxleyi, Limnol. Oceanogr., 47, 120-128, 2002.
Rouco, M., Branson, O., Lebrato, M., and Iglesias-Rodríguez, M. D.: The effect of nitrate and phosphate availability on Emiliania huxleyi (NZEH) physiology under different CO2 scenarios, Frontiers in Microbiology, 4, 2013.
Satoh, M., Iwamoto, K., Suzuki, I., and Shiraiwa, Y.: Cold Stress Stimulates Intracellular Calcification by the Coccolithophore, Emiliania huxleyi (Haptophyceae) Under Phosphate-Deficient Conditions, Mar. Biotechnol., 11, 327-333, 2008.
Schiebel, R., Brupbacher, U., Schmidtko, S., Nausch, G., Waniek, J. J., and Thierstein, H.-R.: Spring coccolithophore production and dispersion in the temperate eastern North Atlantic Ocean, Journal of Geophysical Research, 116, 2011.
Schlüter, L., Lohbeck, K. T., Gutowska, M. A., Gröger, J. P., Riebesell, U., and Reusch, T. B. H.: Adaptation of a globally important coccolithophore to ocean warming and acidification, Nature Climate Change, 2014.
Schroeder, D. C., Biggi, G. F., Hall, M., Davy, J., Martínez, J. M., Richardson, A. J., Malin, G., and Wilson, W. H.: A genetic marker to separate emiliania huxleyi (prymnesiophyceae) morphotypes, Journal of Phycology, 41, 874-879, 2005.
Schulz, K. G., Rost, B., Burkhardt, S., Riebesell, U., Thoms, S., and Wolf-Gladrow, D. A.: The effect of iron availability on the regulation of inorganic carbon acquisition in the coccolithophore Emiliania

huxleyi and the significance of cellular compartmentation for stable carbon isotope fractionation, Geochim. Cosmochim. Acta, 71, 5301-5312, 2007.
Sciandra, A., Harlay, J., Lefevre, D., Lemee, R., Rimmelin, P., Denis, M., and Gattuso, J. P.: Response of coccolithophorid Emiliania huxleyi to elevated partial pressure of CO2 under nitrogen limitation, Marine Ecology-Progress Series, 261, 111-122, 2003.
Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.: Temperature Modulates Coccolithophorid Sensitivity of Growth, Photosynthesis and Calcification to Increasing Seawater pCO2, PLoS ONE, 9, e88308, 2014.
Shutler, J. D., Land, P. E., Brown, C. W., Findlay, H. S., Donlon, C. J., Medland, M., Snooke, R., and Blackford, J. C.: Coccolithophore surface distributions in the North Atlantic and their modulation of the air-sea flux of CO₂ from 10 years of satellite Earth observation data, Biogeosciences, 10, 2699-2709, 2013.
Smith, H. E. K., Tyrrell, T., Charalampopoulou, A., Dumousseaud, C., Legge, O. J., Birchenough, S., Pettit, L. R., Garley, R., Hartman, S. E., Hartman, M. C., Sagoo, N., Daniels, C. J., Achterberg, E. P., and Hydes, D. J.: Predominance of heavily calcified coccolithophores at low CaCO3 saturation during winter in the Bay of Biscay, Proceedings of the National Academy of Sciences, 109, 8845-8849, 2012.
Stolte, W., McCollin, T., and Noordeloos, A. M.: Effect of nitrogen source on the size distribution within marine phytoplankton populations, J. Exp. Mar. Biol. Ecol., 184, 83-97, 1994.
Takahashi, T.: Climatological mean and decadal change in surface ocean pCO2, and net sea–air CO2 flux over the global oceans, Deep Sea Research Part II: Topical Studies in Oceanography, 56, 554-577, 2009.
Thomas, M. K., Kremer, C. T., Klausmeier, C. A., and Litchman, E.: A global pattern of thermazl adaptation in marine phytoplankton, Science, 338, 1085-1088, 2012.
Trimborn, S., Langer, G., and Rost, B.: Effect of varying calcium concentrations and light intensities on calcification and photosynthesis in Emiliania huxleyi, Limnol. Oceanogr., 52, 2285-2293, 2007.
van Rijssel, M. and Gieskes, W. W. C.: Temperature, light, and the dimethylsulfoniopropionate (DMSP) content of Emiliania huxleyi (Prymnesiophyceae), Journal of Sea Research, 48, 2002.
Vanbleijswijk, J. D. L., Kempers, R. S., Veldhuis, M. J., and Westbroek, P.: Cell and Growth-Characteristics of Type-a and Type-B of Emiliania-Huxleyi (Prymnesiophyceae) as Determined by Flow-Cytometry and Chemical-Analyses, Journal of Phycology, 30, 230-241, 1994.
Walsby, A. F. and Raynolds, C. S.: Sinking and floating. In: The physiological ecology of phytoplankton, Morris, I. (Ed.), University of California Press, Berkeley, 1980.
West, G. B., Brown, J. H., and Enquist, B. J.: A general model for the origin of allometric scaling laws in biology, Science, 276, 122-126, 1997.
Wilson, J. D., Barker, S., and Ridgwell, A.: Assessment of the spatial variability in particulate organic matter and mineral sinking fluxes in the ocean interior: Implications for the ballast hypothesis, Glob. Biogeochem. Cycle, 26, 2012.
Winter, A., Henderiks, J., Beaufort, L., Rickaby, R. E. M., and Brown, C. W.: Poleward expansion of the coccolithophore Emiliania huxleyi, Journal of Plankton Research, 36, 316-325, 2013.
Young, J. R. and Henriksen, K.: Biomineralization within vesicles: The calcite of coccoliths, Biomineralization, 54, 189-215, 2003.
Young, J. R., Poulton, A. J., and Tyrrell, T.: Morphology of Emiliania huxleyi coccoliths on the northwestern European shelf – is there an influence of carbonate chemistry?, Biogeosciences, 11, 4771-4782, 2014.
Young, J. R. and Westbroek, P.: Genotypic variation in the coccolitophorid species Emiliania huxleyi, Marine Micropaleontology, 18, 5-23, 1991.
Zondervan, I.: The effects of light, macronutrients, trace metals and CO2 on the production of calcium carbonate and organic carbon in coccolithophores - A review, Deep-Sea Res. Part II-Top. Stud. Oceanogr., 54, 521-537, 2007.

Zondervan, I., Rost, B., and Riebesell, U.: Effect of CO2 concentration on the PIC/POC ratio in the coccolithophore Emiliania huxleyi grown under light-limiting conditions and different daylengths, J. Exp. Mar. Biol. Ecol., 272, 55-70, 2002.

Tables

Table 1 – Entries in the database of coccolitophore metabolism

	Column content	Units/explanation
General information	Literature reference	-
	Coccolitophore species	Species name
	Coccolitophore strain	Strain name
	Experiment type	Batch or chemostat
	Optimal temperature of strain	°C
Experimental conditions	Duration light period	hours
	Duration dark period	hours
	Sampling time	hours from beginning of light period
	Irradiance	μmol photons m^{-2} s^{-1}
	Temperature	°C
	Salinity	g/kg
	pCO2	μatm
	Dissolved inorganic carbon (DIC)	μmol kg^{-1}
	pH_T (total scale)	pH units
	Total alkalinity (TA)	μmol kg^{-1}
	Saturation state (calcite)	-
	Ca	mmol kg^{-1}
	Mg	mmol kg^{-1}
	NO_3	μmol kg^{-1}
	PO_4	μmol kg^{-1}
Experimental results	Organic C quota (POC)	pg_C $cell^{-1}$
	Inorganic C quota (PIC)	pg_C $cell^{-1}$
	Growth rate	day^{-1}
	Photosynthesis rate (RPh)	pgC $cell^{-1}$ day^{-1}
	Calcification rate (RCa)	pgC $cell^{-1}$ day^{-1}
	Coccosphere diameter	μm

1406 Table 2 – Subgroups of experiments and the experimental conditions that define them

Group name	n	Irradiance	pCO_2	TA	PO_4	NO_3	Fe	Ca
		μmol m^{-2} s^{-1}	μatm	mmol kg^{-1}	μmol kg^{-1}	μmol kg^{-1}	nmol kg^{-1}	mmol kg^{-1}
Optimum low pCO_2	85	≥ 80	150 - 550	2.1 – 2.45	≥ 4	≥ 64	replete	9.3-10
Optimum High pCO_2	87	≥ 80	551 - 1311	1.9 – 2.6	≥ 4	≥ 64	replete	9.3 – 11.1
Light-limited	30	< 80	140 – 850	2.0 – 2.56	≥ 4	≥ 64	replete	9.3-10
NO_3-limited	10	≥ 80	200-1200[a]	2.3-4.5	≥ 4	limiting	replete	4-10
PO_4-limited	21	≥ 80	250-1200 [a]	1-4.5	limiting	≥ 64	replete	4-10.6
Fe-limited	1	180	?	~2.35	4	64	limiting	10

1407 [a]The DIC system data presented in the literature does not lend itself to an accurate calculation
1408 of DIC system.
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1410 Table 3 – Changes of cell and coccosphere (sphere) volume for given changes in
1411 environmental conditions in culture experiments

Condition changed	Mean value of environment change	Average cell (sphere) volume change (μm^3)	Average % cell volume (sphere) change	Max cell volume (sphere) change (μm^3)	Max % cell volume (sphere) change
pCO_2	+ 209 μatm	+8.6 (+6.7)	+14.4 (+5.2)	+23.5 (+51.9)	+34.6 (+36.6)
pCO_2	+ 592 μatm	+12.6 (+10.3)	+27.3 (+12.8)	+63.4 (+131.8)	+214.2 (+185.6)
Irradiance	+ 193 μE m^{-2} s^{-1}	+16.7 (+39.5)	+38.0 (+53.0)	+45.0 (+93.0)	+120 (+152.2)
NO_3	Replete to limiting (~ 20 nM)	-14.0 (-33.0)	-22.2 (-22.5)	-26.3 (-80.0)	-33.1 (-36.8)
PO_4	Replete to limiting (~ 0.3 nM)	+50.8 (+73.5)	+43.8 (+58.1)	+77.1 (+93.6)	+67.8 (+120.3)
Fe	Replete to limiting	-32.2	-69.9	-32.2	-69.9
Temperature	+ 7.6 °C (+5.8°C)	-25.8 (-40.4)	-27 (-18.8)	-75.1 (-144.0)	-68 (-57.8)

1412 Table 4 – Changes of cell and coccosphere (sphere) diameter for given changes in
1413 environmental conditions in culture experiments

Condition changed	Mean value of environment change	Average cell (sphere) diameter change (μm^3)	Average % cell (sphere) diameter change	Max cell (sphere) diameter change (μm)	Max % cell (sphere) diameter change
pCO_2	+ 209 μatm	+ 0.2 (+ 0.1)	+ 4.5 (+ 1.6)	+ 0.5 (+ 0.7)	+ 10.4 (+ 11.0)
pCO_2	+ 592 μatm	+ 0.3 (+ 0.2)	+ 7.4 (+ 3.1)	+ 1.8 (+ 2.2)	+ 46.5 (+ 41.9)
Irradiance	+ 193 μE m^{-2} s^{-1}	+ 0.5 (+ 0.7)	+ 10.5 (+ 13.1)	+ 1.3 (+ 1.8)	+ 30.1 (+ 36.1)
NO_3	Replete to limiting (~ 20 nM)	- 0.4 (- 0.5)	- 8.2 (- 8.3)	- 0.7 (- 1.1)	- 12.5 (- 14.2)
PO_4	Replete to limiting (~ 0.3 nM)	+ 0.7 (+ 1.0)	+ 12.7 (+ 16.1)	+ 1.0 (+ 1.5)	+ 18.8 (+ 30.1)
Fe	Replete to limiting	-1.5	-33.3	-1.5	-33.3
Temperature	+ 7.6 °C (+5.8°C)	- 0.65 (-0.54)	-11.7 (-7.4)	-1.9 (-1.95)	-31.7 (-25)

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Figure captions

Figure 1 – Schematic representation of a coccolithophore cell surrounded by a shield of coccoliths. The coccolith bearing-cell is called the coccosphere (modified from Hendericks (2008)).

Figure 2 – Geometirc model used to obtain cell and coccophere geometry from measurements of the particulate organic carbon (POC) and particulate inorganic carbon (PIC) content per cell measured in culture experiments. Panels (a) and (c) show the relationship between POC and PIC and cell geometry (cell and coccosphere diameter) calculated with equation 7. Panels b) and d) show the relationship between the cell and coccosphere diameter calculated with equation 7 and that measured in culture experiments. Notes: panels (a) and (b) present data for *E. huxleyi*, panels (c) and (d) present data from the other coccolithophore species in the database. The filled black dots are the cell diameter, the empty red symbols are the coccosphere diameter and the empty blue symbols are the difference between the coccosphere and cell diameters.

Figure 3 – Allometric relationships between cell volume (equation 7) and photosynthesis rate (a,c) (equation 5), growth rate (b) (equation 3) and calcification rate (d) (equation 6). Notes: in panels (a) and (b) red dots are the experiments from the coccolithophore database carried out in optimum growth conditions and grey dots are published field measurements of metabolic rates for a large number of organisms (Marañón, 2008); in panels (c) and (d) symbols denote coccolithophore species (see legend) and all data refers to optimum growth conditions. The dotted lines are the linear regressions through the experimental coccolitophore data obtained in optimum growth conditions and the field data of Marañón (2008) (see table 2 for definition of optimum growth conditions).

Figure 4 – Effect of sub-optimum growth conditions on allometric relationships for coccolithophores. (a) rate of photosynthesis, (b) growth rate, (c) rate of calcification. Sub-optimum light and nutrient conditions result in cells having reduced metabolic rates compared to cells of equal size grown in optimal growth conditions (see table 2 for definition of growth conditions). The error bars apply only to a limited number of experiments (see text) and correspond to those experiments where the sampling time is not reported.

Figure 5 – Changes in cell-size and metabolic rates of *E. huxleyi* cells (first two columns) and other coccolithophore species (last column) subject to an increase in pCO_2. Note: for *E. huxleyi* symbols denote optimum-low pCO_2 conditions (red circles), optimum-high pCO_2 conditions (red dots), light-limited conditions (blue dots); for the other coccolithophore species symbols denote the species and all conditions are optimum, without distinction of pCO_2 range (see table 2 for definition of growth conditions).

Figure 6 - Changes in cell-size and metabolic rates of *E. huxleyi* cells subject to an increase in irradiance (starting from irradiance-limited conditions), an increase in temperature and a decrease in nutrients (starting from nutrient-replete conditions). The symbols represent the different growth conditions defined in table 2 except for iron for which there is only one datapoint.

Figure 7 – Geometry of Isochrysidales coccospheres along the BIOSOPE transect in the South-East Pacific ocean (Beaufort et al., 2008). (a) geographical location of the BIOSPE transect superimposed on the surface ocean chlorophyll concentration map obtained by satellite observations, (b) distribution of Isochrysidales coccosphere diameter in depth along the BIOSOPE transect determined by the SYRACO automated coccolith analyser system (Beaufort et al., 2008).

Figure 8 – Comparison of the geometry (coccosphere diameter and volume) of *Isochrysidales* coccospheres from the BIOSOPE transect with the geometry of *E. huxleyi* coccospheres from laboratory culture experiments. Histograms in panels (a) and (b) compare BIOSPE field data (grey) with experimental data (red). Horizontal bar graphs in panels (a) and (b) show the average changes in coccosphere geometry observed in *E. huxleyi* culture experiments for given changes in pCO_2, irradiance, temperature, NO_3 and PO_4. (c) Box-whisker plots comparing environmental conditions at the BIOSOPE stations where *Isochrysidales* coccosphere geometry measurements were made (grey) with the range of environmental conditions imposed in laboratory culture experiments with *E. huxleyi* (red). Box-whisker plots show the minimum value, lower quartile, median, upper quartile and maximum value of a given environmental parameter. Note: size data for Fe-limitation is from one experiment in Schultz et al. (2007) and refers to cell-size (not coccosphere-size).

Figure 9 – Simple physiological model of a dividing phytoplankton cell that reproduces the co-variation of metabolic rates and cell-size observed in coccolithophores. Notes: J_X –

assimilation fluxes; J_{GX} – growth fluxes generated from the uptake of nutrient X; J_{MATX} – maturation fluxes generated from the uptake of nutrient X; J_G – total growth flux contributing to the buildup of structure (biomass) M_V; P_R – total maturation flux contributing to the buildup of maturity E_H; SU – synthesizing unit.

Figure 10 – Evolution in time of modeled (a) maturity, (b) cell volume and (c) instantaneous growth rate of a cell undergoing ten successive cycles of growth and division. Notes: the horizontal dashed line in (a) represents the threshold value of accumulated maturity in the cell at which cell division takes place; the horizontal dashed line in (b) is the average cell volume when cell cycles attain steady state, the horizontal dashed line in (c) is the average instantaneous growth rate when cell cycles attain steady state and is numerically equivalent to the growth rate calculated from the generation time (vertical dashed lines) via equation 22 – it is conceptually equivalent to the growth rate measured from cell-counts in culture experiments.

Figure 11 – Effect of changing nitrate and phosphate concentrations on modeled cell volume and growth rate (a, c, d, e and f) and on the maturation flux P_R and the ratio of the growth to maturation fluxes P_G/P_R (b). Notes: the data points in (d) and (f) correspond to the shifts in cell-size and growth rate observed in laboratory cultures with *E. huxleyi* subject to a decrease in nitrate (d) and phosphate (f) concentrations.

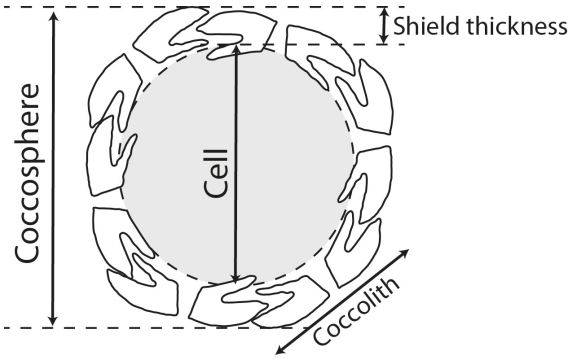

Figure 1

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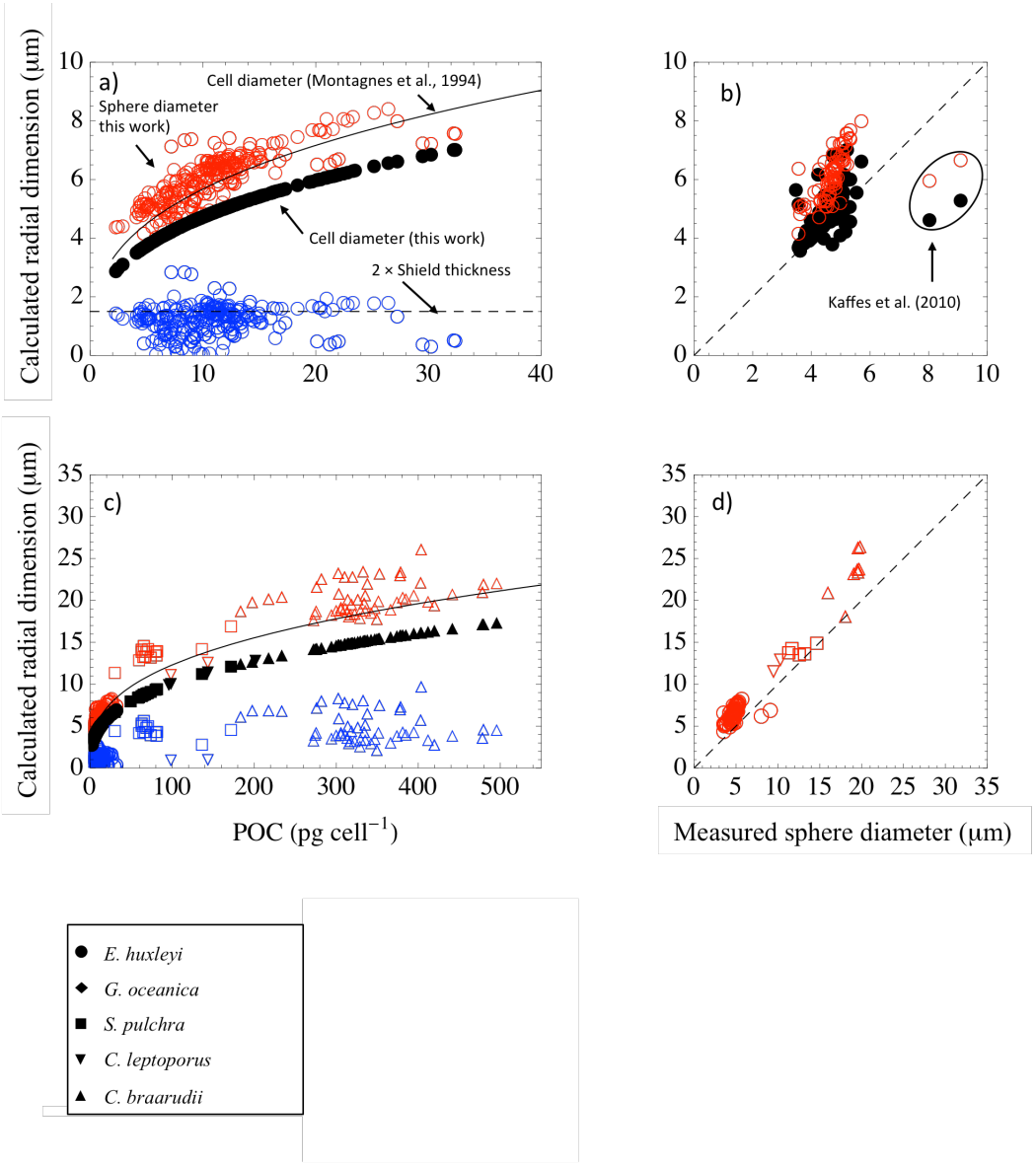
a)
Cell diameter (Montagnes et al., 1994)
Sphere diameter this work)
Cell diameter (this work)
2 × Shield thickness
Calculated radial dimension (µm)
b)
Kaffes et al. (2010)
c)
POC (pg cell^{-1})
d)
Measured sphere diameter (µm)
E. huxleyi
G. oceanica
S. pulchra
C. leptoporus
C. braarudii

Figure 2

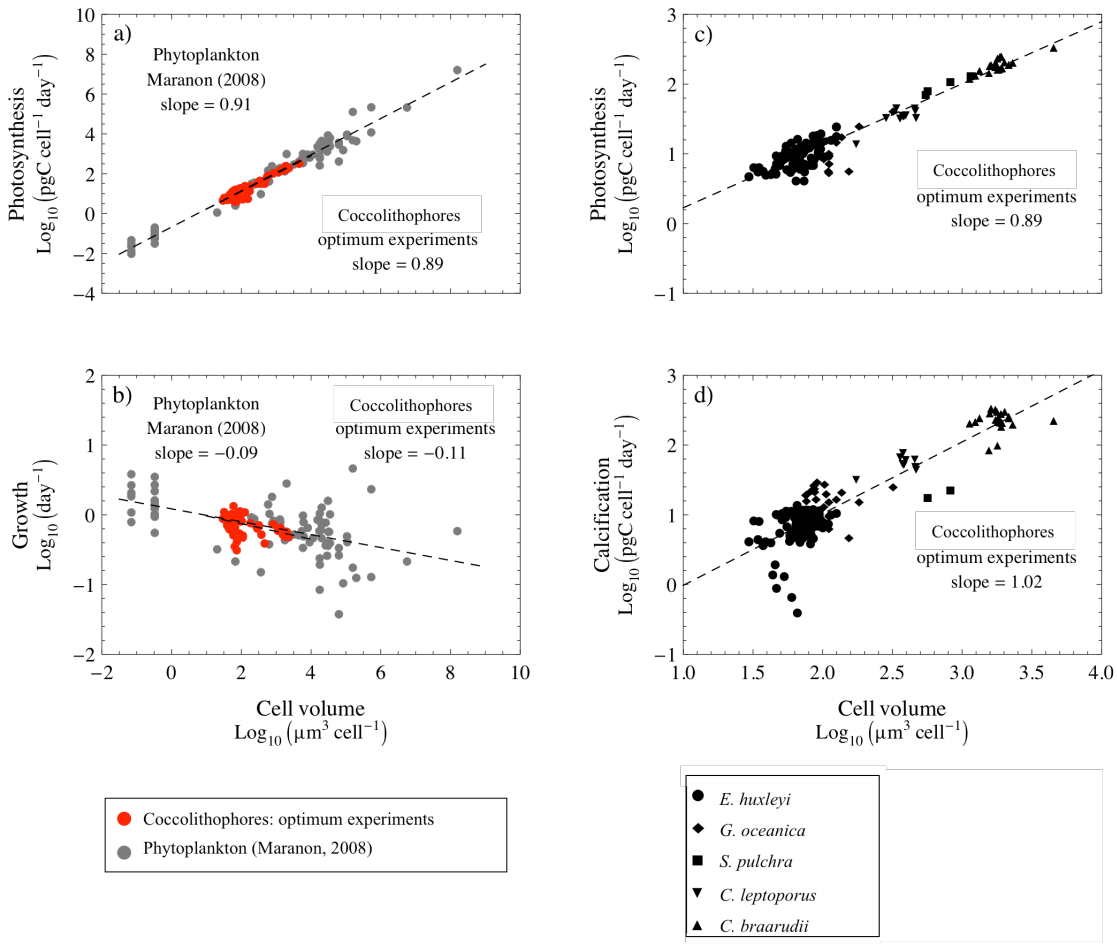
a)
Phytoplankton
Maranon (2008)
slope = 0.91
Coccolithophores
optimum experiments
slope = 0.89
Photosynthesis
Log10 (pgC cell−1 day−1)
c)
Coccolithophores
optimum experiments
slope = 0.89
Photosynthesis
Log10 (pgC cell−1 day−1)
b)
Phytoplankton
Maranon (2008)
slope = −0.09
Coccolithophores
optimum experiments
slope = −0.11
Growth
Log10 (day−1)
d)
Coccolithophores
optimum experiments
slope = 1.02
Calcification
Log10 (pgC cell−1 day−1)
Cell volume
Log10 (µm3 cell−1)
Cell volume
Log10 (µm3 cell−1)
Coccolithophores: optimum experiments
Phytoplankton (Maranon, 2008)
E. huxleyi
G. oceanica
S. pulchra
C. leptoporus
C. braarudii

Figure 3

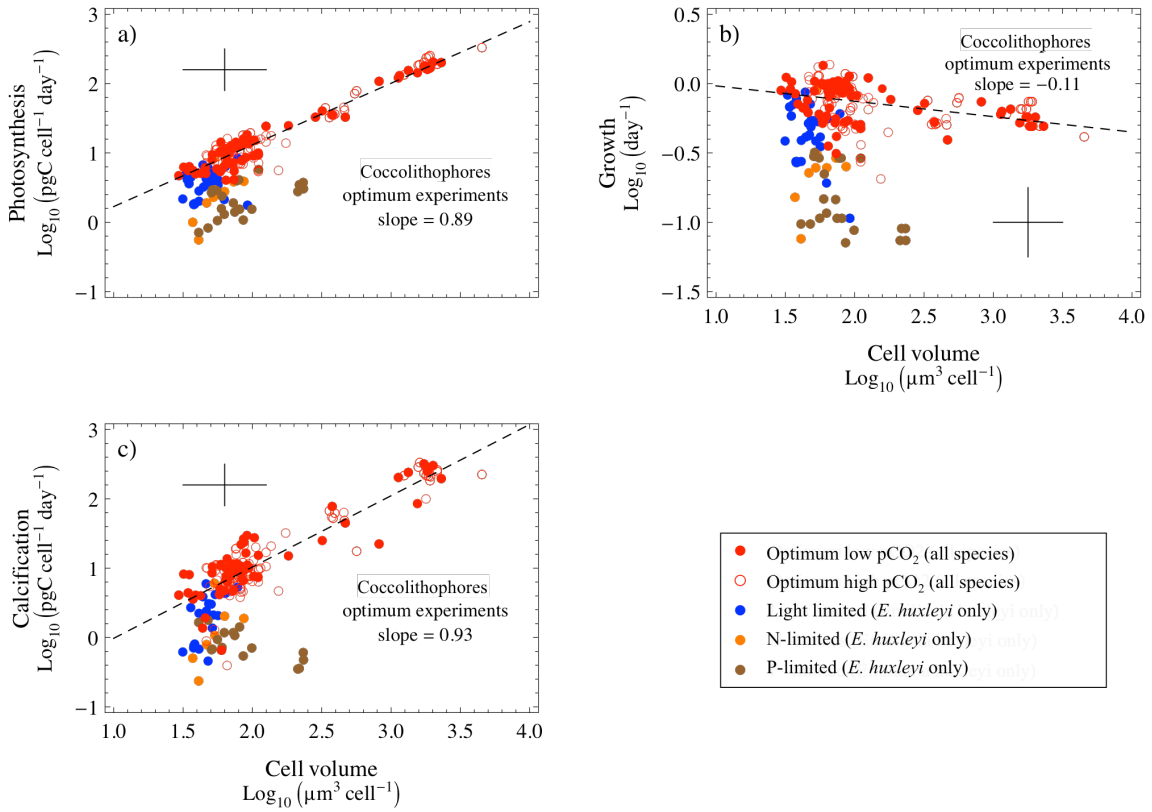
a)
Photosynthesis
Log10 (pgC cell−1 day−1)
Coccolithophores
optimum experiments
slope = 0.89
b)
Growth
Log10 (day−1)
Coccolithophores
optimum experiments
slope = −0.11
Cell volume
Log10 (µm3 cell−1)
c)
Calcification
Log10 (pgC cell−1 day−1)
Coccolithophores
optimum experiments
slope = 0.93
Cell volume
Log10 (µm3 cell−1)
Optimum low pCO2 (all species)
Optimum high pCO2 (all species)
Light limited (E. huxleyi only)
N-limited (E. huxleyi only)
P-limited (E. huxleyi only)

Figure 4

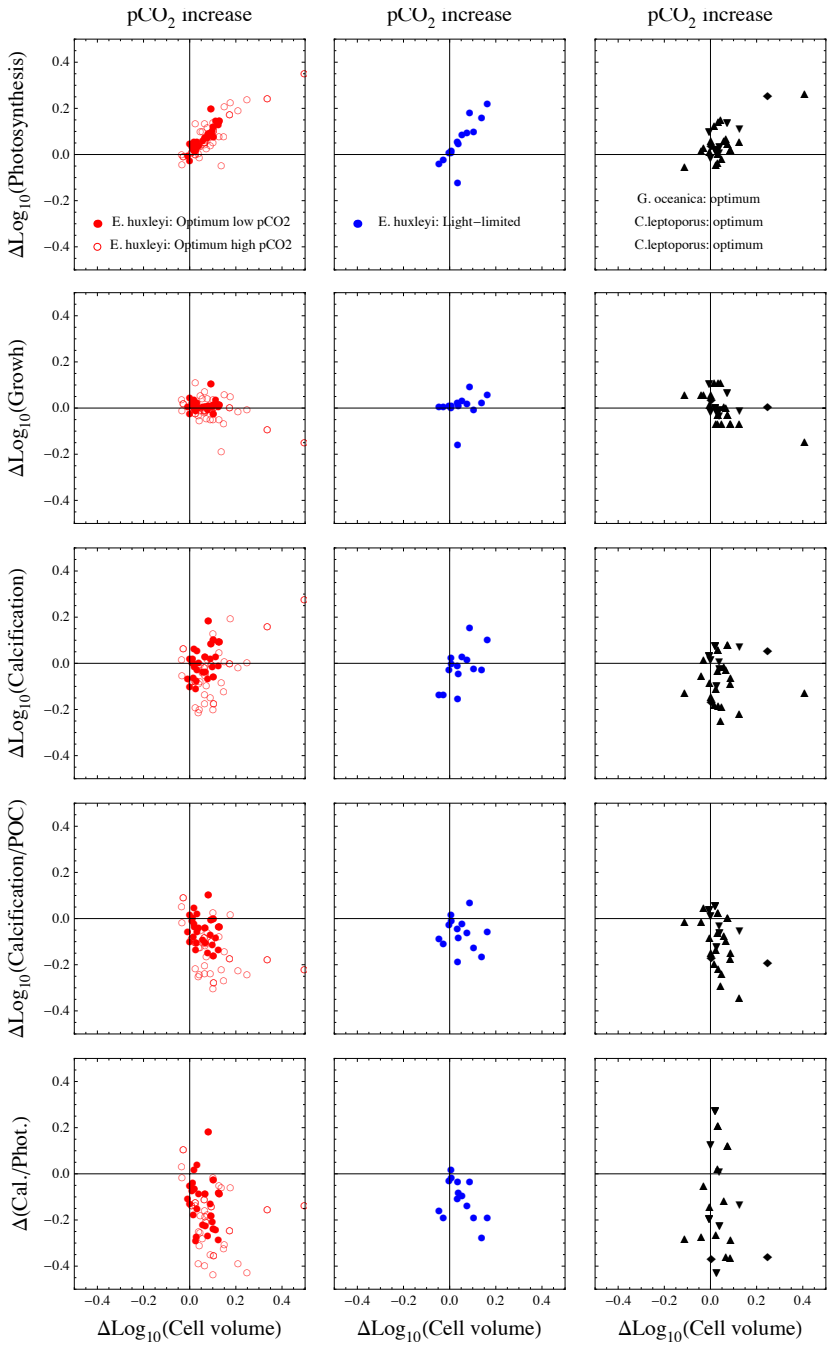
pCO2 increase
pCO2 increase
pCO2 increase
ΔLog10(Photosynthesis)
ΔLog10(Growth)
ΔLog10(Calcification)
ΔLog10(Calcification/POC)
Δ(Cal./Phot.)
E. huxleyi: Optimum low pCO2
E. huxleyi: Optimum high pCO2
E. huxleyi: Light−limited
G. oceanica: optimum
C.leptoporus: optimum
C.leptoporus: optimum
ΔLog10(Cell volume)
ΔLog10(Cell volume)
ΔLog10(Cell volume)

Figure 5

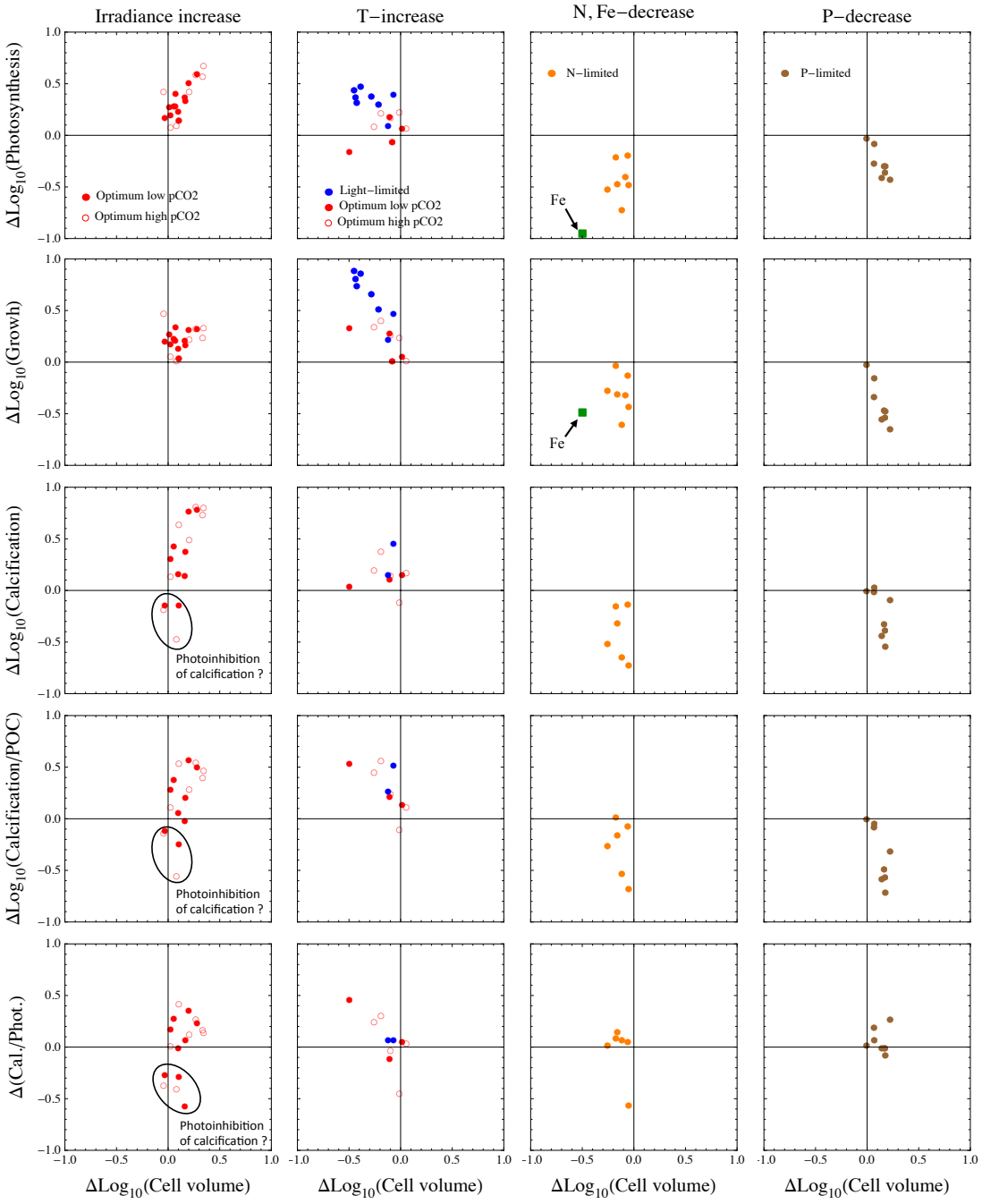

Figure 6

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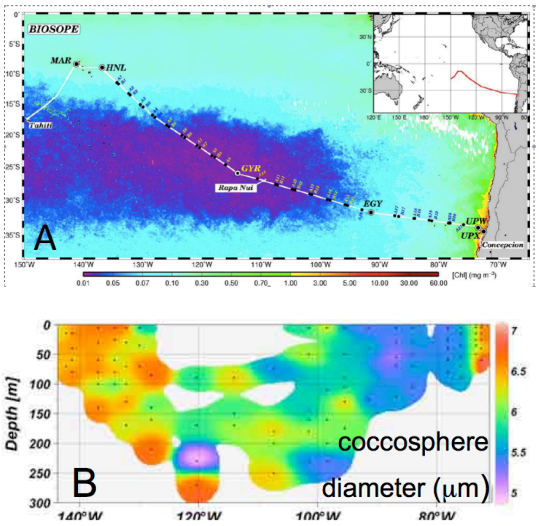

Figure 7

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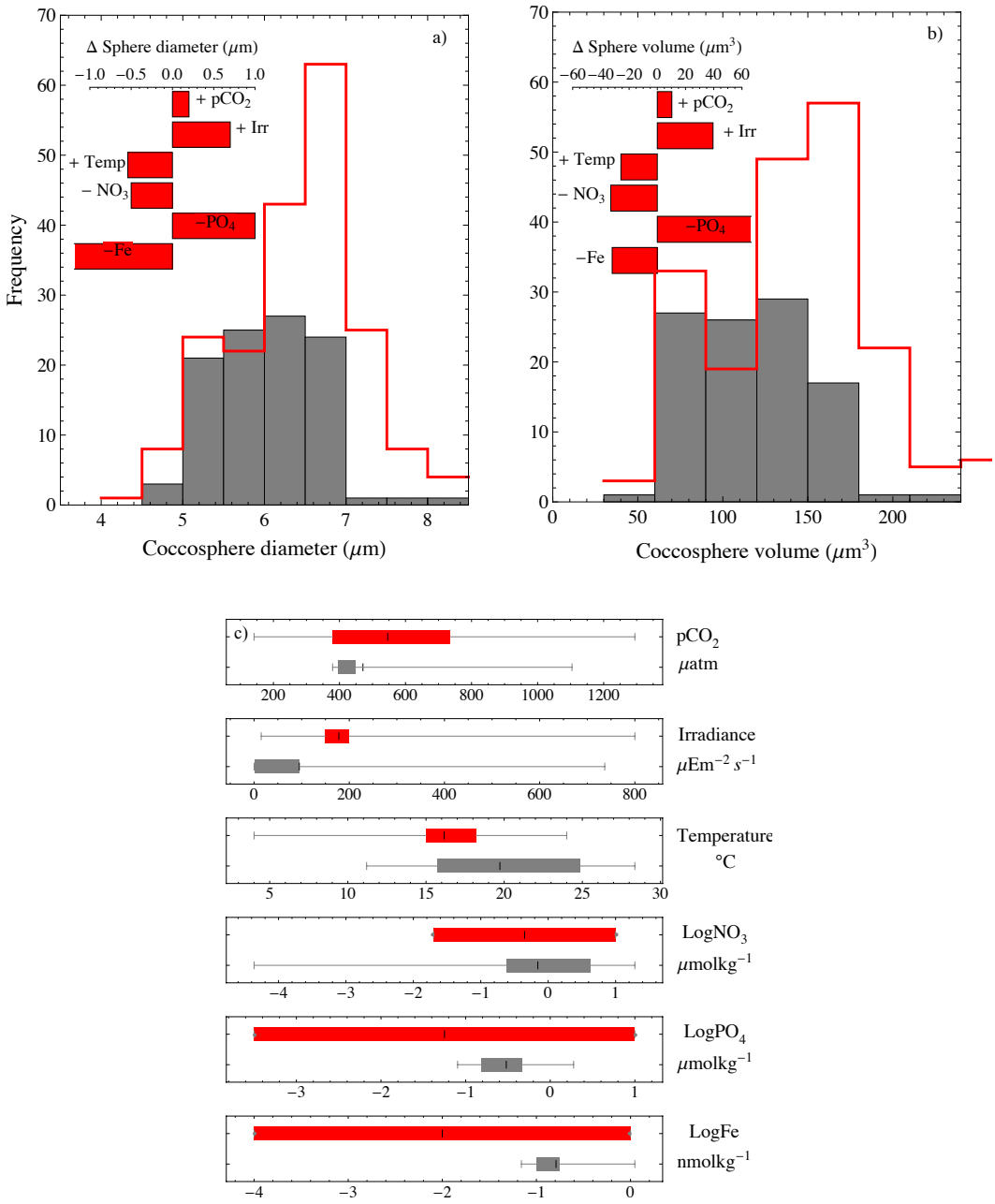

Figure 8

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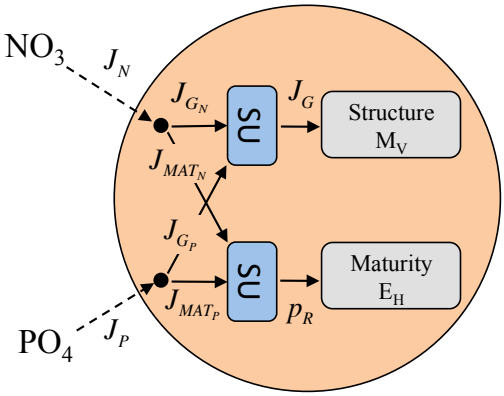

Figure 9

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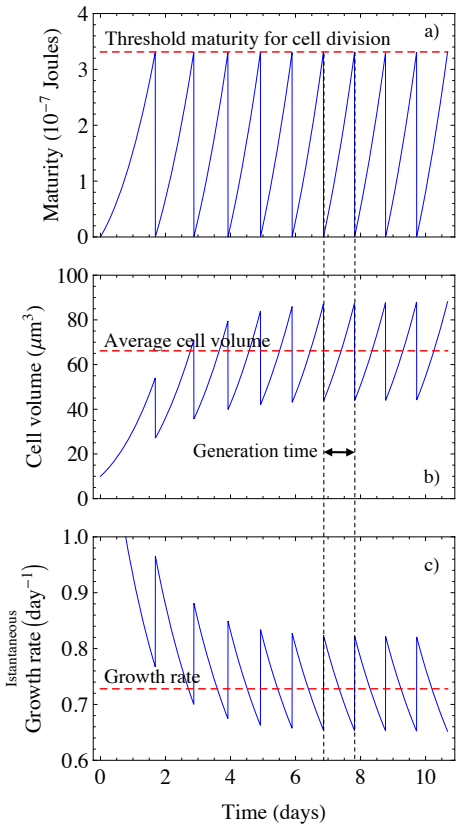

Figure 10

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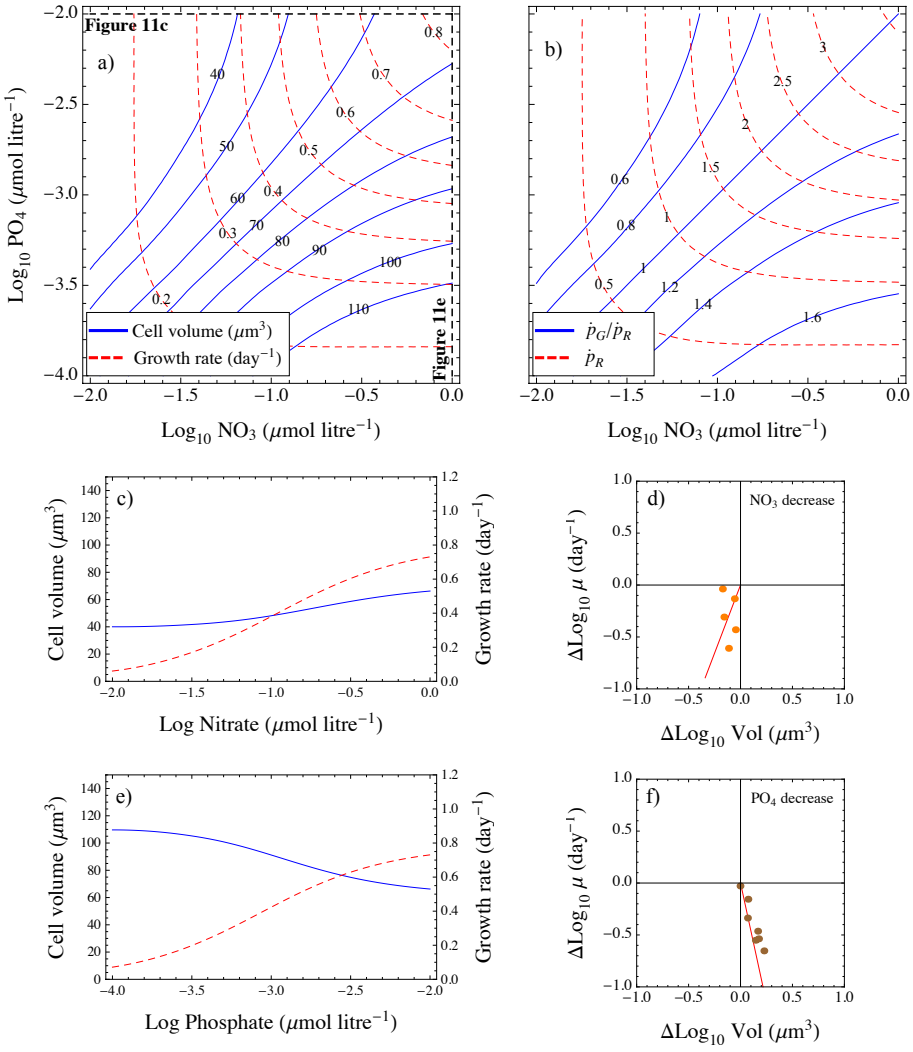

Figure 11

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