Authors response to comments by the Editor

Dear Dr Neftel,

We have added the temperature range information to Fig. 2 as requested by Reviewer 3. We thank the reviewer for their time in reviewing the manuscript a second time.

Kind regards,

David Wilson and co-authors

Authors response to comments from Reviewer 2

This paper adds data on emissions from peat-lands, contributing with 9 sites with thorough measurement program on which models are made for EF construction. The methods and data processing used are excellence. This study will be of large use for compilation of National Inventory reports, and hopefully making the large emissions from drained peat more visible. Thus laying the ground for emission reductions, by rewetting drained peat areas.

However I find it important to discuss more deep the influence of WT on the flux. A reader can get the faulty impression that mainly the temperature is the controlling factor, however the prerequisite is that the site is drained. This study also point to the need for additional measurements, on sites with more variable fertility and drain levels. Also there is a need to measure N2O, which this study did not include.

I find the paper overall interesting and easy to read, why I suggest a minor revision making the text and discussion more clear and fix some small errors.

**Response: We thank Reviewer 2 for the positive review of the manuscript and for the helpful comments and suggestions.**

Specific

Abstract

Line 6, Difficult sentence to read.

**Response: Could the reviewer be more specific as to the difficulty with the sentence?**

2.3 Environmental monitoring

Line 8-9, Explain what PP System and CPY-4 chamber means.

**Response:** **PP System is the manufacturer of the CPY-4 chamber. The chamber is described in the text. We have now provided the manufacturers details to accompany it.**

Soil temperature was recorded at 10 minutes interval, except DP3, where it was hourly intervals. Not clear. How much data gaps?

**Response: Agreed, this is unclear. While weather stations were indeed established at all sites (exception IP5), their data was not used in the calculation of the annual CO2-C balance at the IP sites. We have replaced the original text with**

***“Soil loggers (µ logger, Zeta-tec, UK, Hobo External Data Loggers, Onset Computer Corporation, MA, USA or Comark N2012 Diligence Loggers, Norwich, UK) were established in all the IP sites and recorded soil temperatures at hourly intervals. Weather stations were installed at all the DP sites and recorded photosynthetic photon flux density (PPFD; µmol m-2 s-1) and soil temperatures at 10 minute intervals. At DP3, soil moisture content (%) was also recorded (at 10 min intervals) by the weather station at that site.”***

2.5.1 Field measurements

22. clear acrylic chamber, measuring Reco, must include also photosynthesis and thus NEE? Why not tell?

**Response: Peat extraction had recently ceased at the site (IP6) and as the soil was totally devoid of vegetation, photosynthesis was not likely to occur. As such, NEE = Reco**

2.5.2 Flux calculations

For how long time was the chamber closed?

**Response: This varied between 60-180 secs. The information was given on P7500, L1 but was missing the units (i.e. seconds). This has been amended.**

2.5.4 Annual CO2-C balance

21. Each 12 month period. How many years? One year each site?

**Response: The information is already provided in Table 1 and in Figure 4.**

2.6 Peat fire emissions

Loose Irish moss peat. How decomposed?

**Response: The peat moss is Sphagnum peat ranging between H2-H3 on the von post decomposition scale.**

What influence could drying of the peat before combustion have on the result? In nature, peat fires continue, still it is not fully dry. Burning of peat in nature is not only in the surface, but deep down. How could this influence gases produced?

**Response: We found that in our setup it was not possible to ignite peat that had not been dried beforehand. This would also be the case in natural/managed peatlands, where surface vegetation fires will only spread into the peat when the peat is dry (i.e. during periods of drought) or is dried by the smouldering front moving through the peat. We follow the methodology of other peat fire studies (Christian et al., 2003; Yokelson et al., 1997; Stockwell et al., 2014) who all dry their peat samples before ignition.**

**The main difference between our lab-dried samples and drying in natural/managed peatlands is that the peat would be dry at the surface of open peatlands, but would retain moisture deeper down, whereas our lab samples are dry throughout. In open peatland fires, the combustion of dry peat at the surface may spread into deeper moister layers, but only after these have been dried by the heat produced from the combustion of the surface layer. This is likely to affect the rate of spread into the deeper, moister, peat as energy is used to dry these layers before combustion can commence (Rein et al., 2009). Rein et al. (2009) find that the main resultant effect of increasing peat moisture content on combustion emissions is an increase in the Modified Combustion Efficiency (MCE) with slightly higher (a few percent) CO2 emissions per unit mass of peat burned, whilst CO and CH4 emissions remain unaffected.**

3.2.1 Modelling

8. T is the temperature at which respiration reaches zero,… Should it not be T0?

**Response:** **Yes, the text should be *“T0 is the (minimum) temperature at which respiration reaches zero and is set here at 227.13 K, T is the soil temperature at 5 cm depth”*. This has now been amended.**

For equations 1-3 I have a reflection. For all these soils draining is the prerequisite for soil decomposing. Thus the water table depth >20cm is of need for these equations to be valid. This is why the effect of temperature becomes important, and only in some cases the WT becomes a limitation.

**Response: Agreed.**

3.4 Emission factors

For clearness I suggest here once again to tell the reader on which variable the EF’s are based.

**Response: We feel that this is not necessary, as the basis for the EFs has already been well described.**

3.5 Peat fire emission factors

MCE have not been defined/explained.

**Response: Combustion efficiency is a measure of the amount of fuel carbon released as CO2, and may be approximated using the *Modified Combustion Efficiency* (MCE) formula, which requires only a measurement of CO and CO2 rather than all the carbon containing gases (Yokelson *et al.*, 2008):**

$MCE= \frac{∆CO\_{2}}{∆CO\_{2}+ ∆CO}$ **(1)**

**Where ΔCO2 and ΔCO represent the elevated mixing ratios of these gases (the difference between mixing ratios measured in biomass burn emissions and those in the ambient air). MCE is often expressed as a percentage. Generally, an MCE lower than 0.9 (90%) is considered a low combustion efficiency burn (Lobert *et al.*, 1991; Yokelson *et al.*, 1996).**

MCE typical of smouldering combustion… Reference needed.

**Response: The following references will be cited, all of which publish MCE separately for flaming and smouldering combustion stages (where smouldering stage is typically below MCE of 0.9):**

***“(e.g. Yokelson et al. 1996; Bertschi et al. 2003)”***

4.1 Effects of climate

P7509 L3. Table 2 should be Table 3, Please check the numbering of tables and figures!

**Response: We would disagree. In this sentence we explicitly refer to the annual CO2-C emissions from Site IP6 and direct the reader to *Table 2. Emission factors (t CO2-Cha-1 yr-1) for sites IP1–6 and DP1–3. Uncertainties are 95% confidence intervals.***

L4-6. You say this confirms that soil temperature rather than water table is the main driver of emissions. I am not sure this could be said, since a prerequisite for all sites in this study is a WT level of >-20 cm. And for these types of systems, you show the temperature to be the most influential, which is OK if mentioning the prerequisites. This is confirmed by the wetter conditions and thus lower emission in the DP3 site.

**Response: We discuss the effects of drainage at length in Section 4.2. However, we agree that drainage is a pre-requisite at these sites and have now included this proviso in the text as follow;**

***“Given that all the sites are drained to a similar depth (Fig. 1), the variation in emissions appeared to be controlled largely by differences in soil temperatures between the sites (Fig. 6).”***

L19-. It is not clear how the LAI or PPFD could be drivers for peat decomposition. My suggestion is that the vegetation influences the water content of the soil, by transpiration, making it more aerobic, and thus higher soil CO2 flux. Thus the sunny days are more important than rainy. This also goes for LAI which also influence transpiration. Could this be discussed?

**Response: We did not state that LAI or PPFD were drivers for peat decomposition rather that they were drivers of GPP. Vegetation could stimulate decomposition of the more recalcitrant peat through the addition of labile organic matter (root exudates, leaves etc.). Under higher PPFD and LAI more organic matter is produced by the vegetation and could therefore lead to higher levels of priming in the older peat.**

**Drainage plays a much greater role in determining the water table position (and therefore the zone for aerobic decomposition) than transpiration at these sites. However, given the relatively shallow rooting depth of Calluna vulgaris (the dominant vegetation species at these sites) the effects of transpiration are likely to be confined to the upper 20 cm of the peat profile (**[**Aerts and Heil, 1993**](#_ENREF_1)**), where it may reduce the moisture content in the peat under warm temperatures and at low vapour pressure deficits (open stomata). This is confirmed at DP3, where the addition of the moisture content variable improved the performance of the Reco model at that site. We have now added the following text;**

 ***“During the growing season, the transpiration process is also likely to play a role in determining the moisture content of the peat within the rooting zone (~20cm depth) at these vegetated sites. Moisture losses are likely to be accentuated on sunny days when air and soil temperatures are high, when LAI values are highest (mid-summer) and when vapour pressure deficit is not a limiting factor. As CO2 emissions were closely correlated to soil temperature at 5 cm depth, reduced moisture content in this zone is likely to stimulate aerobic microbial activity.”***

4.3 Peat Characteristics

L10 Interesting that IP4, with C/N lower than 25 had highest emissions. IP2 had similar low C/N however not this high emission. You could have connected the C/N discussion to published similar studies.

**Response: A relationship between C:N ratios and CO2 emissions was not evident at our sites. As such, we endeavoured to concentrate on a discussion of variables that did have a tangible impact on CO2 emissions at these sites.**

4.4 Effects of peat extraction…

L11. It is odd, RH was only measured ad DP1, how could you say it is higher also for DP2 compared to….?

**Response: We provide details at P7506, L15-18 as to how we obtained an estimate for RH at DP2 and DP3.**

***“Estimated emissions from heterotrophic respiration (RH) at DP1 were 344 g CO2-Cm-2 yr-1, which equates to 49% of Reco at that site. Applying this proportional value to the other DP sites, we estimate that RH emissions to be 337 and 213 g CO2-Cm-2 yr-1 at DP2 and DP3 respectively.”***

4.5 Fire emission factors

L4. Here it is said: ‘ the importance of understanding the full suite of trace gas emissions from biomass burning, rather than focussing solely on CO2 and CH4 emissions.” The question then is: Why did you not include N2O in the measurements? In the wetland supplement it is only 4 studies on which the Tier 1 EF is based, temperate extraction sites. Some discussion on why you did not include this in the measurements would be good.

**Response: For the peat extraction areas, we focus solely on CO2 emissions in this discussion paper. CH4 and N2O have been quantified at some of the sites (but not all) and the data is currently being processed with a view to publication in the future.**

**In terms of the fire study, N2O is a difficult gas to measure using our FTIR setup as it can only be determined from spectra with very large enhancements of trace gases. This is because the N2O absorption occurs in a similar wavenumber region to both the CO2 and CO absorption bands (Paton-Walsh et al., 2014). Paton-Walsh et al. (2014) could only determine N2O from two of their five open fires, whilst Smith et al. (2014), who used a similar setup, failed to determine N2O from any of their 21 fires studied. In our study of Irish sphagnum moss peat burns, we found that excess mole fractions of N2O could not be correlated to either CO2 or CO for the determination of emission ratios, precluding the calculation of emission factors. One explanation for this is that N2O is predominantly a product of flaming combustion and is strongly correlated to CO2 (Paton-Walsh et al., 2014). The lack of flaming combustion in our peat burns probably explains our inability to detect significant excess N2O mole fractions.**

4.6 Implications…

L20. Why not say 1.7 t C?

**Response: Given that we found no significant difference between the IP and DP sites, we then used the mean value (1.68 t CO2-C) from all the sites as a single EF. Two decimal points provide a higher level of precision – particularly important for inventory reporting.**

P7515 L1. After ‘6’ I lack the word ‘times’.

**Response: Amended.**

Figures

Figure 3 and Figure 5 do not match. For Figure 3c the NEE show a net uptake of C but the Figure 5 shows NEE as loss. How come? Confusing.

**Response: Indeed. The caption to Fig. 3 should have included the following text**

***“(c) net ecosystem exchange (NEE; mg CO2-C m-2 hr-1) when PPFD>1000 µmol m-2 s-1 at sites DP1-3”.* This has now been added. The letters denoting differences between fluxes were also lost during the uploading process and are now presented below.**

****

**References**

**Aerts, R. and Heil, G. (Eds.): Heathlands: Patterns and processes in a changing environment., Springer Netherlands, 1993.**

**Bertschi , I., Yokelson, R. J., Ward, D. E., Babbitt, R. E., Susott, R. A., Goode, J. G., Hao, W. M. (2003) Trace gas and particle emissions from fires in large diameter and belowground biomass fuels. Journal of Geophysical Research 108: 8472, doi:10.1029/2002JD002100.**

**Christian, T. J., Kleiss, B., Yokelson, R. J., Holzinger, R., Crutzen, P. J., Hao, W. M., Saharjo, B. H. and Ward, D. E. (2003) Comprehensive laboratory measurements of biomass-burning emissions: 1. Emissions from Indonesian, African, and other fuels. Journal of Geophysical Research, 108(D23), 4719.**

**Lobert, J.M., Scharffe, D.H., Hao, W.M., Kuhlbusch, T.A., Seuwen, R., Warneck, P., Crutzen, P.J. (1991) Experimental evaluation of biomass burning emissions: Nitrogen and carbon containing compounds. In: Global Biomass Burning: Atmospheric, Climatic and Biospheric Implications. Ed: Levine, J.S., MIT Press, Cambridge, Massachusetts.**

**Paton-Walsh, C., Smith, T. E. L., Young, E. L., Griffith, D. W. T., and Guérette, É.-A. (2014) New emission factors for Australian vegetation fires measured using open-path Fourier transform infrared spectroscopy – Part 1: Methods and Australian temperate forest fires, Atmos. Chem. Phys., 14, 11313-11333, doi:10.5194/acp-14-11313-2014.**

**Rein, G., Cohen, S., & Simeoni, A. (2009). Carbon emissions from smouldering peat in shallow and strong fronts. Proceedings of the Combustion Institute, 32(2), 2489-2496.**

**Smith, T. E. L., Paton-Walsh, C., Meyer, C. P., Cook, G. D., Maier, S. W., Russell-Smith, J., Wooster, M. J., and Yates, C. P. (2014) New emission factors for Australian vegetation fires measured using open-path Fourier transform infrared spectroscopy – Part 2: Australian tropical savanna fires, Atmos. Chem. Phys., 14, 11335-11352, doi:10.5194/acp-14-11335-2014.**

**Stockwell, C.E., Yokelson, R.J., Kreidenweis, S.M., Robinson, A.L., DeMott, P.J., Sullivan, R.C., Reardon, J., Ryan, K.C., Griffith, D.W.T., Stevens, L. (2014) Trace gas emission from combustion of peat, crop residue, domestic biofuels, grasses, and other fuels: configuration and Fourier transform infrared (FTIR) component of the fourth Fire Lab at Missoula Experiment (FLAME-4) Atmospheric Chemistry and Physics 14: 9727–9754.**

**Yokelson, R.J., Griffith, D.W.T., Ward, D.E. (1996) Open-path Fourier transform infrared studies of large-scale laboratory fires. Journal of Geophysical Research 101(D15): 21067–21080.**

**Yokelson, R. J., Susott, R., Ward, D. E., Reardon, J., & Griffith, D. W. (1997). Emissions from smoldering combustion of biomass measured by open‐path Fourier transform infrared spectroscopy. Journal of Geophysical Research: Atmospheres (1984–2012), 102(D15), 18865-18877.**

**Yokelson, R.J., Christian, T. J., Karl, T. G., Guenther, A. (2008) The Tropical Forest and Fire Emissions Experiment: laboratory fire measurements and synthesis of campaign data. Atmospheric Chemistry and Physics 8: 3509–3527.**

Authors response to comments from Reviewer 3

**General comments**

Peat extraction is a major land use category of peatlands in some countries, and the recent emission factor of the IPCC 2013 Wetland Supplement is indeed based on only a few studies mainly on boreal sites. Therefore, new emission data on peatlands managed for extraction is very welcome and fits well into the scope of Biogeosciences. However, there are some methodological issues which need to be addressed before the manuscript can be published.

**Response: We thank Reviewer 3 for the positive review of the manuscript and for the helpful comments and suggestions.**

**1)** The methods section is very brief and needs to be extended.

2.5.1 Field measurements:

• How often did you measure Reco and GPP at each measurement date?

**Response: Reco: 2-4 measurements per collar per measurement date. NEE (light): 3- 8 measurements per collar per measurement date. This has been added to the text.**

• Did you ensure that the maximum PPFD was reached at each measurement date to avoid an

extrapolation beyond measured values?

**Response: Yes. The measurements were carried out so as to cover the full range of PPFD on a given day (see response below).**

• Did you measure at different PPFD levels by shading or at different times of the day?

**Response: Measurements were carried out between 8 am and 6 pm in the summer and 9am and 3pm in the winter and covered the full range of PPFD on a given day. Artificial shading was used early in the morning to obtain low PPFD levels (<100 µmol m-2 s-1). This has been added to the text.**

• For the NEE measurement: were the chambers also cooled (e.g. by icepacks) and the

temperature measured inside the chamber to avoid more artificial conditions than necessary?

**Response: Yes, the chambers were cooled with a cooling system. The system involves the continual pumping of iced water (from submersion of ice packs/bottles) from a container through a hose pipe into a small car radiator located in the chamber and back to the container via a second hose pipe. Two fans located in the chamber ensure that the cooler air is mixed within the chamber. Air temperature was measured in the chamber continually. The setup is described in detail by (**[**Alm et al., 2007**](#_ENREF_1)**).**

• What kind of chambers have been used for NEE measurements?

**Response: The same polycarbonate chambers (60 x 60 x 33 cm) were used (as described on P7499, L18-19 in the ms). The following information has been added;**

***“At the DP sites, net ecosystem exchange (NEE) was measured under a range of ambient light levels (PPFD; µmol m-2 s-1) prior to Reco measurements with the same polycarbonate chambers described above.”***

• How is the light transmissivity of these chambers (usually, it does not reach 100%) and was this accounted for when modelling GPP? It should be included in the GPP model as otherwise GPP might be underestimated.

**Response: Accounting for the light transmissibility of the chambers is valid if the PPFD sensor is located external to the chambers (e.g.** [**Beyer and Höper, 2015**](#_ENREF_2)**;** [**Beyer et al., 2015**](#_ENREF_3)**). Our PPFD sensors are located within the chamber, so the PPFD recorded during each NEE measurement period is the “attenuated” value (our chambers attenuate light transmissibility by ~12%). When modelling GPP, we used the relationship between fluxes (estimated GPP = measured NEE - measured Reco) and the PPFD values from inside the chamber to produce light response curves.**

**The GPP models were then used with the PPFD time series recorded by the external PPFD sensors on the weather stations to reconstruct the annual CO2 balance.**

**The light response curves are only valid for the range of PPFD values recorded during NEE measurements. This does mean that the highest PPFD value recorded during flux measurements (from inside the chamber) is always likely to be around 12% less than the actual PPFD value (measured by the weather station). However, we feel that this results in minimal underestimation of GPP as (a) it impacts on a very small number of hourly fluxes (<0.1%; i.e. number of occasions in the year where PPFD values recorded by the weather station > than the maximum observed PPFD value in the chamber) and (b) the plots are light saturated at PPFD >1000µmol m-2 s-1, so the difference in NEE at PPFD values >2200µmol m-2 s-1 is likely to be minor.**

• Why didn't you chose a site under ongoing (or recently ceased) industrial extraction?

**Response: A recently ceased extraction site was chosen (IP6), however, for the remainder it was not possible to establish monitoring sites for either logistical or equipment security reasons. Indeed, the decision to locate the sites on abandoned areas with limited access has proven to be sensible given that equipment (weather station batteries, solar panels etc.) at site IP6 have been stolen on a number of occasions.**

2.5.2 Flux calculation

• “GPP was calculated as NEE minus Reco”: Which value of Reco was used; the nearest

value in time or the one calculated by the model? If there was only one Reco measurement

per measurement date, using the actual measured value could potentially induce some

uncertainty as the time lag between the Reco measurement and the first GPP measurement is

not clear and as there is probably a strong temperature-dependent diurnal variability of

Reco.

**Response: The Reco value that was used was always the value closest in time to the NEE measurement. In winter time, diurnal variation in the soil temperature was very small and approximately two Reco flux measurements per plot were taken. In summer, when diurnal changes in soil temperature were very pronounced up to four Reco flux measurements per plot were carried out.**

2.5.3 Modelling

• Modelling GPP should be included in this sub-chapter

• Why was this specific GPP model chosen, and not a Michaelis-Menten type model, which is

frequently used for GPP?

**Response: The basic form of the GPP model used in this study (see Eq. 4) is a Michaelis-Menten type model, which has been used to describe the saturating response of photosynthesis to PPFD in numerous studies (e.g.** [**Tuittila et al., 1999**](#_ENREF_10)**;** [**Byrne et al., 2005**](#_ENREF_4)**;** [**Laine et al., 2006**](#_ENREF_7)**). The Levenberg-Marquardt algorithm described in the manuscript is a multiple non-linear regression technique used to derive model parameters and associated standard errors. The text has been amended as follows;**

***“GPP was related to PPFD using the Michaelis–Menten type relationship that describes the saturating response of photosynthesis to light (***[***Tuittila et al., 1999***](#_ENREF_10)***). GPP model coefficients and associated standard errors were estimated using the Levenberg-Marquardt multiple non-linear regression technique (IBM SPSS Statistics for Windows, Version 21.0. Armonk, NY, USA).”***

• How was “plot-specific LAI” modelled?

**Response: This was described on P7499, L3-12**

***“However, at the DP sites a vegetation component is present and in order to incorporate the seasonal dynamics of the plants into CO2-C exchange models, the leaf area index (LAI) was estimated for each of the collars. This involved accounting for the green photosynthetic area of all vascular plants (leaves and stems) within the collar at monthly intervals. In short, the number of leaves and stems were counted from five subplots (8 x 8cm) within each collar. The size (length, width) of the leaves was measured from sample plants outside the collars. The LAI was then calculated by multiplying the estimated number of leaves by an area estimate of the leaf. Moss and lichen % cover was estimated at the same time. Species-specific model curves were applied to describe the phenological dynamics of the vegetation of each collar, and the models (vascular plants and moss) were summed to produce a plot-specific LAI. For a detailed description of the method see Wilson et al. (2007).”***

**2)** My major concern, however, are the NEE results of the GP sites. The vast majority of the

measured fluxes (Figure 3) show an uptake of CO2, especially at site DP1, but all annual balances (Table 2) show a net release. How do you explain these results?

In my opinion, there are several possibilities:

• The measurements themselves are biased: but as they were regular and rather frequent, I

would suggest that this is not the reason for the surprising results.

• There are problems with either the Reco or the GPP model. Here, I would recommend

checking the following issues:

• Is there extrapolation beyond the range of measured temperature and PPFD data?

• Regarding the GPP model: I doubt whether one “general model” and especially one

(equation 4) not including the LAI could predict correctly the NEE for the whole

year. Therefore, pooling data of several measurement campaigns might be an

alternative if the LAI is not included.

• Given the obvious discrepancy between measured NEE values and modelled sums I

would strongly suggest that the authors check their model by e.g. a cross-validation

approach (leaving one measurement date out at a time and trying to predict both

Reco, GPP and NEE by the remaining data).

**Response: Reviewer 2 has also pointed out this discrepancy. The caption to Fig. 3 should have included the following text**

***“(c) net ecosystem exchange (NEE; mg CO2-C m-2 hr-1) when PPFD>1000 µmol m-2 s-1 at sites DP1-3”.* This has now been added. The letters denoting differences between fluxes were also lost during the uploading process and are now presented below.**

****

3) Measuring emissions from burning peat is a valuable addition to the manuscript. However, I'm not sure whether these numbers are to be used for Lfire (Wetlands Supplement) – is fire an issue for non-vegetated peat extraction sites in Ireland and the UK? Otherwise, wouldn't be burning of peat reported in the energy sector (and how is it done now if there are no numbers available)? These issues should be made clearer especially for those not familiar with reporting methodologies.

**Response: Emissions associated with the burning of peat are reported under the Energy sector. The peat burning EFs in this study (Table 3) are primarily to be utilised for Lfire (i.e. on site). Fires do occur on non-vegetated sites in ROI and the UK, particularly in very dry years. For clarity, the following text has been added to the introduction:**

***“Emissions associated with off-site peat combustion are reported under the Energy sector and are not considered further here.”***

**Specific comments**

**Abstract**

I have recently experienced some discussions during which the relatively low EFs for peat

extraction sites (at least compared to agriculture) tended to raise the rather questionable opinion that peat extraction is a climate-friendly activity in peatlands. Therefore, it would be helpful to clearly include a statement on the system boundaries of your study, especially as you included peat burning but no horticulture.

**Response: We clearly state the boundaries of our study in the discussion section 4.7 but have added the following text in the abstract:**

***“Drainage related methane (CH4) and nitrous oxide (N2O) emissions, as well as CO2-Cemissions associated with the off-site decomposition of horticultural peat were not included.”***

**Results**

Both WT and VMC enter the Reco-equations without any additional model parameter. That

suggests that a) there is no optimum water level or moisture for respiration which I would expect to exist and b) respiration is highest at that highest VMC, i.e. at saturation which is rather surprising.

Could you comment on this?

**Response: Both points made by the Reviewer above are valid. The Reco models in this study are controlled by soil temperature. While the addition of WT and VMC improved the performance of the models at some of the sites, the improvement was only slight. We feel that this is due to the fairly narrow range of WT/VMC values recorded over the course of the 12 month study (e.g. the range in VMC values at DP3 only ranged between 56-64%). Therefore, optimum WT /VMC levels for respiration may not have been encountered. The Reco models used here are only valid for the data that was measured over the course of the study at each site and should not be extrapolated beyond the range of that data.**

**Discussion**

Generally, I do not really understand why the emissions are lower than in other studies: Your study areas are relatively warm with mild winters (at least compared to boreal sites), the physical environment is with maximum soil temperatures of 28°C not too extreme, and at least some of the sites are characterised by rather narrow CN ratios and sub-neutral pH-values. How is the WT compared to previous studies?

**Response: The WT levels are our sites are similar to the other studies in Fig. 7 (where a WT is reported).**

In my opinion, the choice of sites might be a reason why the emissions are lower than in other studies: Easily degradable organic matter would have been already gone, while during abstraction there would have been also less decomposed “fresh” peat at the surface during certain periods of time. Furthermore, your chose “unvegetated microsites” for the measurements, which suggests that parts of the peatlands are already re-vegetated. Probably, these microsites are unvegetated for a reason (peat quality, water repellency,...), and these conditions might also limit microbial activity.

**Response: The CO2 emissions from our sites are slightly higher than those reported for Fenno-Scandia but lower than emissions from Canadian sites. We believe that this is due to the peat end-use requirement in Canada (i.e. horticultural peat). As the Reviewer has stated, this latter peat is likely to be more fibric, less decomposed and produce higher CO2 emissions. In contrast, the residual peat at our IP sites and at the majority of Fenno-Scandia sites is utilised for energy production as it is more decomposed. We have discussed this at length in section 4.4.**

**In regard to the “unvegetated microsites”, areas around the periphery of the industrial sites may be vegetated, as they are often close to a seed source, while the remainder can remain largely bare and unvegetated for decades after the cessation of peat extraction. Even where a seed source is available plant establishment and survival are made more difficult by the edaphic conditions that may exist in the upper layers of the peat surface. As the reviewer has mentioned the lack of plant colonisation could be due to an unsuitable nutrient status (**[**Wind-Mulder et al., 1996**](#_ENREF_13)**) but could also be caused by the instability of the peat surface (**[**Campbell et al., 2002**](#_ENREF_5)**), water table fluctuations (**[**Price, 1997**](#_ENREF_9)**), high evaporative losses (**[**Waddington and Price, 2000**](#_ENREF_11)**) and high peat temperatures in mid-summer (**[**Waddington and Warner, 2001**](#_ENREF_12)**).**

Are there any obvious differences between the vegetated and non-vegetated sites?

**Response: In the ROI sites, peat type would appear to be an obvious difference between DP and IP sites, however this division does not hold up when the UK sites are included.**

Effects of drainage level: To my understanding, the effect of the WT on single fluxes (i.e. the Reco dynamics) and the effect on the general emission level shouldn't be mixed up. While at the scale of single fluxes, effects of the WT might be obscured by a co-variance between WT and temperature, or the activity of the vegetation, the general height of the emission might indeed be influenced by the WT (which seems not need to be the case in the study). However, at the scale of single fluxed, I do not think that concluding that there is generally (nearly) no effect of the WT is not valid unless fluxes from all sites are combined into one model.

**Response: We do not think that we state that there is no effect of WT on CO2 emissions at the site level. We have stated that, based on our results, soil temperature at 5 cm depth is the strongest determinant of fluxes but that WT depth (and VMC) also play a role in some sites. Clearly, there is a certain element of co-variance between soil temperature and WT at play. However, the Reviewer is correct to point out that the effect of WT on the general emission level is different. Our sites are all drained to varying degrees (deeper than -20cm) and as such the emissions that we have measured are all a function of the drainage. While we did not find a close relationship between annual CO2-C and any of the WT parameters (e.g. mean, max or min) across the sites, this is likely to change if rewetted sites (WT shallower than -20cm), for example, were included in the analysis. We have added the following text in the Discussion;**

***“Given that all the sites are drained to a similar depth (Fig. 1), the variation in emissions appeared to be controlled largely by differences in soil temperatures between the sites (Fig. 6).”***

How do you differentiate between areas influenced by domestic peat cutting and otherwise

disturbed peatlands with similar WT or vegetation which not used for agriculture or forestry? To do so, you would probably need to define a zone of influence. You briefly mention this problem in the discussion section, and I agree that there will be a problem with the activity data. Do you see any way forward to identify domestic peat cutting areas?

**Response: Activity data for domestic peat cutting is highly problematic for both jurisdictions. In the ROI, there are potentially 600,000 ha of peatlands (~30% of the total peatland area) affected to some degree by domestic peat extraction (**[**Malone and O'Connell, 2009**](#_ENREF_8)**). We have added the following text to the discussion of activity data;**

***“Determining to what degree that peatlands have been affected by domestic peat extraction and how far those impacts extend into the main peatland area are obvious challenges facing future research. The use of remote sensing platforms could provide high resolution data that will be able to differentiate between domestic peat extraction and other disturbed peatlands. In particular, the use of Unmanned Aerial Vehicles (i.e. drones), which have been used to map individual peatlands at a very high resolution (***[***e.g. Knoth et al., 2013***](#_ENREF_37)***) offer considerable potential for more detailed mapping of domestic peatlands at the national scale***.***”***

**Tables and Figures**

The tables and figures are generally of good quality.

Table 1: Please include the WT and the vegetation at the DP sites.

**Response: WT values are already presented in Fig. 1. We have added vegetation to Table 1 as suggested.**

Figure 2: I don't think this figure is really necessary. If you should chose to keep it, please use

percentages instead of absolute counts as due to the different lengths of the study periods the sites are hard to compare by absolute counts. In this case, please add the range of temperatures at which measurements took place.

**Response: We feel that this figure is important as it clearly shows which sites are “extreme” in terms of soil temperature and allows for comparison between the sites. We have made the changes as suggested (i.e. percentages and range of temperatures).**

Overall, the manuscript is clear and well-written, but, in some cases, uses IPCC-related jargon. Therefore, I would suggest to have the manuscript read by a scientist not familiar with National inventories or reporting issues. Similarly, the discussion should focus a bit stronger on those results interesting for scientists not involved with emission reporting.

**Response: We have deleted IPCC jargon text at**

**P7495, L15-16, 19, 20**

**P7514, L15-18**

**We feel that there is a good balance in the discussion as it is. The discussion is composed of seven sections, five of which are devoted to non-emission reporting results.**

**References**

**Alm, J., Shurpali, N. J., Tuittila, E.-S., Laurila, T., Maljanen, M., Saarnio, S., and Minkkinen, K.: Methods for determining emission factors for the use of peat and peatlands - flux measurements and modelling, Boreal Environment Research, 12, 85-100, 2007.**

**Beyer, C. and Höper, H.: Greenhouse gas exchange of rewetted bog peat extraction sites and a Sphagnum cultivation site in northwest Germany, Biogeosciences, 12, 2101-2117, 2015.**

**Beyer, C., Liebersbach, H., and Höper, H.: Multiyear greenhouse gas flux measurements on a temperate fen soil used for cropland or grassland, Journal of Plant Nutrition and Soil Science, 2015. 10.1002/jpln.201300396, 2015.**

**Byrne, K. A., Kiely, G., and Leahy, P.: CO2 fluxes in adjacent new and permanent temperate grasslands, Agricultural and Forest Meteorology, 135, 82-92, 2005.**

**Campbell, D. R., Lavoie, C., and Rochefort, L.: Wind erosion and surface stability in abandoned milled peatlands, Canadian Journal of Soil Science, 82, 85-95, 2002.**

**Knoth, C., Klein, B., Prinz, T., and Kleinebecker, T.: Unmanned aerial vehicles as innovative remote sensing platforms for high-resolution infrared imagery to support restoration monitoring in cut-over bogs, Applied Vegetation Science, 16, 509-517, 2013.**

**Laine, A., Sottocornola, M., Kiely, G., Byrne, K. A., Wilson, D., and Tuittila, E.-S.: Estimating net ecosystem exchange in a patterned ecosystem: Example from blanket bog, Agricultural and Forest Meteorology, 138, 231-243, 2006.**

**Malone, S. and O'Connell, C.: Irelands Peatland Conservation Action Plan 2020 - Halting the loss of biodiversity, Irish Peatland Conservation Council, 152 pp., 2009.**

**Price, J.: Soil moisture, water tension and water table relationships in a managed cutover bog, Journal of Hydrology, 202, 21-32, 1997.**

**Tuittila, E.-S., Komulainen, V.-M., Vasander, H., and Laine, J.: Restored cut-away peatland as a sink for atmospheric CO2, Oecologia, 120, 563 - 574, 1999.**

**Waddington, J. M. and Price, J. S.: Effect of peatland drainage, harvesting and restoration on atmospheric water and carbon exchange, Physical Geography, 21, 433-451, 2000.**

**Waddington, J. M. and Warner, K. D.: Atmospheric CO2 sequestration in restored mined peatlands, Ecoscience, 8, 359-368, 2001.**

**Wind-Mulder, H. L., Rochefort, L., and Vitt, D. H.: Water and peat chemistry comparisons of natural and post-harvested peatlands across Canada and their relevance to peatland restoration, Ecological Engineering, 7, 161-181, 1996.**

**Derivation of Greenhouse Gas emission factors for peatlands managed for extraction in the Republic of Ireland and the United Kingdom**

**D. Wilson1, S.D. Dixon2, R.R.E. Artz3, T. E.L. Smith4, C. D. Evans5, H.J.F. Owen4, E. Archer2 and F. Renou-Wilson6**

1 Earthy Matters Environmental Consultants, Glenvar, Co. Donegal, Ireland. + 353 74 9177613, david.wilson@earthymatters.ie

2 Department of Earth Sciences, University of Durham, Durham, United Kingdom +44 191 334 2356

3 The James Hutton Institute, Aberdeen, Scotland. + 44 844 928 5428

4 King’s College London, Department of Geography, Strand, London, United Kingdom, + 44 207 8482525

5 Centre for Ecology and Hydrology, Bangor, Wales + 44 124 8374500

6 School of Biology & Environmental Science, University College Dublin, Ireland + 353 1 716 2440

**Abstract**

Drained peatlands are significant hotspots of carbon dioxide (CO2) emissions and may also be more vulnerable to fire with its associated gaseous emissions. Under the United Nations Framework Convention on Climate Change (UNFCCC) and the Kyoto Protocol, greenhouse gas (GHG) emissions from peatlands managed for extraction are reported on an annual basis. However, the Tier 1 (default) emission factors (EFs) provided in the IPCC 2013 Wetlands Supplement for this land use category may not be representative in all cases and countries are encouraged to move to higher Tier reporting levels with reduced uncertainty levels based on country or regional specific data. In this study, we quantified (1) CO2-Cemissions from 9 peat extraction sites in the Republic of Ireland and the United Kingdom, which were initially disaggregated by land use type (industrial versus domestic peat extraction), and (2) a range of GHGs that are released to the atmosphere with the burning of peat. Drainage related methane (CH4) and nitrous oxide (N2O) emissions, as well as CO2-Cemissions associated with the off-site decomposition of horticultural peat were not included here. Our results show that net CO2-Cemissions were strongly controlled by soil temperature at the industrial sites (bare peat), and by soil temperature and leaf area index at the vegetated domestic sites. Our derived EFs of 1.70 (±0.47) and 1.64 (±0.44) t CO2-C ha-1 yr-1 for the industrial and domestic sites respectively, are considerably lower than the Tier 1 EF (2.8±1.7 t CO2-C ha-1 yr-1) provided in the Wetlands Supplement. We propose that the difference between our derived values and the Wetlands Supplement value is due to differences in peat quality and, consequently, decomposition rates. Emissions from burning of the peat (g kg-1 dry fuel burned) were estimated to be approximately 1346 (CO2), 8.35 (methane, CH4), 218 (carbon monoxide, CO), 1.53 (ethane,C2H6)), 1.74 (ethylene,C2H4)), 0.60 (methanol,CH3OH)), 2.21 (hydrogen cyanide, HCN) and 0.73 (ammonia, (NH3) and emphasises the importance of understanding the full suite of trace gas emissions from biomass burning. Our results highlight the importance of generating reliable Tier 2 values for different regions and land-use categories. Furthermore, given that the IPCC Tier 1 EF was only based on 20 sites (all from Canada/Fenno-Scandia) we suggest that data from another 9 sites significantly expands the global dataset, as well as adding a new region.

**1 Introduction**

Greenhouse gas (GHG) emissions to the atmosphere have increased significantly since pre-industrial times as a direct result of human activities, such as fossil fuel burning, cement production and land use changes (IPCC, 2013). The Intergovernmental Panel on Climate Change (IPCC) have estimated in their Fifth Assessment Report (AR5) that around one third of all anthropogenic emissions of carbon dioxide (CO2) for the period 1750-2011, were caused by land use changes (IPCC, 2013). From 2000-2009, the Agriculture, Forestry and Other Land-Use (AFOLU) sector accounted for 24% of all global GHG emissions (around 10 Gt CO2-eq yr-1 ), with emissions from peatland drainage and burning alone estimated at around 0.9 Gt CO2-eq yr-1.

Natural (i.e. undrained) peatlands function as long term carbon (C) stores as the sequestration of CO2 over time is greater than the amount of C that is emitted from the peatland as methane (CH4) and leached in waterborne exports (Roulet et al., 2007; Nilsson et al., 2008; Koehler et al., 2011; Gažovič et al., 2013). Key to this role is the position of the water table, which largely dictates the rate of decomposition within the peatland. When the water table is positioned close to the peat surface, the breakdown and degradation of organic matter typically proceeds very slowly in the absence of oxygen. As a consequence, there is an accumulation of peat (and C within) (Dise, 2009).

In the Republic of Ireland (ROI) and the United Kingdom (UK), peat has been extracted for energy use for many centuries (Chapman et al., 2003; Renou et al., 2006). Traditionally, this involved the manual removal of the peat i.e. hand cutting, however this has been largely superseded by highly mechanised methods to extract the peat for both energy and horticulture requirements. In the ROI, over 4 million tonnes of peat per annum are industrially extracted from approximately 50,000 ha to provide ca. 5.5% of primary energy requirements (Howley et al., 2012) and for use in horticulture. A further 0.4 million tonnes per year is likely burned for domestic heating (Duffy et al., 2014) and may impact as much as 600,000 ha of peatlands (Wilson et al., 2013b). Although peat extraction areas in the UK have generally declined over the last few decades, approximately 0.8 million tonnes of peat is still extracted each year in England and Scotland (Webb et al., 2014), although it is UK Government policy to phase out peat extraction in England by 2030 (Department of Environment Food and Rural Affairs, 2011). Peat extraction areas in Wales are small (482 ha) and have remained unchanged in the 1991-2010 period (Webb et al., 2014). In Northern Ireland, the area of peatland utilised for fuel (mechanical and hand cutting) has declined considerably in the 1990-2008 period, although a slight increase in the areas used for horticulture have been recorded (Tomlinson, 2010).

In industrial peatlands, the extraction of peat is facilitated by the installation of drainage ditches at regular (typically 15-30m) intervals across the peatland. For peat used for horticultural purposes, the more fibrous upper layers (e.g. *Sphagnum* peat) are extracted and utilised. If the peat is to be used for energy production the more highly decomposed peat is milled, dried in the production fields and removed for immediate use or stockpiled for later requirements. Peat extraction ceases for energy production when either the sub-peat mineral soil is reached, large quantities of fossilised timber are encountered or drainage is no longer practical (Farrell and Doyle, 2003). For peatlands used for the provision of domestic heating, the peat is either removed by a digger from the margins of peatlands, placed in a tractor mounted hopper and extruded onto the surface of the peatland, or the peat is extruded onto the surface of the peatland from openings made in the peat by a chain cutter. Over a period of weeks the peat is dried *in situ* and removed from the site. The effect of peat extraction on the hydrological functioning is marked by a large fall in the water level either throughout the peatland (industrial) or at the margins of the peatland (domestic). In the latter, significant water level drawdown is also experienced further inward towards the centre of the peatland (Schouten, 2002).

The impact of drainage on C cycling in peatlands has been widely documented. In general, a lowering of the water table leads to increased CO2 emissions (Silvola et al., 1996; Salm et al., 2012; Haddaway et al., 2014) as the aerobic layer is deepened and mineralisation rates are accentuated. Concurrently, CH4 emissions (with the exception of ditches) may decrease or cease (Salm et al., 2012; Turetsky et al., 2014), waterborne C exports may increase (Strack et al., 2008; Evans et al., 2015) and there may be a heightened risk of C loss through fire (Turetsky et al., 2015). In the case of peat extraction, C cycling may be further altered by the removal of vegetation (Waddington and Price, 2000), and losses of windblown particulate organic carbon (POC) may be exacerbated from the bare peat surfaces (Lindsay, 2010).

Under the United Nations Framework Convention on Climate Change (UNFCCC) and the Kyoto Protocol, “Annex 1” countries (i.e. countries that have committed to targets that limit or reduce emissions) are obligated to prepare annual National Inventory Reports (NIR) and up-to-date annual inventories, detailing GHG emissions and removals from six different sectors. Emissions associated with off-site peat combustion are reported under the Energy sector and are not considered further here. The recent IPCC Wetlands Supplement (IPCC, 2014) to the 2006 Good Practice Guidance (GPG) (IPCC, 2006) derived new Tier 1 emission factors (EFs) for drained organic soils that differentiated between on-site emissions (e.g. CO2-Con-site, fire) and off-site losses ( e.g. leaching of waterborne C). In the case of peatlands managed for extraction in the temperate climate zone, the CO2-Con-site values have increased from 0.2 (nutrient poor/bogs) and 1.1 (nutrient rich/fens) t CO2-C ha-1 yr-1 in the 2006 GPG to a single higher EF of 2.8 t CO2-C ha-1 yr-1 (covering the entire boreal and temperate regions) in the Wetlands Supplement. On-site burning directly consumes aboveground C stocks (prescribed and wildfire burning) and the underlying peat C store (wildfire burning), and rapidly releases both gases (e.g. CO2, CH4) and particulates (e.g. black carbon) to the atmosphere. In the Wetlands Supplement, an EF for GHG emissions from prescribed fire on drained peatlands is not provided due to a paucity of published data at present. However, emissions from wildfires are addressed and EFs of 362, 9 and 207 g kg-1 dry fuel burned is provided for CO2-C, CH4 and CO respectively with a proviso that they were derived from a very small dataset.

Given the relatively large areas under peat extraction in both the ROI and the UK, a move from Tier 1 to higher reporting levels is desirable, particularly as (a) a wide range in uncertainty is associated with the IPCC Tier 1 values (1.1 - 4.2 t CO2-C ha-1 yr-1), which reflects the disparity in emissions from drained peatlands from different climate zones and nutrient composition, (b) the most recently published annual CO2 flux estimates (not included in the derivation of IPCC Tier 1 values) also display a very wide amplitude (cf. Järveoja et al., 2012; Mander et al., 2012; Salm et al., 2012; Strack et al., 2014), (c) no data from ROI or UK peatlands were included in the IPCC derivation, which might mean that the Tier 1 value may not be appropriate for these countries, and (d) no distinction is made between industrial or domestic extraction sites, despite large differences in their drainage, vegetation cover and management characteristics. In addition, previous studies of peatland fire EFs have focused on the boreal peatlands of Alaska (Yokelson et al., 1997) and Canada (Stockwell et al., 2014); and the temperate peatlands of Minnesota (Yokelson et al., 1997) and North Carolina (Stockwell et al., 2014). These studies found that the smouldering combustion of peats associated with low combustion efficiency leads to relatively lower CO2 emissions (compared with other ecosystems), and much higher carbon monoxide (CO), CH4, and other non- CH4 hydrocarbon emissions. Therefore, it is important to quantify emissions of these gases as they include strong GHGs (e.g. CH4) and reactive gases responsible for tropospheric ozone formation and poor air quality (e.g. CO, ammonia (NH3), hydrogen cyanide (HCN)).

The objectives of the study are (1) to provide estimates of the annual CO2-C exchange (i.e. CO2-Con-site) for 9 peat extraction sites in the ROI and the UK, (2) to derive regional specific CO2-C EFs for drained peat extraction areas that would permit ROI and the UK to progress to the Tier 2 reporting level, (3) analyse the factors that influence CO2-C dynamics in this region (i.e. land use, climate etc.), and (4) to report GHG emissions associated with the burning of Irish *Sphagnum* moss peat in the first laboratory study to investigate fire emissions from European temperate peats.

**2 Materials and Methods**

**2.1 Study sites**

The study sites were located at 9 peat extraction areas in the ROI and the UK with a history of either industrial peat (IP) or domestic peat (DP) extraction (Table 1). Boora (IP1), Blackwater (IP2), Bellacorick (IP3), Turraun (IP4), Middlemuir Moss (IP5) and Little Woolden Hall Moss (IP6) are industrial cutaway peatlands where significant areas of bare peat (i.e. unvegetated microsites) have remained following the cessation of milled peat extraction. At IP6, milled peat is currently extracted from areas close (<150m) to the study site. The IP sites are former raised bogs with the exception of IP3, which is a former Atlantic blanket bog. At all sites, the drainage ditches have remained functional. Here we define “drained” as a mean annual water table position deeper than -20cm (Couwenberg and Fritze, 2012; Strack et al., 2014). Physico-chemical characteristics of all the sites are detailed in Table 1.

At Clara (DP1), Glenlahan (DP2) and Moyarwood (DP3) the peat has been extracted from the margins of the sites for use in domestic heating. In the case of Clara, peat extraction was an ongoing activity at the time of our study despite the designation of the site as a Special Area of Conservation (SAC). DP1 and DP3 are raised bogs and DP2 is a mountain blanket bog. The vegetation component at all the sites is species poor and is composed mainly of ling heather (*Calluna vulgaris*), cross leaved heather (*Erica tetralix*) and lichens (*Cladonia* spp.) A continuous water table level was not observed at DP2, as the relatively shallow peat deposit (~40cm) over bedrock at that site was prone to drying out at various times throughout the study.

**2.2 Climatic conditions**

All the sites are located within the temperate zone as defined by IPCC (2006), and are characterised by an oceanic climate with prevailing south-west winds, mild mean annual air temperatures (8 to 10.3°C) and moderate to high annual rainfall (804 to 1245 mm) (Table 1).

**2.3 Environmental monitoring**

At each site, 3-9 aluminium square collars (60 x 60 cm) were inserted to a depth of 30cm into the peat. At IP6, smaller circular plastic collars were used (15cm diameter) to facilitate the use of the CPY-4 chamber (PP Systems, UK) at that site. Soil loggers (µlogger; Zeta-tec, UK, Hobo External Data Loggers; Onset Computer Corporation, MA, USA or Comark N2012 Diligence Loggers, Norwich, UK) were established in all the IP sites and recorded soil temperatures (°C) at hourly intervals. Weather stations were installed at all the DP sites and recorded photosynthetic photon flux density (PPFD; µmol m-2 s-1) and soil temperatures (5 and 10cm depths) at 10 minute intervals. At DP3, soil volumetric moisture content (VMC, %) was also recorded (at 10 min intervals) by the weather station at that site. At sites IP5 and IP6, soil temperature was only measured manually during CO2 flux measurements. In order to estimate soil temperature at times where data was lacking at these two sites, a regression based approach between manually recorded T5cm and air temperature recorded at 15 min intervals by a logger on the site was used to gap fill the data (r2 = 88.7 %). Water table level (WT) was manually measured from dipwells (internal diameter 2 cm) inserted adjacent to each collar. Wooden boardwalks were established at each site (exception IP6).

**2.4 Leaf area index (LAI)**

At the IP sites, the vegetation had been removed prior to the commencement of peat extraction and virtually no natural recolonization has taken place following cessation of peat extraction. However, at the DP sites a vegetation component was present and in order to incorporate the seasonal dynamics of the plants into CO2-C exchange models, the leaf area index (LAI) was estimated for each of the collars. This involved accounting for the green photosynthetic area of all vascular plants (leaves and stems) within the collar at monthly intervals. In short, the number of leaves and stems were counted from five subplots (8 x 8cm) within each collar. The size (length, width) of the leaves was measured from sample plants outside the collars. The LAI was then calculated by multiplying the estimated number of leaves by an area estimate of the leaf. Moss and lichen % cover was estimated at the same time. Species-specific model curves were applied to describe the phenological dynamics of the vegetation of each collar, and the models (vascular plants and moss) were summed to produce a plot-specific LAI. For a detailed description of the method see Wilson et al. (2007). At site DP1 only, the vegetation was removed by regular clipping from one third of the collars, in order to provide an estimate of the heterotrophic contribution (RH) to ecosystem respiration (Reco).

**2.5 On site carbon dioxide flux estimation**

**2.5.1 Field measurements**

At sites IP1-5 and DP1-3, Reco was measured with a static polycarbonate chamber (60 x 60 x 33 cm) equipped with two internal fans to ensure mixing of the air within the chamber, and a cooling system (submerged ice packs, and pumped water to a radiator located within the chamber) to maintain the temperature within the chamber close to the ambient air temperature (for a more detailed description see Alm et al., 2007b). At IP6, Reco was measured with a CPY-4 (PP Systems, UK) clear acrylic chamber (14.6 cm diameter, 14.5 cm height). The CPY-4 chamber was equipped with an internal fan, PPFD sensor and thermistor. Sampling was carried out at fortnightly or monthly (winter) intervals (2-4 measurements per collar per measurement day). For each Reco flux measurement, the chamber was placed in a water-filled channel at the top of the collar or connected with a rubber gasket (IP5), covered with an opaque cover and the CO2 concentration (ppmv) in the chamber headspace was measured at 15-second (5-second at IP6) intervals over a period of 60-180 seconds using a portable CO2 analyser (EGM-4; PP Systems, UK). Concurrently, air temperature (°C) within the chamber and soil temperatures at 5, 10 and 20 cm depths were recorded at each collar (soil temperature probe; ELE International, UK). The WT position relative to the soil surface was manually measured with a water level probe (Eijkelkamp Agrisearch Equipment, The Netherlands). At the DP sites, net ecosystem exchange (NEE) was measured with the same polycarbonate chambers described above under a range of ambient light levels (PPFD; µmol m-2 s-1) prior to Reco measurements. NEE measurements were carried out between 8 am and 6pm in the summer and between 9am and 3pm in the winter (3 to 8 measurements per collar per measurement day) to ensure that the maximum PPFD was reached at each measurement date. Artificial shading was used in the early morning to obtain low PPFD levels (<100 µmol m-2 s-1). PPFD was recorded from a sensor (PAR-1. PP Systems) located within the chamber. The portable CO2 analysers were regularly calibrated with a CO2 standard gas.

**2.5.2 Flux calculations**

Flux rates (mg CO2-Cm-2 h-1) were calculated as the linear slope of the CO2 concentration in the chamber headspace over time, with respect to the chamber volume, collar area and air temperature. A flux was accepted if the coefficient of determination (*r2*) was at least 0.90. An exception was made in cases where the flux was close to zero (mainly in winter time where soil processes are typically slower) and the *r2* is always low (Alm et al., 2007b). In these cases the flux data were examined graphically and fluxes with obvious non-linearity (due to chamber leakage, fan malfunction etc.) were discarded. The remainder were accepted provided that some of the environmental variables measured at the same time (e.g. soil temperature) were sufficiently low to account for the low flux values (Wilson et al., 2013a). In this study, we followed the sign convention whereby positive values indicated a CO2-C flux from the peatland to the atmosphere (source) and negative values indicated a flux from the atmosphere to the peatland (sink). Gross primary production (GPP) was calculated as NEE minus Reco (Alm et al., 2007b), and the closest Reco flux value in time to a NEE flux value was used.

**2.5.3 Modelling**

Statistical and physiological response models (Alm et al., 2007b) were constructed and parameterised for each study site. Model evaluation was based on the following criteria; (a) statistically significant model parameters (p<0.05), (b) lowest possible standard error of the model parameters and (c) highest possible coefficient of determination (adjusted r2) (see Laine et al., 2009). The basic Reco models, based upon the Arrhenius equation (Lloyd and Taylor, 1994), are non-linear models related to soil temperature. GPP was related to PPFD using the Michaelis–Menten type relationship that describes the saturating response of photosynthesis to light (Tuittila et al., 1999). GPP model coefficients and associated standard errors were estimated using the Levenberg-Marquardt multiple non-linear regression technique (IBM SPSS Statistics for Windows, Version 21.0. Armonk, NY, USA). During model construction, the relationship between Reco or GPP and a range of independent environmental variables (recorded in conjunction with flux measurements) was tested. Only variables that increased the explanatory power of the model (i.e. improved r2 values) were included. The models were accepted if the residuals were evenly scattered around zero.

**2.5.4 Annual CO2 –C balance**

The response functions estimated for Reco and GPP were used for the reconstruction of the annual CO2 -C balance. Reco fluxes were reconstructed for each collar in combination with an hourly time series of (1) T5cm, (2) VMC (at DP3) recorded by the data loggers or (3) WT depths linearly interpolated from weekly measurements. The annual CO2-C balance (g C m-2 yr-1) was calculated for each sample plot by integrating thehourly Reco values over each 12-month period. (Note: integration periods vary between study sites; see Table 1). At the DP sites, GPP was reconstructed in combination with (1) PPFD values recorded by the weather station, (2) plot specific modelled LAI and (3) an hourly time series of T5cm (DP1only). At the DP sites, annual NEE was calculated as annual GPP + annual Reco.

**2.5.5 Statistical analysis**

The CO2-C flux data (Reco for the IP sites, and Reco and GPP for the DP sites) had a non-normal distribution, so the non-parametric Kruskal-Wallis (*p=0.05*) and Mann-Whitney tests were used to test for differences between sites. Uncertainty in reconstructed annual *R*eco and GPP was calculated by summing up the maximum and minimum standard errors associated with each of the model parameters (e.g. Drösler, 2005; Elsgaard et al., 2012; Renou-Wilson et al., 2014). Uncertainty in the annual Reco or NEE estimate was calculated following the law of error propagation as the square root of the sum of the squared standard errors of GPP and Reco (IPCC, 2006).

**2.6 Peat fire emissions**

Around 5 kg (dry mass) of loose Irish *Sphagnum* moss peat (H2-H3 on the von post decomposition scale) was used for measuring fire EFs. Subsamples of the peat were taken and placed into a 22 x 12 x 10 cm open-topped insulated chamber. The chamber was constructed from lightweight Celcon insulation blocks and was used to replicate natural surface combustion conditions, leaving only one surface of the peat exposed to open air thereby reducing heat loss and oxygen exchange from the other surfaces, in accordance with the suggested peat combustion methodology of Rein et al. (2009). Each sample was dried in an oven overnight at 60°C. In order to produce comparable replicates, the samples for the burning experiment had to be dried to an absolute dry base to increase ignition probability (Frandsen, 1997) and encourage pyrolysis (Rein et al., 2009). Following drying, the chamber and sample were placed in a fume cupboard under controlled air flow conditions and the peat was ignited using a coiled nichrome wire heated to ~600°C and placed in contact with the surface of the peat. This also best represents natural ignition conditions (e.g. from a surface shrub fire), also in accordance with the methodology of Rein et al. (2009). Once ignited, each 1 kg sample proceeded to burn for ~90 minutes. The resulting smoke was continuously sampled using a pump and a 90 cm sample line with a funnel held ~12 cm above the smouldering peat. The smoke was sampled into an 8.5 litre infrared White (multipass) cell (Infrared Analysis, Inc.) where infrared spectra were collected using a Fourier Transform Infrared (FTIR) spectrometer. Analysis of the FTIR spectra was performed using the Multi Atmospheric Layer Transmission (MALT) software (Griffith, 1996), yielding trace gas mole fractions inside the White cell, from which emissions factors may be calculated. A full description of how EFs may be calculated from FTIR measurements of gas mole fractions is given in Paton-Walsh et al. (2014) and Smith et al. (2014). Here we use the C mass balance approach to calculate EFs for CO2 and CO (Eq. 1 in Paton-Walsh et al., 2014). The C content of the peat (required for calculating EFs via the C mass balance approach) is assumed to be 53.3%, as measured in Scottish sphagnum moss peat (Cancellieri et al., 2012). For all other gas species considered in the study; CH4, ethylene (C2H4), ethane (C2H6), methanol (CH3OH), HCN, NH3), we use their respective emission ratios to CO and the EF for CO to calculate EFs (via Eq. 5 in Paton-Walsh et al., 2014).

Combustion efficiency is a measure of the amount of fuel carbon released as CO2, and may be approximated using the Modified Combustion Efficiency (MCE) formula, which requires only a measurement of CO and CO2 rather than all the C containing gases (Yokelson et al., 2008):

$MCE= \frac{∆CO\_{2}}{∆CO\_{2}+ ∆CO}$ (1)

where ΔCO2 and ΔCO represent the elevated mixing ratios of these gases (the difference between mixing ratios measured in biomass burn emissions and those in the ambient air). MCE is often expressed as a percentage. Generally, an MCE lower than 0.9 (90%) is considered a low combustion efficiency burn (Lobert et al., 1991; Yokelson et al., 1996).

**3 Results**

**3.1 Environmental variables**

Annual rainfall varied between sites and between years (Fig. 1). The wettest site was DP3 (1390 mm), and the driest was IP6 (746 mm) in the first year of measurements at that site. All multi-year sites displayed inter-annual variation in rainfall with the largest differences observed in IP4 (210 mm difference in annual rainfall between years). Annual rainfall at IP2, IP5, DP1, DP2 and DP3 was above the long-term average in all years. IP1 and IP4 were wetter than the long-term average in one of the years and drier in the other. IP3 and IP6 were drier than the long-term average. The mean annual water table was below -20cm at all sites in all years (Fig. 1). The deepest mean annual values were at IP1 (-60cm) and the shallowest at IP3, 4 and 5 (-25cm). Mean water table position tracked annual rainfall (i.e. higher rainfall resulted in higher water table positions) in all multi-year sites with the exception of IP1.

The highest mean annual soil temperature (T5cm) value (12.7°C) was recorded at IP4 and the lowest at IP5 (6.7°C) and inter-annual variation was evident in the multi-year sites (Fig. 1). The lowest hourly T5cm value (-12.9°C) was recorded at IP5 and the highest (28.4°C) at IP4 (Fig. 2). The proportion of hourly T5cm values less than 0°C ranged from 0% (IP3) to 13.8% (IP5), and the proportion of values greater than 20°C ranged from 0.2% (IP5) to 5.3% (IP2) (Fig.2).

**3.2 On-site carbon dioxide fluxes**

At the IP sites, Reco fluxes ranged from 0 to 133 mg CO2-C m-2 hr-1 and differed significantly between sites (Fig. 3a Kruskal-Wallis, *H*=98.59). Site IP4 had significantly higher Reco flux values than all the other IP sites (Mann Whitney p<0.001) and IP5 had significantly lower flux values than IP2, IP4 and IP6 (Mann Whitney *p*<0.001) but not IP1 and IP3 (Mann Whitney *p*=0.31). At the DP sites, Reco fluxes ranged from 12 to 200 mg CO2-C m-2 hr-1 and there was a significant difference in Reco fluxes between the DP sites (Fig. 3b Kruskal-Wallis, *H*=37.52) but no significant difference between DP1 and DP2 (Mann Whitney *p*=0.075). Reco values differed significantly between the IP and DP sites (Kruskal-Wallis, *H*=395.22). Measured NEE (at PPFD>1000 µmol m-2 s-1) ranged from 60 to -325 mg CO2-C m-2 hr-1 at the DP sites and values differed significantly between sites (Fig. 3c Kruskal-Wallis, *H*=90.82).

**3.2.1 Modelling**

At sites IP6 and DP2, T5cm was the sole explanatory variable in the Reco models (Eq.2) and explained 32% and 42% respectively of the variability in fluxes. The addition of water table to the Reco model (Eq.3) slightly improved the explanatory power and the model explained between 55% and 85% of the variability at IP1-4 and 69% at DP1. No relationship between Reco and WT was observed at DP3, but the addition of VMC (Eq. 4) also slightly improved the explanatory power of the model (78%). At IP5, the data were too limited (n=22) to construct a reliable model that satisfied the criteria outlined in section 2.5.3. Instead, we calculated monthly mean values and integrated these values over the 12 month study period.

$R\_{eco }= a\*exp\*\left[b\left(\frac{1}{T\_{REF-}T\_{0}}- \frac{1}{T- T\_{0}}\right)\right]$ (2)

$R\_{eco }= a\*exp\*\left[b\left(\frac{1}{T\_{REF-}T\_{0}}- \frac{1}{T- T\_{0}}\right)\right]\* WT$ (3)

$R\_{eco }= a\*exp\*\left[b\left(\frac{1}{T\_{REF-}T\_{0}}- \frac{1}{T- T\_{0}}\right)\right]\*VMC$ (4)

where Reco is ecosystem respiration, *TREF* is reference temperature set at 283.15 K, *T0* is the (minimum) temperature at which respiration reaches zero and is set here at 227.13 K, *T* is the soil temperature at 5 cm depth, WT is water table depth, VMC is volumetric moisture content, *a* and *b* are fitted model parameters.

A strong relationship was observed between GPP and PPFD at the DP sites. It was the sole explaining variable at DP2 (Eq. 5) where it accounted for 70% of the variation. The addition of LAI (Eq. 6) increased the explanatory power of the GPP model at DP3 (59%) and the addition of LAI and T5cm resulted in 62% of the variation explained at DP1.

$GPP= P\_{max }\left(\frac{PPFD}{PPFD+ k\_{PPFD}}\right)$ (5)

$GPP= P\_{max }\left(\frac{PPFD}{PPFD+ k\_{PPFD}}\right)\*LAI$ (6)

$GPP= P\_{max }\left(\frac{PPFD}{PPFD+ k\_{PPFD}}\right)\*LAI\*T5cm$ (7)

where*GPP*  is gross primary productivity, *Pmax* is maximum photosynthesis, PPFD is photosynthetic photon flux density, *kPPFD* is the PPFD value at which *GPP* reaches half its maximum, LAI is leaf area index, T5cm is soil temperature at depth of 5 cm.

**3.2.2 Annual CO2-C balance**

The annual CO2-C balance varied both spatially (between sites) and temporally (multi-year sites) (Figs. 4 and 5). In the IP sites, emissions ranged from 93 g CO2-C m-2 yr-1 (IP5) to 304 g CO2-C m-2 yr-1 (IP4). Annual emissions varied considerably within the multi-year sites, where coefficient of variation values ranged from 4% (IP1) to 20% (IP2). As would be expected given the close relationship observed between soil temperature and CO2-C fluxes, a noticeable increase in modelled CO2-C emissions was observed during the summer months at all sites (Fig. 4), although the rate of the increase varied somewhat in strength between years in the multi-year sites as a function of measured T5cm and WT (where applicable). In the DP sites (Fig. 5), annual GPP and Reco were highest in DP1 (-526 and 702 g CO2-C m-2 yr-1 respectively), intermediate in DP2 (-484 and 687 g CO2-C m-2 yr-1 respectively) and lowest in DP3 (-319 and 434 g CO2-C m-2 yr-1 respectively). The DP sites were a net annual CO2-C source with the highest emissions observed at DP2 (203 g CO2-C m-2 yr-1), intermediate at DP1 (176 g CO2-C m-2 yr-1) and lowest at DP3 (114 g CO2-C m-2 yr-1). Estimated emissions from heterotrophic respiration (RH) at DP1 were 344 g CO2-C m-2 yr-1, which equates to 49% of Reco at that site. Applying this proportional value to the other DP sites, we estimate that RH emissions to be 337 and 213 g CO2-C m-2 yr-1 at DP2 and DP3 respectively.

**3.3 Drivers of annual CO2-Con site**

No relationships were observed between annual CO2-C balances (NEE) and nutrient concentrations, water table levels (average, maximum or minimum) or the von Post scale at either the IP or DP (p>0.05) sites. A strong relationship (r2=0.63) between average soil temperature at 5cm depth and Reco was very evident across the IP sites (Fig. 6); the highest annual emissions and highest average soil temperatures were associated with IP4 and the lowest at IP5. The variation in NEE between the DP sites appeared to be related to differences in LAI (Fig. 6), however the number of sites was very small (n=3) and some caution must be used in this regard.

**3.4 Emission factors**

Using a single mean value for each multi-year site and for its associated uncertainty (IPCC, 2014), an EF was calculated for each land use category. The derived EFs for the IP and DP sites were 1.70 (±0.47) and 1.64 (±0.44) t CO2-C ha-1 yr-1 respectively (Table 2). The 95% confidence intervals associated with the derived EFs were ±28% and ±26% for the IP and the DP sites respectively. There was no significant difference in the EF values between the IP and DP sites (*p=0.90*).

**3.5 Peat fire emission factors**

Mean modified combustion efficiency (MCE) and EFs with their standard deviations for eight trace gas species were calculated from measurements of five Irish sphagnum moss peat samples (Table 3). The peat burned with a mean MCE of 0.837 (±0.019) typical of smouldering combustion (e.g. Yokelson et al., 1996; Bertschi et al., 2003). Emissions of CO2 amounted to 1,346 (±31) g CO2 kg-1 of dry fuel burned or 342 (±8) g CO2-C. Other carbonaceous emissions amounted to 218 g CO kg‑1; 8.35 g CH4 kg-1; 1.74 g C2H4 kg-1; 1.53 g C2H6 kg-1; and 0.60 g CH3OH kg-1 of dry fuel burned. Emissions of the nitrogenous compounds amounted to 2.21 g HCN kg-1; and 0.73 g NH3 kg-1.

**4 Discussion**

There is a very wide range in reported CO2 emissions from both active and abandoned peat extraction areas in the literature (Figure 7). Much of this variation can be attributed to differences in climate, drainage level, peat type, peat extraction methods and the end use of the peat and, as such, provides a useful framework to examine the variations in this study.

**4.1 Effects of climate**

While the study sites in this paper are all located within the temperate zone, considerable variation in CO2-C emissions was evident. Given that all the sites are drained to a similar depth (Fig. 1), it is not surprising that the variation in emissions appeared to be controlled largely by differences in soil temperatures between the sites (Fig. 6). The coldest site in terms of mean soil temperatures and lowest in terms of annual emissions was Muirhead Moss (IP5) in North-Eastern Scotland. Although rainfall and site water table levels were similar to the other sites, soil temperatures at this site remained below 0°C for a high proportion (~14%) of the year, and are likely to have resulted in a slowdown of extracellular enzymatic diffusion (Davidson and Janssens, 2006), reduced microbial activity (Fenner et al., 2005) and consequently lower rates of CO2 production (Basiliko et al., 2007). Indeed, it is likely that our value of 0.93 t CO2-C ha-1 yr-1 at this site may be an overestimation given that it was calculated from monthly mean values that were measured during day time hours (highest daily temperatures). As much of the peatlands in Scotland fall within the same temperature regime (Chapman and Thurlow, 1998), CO2-C emissions data from a wider range of peat extraction sites in this region might significantly refine our EF derivation.

At the other end of the spectrum, the highest emissions and soil temperatures were observed at Turraun (IP4) in the Irish Midlands. Data from this site had been previously reported by Wilson et al. (2007). In this study, we only utilised CO2-C flux data from plots where the mean annual water table position was deeper than -20cm. This resulted in a higher mean value (taken over two years) in this current study. Three of IP sites in the ROI are located in the Midlands where more “extremes” in climate are generally experienced (lower winter temperatures, higher summer temperatures) than along the Western coast (IP3). However, during this study, winter temperatures at all the ROI sites seldom decreased below 0°C (Fig. 3) and the proportion of hourly temperatures higher than 20°C were somewhat similar between the sites. Although, Little Woolden Moss (IP6) received the lowest annual rainfall of all sites in year 1 of the study at that site (Fig.1), mean annual soil temperatures were in the mid-range of the 9 study sites, hourly T5cm values were normally distributed (Fig.3) and CO2-Con site emissions were close to the derived EF value of 1.70 t CO2-C ha-1 yr-1 (Table 2).

The DP sites are all located in the ROI and within a 35km radius, but considerable variation in annual rainfall was apparent during this study (Fig. 1), with DP3 (the furthest west) receiving the highest rainfall of all sites in the study (on average 34% more rainfall than the other DP sites). The east-west rainfall gradient in the ROI is well documented and coincides with a change in peatland types (i.e. raised bogs to Atlantic blanket bogs). This climatic variation is reflected in the annual Reco values, which were similar between DP1 and DP2 but much lower in DP3 (Fig. 5). There is an established relationship between rainfall amount and the moisture content of peat (Price and Schlotzhauer, 1999; Strack and Price, 2009). For the sites located in high rainfall areas, such as DP3, there may be a suppression of aerobic microbial activity within the peat matrix, and as a consequence Reco values may be lower than would be expected for a drained peat soil. Indeed, at some of these sites, occult precipitation (e.g. dew and fog droplets) may also contribute significantly to higher levels of soil moisture (Lindsay et al., 2014). During the growing season, the transpiration process is also likely to play a role in determining the moisture content of the peat within the rooting zone (~20cm depth) at these vegetated sites. Moisture losses are likely to be accentuated on sunny days when air and soil temperatures are high, when LAI values are highest (mid-summer) and when vapour pressure deficit is not a limiting factor. As CO2 emissions were closely correlated to soil temperature at 5 cm depth, reduced moisture content in this zone is likely to stimulate aerobic microbial activity. Annual GPP showed a similar trend to annual Reco in these vegetated DP sites. GPP is strongly controlled by the amount of light received by the plants (i.e. PPFD levels and LAI) and the efficiency with which the plants use it. PPFD values (data not shown) and the vegetation composition were broadly similar during the sampling periods, which would seem to indicate that LAI is the driver of both productivity and therefore NEE at these sites (Fig. 6). However, variations in LAI are likely to be the result of subtle differences in a number of other variables (e.g. nutrient status, site management) that were not captured in our measurements.

**4.2 Effects of drainage level**

While a close relationship between WT position and CO2-C emissions has been established in some peatland studies (Silvola et al., 1996; Blodau and Moore, 2003; Blodau et al., 2004), soil temperature proved to be the strongest determinant of CO2-Con- site emissions at our sites and this relationship has also been observed by other studies in peat extraction areas (e.g. Shurpali et al., 2008; Mander et al., 2012; Salm et al., 2012)**.** While the addition of WT or VMC improved the performance of the Reco models at some of the sites, the improvement was only slight and this is likely due to the fairly narrow range of WT/VMC values recorded over the course of the 12 month study (e.g. the range in VMC values at DP3 only ranged between 56-64%). Therefore, optimum WT /VMC levels for respiration may not have been encountered. The Reco models used here are only valid for the data that was measured over the course of the study at each site and cannot be readily extrapolated beyond the range of that data. For those sites where water table did not appear to influence Reco dynamics it may be that fluctuations in WT level were missed with the interpolation approach and CO2-C flux measurement regimes that we employed here, although these methodologies have been widely used elsewhere (Riutta et al., 2007; Soini et al., 2010; Renou-Wilson et al., 2014). Instead, it is probable that our results reflect the complexity of the relationship between Reco and WT in very dry soils as outlined by Lafleur et al.(2005), where factors such as a stable, low surface soil moisture content, and decreased porosity (i.e. limited oxygen availability) at the depths that the WT is mainly located, ensure that when CO2-C fluxes are measured, the WT is deeper than the zone where it has a discernible impact on Reco (Juszczak et al., 2013). As such, the soil temperature regime in these sites may act as a “proxy” for drainage level (i.e. higher soil temperatures are likely to occur in conjunction with deeper water table levels and vice versa) (Mäkiranta et al., 2009).

**4.3 Peat characteristics**

Industrial peat extraction involves the removal of surface vegetation and results in the exposure of decomposed peat at the surface. The level of decomposition in the peat is related to depth and as extraction proceeds, the more highly decomposed peat is exposed. The peat in industrial extraction sites tends to have a lower aerobic CO2 production potential than natural sites for example, due to differences in substrate and nutrient availability, a more extreme physical environment (Glatzel et al., 2004) and reduced labile organic matter supply in the absence of plant communities (i.e. priming). In our study, the C content (with the exception of DP2) was similar across all sites (Table 1). Although, Glatzel et al. (2004) noted that CO2 production was negatively correlated with the von Post scale of decomposition, no correlation with annual CO2-C emissions was evident in our study (p>0.05). Similarly, despite obvious difference in nitrogen content and pH values between IP sites, no relationships with CO2 fluxes were discerned. However, the residual peat at IP4 is strongly influenced by the close proximity of limestone parent material, as evidenced by high pH values and the lowest C:N ratio (Table 1), and is highly minerotrophic. Given the high CO2-C emissions associated with this site, consideration should be given to disaggregation by nutrient type should more data become available in the future.

Organic matter quality has been closely linked to the soil respiration rate, with lower emission rates associated with the poorer quality organic matter found at depth in drained peatlands (Leifeld et al., 2012). The lowest emissions at our sites occurred where the residual peat was either of Cyperaceous (IP3) or *Sphagnum* / Cyperaceous (IP5) origin. However, while the slow decomposition rate of *Sphagnum* litter in comparison to other plant litter has been well documented (Verhoeven and Toth, 1995; Bragazza et al., 2007), there is insufficient data from our study sites to determine whether the limited relationship observed here between peat type and CO2-C emissions in our study sites is coincidental rather than causal.

**4.4 Effects of peat extraction methods and peat end use**

For peat utilised for horticulture, the more fibrous peat layers nearer the surface are extracted. This may result in the oxidation of more labile organic matter and may account for the very high emissions associated with Canadian peatlands for example (Fig. 7) in comparison to countries where the deeper peat layers are extracted (Mander et al., 2012). However, the IP sites in this study are highly decomposed peat and have been abandoned for 30 years or more in some cases (e.g. IP4) and have remained unvegetated. It is possible that CO2-C emissions from active extraction areas may be higher than those derived in this study given that over the summer period the surface of the peat is regularly scarified and aerated. However, Salm et al. (2012) reported higher emissions from abandoned areas in comparison to active areas, although colonisation by vegetation in the former may have accentuated respiration losses. High annual CO2-C emissions following abandonment and recolonization have also been reported by Strack and Zuback (2013) and are in close agreement with the Reco values reported here for the DP sites (Fig. 5).

We have estimated the contribution of heterotrophic respiration (RH) to Reco at 49%. Although, this is based on measurements at a single site (DP1), it is within the range reported by other studies (Frolking et al., 2002; Moore et al., 2002; Shurpali et al., 2008). The RH values measured at DP1 (Fig.5) and estimated at DP2 are higher than the Reco values at the IP sites, which would indicate that decomposition of the belowground biomass (following clipping) and subsequent “priming” effects may contribute significantly to CO2-C dynamics at vegetated extraction sites. Furthermore, the methods employed to extract the peat at some of the DP sites (the peat is extruded onto the surface of the peatland from narrow openings made in the peat by a chain cutter) has led to the formation of deep fissures (ca. 4 cm wide and > 2m deep) within the peat that may enhance oxidation throughout the peat profile. Nonetheless, fissures (ca. 10 cm wide and > 1m deep) formed in the peat during climatically dry years and that were partially filled in during wetter/windier years were also observed at IP5 where the lowest annual emissions were observed.

**4.5 Fire emission factors**

The mean MCE reported here (0.837) is typical of smouldering combustion (e.g. Yokelson et al., 1996; Bertschi et al., 2003) and comparable with the reported range of MCE in other studies of high latitude peats (Yokelson et al., 1997; Stockwell et al., 2014). Emission factors for CO2 and CO are also typical of smouldering combustion and similar to those from other peat studies, particularly Yokelson et al. (1997). As found in other studies of peat fire emissions, our measurements confirm that the CH4 EF for Irish peat is particularly high (8.35 g kg-1 dry fuel burned) when compared with other forms of biomass burning. Given the high Global Warming Potential, where each gram of emitted CH4 is equivalent to 34 g of CO2 (100 year time horizon, IPCC, 2013), the CH4 emissions from Irish peat fires may account for over 12% of the CO2-equivalent emissions. This result emphasises the importance of understanding the full suite of trace gas emissions from biomass burning, rather than focussing solely on CO2 and CH4 emissions. In general, the other EFs reported here lie within the range of variability observed by other peat burning studies, with the exception of NH3, which is particularly low, possibly as a result of the nitrogen-poor soils that are typical of Irish and UK blanket bogs. Here, we also report the first C2H6 EF for peat (1.53 ± 0.17 g kg-1 dry fuel burned), similar in magnitude to C2H6 emissions from boreal forests (1.77 g kg-1 dry fuel burned) according to Akagi et al. (2011). However, the use of prescribed fire in the UK to burn off old heather growth to encourage new growth (e.g. the muirburn practice) may not impact the underlying peat to any great extent, given that the practice is restricted to the October-April period when soil moisture conditions are highest. Emissions result from the burning of the woody aboveground biomass, and the underlying peat is generally unaffected. In contrast, wildfires typically occur during the summer months when temperatures are highest and moisture levels are low, resulting in burning of both the vegetation and the peat itself. Indeed, recent work by Kettridge et al. (2015) has highlighted the vulnerability of drained peatlands, even at high latitudes, to increased risk of wildfire and subsequent vegetation changes.

**4.6 Implications for National Inventory reporting**

The ROI currently employs the 2006 GPG default value of 0.2 t CO2-C ha-1 (nutrient poor) in reporting of all peat extraction areas, and estimated emissions for 2012 (the most recent assessment year) were 9,312 t CO2-C yr-1 (Table 4). In contrast, the approach in the UK has been to differentiate between peat extracted for fuel and horticulture and then applying the default EFs for nutrient rich (1.1 t CO2-C ha-1) and nutrient poor peat (0.2 t CO2-C ha-1) respectively. For 2012, CO2-C emissions from UK extraction peatlands were estimated at 2,118 t CO2-C yr-1 (Table 4).

Reported annual emissions are likely to increase considerably if the Tier 1 values in the IPCC Wetlands Supplement are adopted by inventory compilers. We estimate that emissions from peatlands managed for extraction will be approximately 16 and 10 times higher for the ROI and UK respectively (Table 4). The EFs derived in this study for CO2-Con site for both industrial and domestic peatlands (Table 2) are considerably lower than the Tier 1 value of 2.8 tonnes CO2-C ha-1 yr-1 provided in the IPCC Wetlands Supplement (2014). Although the EFs derived in this study fall within the lower confidence margin of the Tier 1 range, our new EFs have a marked reduction in associated uncertainty. As the Tier 1 is a generic value based on published literature rather than a targeted measurement programme, it is naturally subject to a certain level of bias, which result when the underlying studies are not representative of management practices, climatic zones, or soil types in a particular region (Ogle et al., 2004), and may lead to either an over- or underestimation of CO2-C emissions. Given that no significant difference exists between the EFs derived for the IP and DP sites in this study, we propose a single EF for CO2-Con-site of 1.68 t CO2-C ha-1 yr-1 to be applied to peatlands managed for extraction in the ROI and UK regardless of peat type. This EF value could be further disaggregated by regional climate, domestic peat extraction intensity (based on extraction rates) or by end use of the peat (horticulture or energy) if more data becomes available. For the latter, it would be highly useful to determine quantitatively whether CO2-Con-site emissions vary between the less decomposed residual peat utilised for horticulture and the more decomposed residual peat used for energy production. As the EFs derived in this study have come from sites located within the same “climatic” region, we feel that they are more appropriate for the ROI and the UK inventory purposes than either the 2006 GPG or the 2013 Wetlands Supplement. If the CO2-Con site EFs derived from this study are used in annual NIRs, we estimate that annual emissions would be 9.5 and 6 times higher for the ROI and UK respectively, in comparison to the emissions calculated with the 2006 GPG Tier 1 value, and 40% lower than emissions calculated with the Wetlands Supplement EF.

As reported CO2-Con-site emissions are henceforth likely to be much higher for any country that moves from the 2006 GPG to the 2013 Wetlands Supplement, some consideration of potential mitigation measures is required. Wetland Drainage and Rewetting is a new elective activity under Article 3.4 of the Kyoto Protocol (second commitment period) andapplies to all lands that have been drained since 1990 and to all lands that have been rewetted since 1990. Countries that elect to report under this activity will also be able to claim C benefits from the rewetting of drained peatlands. In theory, this should provide an impetus for the rewetting of high emitting land use categories such as peatlands managed for extraction, particularly as these areas will remain persistent long term emission hotspots in the absence of rewetting actions (Waddington et al., 2002).

**4.7 Information gaps**

Greenhouse gas emissions from peatlands used for extraction are composed of (a) on-site emissions (i.e. from peat extraction areas, ditches and stockpiles) and (b) off-site emissions associated with water borne losses and the use of the peat for energy or horticulture. In this paper, we have focused solely on the on-site CO2-C emissions from the peat extraction areas, and GHG emissions from fire. However, C losses from other pathways may also be substantial. Research has shown that GHG emissions from on-site peat stockpiles and ditches are considerable (Alm et al., 2007a and references therein). Currently, emissions data from stockpiles in the temperate zone are not available and the IPCC Wetlands Supplement does not provide a Tier 1 value, and instead encourages countries to move to higher Tiers in terms of reporting (IPCC, 2014). However, countries such as Finlandhave developed a Tier 2 approach in which EFs (incl. CH4 and N2O) depend on regional weather and in which emissions from ditches and stockpiles are taken into account (Alm et al., 2007a; Lapveteläinen et al., 2007). The IPCC Wetlands Supplement provides Tier 1 EFs for CH4 emissions from both peat extraction areas and from ditches. The value for the latter is particularly high (542 kg CH4 ha-1 yr-1­, expressed per unit area of ditch surface) and indicates the importance of this pathway in the full GHG balance (Evans et al., 2015). Similarly, N2O emissions have been shown to be significant from drained peatlands (Regina et al., 1996) yet despite this, there are only a small number of published studies and more research is critical in order to provide regional specific EFs. While CH4 and N2O fluxes have been quantified at some of the sites, the data is currently being processed with a view to publication in the future. In terms of the fire study, N2O is a difficult gas to measure using the FTIR setup employed in this study, as it can only be determined from spectra with very large enhancements of trace gases. This is because the N2O absorption occurs in a similar wave number region to both the CO2 and CO absorption bands (Paton-Walsh et al., 2014). Paton-Walsh et al. (2014) could only determine N2O from two of their five open fires, whilst Smith et al. (2014), who used a similar setup, failed to determine N2O from any of their 21 fires studied. In our study, we found that excess mole fractions of N2O could not be correlated to either CO2 or CO for the determination of emission ratios, precluding the calculation of EFs. One explanation for this is that N2O is predominantly a product of flaming combustion and is strongly correlated to CO2 (Paton-Walsh et al., 2014). The lack of flaming combustion in our peat burns probably explains our inability to detect significant excess N2O mole fractions.

Other pathways may be of equal importance. For example, the loss of POC from bare peat surfaces may be considerable where the surface is exposed and subject to wind or water erosion (Evans et al., 2006; Lindsay, 2010). While some of the windborne POC is likely to be deposited within the extraction field itself, a proportion undoubtedly leaves the peatland, although there are currently few data available to quantify losses from either wind or water erosion, or the extent to which POC is converted to CO2 (IPCC, 2014). In addition, high losses of DOC from drained peatlands have been reported (Evans et al., 2015 and references therein). Although a Tier 1 EF value for DOC is provided in the IPCC Wetlands Supplement, disaggregated by climate zone, with the assumption that 90% of the exported DOC is converted to CO2, there is an obvious need to quantify these losses on a regional basis given the high precipitation loads experienced by the ROI and the UK, and associated differences in peat type (Evans et al., 2015). Emissions from burning are not currently reported in either the ROI or UK inventory reports. The EF provided in the IPCC Wetlands Supplement for CO2 emissions associated with wildfire burning is similar to our value here (Table 3). Furthermore, given the high CH4 emissions associated with the burning of the peat that we have reported here (Table 3), and taking cognisance of the strong GWP of CH4, more research is urgently required to quantify this emission pathway, particularly under field conditions.

The provision of activity data for inventory reporting varies between the ROI and the UK, with the peat extraction industry the source of data in the former (Duffy et al., 2014), and a multi-source approach (Directory of Mines and Quarries point locations with Google Earth imagery, scientific reports/papers) used in the latter (Webb et al., 2014). However, CO2 emissions from domestic peat extraction in the ROI are not currently reported due to a lack of activity data and could potentially be very high (Wilson et al., 2013b). In the UK, areas under domestic extraction are included in the Grassland category but may be moved as the UK considers changes post-Wetlands Supplement. Determining to what degree that peatlands have been affected by domestic peat extraction and how far those impacts extend into the main peatland area are obvious challenges facing future research.The use of remote sensing platforms could provide high resolution data that will be able to differentiate between domestic peat extraction and other types of disturbed peatlands. In particular, the use of Unmanned Aerial Vehicles (i.e. drones), which have been used to map individual peatlands at a very high resolution (e.g. Knoth et al., 2013) offer considerable potential for more detailed mapping of domestic peatlands at the national scale.

**5 Conclusion**

Peatlands managed for extraction are a substantial CO2-Cemissions hotspot at the landscape scale and further contribute to climate change through significant GHG emissions when the peat is burned or utilised in horticulture. This study, which measured and modelled emissions from a range of sites across the ROI and the UK, has highlighted the importance of generating robust Tier 2 values for different regions and land-use categories. Given that the IPCC Tier 1 EF was only based on 20 sites (all from Canada/Fenno-Scandia) we suggest that data from another 9 sites significantly expands the global dataset, as well as adding a new region.

**Acknowledgements**

Funding to DW and FRW by the Environment Protection Agency (Ireland) and by Bord na Móna is acknowledged. Thanks to Ignatius Kelly, Bord na Móna for rainfall data. Funding to SDD and EA (Site IP6) was provided by the Department for Environment, Food and Rural Affairs (UK) under research grant number SP1210. Thanks to Chris Millar, Lancashire Wildlife Trust and Phil Jepson for access to IP6, and to Cat Moody for fieldwork assistance. Experimental work at IP5 was carried out as part of the RECIPE project, with funding through the European Union Fifth Framework Programme. We would like to thank Dr Christopher MacLellan and Dr Alasdair MacArthur of NERC FSF for their wide ranging support, including the loan of the FTIR spectrometer and infrared White cell used here. We thank Dr Bruce Main and Bill Luckhurst of King’s College London for their technical support at the laboratory burns. Grateful thanks to Eeva-Stiina Tuittila and Anna Laine for useful modelling discussions. We thank the two anonymous reviewers for their insightful comments and suggestions on the earlier draft of the manuscript.

**References**

Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D., and Wennberg, P. O.: Emission factors for open and domestic biomass burning for use in atmospheric models, Atmos. Chem. Phys., 11, 4039-4072, 10.5194/acp-11-4039-2011, 2011.

Alm, J., Shurpali, N. J., Minkkinen, K., Aro, L., Hytönen, J., Laurila, T., Lohila, A., Maljanen, M., Martikainen, P. J., Mäkiranta, P., Penttilä, T., Saarnio, S., Silvan, N., Tuittila, E.-S., and Laine, J.: Emission factors and their uncertainty for the exchange of CO2, CH4 and N2O in Finnish managed peatlands., Boreal Environment Research, 12, 191-209, 2007a.

Alm, J., Shurpali, N. J., Tuittila, E.-S., Laurila, T., Maljanen, M., Saarnio, S., and Minkkinen, K.: Methods for determining emission factors for the use of peat and peatlands - flux measurements and modelling, Boreal Environment Research, 12, 85-100, 2007b.

Basiliko, N., Blodau, C., Roehm, C., Bengtson, P., and Moore, T.: Regulation of decomposition and methane dynamics across natural, commercially mined, and restored northern peatlands, Ecosystems, 10, 1148-1165, 10.1007/s10021-007-9083-2, 2007.

Bertschi, I., Yokelson, R. J., Ward, D. E., Babbitt, R. E., Susott, R. A., Goode, J. G., and Hao, W. M.: Trace gas and particle emissions from fires in large diameter and belowground biomass fuels, Journal of Geophysical Research: Atmospheres, 108, 10.1029/2002jd002100, 2003.

Blodau, C., and Moore, T. R.: Experimental response of peatland carbon dynamics to a water table fluctuation., Aquatic Sciences, 65, 47-62, 2003.

Blodau, C., Basiliko, N., and Moore, T.: Carbon turnover in peatland mesocosms exposed to different water table levels., Biogeochemistry, 67, 331-351, 2004.

Bragazza, L., Siffi, C., Iacumin, P., and Gerdol, R.: Mass loss and nutrient release during litter decay in peatland: The role of microbial adaptability to litter chemistry, Soil Biology & Biochemistry, 39, 257–267, 2007.

Cancellieri, D., Leroy-Cancellieri, V., Leoni, E., Simeoni, A., Kuzin, А. Y., Filkov, А. I., and Rein, G.: Kinetic investigation on the smouldering combustion of boreal peat, Fuel, 93, 479-485, http://dx.doi.org/10.1016/j.fuel.2011.09.052, 2012.

Chapman, S., Buttler, A., Francez, A.-J., Laggoun - Défarge, F., Vasander, H., Schloter, M., Combe, J., Grosvernier, P., Harms, H., Epron, D., Gilbert, D., and Mitchell, E.: Exploitation of northern peatlands and biodiversity maintenance: a conflict between economy and ecology., Front Ecol Environ, 1, 525-532, 2003.

Chapman, S. J., and Thurlow, M.: Peat respiration at low temperatures., Soil Biology and Biochemistry, 30, 1013-1021, 1998.

Cleary, J., Roulet, N. T., and Moore, T. R.: Greenhouse gas emissions from Canadian peat extraction, 1990-2000: a life cycle analysis., Ambio, 34, 456-461, 2005.

Couwenberg, J., and Fritze, C.: Towards developing IPCC methane ‘emission factors’ for

peatlands (organic soils) available at: <http://www.mires-and> peat.net/pages/volumes/map10/map1003.php (last access: 18 May 2015), Mires and Peat, 10, 1–17, 2012.

Davidson, E., and Janssens, I. A.: Temperature sensitivity of soil carbon decomposition and feedbacks to climate change., Nature, 440, 165-173, 2006.

Department of Environment Food and Rural Affairs: The Natural Choice: securing the value of nature, www.official-documents.gov.uk, (last access: 18 May 2015), 2011.

Dise, N. B.: Peatland response to global change, Science, 326, 810-811, 2009.

Drösler, M.: Trace gas exchange and climatic relevance of bog ecosystems, southern Germany., Department für Ökologie, Universität München, Munich, Germany , 182 pp., 2005.

Duffy, P., Hanley, E., Hyde, B., O'Brien, P., Ponzi, J., Cotter, E., and Black, K.: National

Inventory Report 2014. Greenhouse gas emissions 1990–2012 reported to the United Nations Framework Convention on Climate Change, 469, Environmental Protection Agency (Ireland), Co. Wexford, Ireland, 2014.

Elsgaard, L., Gorres, C.-M., Hoffmann, C. C., Blicher-Mathiesen, G., Schelde, K., and Petersen, S. O.: Net ecosystem exchange of CO2 and carbon balance for eight temperate organic soils under agricultural management, Agriculture, Ecosystems & Environment, 162, 52-67, 2012.

Evans, C., Renou-Wilson, F., and Strack, M.: The role of waterborne carbon in the greenhouse gas balance of drained and re-wetted peatlands, Aquatic Sciences, In press, 2015.

Evans, M., Warburton, J., and Yang, J.: Eroding blanket peat catchments: Global and local implications of upland organic sediment budgets., Geomorphology, 79, 45-57, 2006.

Farrell, C. A., and Doyle, G. J.: Rehabilitation of industrial cutaway Atlantic blanket bog in County Mayo, North-West Ireland, Wetlands Ecology and Management, 11, 21-35, 2003.

Fenner, N., Freeman, C., and Reynolds, B.: Observations of a seasonally shifting thermal optimum in peatland carbon-cycling processes; implications for the global carbon cycle and soil enzyme methodologies Soil Biology and Biochemistry, 37, 1814-1821, 2005.

Frandsen, W. H.: Ignition probability of organic soils, Canadian Journal of Forest Research, 27, 1471-1477, 10.1139/x97-106, 1997.

Frolking, S., Roulet, N. T., Moore, T. R., Lafleur, P. M., Bubier, J. L., and Crill, P. M.: Modeling seasonal to annual carbon balance of Mer Bleue Bog, Ontario, Canada, Global Biogeochemical Cycles, 16, 4-1-4-21, 10.1029/2001gb001457, 2002.

Gažovič, M., Forbrich, I., Jager, D. F., Kutzbach, L., Wille, C., and Wilmking, M.: Hydrology-driven ecosystem respiration determines the carbon balance of a boreal peatland, Science of the Total Environment, 463–464, 675-682, http://dx.doi.org/10.1016/j.scitotenv.2013.06.077, 2013.

Glatzel, S., Kalbitz, K., Dalva, M., and Moore, T.: Dissolved organic matter properties and their relationship to carbon dioxide efflux from restored peat bogs., Geoderma, 113, 397-411, 2003.

Glatzel, S., Basiliko, N., and Moore, T.: Carbon dioxide and methane production potentials of peats from natural, harvested and restored sites, eastern Québec, Canada, Wetlands., 24, 261-267, 2004.

Griffith, D. W. T.: Synthetic calibration and quantitative analysis of gas-phase FT-IR spectra, Appl. Spectrosc., 50, 59-70, 1996.

Haddaway, N. R., Burden, A., Evans, C. D., Healey, J. R., Jones, D. L., Dalrymple, S. E.,

and Pullin, A. S.: Evaluating effects of land management on greenhouse gas fluxes and carbon balances in boreo-temperate lowland peatland systems, available at: http://www.environmentalevidencejournal.org/content/3/1/5 (last access: 18 May 2015), Environmental Evidence, 3, 30, doi:10.1186/2047-2382-3-5, 2014.

Howley, M., Dennehy, E., O'Gallachoir, B., and Holland, M.: Energy in Ireland 1990-2011, Sustainable Energy Authority of Ireland, Dublin, 92, 2012.

IPCC: IPCC Guidelines for National Greenhouse Gas Inventories, Prepared by the National

Greenhouse Gas Inventories programme, IGES, Published by the Institute for Global Environmental Strategies (IGES), Hayama, Japan, 2006.

IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge, United Kingdom and New York, NY, USA, 1535, 2013.

IPCC: 2013 Supplement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories: Wetlands, edited by: Hiraishi, T., Krug, T., Tanabe, K., Srivastava, N., Baasansuren, J., Fukuda, M., and Troxler, T. G., IPCC, Switzerland, 2014.

Järveoja, J., Laht, J., Maddison, M., Soosaar, K., Ostonen, I., and Mander, Ü.: Mitigation of

greenhouse gas emissions from an abandoned Baltic peat extraction area by growing reed canary grass: life-cycle assessment, Reg. Environ. Change, 13, 781–795, doi:10.1007/s10113-012-0355-9, 2012.

Juszczak, R., Humphreys, E., Acosta, M., Michalak-Galczewska, M., Kayzer, D., and Olejnik, J.: Ecosystem respiration in a heterogeneous temperate peatland and its sensitivity to peat temperature and water table depth, Plant and Soil, 366, 505-520, 10.1007/s11104-012-1441-y, 2013.

Kettridge, N., Turetsky, M. R., Sherwood, J. H., Thompson, D. K., Miller, C. A., Benscoter,

B. W., Flannigan, M. D., Wotton, B. M., and Waddington, J. M.: Moderate drop in

water table increases peatland vulnerability to post-fire regime shift, Sci. Rep., 5, 8063,

doi:10.1038/srep08063, 2015.

Knoth, C., Klein, B., Prinz, T., and Kleinebecker, T.: Unmanned aerial vehicles as innovative remote sensing platforms for high-resolution infrared imagery to support restoration monitoring in cut-over bogs, Applied Vegetation Science, 16, 509-517, 10.1111/avsc.12024, 2013.

Koehler, A.-K., Sottocornola, M., and Kiely, G.: How strong is the current carbon sequestration of an Atlantic blanket bog?, Global Change Biology, 17, 309-319, 2011.

Lafleur, P. M., Moore, T. R., Roulet, N. T., and Frolking, S.: Ecosystem respiration in a cool temperate bog depends on peat temperature but not water table., Ecosystems, 8, 619-629, 2005.

Laine, A., Riutta, T., Juutinen, S., Väliranta, M., and Tuittila, E.-S.: Acknowledging the spatial heterogeneity in modelling / reconstructing carbon dioxide exchange in a northern aapa mire, Ecological Modelling, 220, 2646-2655, 2009.

Lapveteläinen, T., Regina, K., and Perälä, P.: Peat-based emissions in Finland's national greenhouse gas inventory, Boreal Environment Research, 12, 225-236, 2007.

Leifeld, J., Steffens, M., and Galego-Sala, A.: Sensitivity of peatland carbon loss to organic

matter quality, Geophys. Res. Lett., 39, L14704, doi:10.1029/2012GL051856, 2012.

Lindsay, R.: Peatbogs and Carbon - A critical synthesis, University of East London and RSPB Scotland, London, 2010.

Lloyd, J., and Taylor, J. A.: On the temperature dependence of soil respiration., Functional Ecology, 8, 315-323, 1994.

Lobert, J. M., Scharffe, D. H., Hao, W. M., Kuhlbusch, T. A., Seuwen, R., Warneck, P., and Crutzen, P. J.: Experimental evaluation of biomass burning emissions: Nitrogen and carbon containing compounds., in: Global Biomass Burning: Atmospheric, Climatic and Biospheric Implications, edited by: Levine, J. S., MIT Press, Cambridge, Massachusetts, USA, 1991.

Mäkiranta, P., Laiho, R., Fritze, H., Hytönen, J., Laine, J., and Minkkinen, K.: Indirect regulation of heterotrophic peat soil respiration by water level via microbial community structure and temperature sensitivity, Soil Biology and Biochemistry, 41, 695-703, http://dx.doi.org/10.1016/j.soilbio.2009.01.004, 2009.

Mander, Ü., Järveoja, J., Maddison, M., Soosaar, K., Aavola, R., Ostonen, I., and Salm, J.-O.: Reed canary grass cultivation mitigates greenhouse gas emissions from abandoned peat extraction areas, GCB Bioenergy, 4, 462-474, 10.1111/j.1757-1707.2011.01138.x, 2012.

McNeil, P., and Waddington, J. M.: Moisture controls on *Sphagnum* growth and CO2 exchange on a cutover bog., Journal of Applied Ecology, 40, 354-367, 2003.

Moore, T. R., Bubier, J. L., Frolking, S. E., Lafleur, P. M., and Roulet, N. T.: Plant biomass and production and CO2 exchange in an ombrotrophic bog, Journal of Ecology, 90, 25-26, 2002.

Nilsson, M., Sagerfors, J., Buffam, I., Laudon, H., Eriksson, T., Grelle, A., Klemedtsson, L., Weslien, P. E. R., and Lindroth, A.: Contemporary carbon accumulation in a boreal oligotrophic minerogenic mire – a significant sink after accounting for all C-fluxes, Global Change Biology, 14, 2317-2332, 10.1111/j.1365-2486.2008.01654.x, 2008.

Ogle, S. M., Conant, R. T., and Paustian, K.: Deriving grassland management factors for a carbon accounting method developed by the Intergovernmental Panel on Climate Change, Environmental Management, 33, 474-484, 2004.

Paton-Walsh, C., Smith, T. E. L., Young, E. L., Griffith, D. W. T., and Guérette, É.-A.: New

emission factors for Australian vegetation fires measured using open-path Fourier transform infrared spectroscopy – Part 1: Methods and Australian temperate forest fires, Atmos. Chem.Phys., 14, 11313–11333, doi:10.5194/acp-14-11313-2014, 2014.

Price, J. S., and Schlotzhauer, S. M.: Importance of shrinkage and compression in determining water storage changes in peat: the case of a mined peatland, Hydrological Processes, 13, 2591-2601, 1999.

Regina, K., Nykänen, H., Silvola, J., and Martikainen, P.: Fluxes of nitrous oxide from boreal peatlands as affected by peatland type, water table level and nitrification capacity, Biogeochemistry, 35, 401-418, 10.1007/bf02183033, 1996.

Rein, G., Cohen, S., and Simeoni, A.: Carbon emissions from smouldering peat in shallow and strong fronts, Proceedings of the Combustion Institute 32 2489-2496, 2009.

Renou-Wilson, F., Barry, C., Müller, C., and Wilson, D.: The impacts of drainage, nutrient status and management practice on the full carbon balance of grasslands on organic soils in a maritime temperate zone, Biogeosciences, 11, 4361-4379, doi:10.5194/bg-11-4361-2014, 2014.

Renou, F., Egan, T., and Wilson, D.: Tomorrow's landscapes: studies in the after-uses of industrial cutaway peatlands in Ireland, Suo, 57, 97-107, 2006.

Riutta, T., Laine, J., and Tuittila, E.-S.: Sensitivity of CO2 exchange of fen ecosystem components to water level variation, Ecosystems, 10, 718-733, 10.1007/s10021-007-9046-7, 2007.

Roulet, N. T., Lafleur, P. M., Richard, P. J. H., Moore, T., Humphreys, E. R., and Bubier, J.:

Contemporary carbon balance and late Holocene carbon accumulation in a northern peatland, Global Change Biol., 13, 397–411, doi:10.1111/j.1365-2486.2006.01292.x, 2007.

Saarnio, S., Morero, M., Shurpali, N. J., Tuittila, E.-S., Mäkilä, M., and Alm, J.: Annual CO2 and CH4 fluxes of pristine boreal mires as a background for the lifecycle analyses of peat energy, Boreal Environment Research, 12, 101-113, 2007.

Salm, J.-O., Maddison, M., Tammik, S., Soosaar, K., Truu, J., and Mander, Ü.: Emissions of CO2, CH4 and N2O from undisturbed, drained and mined peatlands in Estonia, Hydrobiologia, 692, 41-55, 10.1007/s10750-011-0934-7, 2012.

Schouten, M. G. C.: Conservation and restoration of raised bogs: geological, hydrological

and ecological studies, Department of the Environment and Local Government / Staatsboshbeheer, Dublin, Ireland, 2002.

Shurpali, N. J., Hyvönen, N., Huttunen, J. T., Biasi, C., Nykanen, H., Pekkarinen, N., and

Martikainen, P. J.: Bare soil and reed canary grass ecosystem respiration in peat extraction sites in Eastern Finland, Tellus, 60B, 200-209, 2008.

Silvola, J., Alm, J., Ahlholm, U., Nykänen, H., and Martikainen, P. J.: CO2 fluxes from peat in boreal mires under varying temperature and moisture conditions, Journal of Ecology, 84, 219-228, 1996.

Smith, T. E. L., Paton-Walsh, C., Meyer, C. P., Cook, G. D., Maier, S. W., Russell-Smith, J., Wooster, M. J., and Yates, C. P.: New emission factors for Australian vegetation fires measured using open-path Fourier transform infrared spectroscopy – Part 2: Australian tropical savanna fires, Atmos. Chem. Phys., 14, 11335-11352, 10.5194/acp-14-11335-2014, 2014.

Soini, P., Riutta, T., Yli-Petäys, M., and Vasander, H.: Comparison of vegetation and CO2 dynamics between a restored cut-way peatland and a pristine fen: evaluation of the restoration success, Restoration Ecology, 18, 894-903, 2010.

Stockwell, C. E., Yokelson, R. J., Kreidenweis, S. M., Robinson, A. L., DeMott, P. J., Sullivan, R. C., Reardon, J., Ryan, K. C., Griffith, D. W. T., and Stevens, L.: Trace gas emissions from combustion of peat, crop residue, domestic biofuels, grasses, and other fuels: configuration and Fourier transform infrared (FTIR) component of the fourth Fire Lab at Missoula Experiment (FLAME-4), Atmos. Chem. Phys., 14, 9727-9754, 10.5194/acp-14-9727-2014, 2014.

Strack, M., Waddington, J. A., Bourbonniere, R. A., Buckton, E. L., Shaw, K., Whittington, P., and Price, J. S.: Effect of water table drawdown on peatland dissolved organic carbon export and dynamics, Hydrological Processes, 22, 3373-3385, DOI: 3310.1002/hyp.6931, 2008.

Strack, M., and Price, J. S.: Moisture controls on carbon dioxide dynamics of peat-Sphagnum monoliths, Ecohydrology, 2, 34-41, 10.1002/eco.36, 2009.

Strack, M., and Zuback, Y. C. A.: Annual carbon balance of a peatland 10 yr following restoration, Biogeosciences 10, 2885-2896, 2013.

Strack, M., Keith, A. M., and Zu, B.: Growing season carbon dioxide and methane exchange at a restored peatland on the Western Boreal Plain, Ecological Engineering, 64, 231-239, 2014.

Sundh, I., Nilsson, M., Mikkelä, C., Granberg, G., and Svensson, B. H.: Fluxes of methane and carbon dioxide on peat-mining areas in Sweden, Ambio, 29, 499-503, 2000.

Tomlinson, R. W.: Changes in the extent of peat extraction in Northen Ireland 1990-2008 and associated changes in carbon loss, Applied Geography, 30, 294-301, 2010.

Tuittila, E.-S., Komulainen, V.-M., Vasander, H., and Laine, J.: Restored cut-away peatland as a sink for atmospheric CO2, Oecologia, 120, 563 - 574, 1999.

Tuittila, E.-S., Vasander, H., and Laine, J.: Sensitivity of carbon sequestration in reintroduced *Sphagnum* to water-level variation in a cutaway peatland, Restoration Ecology, 12, 482-492, 2004.

Tuittila, E. S., and Komulainen, V. M.: Vegetation and CO2 balance in an abandoned harvested peatland in Aitoneva, southern Finland, Suo, 46, 69-80, 1995.

Turetsky, M. R., Kotowska, A., Bubier, J., Dise, N. B., Crill, P., Hornibrook, E. R. C., Minkkinen, K., Moore, T. R., Myers-Smith, I. H., Nykänen, H., Olefeldt, D., Rinne, J., Saarnio, S., Shurpali, N., Tuittila, E.-S., Waddington, J. M., White, J. R., Wickland, K. P., and Wilmking, M.: A synthesis of methane emissions from 71 northern, temperate, and subtropical wetlands, Glob. Change Biol., 20, 2183–2197, doi:10.1111/gcb.12580, 2014.

Turetsky, M. R., Benscoter, B., Page, S., Rein, G., van der Werf, G. R., and Watts, A.: Global vulnerability of peatlands to fire and carbon loss, Nature Geoscience, 8, 11-14, 10.1038/ngeo2325.http://www.nature.com/ngeo/journal/v8/n1/abs/ngeo2325.html#supplementary-information, 2015.

Verhoeven, J. T. A., and Toth, E.: Decomposition of Carex and Sphagnum litter in fens: Effect of litter quality and inhibition by living tissue homogenates, Soil Biology and Biochemistry, 27, 271-275, http://dx.doi.org/10.1016/0038-0717(94)00183-2, 1995.

Waddington, J. M., and Price, J. S.: Effect of peatland drainage, harvesting and restoration on atmospheric water and carbon exchange, Physical Geography, 21, 433-451, 2000.

Waddington, J. M., Warner, K. D., and Kennedy, G. W.: Cutover peatlands: a persistent source of atmospheric CO2, Global Biogeochemical Cycles, 16, 21-27, 2002.

Waddington, J. M., Strack, M., and Greenwood, M. J.: Toward restoring the net carbon sink function of degraded peatlands: Short-term response in CO2 exchange to ecosystem-scale restoration, J. Geophys. Res., 115, G01008, doi:10.1029/2009JG001090, 2010.

Webb, N., Broomfield, M., Brown, P., Buys, G., Cardenas, L., Murrells, T., Pang, Y., Passant, N., Thistlewaite, G., and Watterson, J.: UK Greenhouse Gas Inventory, 1990 to 2012. Annual Report for Submission under the Framework Convention on Climate Change, Compiled on Behalf of the UK Department of Energy and Climate Change (DECC) Science Division by Ricardo-AEA, 594, Ricardo-AEA, Oxfordshire, UK, 2014.

Wilson, D., Tuittila, E.-S., Alm, J., Laine, J., Farrell, E. P., and Byrne, K. A.: Carbon dioxide dynamics of a restored maritime peatland, Ecoscience, 14, 71-80, 2007.

Wilson, D., Farrell, C., A., Müller, C., Hepp, S., and Renou-Wilson, F.: Rewetted industrial cutaway peatlands in western Ireland: prime location for climate change mitigation?, available at: http://mires-and-peat.net/pages/volumes/map11/map1101.php (last access: 18 May 2015), Mires and Peat, 11, 1–22, 2013a.

Wilson, D., Müller, C., and Renou-Wilson, F.: Carbon emissions and removals from Irish peatlands: current trends and future mitigation measures, Irish Geography, 46, 1-23, 2013b.

Yokelson, R. J., Griffith, D. W. T., and Ward, D. E.: Open-path Fourier transform infrared studies of large-scale laboratory biomass fires, Journal of Geophysical Research: Atmospheres, 101, 21067-21080, 10.1029/96jd01800, 1996.

Yokelson, R. J., Susott, R., Ward, D. E., Reardon, J., and Griffith, D. W. T.: Emissions from smoldering combustion of biomass measured by open-path Fourier transform infrared spectroscopy, Journal of Geophysical Research: Atmospheres, 102, 18865-18877, 10.1029/97jd00852, 1997.

Yokelson, R. J., Susott, R., Ward, D. E., Reardon, J., and Griffith, D. W.: Emissions from smoldering combustion of biomass measured by open-path Fourier transform infrared spectroscopy, Journal of Geophysical Research: Atmospheres (1984-2012), 102 (D15), 18865-18877, 2008.

Table 1. Site characteristics. Mean annual air temperature (°C) and mean annual rainfall (mm yr-1) are long-term values (1981-2010; Met Éireann http://www.met.ie/ and Met Office UK; http://www.metoffice.gov.uk). \*Time between cessation of peat extraction and the study period.

|  |  |  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- | --- | --- |
| Site name | Boora | Blackwater | Bellacorick | Turraun | MiddlemuirMoss | LittleWoolden | Clara | Glenlahan | Moyarwood |
| Site code | IP1 | IP2 | IP3 | IP4 | IP5 | IP6 | DP1 | DP2 | DP3 |
| Time since last extraction\*  | >20 years | >25 years | >10 years | >30 years | >10 years | ca. 1 | 0 | >20 years | >20 years |
| Study period | 1/9/2007: 30/8/2009 | 1/5/2011: 30/4/2014 | 1/1/2012: 31/12/2013 | 1/1/2002: 31/12/2003 | 1/11/2003: 31/10/2004 | 1/1/2013: 31/12/2014 | 1/4/2006: 31/3/2007 | 1/4/2006: 31/3/2007 | 1/4/2013: 31/3/2014 |
| Latitude Longitude | 53.203-7.726 | 53.297-7.965 | 54.128-9.556 | 53.260-7.720 | 57.60-2.15 | 53.451-2.468 | 53.316-7.647 | 53.103-7.538 | 53.346-8.514 |
| Sub-region | IrishMidlands | Irish Midlands | North-West Ireland | Irish Midlands | North-EastScotland | Northern England | Irish Midlands | Irish Midlands | Western Ireland |
| Mean annual air temperature (°C)  | 9.3 | 9.8 | 10.3 | 9.3 | 8.0 | 10.2 | 9.3 | 9.3 | 10.0 |
| Mean annual rainfall (mm yr-1) | 970 | 907 | 1245 | 807 | 851 | 867 | 970 | 804 | 1193 |
| Vegetation |  |  |  |  |  |  | *Calluna vulgaris*, *Erica. tetralix*, *Cladonia* sp. |
| Peat type | Phragmites | Phragmites | Cyperaceous | Phragmites | Sphagnum/ Cyperaceous | Sphagnum/ Cyperaceous | Sphagnum | Ericaceous | Sphagnum |
| von Post scale | H7 | H7 | H5 to 6 | H7 | H8 | H6 to 7 | H6 | H6 | H6 |
| Parent material | Limestone | Limestone | Shale | Limestone | Granite drifts and rocks | Triassic Sandstone | Limestone | Old Red Sandstone | Limestone |
| Peat depth (m) | 1.0 | 1.5 | 0.5 | 0.5-1.8 | 0.7-3.1 | 0.5-1.75 | 4 | 0.4 | 4.4 |
| pH | 4.3 | 4.9 | 3.8 | 6.3 | 3.6-4.1 | 2.9 | 4.0 | 3.8 | 4.4 |
| C (%) | 50 | 52.4 | 56 | 52 | 52 | 49.1 | 49.8 | 29.1 | 51.5 |
| N (%) | 1.09 | 2.14 | 0.97 | 2.1 | 1.4 | 1.34 | 1.46 | 0.69 | 1.32 |
| C:N | 45.9 | 24.5 | 57.7 | 24.8 | 37 | 36.6 | 34.1 | 42.2 | 39 |
|  |  |  |  |  |  |  |  |  |  |

Table 2. Emission factors (tonnes CO2-C ha-1 yr-1) for sites IP1-6 and DP1-3.

Uncertainties are 95% confidence intervals.

|  |  |  |
| --- | --- | --- |
|  | CO2-C | 95% confidence interval |
| Site | (t ha-1 yr-1) |  (t ha-1 yr-1) |
| IP1  | 1.82 | 1.75 | 1.89 |
| IP2  | 1.53 | 1.37 | 1.60 |
| IP3  | 1.38 | 1.25 | 1.52 |
| IP4  | 2.86 | 2.65 | 3.06 |
| IP5  | 0.93 | 0.59 | 1.27 |
| IP6  | 1.70 | 1.43 | 1.98 |
| **Emission factor** | **1.70** | **1.23** | **2.17** |
|  |  |  |  |
| DP1  | 1.76 | 1.59 | 1.99 |
| DP2  | 2.03 | 1.73 | 2.30 |
| DP3  | 1.14 | 0.85 | 1.41 |
| **Emission factor** | **1.64** | **1.22** | **2.06** |

Table 3.Mean modified combustion efficiency (MCE) and emission factors (g kg-1 dry fuel burned) reported by this study and those for the same trace gases reported by previous studies of temperate or boreal peat (Yokelson et al. 1997; Stockwell et al. 2014). The mean and standard deviation of the emission factor is calculated from individual sample burns. *nr*=not reported.

|  |  |
| --- | --- |
|  | Emission Factor (g kg-1 dry fuel burned) |
| Trace Gas | Irish sphagnum moss peat (this study) | Canadian boreal peat(Stockwell et al. 2014) | North Carolina temperate peat (Stockwell et al. 2014) | Alaska/Minnesota peat(Yokelson et al. 1997) |
| MCE | 0.837 ± 0.019 | 0.805 ± 0.009 | 0.726 ± 0.009 | 0.809 ± 0.033 |
| CO2 | 1346 ± 31 | 1274 ± 19 | 1066 ± 287 | 1395 ± 52 |
| CO | 218 ± 22 | 197 ± 9 | 276 ± 139 | 209 ± 68 |
| CH4 | 8.35 ± 1.3 | 6.25 ± 2.17 | 10.9 ± 5.3 | 6.85 ± 5.66 |
| C2H4 | 1.74 ± 0.23 | 0.81 ± 0.29 | 1.27 ± 0.51 | 1.37 ± 0.51 |
| C2H6 | 1.53 ± 0.17 | *nr* | *nr* | *nr* |
| CH3OH | 0.60 ± 0.87 | 0.75 ± 0.35 | 2.83 ±2.87 | 4.04 ± 3.43 |
| HCN | 2.21 ± 0.35 | 1.77 ± 0.55 | 4.45 ± 3.02 | 5.09 ± 5.64 |
| NH3 | 0.73 ± 0.50 | 2.21 ± 0.24 | 1.87 ± 0.37 | 8.76 ± 13.76 |

Table 4.Annual CO2-C emissions (tonnes CO2-C yr-1) from peatlands managed for extraction in the ROI and UK calculated using the IPCC 2006 Good Practice Guidance (Tier 1 value: 0.2 and 1.1 t CO2-C ha-1 yr-1for nutrient poor and nutrient rich peatlands respectively), the IPCC 2013 Wetlands Supplement (Tier 1 value: 2.8 t CO2-C ha-1 yr-1) and the Emission Factors derived in this study (Table 2). Areas (ha) and CO2-C emissions using the IPCC 2006 Good Practice Guidance values are taken from the 2014 National Inventory Reports (NIR) for the ROI (Duffy et al., 2014) and the UK (Webb et al., 2014).

|  |  |  |
| --- | --- | --- |
| Country | Area (ha) | Emissions (tonnes CO2-C yr-1) |
|  |  | IPCC 2006 | IPCC 2013 | This study |
| ROI | 52,422  | 9,312 | 146,782 | 88,069 |
| England | 4,790 | 960 | 13,412 |  8,047 |
| Scotland | 1,610 |  545 | 4,508 | 2,705 |
| Wales | 482 |  95 | 1,350 |  810 |
| N. Ireland | 1,030 | 518 | 2,884 |  1,730 |
| UK | 7,912 | 2,118 | 22,154 | 13,292 |



Figure 1. Annual rainfall (mm), mean annual water tables (cm), mean annual temperature (°C) at 5 cm depths (T5cm) at sites IP1 (two years), IP2 (three years), IP3 (two years), IP4 (two years), IP5 (one year), IP6 (two years), DP1 (one year), DP2 (one year) and DP3 (one year). Dotted horizontal line indicates 30 year mean rainfall at each site (1981-2010; Met Éireann http://www.met.ie/ and Met Office UK; http://www.metoffice.gov.uk). Error bars are standard deviations. Negative water table values indicate water level below the soil surface.



Figure 2. Frequency distribution of soil temperature at 5cm depth (T5cm) at sites IP1-6 shown as a percentage (%) of total count.

****

Figure 3. (a) Ecosystem respiration (Reco; mg CO2-C m-2 hr-1) at sites IP1-6, (b) Reco (mg CO2-C m-2 hr-1) at sites DP1-3 and (c) net ecosystem exchange (NEE; mg CO2-C m-2 hr-1) when PPFD>1000 µmol m-2 s-1at sites DP1-3. Positive values indicate CO2-C flux from the peatland to the atmosphere (source) and negative values indicate CO2-C flux from the atmosphere to the peatland (sink). The 10th and 90th percentile are indicated by the bars, the 25th and 75th percentiles with the top and bottom of the box and the median value by the centre line. Different letters indicate significant differences in the *post-hoc* test for multiple comparisons.



Figure 4. Annual cumulative ecosystem respiration (Reco: g CO2-C m-2) at sites IP1-6. Positive values indicate CO2-C flux from the peatland to the atmosphere (source). Value at end of the curve indicates the total annual Reco value. Brown line indicates year 1, black line year 2 and grey line year 3 of the study at the individual sites. Note the differences in integration period between sites (x axis).

****

Figure 5. Annual cumulative gross primary productivity (GPP: g CO2-C m-2), ecosystem respiration (Reco: g CO2-C m-2), heterotrophic respiration (RH: Site DP1 only) and net ecosystem exchange (NEE: g CO2-C m-2) at sites DP1-3. Positive values indicate CO2-C flux from the peatland to the atmosphere (source) and negative values indicate CO2-C flux from the atmosphere to the peatland (sink). Value at end of the curve indicates the total annual value for each component. Note the differences in integration period between sites (x axis).

****

Figure 6. Relationship between (a) ecosystem respiration (Reco: t CO2-C ha-1 yr-1) and mean soil temperature (°C) at 5 cm depth at the IP sites and (b) net ecosystem exchange (NEE: t CO2-C ha-1 yr-1) and leaf area index (LAI: m2 m-2). Circles indicate an annual value.

****

Figure 7. Carbon dioxide emissions (t CO2-C ha-1 yr-1) from peatlands managed for extraction in Canada, ROI/UK (this study) and Fenno-scandinavia. The 10th and 90th percentile are indicated by the bars, the 25th and 75th percentiles with the top and bottom of the box and the median value by the centre line.

(Data for Canada and Fenno-Scandia taken from the following studies; Tuittila and Komulainen, 1995; Sundh et al., 2000; Waddington et al., 2002; Glatzel et al., 2003; McNeil and Waddington, 2003; Tuittila et al., 2004; Cleary et al., 2005; Alm et al., 2007a; Shurpali et al., 2008; Waddington et al., 2010; Järveoja et al., 2012; Mander et al., 2012; Salm et al., 2012; Strack et al., 2014). Where studies reported seasonal fluxes (typically May to October), these were converted to annual fluxes by assuming that 15% of the flux occurs in the non-growing season (Saarnio et al., 2007).