

1 **Referee #2**

2
3 *General comment:*

4 *The manuscript improved significantly after the previous revisions. Although insufficient*
5 *sequence number of clone library is the major limitation of this study, addition of qPCR of*
6 *some nifH phylotypes is helpful in compensating the limitation. As limited amount of relevant*
7 *studies have been done in temperate regions of North Pacific Ocean, it is difficult for the*
8 *authors to further elaborate and discuss with this set of data in a convincing way. On another*
9 *hand, the data set could provide useful information for the related studies in the future.*
10 *Therefore, I think this work can be accepted, after some minor revisions. There are some*
11 *grammatical mistakes found in the manuscript, and the authors should do proof-reading more*
12 *carefully.*

13
14 We have checked the manuscript thoroughly and have revised unclear wordings.
15 Furthermore, the English in the revised manuscript has been checked by a professional editor,
16 a native speaker of English.

17
18 *Specific comments:*

19 *L.141, What does “when nitrogen fixation was not detected” mean? Did the authors mean*
20 *“when nitrogen fixation was undetectable” or “not measured”? If the nitrogen fixation rate*
21 *was not measured, the authors should not assume the missing data to be zero. The authors*
22 *should clarify their meaning here.*

23
24 We have corrected the sentence to “when nitrogen fixation was undetectable”. (L140)

25
26 *L.199, “underestimates” should be “underestimated”.*

27
28 We have corrected as suggested. (L195)

29
30 *L.401, “one or more the factors” should be “one or more factors”*

31
32 We have corrected as suggested. (L384)

33
34 *L. 432, It is suggested to use “undetectable” to replace “disappear”*

35
36 We have changed the sentence as follows. (L411)

37 “It appears that UCYN-A abundance decreased with decreasing temperature from fall to
38 winter, and then became undetectable in spring.”

39

40 *L.455-457, The logic here is not clear enough. Since the author did not use qPCR to quantify*
41 *the P. stutzeri-like nifH gene, there is no reason to say “P. stutzeri could not be a major*
42 *diazotroph in this study region”. Also, “ γ -24774A11 was not detected on that occasion by*
43 *qPCR analysis” should not be the evident suggesting “ γ -24774A11 was not quantified as P.*
44 *stutzeri”. The authors can simply compare the sequences of γ -24774A11 and P. stutzeri nifH*
45 *gene recovered in this study.*

46

47 We have corrected the sentence as follows. (L436-437)

48 “ γ -24774A11 was not detected on that occasion by qPCR analysis probably due to the
49 difference in the sequence between γ -24774A11 and *P. stutzeri*.”

50

51 *L.496, It is suggested to add “coastal” after “temperate”, as the study was conducted in*
52 *coastal area.*

53

54 We have added “coastal” as suggested. (L470)

55

1 **Nitrogen fixation and the diazotroph community in the**
2 **temperate coastal region of the northwestern North Pacific**

3

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10

11 **Abstract**

12 Nitrogen fixation in temperate oceans is a potentially important, but poorly understood
13 process that may influence the marine nitrogen budget. This study determined seasonal
14 variations in nitrogen fixation and the diazotroph community within the euphotic zone in the
15 temperate coastal region of the northwestern North Pacific. Nitrogen fixation as high as
16 $13.6 \text{ nmol N L}^{-1} \text{ d}^{-1}$ was measured from early summer to fall when the surface temperature
17 exceeded 14.2°C (but was lower than 24.3°C) and the surface nitrate concentration was
18 low ($\leq 0.30 \text{ }\mu\text{M}$), although we also detected nitrogen fixation in subsurface layers (42–62 m)

19 where nitrate concentrations were high ($>1 \mu\text{M}$). Clone library analysis results indicated
20 that *nifH* gene sequences were omnipresent throughout the investigation period. During the
21 period when nitrogen fixation was detected (early summer to fall), the genes affiliated with
22 UCYN-A, *Trichodesmium*, and γ -proteobacterial phylotype γ -24774A11 were frequently
23 recovered. In contrast, when nitrogen fixation was undetectable (winter to spring), many
24 sequences affiliated with Cluster III diazotrophs (putative anaerobic bacteria) were recovered.
25 Quantitative PCR analysis revealed that UCYN-A was relatively abundant from early to late
26 summer compared with *Trichodesmium* and γ -24774A11, whereas *Trichodesmium* abundance
27 was the highest among the three groups during fall.

28

29 **1. Introduction**

30 The amount of bioavailable nitrogen introduced into the global ocean via nitrogen fixation is
31 considered to be roughly balanced at the large spatiotemporal scale by nitrogen loss through
32 denitrification, as indicated by the sedimentary nitrogen isotope record during the Holocene
33 epoch (Brandes and Devol, 2002; Deutsch et al., 2004). However, rate measurement data
34 have revealed that denitrification far exceeds nitrogen fixation (Codispoti, 2007). This
35 discrepancy in the nitrogen balance has raised the possibility that the current estimate of
36 marine nitrogen fixation, which is primarily based on data collected in tropical and

37 subtropical oceans where large cyanobacterial diazotrophs (e.g., *Trichodesmium* spp. and
38 *Richelia intracellularis*) are considered to be ~~mainly~~primarily responsible for nitrogen
39 fixation (e.g., Capone et al., 1997), might be too low (Codispoti, 2007). This is supported
40 by the results of recent studies using molecular approaches that have increasingly revealed
41 that marine diazotrophs are more diverse and widespread than previously thought (Riemann
42 et al., 2010; Zehr, 2011). Recently discovered marine diazotrophic taxa, including those
43 belonging to unicellular cyanobacteria and heterotrophic bacteria, are abundant in oceanic
44 regions where large cyanobacterial diazotrophs are scarce (Needoba et al., 2007; Moisaner
45 et al., 2010; Halm et al., 2012; Bonnet et al., 2013; Rahav et al., 2013; Shiozaki et al., 2014a),
46 suggesting that a failure to account for nitrogen fixation mediated by these diazotrophs might
47 result in underestimation of marine nitrogen fixation.

48 The temperate coastal ocean is one of the regions where nitrogen fixation rates have been
49 understudied and potentially underestimated. Conventionally, nitrogen fixation in temperate
50 oceans has been assumed to be low because of the relatively low temperatures (< ~20°C),
51 which generally inhibit the growth of large cyanobacterial diazotrophs (Breitbarth et al.,
52 2007), and development of high dissolved inorganic nitrogen (DIN) concentrations (>1 µM).
53 High DIN concentrations are generally regarded to inhibit nitrogen fixation (Falkowski,
54 1983), especially during mixing periods. However, recent studies have indicated that

55 nitrogen fixation, presumably mediated by unicellular cyanobacteria and heterotrophic
56 bacteria, is detectable even in the relatively cold ($<10^{\circ}\text{C}$) and DIN-rich waters ($>1\ \mu\text{M}$) of
57 the Atlantic coast (Mulholland et al., 2012) and the Baltic Sea estuaries (Bentzon-Tilia et al.,
58 2015). These results highlight the necessity of re-evaluating the extent, variation, and
59 control mechanisms of nitrogen fixation in temperate oceans, with recognition of the
60 widespread occurrence of diverse diazotrophic microbes.

61 This study examined the seasonal variation in nitrogen fixation—along two inshore-offshore
62 transects in the temperate inside bays and open ocean located in the interfrontal zone of the
63 northwestern North Pacific.—In this temperate region, physical, chemical, and biological
64 properties vary widely between seasons (Shiozaki et al., 2014b) due to the confluence of
65 three currents: the Kuroshio (warm current), the Tsugaru Warm Current, and Oyashio (cold
66 current). Data on nitrogen fixation rates in the temperate Pacific are limited (Needoba et al.,
67 2007), and to the best of our knowledge, the present study is the first to examine diazotrophy
68 during all seasons in the temperate ocean. This study was conducted as part of a project to
69 monitor the dynamics of the coastal ecosystem and the recovery thereof after the 2011
70 Tohoku-oki tsunami, which struck the region on 11 March 2011.

71

72 **2. Materials and Methods**

73 The experiments were conducted during six cruises in the temperate coastal region of the
74 western North Pacific. These cruises covered a full seasonal cycle, including spring
75 (KS-14-2_Mar, 14–19 March 2014), early summer (KK-13-1_Jun, 24–29 June 2013),
76 summer (KT-12-20_Aug, 7–12 August 2012), late summer (KK-13-6_Sep, 14–21 September
77 2013), fall (KT-12-27_Oct, 15–22 October 2012), and winter (KT-13-2_Jan, 19–25 January
78 2013). Sampling stations were located along the transect lines OT ($39^{\circ}20'N$,
79 $141^{\circ}56'–142^{\circ}50'E$) and ON ($38^{\circ}25'N$, $141^{\circ}29'–142^{\circ}20'E$). Eight stations were located
80 offshore (OT4–6, ON4–8), while two stations were deployed in the Otsuchi (OT1) and
81 Onagawa (ON1) bays (Fig. 1). Just before the KK-13-6_Sep cruise, Typhoon Man-yi
82 passed from southwest to northeast in the study area (Fig. S1).

83 Temperature, salinity, and dissolved oxygen profiles of regions near the bottom floor were
84 measured using a SBE 911-plus conductivity-temperature-pressure (CTD) system (Sea-bird
85 Electronics, Bellevue, WA, USA). Water samples were collected in an acid-cleaned bucket
86 and Niskin-X bottles. At offshore stations, samples for nutrient analysis were collected
87 from 7–15 different depths in the upper 200 m, while at shallower (<200 m) bay stations,
88 samples were collected from 4–9 different depths in the entire water column, except at Stn.
89 OT1 where only surface water samples were collected. Samples for DNA analysis and
90 incubation experiments were collected from the surface at almost every station, and from

91 depths corresponding to 10% and 1% of the surface light intensities at Stns. OT4 and ON5.

92 Light attenuation was determined using a submersible PAR sensor.

93 **2.1. Nutrients**

94 Samples for nutrient analysis were stored in 10-mL ~~polyethylene-acrylic~~ tubes and kept

95 frozen until onshore analyses. Nitrate, nitrite, ammonium, and phosphate concentrations

96 were determined using an AACSII auto-analyzer (Bran+Luebbe, Norderstedt, Germany).

97 The detection limits of nitrate, nitrite, ammonium, and phosphate ranged from 0.01–0.04 μM ,

98 0.01–0.02 μM , 0.01–0.03 μM , and 0.01–0.02 μM , respectively. The nitracline was defined

99 as the depth where nitrate concentrations increased above 1 μM .

100 **2.2. Nitrogen fixation activity and mannitol enrichment experiment**

101 Nitrogen fixation was determined by the $^{15}\text{N}_2$ gas bubble method (hereafter, the bubble

102 method; Montoya et al., 1996). Samples for incubation were collected in duplicate

103 acid-cleaned 2-L polycarbonate (PC) bottles. The time-zero samples (n=1) were

104 immediately filtered onto precombusted GF/F filters. Two milliliters of $^{15}\text{N}_2$ gas [SI

105 Science Co. Japan, for this gas, contaminations of nitrate, nitrite, and ammonium were

106 determined to be low (< nM level), indicating that the overestimation of nitrogen fixation

107 rates due to the uptake of ^{15}N -labeled contaminants (Dabundo et al. 2014) was minimal

108 (Shiozaki et al., unpublished data)] were injected directly into the incubation bottles through

109 a septum using a gastight syringe. The tracer-added samples were covered with
110 neutral-density screens to adjust the light level and incubated for 24 h in an on-deck incubator
111 filled with flowing surface seawater. After the incubation, the samples were filtered onto
112 precombusted GF/F filters. The isotopic analyses were performed as described previously
113 (Shiozaki et al., 2009). The rate of nitrogen fixation was calculated using the equations of
114 Montoya et al. (1996).

115 To examine the possibility of underestimation of nitrogen fixation as determined by the
116 bubble method (Mohr et al., 2010; Großkopf et al., 2012), we compared the nitrogen fixation
117 rates determined using the $^{15}\text{N}_2$ gas dissolution method (hereafter, the dissolution method;
118 Mohr et al., 2010) with those determined using the bubble method (see above) during the
119 KK-13-6_Sep and KS-14-2_Mar cruises. For the dissolution method, $^{15}\text{N}_2$ -enriched
120 seawater was prepared according to Mohr et al. (2010) and Großkopf et al. (2012). Briefly,
121 filtered seawater was degassed using a Sterapore membrane unit (20M1500A: Mitsubishi
122 Rayon Co., Ltd., Tokyo, Japan) at a flow rate of $\sim 500 \text{ mL min}^{-1}$ (recirculation period, 10 min).
123 Degassed seawater was stored in 1-L Tedlar bags without headspaces and $^{15}\text{N}_2$ gas was added
124 at a ratio of 10 ml $^{15}\text{N}_2$ per 1L seawater. After complete dissolution, the $^{15}\text{N}_2$ -enriched
125 seawater was added to seawater samples contained in 2-L PC bottles, which were incubated
126 and used for isotopic analyses as described above. The $^{15}\text{N}_2$ -enriched seawater was

127 prepared at each station, and was added to the incubation bottles within 1 h after preparation.
128 The nitrogen fixation rate was calculated according to Mohr et al. (2010). For this
129 comparison, triplicate samples were used for both the dissolution and bubble methods.
130 To examine if sugar addition affected nitrogen fixation rates (Bonnet et al., 2013; Rahav et al.,
131 2013; Moisander et al., 2011), we determined nitrogen fixation rates (the ¹⁵N₂-~~gas~~-bubble
132 method, see above) for surface seawater samples (stations ON4 and OT6 during the
133 KS-14-2_Mar cruise) with and without addition of mannitol (final conc. 0.8 μM) (n=3).

134 **2.3. Statistical analysis**

135 Pearson's correlation coefficient was used to examine the relationships between nitrogen
136 fixation activities and environmental variables including temperature, nitrate, ammonium,
137 phosphate, and the ratio of nitrate + nitrite + ammonium to phosphate (N/P ratio) in the entire
138 water column (the data used for the calculation ~~were~~are shown in Table S1). When the
139 nutrient concentration was below the detection limit, the value of the detection limit was used
140 for the analysis. When nitrogen fixation was ~~not detected~~undetectable, the value was
141 assumed to be zero.

142 **2.4. DNA analysis**

143 2.4.1. DNA extraction, sequencing, and phylogenetic analysis

144 Samples (0.38–1 L) for DNA analysis were filtered through 0.2-μm-pore-sized Nuclepore

145 filters and stored in a deep freezer (-80[°]C) until onshore analysis. Total DNA was
146 extracted using a ChargeSwitch Forensic DNA Purification Kit (Invitrogen, Carlsbad, CA,
147 USA) with slight modification of the manufacturer's protocol (Shiozaki et al., 2014a).
148 Partial *nifH* fragments were amplified using a nested PCR strategy (Zehr and Turner, 2001)
149 from samples collected from surface water at Stns. OT4, ON1, ON5, and ON7 during the
150 KT-12-20_Aug and KT-12-27_Oct cruises, at Stns. OT4, ON1, and ON5 during the
151 KT-13-2_Jan and KS-14-2_Mar cruises, at Stns. OT4, ON1, ON5, and ON8 during the
152 KK-13-1_Jun cruise, and at Stns OT4, ON5, ON7 during the KK-13-6_Sep cruise (Table 1).
153 PCR reagents were applied as described by Shiozaki et al. (2014a). The first and second
154 PCRs were run using the same cycling conditions: 95[°]C for 30 s followed by 30 cycles of
155 98[°]C for 10 s, 52[°]C for 30 s, and 72[°]C for 30 s; followed by a final extension at 72[°]C
156 for 7 min. Sterile distilled water was used as the negative control. After PCR analysis, we
157 confirmed ~~there was no band in agarose gel of electrophoresis from that~~ the negative control
158 showed no bands in the gel. The PCR products were cloned and sequenced according to
159 Shiozaki et al. (2014a). The present study obtained 197 *nifH* sequences in total. The *nifH*
160 sequences were translated into amino acid sequences and searched against the protein
161 database of the National Center for Biotechnology Information using the BLASTp algorithm.
162 Clones with 100% amino acid sequence similarity were defined as the same operational

163 taxonomic unit (OTU) using the CD-HIT suite (Huang et al., 2010). The amino acid
164 sequences were aligned using multiple sequence comparisons by the log-expectation
165 (MUSCLE) module in the MEGA5 package (Tamura et al., 2011). A phylogenetic tree was
166 constructed using the maximum likelihood method employing the Dayhoff matrix-based
167 mode, and 1,000 bootstrap replicates were run. The obtained sequences were assigned to
168 bacterial groups based on known sequences included in a cluster within the phylogenetic tree
169 (Zehr et al., 2003a). The sequences from this study were deposited in the DNA Data Bank
170 of Japan (DDBJ) as accession numbers LC013480 to LC013676.

171 2.4.2. Quantitative PCR (qPCR) analysis

172 The clone library analysis showed that UCYN-A, *Trichodesmium*, and γ -proteobacterial
173 phylotype γ -24774A11 (hereafter γ -24774A11) were likely important diazotrophs from early
174 summer to fall when nitrogen fixation occurred (see below). Therefore, the present study
175 quantified these *nifH* phylotypes by qPCR analysis to examine their relative importance
176 ~~during~~ during these seasons. In addition, UCYN-B which is considered to be a major
177 diazotroph in the tropical and subtropical oligotrophic ocean (Moisander et al., 2010), was
178 quantified. TaqMan primer and probe sets previously designed for these four *nifH*
179 phylotypes were used for quantification (Shiozaki et al., 2014a,c; [Moisander et al., 2014](#)).
180 The ~~20-20~~ μ L qPCR reactions contained 10 μ L $2 \times$ Premix Ex Taq (Probe qPCR; Takara), 5.6

181 μL of nuclease-free water, 1 μL each of the forward and reverse primers, 0.4 μL of TaqMan
182 probe, and 2 μL of template DNA. The qPCR assays were performed using LightCycler
183 480 System (Roche Applied Science, Germany). The qPCR assays were run in triplicate
184 reactions. Linear regression r^2 values for the standard curves were >0.99 for all reactions.
185 The efficiency of the qPCR assays ranged from 90.9 to 98.4%, with an average of 95.1%.
186 As the negative control, sterile distilled water was used, from which no amplification signals
187 were detected. The detection limit was 75 copies L^{-1} .

188 **3. RESULTS**

189 **3.1. Comparison of the bubble method and the dissolution method**

190 Nitrogen fixation rates determined by the bubble and dissolution methods were compared
191 during the KK-13-6_Sep and KS-14-2_Mar cruises (Fig. 2). Both methods failed to detect
192 nitrogen fixation in samples collected during the KS-14-2 cruise. During the KK-13-6_Sep
193 cruise, the nitrogen fixation rates determined by the dissolution method were significantly
194 higher (1.5–2.2 fold) than those determined by the bubble method at Stns. OT6 and ON5 (p
195 <0.05). At Stns. OT4 and ON7, the nitrogen fixation rates determined by the two methods
196 did not differ significantly. Thus, the bubble method may have significantly underestimated
197 the nitrogen fixation rates in some, if not all, of the samples that we analyzed. Although the
198 nitrogen fixation rates reported in the rest of this paper are those obtained by using the bubble

199 method, which was used as athe standard protocol during all ~~the~~ cruises, the possibility that
200 some of these rates could be underestimated must be kept in mind. ~~The following nitrogen~~
201 ~~fixation results were obtained by the bubble method; we sought to standardize values among~~
202 ~~all cruises and to compare them with previous results. Hence, the levels could be~~
203 ~~underestimates~~underestimated.

204

205 **3.2. Seasonal variations in nitrogen fixation rates**

206 According to the temperature-salinity (TS) diagram proposed by Hanawa and Mitsudera
207 (1987), both the offshore and bay waters collected during this investigation mostly belonged
208 mostly to either the surface layer water system (SW) or the Tsugaru Warm Current water
209 system (TW) (Fig. 3). ~~Exceptions included the, with the exception of~~ waters collected from
210 the 1% light depth (119 m) at Stn. ON5 during the KT-13-2_Jan cruise (classified as the
211 Oyashio water system (OW)) and those collected at the surface of OT5 during the
212 KS-14-2_Mar cruise (classified as, ~~which were classified as belonging to the Oyashio water~~
213 ~~system (OW) and~~ the Coastal Oyashio water system (CO), ~~respectively~~). These water
214 classifications based on the TS diagram were generally consistent with the geostrophic
215 current field of the investigated region (Fig.S1). Based on these results, it was ~~considered~~
216 assumed that surface waters collected during the same cruise in a particular season generally

217 belonged to the same water system that was prevalent in the investigated region at the time of
218 our sampling.

219 Sea surface temperatures (SSTs) (range, 1.5 ~~to~~ 24.3 °C) (Figs. 4a and S1) and surface
220 nitrate and phosphate concentrations determined during each cruise were averaged to

221 ~~emphasize~~ indicate the seasonal variability of these parameters (Fig. 4b). In general, surface
222 nitrate and phosphate concentrations were low ($\leq 0.07 \mu\text{M}$ and $\leq 0.20 \mu\text{M}$, respectively) in the

223 warmer seawaters (14.2–24.3° C) sampled in early summer (KK-13-1_Jun), summer
224 (KT-12-20_Aug), and fall (KT-12-27_Oct), whereas they were relatively high ($\geq 0.75 \mu\text{M}$ and

225 $\geq 0.28 \mu\text{M}$, respectively) in the colder seawaters (1.5–9.8° C) sampled during winter
226 (KT-13-2_Jan), and spring (KS-14-2_Mar). During the KK-13-6_Sep cruise (late summer),

227 the nitrate concentrations were relatively high and variable (mean \pm SD; $2.92 \pm 7.90 \mu\text{M}$).
228 This was because the highest nitrate concentration (22.6 μM) was determined at the

229 near-shore Stn. OT1 (Fig. S2). Similar to nitrate, surface phosphate concentrations tended
230 to be high during winter (KT-13-2_Jan) and spring (KS-14-2_Mar), while they were low

231 during the warmer seasons. ~~The seasonal variation pattern of the average~~ By contrast,
232 surface ammonium concentrations were generally low ~~at the surface differed from those of~~

233 ~~nitrate and phosphate concentrations (Fig. 4b), characterized by low concentrations~~ ($\leq \sim 1$
234 μM) throughout the year (Fig. 4b), except for the high ammonium concentration determined

235 at Stn. OT1 (1.41 μM) during. ~~The high variation in surface ammonium concentration~~
236 ~~during the KK-13-6_Sep cruises were due to relatively high ammonium concentrations at Stn.~~
237 ~~OT1 (1.41 μM) (Fig. S2).~~

238 ~~Nitrogen fixation was detected in the surface waters of most samples collected~~ During the
239 four cruises conducted in early summer (KK-13-1_Jun), summer (KT-12-20_Aug), late
240 summer (KK-13-6_Sep), and fall (KT-12-27_Oct), nitrogen fixation was measurable in most
241 of the samples collected from surface waters: the nitrogen fixation rates varied and varied in
242 the range of 0.33–13.6 $\text{nmol N L}^{-1} \text{d}^{-1}$ (Figs. 4c and S2). Relatively high nitrogen fixation

243 rates were determined for samples collected during the KT-12-20_Aug cruise, although the
244 highest value was obtained at Stn. ON7 during the KK-13-6_Sep cruise. Nitrogen fixation
245 was ~~not detected~~ below the detection limit in seawater samples collected during the winter and
246 spring cruises. For those samples, nitrogen fixation was unmeasurable undetectable, even after
247 the addition of mannitol (KS-14-2_Mar). ~~Furthermore, n~~ Also, nitrogen fixation was not
248 detected unmeasurable undetectable in ~~DIN-replete a DIN-replete~~ water collected at Stn. OT1
249 in late summer (KK-13-6_Sep).

250 ~~The rates of nitrogen fixation in~~ Nitrogen fixation rates were determined for samples collected
251 from different depths (0–119 m) ~~were examined~~ at Stns. OT4 and ON5 (Fig. 5). Nitrogen
252 fixation was detectable in surface and deeper layers only during ~~the~~ four cruises conducted

253 in early summer (KK-13-1_Jun), summer (KT-12-20_Aug), late summer (KK-13-6_Sep), and
254 fall (KT-12-27_Oct), ~~the same seasons during which surface nitrogen fixation was observed~~
255 (Fig. 4). Nitrogen fixation rates tended to be higher at the surface than in the deeper layers
256 during summer (KT-12-20_Aug) and late summer (KK-13-6_Sep (at Stn. OT4)), whereas this
257 vertical trend was less evident during fall (KT-12-27_Oct) and early summer (KK-13-1_Jun).
258 At Stn. OT4, nitrogen fixation was detectable ~~even in deeper layers~~ in the layers below
259 the nitracline (KT-12-27_Oct, depth = 62 m; KK-13-1_Jun, depth = 42 m). During
260 KK-13-1_Jun cruise, the nitrogen fixation rate determined at the depth of 42 m (1.56 nmol N
261 L⁻¹ d⁻¹) was 1.8 fold higher than the corresponding rate at the surface (0.87 nmol N L⁻¹ d⁻¹).
262 The concentrations of nitrate and ammonium in these layers varied in the range of
263 <0.02–22.5 μM and <0.01–1.41 μM, respectively. ~~where nitrate concentrations were~~
264 ~~relatively high (KT 12 27_Oct, depth = 62 m; KK 13 1_Jun, depth = 42 m). In this layer,~~
265 ~~the ammonium concentrations were 0.05 μM (KT 12 27_Oct) and 0.62 μM (KK 13 1_Jun).~~
266 ~~The nitrogen fixation rate below the nitracline (1.56 nmol N L⁻¹ d⁻¹) was higher than that at~~
267 ~~the surface (0.87 nmol N L⁻¹ d⁻¹) during the KK-13-1_Jun cruise.~~ The maximum
268 depth-integrated nitrogen fixation (294 μmol N m⁻² d⁻¹) was ~~observed~~ found at Stn. OT4
269 during summer (KT-12-20_Aug).

270 3.3. Relationship between nitrogen fixation rates and environmental

271 **variables**

272 Nitrogen fixation rates tended to increase with temperature ($p < 0.01$) (Fig. 6a and Table 2).

273 Nitrogen fixation was detected only when seawater temperatures exceeded 11.7° C, with

274 higher rates ($>6 \text{ nmol N L}^{-1} \text{ d}^{-1}$) noted in waters warmer than 19.5° C. However, there were

275 exceptions to this general relationship between the nitrogen fixation rate and temperature.

276 For example, from the data collected during the KK-13-1_Jun cruise the nitrogen fixation

277 rate was highest at 15.4° C, while it was low (~~undetectable~~below the detection limit) at higher

278 temperatures.

279 Nitrogen fixation rates were negatively correlated with nitrate and phosphate concentrations

280 ($p < 0.01$) (Table 2), ~~whereas they were not significantly correlated with~~ . ~~There was no~~

281 ~~significant correlation between nitrogen fixation rates and ammonium concentrations~~ ($p >$

282 0.05) (Table 2). We also found no significant correlation between nitrogen fixation rates

283 and the ratio of total inorganic nitrogen (nitrate + nitrite + ammonium) to phosphate (Table 2).

284 ~~A plot of the nitrogen fixation against nitrate concentrations indicated that ni~~Nitrogen fixation

285 was generally detectable only when nitrate was depleted (Fig. 6b), except that relatively high

286 nitrogen fixation rates were determined in the subsurface layer of Stn. OT4 (KT-12-27_Oct

287 and KK-13-1_Jun). ~~Active~~High nitrogen fixation rates tended to be detected when

288 ammonium concentrations were low (~~occur at low ammonium concentration~~ $\leq \sim 0.1 \mu\text{M}$),

289 although there was no statistically significant relationship between nitrogen fixation rates and
290 ammonium concentrations. ~~However, seasonal variation in ammonium concentration was~~
291 ~~small and no statistically significant relationship with nitrogen fixation was observed (Fig.~~
292 ~~6e).~~

293

294 **3.4. Seasonal variation in the diazotroph community**

295 3.4.1. Diazotroph community

296 PCR reagents have been suggested to be a potential source of *nifH* genes during analysis of
297 the diazotroph community (Zehr et al., 2003b). Although we confirmed the absence of any
298 bands from the negative control in agarose gel electrophoresis, some sequences recovered
299 from the samples obtained during the KK-13-6 Sep and KS-14-2 Mar cruises (10 clones in
300 total) were judged to be the contaminants in PCR reagents (with similarity (>97% similarity)
301 at the amino acid level was used as a criterion). ~~to contaminants in PCR reagents were~~
302 ~~recovered from samples obtained during the KK-13-6 Sep and KS-14-2 Mar cruises (10~~
303 ~~clones in total).~~—We did not include these sequences in our data analysis.

304 The *nifH* gene was recovered from all the samples that we collected during this study across
305 different stations and seasons (Table 1). Sixty-one OTUs were grouped from 187 *nifH*
306 clones, based on 100% amino acid sequence similarity. The OTUs were assigned to

307 cyanobacteria, α -, β -, γ -, and δ -proteobacteria, and Cluster III diazotrophs (Zehr et al., 2003a)
308 (Figs. S3, S4, and S5).

309 The ~~detected~~recovered cyanobacterial sequences belonged to *Trichodesmium*, UCYN-A, and
310 *Leptolyngbya*. The *nifH* sequences of UCYN-B, UCYN-C, and *Richelia intracellularis*
311 were not recovered. The *nifH* sequence of *Trichodesmium* was recovered only during the
312 KT-12-27_Oct cruise (Table 1). UCYN-A was generally ~~observed~~recovered from early
313 summer to fall, while *nifH* of *Leptolyngbya* was ~~detected~~recovered during winter. The
314 present study detected the sequences of γ -24774A11 during the KT-12-27_Oct and
315 KK-13-6 Sep cruises. This heterotrophic bacterial phylotype is considered to significantly
316 contribute to nitrogen fixation in a wide range of oceanic environments –(Moisander et al.,
317 2014). During the KS-14-2_Mar cruise, all ~~recovered~~of the sequences that we recovered
318 were derived from heterotrophic bacteria, and were dominated by Cluster III diazotrophs at
319 Stns. OT4 and ON5. The Cluster III diazotroph *nifH* sequences were recovered ~~during~~on all
320 cruises except for the KK-13-1_Jun cruise. Note that 58 out of 187 sequences displayed
321 >97% similarity, at the amino acid level, to terrestrial diazotroph sequences derived from soil,
322 mudflats, and lakes (Fig. S3, S4, and S5). These sequences were mainly affiliated with α -
323 and δ -proteobacterial diazotrophs, with 29 of 39 α -proteobacterial sequences and 22 of 24
324 δ -proteobacterial sequences being similar to terrestrial diazotroph sequences.

325 3.4.2. Diazotrophs abundances

326 The *nifH* sequence of *Trichodesmium* was detected by qPCR assay during the KT-12-27_Oct
327 and KK-13-6_Sep cruises (Fig. 7 and 8). During these two cruises, the abundance of
328 *Trichodesmium* ranged from below the detection limit to 8.7×10^4 copies L⁻¹ at all depths.
329 *Trichodesmium* abundance at the surface was higher than those of UCYN-A, UCYN-B, and
330 γ -24774A11 at most stations during the KT-12-27_Oct cruise (Fig. 7 and S6). UCYN-A
331 was detected on all cruises except for the KS-14-2_Mar cruise (Fig. 7 and 8). The
332 maximum abundance of UCYN-A generally occurred at the surface except at Stn. OT4
333 during the KK-13-6_Sep cruise where the peak (1.2×10^3 copies L⁻¹) was observed at 72 m
334 (Fig. 8). The abundance of UCYN-A varied from below the detection limit to 2.6×10^5
335 copies L⁻¹ at all depths. At the surface, UCYN-A was the most abundant among the four
336 groups at most of the stations investigated during the KT-12-20_Aug, KT-13-2_Jan,
337 KK-13-1_Jun, and KK-13-6_Sep cruises (Fig. 7 and S6). UCYN-B was detected only at
338 Stn. ON7 during the KK-13-6_Sep cruise (Fig. 7, 8, and S6). γ -24774A11 was detected
339 during all cruises except for the KS-14-2_Mar cruise (Fig. 7 and 8). The abundance of
340 γ -24774A11 ranged from below the detection limit to 1.8×10^4 copies L⁻¹, with a tendency of
341 a subsurface peak at both stations (Fig. 8).

342

343 **4. DISCUSSION**

344 **4.1. Seasonal variations in nitrogen fixation rates in the temperate coastal**
345 **ocean**

346 Nitrogen fixation rates were measurable mainly from early summer to fall when nitrate was
347 generally depleted in sample seawaters, although there were some exceptions. Our
348 estimates of the nitrogen fixation rates ($0.33\text{--}13.6 \text{ nmol N L}^{-1} \text{ d}^{-1}$) were significantly (p
349 <0.05) higher than the corresponding values previously reported in the temperate region of
350 the eastern North Pacific ($0.15\text{--}0.31 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Needoba et al., 2007) and the
351 oligotrophic region of the western and central North Pacific ($0.17\text{--}3.62 \text{ nmol N L}^{-1} \text{ d}^{-1}$;
352 Shiozaki et al., 2010), whereas they were comparable to those determined in the Kuroshio
353 ($0.54\text{--}28 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Shiozaki et al., 2010) and the western Atlantic coastal regions
354 ($1.3\text{--}49.8 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Mulholland et al., 2012). Higher nitrogen fixation rates have been
355 determined in other temperate oceans, including the western English Channel (18.9 ± 0.01 and
356 $20.0 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Rees et al., 2009) and the Baltic Sea estuaries ($47\text{--}83 \text{ nmol N L}^{-1} \text{ d}^{-1}$;
357 Bentzon-Tilia et al., 2015).

358 In our study, spatiotemporal variability in nitrogen fixation rates appeared to be partly related
359 to the Tsugaru Warm Current path. This current, which flows from the north (after passage
360 through the Tsugaru Strait) to the study region (Fig. S1), may carry active diazotrophs and

361 therefore enhance nitrogen fixation in our study region. This is supported by the fact that
362 nitrogen fixation rates during individual cruises tended to be higher at Stn. OT4 than at Stn.
363 ON5. These stations were located up- and down-stream of the Tsugaru Warm Current,
364 respectively. In addition, variations in nitrogen fixation rates among stations and seasons
365 might also be related to the extent of vertical mixing in the Tsugaru Warm Current. It has
366 been suggested that vertical mixing may introduce iron-rich subsurface water to the surface
367 of the Tsugaru Strait (Saitoh et al., 2008). Such input of iron may enhance nitrogen fixation
368 rates. Consistent with this notion, our results showed that the nitrogen fixation rate was
369 relatively high at Stn. OT4, where the nitracline was relatively deep.

370 Blais et al. (2012) proposed that nitrogen fixation can proceed even in nutrient-replete waters,
371 if large amounts of iron and organic materials are available for consumption by bacterial
372 diazotrophs. In the present study, this possibility was examined by conducting mannitol
373 addition experiments using surface seawaters collected during spring. These waters,
374 belonging to the Oyashio Current system (Nishioka et al., 2007, 2011; Shiozaki et al., 2014b),
375 were considered to be rich in iron during spring, as indicated by a previous study (iron conc.,
376 0.79–8.46 nM; Nishioka et al. 2007). Despite potentially high iron concentrations, our
377 results showed that nitrogen fixation was undetectable even after the mannitol addition,
378 suggesting that, contrary to the Blais et al. proposition, diazotrophs remained inactive under

379 our experimental settings.

380 Our data showed that nitrogen fixation rates were below the detection limit during winter,
381 spring, and late summer (KK-13-6_Sep), when nitrate concentrations were high. These
382 results were consistent with the results of previous studies in the Pacific Ocean, which
383 indicated that nitrogen fixation rates were low or undetectable in DIN-replete waters
384 (Shiozaki et al., 2010). In contrast, Mulholland et al. (2012) reported that, in temperate
385 regions of the Atlantic Ocean, nitrogen fixation rates were high even in DIN-replete ($>1 \mu\text{M}$)
386 and cold ($<10^\circ\text{C}$) surface seawaters. Their study was conducted downstream of the Gulf
387 Stream, where diazotrophs could be delivered from subtropical oceans where DIN is depleted.
388 Previous studies have suggested that cyanobacterial diazotrophs can travel over long
389 distances ($>1,000 \text{ km}$) in currents, without losing their capacity for N_2 fixation (Shiozaki et
390 al., 2013), and that activity is not lost immediately even after mixing with DIN-replete
391 seawaters (Holl and Montoya, 2005; DeKaetzmacker and Bonnet, 2011). In our region,
392 because the Tsugaru Warm Current flows from north to south, diazotrophs entrained by the
393 current have little chance of meeting high-DIN water at the surface. DIN-replete water
394 during summer was observed at the inside bay station OT1 (Fig. S2). Concomitantly,
395 low-salinity surface waters spread offshore along the OT transect line (Fig. S7), suggesting
396 that anomalously high DIN concentrations were likely attributable to terrestrial surface

397 discharge enhanced by Typhoon Man-yi, which passed over the region immediately before
398 the cruise. Subramaniam et al. (2008) reported that nitrogen fixation rates near the Amazon
399 River estuary, with low salinity and high nitrate levels, were fairly low. Their results are
400 consistent with ours. Ammonium inhibits nitrogen fixation, especially when ammonium
401 concentrations exceed 1 μM , as demonstrated ~~for~~by *Trichodesmium* (Mulholland et al. 2001).
402 In our study, ammonium concentrations were generally low ($\leq \sim 1 \mu\text{M}$) throughout the
403 investigation, and no negative relationship between nitrogen fixation and ammonium
404 concentration was found. ~~—This can likely be explained by relatively low ammonium~~
405 ~~concentrations ($\leq \sim 1 \mu\text{M}$) throughout the year and across the investigated region.—~~ Our data
406 showing that nitrogen fixation rates were negatively correlated with nitrate concentrations
407 (Table 2) are consistent with the general notion that nitrogen fixation rates are generally low
408 in nitrate replete waters (Falkowski, 1983). Our data also showed nitrogen fixation rates
409 tended to increase with increasing temperature and with decreasing phosphate concentrations
410 ~~(Nitrogen fixation was also negatively correlated with phosphate as with nitrate, and~~
411 ~~positively correlated with temperature (Table 2). Since correlation between~~ Because ~~t-nitrate,~~
412 ~~phosphate, and temperature~~ and phosphate concentrations were correlated with nitrate
413 concentrations, —was significant, all these factors would not necessarily influence nitrogen
414 fixation directly. Rather, one or more ~~the~~ factors that varied with nitrate could

415 synergistically influence nitrogen fixation.

416

417 **4.2. Seasonal variation in the diazotroph community in the temperate**
418 **coastal ocean**

419 The qPCR analysis demonstrated that the target groups were quantifiable even at stations at
420 which their sequences were not recovered by the clone library analysis, suggesting that the
421 number of clones was not sufficient to capture the diazotroph community structure on each
422 cruise. Despite this limitation, the sequences more frequently recovered in the clone library
423 generally corresponded to the most abundant group revealed by the qPCR analysis. For
424 example, UCYN-A was frequently recovered in the library during the KT-12-20_Aug,
425 KK-13-1_Jun, and KK-13-6_Sep cruises; for these samples, the qPCR results showed that
426 UCYN-A was the most abundant group among the four examined. Similarly, qPCR data
427 indicated that *Trichodesmium* was the most abundant group during fall, when this group was
428 frequently recovered in the library (during the KT-12-27_Oct cruise). Therefore, the
429 diazotrophs targeted by the qPCR analysis were likely important for nitrogen fixation in this
430 study region. In the discussion below, we mainly discuss possible factors responsible for
431 seasonal variation in the diazotrophs targeted by the qPCR analysis.

432 UCYN-A was detected ~~by qPCR~~ in all seasons except spring (KS-14-2_Mar), suggesting that

433 this group of diazotrophs could be important agents of nitrogen fixation in this region.

434 Especially from early to late summer, the abundance of UCYN-A was generally higher than

435 that of *Trichodesmium*, UCYN-B, and γ -24774A11. UCYN-A has been widely detected in

436 temperate regions, and is considered to be one of the major diazotrophs of these locations

437 (Needoba et al., 2007; Rees et al., 2009; Mulholland et al., 2012; Bentzon-Tilia et al., 2015).

438 UCYN-A is known to be most abundant in relatively warm waters around $\sim 20^\circ\text{C}$ (Needoba

439 et al., 2007; Moisander et al., 2010). In our study, UCYN-A was detected~~by qPCR even~~

440 ~~during winter at some stations. It appears that , yet, was not observed during spring. This~~

441 ~~could be because~~ UCYN-A abundance decreased with decreasing temperature from fall to

442 winter, and then became undetectable~~decreased from fall to winter in spring. with~~

443 ~~decreasing temperatures, eventually disappeared in spring.~~

444 *Trichodesmium* was detected from late summer to fall~~by qPCR analysis~~, when water

445 temperatures ranged from 19.1 to 23.4 $^\circ\text{C}$ at the surface. Given that the optimal growth

446 temperature for *Trichodesmium* has been reported to be high (24–30 $^\circ\text{C}$) (Breitbarth et al.,

447 2007), *Trichodesmium* detected in the investigated region likely existed under suboptimum

448 conditions. The relatively high abundance of *Trichodesmium* observed during fall, despite

449 the suboptimal temperature conditions, might indicate that *Trichodesmium* was transported

450 from the adjacent subtropical region where seawater temperatures were high ($>24^\circ\text{C}$). In

451 the western North Pacific subtropical region, *Trichodesmium* is abundant from July to
452 September (Marumo and Nagasawa, 1976; Chen et al., 2008). *Trichodesmium* that
453 flourished in the subtropical region during summer could be transported by the Tsugaru
454 Warm Current, displaying peak abundance during fall in the investigated region. This could
455 support the above discussion that waters containing active nitrogen fixation were delivered to
456 this region by the Tsugaru Warm Current.

457 We ~~observed~~ detected γ -24774A11 ~~by qPCR analysis~~ during all cruises except for the
458 KS-14-2_Mar cruise. γ -24774A11 is considered to be one of the most important
459 heterotrophic diazotrophs in the tropical and subtropical oligotrophic ocean (Moisander et al.,
460 2014). However, ~~This~~ phylotype γ -24774A11 sequence has not been ~~detected~~ reported
461 previously in other temperate oceans (Needoba et al., 2007; Rees et al., 2009; Mulholland et
462 al., 2012). ~~The~~ The ~~*nifH* sequence of~~ γ -24774A11 sequence was similar (94% similarity at
463 the amino acid level) to the *nifH* sequence ~~at~~ of *Pseudomonas stutzeri*, which ~~was~~
464 ~~observed~~ has been reported to be present in temperate estuaries ~~waters including temperate~~
465 ~~regions~~ (Bentzon-Tilia et al., 2015). Bentzon-Tilia et al. (2015) reported that *P. stutzeri*-like
466 *nifH* genes (99% similarity at the nucleotide level) were the most abundant sequences among
467 their samples collected from the ~~temperate~~ Baltic Sea estuary. In the present study, we
468 recovered *P. stutzeri*-like *nifH* genes (>97% similarity at the amino acid level) only at ~~from~~

469 Stn. OT4 during the KT-13-2_Jan cruise by the clone library analysis, and γ -24774A11 was
470 not detected on that occasion by qPCR analysis probably due to the difference in the
471 sequence between γ -24774A11 and *P. stutzeri*. ~~suggesting that γ -24774A11 was not~~
472 ~~quantified as *P. stutzeri* and that *P. stutzeri* could not be a major diazotroph in this study~~
473 ~~region~~ The ecology of γ -24774A11 is still fairly unknown. It remains to be seen, in future
474 studies whether this phylotype contributes to the nitrogen fixation in this region.

475 UCYN-B was not detected ~~by qPCR~~ except at one station. This result is consistent with
476 previous knowledge. UCYN-B becomes abundant with increasing temperature, similar to
477 *Trichodesmium* (Moisander et al., 2010), and is rarely observed in the temperate region
478 (Needoba et al., 2007; Rees et al., 2009; Mulholland et al., 2012; Bentzon-Tilia et al., 2015).
479 Furthermore, UCYN-B abundance is low in shallow nitracline regions (Shiozaki et al.,
480 2014a,c). The nitracline depth in this region (≤ 60 m) was shallower than that of >100 -m
481 depths of regions where UCYN-B is abundant (Shiozaki et al., 2014a). Therefore, although
482 UCYN-B might also have been delivered from subtropical region, it could not survive in the
483 shallower nitracline region.

484 In nitrate-rich water during winter and spring, Cluster III diazotrophs were detected at most
485 of the stations. Furthermore, from early summer to fall, *nifH* sequences of Cluster III
486 diazotrophs were recovered by the clone library analysis in samples from all cruises (except

487 KK-13-1_Jan). Therefore, Cluster III diazotrophs ~~likely presented~~appeared to be present
488 throughout the investigation period~~a year~~. Cluster III diazotrophs are putative anaerobes
489 (Hamersley et al., 2011; Farnelid et al., 2013; Bentzon-Tilia et al., 2014), and hence, they are
490 usually dominant in the diazotrophic community of oxygen-depleted waters (Hamersley et al.,
491 2011; Farnelid et al., 2013) or marine sediments (Bertics et al., 2013). In this study,
492 dissolved oxygen was not depleted ($>3.16 \text{ ml L}^{-1}$) in the upper winter maximum mixed layer
493 depth in this region (~200 m; Shiozaki et al., 2014b) (Fig. S8). Therefore, the Cluster III
494 activity was likely strongly suppressed in the water column because of the high oxygen
495 concentration.

496 Many *nifH* sequences recovered by the clone library analysis were similar to terrestrially
497 derived sequences. These results agree with previous data collected in coastal regions,
498 where terrestrially derived *nifH* sequences were also found (Rees et al., 2009; Mulholland et
499 al., 2012; Blais et al., 2012). We obtained a *Leptolyngbya*-like *nifH* gene during the
500 KT-13-2_Jan cruise. The organism has been found on beaches or coastal land areas (Brito
501 et al. 2012), but not in the open ocean. Because nitrogen fixation was not detected during
502 the KT-13-2_Jan cruise, the organism was considered not to perform nitrogen fixation.

503

504 **5. CONCLUSION**

505 This study demonstrated that nitrogen fixation can and does proceed at high rates, depending
506 on the season, in the temperate coastal region of the northwestern North Pacific, although we
507 failed to detect nitrogen fixation in DIN-replete cold waters. *nifH* sequences were
508 omnipresent and recovered throughout the year, displaying a marked seasonality in their
509 composition. UCYN-A was a major diazotroph during summer, whereas *Trichodesmium*
510 was abundant during fall, despite low temperatures. It has been suggested that
511 *Trichodesmium* was laterally transported from the adjacent subtropical region, which displays
512 high temperatures. Although the Cluster III diazotrophs were recovered almost throughout
513 a year, they were considered to be inactivated in oxic water columns.

514

515 **Author Contributions**

516 T.S., T.N., and K.F. designed the experiment and T.S. collected the samples at sea. T.S.
517 determined nitrogen fixation and nutrient concentrations and analyzed satellite datasets. T.S.
518 and M.I. conducted the genetic analyses. T.S. prepared the manuscript with contributions
519 from all co-authors.

520

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535

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687

688 Table 1. Summary of recovered *nifH* sequences belonging to *Trichodesmium* (Tri), UCYN-A
689 (UA), *Leptolyngbya* (Lep), α -proteobacteria (α -Pro), β -proteobacteria (β -Pro),
690 γ -proteobacteria (γ -Pro), δ -proteobacteria (δ -Pro), and Cluster III (CIII)

Cruise	Station	No. of clones	Cyanobacteria			α -Pro	β -Pro	γ -Pro	δ -Pro	CIII
			Tri	UA	Lep					
KT-12-20_Aug summer	OT4	12		9		3				
	ON1	5		2					3	
	ON5	8		8						
	ON7	7		1		6				
Total		32	0	20	0	9	0	0	3	
KT-12-27_Oct fall	OT4	7	1						6	
	ON1	9						4(2)	5(5)	
	ON5	6					1		5	
	ON7	13	6	1		5(5)		1(1)		
Total		35	7	1	0	5(5)	0	2(1)	4(2)	
KT-13-2_Jan winter	OT4	11			10		1			
	ON1	1							1	
	ON5	14				5(5)		2(2)	7	
Total		26	0	0	10	5(5)	0	1	2(2)	
KK-13-1_Jun early summer	OT4	10		2		8(8)				
	ON1	15		3			2	10(10)		
	ON5	11		4		7(7)				
	ON8	1					1			
Total		37	0	9	0	15(15)	1	2	10(10)	
KK-13-6_Sep late summer	OT4	7						4(4)	1	
	ON5	11		11						
	ON7	10		2		1		7		
Total		28	0	13	0	1	0	7	4(4)	
KS-14-2_Mar spring	OT4	10						1(1)	9	
	ON1	13				3(3)	3	1(1)	3(3)	
	ON5	15				2(2)			9	
Total		38	0	0	0	5(5)	3	1(1)	4(4)	

691 Numbers in parentheses indicate the number of sequences with >97% similarity at the amino
692 acid level to terrestrial diazotroph sequences.

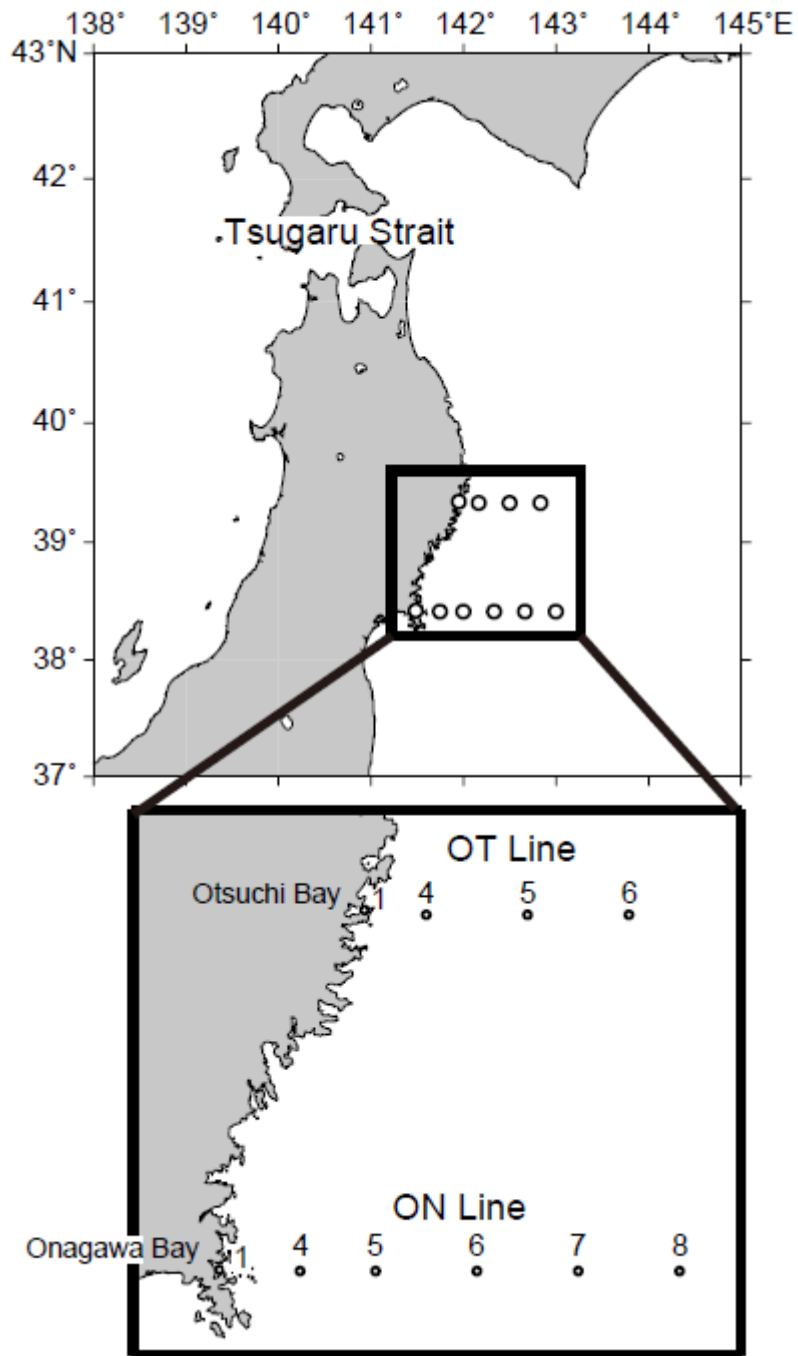
693 Table 2 Pearson's correlation matrix of N₂ fixation rates and water properties in the entire
 694 water column (n=73).

	Temperature	Nitrate	Ammonium	Phosphate	N/P ratio	N ₂ fixation
Temperature	1					
Nitrate	-0.722**	1				
Ammonium	-0.036	0.439**	1			
Phosphate	-0.880**	0.881**	0.119	1		
N/P ratio	-0.266*	0.722**	0.751**	0.349**	1	
N ₂ fixation	0.435**	-0.325**	-0.122	-0.351**	-0.219	1

695 * $p < 0.05$, ** $p < 0.01$

696 N/P ratio denotes the ratio of (nitrate + nitrite + ammonium) to phosphate

697

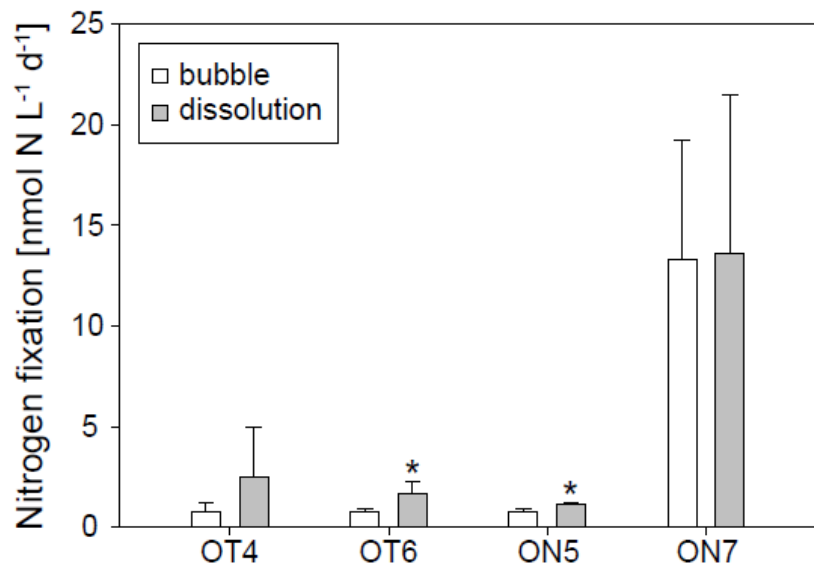


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700 Fig. 1. Sampling locations in the northwestern North Pacific Ocean.

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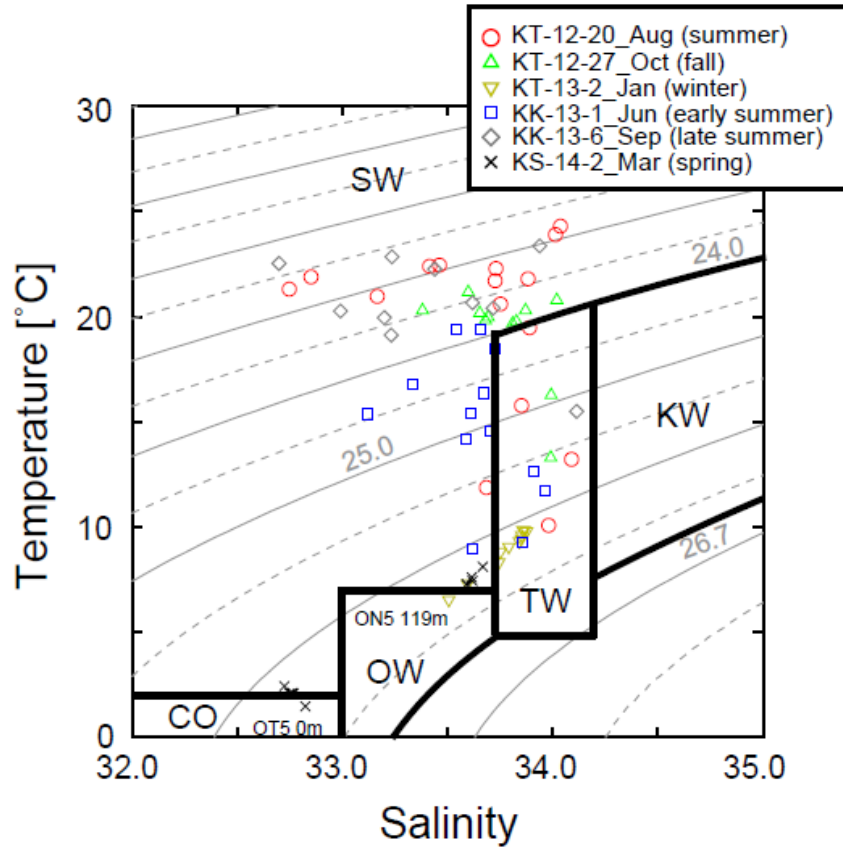
703

704 Fig.2. Nitrogen fixation rates estimated simultaneously by the ¹⁵N₂ gas bubble and

705 dissolution methods during the KK-13-6_Sep cruise. An asterisk indicates a significant

706 difference between the two methods ($p < 0.05$).

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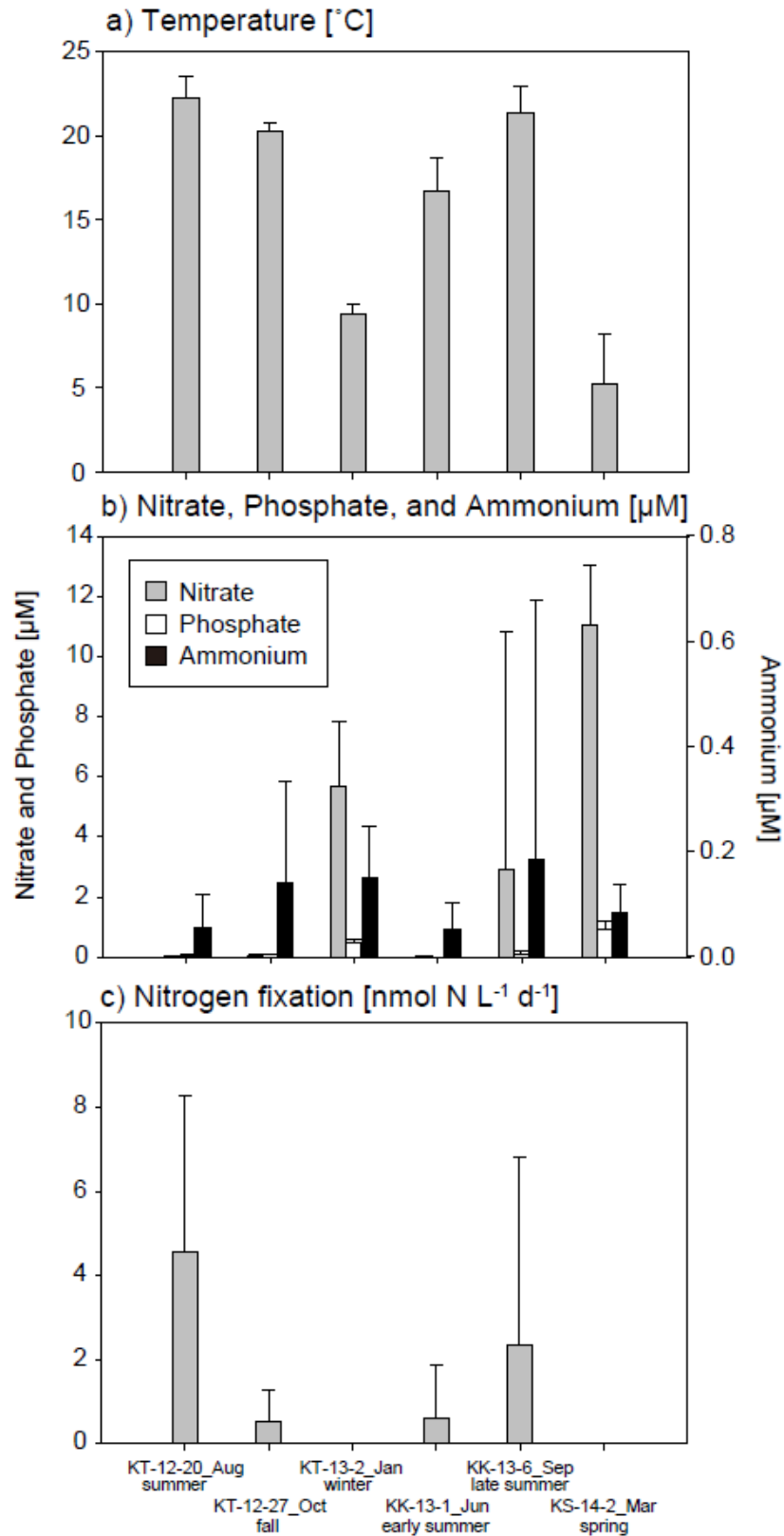


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710 Fig. 3. Temperature-salinity diagram at each sampling point. The water classification was
 711 defined by Hanawa and Mitsudera (1986). SW, KW, TW, OW, and CO denote the surface
 712 layer water system, Kuroshio water system, Tsugaru Warm Current water system, Oyashio
 713 water system, and Coastal Oyashio water system, respectively.

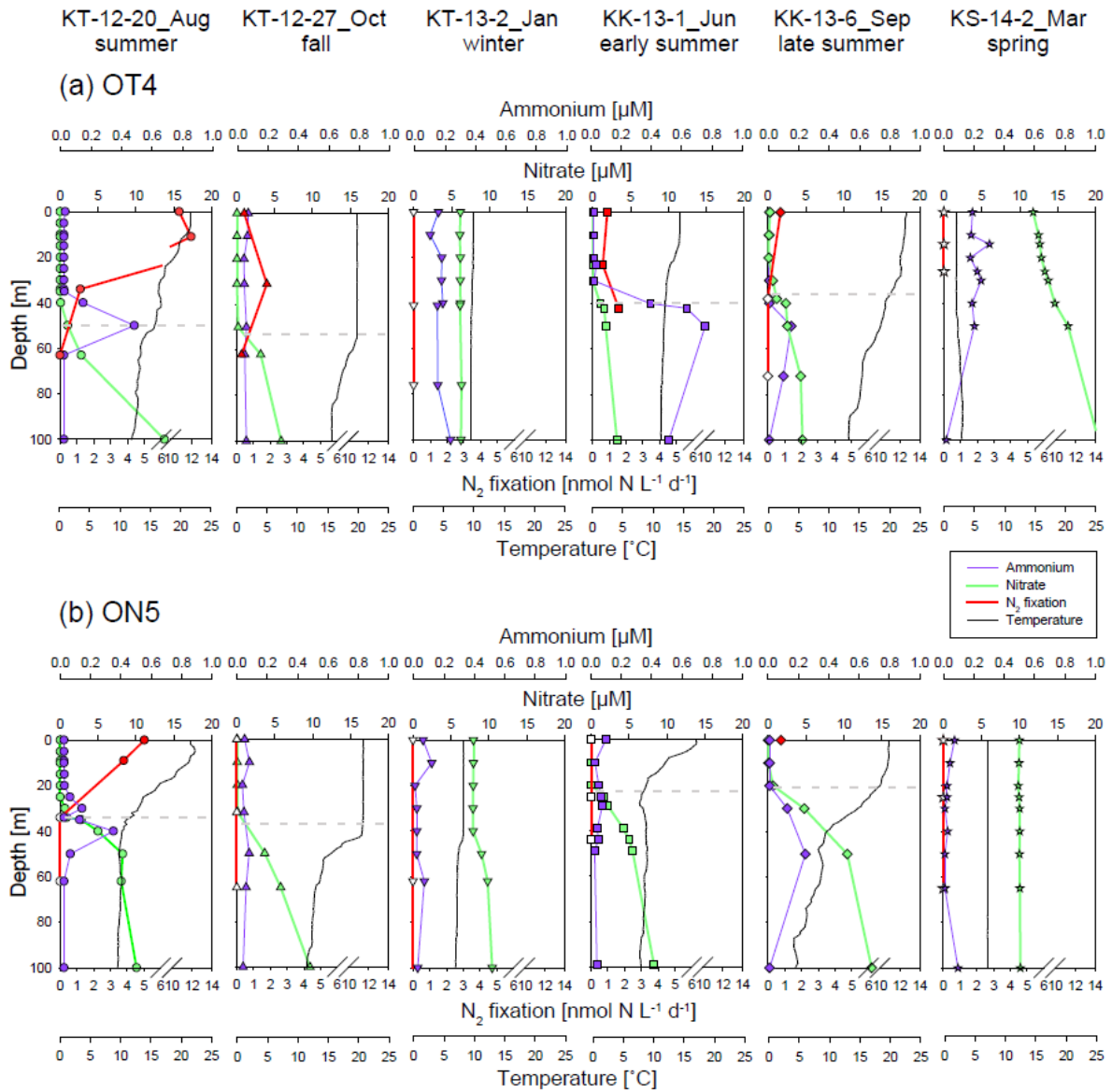
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716 Fig. 4. Average (a) temperature [°C], (b) nitrate, phosphate, and ammonium concentrations

717 [μM], and (c) nitrogen fixation [nmol N L⁻¹ d⁻¹] at the surface during each cruise.



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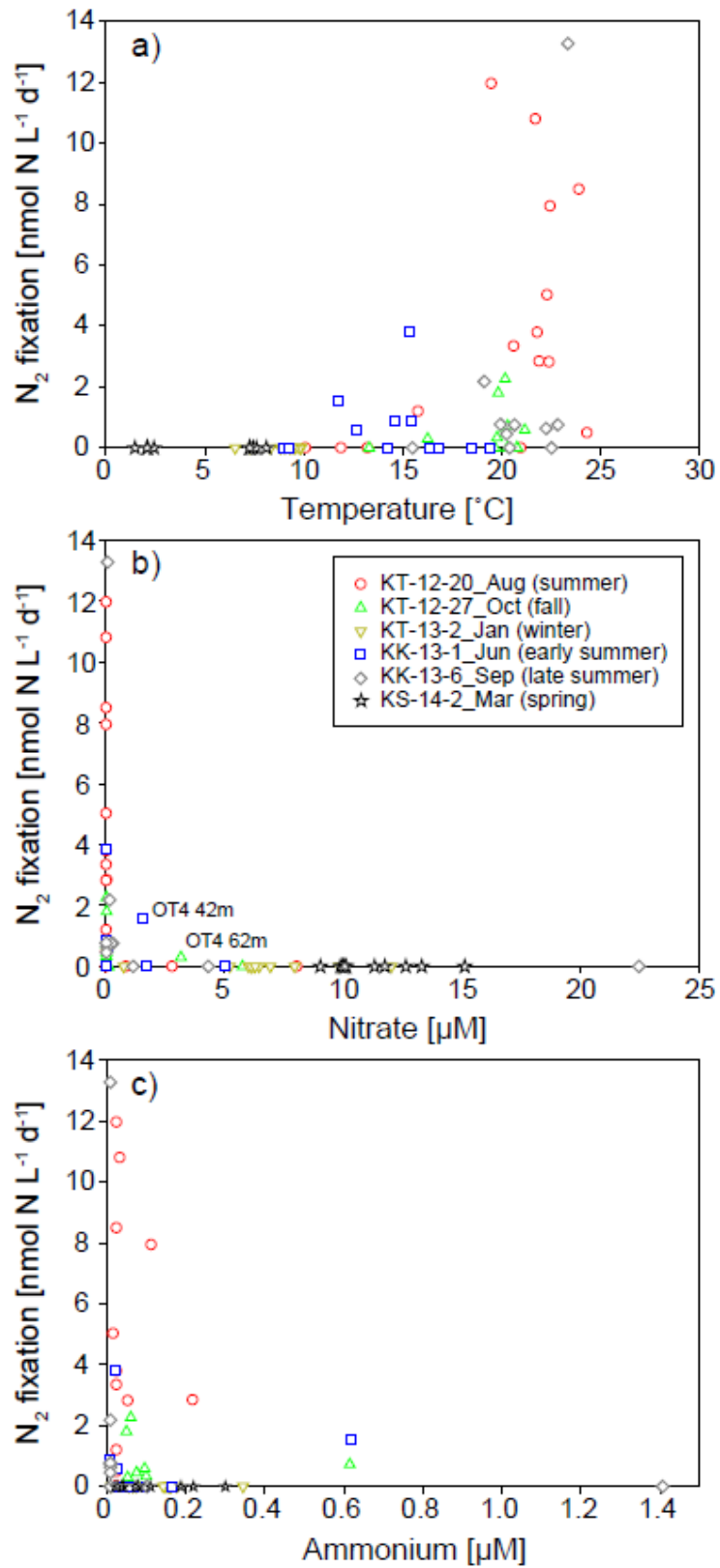
720 Fig. 5. Time-series variations in the vertical profiles of temperature [°C] (black),

721 ammonium (purple) and nitrate (green) concentration [μM], and nitrogen fixation (red) [nmol

722 N L⁻¹ d⁻¹] at Stns (a) OT4 and (b) ON5. Open symbols indicate that nitrogen fixation was

723 not detected. The horizontal dashed line indicates the nitracline depth. The straight lines of

724 temperature and nitrate were ascribable to strong mixing.

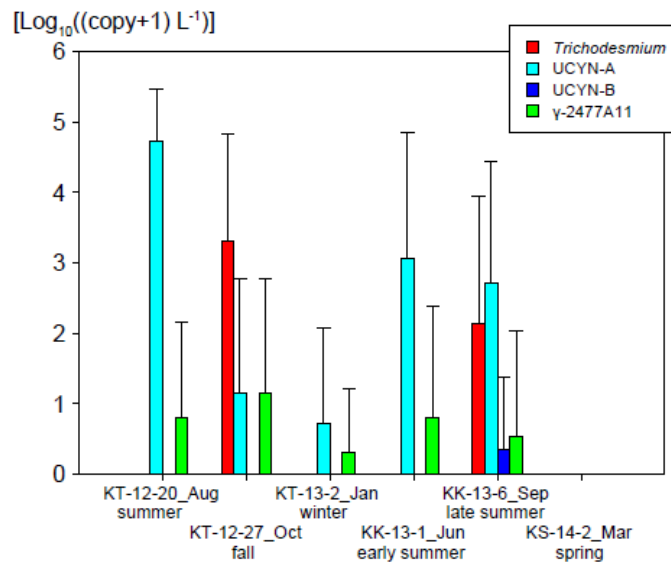


725

726 Fig. 6. Relationship between nitrogen fixation [$\text{nmol N L}^{-1} \text{d}^{-1}$] and (a) temperature [$^{\circ}\text{C}$], (b)

727 nitrate [μM], and (c) ammonium [μM] for all six cruises.

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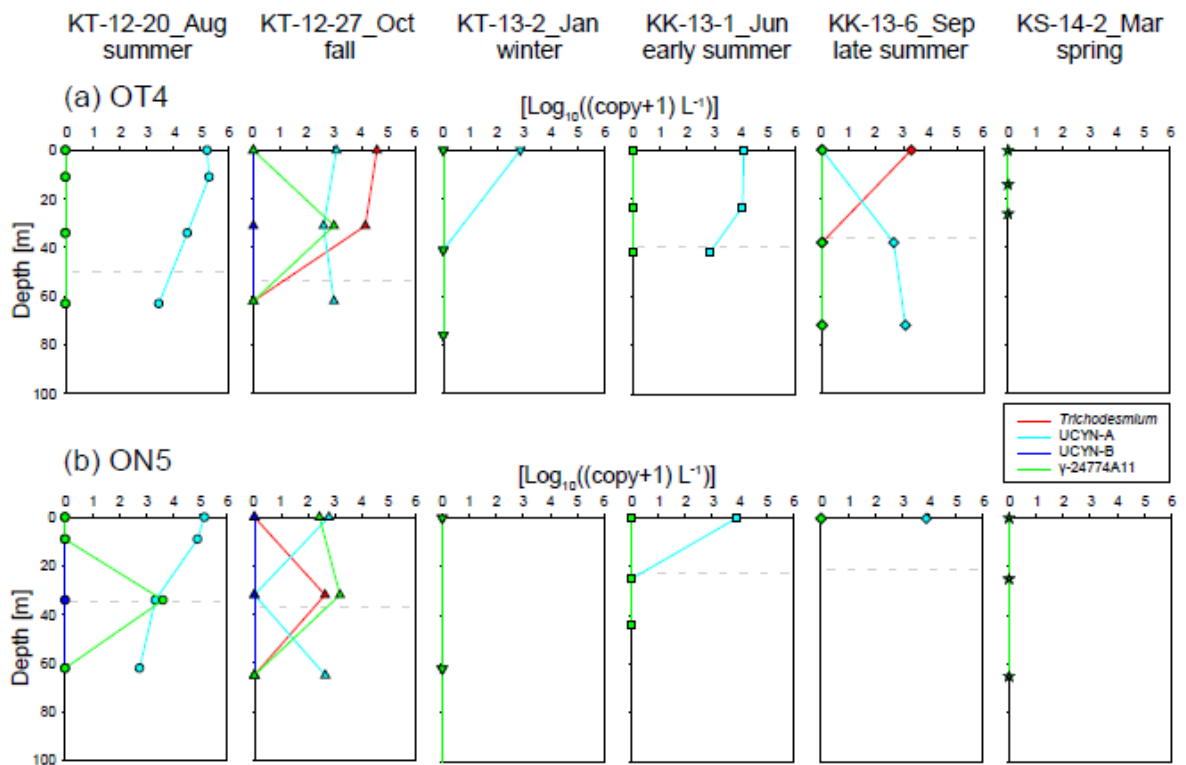
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731 Fig. 7. Average abundances of *Trichodesmium* (red), UCYN-A (light blue), UCYN-B

732 (blue), and γ-24774A11 (green) [$\text{Log}_{10}(\text{copy}+1) \text{ L}^{-1}$] at the surface during each cruise.

733 When the target *nifH* gene was not detected, the copy number was assumed to be zero.

734



735

736

737 Fig. 8. Time-series variations in the vertical profiles of *Trichodesmium* (red), UCYN-A

738 (light blue), UCYN-B (blue), and γ -24774A11 (green) [$\text{Log}_{10}(\text{copy}+1) \text{L}^{-1}$] at Stns. (a) OT4

739 and (b) ON5. The horizontal dashed line indicates the nitracline depth.