Dear Dr. Jens-Arne Subke,

Many thanks to you, all anonymous referees and public reviewer Dr. Ge, for your helpful and valuable comments which we appreciate and have been considered carefully in the revised manuscript. We here submit a revised manuscript that incorporates and highlights all reviewers' comments. The authors' detailed responses to those comments were listed below and italicized in blue color (please note that the page numbers in the response refer to those in the revised manuscript with changes marked-up).

With best regards, Tianshan and all co-authors

## Author response (bgd-12-9465-2015) Response to Dr. Ge's comments (SC C3400):

General comments:

Wang B. and the co-authors do a good job with the high quality dataset of microtopographic variation in soil respiration in a desert-shrub ecosystem. This might be the first such study in terms of regulation of plant phenology in spatial variation in soil respiration. Accordingly, their findings are important to our understanding of biophysical control of soil respiration in the dune. Their results are robust, and the methods are well described.

I have only a few concerns for the manuscript. Firstly, I think it would be good to include in the introduction the references of Tamai (2010), Kang et al. (2003) and Yuste et al. (2004), whose reports showed the effects of environmental factors, soil properties and plant phenological patterns on topographic variations of soil respiration in different ecosystems. Secondly, I cannot understand why the dataset for Phase I was missing in some result sections, such as Fig. 3, Fig. 4, Fig 5, Fig. 6 and Fig. 8. Some related explanations are needed. I suggest combining Fig. 5 and Fig. 6 as a clear compared list. In Table 1, the Phases need to be defined. In Fig. 3, the unit of LAI is incorrect. In Fig. 7, the legend of (a) and (b) need to be defined.

References: Tamai K. 2010. Effects of environmental factors and soil properties on topographic variations of soil respiration. Biogeosciences, 7, 1133–1142. Kang S, Doh S, Lee D,L Dowon, Jin VL, Kimball JS. 2003. Topographic and climatic controls on soil respiration in six temperate mixed-hardwood forest slopes, Korea. Global Change Biology, 9, 1427–1437. Yuste JC, Janssens IA, Carrara A, Ceulemans R. 2004. Annual Q<sub>10</sub> of soil respiration reïn<sup>°</sup>C' ects plant phenological patterns as well as temperature sensitivity. Global Change Biology, 10, 161–169.

### Specific comments:

Firstly, I think it would be good to include in the introduction the references of Tamai (2010), Kang et al. (2003) and Yuste et al. (2004), whose reports showed the effects

of environmental factors, soil properties and plant phenological patterns on topographic variations of soil respiration in different ecosystems.

Answer: We agree. We added the three references in Page 4, Line 7-8 in the revised manuscript.

Secondly, I cannot understand why the dataset for Phase I was missing in some result sections, such as Fig. 3, Fig. 4, Fig 5, Fig. 6 and Fig. 8. Some related explanations are needed.

Answer: Our measurements in Phase I were only 9 days, we think it cannot well represent for whole Phase I, thus we did not show dataset for phase I in our analyses. We added 'Due to fewer measurements in Phase I, data analyses were mainly concentrated on Phase II and Phase III.' in Page 7, Line 17-18 in the revised manuscript.

I suggest combining Fig. 5 and Fig. 6 as a clear compared list.

Answer: We agree. We combined Fig. 5 and Fig. 6. See Page 32 in the revised manuscript.

In Table 1, the Phases need to be defined.

Answer: We agree. We added the definition of Phases as 'Phase I: leaf expansion; Phase II: flowering-bearing; Phase III: coloration-defoliation' at the end of caption of Table 1. See Page 23, Line 6-7 in the revised manuscript.

In Fig. 3, the unit of LAI is incorrect.

Answer: We corrected it with  $m^2 m^{-2}$  in Fig.3b. See Page 29 in the revised manuscript.

In Fig. 7, the legend of (a) and (b) need to be defined.

Answer: We agree. We added legend (a) and (b) to caption of Fig.7. See Page 34 in the revised manuscript.

#### **Response to referee # 1's comments (RC C4050):**

General comments:

The manuscript of Wang et al., titled "Micro-topographic variation in soil respiration and its controlling factors vary with plant phenophases in desert-shrub ecosystems": is an interesting observational study that investigated the spatial variability of Rs in desert ecosystem in relation to plants. While the sampling is limited, due to logistical constraints, the observations are still valuable. I agree with the other referee that the missing Phase I data should be excluded, it does not add much to your results and discussion, where you discuss only Phase II and III of the study. I would also suggest in the future to use portable IRGA system to sample and measure soil respiration in replicate dunes or increasing the sample points at this dune. The LI-8100 system is good for high frequency temporal data, but in this study you do not utilize that capacity you focus on spatial variability and relating fluxes to explanatory variables (ex. LAI, root biomass) that were measured on biweekly or longer timescales. So you could've sampled Rs also on biweekly timescale using manual portable system to increase spatial area of study. It would've been nice to have a larger sample size in this study, not to take away from this one. The current study is still informative - a synthesis of multiple variables measured concurrently at the site, as the authors explore the causes of spacial variability in Rs.

Answer: We agree with you. We excluded Phase I data in our study (see Page 23 in the revised manuscript). Thank you for your previous advices, you suggestions of how to increase sampling will be considered in our future field measurements.

#### **Response to referee # 3's comments (RC C4554):**

### General comments:

It is a very written and concise paper, addressing micro-topographic variations in soil respiration and its controlling factors. This paper uses a well-designed and collected dataset to illustrate the variations caused by the plant biomass variation across sand dune. The information will be useful to the stabilization of the sand dune in the desert region. My minor suggestion to improve this paper is to ask the authors to further exemplify the potential application of this research in practical aspects. For example, how this study can be helpful in relating to the big issue - desertification.

In term of the microsite around a sand dune, the depression area around the sand dune is completely different from the sand dune slope. For example, there is a higher water level and different pH compared with sand slope. Soil texture is completely different. Therefore, the natural vegetation shows great variations across sand dune, which results from multiple processes (wind blowing, seed transportation, soil water movement during a year, and soil nutrient dynamics, soil crust formation, biological processes within the soil). In particular, the soil crust in the desert could change the gas release pattern of soil. The plant coverage can also have critical feedback impacts on the crust depth and structure. Therefore, the soil respiration difference could be a combined result of multiple processes. Do you have any finding which could illustrate the relationships between the crust depth and plant coverage or partitioning the component soil respiration.

A few other suggestions are attached in the supplement file.

Answer: The main purpose of our study is to investigate the controls of micro-spatial variation in soil respiration across a sand dune, thus we exemplify the potential application of our results in C cycle models but not for desertification. However, our results showed much higher plant cover on leeward slope than other slopes, which potentially indicated a better condition for plant growth on this slope, and suggests revegetation in desert ecosystem may consider this specific slope of sand dunes.

We agree with you that soil crust in the desert could change the gas release pattern of soil. However, the microsites we measured were with no crust, thus the influences of crust on soil respiration were not considered in our paper.

Specific comments (bgd-12-C4554-2015-supplement):

P9470 L9-12: This does not make any sense because only individual average size

without density does not tell the coverage. In the desert, the vegetation coverage has a critical impacts on the soil crust which has critical effects on the soil  $CO_2$  efflux. I would convert a coverage percentage.

Answer: We agree. We added 'The coverage percentage of plant in this area ranged from 30 - 60%.' next to this sentence. (see Page 5, Line 13-14 in the revised manuscript)

P9471, L3: There is normally a depression area between sand dunes. There is a significant differences between the depression areas and sand dune slope. Do you have any measurement point(s) to address this?

Answer: No, we did not have the measurements on depression area between sand dunes. This point will be considered in our future measurements at larger special scale, e.g. inter-dune variation in soil respiration.

### P 9473, L 10: Is the soil nitrogen plant related factor?

Answer: In our study, we found good relationship between soil nitrogen and litter fall (Figure 7b), which suggested the decomposition processes of litterfall were the important sources of soil nitrogen. Therefore, we attributed the soil nitrogen to plant related factor in our site.

#### P 9473, L 15: I would say locations instead of slopes.

Answer: The slopes here represent for different orientations of the sand dune. Therefore, we prefer using word 'slopes'.

#### P 9473, L 16: significant?

Answer: Yes, it is statistically significant. We added the p value (p < 0.001) there. (see P 9473, L 16 in the revised manuscript)

### P 9475, L 17: space?

Answer: We agree. We corrected 'earlierthat' to 'earlier that'. (see Page 9, Line 2 in the revised manuscript)

#### P 9476, L 16-20: provide some justifications about why you have this comparison.

Answer: Our results show the contributor to the spatial variation in soil variation in desert was different from that in forests and grassland. We suggested future studies on spatially scaling up soil respiration in desert ecosystems should consider the spatial variation in substrate supply more than temperature. This statement was added in the revised manuscript.(see Page 12, Line 5-6 in the revised manuscript)

#### P 9476, L 25: What are they?

Answer: *Here, the photosynthesis-related parameters refer to root biomass and litterfall. We clarified the sentence (Page 12, Line 11) in the revised manuscript.* 

P9477, L 10: Any concerns about the reason to cause the variations in plant distribution? To my knowledge, it is the soil water, nutrient and the microsite that has led to the distribution of plants along a sand dune.

Answer: We agree that soil water, nutrient and the microsite that has led to the

distribution of plants at earlier stage of plant formation on a sand dune, thus plants subsequently affecting spatial variation of soil respiration.

P 9478, L 3: ? Answer: We deleted 'to'. (see Page 13, Line 17 in the revised manuscript)

#### P9485: what are the time scale for the means?

Answer: The time scale for the means was showing in Figure 2 in our study. We added this statement (see Page 23, Line 2-3) in the revised manuscript.

P9486: Time scale?

Answer: we added time scale 'over the measurement period' at the end of this sentence. (see Page 24, Line 2 in the revised manuscript)

P9486: Leeward has higher biomass and root mass, of course, higher soil respiration, even more soil biota activity as well because of the higher soil moisture, relatively lower soil temperature which are the critical factor in the desert ecosystem. It is the combination of relative.

Answer: we agree with you.

P9486: north-facing has a higher temperature than south-facing, any reason? What is the time scale for the mean?

Answer: In our study, plant distribute sparsely in both north- and south-facing. However, litterfall in south-facing was about three times higher than that in north-facing, which may exert stronger shading effects in south-facing, thus resulting in lower soil temperature in south-facing. The time scale for the mean is the whole measurement period in our study.

P9487: the "x" in the equation is not needed in my opinion.

Answer: we agree. The " $\times$ " in the equation was excluded. (see Page 25 in the revised manuscript)

P9488: Is there any site located in the depression area between sand dune?

Answer: No, we don't have site in our study located in the depression area between sand dunes. The between-dune sites will be considered in our future measurements at larger special scale, e.g. inter-dune variation in soil respiration.

P9490: Time scale?

Answer: Time scales of each phenophases were showed in figure 2. We added this statement (see Page 29, Line 6) in the revised manuscript.

#### P9490: represents?

Answer: we agree. We changed 'represent' to 'represents'. (see Page 29, Line 9 in the revised manuscript)

P9491: Table 3 can be combined into this figure, making the figure easy to understand.

Answer: The relationships between soil respiration and other variables are also

showing in Table 3, but no in Figure 4. Thus we would like to keep them both.

P9495:' thresholds'

Answer: we changed 'thresholds' into 'critical values'. (see Page 35, Line 5 in the revised manuscript).

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This discussion paper is/has been under review for the journal Biogeosciences (BG). Please refer to the corresponding final paper in BG if available.

## Micro-topographic variation in soil respiration and its controlling factors vary with plant phenophases in a desert-shrub ecosystem

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## Abstract

Soil respiration ( $R_s$ ) and its biophysical controls were measured over a fixed sand dune in a desert-shrub ecosystem in northwest China in 2012 to explore the mechanisms controlling the spatial heterogeneity in  $R_s$  and to understand the plant effects on the spatial variation in  $R_s$  in different phenophases. The measurements were car-

- <sup>5</sup> on the spatial variation in  $R_s$  in different phenophases. The measurements were carried out on four slope orientations (i.e., windward, leeward, north- and south-face) and three height positions on each slope (i.e., lower, upper, and top) across the phenophases of the dominant shrub species (*Artemisia ordosica*). Coefficient of variation (i.e., standard deviation/mean) of  $R_s$  across the 11 microsites over our measure-
- <sup>10</sup> ment period was 23.5%. Soil respiration was highest on the leeward slope, but lowest on the windward slope. Over the measurement period, plant-related factors, rather than micro-hydrometeorological factors, affected the topographic variation in  $R_s$ . During the flowering-bearing phase, root biomass affected  $R_s$  most, explaining 72% of the total variation. During the leaf coloration-defoliation phase, soil nitrogen content affected  $R_s$
- the most, explaining 56 % of the total variation. Our findings highlight that spatial pattern in  $R_s$  was dependent on plant distribution over a desert sand dune, and plant-related factors largely regulated topographic variation in  $R_s$ , and such regulations varied with plant phenology.

### 1 Introduction

- <sup>20</sup> Soil respiration ( $R_s$ ) plays an important role in the global carbon (C) cycle (Bond-Lamberty and Thomson, 2010a). Even relatively small increases in  $R_s$  may have a profound impact on atmospheric CO<sub>2</sub> concentrations, exerting a positive feedback to global warming (Schlesinger and Andrews, 2000; Davidson and Janssens, 2006; Luo, 2007). However, estimating soil C release is usually difficult at large spatiotemporal scales, mainly due to the strong heterogeneity in  $R_s$  (Bond-Lamberty and Thomson,
- 2010b; Luo et al., 2012; Jia et al., 2013). Such heterogeneity depends strongly on the



variability of substrate of respiration (e.g., soil organic matter and living biomass (Martin and Bolstad, 2009; Geng et al., 2012; Ngao et al., 2012)) and hydrometeorological factors (e.g., soil water content (SWC) and soil temperature ( $T_s$ ); Marrin and Bolstad, 2009; Ngao et al., 2012).

- <sup>5</sup> Desert ecosystems cover up to 45% of the global terrestrial surface (Asner et al., 2003). The C cycle in desert ecosystems cannot be ignored in the global C cycle (Austin, 2011; Poulter et al., 2014). *R*<sub>s</sub> represents the second largest C flux between ecosystems and the atmosphere (Raich and Schlesinger, 1992; Raich et al., 2002), and exhibits high spatial variation in desert ecosystems (Mahall et al., 1992; Gold et al.,
- <sup>10</sup> 1995; Xie et al., 2001). In addition, active, semi-active aeolian, and fixed sand dunes are the primary features of desert landscapes. Since the greatest amount of living biomass occurs on fixed sand dunes (Roles et al., 2001), its  $R_s$  contributes significantly to C emissions to the atmosphere. Thus, for accurate regional C estimates, there is an urgent need to understand the mechanisms driving spatial pattern in  $R_s$  over fixed sand
- 15 dunes.

Generally, topographic variation in  $R_s$  in arid and semi-arid areas can be explained by hydrometeorological factors alone, especially by SWC (Xu and Wan, 2008; and Liu et al., 2010). This is because  $T_s$  and SWC are reported to have a strong primary control on the spatiotemporal variation in  $R_s$  (Lloyd and Taylor, 1994; Davidson et al., 1998;

- <sup>20</sup> Wang et al., 2014), as they both exhibit strong topographic heterogeneity as a result of the influence of topography in governing the amount of solar radiation received and the re-distribution of surface and shallow subsurface water (Kang et al., 2003; Liu et al., 2010). Heterogeneity in micro-hydrometeorological factors also has an explicit role in defining the formation and patchy distribution of vegetation growing in an area (Rich-
- erson and Lum, 1980; Parker, 1991). Distribution of vegetation, in turn, constrains the spatial variation in  $R_s$  through its patchy supply of substrate to  $R_s$ . In desert ecosystems, such plant effects usually refer to a phenomenon called "fertility islands". This phenomenon is observed by the enrichment of water, nutrient, and root biomass in vicinity of plant canopies (Hook et al., 1991; Schlesinger and Raikes, 1996), resulting



from several plant-scale processes, including plant-facilitated entrapment of soil particles and organic matter propelled by wind or rain droplets, canopy shading, hydraulic lift, and accumulation and decomposition of litterfall (Schlesinger and Raikes, 1996; Stock et al., 1999; Thompson et al., 2005; Li et al., 2007). Consequently, plant controls
 on *R*<sub>s</sub> could be nested within micro-hydrometeorological controls, leading to greater complexity in plant-soil interaction.

Moreover, both micro-hydrometeorological and plant control could vary with plant phenophases (Fu et al., 2002; Dungan et al., 2003; Tang et al., 2005; Asaeda and Rashid, 2015; Osono, 2014), as current photosynthate supply and decomposition and supplementation of litterfall changes over time. For example, root activity and rhizo-

- <sup>10</sup> supplementation of litterfall changes over time. For example, root activity and rhizosphere processes could vary with plant phenophases (Fu et al., 2002), and potentially alter the relationship between  $R_s$  and  $T_s$  (Yuste et al., 2004; DeForest et al., 2006). Previous studies have focused on the influences of plant phenology on the temporal dynamics of  $R_s$ . However, understanding the temporal changes in effects of plants on the spatial variation in  $R_s$  is still an important gap in our knowledge concerning plant–
- the spatial variation in  $R_s$  is still an important gap in our knowledge concerning plant soil interactions in desert environments.

We hypothesized that over a fixed sand dune: (1) topographic heterogeneity in  $R_s$  is controlled to a large measure by the distribution of substrate of respiration supplied by the annual casting of plant photosynthates and nitrogen compounds during leaf-fall,

and (2) drivers of topographic variation in  $R_s$  vary with plant phenology. To test our hypothesis,  $R_s$  and associated micro-hydrometeorological and plant-related factors were measured over a fixed sand dune in a desert-shrub ecosystem in northwest China.

#### 2 Materials and methods

#### 2.1 Site description

The study was conducted on a fixed sand dune at Yanchi Research station of Beijing Forestry University, Ningxia, northwest China (107.20–107.26° E, 37.68–37.73° N,



1550 ma.s.l.). The site is located at the edge of the Mu Us desert in a transition zone between arid and semi-arid climatic conditions. Based on 51 years (1954–2004) of data from the meteorological station of Yanchi County, the mean annual temperature is 8.1 °C. The mean annual precipitation is 292 mm (with a range of 250–350 mm), of which 62 % falls between July and September, and the mean annual total potential evaporation is 2024 mm (Wang et al., 2014). The sandy soil has a bulk density of 1.6 ± 0.02 g cm<sup>-3</sup>. The water-filled pore space across all the microsites was 23.42±0.92 %. Soil organic matter and CaCO<sub>3</sub> content, and pH were 0.21–2.14, 0.23–0.54 g kg<sup>-1</sup>, and 7.76–9.08, respectively (Feng et al., 2013). The vegetation at our study site was regenerated since aerial seeding in 1998 and was dominated by semi-shrub species of is *Artemisia ordosica*, which averaged 50 cm tall with a canopy size of about 80 cm × 60 cm.

#### 2.2 Measurements of soil CO<sub>2</sub> efflux and micro-hydrometeorological factors

Soil CO<sub>2</sub> efflux ( $R_s$ , µmolCO<sub>2</sub>m<sup>-2</sup>s<sup>-1</sup>) was measured with 11 automated chambers from June to October 2012. The measurements were taken on different slope orientations (i.e., windward, leeward, and north- and south-facing slopes) and positions (i.e., lower, upper, and top positions) on a typical shrub-dominated sand dune. The sand dune was of typical size for the study area, with a (i) height of 2.9 m, (ii) 7.6 m long leeward slope, (iii) 15.6 m long windward slope, (iv) 13.3 m long north-facing slope, and

- (v) 13.4 m long south-facing slope. Soil respiration was measured in situ using an automated chamber system (model Li-8100A, equipped with Li-8150 multiplexer, LI-COR, Nebraska, USA). Measurements were made along two 26 m line transects that intersected at perpendicular angles at the centre of the sand dune (Fig. 1). The R<sub>s</sub> system included the installation of 11 permanent opaque chambers set on PVC collars (model
- Li-104, LI-COR, Nebraska, USA). The collars, with a diameter of 20.3 cm and a height of 10 cm, were inserted into the soil at a 7 cm depth. The gas chambers were located at about 3 to 5 m apart along the line transects, each placement representing different sand-dune microsites: top (position 1); upper parts of both the windward (position 2)



and leeward slopes (position 7); lower parts of both the windward (position 3) and leeward slopes (position 8); and the microsites from top to the bottom of both the south-(position 4, 5, 6) and north-facing slopes (position 9, 10, 11; Fig. 1).

Soil respiration measurement time for each chamber was 3 min and 15 s, including a 30 s pre-purge, a 45 s post-purge, and a 2 min observation period. Any plant re-growth within the measurement collar was manually removed in a timely manner. Hourly soil temperature ( $T_s$ , °C) and volumetric soil water content (SWC, m<sup>3</sup> m<sup>-3</sup>) were measured simultaneously at a 10 cm depth outside of each chamber using 8150–203 soil temperature and EC<sub>H<sub>2</sub>O</sub> soil moisture sensors (LI-COR, Nebraska, USA). For more details concerning measurement protocol for  $R_s$ , SWC, and  $T_s$ , see Wang et al. (2014).

#### 2.3 Measurements of plant-related factors

Plant-related factors considered in this work were root biomass (kgm<sup>-3</sup>), litterfall (kgm<sup>-2</sup>), total soil nitrogen content (gkg<sup>-1</sup>) within the first 25 cm of the soil, leaf area index (LAI, m<sup>3</sup>m<sup>-3</sup>), and plant phenophases. Usually soil samples were collected near
each R<sub>s</sub> chamber on the 15th and 30th day of each month using a soil corer (25 cm in height and 5 cm in diameter). However, collection of soil samples was delayed for several days after rain events to avoid its effects on soil nitrogen content. Three soil replicates were taken near each collar (within 0.5 m), and mixed before they were air dried and sieved through 1, 0.5, 0.25 mm meshes. During sieving, roots and litterfall
were first picked up manually, and scraps of litterfall remaining were separated from the sand by washing the sample with clean water. Dead and alive roots were sorted out by color. The black ones were dead roots and mixed with litterfall as the litterfall sample. The yellow ones were alive roots. Both alive roots and litterfall samples were oven dried at 70 °C to a constant weight. Then the soil samples were sieved through a 0.25 mm mesh for total soil nitrogen measurements. Total nitrogen in the soil was

<sup>25</sup> a 0.25 mm mesh for total soil nitrogen measurements. Iotal nitrogen in the soil was determined with a Kjeldahl Total Nitrogen Apparatus (FOSS 2200, Foss, Denmark).

LAI was measured near each chamber within one hour immediately after sunset using LI-COR 2000 (LI-COR, Nebraska, USA) with a 90° view cap, twice or three times



every month during the measurement period. At each measurement time, LAI was measured close to ground surface in eight directions, which were roughly uniformly distributed within a 0.5 m radius around the chamber, in order to well reflect the effect of vegetation around. Final output of LAI value was mean of the eight records of the eight directions. In addition, the phenological phases of *Artemisia ordosica* were

<sup>5</sup> the eight directions. In addition, the phenological phases of *Artemisia ordosica* were observed and recorded over the growing season at weekly intervals. Three obvious phases identified were leaf expansion (phase I), flowering-bearing (phase II), and leaf coloration-defoliation (phase III; Fig. 2).

## 2.4 Data analysis

- <sup>10</sup> Hourly  $R_s$  data were first screened using limit checking, in which values beyond the range of -1 to 15 µmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup> (mean + 5 standard deviation) were removed from the dataset as spurious values (Wang et al., 2014). Instrument failure and quality control procedures resulted in the removal of about 3% of the data from 4 June to 22 October 2012. The days of rainfall were also excluded during our analyses. Daily mean <sup>15</sup> values (total n = 92) of  $R_s$ ,  $T_s$  and SWC were used in further analyses. We assumed these biweekly measurements of LAI (total n = 8), roots (total n = 7), litterfall (total n = 7) and soil nitrogen (total n = 92) represent or reflect the average variation over the defined time period, because those variables are slow responsive factors to environ-
- mental changes and remain relatively stable over two-week period. The point values in the regression for target variables represent the average characteristics over the defined time period. We calculated the daily coefficient of variation (CV, standard deviation/mean) in  $R_s$  across the 11 positions as a metric of spatial variability. CV for each of phenophases or whole measurement period was calculated as a mean of daily CV over the corresponding period.
- Pearson correlation and regression analyses were used to examine the relationship between  $R_s$  and micro-hydrometeorological and plant-related factors. Class and Regression Tree analysis (CART) were used to explore the main factors controlling spatial variability in  $R_s$  (Breiman et al., 1984). We excluded measurements during rainfall



events in both the correlation and CART analyses, due to the prevalence of CO<sub>2</sub>\_ flux pulses during these events (Fig. 5 in Wang et al., 2014). In CART analyses, we first set the split value to 1, and then used cost-complexity tree pruning based on a 10-fold cross-validation (Venables and Ripley, 2002). All statistical analyses were performed in Matlab (R2010b, Mathworks Inc., Natick MA, USA); the significance level was set at 0.05.

## 3 Results

## 3.1 Topographic heterogeneity in plant-related and micro-hydrometeorological factors, and in soil respiration

- Plant-related factors (i.e., litterfall, root biomass, and soil nitrogen) except LAI, exhibited large spatial heterogeneity (Table 1, Fig. 3). The mean daily CVs of litterfall, root biomass, and soil nitrogen were 82.0, 64.7, and 49.1 %, respectively (Table 1). In contrast, among the four slopes, the micro-hydrometeorological factors exhibited nominal spatial variation (Table 1, Fig. 3). The mean daily CVs of *T*<sub>s</sub> and SWC were only 4.6 and
   22.0 %, respectively (Table 1). Among the four slopes, mean SWC changed very little (0.01–0.03 m<sup>3</sup> m<sup>-3</sup>; Table 2). Soil temperature (*T*<sub>s</sub>) differed among the four slopes, with mean differences ranging from 0.5 to 2.4 °C (Table 2). Among the three height positions (lower, upper, and top positions), there were no consistent pattern in both plant-related and micro-hydrometeorological factors (Fig. 3).
- Soil respiration on the four slopes exhibited obvious differences (Table 2). The lowest averaged  $R_{\rm s}$  on the windward slope was 58 % of the highest  $R_{\rm s}$  on the leeward slope (Table 2). Mean daily CV for  $R_{\rm s}$  across the 11 microsites was 23.5 % over the entire measurement period (Table 1). Among the three height positions for all slopes, no consistent pattern existed in  $R_{\rm s}$  (Fig. 3a).



# 3.2 Relationships between soil respiration, micro-hydrometeorological, and plant-related factors over the measurement period

Topographic heterogeneity in  $R_s$  was correlated positively with the plant-related factors, like root biomass (Table 3, Fig. 4a), litterfall (Table 3, Fig. 4b), and soil nitrogen over the measurement period (Table 3, Fig. 4c), but not with the micro-hydrometeorological factors (i.e.,  $T_s$ ; p = 0.061; Table 3 and SWC; p = 0.852; Table 3). Aside from  $R_s$ , both root biomass and soil nitrogen were correlated positively with litterfall (Fig. 7a and b).

## 3.3 Relationship between micro-hydrometeorological factors, plant-related factors, and soil respiration for different phenophases

- <sup>10</sup> During phase II (Fig. 2), the spatial variation in  $R_s$  correlated with the spatial variation in root biomass, litterfall, and soil nitrogen (Table 3, Figs. 4d–f and 5d, f, g). Variation in root biomass, litterfall, and soil nitrogen explained the majority (> 61 %) of the variation in  $R_s$  among the 11 microsites (Table 3, Fig. 4). Regression slopes between  $R_s$  and root biomass, litterfall, and soil nitrogen were 0.94, 0.51 and 0.77, respectively (Table 3).
- <sup>15</sup> Based on CART analysis, the root biomass was the most significant factor affecting the spatial variation in  $R_s$  during phase II (Fig. 8a).

During phase III (Fig. 2), spatial variation in  $R_s$  was strongly correlated to the spatial variation in root biomass and soil nitrogen content (Table 3, Figs. 4d, f and 6d, g). Variation in soil nitrogen content and root biomass explained 56 and 39% of the variation

<sup>20</sup> in  $R_s$  among the 11 microsites, respectively (Table 3, Fig. 4d and f). Regression slopes between  $R_s$  and root biomass and  $R_s$  and soil nitrogen content were 0.14 and 0.63 (Table 3). Based on CART analysis, soil nitrogen content affected the spatial variation in  $R_s$  the most during phase III (Fig. 8b).



### 4 Discussion

## 4.1 Spatial patterns in soil respiration with respect to plant patchiness

The observed pattern that  $R_s$  on the leeward slope was highest among the four slope orientations, was consistent with the findings of Chai et al. (2012), who found that  $R_s$  on

- <sup>5</sup> leeward slopes was over two times greater than  $R_s$  on windward slopes of a sand dune. The greater  $R_s$  on leeward slopes may be explained by the presence of larger patches of vegetation, most likely supported by the better soil conditions induced by the "fertility island" effect, addressed earlier (Table 2). Associated with the "fertility island" effect is the enrichment of the soil near the plants by the annual casting of photosynthates
- <sup>10</sup> and nitrogen compounds with annual litterfall. In contrast, on the windward slope the level of soil enrichment is limited as a result of smaller vegetation patches, leading to suppression of  $R_s$ . Higher supply of photosynthetic products and soil nitrogen in vicinity to the plants was reported in previous studies by Scott-Denton et al. (2003) and Tang et al. (2005).
- <sup>15</sup> However, we did not find consistent patterns in  $R_s$  between height positions. Thus, we could not draw the conclusion that  $R_s$  at the top position was higher than that at low positions or vice versa. As a comparison, Chai et al., 2012 observed earlierthat  $R_s$  was lower in the lower positions, whereas Xu and Wan (2008) and Liu et al. (2010) showed opposite result. The magnitude of  $R_s$  at a given slope position was consistent <sup>20</sup> with those of plant-related factors (Table 3, Fig. 4),  $R_s$  being highest at positions closest to the plants regardless of height gradients along the slopes.

# 4.2 Plant drivers of the topographic variation in soil respiration over the measurement period

Our present finding that plant-related factors drive the topographic variation in  $R_s$  over a sand dune (Table 3, Fig. 4) was consistent with those of Sponseller and Fisher (2008).

<sup>25</sup> a sand dune (Table 3, Fig. 4) was consistent with those of Sponseller and Fisher (2008). They reported that spatial patterns in  $R_s$  in the Sonoran Desert were related to plant



size and productivity, but unrelated to soil water. Desert ecosystems are generally limited with substrate supply due to their low living biomass and accumulated soil organic matter, but feature large spatial heterogeneity in respiration substrate due to the "fertility-island" effect (Gold et al., 1995; Xie et al., 2001). This can be supported by

- <sup>5</sup> large CVs (49.1, 64.7 and 82.0%) in plant-related factors of our study. In contrast, CVs for the hydrometeorological factors were much smaller (4.6 and 22.0%). Positive relationship between plant-related factors (litterfall, soil nitrogen, and root biomass), together with their positive effects on  $R_s$  (Figs. 4 and 7), emphasized that spatial variation in supply of respiration substrate regulated spatial variation in  $R_s$ . Current pho-
- <sup>10</sup> tosynthate is supplied for root and rhizosphere respiration (Högberg et al., 2001; Tang et al., 2005; Han et al., 2014). Litterfall is supplied for heterotrophic microbe respiration (Reichstein et al., 2002; Zhou et al., 2013). These processes are influenced by soil nitrogen content (Allison et al., 2008; Deng et al., 2010). These findings support our first hypothesis that topographic heterogeneity in  $R_s$  was controlled by the respiration substrate supply provided by plants.

In contrast, many studies in forests and grasslands report that topographic heterogeneity in  $R_s$  can be suitably explained by  $T_s$  and SWC (Kang et al., 2003; Liu et al., 2010). These forests and grasslands are generally not limited by substrate supply, thus more litterfall and heterotrophic respiration. Spatial variation in  $T_s$  and SWC is reported to contribute to the spatial variation in  $R_s$  in forest ecosystems.

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Our finding that plants controlled the topographic variation in  $R_s$  over a fixed sand dune, along with similar findings by other studies (Moyano et al., 2008; Kuzyakov and Gavrichkova, 2010; Han et al., 2014), challenge the legitimacy of empirical models of  $R_s$  based on inputs of  $T_s$  and VWC at the sand-dune scale. As an alternative, we suggest that photosynthesis-related parameters (or variables) ought to be incorporated into these models for an improved characterization of soil–plant relationships in desert ecosystems.



# 4.3 Drivers of topographic heterogeneity in soil respiration as a function of plant phenology

In our study, topographic heterogeneity in R<sub>s</sub> at the flowering and bearing phase (phase II) was significantly related to root biomass, litterfall, and soil nitrogen content. But
 with consideration of interactions between these plant-related factors (Fig. 7), only root biomass was identified as the most important driving factor. In contrast, topographic heterogeneity in R<sub>s</sub> at the leaf coloration and defoliation phase (phase III) was significantly related to root biomass and soil nitrogen content, but only soil nitrogen content was identified as important. Our results support our second hypothesis that drivers of topographic variation in R<sub>s</sub> varies with plant phenophases.

From phase II to III, photosynthetic pigments and leaf photosynthesis declined, and induced large decreases in current photosynthate supply for roots and rhizosphere microorganisms (Hanson et al., 2000; Fu et al., 2002; Sey et al., 2010). Decomposition of litterfall results in a high percentage of recalcitrant substrates during phase II, but

- <sup>15</sup> addition of fresh litterfall during phase III increases the percentage of liable substrates (Lebret et al., 2001; Regina, 2001; Ochieng and Erftemeijer, 2002). Thus, changes in substrate supply between the two phenophases could result in seasonal changes in the relative contribution of autotrophic and heterotrophic respiration to total  $R_s$ . Root respiration contributes largely to total  $R_s$  during the growing season, but heterotrophic <sup>20</sup> respiration contributes largely during the senescence season (Fu et al., 2002; Sey
- et al., 2010; Mauritz and Lipson, 2013). Specifically, in our study, autotrophic respiration may contribute largely to the total  $R_s$  during the phase II, but heterotrophic respiration may contribute largely during the phase III. We could not test this inference directly due to lack of independent measurements of different  $R_s$  components. However, some of our results can indirectly confirm this inference.

In our study,  $R_s$  was sensitive to root biomass during the phase II, but less sensitive during the phase III, indicating a decrease in relative contribution of root respiration to  $R_s$  during the phase III.  $R_s$  was related to litterfall during the phase II, unlike during the



phase III, indicating the switched substrate pool and disproportionate addition of fresh litterfall. However, the controlling factors of autotrophic and heterotrophic respiration can be quite different. Autotrophic respiration is shown in others to studies be largely controlled by root biomass and soil nitrogen (Tu et al., 2013; Zhou et al., 2014), and heterotrophic respiration by the amount of litterfall and soil nitrogen (Yan et al., 2010; Sayer et al., 2011). Thus, the key factor controlling the variation in  $R_s$  can be different for the different plant phenophases.

There has been a clear lack of understanding on the effects of plant phenology on the driving forces of topographic heterogeneity in  $R_s$  (Janssens et al., 2004; Bahn et al.,

<sup>10</sup> 2008; Talmon et al., 2011). In this sense, our work provides new knowledge of driving mechanisms of topographic heterogeneity in  $R_s$  as affected by plant phenology. Based on our work, we suggest that the influences of plant phenology on both current photosynthate and litterfall supply should not be neglected, and partitioning  $R_s$  is necessary for future studies on the spatial heterogeneity in  $R_s$ .

#### 15 5 Conclusions

Soil respiration was the highest on the leeward slope, but the lowest on the windward slope. Topographic heterogeneity in  $R_s$  was related to the variation in supply of downed photosynthetic products, indicating that plants largely control the topographic heterogeneity in  $R_s$  over a sand dune. Due to the "fertility island" effect, spatial patterns of  $R_s$  on different slopes were related to the distance from plant patches, regardless of micro-topographic relief. Drivers of topographic heterogeneity in  $R_s$  varied with plant phenophases. Our findings highlight the importance of plants in controlling the spatial variation in  $R_s$  in desert ecosystems. We suggest that future studies on the spatial variation in  $R_s$  should consider both the spatial and the temporal effects of vegetation, especially in desert environments.

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#### References

- Allison, S. D., Czimczik, C. I., and Treseder, K. K.: Microbial activity and soil respiration under nitrogen addition in Alaskan boreal forest, Glob. Change Biol., 14, 1-13, 2008.
- Asaeda, T. and Rashid, M. H.: Nutrient retention associated with phenological features in Spar-10 ganium erectum stands in a lowland stream. River Res. Appl., 31, 207–215, 2015.
  - Asner, G. P., Archer, S., and Hughes, R. F.: Net changes in regional woody vegetation cover and carbon storage in Texas Dry-lands, 1937–1999, Glob. Change Biol., 9, 316–335, 2003. Austin, A. T.: Has water limited our imagination for aridland biogeochemistry?, Trends Ecol.
- 15
- Evol., 26, 229-235, 2011.
  - Bahn, M., Rodeghiero, M., Anderson-Dunn, M., Dore, S., Gimeno, C., Drösler, M., Williams, M., Ammann, C., Berninger, F., Flechard, C., Jones, S., Balzarolo, M., Kumar, S., Newesely, C., Priwitzer, T., Raschi, A., Siegwolf, R., Susiluoto, S., Tenhunen, J., Wohlfahrt, G., and Gernusca, A.: Soil respiration in European grasslands in relation to climate and assimilate sup-
- ply, Ecosystems, 11, 1353-1367, 2008. 20
- Bond-Lamberty, B. and Thomson, A.: Temperature-associated increases in the global soil respiration record, Nature, 464, 579-582, 2010a.
  - Bond-Lamberty, B. and Thomson, A.: A global database of soil respiration data, Biogeosciences, 7, 1915-1926, doi:10.5194/bg-7-1915-2010, 2010b.
- Breimann, L., Friedman, J., Stone, C. J., and Olshen, R. A.: Classification and regression trees. Chapman & Hall/CRC, Wadsworth, Belmont, California, 1984.
  - Chai, H., Feng, J., and Jing, Y.: Spatial distribution and change trait of soil respiration at the dunes in Hogin Sand Land, Arid Land Geography, 3, 465-472, 2012.
- Davidson, E. A. and Janssens, I. A.: Temperature sensitivity of soil carbon decomposition and feedbacks to climate change, Nature, 440, 165–173, 2006. 30



- 9480
- <sup>30</sup> Hook, P. B., Burke, I. C., and Lauenroth, W. K.: Heterogeneity of soil and plant N and C associated with individual plants and openings in North American short grass steppe, Plant Soil, 138.247-256.1991.
- Han, G., Luo, Y., Li, D., Xia, J., Xing, Q., and Yu, J.: Ecosystem photosynthesis regulates soil respiration on a diurnal scale with a short-term time lag in a coastal wetland, Soil Biol. Biochem., 68, 85-94, 2014. Hanson, P. J., Edwards, N. T., Garten, C. T., and Andrews, J. A.: Separating root and soil micro-

bial contributions to soil respiration: a review of methods and observations, Biogeochemistry,

- doi:10.1371/journal.pone.0034968, 2012. Gold, W. G. and Bliss, L. C.: Water limitations and plant community development in a polar desert, Ecology, 76, 1558-1568, 1995.
- Geng, Y., Wang, Y., Yang, K., Wang, S., Zeng, H., Baumann, F., Kuehn, P., Scholten, T., and He, J.: Soil respiration in Tibetan Alpine Grasslands: belowground biomass and soil moisture, but not soil temperature, best explain the large-scale patterns, PLoS ONE, 7, e34968, 20

nology: a greenhouse pot experiment, Plant Soil, 239, 133-140, 2002.

Feng, W., Zhang, Y., Wu, B., Zha, T., Jia, X., Qin, S., Shao, C., Liu, J., Lai, Z., and Fa, K.: Influence of disturbance on soil respiration in biologically crusted soil during the dry season, The Scientific World J., 2013, 408560, doi:10.1155/2013/408560, 2013. Fu, S., Cheng, W., and Susfalk, R.: Rhizosphere respiration varies with plant species and phe-

15

25

48, 115-146, 2000.

- China, Biogeosciences, 7, 315-328, doi:10.5194/bg-7-315-2010, 2010. Dungan, R. J., Whitehead, D., and Duncan, R. P.: Seasonal and temperature dependence of 10 photosynthesis and respiration for two co-occurring broad-leaved tree species with contrasting leaf phenology, Tree Physiol., 23, 561-568, 2003.
- or confounded factors controlling soil respiration in a temperate mixed hardwood forest, Glob. Change Biol., 4, 217–227, 1998. DeForest, J. L., Noormets, A., McNulty, S. G., Sun, G., Tenney, G., and Chen, J.: Phenophases alter the soil respiration-temperature relationship in an oak-dominated forest, Int. J. Biome-5 teorol., 51, 135-144, 2006.

Deng, Q., Zhou, G., Liu, J., Liu, S., Duan, H., and Zhang, D.: Responses of soil respiration

to elevated carbon dioxide and nitrogen addition in young subtropical forest ecosystems in

Davidson, E. A., Belk, E., and Boone, R. D.: Soil water content and temperature as independent

**BGD** 12, 9465–9495, 2015

Discussion

Paper

Discussion Paper

**Discussion** Paper

**Micro-topographic** variation in soil respiration and its controlling factors

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- Högberg, P., Nordgren, A., Buchmann, N., Taylor, A. F. S., Ekblad, A., Högberg, M. N., Nyberg, G., Ottosson-Löfvenius, M., and Read, D. J.: Large-scale forest girdling shows that current photosynthesis drives soil respiration, Nature, 411, 789–792, 2001.
- Janssens, I. A., Carrara, A., and Ceulemans, R.: Annual *Q*<sub>10</sub> of soil respiration reflects plant phenological patterns as well as temperature sensitivity, Glob. Change Biol., 10, 161–169, 2004.
  - Jia, X., Zhan, T., Wu, B., Zhang, Y., Chen, W., Wang, X., Yu, H., and He, G.: Temperature response of soil respiration in a Chinese pine plantation: hysteresis and seasonal vs. diel  $Q_{10}$ , PLoS ONE, 8, e57858, doi:10.1371/journal.pone.0057858, 2013.
- Kang, S., Doh, S., Lee, D. S., Lee, D., Jin, V. L., and Kimball, J. S.: Topographic and climatic controls on soil respiration in six temperate mixed-hardwood forest slopes, Korea, Glob. Change Biol., 9, 1427–1437, 2003.
  - Kuzyakov, Y. and Gavrichkova, O.: Time lag between photosynthesis and carbon dioxide efflux from soil: a review of mechanisms and controls, Glob. Change Biol., 16, 3386–3406, 2010.
- Lebret, M., Nys, C., and Forgeard, F.: Litter production in a Atlantic beech (*Fagus sylvatica* L.) time sequence, Ann. For. Sci., 58, 755–768, 2001.
  - Lloyd, J. and Taylor, J. A.: On the temperature dependence of soil respiration, Funct. Ecol., 8, 315–323, 1994.
  - Li, J., Zhao, C., Zhu, H., Li, Y., and Wang, F.: Effect of plant species on shrub fertile island at
- an oasis-desert ecotone in the South Junggar Basin, China, J. Arid Environ., 71, 350–361, 2007.
  - Li, P. X., Wang, N., He, W. M., Krüsi, B. O., Gao, S. Q., Zhang, S. M., Yu, F. H., and Dong, M.: Fertile islands under Artemisia ordosica in inland dunes of northern China: effects of habitats and plant developmental stages, J. Arid Environ., 72, 953–963, 2008.
- Liu, W., Xu, W., Hong, J., and Wan, S.: Interannual variability of soil microbial biomass and respiration in responses to topography, annual burning and N addition a semiarid temperature steppe, Geoderma, 158, 259–267, 2010.
  - Luo, J., Chen, Y., Wu, Y., Shi, P., She, J., and Zhou, P.: Temporal-spatial variation and controls of soil respiration in different primary succession stage on glacier forehead in Gongga
- Mountain, China, PLoS ONE, 7, e42354, doi:10.1371/journal.pone.0042354, 2012. Luo, Y. Q.: Terrestrial carbon-cycle feedback to climate warming, Annu. Rev. Ecol. Evol. S., 38, 683–712, 2007.



- Mahall, B. E. and Callaway, R. M.: Root communication mechanisms and intracommunity distributions of two Mojave Desert shrubs, Ecology, 73, 2145–2151, 1992.
- Martin, J. G. and Bolstad, P. V.: Variation of soil respiration at three spatial scales: components within measurements, intra-site variation and patterns on the landscape, Soil Biol. Biochem.,
- <sup>5</sup> 41, 530–543, 2009.

20

30

- Mauritz, M. and Lipson, D. L.: Altered phenology and temperature sensitivity of invasive annual grasses and forbs changes autotrophic and heterotrophic respiration rates in a semiarid shrub community, Biogeosciences Discuss., 10, 6335–6375, doi:10.5194/bgd-10-6335-2013, 2013.
- Moyano, F. E., Kutsch, W. L., and Rebmann, C.: Soil respiration fluxes in relation to photosynthetic activity in broad-leaf and needle-leaf forest stands, Agr. Forest Meteorol., 148, 135– 143, 2008.
  - Ngao, J., Epron, D., Delpierre, N., Bréda, N., Granier, A., and Longdoz, B.: Spatial variability of soil CO<sub>2</sub> efflux linked to soil parameters and ecosystem characteristics in a temperate beech forest. Agr. Ecrost Motoorol. 154, 136–146, 2012
- <sup>15</sup> forest, Agr. Forest Meteorol., 154, 136–146, 2012.
  - Ochieng, C. A. and Erftemeijer, P. L.: Phenology, litterfall and nutrient resorption in *Avicennia marina* (Forssk.) Vierh in Gazi Bay, Kenya, Trees, 16, 167–171, 2002.
    - Osono, T.: Diversity, resource utilization, and phenology of fruiting bodies of litter-decomposing macrofungi in subtropical, temperate, and subalpine forest, J. Forest Res.-Jpn., 20, 60–68, 2014.
  - Parker, K. C.: Topography, substrate, and vegetation patterns in the northern Sonoran Desert, J. Biogeogr., 18, 151–163, 1991.
  - Poulter, B., Frank, D., Ciais, P., Myneni, R. B., Andela, N., Bi, J., Broquet, G., Canadell, J. G., Chevallier, F., Liu, Y. Y., Running, S. W., Sitch, S., and Van der Werf, G. R.: Contribution
- of semi-arid ecosystems to interannual variability of the global carbon cycle, Nature, 509, 600–603, 2014.
  - Raich, J. W. and Schlesinger, W. H.: The global carbon dioxide flux in soil respiration and its relationship to vegetation and climate, Tellus B, 44, 81–99, 1992.
  - Raich, J. W., Potter, C. S., and Bhagawati, D.: Interannual variability in global soil respiration, 1980–94, Glob. Change Biol., 8, 800–812, 2002.
  - Regina, I. S.: Litter fall, decomposition and nutrient release in three semi-arid forests of Duero basin, Spain, Foresty, 74, 347–358, 2001.



- Reichstein, M., Tenhunen, J. D., Roupsard, O., Ourcival, J.-M., Rambal, S., Miglietta, F., Peressotti, A., Pecchiari, M., Tirone, G., and Valentini, R.: Severe drought effects on ecosystem CO<sub>2</sub> and H<sub>2</sub>O fluxes at three Mediterranean evergreen sites: revision of current hypotheses?, Glob. Change Biol., 8, 999–1017, 2002.
- <sup>5</sup> Richerson, P. J. and Lum, K.: Patterns of plant species diversity in California: relation to weather and topography, Am. Nat., 116, 504–536, 1980.

Roles, B., Donders, S., Werger, M. J., and Dong, M.: Relation of wind-induced sand displacement to plant biomass and plant sand-binding capacity, Acta Bot. Sin., 43, 979–982, 2001.

- Sayer, E. J., Heard, M. S., Grant, H. K., Marthews, T. R., and Tanner, E. V. J.: Soil carbon
- release enhanced by increased tropical forest litterfall, Nature Climate Change, 1, 304–307, 2011.

Schlesinger, W. H. and Andrews, J. A.: Soil respiration and the global carbon cycle, Biogeochemistry, 48, 7–20, 2000.

#### Schlesinger, W. H. and Raikes, J. A.: On the spatial pattern of soil nutrients in desert ecosys-

tems, Ecology, 77, 364–374, 1996.

15

20

25

Scott-Denton, L. E., Sparks, K. L., and Monson, R. K.: Spatial and temporal controls of soil respiration rate in a high-elevation, subalpine forest, Soil Biol. Biochem., 35, 525–534, 2003.

Sey, B. K., Manceur, A. M., Whalen, J. K., Gregorich, E. G., and Rochette, P.: Root-derived respiration and nitrous oxide production as affected by crop phenology and nitrogen fertilization, Plant Soil, 326, 369–379, 2010.

- Sponseller, R. A. and Fisher, S. G.: The influence of drainage networks on patterns of soil respiration in a desert catchment, Ecology, 89, 1089–1100, 2008.
- Stock, W. D., Dlamini, T. S., and Cowling, R. M.: Plant induced fertile islands as possible indicators of desertification in a succulent desert ecosystem in northern Namaqualand, South Africa, Plant Ecol., 142, 161–176, 1999.
- Tang, J., Baldocchi, D. D., and Xu, L.: Tree photosynthesis modulates soil respiration on a diurnal time scale, Glob. Change Biol., 11, 1298–1304, 2005.

Thompson, D. B., Walker, L. R., Landau, F. H., and Stark, L. R.: The influence of elevation, shrub species, and biological soil crust on fertile islands in the Mojave Desert, USA, J. Arid

- <sup>30</sup> Environ., 61, 609–629, 2005.
  - Tu, L., Hu, X., Zhang, J., Li, X., Hu, H., Liu, L., and Xiao, Y.: Nitrogen addition simulates different components of soil respiration in a subtropical bamboo ecosystem, Soil Biol. Biochem., 58, 255–264, 2013.



- Venables, W. N. and Ripley, B. D.: Modern Applied Statistics with S-Plus, Springer Press, New York, 2002.
- Wang, B., Zha, T. S., Jia, X., Wu, B., Zhang, Y. Q., and Qin, S. G.: Soil moisture modifies the response of soil respiration to temperature in a desert shrub ecosystem, Biogeosciences,
- 11, 259–268, doi:10.5194/bg-11-259-2014, 2014. 5
  - Xie, G. and Steinberger, Y.: Temporal patterns of C and N under shrub canopy in a loessial soil desert ecosystem, Soil Biol. Biochem., 33, 1371-1379, 2001.
  - Xu, W. and Wan, S.: Water- and plant-mediated responses of soil respiration to topography, fire, and nitrogen fertilization in a semiarid grassland in north China, Soil Biol. Biochem., 40, 679-687, 2008.
- 10
  - Yan, L., Chen, S., Huang, J., and Lin, G.: Different responses of auto- and heterotrophic soil respiration to water and nitrogen addition in a semiarid temperate steppe. Glob. Change Biol., 16, 2345-2357, 2010.

Yuste, J. C., Janssens, I. A., Carrara, A., and Ceulemans, R.: Annual Q<sub>10</sub> of soil respiration

- reflects plant phenological patterns as well as temperature sensitivity, Glob. Change Biol., 15 10, 161–169, 2004.
  - Zhou, L., Zhou, X., Zhang, B., Lu, M., Luo, Y., Liu, L., and Li, B.: Different responses of soil respiration and its components to nitrogen addition among biomes: a meta-analysis, Glob. Change Biol., 7, 2332–2343, 2014.
- Zhou, Z., Zhang, Z., Zha, T., Luo, Z., Zheng, J., and Sun, J.: Predicting soil respiration using 20 carbon stock in roots, litter and soil organic matter in forest of Loss Plateau in China, Soil Biol. Biochem., 57, 135–143, 2013.



**Table 1.** Average values (standard error in brackets) and CV (%) for micro-hydrometeorological factors and plant-related factors over the three plant phenophases and measurement period,  $R_s$ : soil respiration (µmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>);  $T_s$ : soil temperature at 10 cm depth (°C); SWC: volumetric soil water content (m<sup>3</sup> m<sup>-3</sup>); Litter: litterfall (kgm<sup>-2</sup>); Root: root biomass (kgm<sup>-3</sup>); LAI: leaf area index (m<sup>2</sup> m<sup>-2</sup>); N: soil nitrogen content at 0–25 cm depth (gkg<sup>-1</sup>).

Variale	Phase I		Phase II		Phase III		Measurement period	
	<del>mean (SE)</del>	<mark>€V%</mark>	mean (SE)	CV%	mean (SE)	CV%	mean (SE)	CV%
R <sub>s</sub>	<del>0.94 (0.22)</del>	<del>23.6</del>	1.49 (0.11)	25.2	0.90 (0.06)	21.3	1.21 (0.09)	23.5
Ts	<del>21.73 (0.92)</del>	<del>4.2</del>	23.82 (0.24)	3.4	14.90 (0.33)	7.4	19.74 (0.27)	4.6
SWC	<del>0.07 (0.01)</del>	<del>14.3</del>	0.09 (0.01)	20.8	0.10 (0.01)	24.0	0.09 (0.01)	22.0
Litter			0.74 (0.17)	78.5	0.26 (0.05)	67.7	0.63 (0.16)	82.0
Root			0.59 (0.15)	84.2	0.90 (0.25)	92.2	0.75 (0.14)	64.7
LAI			1.76 (0.12)	22.7	1.72 (0.12)	23.9	1.76 (0.12)	21.8
Ν			0.18 (0.02)	43.5	0.13 (0.02)	53.3	0.14 (0.02)	49.1



**Table 2.** Mean values (standard error in brackets) of micro-hydrometeorological and plantrelated factors on different slope orientations.  $R_s$ : soil respiration (µmol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>);  $T_s$ : soil temperature (°C); SWC: volumetric soil water content (m<sup>3</sup> m<sup>-3</sup>); Litter: litterfall (kg m<sup>-2</sup>); Root: root biomass (kg m<sup>-3</sup>); LAI: leaf area index (m<sup>2</sup> m<sup>-2</sup>); N: soil nitrogen content at 0–25 cm soil depths (g kg<sup>-1</sup>).

	Windward	Leeward	North-facing	South-facing
R <sub>s</sub>	0.96 (0.05)	1.65 (0.03)	1.13 (0.27)	1.27 (0.19)
$T_{\rm s}$	19.94 (0.60)	18.36 (0.08)	20.77 (0.23)	19.44 (0.09)
SWC	0.10 (0.02)	0.11 (0.04)	0.08 (0.02)	0.09 (0.02)
Litter	0.49 (0.29)	1.53 (0.37)	0.23 (0.15)	0.67 (0.28)
Root	0.31 (0.13)	1.41 (0.02)	0.46 (0.03)	0.32 (0.17)
LAI	1.80 (0.48)	2.22 (0.43)	1.70 (0.05)	1.47 (0.25)
Ν	0.11 (0.03)	0.25 (0.04)	0.09 0.02	0.14 (0.03)



**Table 3.** Regression equations between  $R_s$  (*y*) and different micro-hydrometeorological and plant-related factors (*x*) for the measurement period and flowering-bearing (phase II) and leaf coloration-defoliation phase (phase III).  $T_s$ : soil temperature (°C); SWC: volumetric soil water content (m<sup>3</sup> m<sup>-3</sup>); Litter: litterfall (kgm<sup>-2</sup>); Root: root biomass (kgm<sup>-3</sup>); LAI: leaf area index (m<sup>2</sup> m<sup>-2</sup>); N: soil nitrogen content at 0–25 cm soil depths (gkg<sup>-1</sup>).

Phase	Factors	Equation	$R^2$	p
Whole period	Root biomass Litterfall	y = 0.51 + x + 0.93 y = 0.43 + x + 0.95	0.61 0.59	0.004 0.005
	Ts	y = -0.18 + x + 4.84	0.34	0.061
	SWC	$y = 0.89 \times x + 1.13$	< 0.01	0.852
	LAI	$y = 0.28 \times x + 0.73$	0.14	0.256
	Ν	$y = 3.48 \times x + 0.74$	0.68	0.002
Phase II	Root biomass	y = 0.64 + x + 1.11	0.72	0.001
	Litterfall	$y = 0.51 \times x + 1.11$	0.61	0.004
	T <sub>s</sub>	y = -0.23 + x + 7.00	0.25	0.118
	SWC	y = -0.27 + x + 1.51	< 0.01	0.969
	LAI	$y = 0.48 \times x + 0.64$	0.26	0.108
	Ν	$y = 5.04 \times x + 0.77$	0.68	0.001
Phase III	Root biomass	y = 0.14 + x + 0.77	0.37	0.048
	Litterfall	$y = 0.55 \times x + 0.76$	0.26	0.112
	T <sub>s</sub>	y = -0.09 + x + 2.24	0.26	0.105
	SWC	$y = 0.39 \times x + 0.86$	< 0.01	0.887
	LAI	y = 0.11 + x + 0.71	0.05	0.492
	Ν	$y = 1.99 \times x + 0.63$	0.56	0.008

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**Figure 2.** Phenophases of *Artemisia ordosica* over the measurement period. Three phases considered, included leaf expansion (phase I), flowering-bearing (phase II), and leaf coloration-defoliation (phase III) phases.





**Figure 3.** Mean values of soil respiration ( $R_s$ ), micro-hydrometeorological ( $T_s$  and SWC) and plant-related factors (litter, root, soil N and LAI) at 11 positions over the measurement period (whole period), flowering-bearing phase (phase II), and leaf coloration-defoliation phase (phase III) from June to October 2012.  $R_s$ : soil respiration ( $\mu$ mol CO<sub>2</sub> m<sup>-2</sup> s<sup>-1</sup>);  $T_s$ : soil temperature (°C); SWC: volumetric soil water content (m<sup>3</sup> m<sup>-3</sup>); Litter: litterfall (kgm<sup>-2</sup>); Root: root biomass (kgm<sup>-3</sup>); LAI: leaf area index (m<sup>2</sup> m<sup>-2</sup>); Soil N: soil nitrogen content at 0–25 cm soil depths (gkg<sup>-1</sup>). Error bar represent standard error. Black color represents for whole period. Red color represents for flowering-bearing phase. Bule color represents for leaf coloration-defoliation defoliation phase.





**Figure 4.** Relationships between soil respiration ( $R_s$ ) and root biomass, and litterfall and soil nitrogen over the measurement period (open circles) and during flowering-bearing (phase II; gray circles) and leaf coloration-defoliation (phase III; black circles) phase. Equations,  $R^2$ , and p values are given in Table 3.











**Figure 6.** Spatial patterns of soil respiration ( $R_s$ ), micro-hydro meteorological ( $T_s$  and SWC) and plant-related factors (Litter, Root, Soil N and LAI) over the sand dune in coloration-defoliation phase. Data values for all variables were normalized into the range of 0–1 using feature scaling method ( $X_{normalized} = (X - X_{min})/(X_{max} - X_{min})$ ). Black dots represent the measurement positions as showed in Fig. 1.











**Figure 8.** Results of CART analysis with key factors explaining spatial variability in soil respiration at the flowering-bearing (a) and coloration-defoliation phases (b). Predictor variables are depicted at the top of each branch. Their thresholds are shown at the side of each branch and the mean soil  $CO_2$  efflux (including number of observations in parentheses) is reported below the terminal nodes.

