Dynamics of N₂ fixation and fate of diazotroph-derived
 nitrogen in a Low Nutrient Low Chlorophyll ecosystem:
 results from the VAHINE mesocosm experiment (New
 Caledonia)

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    S. Bonnet<sup>1,2</sup>, H. Berthelot<sup>1</sup>, K. Turk-Kubo<sup>3</sup>, S. Fawcett<sup>4,5</sup>, E. Rahav<sup>6,7</sup>, S. L'Helguen<sup>8</sup>,
    I. Berman-Frank<sup>6</sup>
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9 [1] {IRD, Aix Marseille Université, CNRS/INSU, Université de Toulon, Mediterranean Institute
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10 of Oceanography (MIO) UM 110, 13288, Marseille-Noumea, France-New Caledonia}
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- 11 [2] {Mediterranean Institute of Oceanography (MIO) IRD/CNRS/Aix-Marseille University
- 12 IRD Noumea, 101 Promenade R. Laroque, BPA5, 98848, Noumea cedex, New Caledonia}
- 13 [3] {Ocean Sciences Department, University of California, Santa Cruz, USA}
- [4] {Department of Geosciences, M45 Guyot Hall, Princeton University, Princeton, New Jersey
 08544, USA}
- 16 [5] {Department of Oceanography, University of Cape Town, Rondebosch, 7701, South Africa}
- 17 [6] {Mina and Everard Goodman Faculty of Life Sciences, Bar-Ilan University, Ramat Gan,18 Israel}
- 19 [7] {National Institute of Oceanography, Israel Oceanographic and Limnological Research, Haifa,
- 20 Israel}
- 21 [8] {Université de Brest, CNRS/IRD, UMR6539, Laboratoire des Sciences de l'Environnement
- 22 Marin, OSU-IUEM, 29280 Plouzané, France}
- 23
- 24 Correspondence to: S. Bonnet (<u>sophie.bonnet@ird.fr</u>)

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1 Abstract

 N_2 fixation rates were measured daily in large (~50 m³) mesocosms deployed in the tropical 2 South West Pacific coastal ocean (New Caledonia) to investigate the temporal variability in N₂ 3 fixation rates in relation with environmental parameters and study the fate of diazotroph-derived 4 nitrogen (DDN) in a low nutrient, low chlorophyll ecosystem. The mesocosms were fertilized 5 with ~0.8 µM dissolved inorganic phosphorus (DIP) to stimulate diazotrophy. Bulk N₂ fixation 6 rates were replicable between the three mesocosms, averaged 18.5 ± 1.1 nmol N L⁻¹ d⁻¹ over the 23 7 days, and increased by a factor of two during the second half of the experiment (days 15 to 23) to 8 reach 27.3±1.0 nmol N L⁻¹ d⁻¹. These later rates measured after the DIP fertilization are higher 9 than the upper range reported for the global ocean. During the 23-days of the experiment, N_2 10 11 fixation rates were positively correlated with seawater temperature, primary production, bacterial 12 production, standing stocks of particulate organic carbon (POC), nitrogen (PON) and phosphorus 13 (POP), and alkaline phosphatase activity, and negatively correlated with DIP concentrations, DIP turnover time, nitrate, and dissolved organic nitrogen and phosphorus concentrations. The fate of 14 15 DDN was investigated during a bloom of the unicellular diazotroph UCYN-C that occurred during the second half of the experiment. Quantification of diazotrophs in the sediment traps 16 indicates that ~10 % of UCYN-C from the water column was exported daily to the traps, 17 representing as much as 22.4 ± 5.5 % of the total POC exported at the height of the UCYN-C 18 19 bloom. This export was mainly due to the aggregation of small (5.7±0.8 µm) UCYN-C cells into large (100-500 µm) aggregates. During the same time period, a DDN transfer experiment based 20 21 on high-resolution nanometer scale secondary ion mass spectrometry (nanoSIMS) coupled with $^{15}N_2$ isotopic labelling revealed that 16±6 % of the DDN was released to the dissolved pool and 22 21±4 % was transferred to non-diazotrophic plankton, mainly picoplankton (18±4 %) followed 23 by diatoms $(3\pm 2\%)$. This is consistent with the observed dramatic increase in picoplankton and 24 diatom abundances, primary production, bacterial production, and standing stocks of POC, PON, 25 and POP during the second half of the experiment. These results offer insights into the fate of 26 DDN during a bloom of UCYN-C in low nutrient, low chlorophyll ecosystems. 27

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1 **1 Introduction**

Next to light, nitrogen (N) is the major limiting factor for primary productivity in much of the low-latitude surface ocean (Falkowski, 1997; Moore et al., 2013). Nitrate (NO_3^{-}) is the dominant form of fixed nitrogen (N) in seawater and derives from the remineralization of sinking organic N in the dark ocean. NO_3^{-} is supplied to photic waters by upward mixing and transport, and constitutes the main source of fixed N for photosynthetic organisms in the temperate and high latitude ocean. In the oligotrophic tropical and subtropical oceans, vertical mixing and transport of NO_3^{-} is generally low and surface waters are often depleted in NO_3^{-} .

In these ocean deserts, specialized organisms termed N₂-fixers (or diazotrophs) are able to use N in its simplest and most abundant form on Earth and in seawater, namely dinitrogen (N₂). Diazotrophs possess the nitrogenase enzyme, which cleaves the strong triple bond of the N₂ molecule to form bioavailable ammonium (NH₄⁺) which is assimilated as amino acids enabling biomass growth and division. N₂ fixation thus introduces a source of new bioavailable N to surface waters, and is considered to be the most important external source of N to the ocean, more significant than atmospheric and riverine inputs (Gruber, 2004).

The dynamics of microbial communities such as diazotrophs can change abruptly in the ocean in 16 response to small perturbations or environmental stressors. In particular, N2 fixation has been 17 described as a very 'patchy' process in the ocean (Bombar et al., 2015). Many factors control the 18 19 distribution and activity of diazotrophs such as temperature (Bonnet et al., 2015; Moisander et al., 2010; Raveh et al., 2015; Staal et al., 2003), nutrient availability (mainly phosphate and iron) 20 (e.g., (Mills et al., 2004)), pCO₂ ((e.g. (Levitan et al., 2007)), ambient concentrations of fixed N 21 $(NO_3^- and NH_4^+)$ (e.g., (Knapp et al., 2012), as well as physical forcing (e.g., (Fong et al., 2008)). 22 Most studies dedicated to understanding the controls on marine N₂ fixation have been undertaken 23 24 along large oceanic transects; these are particularly valuable and have recently led to the compilation of a global ocean database of diazotrophy (Luo et al., 2012). Spatial variability in N_2 25 fixation is thus far better documented and understood than temporal variability, despite the 26 intimate connections between time and space scales in the ocean. Time-series stations with near-27 28 monthly observations set up in the late 1980's under the international JGOFS program in the subtropical North Atlantic, Pacific, and Mediterranean Sea have provided valuable data regarding 29 the controls on N_2 fixation and its role in biogeochemical cycles on seasonal and inter-annual 30 timescales (Dore et al., 2008; Garcia et al., 2006; Grabowski et al., 2008; Karl et al., 2012; Knapp 31

et al., 2005; Orcutt et al., 2001), and have also revealed novel diazotrophic microorganisms (Zehr
et al., 2008) with unexpected metabolic strategies such as UCYN-A cyanobacteria that lack the
oxygen-producing photosystem II complex (Tripp et al., 2010). However, fairly little attention
has been paid to sub-seasonal variability in N₂ fixation and its biogeochemical drivers and
consequences.

In the framework of the VAHINE (VAriability of vertical and tropHIc transfer of diazotroph 6 7 derived N in the south wEst Pacific) project, we deployed three large volume mesocosms (~50 m³, Fig. 1) in the tropical South West Pacific coastal ocean, a region known to support 8 9 diazotrophy during the austral summer (Dupouy et al., 2000; Rodier and Le Borgne, 2010, 2008). Our goal was to study the high frequency temporal dynamics of N₂ fixation over short time scales 10 (sampling every day for 23 days), in relation to hydrological parameters, biogeochemical stocks 11 12 and fluxes, and the dynamics of phytoplanktonic and bacterial communities in the same water 13 mass.

The mesocosm approach allowed us to investigate the fate of the recently fixed N₂ and its transfer 14 15 from diazotrophs to non-diazotrophic organisms in this oligotrophic marine ecosystem. Diazotrophs can typically release from 10 to 50 % of their recently fixed N₂ (or diazotroph 16 derived N, hereafter called DDN) as dissolved organic N (DON) and NH₄⁺ (Glibert and Bronk, 17 1994; Meador et al., 2007; Mulholland et al., 2006). This exudate is potentially available for 18 19 assimilation by the surrounding planktonic communities. However, such transfer of DDN to the 20 surrounding planktonic community and its potential impact on export production is poorly understood and rarely quantified. 21

22 Over the course of this 23-day mesocosm experiment, diatom-diazotroph associations (DDAs) were the most abundant N₂ fixers during the first half of the experiment (days 2 to 14), while a 23 bloom of the unicellular N₂-fixing cyanobacteria from Group C (UCYN-C) occurred during the 24 25 second half of the experiment (days 15 to 23) (Turk-Kubo et al., 2015). Berthelot et al. (2015b) described the evolution of the C, N, and P pools and fluxes during the experiment and 26 investigated the contribution of N2 fixation and DON uptake to primary production and particle 27 28 export. They also explored the fate of the freshly produced particulate organic N (PON), i.e., whether it was preferentially accumulated and recycled in the water column or exported out of 29 the system. Complementary to this approach, Knapp et al. (2015) report the results of a $\delta^{15}N$ 30 budget performed in the mesocosms to assess the dominant source of N (i.e., NO_3^- versus N_2 31

fixation) fueling export production during the 23-day experiment. In the present study, we focus specifically on the fate of DDN in the ecosystem during the UCYN-C bloom by studying i) the direct export of diazotrophs into the sediment traps, and ii) the transfer of DDN to nondiazotrophic plankton using high-resolution nanometer scale secondary ion mass spectrometry (nanoSIMS) coupled with $^{15}N_2$ isotopic labelling during a 72 h-process experiment.

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7 2 Methods

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9 2.1 Mesocosm description and sampling strategy

Three replicate large-volume mesocosms (surface 4.15 m^2 , volume ~50 m^3 , Fig. 1) were 10 deployed in the oligotrophic New Caledonian lagoon, 28 km off the coast of Noumea (latitude: 11 22°28,855' S; longitude: 166°26,724' E) from January 13th to February 6th 2013. They consisted 12 13 of large enclosures open to the air made of two 500 µm-thick films of polyethylene (PE) and vinyl acetate (EVA, 19%), with nylon meshing in between to allow for maximum resistance and 14 15 light penetration (produced by HAIKONENE KY, Finland). The mesocosm bags were 2.3 m in diameter and 15 m in height, and were equipped with removable sediment traps that enabled the 16 collection of sinking material once a day (Fig. 1b). To alleviate any potential phosphorus 17 limitation of diazotrophy in the mesocosms, the bags were intentionally fertilized with ~0.8 µmol 18 L^{-1} of dissolved inorganic phosphorus (DIP) four days after the start of the experiment. A more 19 detailed description of the mesocosms setup, the selection of the study site, and the deployment 20 strategy can be found in Bonnet et al. (2016). 21

Vertical CTD profiles were performed every morning in each of the three mesocosms (hereafter referred to as M1, M2, and M3) and in the surrounding waters (hereafter referred to as lagoon waters) using a SBE Seabird CTD. All discrete samples for the parameters described below were collected daily at 7 am at three depths (1, 6, and 12 m) in each mesocosm and in the lagoon waters using braided PVC tubing (Holzelock-Tricoflex, inner diameter = 9.5 mm) connected to a Teflon PFA pump (St-Gobain Performance Plastics) activated by pressurized air. Finally, sediment trap samples were collected daily from each mesocosm by SCUBA divers.

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1 2.2 Experimental procedures

2 2.2.1 N₂ fixation measurements within the mesocosms and methods 3 intercomparison

Seawater samples for N_2 fixation rate measurements were dispensed into HCl-washed 4.5 L 4 polycarbonate bottles that were sealed with septa and amended with ¹⁵N₂-enriched seawater 5 (Mohr et al., 2010; Wilson et al., 2012), hereafter called the ¹⁵N₂ dissolution method. Briefly, the 6 $^{15}N_2$ -enriched seawater was prepared from 0.2 μ m-filtered seawater (Sartobrand (Sartorius) 7 cartridges) collected from the same site in a 4.5 L polycarbonate HCl-washed bottle. Seawater 8 was first degassed through a degassing membrane (Membrana, Minimodule®, flow rate fixed at 9 450 mL min⁻¹) connected to a vacuum pump (<200 mbar) for at least 1 h. The bottle was then 10 closed with a septum cap and amended with 1 mL of ${}^{15}N_2$ (98.9 atom% ${}^{15}N$, Cambridge Isotopes 11 Laboratories, Inc) per 100 mL of seawater. The bottle was shaken vigorously to fragment the ${}^{15}N_2$ 12 bubble, and incubated overnight at 20 m depth at the study site (3 bars) to promote ${}^{15}N_2$ 13 dissolution. The experimental bottles were amended with 5 % vol:vol ¹⁵N₂ enriched seawater 14 15 (i.e., 225 mL), sealed without headspace with silicon septum caps, and incubated for 24 h on an *in situ* mooring line located close to the mesocosms at the appropriate sampling depth. After 24 h, 16 12 mL of the incubated seawater were subsampled into Exetainers®. These were preserved 17 upside down in the dark at 4 °C and analyzed less than 6 months after the experiment using a 18 Membrane Inlet Mass Spectrometer (MIMS) (Kana et al., 1994) to quantify the ¹⁵N enrichment of 19 the N_2 pool in the incubation bottles. The MIMS analyses yielded an average atom% ¹⁵N for the 20 N₂ pool of 2.4±0.2 (n=10). After collection of the Exetainer® subsamples, 2.2 L from each 21 experiment bottle were filtered under low vacuum pressure (<100 mm Hg) onto a pre-combusted 22 (4 h at 450 °C) GF/F filter (25 mm diameter, 0.7 µm nominal porosity) for 'bulk' N₂ fixation rate 23 determination. The remaining volume (2.2 L) was pre-filtered through a 10 µm pore-size 24 polycarbonate filter, and collected on a pre-combusted GF/F filter for analysis of the pico- and 25 nanoplanktonic (<10 µm) N₂ fixation rates. Filters were stored at -20 °C until the end of the 26 VAHINE experiment, then dried for 24 h at 60 °C before mass spectrometric analysis (see 27 section 'Mass spectrometry analyses' below). Every day, an extra 2.2 L bottle was filled with 28 mesocosm surface water (from ~1 m), spiked with ${}^{15}N_2$, and immediately filtered to determine 29 the natural ¹⁵N enrichment of the PON, which is required for calculations of N₂ fixation rates. 30

In the present study, we decided to use the ${}^{15}N_2$ dissolution method to measure N₂ fixation rates 1 as several authors (Großkopf et al., 2012; Mohr et al., 2010; Rahav et al., 2013; Wilson et al., 2 2012) have reported an underestimation of rates when using the bubble method (i.e., when the 3 $^{15}N_2$ gas is injected directly in the incubation bottle using a syringe, see below) due to incomplete 4 equilibration of the ${}^{15}N_2$ gas between the headspace and the seawater in the incubation bottles 5 compared to theoretical calculations. However, the differences observed between the two 6 7 methods appear to depend on the environmental conditions (Shiozaki et al., 2015). Here, we performed an inter-comparison of both methods on day 11 in surface waters (from ~1 m) 8 collected from M1. Briefly, seawater samples from M1 were dispensed into twelve HCl-washed 9 4.5 L polycarbonate bottles as described above and closed with septum caps. Six bottles were 10 spiked with 4 mL ¹⁵N₂ (98.9 atom% ¹⁵N, Cambridge isotopes Laboratories, Inc) via a gas-tight 11 syringe, hereafter called the bubble method. Each bottle was shaken 20 times to fragment the ${}^{15}N_2$ 12 13 bubble and facilitate its dissolution. The six remaining bottles were treated as described above for the dissolution method. All twelve bottles were then incubated for 24 h in an on-deck incubator at 14 15 irradiances corresponding to the sampling depth using screening, and cooled with circulating surface seawater. 16

A recent study (Dabundo et al., 2014) reports potential contamination of some commercial ${}^{15}N_2$ 17 gas stocks with ¹⁵N-enriched NH_4^+ , NO_3^- and/or nitrite (NO_2^-), and nitrous oxide (N_2O). Dabundo 18 et al. (2014) analysed various brands of ¹⁵N₂ gas and found that the Cambridge Isotopes stock 19 (i.e., the one used in this study) contained low concentrations of ¹⁵N contaminants, and the 20 potential overestimation of N₂ fixation rates modeled using this contamination level would range 21 from undetectable to 0.02 nmol N L^{-1} d⁻¹. The rates measured in this study ranged from 0.5 to 22 69.6 nmol N $L^{-1} d^{-1}$ suggesting that, if present, stock contamination of the magnitude reported by 23 (Dabundo et al., 2014) would be too low to affect the results described here. To verify this, one of 24 our ¹⁵N₂ Cambridge Isotopes batches (18/061501) was checked for contamination following the 25 method described in Dabundo et al. (2014); it was 1.4×10^{-8} mol of ${}^{15}NO_3^{-1}$ per mol of ${}^{15}N_2$ and 26 1.1 x 10^{-8} mol NH₄⁺ per mol of ${}^{15}N_2$. The application of this contamination level to our samples 27 using the model provided by Dabundo et al. (2014) indicates that our rates may only be 28 overestimated by ~0.05 %, confirming that our present results were unaffected by possible ${}^{15}N_2$ 29 stock contamination. 30

2.2.2 Phenotypic characterization of UCYN in the water column and the sediment traps

To investigate the direct export of UCYN-C cells during the bloom of UCYN-C that occurred in the second half of the experiment, a detailed phenotypic characterization of UCYN-C was performed at the height of the bloom (days 17 and 19), both in the water column and in the sediment traps. In parallel, UCYN-C and other diazotroph phylotypes were quantified in the sediment traps on days 17 and 19 (analytical protocols are detailed below in section 2.3).

Seawater samples for microscopic analyses were collected every day from 1, 6, and 12 m in each 8 9 mesocosm in 4.5 L polycarbonate bottles as described above. Samples were immediately filtered onto 2 µm 47 mm polycarbonate filters that were fixed with paraformaldehyde (4 % final 10 11 concentration) and incubated for 15 minutes at room temperature, then stored at -80 °C until microscopic analysis. Formalin-fixed sediment trap samples were homogenized and 2 ml were 12 13 filtered onto 2 µm polycarbonate filters for further microscopic analyses. To characterize the phenotype of UCYN (free living cells versus colonies) in the mesocosms as a function of depth, 14 15 we performed a detailed microscopic analysis on days 17 and 19 in M2. Note that UCYN-A cannot be observed by standard epifluorescent microscopy. Filtered samples from each depth (1, 16 6, and 12 m) and from the sediment traps (~15 m) were visualized using a Zeiss Axioplan (Zeiss, 17 Jena, Germany) epifluorescence microscope fitted with a green (510-560 nm) excitation filter, 18 19 which targeted the UCYN phycoerythrin-rich cells. For each filter, 47 photographs of various sections of the filter were taken at random. Each fluorescent particle was automatically delimited 20 as a region of interest (ROI) using an in-house imageJ script. The photographs were scanned 21 22 visually to remove ROIs that did not correspond to UCYN cells or UCYN aggregated cells. The area of each ROI was converted to equivalent volume assuming a spherical shape for all the 23 aggregates. The volume of individual cells was determined from the average volume of the ROI 24 represented by only one cell. The resultant cell volume was then used to compute the number of 25 cells in each aggregate. 26

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28 2.2.3 DDN transfer experiment

The fate of the fixed N_2 during the UCYN-C bloom (that occurred from days 15 to 23) was investigated on days 17 to 20 in M2 at 6 m. In addition to N_2 fixation measurements, seawater was sampled as described above into twelve additional 4.5 L HCl-washed polycarbonate bottles

equipped with septum caps. Full bottles were immediately amended with the dissolved ${}^{15}N_2$ gas 1 2 (98.9 atom% ¹⁵N, Cambridge Isotopes Laboratories, Inc) as described above (dissolution method), and with 1 mL of 80 g L⁻¹ NaH¹³CO₃ solution (99 atom% ¹³C, Cambridge Isotopes 3 Laboratories, Inc) and incubated *in situ* on the mooring line at 6 m-depth close to the mesocosms. 4 After 24 h, 36 h, and 72 h of incubation (hereafter referred to as T24 h, T36 h, and T72 h), three 5 replicate ¹⁵N₂ labelled bottles were recovered from the mooring line and subsampled for the 6 analysis of bulk N2 fixation rates, DDN released to the dissolved pool, abundance of targeted 7 diazotrophs using qPCR, picophytoplankton and bacterial counts, and nanoSIMS analyses on 8 UCYN-C and non-diazotrophs (diatoms and the 0.2-2 µm fraction) to assess the DD¹⁵N transfer 9 from diazotrophs to non-diazotrophs. All analytical protocols are detailed below in section 2.3. 10 Three 4.5 L bottles were kept as unamended controls (i.e., without ${}^{15}N_2$ addition) and were 11 immediately subsampled for the same parameters. 12

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14 2.3 Analytical protocols

15 2.3.1 Mass spectrometry analyses

PON content and PON ¹⁵N enrichment of samples collected for N_2 fixation rates determination were measured using a Delta Plus Thermo Fisher Scientific isotope ratio mass spectrometer (Bremen, Germany) coupled with an elemental analyzer (Flash EA, ThermoFisher Scientific). N_2 fixation rates were calculated according to the equations detailed in Montoya et al. (1996). Rates were considered significant when the ¹⁵N enrichment of the PON was higher than three times the standard deviation obtained from T0 samples. The standard deviation was 0.004 µmol L⁻¹ for PON and 0.0001 atom% for the ¹⁵N enrichment.

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24 2.3.2 Quantification of diazotrophs using qPCR in sediment traps and during the 25 DDN transfer experiment

During the bloom of UCYN-C (days 17 and 19), immediately after sediment trap samples were collected and prior to their fixation with formalin, trap material was homogenized and fresh aliquots of 1 mL were subsampled from each jar (trap from M1, M2, and M3) and filtered onto 0.2 μ m Supor (Pall-Gelman) filters, flash frozen in liquid N₂, and stored at -80 °C until analysis. For the DDN transfer experiment, after each incubation period, 2 L from each triplicate ¹³C and ¹⁵N₂-labeled 4.5 L bottle were subsampled and filtered through 0.2 μ m Supor (Pall-Gelman)

filters using gentle peristaltic pumping, and stored as described above. The abundance of eight 1 diazotrophic phylotypes was determined using Taqman® qPCR assays: unicellular 2 cyanobacterial groups A1 (UCYN-A1; (Church et al., 2005)), A2 (UCYN-A2; (Thompson et al., 3 2014)), B (UCYN-B or Crocosphaera spp.; (Moisander et al., 2010)), and C (UCYN-C; (Foster 4 et al., 2007)), the filamentous, colonial cyanobacteria *Trichodesmium* spp. (Church et al., 2005), 5 the two DDAs Richelia associated with both Rhizosolenia (het-1; (Church et al., 2005)) and 6 Hemiaulus (het-2; (Foster et al., 2007)) diatoms, Calothrix associated with Chaetoceros (het-3; 7 (Foster et al., 2007)), as well as a heterotrophic phylotype of gamma proteobacteria (γ -24474A11; 8 (Moisander et al., 2008)). All procedures are described extensively in the companion paper by 9 (Turk-Kubo et al., 2015). Briefly, DNA was extracted using a Oiagen DNeasy kit with 10 modifications to recover high quality genomic DNA from cyanobacteria including a freeze thaw 11 12 step, agitation and a proteinase K digestion. Extracts were tested for the presences of PCR inhibitors, compounds sometimes present in DNA extracts from the environment or introduced in 13 the extraction process that reduce PCR efficiency, using either the UCYN-B or the UCYN-C 14 15 assay. If recovery of the spiked standard template in the sample extract was <98%, the sample 16 was considered inhibited, and diluted 1:10 with 5 kD filtered milliQ water. All extracts from the sediment traps showed inhibition when undiluted, and no inhibition when diluted 1:10. DNA 17 18 extracts from the DDN transfer experiment showed no inhibition. All qPCR reactions were carried out on diluted extracts as described in (Goebel et al., 2010). The limit of detection (LOD) 19 and limit of quantitation (LOQ) was 250 and 2000 nifH copies mL⁻¹, respectively, for the 20 sediment trap samples. The LOD and LOQ for DDN transfer experiment samples was 29 and 229 21 *nif*H copies L^{-1} , respectively. 22

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24 2.3.3 Quantification of the net release of DDN to the dissolved pool during the 25 DDN transfer experiment

After each incubation period, 60 mL from each ${}^{15}N_2$ -labeled 4.5 L bottle were subsampled and filtered through pre-combusted (4 h, 450 °C) GF/F filters and immediately frozen for later quantification of ${}^{15}N$ release (i.e., DDN release) to the total dissolved N pool (TDN; i.e., the sum of NO₂⁻, NO₃⁻, NH₄⁺, and DON). The dissolved N was oxidized to NO₃⁻ using the persulfate oxidation method of Knapp et al. (2005) with the amendments of Fawcett et al. (2011). Briefly, 1 mL of potassium persulfate oxidizing reagent (POR) was added to duplicate 5 mL aliquots of

each subsample in 12 mL pre-combusted glass Wheaton vials, and to triplicate vials containing 1 varying quantities of two L-glutamic acid standards, USGS-40 and USGS-41 (Qi et al., 2003) 2 used to ensure complete oxidation and quantify the POR-associated N blank. The POR was made 3 by dissolving 6 g of sodium hydroxide and 6 g of four-times recrystallized, methanol-rinsed 4 potassium persulfate in 100 mL of ultra-high purity water (DIW). Sample vials were capped 5 tightly after POR addition, and autoclaved at 121°C for 55 minutes on a slow-vent setting. The 6 entire oxidation protocol was performed in duplicate (yielding a total of 4 oxidized aliquots for 7 each subsample). 8

9 The concentration of the resultant NO_3^- (i.e., TDN + the POR-associated N blank) was measured by chemiluminescence (Braman and Hendrix, 1989), after which the TDN isotopic composition 10 11 was determined using the 'denitrifier method', wherein denitrifying bacteria that lack N₂O reductase quantitatively convert sample NO_3^- to N_2O (Casciotti et al., 2002; Sigman et al., 2001). 12 13 The denitrifying bacteria (see below) are extremely sensitive to pH; care was thus taken to lower sample pH to 7-8 after POR oxidation via the addition of 12N ACS grade HCl. The $^{15}\mathrm{N}$ 14 enrichment of the N2O was measured by GC-IRMS using a Delta V isotope ratio mass 15 spectrometer and custom-built on-line N₂O extraction and purification system. The international 16 reference materials, IAEA-N3, USGS-34, USGS-32, and an in-house N₂O standard were run in 17 parallel to monitor bacterial conversion and mass spectrometry, and each oxidized sample was 18 analyzed twice. The final TDN concentration and ¹⁵N atom% were corrected for the N blank 19 associated with the POR. The DDN released to the TDN pool was calculated according to: ¹⁵N 20 release (nmol L⁻¹ d⁻¹) = (${}^{15}N_{ex} \times TDN_{con}$)/N_{sr}, where ${}^{15}N_{ex}$ is the atom% excess of the TDN for a 21 given time point; the TDN_{con} is the TDN concentration measured at each time point, and N_{sr} is 22 the ${}^{15}N$ enrichment of the source pool (N₂) in the experimental bottles (i.e., 2.4±0.2 atom% ${}^{15}N$; 23 see above). 24

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26 **2.3.4** Picophytoplankton and bacteria counts during the DDN transfer experiment

After each incubation period, 3.6 mL from each ¹⁵N₂-lableled 4.5 L bottle were subsampled into cryotubes, fixed with paraformaldehyde (2 % final concentration), flash frozen in liquid N₂, and stored at -80°C until analysis. Picoplanktonanalyses were carried out at the PRECYM flow cytometry platform (https://precym.mio.univ-amu.fr/). Samples were analyzed using a FACSCalibur (BD Biosciences, San Jose, CA). For heterotrophic bacterial abundance (BA), 1.8

mL of seawater was fixed with formaldehyde (2 % final concentration, 15 minutes incubation at 1 room temperature in the dark), frozen and_stored in liquid N₂ until analysis in the laboratory. 2 After thawing, 0.3 mL of each samples was incubated with SYBR Green II (Molecular Probes, 3 final conc. 0.05 % [v / v], for 15 minutes at room temperature in the dark), for the nucleic acid 4 staining, according to Marie et al. (2000). Cells were characterized by 2 main optical signals: side 5 6 scatter (SSC), related to cell size, and green fluorescence (530/40), related to nucleic acids 7 staining. Based on these criteria, two subsets of bacteria (referred to low- and high nucleic acidcontaining, or LNA and HNA, respectively) were optically resolved in all samples based on their 8 green fluorescence intensity (Gasol et al., 1999). Just before analysis, 2 µm beads (Fluoresbrite 9 YG, Polyscience), used as an internal control, and TruCount beads (BD Biosciences), used to 10 11 determine the volume analyzed, were added to the samples. To assess autotrophic picoplankton 12 abundances, the red fluorescence (670LP, related to chlorophyll a content) was used as trigger 13 signal and phytoplankton cells were characterized by 3 other optical signals: forward scatter (FSC, related to cell size), side scatter (SSC, related to cell structure), and the orange 14 15 fluorescence (580/30, related to phycoerythrin content). The 2 µm beads (Fluoresbrite YG, Polyscience) were also used to discriminate picoplankton (< $2 \mu m$) from nanoplankton (> $2 \mu m$) 16 populations. The flow rate was estimated by weighing 3 tubes of samples before and after a 3 17 minutes run of the cytometer. The cells concentration was determined from both Trucount beads 18 19 and flow rate measurements. All data were collected in log scale and stored in list mode using the CellQuest software (BD Biosciences). Data analysis was performed a posteriori using SUMMIT 20 v4.3 software (Dako). on a FACScalibur flow cytometer (BD Biosciences, Franklin Lakes, NJ) at 21 22 the Regional Flow Cytometry Platform for Microbiology (PRECYM) (https://precym.mio.univamu.fr/). Standard protocols (Marie et al., 1999) were used to enumerate phytoplankton and 23 24 heterotrophic prokaryotes. Samples were thawed at room temperature in the dark, homogenized by gentle shaking, and filtered through 20 µm strainers in order to avoid large aggregates 25 clogging the instrument fluidics. Just before analysis, 1 mL of sample was transferred into a flow 26 cytometry tube, and 10 µL of a 2 µm fluorosphere (FluoresbryteTM, Polysciences) solution were 27 28 added. These beads were used both as an internal control, and to discriminate cell clusters. Flow cytometric analyses of heterotrophic prokaryotes required pre-staining with a fluorescent nucleic 29 acid probe, SYBR® Green I (Sigma, Germany), at a 1:1000 v/v final dilution of the commercial 30 solution (excitation 488 nm/emission 530 nm). Prior to analysis, samples were incubated with 31

SYBR® Green I for 15 minutes in the dark at room temperature. Side scatter (SSC) was used as
 trigger signal and SYBR® Green I green fluorescence was collected in the green range of 510 550 nm. Combining SYBR® Green I fluorescence and light scattering unambiguously
 distinguishes cells from inorganic particles, detritus, and free DNA (Marie et al., 1999).

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6 2.3.5 Microscopic cell counts during the DDN transfer experiment

Microscopic cell counts. In parallel with the picoplankton counts, diatoms, dinoflagellates, and ciliates were enumerated from 100 mL subsamples collected from each mesocosm that were preserved in Lugol's solution following the Utermöhl method (Hasle, 1978). Cells were counted on a Nikon Eclipse TE2000-E inverted microscope equipped with phase-contrast and a long distance condenser. All groups were quantified in each sample, and diatoms were identified to the lowest possible taxonomic level to examine potential community composition changes and help us to prioritize nanoSIMS analyses.

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15 2.3.6 NanoSIMS analyses and ¹³C and ¹⁵N assimilation rates during the DDN 16 transfer experiment

nanoSIMS analyses. After each incubation period (24, 36 and 72 h), 250 mL from each labeled 17 4.5 L bottle were subsampled, fixed with 25 mL of paraformaldehyde (2 % final concentration) 18 19 and incubated for 24 h at 4 °C, then filtered successively through 25 mm diameter 10 µm, 2 µm, 20 and 0.2 µm pore size polycarbonate filters and rinsed with 0.2 µm filtered seawater. All filters were then sputtered with gold and palladium to ensure conductivity prior to nanoSIMS analyses. 21 22 Diatoms and UCYN-C were analysed on the 10 μ m filters, and the picoplanktonic (0.2-2 μ m) fraction was analysed on the 0.2 µm filters. Diatoms were easily recognized on the CCD (charge 23 coupled device) camera of the nanoSIMS, as were UCYN-C that formed large aggregates of 24 cells, facilitating their recognition for nanoSIMS targeted analyses. However, we cannot exclude 25 the possibility that some UCYN-B were analysed, despite being present at very low abundances, 26 i.e., almost two orders of magnitude less abundant than UCYN-C (Fig. 5) in the analysed 27 28 samples. Several analyses were performed for each group of cells of interest (an average of ~ 25 cells analysed for UCYN-C and diatoms, and between 62 and 140 cells analysed for the 0.2-2 µm 29 30 fraction per time point) to assess the variability of their isotopic composition. A total of ~400 individual cells were analysed by nanoSIMS in this experiment to ensure the robustness of the 31

data. NanoSIMS analyses were performed on a N50 (Cameca, Gennevilliers France) at the 1 French National Ion MicroProbe Facility according to methods previously described (Bonnet et 2 al., Accepted). A 1.3-3 pA 16 keV Cesium (Cs⁺) primary beam focused onto a ~100 nm spot 3 diameter was scanned on a 256 x 256 or 512 x 512 pixel raster (depending on the raster areas, 4 which ranged from 15 μ m x 15 μ m to 50 μ m x 50 μ m) with a counting time of 1 ms per pixel. 5 Samples were implanted with Cs⁺ prior to analysis to remove surface contaminants and increase 6 conductivity. For diatoms, the pre-implant was longer and with higher voltage (2-5 min, 17 pA) 7 to penetrate the silica shell. Negative secondary ions ¹²C⁻, ¹³C⁻, ¹²C¹⁴N⁻, ¹²C¹⁵N⁻, and ²⁸Si⁻ were 8 detected with electron multiplier detectors, and secondary electrons were imaged simultaneously. 9 Ten to fifty serial quantitative secondary ion mass planes were generated and accumulated in the 10 11 final image. Mass resolving power was ~8000 in order to resolve isobaric interferences. Data were processed using the look@nanosims software package (Polerecky et al., 2012). All scans 12 13 were first corrected for any drift of the beam during acquisition, and C and N isotope ratio images were created by adding the secondary ion counts for each recorded secondary ion for each pixel 14 15 over all recorded planes and dividing the total counts by the total counts of a selected reference mass. Individual cells were easily identified in nanoSIMS secondary electron, ¹²C⁻, ¹²C¹⁴N⁻, and 16 ²⁸Si images that were used to define regions of interest (ROI) around individual cells (²⁸Si data 17 are not presented here). For each ROI, the ¹⁵N and ¹³C enrichments were calculated. 18

¹⁵N assimilation rates were calculated for individual cells analysed by nanoSIMS. Our goal was 19 to determine the biological compartment to which the ¹⁵N had been transferred. These were 20 performed after 24 h of incubation. Calculations were performed as follows (Foster et al., 2011; 21 Foster et al., 2013): Assimilation (mol N cell⁻¹ d⁻¹) = $({}^{15}N_{ex} \times N_{con})/N_{sr}$, where ${}^{15}N_{ex}$ is the excess 22 atom% of the individual cells measured by nanoSIMS after 24 h of incubation; the N_{con} is the N 23 content of each cell determined as described below, and N_{sr} is the ¹⁵N enrichment of the source 24 pool (N₂) in the experimental bottles (i.e. 2.4 ± 0.2 atom% ¹⁵N in this experiment). The cell-25 specific N assimilation rate was then multiplied by the cell number enumerated for each group of 26 phytoplankton and bacteria by microscopy and flow cytometry. Standard deviations were 27 calculated using the variability of ¹⁵N enrichment measured by nanoSIMS on replicate cells and 28 the standard deviation of the estimated cellular N content (see below) of UCYN-C, non-29 diazotrophic phytoplankton, and bacteria. Final standard deviations were calculated according to 30 propagation of errors laws. 31

To determine the N_{con} of diatoms, cell cross section, apical and transapical dimensions were 1 measured on the dominant diatom species present in the mesocosms and analysed by nanoSIMS 2 to calculate biovolumes. All dimensions were measured on at least 20 cells using a Nikon Eclipse 3 TE2000-E inverted microscope equipped with phase-contrast and a long distance condenser. 4 Dimensions were entered into the international diatom data base (Leblanc et al., 2012) in which 5 bio-volumes are calculated following the geometric model of each cell type as described in (Sun 6 and Liu, 2003). Carbon (C) content (C_{con}) was then calculated for the species of interest using the 7 equations of (Eppley et al., 1970) and (Smayda, 1978). For Synechococcus spp. and 8 picoeukaryotes, we used C_{con} data from Fu et al. (2007) (249±21 fg C cell⁻¹) and Yentsch and 9 Phinney (1985) (2100 fg cell⁻¹), respectively. C_{con} was then converted to N_{con} using the Redfield 10 ratio of 6.6:1 (Redfield, 1934). For bacteria, an average N_{con} of 5.8±1.5 fg N cell⁻¹ (Fukuda et al., 11 1998) was used. For UCYN-C, cell dimensions were measured and the bio-volume was 12 13 calculated based on the equations reported in Sun and Liu (2003). C_{con} was then calculated using the relationship between bio-volume and C_{con} (Verity et al., 1992) (22 pg cell⁻¹). C_{con} was then 14 converted to N_{con} (2.3 pg cell⁻¹) using a ratio of 8.5:1 (Berthelot et al., 2015a). 15

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17 2.4 Statistical analyses

18 Spearman correlation coefficients were used to examine the relationships between N₂ fixation 19 rates, hydrological, biogeochemical, and biological variables in the mesocosms (n=57 to 61, 20 α =0.05). The methods used to analyze the parameters reported in the correlation table are 21 described in detail in companion papers in this issue (Berthelot et al., 2015b; Bonnet et al., 2016; 22 Leblanc et al., 2016; Turk-Kubo et al., 2015).

A non-parametric Mann-Whitney test (α =0.05) was used to compare the means of N₂ fixation rates obtained using the dissolution and the bubble method, as well as to compare the means of N₂ fixation between the different phases of the experiment, mean isotopic ratios between ¹⁵N₂enriched and natural abundance of N (0.366 atom%), and mean isotopic ratios between T24 h and T72 h in the DDN transfer experiment.

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1 3 Results

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3 3.1 N₂ fixation rates in the mesocosms

Bulk N₂ fixation rates averaged 18.5 \pm 1.1 nmol N L⁻¹ d⁻¹ throughout the 23 days of the experiment 4 in the three mesocosms (all depths averaged together) (Table 1). The variance between the three 5 6 mesocosms was low, and the temporal dynamics of the rates were similar (Fig. 2, Table 1), indicating good replicability between the mesocosms. Based on our data on N2 fixation 7 dynamics, we could identify three main periods during the experiments. These three periods were 8 9 also defined by Berthelot et al. (2015b) based on biogeochemical characteristics and by Turk-Kubo et al. (2015) based on changes in abundances of targeted diazotrophs. During the first 10 period (P0; from day 2 to 4, i.e., prior to the DIP fertilization), the average bulk N₂ fixation rate 11 for the three mesocosms was 17.9 ± 2.5 nmol N L⁻¹ d⁻¹ (Fig. 2a). These N₂ fixation rates decreased 12 significantly (p<0.05) by ~40 % from day 5 to ~15 (hereafter called P1) to 10.1 ± 1.3 nmol N L⁻¹ 13 d^{-1} , then increased significantly (p<0.05) from day 15 until the end of the experiment (day 15 to 14 23, hereafter called P2) to an average of 27.3 ± 1.0 nmol N L⁻¹ d⁻¹ (Fig. 2a). Maximum rates were 15 reached during P2 (between days 18 and 21) with 69.7, 67.7 and 60.4 nmol N $L^{-1} d^{-1}$ in M1 (12 16 m), M2 (6 m) and M3 (12 m), respectively. From day ~15 to 21, N₂ fixation rates were higher at 17 12 m depth than in the surface. The difference was significant in M2 and M3 (p<0.05), but not in 18 19 M1 (p>0.05). Size fractionation experiments indicate that 37 ± 7 % of the measured N₂ fixation was associated with the <10 µm size fraction (Fig. 2b), and N₂ fixation rates in this fraction 20 followed the same temporal trend as bulk N₂ fixation. These data indicate that for the experiment 21 as a whole, the majority (~63 %) of the N₂ fixation was associated with the >10 μ m fraction. N₂ 22 fixation rates measured in the lagoon waters were half those measured in the mesocosms, and 23 were on average 9.2 \pm 4.7 nmol N L⁻¹ d⁻¹ over the 23 days of the experiment. 24

The Spearman correlation matrix (Table 2) indicates that N_2 fixation was positively correlated with seawater temperature in the mesocosms, which was not the case in lagoon waters, although temperature was exactly the same inside and outside the mesocosms (from 25.4°C to 26.8°C) (Bonnet et al., 2016). N₂ fixation in the mesocosms was also positively correlated with particulate organic carbon (POC), PON, and particulate organic phosphorus (POP) (except in M2) concentrations, Chl *a* concentrations, primary production, bacterial production, alkaline phosphatase activity (APA), and *Synechococcus*, picoeukaryote and nanoeukaryote (except in M2) abundances. N₂ fixation was negatively correlated with NO₃⁻, DIP, DON, dissolved organic
 phosphorus (DOP) (except in M2) concentrations and DIP turn-over time.

The intercomparison between the bubble and dissolution methods performed on day 11 in M2 indicates that rates determined for the 6 replicates were 7.2 ± 0.8 nmol N L⁻¹ d⁻¹ and 6.4 ± 2.0 nmol N L⁻¹ d⁻¹ for the dissolution method and the bubble method, respectively, demonstrating that, at least in this study, N₂ fixation rates were not significantly different (p>0.05) between the two methods.

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9 3.2 Phenotypic characterization of UCYN by microscopy

The average size of the UCYN-C cells present in the mesocosms was $5.7\pm0.8 \ \mu m$ (n=17). Both 10 11 free-living and aggregated UCYN-C cells were observed in the water columns of the mesocosms. However, the detailed microscopic analysis performed on day 17 and day 19 in M2 (during the 12 13 bloom of UCYN-C) (Fig. 3) indicates that the proportion of free-living cells (ROI characterized by one cell or two cells defined as dividing cells) was low (<1 % on day 17 and <5 % on day 19). 14 15 The average number of UCYN-C cells per aggregate increased with depth (Fig. 3a), with the size of the aggregates reaching 50-100 µm at 6 m and 100-500 µm at 12 m depth. On day 17, the 16 number of cells per aggregate averaged 162, 74, and 1273 at 1, 6, and 12 m, respectively. On day 17 19, the aggregates were much smaller (\sim 50 µm) with only 4, 11, and 19 cells per aggregate. The 18 19 sediment traps contained extremely high densities of UCYN-C cells with the average number of cells per aggregate 60 to 50,000 times higher than that measured in the water column aggregates 20 (Fig. 3b-e). 21

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3.3 Quantification of diazotrophs in sediment traps

qPCR analysis confirmed that UCYN-C was the most abundant diazotroph in the sediment traps 24 on days 17 and 19, with abundances reaching 2.7 x 10^8 to 4 x 10^9 nifH copies L⁻¹ (Fig. 4a). 25 UCYN-C accounted for 97.4 to 99.2 % of the total *nif*H pool quantified in the traps. Abundances 26 were higher in M2 and M3 (1.8 x 10^9 in M2 and 3 x 10^9 *nif*H copies L⁻¹ in M3) compared to M1 27 $(2.5 \times 10^8 \text{ nifH copies L}^{-1})$ on day 19. Het-1 and het-3 were always recovered in the sediment 28 traps, albeit at lower abundances (1.8 to 8.6 x 10^6 *nif*H copies L⁻¹ for het-1 and 4.9 x 10^6 to 2.8 x 29 10^7 nifH copies L⁻¹ for het-3) (Fig. 4b). They represented between 0.1 and 1.8 % of the targeted 30 nifH pool. UCYN-B was detected in all mesocosms on both days (except in M1 on day 19), and 31

UCYN-A2 and *Trichodesmium* were detected in M2 on day 17 but at low abundances (0.05 % of
 the total *nif*H pool) compared to the other phylotypes. Het-2 was never detected in the traps, and
 neither was γ24774A11 or UCYN-A1.

Using the volume of each mesocosm (Bonnet et al., 2016) and the total nifH copies for each 4 diazotroph phylotype in the sedimenting material and in the water column the day before the 5 collection of the sediment traps (Turk-Kubo et al., 2015) (assuming a sinking velocity of the 6 exported material of ~ 10 m day⁻¹, Gimenez et al. (2016)), we estimated the export efficiency for 7 each phylotype. For UCYN-C, 4.6 % and 6.5 % of the cells present in the water column were 8 exported to the traps per 24 h on day 17 and 19, respectively (assuming one *nif*H copy per cell). 9 For het-1, 0.3 and 0.4 % of cells were exported into the traps on day 17 and 19, for het-3, 15.5 % 10 and 10.5 % were exported, and for UCYN-B, 37.1 % and 15.5 % of UCYN-B were exported on 11 12 day 17 and 19, respectively.

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14 **3.4.** DDN transfer experiment performed on day 17

Net ¹⁵N₂ uptake was 24.1±2.8 nmol N L⁻¹ during the first 24 h of the DDN transfer experiment performed from days 17 to 20 (Fig. 5a). As expected, integrated ¹⁵N₂ uptake increased over the course of the experiment to reach 28.8±4.3 nmol N L⁻¹ at T48 h and 126.8±35.5 nmol N L⁻¹ at T72 h. The DDN quantified in the TDN pool ranged from 6.2±2.4 nmol N L⁻¹ at T24 h to 9.6±1.6 nmol N L⁻¹ at T72 h. Considering gross N₂ fixation as the sum of net N₂ fixation and DDN release (Mulholland et al., 2004), the DDN released to the TDN pool accounted for 7.1±1.2 to 20.6±8.1 % of gross N₂ fixation.

During the 72 h targeted experiment (Fig. 5b) the diazotroph assemblage reflected that of the mesocosms from which they were sampled: UCYN-C dominated the diazotrophic community, comprising on average 62 % of the total *nifH* pool. The other most abundant phylotypes were UCYN-A2 and het-2, which represented 18 and 13 % of the total *nifH* pool, respectively. UCYN-A1, UCYN-B, het-1, het-3, and *Trichodesmium* were also detected but together they comprised less than 8 % of the total targeted community. Phylotype abundances remained relatively stable throughout the 72 h of the experiment.

NanoSIMS analyses performed on individual UCYN-C at 24 h (Fig. 6) revealed significant $(p<0.05)^{13}C$ (1.477±0.542 atom%, n=35) and ^{15}N (1.515±0.370 atom%, n=35) enrichments

31 relative to natural abundance, indicating that UCYN-C were actively photosynthesizing and

fixing N₂. The correlation between ${}^{13}C$ enrichment and ${}^{15}N$ enrichment was significant (r=0.85, 1 p<0.01, Fig. 6b). NanoSIMS analyses performed on diatoms and picoplankton (Fig. 5c) also 2 revealed significant (p<0.05)¹⁵N enrichment of non-diazotrophic plankton, demonstrating a 3 transfer of DDN from the diazotrophs to other phytoplankton. Both diatoms and picoplanktonic 4 cells were significantly (p<0.05) more enriched at the end of the experiment (T72 h) 5 (0.489±0.137 atom%, n=12 for diatoms; 0.457±0.077 atom%, n=96 for picopankton) than after 6 the first 24 h (0.408±0.052 atom%, n=23 for diatoms; 0.389±0.014 atom%, n=63 for 7 picoplankton). Finally, the ¹⁵N enrichment of picoplankton and diatoms was not significantly 8 different (p>0.05) during the DDN experiment. 9

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11 4 Discussion

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13 4.1 The bubble vs. the dissolution method: an intercomparison experiment

The inter-comparison experiment performed on day 11 reveals slightly lower, yet insignificantly 14 15 different (p>0.05) average N₂ fixation rates when using the bubble method compared to the dissolution method. This result is in accordance with some comparisons made by Shiozaki et al. 16 (2015) in temperate waters of the North Pacific. However, a lower degree of dissolution of the 17 ¹⁵N₂ bubble may occur in warm tropical waters such as those near New Caledonia compared to 18 19 the cooler, temperate North Pacific waters. In calculating N_2 fixation rates using the dissolution method, we used the value of 2.4 ± 0.2 atom% for the ^{15}N enrichment of the N₂ pool as measured 20 by MIMS. For the bubble method, we used the theoretical value of 8.4 atom% calculated for 21 seawater with a temperature of 25.5 °C and salinity of 35.3 (as was the case on day 11). If we 22 assume that equilibration was incomplete in our experiment using the bubble method, i.e., 75 % 23 instead of 100 % as shown by Mohr et al. (2010), we calculate higher, albeit still insignificantly 24 (p>0.05) N₂ fixation rates for the bubble method (8.3 \pm 2.8 nmol N L⁻¹ d⁻¹) compared to the 25 dissolution method (7.2 \pm 0.8 nmol N L⁻¹ d⁻¹), confirming that equivalent results are obtained with 26 both methods in this ecosystem. 27

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29 4.2 The temporal dynamics of N₂ fixation in the mesocosms

Average N₂ fixation rates measured in the lagoon waters (outside the mesocosms, 9.2 ± 4.7 nmol N L⁻¹ d⁻¹, Table 1) are of the same order of magnitude as those reported for the Noumea lagoon during austral summer conditions (Biegala and Raimbault, 2008). They are within the upper
range of rates reported in the global ocean database (Luo et al., 2012). Indeed, open ocean cruises
performed offshore of New Caledonia in the Coral and Solomon Seas (e.g., (Bonnet et al., 2015;
Garcia et al., 2007) also suggest that the South West Pacific Ocean is one of the areas with the
highest N₂ fixation rates in the global ocean.

Averaged over the 23 days of the experiment, N_2 fixation rates in the mesocosms were ~ 2 fold 6 higher (18.5 \pm 1.1 nmol N L⁻¹ d⁻¹) than those measured in lagoon waters (9.2 \pm 4.7 nmol N L⁻¹ d⁻¹). 7 The maximum observed rates of >60 nmol N $L^{-1} d^{-1}$ from days 18-21 are among the highest 8 reported for marine waters (Luo et al., 2012). DIP concentration was the predominant difference 9 between the ambient lagoon waters and those of the mesocosms. The mesocosms were fertilized 10 with DIP on day 4, reaching ambient concentrations of ~ 0.8μ mol L⁻¹ compared to lagoon waters 11 in which DIP concentrations were typically <0.05 µmol L⁻¹. According to our experimental 12 13 assumption, diazotrophy would be promoted by high concentrations of DIP. Yet, in all three mesocosms, N₂ fixation rates were negatively correlated with DIP concentrations and DIP 14 15 turnover time and positively correlated with APA (Table 2). Below, we describe the scenario that 16 likely occurred in the mesocosms, which likely explains these correlations.

During P0 (day 2 to 4), N₂ fixation rates were higher in the mesocosms than in the lagoon waters, 17 possibly due to the reduction of turbulence in the water column facilitated by the closing of the 18 19 mesocosms (Moisander et al., 1997) and/or to the reduction of the grazing pressure in the mesocosms as total zooplankton abundances were slightly lower (by a factor of 1.6) in the 20 mesocosms compared to the lagoon waters (Hunt et al., 2016). The most abundant diazotrophs in 21 22 the mesocosms at P0 were het-1 and *Trichodesmium*, which were probably the most competitive groups under the initial conditions, i.e., NO₃⁻ depletion (concentrations were $0.04\pm0.02 \mu$ mol L⁻¹, 23 Table 3) and low DIP concentrations (0.03 \pm 0.01 µmol L⁻¹, Table 3). *Trichodesmium* is able to 24 use organic P substrates (DOP pool) under conditions of DIP deficiency (Dyhrman et al., 2006; 25 Sohm and Capone, 2006). 24 h after the DIP fertilization (day 5), N₂ fixation rates in the 26 mesocosms decreased by ~40 %, reaching rates comparable to those measured in lagoon waters 27 28 during P1 (day 5 to 14). Enhanced DIP availability likely enabled non-diazotrophic organisms with lower energetic requirements and higher growth rates to outcompete the diazotrophs in the 29 mesocosms via utilization of recycled N derived from recent N₂ fixation. This is supported by the 30 observation that nano-eukaryotes and non-diazotrophic cyanobacteria such as Prochlorococcus 31

sp. increased in abundance during P1 (Leblanc et al., 2016) in the three mesocosms when N₂
fixation rates declined (Fig. 2).

3 During P2 (day 15 to 23), N₂ fixation rates increased dramatically in all three mesocosms. This period was defined by a high abundance of UCYN-C, which were present in low numbers in the 4 5 lagoon and within the mesocosms during P0 and P1 (Turk-Kubo et al., 2015). The increase in UCYN-C abundance was synchronous with a decrease in DIP concentrations in the mesocosms 6 7 (Turk-Kubo et al., 2015): UCYN-C abundance first increased in M1 (day 11), then in M2 (day 13), and finally in M3 (day 15). In all cases, the increase in UCYN-C abundance coincided with 8 9 low DIP turnover time, indicative of DIP deficiency (Berthelot et al., 2015b; Moutin et al., 2005). Under NO₃⁻ depletion and low DIP availability, UCYN-C appeared to be the most competitive 10 diazotroph in the mesocosms, as they exhibited the highest maximum growth rates compared to 11 those calculated for the other diazotrophic phylotypes for the same period (Turk-Kubo et al., 12 13 2015). Some Cyanothece strains possess the genes required for utilization of organic P substrates such as phosphonates (Bandyopadhyay, 2011). Thus, UCYN-C, which were the major 14 15 contributors to N₂ fixation during P2 (see below), may have used DOP as a P source during this period, consistent with the negative correlation observed between N₂ fixation rates and DOP 16 concentrations (except in M2, Table 2), and driving the significant decline in DOP concentrations 17 18 observed in all three mesocosms during P2 (Berthelot et al., 2015b; Moutin et al., 2005).

19 While temperature was not correlated with N₂ fixation in the lagoon, we observed a significant positive correlation between these parameters in the mesocosms (Table 2), probably because 20 some diazotrophic phylotypes present in the mesocosms and absent in the lagoon waters were 21 22 particularly sensitive to seawater temperature. UCYN-C reached high abundances inside the mesocosms, but was virtually absent in the lagoon waters outside the mesocosms. Turk-Kubo et 23 24 al. (2015) showed that UCYN-C abundance was positively correlated with seawater temperature, suggesting that the optimal temperature for UCYN-C growth is above 25.6 °C. This result is 25 consistent with culture studies performed using three UCYN-C isolates from the Noumea lagoon 26 that are closely related to the UCYN-C observed here, indicating maximum growth rates at 27 around 30°C and no growth below 25 °C (Camps, Turk-Kubo, Bonnet, Pers. comm.). 28 Temperature above 25.6 °C and up to 26.7°C were reached on day 12 and were maintained 29 through to the end of the mesocosm experiment, possibly explaining why UCYN-C was not 30

observed during P0 (when temperature was 25.4°C) even though DIP turn-over time was low
(below ~1 d) (Berthelot et al., 2015b; Moutin et al., 2005).

If low DIP concentrations and seawater temperatures greater than 25.6 °C are prerequisites for 3 UCYN-C growth, an obvious question is why they did not thrive (despite being present at low 4 abundances) in the lagoon waters during P2 when similar conditions prevailed. We consider three 5 possible explanations that are discussed extensively in Turk-Kubo et al. (2015): first, it is 6 7 possible that UCYN-C are sensitive to turbulence, which was likely reduced in the mesocosms compared to the lagoon waters that are susceptible to trade winds and tides. Second, grazing 8 pressures on UCYN-C may have been reduced as total zooplankton abundances were slightly 9 lower (by a factor of 1.6) in the mesocosms compared to those in the lagoon waters (Hunt et al., 10 11 2016). Third, the water masses outside the mesocosms changed with tides and winds; thus, it is possible that UCYN-C were absent from the water mass encountered outside the mesocosms 12 13 when we sampled for this experiment.

In the mesocosms, the cell specific ${}^{15}N_2$ fixation rate measured on day 17 (M2) for UCYN-C was 14 $6.3\pm2.0 \times 10^{-17}$ mol N cell⁻¹ d⁻¹. Multiplying this rate by the abundance of UCYN-C indicates that 15 UCYN-C accounted for 90±29 % of bulk N₂ fixation during that period. This is consistent with 16 the positive correlation observed between N₂ fixation rates and UCYN-C abundances in M2 17 (Table 2). In M1 and M3, the correlation was also positive yet insignificant. This may have been 18 19 due to the low number of UCYN-C data points, thus decreasing the sensitivity of the statistical test. Coupling between UCYN-C 13 C and 15 N incorporation was significant (r=0.85, p<0.01) (Fig. 20 6b) and contrasts with results reported by Berthelot et al. (2016) for UCYN-C, in which ¹³C and 21 ¹⁵N enrichment (and thus inorganic C and N₂ fixation) was uncoupled in the cells. Based on their 22 observations, these authors suggest that the heterogeneity in the ¹⁵N and ¹³C enrichments can be 23 explained by a specialization of some cells that induces variability in cell-specific ¹⁵N-enrichment 24 e.g., diazocytes that contain the nitrogenase enzyme in the colonial filamentous Trichodesmium 25 sp. Spatial partitioning of N₂ and C fixation by colonial unicellular types was also evidenced for 26 diazocyte-like formation in colonial Crocosphaera.watsonii-like (UCYN-B) cells (Foster et al., 27 2013). Here, UCYN-C cells fixed both ¹³C and ¹⁵N proportionally, which suggests they did not 28 utilize diazocytes to separate diazotrophy from photosynthesis in our experiments. 29

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31 4.3 UCYN aggregation and export

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Throughout the 23 days of the experiment, the majority of N2 fixation (63 %) occurred in the >101 μm size fraction, even during P2 when the small (5.7±0.8 μm) unicellular UCYN-C dominated 2 3 the mesocosm diazotrophic community. These findings can be explained by the aggregation of UCYN-C cells into large (>10 μ m) aggregates (Fig. 7) that were retained on 10 μ m filters (Fig. 4 3). These large UCYN-C aggregates probably formed in part due to the presence of sticky TEP 5 (Berman-Frank et al., 2016) or other extracellularly-released proteins, and were characterized by 6 7 a high sinking velocity due to their large size (up to 500 µm in diameter) and a density greater than that of seawater (Azam and Malfatti, 2007). Their aggregation and subsequent sinking 8 9 within the mesocosms likely explains why volumetric N₂ fixation rates were higher at 12 m than at the surface during P2, as well as why the size of the aggregates increased with depth, and why 10 11 numerous large-size aggregates and extremely high abundances of UCYN-C were recovered in the sediment traps. Aggregation processes may have been favored by the low turbulence in the 12 13 mesocosms and it would be necessary to confirm that such processes also occur in the open 14 ocean.

15 Colonial phenotypes of UCYN (UCYN-B) have been observed in the water column of the North Tropical Pacific (ALOHA station) (Foster et al., 2013), but to our knowledge, this is the first time 16 that UCYN have been detected in sediment traps. Contrary to published data (e.g. (White et al., 17 2012)), here we demonstrate a greater export efficiency of UCYN (~10 % exported to the traps 18 19 within 24 h) compared to the export of DDAs (efficiency of 0.24 to 4.7 %). Diatoms sink rapidly 20 and DDAs have been found in sediment traps at Station ALOHA (Karl et al., 2012; Karl et al., 1997; Scharek et al., 1999a; Sharek et al., 1999b), in the Gulf of California (White et al., 2012), 21 and in the Amazon River plume (Subramaniam et al., 2008). In our study, we observed limited 22 export of het-1 (Richelia in association with Rhizosolenia) and het-3 (Calothrix) during P2, while 23 24 het-2 (Richelia associated with Hemiaulus) was never recovered in the sediment traps. This is likely because *Hemiaulus* has a lower sinking rate than *Rhizosolenia* due do its smaller size, or 25 may be more easily grazed by zooplankton than *Rhizosolenia* or *Calothrix*, which are known to 26 be toxic to crustaceans (Höckelmann et al., 2009). We observed only rare occurrences of 27 28 Trichodesmium export in this study probably due to its extremely limited presence and low growth rates in the mesocosms. Direct comparisons of our export results with findings from open 29 30 ocean studies should be made cautiously as our mesocosms were shallower (15 m) than typical

oceanic export studies (>100 m) and were also probably characterized by reduced turbulence
 (Moisander et al., 1997).

3 We estimate that the direct export of UCYN-C accounted for 22.4 ± 5.5 % of the total POC exported in each mesocosm at the height of the UCYN-C bloom (day 17) and decreased to 4 4.1±0.8 % on day 19 (Fig. 4c, Fig. 7). This calculation is based on the total POC content 5 measured in the sediment traps (Berthelot et al., 2015b), our C_{con} for UCYN-C estimated as 6 described above, and published C_{con} for other diazotrophs. The corresponding export of het-1, 7 het-3, Trichodesmium, and UCYN-B on day 17 based on published C_{con} (Leblanc et al., 2012; 8 9 Luo et al., 2012), and using an average of three *Richelia* and *Calothrix* symbionts per diatom, accounted for 6.8±0.5, 0.5±0.02, 0.3±0.3, and 0.1±0.01 % of the POC export on day 17, 10 11 respectively, and for 4.2±1.7, 0.04±0.03 of the POC export on day 19 (the contribution of Trichodesmium and UCYN-B to POC export on day 19 was negligible). Thus, our data 12 13 emphasize that despite their small size relative to DDAs, UCYN-C are able to directly export organic matter to depth by forming densely-populated aggregates that can rapidly sink. This 14 15 observation is further confirmed by the *e* ratio, which quantifies the efficiency of a system to export POC relative to primary production (e ratio = POC export/PP) and was significantly 16 higher (p<0.05) during P2 (i.e., during the UCYN-C bloom; 39.7±24.9 %) than during P1 (i.e., 17 when DDAs dominated the diazotrophic community; 23.9±20.2 %) (Berthelot et al., 2015b). It is 18 19 also consistent with the significantly (p<0.05) higher contribution of N₂ fixation to export production during P2 (56±24 %, and up to 80 % at the end of the experiment) compared to P1 20 (47±6 %, and never exceeded 60 %) as estimated by Knapp et al. (2015) using a $\delta^{15}N$ budget for 21 the mesocosms. Our calculated contribution of N₂ fixation to export production is very high 22 compared to other tropical and subtropical regions where diazotrophs are present (10 to 25 %; 23 e.g., (Altabet, 1988; Knapp et al., 2005)). However, it is consistent with the high rates of N_2 24 fixation measured in the enclosed mesocosms compared to those from the lagoon and other 25 tropical pelagic studies (Luo et al., 2012). The direct export of UCYN-C and other diazotrophs 26 cannot solely explain the high e ratio estimated for P2. We thus hypothesize that a fraction of the 27 28 DDN export that occurred during P2 was transferred indirectly via primary utilization by nondiazotrophic plankton cells that were eventually exported to the sediment traps (Fig. 7). 29

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1 **4.4 DDN transfer to non-diazotrophic phytoplankton and ecological** 2 implications

3 The amount of DDN measured in the TDN pool during the 72 h DDN transfer experiment is higher than that reported for culture studies of Cyanothece populations $(1.0\pm0.3 \text{ to } 1.3\pm0.2 \text{ \% of}$ 4 gross N₂ fixation (Benavides et al., 2013; Berthelot et al., 2015a)). The DDN measured in the 5 TDN pool reflects the DDN release by diazotrophs during N₂ fixation and is likely 6 underestimated here as a fraction of this DDN has been taken up by surrounding planktonic 7 communities. In our experiment, other diazotrophs were present in addition to Cyanothece, and 8 9 they may have also contributed to the dissolved pool. Moreover, unlike in culture studies, field experiments are also impacted by other exogenous factors such as viral lysis (Fuhrman, 1999) 10 11 and sloppy feeding (O'Neil and Roman, 1992; Vincent et al., 2007), which may enhance N release. 12

13 This DDN release plays a critical role in the N transfer between diazotrophs and non-diazotrophs. The cell-specific uptake rates of DDN during the DDN transfer experiment were calculated for 14 15 each cell analysed by nanoSIMS (diatoms and cells from the 0.2-2 µm fraction). By multiplying cell-specific N uptake rates by the cellular abundance of each group on a particular day, we could 16 identify the specific pool (diazotrophs, dissolved pool, non-diazotrophs) into which the DD¹⁵N 17 was transferred after 24 h, and the extent to which this ¹⁵N₂ accumulated. The results are 18 summarized in Fig. 7. After 24 h, 52 ± 17 % of the newly fixed ${}^{15}N_2$ remained in the UCYN-C 19 biomass, 16±6 % had accumulated in the dissolved N pool, and 21±4 % had been transferred to 20 non-diazotrophic plankton. In addition, 11 % of the newly fixed ¹⁵N₂ accumulated in a pool that 21 we refer to as 'others' (corresponding to diazotrophs other than UCYN-C and potential non-22 diazotrophs to which ¹⁵N₂ was transferred; these cells were not analysed by nanoSIMS due to 23 their very low abundance). Uncertainties take into account both the variability of the ¹⁵N 24 enrichment determined on ~25 cells per group by nanoSIMS, and the uncertainty in the N content 25 per cell measured or taken from the literature. 26

Within the fraction of DDN transferred to the non-diazotrophs after 24 h (21 %), we calculate that 18 ± 4 % was transferred to picoplankton, and only 3 ± 2 % was transferred to diatoms (Fig. 7). The ¹⁵N enrichment of picoplankton and diatoms was not significantly different (p>0.05) in this study, but as picoplankton dominated the planktonic community in the mesocosms at the time of the DDN transfer experiment, they were the primary beneficiaries of the DDN. This is consistent

with the positive correlation between N₂ fixation rates, Synechococcus, and pico-eukaryote 1 abundances in the mesocosms (Table 2), as well as with the observed dramatic increase in 2 3 Synechococcus and pico-eukaryotes abundances (by a factor of >2 between P1 and P2) (Leblanc et al., 2016). Diatom abundances also increased in the mesocosms by a factor of 2 between P1 4 and P2 (largely driven by Cylindrotheca closterium), but this increase occurred earlier than the 5 6 picoplankton increase, i.e., at the end of P1 (days 11-12). Maximum diatom abundances were 7 reached on day 15-16 at the very beginning of P2, and then declined by day 18 to reach abundances similar to those observed during P1. These results suggest that diatoms were the 8 primary beneficiaries of DDN in the mesocosms at the start of P2, when N₂ fixation rates and 9 UCYN-C abundances increased dramatically. This is consistent with a previous DDN transfer 10 11 study performed in New Caledonia (Bonnet et al., Accepted) during which diatoms (mainly Cylindrotheca closterium) advantageously competed and utilized DDN released during 12 13 Trichodesmium blooms. When the present DDN transfer experiment was performed (days 17 to 20), diatom abundances had already declined, likely due to DIP limitation (DIP turnover time 14 15 was low, i.e below 1d). We hypothesize that picoplankton were more competitive for DDN under low DIP conditions as small cells with high surface to volume ratios are known to outcompete 16 larger cells for the available DIP (Moutin et al., 2002). Moreover, some prokaryotes from the 0.2-17 2 µm size-fraction can utilize DOP compounds (Duhamel et al., 2012). In this study, we could 18 19 not discriminate the DDN transfer to pico-autotrophs from that to pico-heterotrophs, but it is likely that both communities took advantage of the DDN, as both primary production (Berthelot 20 et al., 2015b) and bacterial production (Van Wambeke et al., 2015) were positively correlated 21 with N₂ fixation rates (Table 2) and increased dramatically following the increase in N₂ fixation 22 during P2. The standing stocks of POC, PON, and POP were also positively correlated with N₂ 23 24 fixation rates, suggesting that DDN sustained productivity in the studied system.

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26 **5 Conclusions**

While studies on the fate of DDN in the ocean are rare, the contribution of DDN to particle export based on the δ^{15} N signatures of exported material indicate that N₂ fixation can efficiently contribute to export production in the oligotrophic ocean (Dore et al., 2008). The export of DDN may either be direct, through the sinking of diazotrophs, or indirect, through the transfer of DDN to non-diazotrophic plankton in the photic zone that are subsequently exported. 1 *Trichodesmium* is rarely recovered in sediment traps (Walsby, 1992) and most of the research 2 dedicated to the export of diazotrophs has focused on DDAs (Karl et al., 2012) due to their high 3 sinking velocity. Here, we demonstrate for the first time that UCYN can efficiently contribute to 4 POC export in oligotrophic systems, predominantly due to the aggregation of small ($5.7\pm0.8 \mu m$) 5 UCYN-C cells into large aggregates, which increase in size (up to 500 μm) with depth. Our 6 results suggest that these small (typically 3-7 μm) organisms should be considered in future 7 studies to confirm if processes observed in mesocosms are applicable to open ocean systems.

Moreover, the experimental and analytical approach used in this study allowed for the 8 9 quantification of the actual transfer of DDN to different groups of non-diazotrophic plankton in the oligotrophic ocean. Our nanoSIMS results coupled with ¹⁵N₂ isotopic labelling revealed that a 10 11 significant fraction of DDN (21±4 %) is quickly (within 24 h) transferred to non-diazotrophic plankton, which increased in abundance simultaneously with N2 fixation rates. A similar 12 nanoSIMS study performed during a Trichodesmium bloom (Bonnet et al., Accepted) revealed 13 that diatoms were the primary beneficiaries of DDN and developed extensively during and after 14 15 Trichodesmium spp. blooms. Diatoms are efficient exporters of organic matter to depth (Nelson et al., 1995). These studies show that plankton grown on DDN in the oligotrophic ocean drive 16 indirect export of organic matter out of the photic zone, thus revealing a previously unaccounted 17 for conduit between N₂ fixation and the eventual export to depth of DDN from the photic zone. 18

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Author contribution: S. Bonnet designed the experiments and S. Bonnet and H. Berthelot carried them out with help from E. Rahav. S. Bonnet, H. Berthelot, K.A. Turk-Kubo, S. Fawcett and S. L'Helguen analyzed the samples. I. Berman-Frank took part in experimental planning, preparation, and implementation of the project. S. Bonnet prepared the manuscript with contributions from all co-authors.

6

7 Figure captions

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9 Figure 1. (a) Mesocosms (~50 m³) deployed in the framework of the VAHINE project. (b)
10 Sediment traps screwed onto the base of the mesocosms and were sampled daily by SCUBA
11 divers.

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Figure 2. (a) Horizontal and vertical distributions of bulk N_2 fixation rates (nmol N L⁻¹ d⁻¹), and (b) <10 µm N_2 fixation rates (nmol N L⁻¹ d⁻¹) in M1, M2, M3, and lagoon waters. Note that N_2 fixation rates in the <10 µm fraction were not measured (lower right panel). The grey bars indicate the timing of the DIP addition on day 4.

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Figure 3. (a) UCYN-C cells per aggregate in M2 on day 17 and 19. (b to e) Green excitation (510-560 nm) epifluorescent replicate micrographs of UCYN-C on day 17 taken at 1 m depth (x40) (b), 6 m depth (x40) (c), 12 m depth (x40) (d), and in the sediment traps (x10) (e). Scale bar 20 μ m (b to d) and 100 μ m (e).

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Figure 4. (a) Abundance of UCYN-C (*nif*H copies L^{-1}) and (b) other *nif*H phylotypes (UCYN-A2, UCYN-B, *Trichodesmium*, het-1, het-3) (*nif*H copies L^{-1}) recovered in the sediment trap on day 17 and 19. (c) Proportion of POC export associated with diazotrophs in the sediment traps on day 17 in M2 (height of UCYN-C bloom).

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Figure 5. Results from the DDN transfer experiment performed from day 17 to 20 in M2. (a) Temporal changes in ${}^{15}N_2$ uptake (white, nmol N L⁻¹) and quantification of DDN in the dissolved pool (grey) over the course of the experiment. Error bars represent the standard deviation of three independent replicate incubations. (b) Temporal changes in diazotroph abundance determined by QPCR (*nif*H gene copies L⁻¹) during the same experiment. Error bars represent the standard deviation of triplicate incubations. (c) Summary of the nanoSIMS analyses. Measured ¹³C and ¹⁵N atom% values of non-diazotrophic diatoms (white) and picoplankton (grey) as a function of incubation time. The horizontal dashed line indicates the natural abundance of ¹⁵N (0.366 atom%), and the error bars represent the standard deviation for the several cells analysed by nanoSIMS.

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Figure 6. (a) Green excitation (510-560 nm) epifluorescent micrographs of UCYN-C, (b) 13 C and ¹⁵N isotopic enrichment (atom%) in individual UCYN-C cells on day 17 in M2, (c, d) nanoSIMS images showing the 13 C (c) and 15 N (d) enrichment of individual UCYN-C cells after 24 h of incubation. The white outlines show regions of interest (ROIs), which were used to estimate the ¹³C/ 12 C and 15 N/ 14 N ratios.

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Figure 7. Cartoon summary of the simplified pathways of N transfer in the first trophic level of the food web and the potential impact on the sinking POC flux at the height of the UCYN-C bloom in the VAHINE mesocosm experiment.

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1 References

- 2 Altabet, M. A.: Variations in Nitrogen Isotopic Composition between Sinking and Suspended Particles -
- Implications for Nitrogen Cycling and Particle Transformation in the Open Ocean, Deep Sea Research,
 35, 535-554, 1988.
- 5 Aminot, A. and Kerouel, R.: Dosage automatique des nutriments dans les eaux marines. Ifremer (Ed.), 6 2007.
- Azam, F. and Malfatti, F.: Microbial structuring of marine ecosystems, Nature Reviews Microbiology 5,
 782-791, 2007.
- Bandyopadhyay, A., Elvitigala, T., Welsh, E., Stöckel, J., Liberton, M., Min, H., Sherman, L. A. and
 Pakrasi, H. B.: Novel metabolic attributes of the genus cyanothece, comprising a group of unicellular
 nitrogen-fixing Cyanothece, MBio, 2, 2011.
- 12 Benavides, M., Agawin, N., Arístegui, J., Peene, J., and Stal, L.: Dissolved organic nitrogen and carbon
- release by a marine unicellular diazotrophic cyanobacterium, Aquatic microbial ecology, 69, 69-80,
 2013.
- Berman-Frank, I., Spungin, D., Rahav, E., F., V. W., Turk-Kubo, K., and Moutin, T.: Dynamics of transparent exopolymer particles (TEP) during the VAHINE mesocosm experiment in the New Caledonia lagoon, Biogeosciences Discussions, doi: doi:10.5194/bg-2015-612, 2016. 2016.
- 18 Berthelot, H., Bonnet, S., Camps, M., Grosso, O., and Moutin, T.: Assessment of the dinitrogen released 19 as ammonium and dissolved organic nitrogen by unicellular and filamentous marine diazotrophic 20 cyanobacteria grown in culture, Frontiers in Marine Science, 2, 2015a.
- Berthelot, H., Bonnet, S., Grosso, O., Cornet, V., and Barani, A.: Transfer of diazotroph derived nitrogen
 towards non-diazotrophic planktonic communities: a comparative study between *Trichodesmium erythraeum*, *Crocosphaera watsonii* and *Cyanothece* sp., Biogeosciences Discussions, doi:
- 24 doi:10.5194/bg-2015-607, 2016. 2016.
- Berthelot, H., Moutin, T., L'Helguen, S., Leblanc, K., Hélias, S., Grosso, O., Leblond, N., Charrière, B.,
 and Bonnet, S.: Dinitrogen fixation and dissolved organic nitrogen fueled primary production and
 particulate export during the VAHINE mesocosm experiment (New Caledonia lagoon), Biogeosciences,
 12, 4099-4112, 2015b.
- Biegala, I. C. and Raimbault, P.: High abundance of diazotrophic picocyanobacteria (< 3 μm) in a
 Southwest Pacific coral lagoon, Aquatic Microbial Ecology, 51, 45-53, 2008.
- Bombar, D., Taylor, C. D., Wilson, S. T., Robidart, J. C., Rabines, A., Turk-Kubo, K. A., Kemp, J. N., Karl,
 D. M., and Zehr, J. P.: Measurements of nitrogen fixation in the oligotrophic North Pacific Subtropical
 Gyre using a free-drifting submersible incubation device, Journal of Plankton Research, 37, 727–739,
- 34 **2015.**
- Bonnet, S., Berthelot, H., Turk-Kubo, K., Cornet-Bartaux, V., Fawcett, S. E., Berman-Frank, I., Barani, A.,
 Dekaezemacker, J., Benavides, M., Charriere, B., and Capone, D. G.: Diazotroph derived nitrogen
 supports diatoms growth in the South West Pacific: a quantitative study using nanoSIMS, Limnology
- 38 and Oceanography, Accepted. Accepted.
- 39 Bonnet, S., Moutin, T., Rodier, M., Grisoni, J. M., Louis, F., Folcher, E., Bourgeois, B., Boré, J. M., and
- 40 Renaud, A.: Introduction to the project VAHINE: VAriability of vertical and tropHIc transfer of
- 41 diazotroph derived N in the south wEst Pacific, Biogeosciences Discussions, doi: doi:10.5194/bg-2015-
- 42 **615, 2016. 2016.**

- 1 Bonnet, S., Rodier, M., Turk, K., K., Germineaud, C., Menkes, C., Ganachaud, A., Cravatte, S.,
- 2 Raimbault, P., Campbell, E., Quéroué, F., Sarthou, G., Desnues, A., Maes, C., and Eldin, G.: Contrasted
- 3 geographical distribution of N2 fixation rates and nifH phylotypes in the Coral and Solomon Seas
- 4 (South-Western Pacific) during austral winter conditions, Global Biogeochemical Cycles, 29, 2015.
- 5 Braman, R. S. and Hendrix, S. A.: Nanogram nitrite and nitrate determination in environmental and 6 biological materials by vanadium (III) reduction with chemiluminescence detection, Analytical 7 Chemistry, 61, 2715–2718, 1989.
- 8 Casciotti, K. L., Sigman, D. M., Hastings, M. G., Böhlke, J. K., and Hilkert, A.: Measurement of the 9 Oxygen Isotopic Composition of Nitrate in Seawater and Freshwater Using the Denitrifier Method, 10 Analytical Chemistry, 74, 4905–4912, 2002.
- 11 Church, M. J., Jenkins, B. D., Karl, D. M., and Zehr, J. P.: Vertical distributions of nitrogen-fixing 12 phylotypes at Stn ALOHA in the oligotrophic North Pacific Ocean, Aquatic Microbial Ecology, 38, 3-14,
- 13 **2005.**
- 14 Dabundo, R., Lehmann, M. F., Treibergs, L., Tobias, C. R., Altabet, M. A., Moisander, A. M., and
- 15 Granger, J.: The Contamination of Commercial ¹⁵N₂ Gas Stocks with ¹⁵N–Labeled Nitrate and 16 Ammonium and Consequences for Nitrogen Fixation Measurements, PloS one, 9, e110335, 2014.
- 17 Dore, J. E., Letelier, R. M., Church, M. J., Lukas, R., and Karl, D. M.: Summer phytoplankton blooms in
- the oligotrophic North Pacific Subtropical Gyre: Historical perspective and recent observations,
 Progress in Oceanography, 76, 2-38, 2008.
- 20 Duhamel, S., Björkman, K. M., and Karl, D. M.: Light dependence of phosphorus uptake by 21 microorganisms in the North and South Pacific subtropical gyres, Aquatic Microbial Ecology, doi: doi: 22 10.3354/ame01593, 2012. 2012.
- Dupouy, C., Neveux, J., Subramaniam, A., Mulholland, M. R., Montoya, J. P., Campbell, L., Carpenter, E.
 J., and Capone, D. G.: Satellite captures trichodesmium blooms in the southwestern tropical Pacific,
 EOS, 81, 13-16, 2000.
- 26 Dyhrman, S. T., Chappell, D., Haley, S. T., Moffett, J. W., Orchard, E. D., Waterbury, J. B., and Webb, E.
- A.: Phosphonate utilization by the globally important marine diazotroph *Trichodesmium*, Nature, 439,
 68-71, 2006.
- Eppley, R. W., Reid, F. M. H., and Strickland, J. D. H.: Estimates of phytoplankton crop size, growth
 rate, and primary production, Bull. Scripps Inst. Oceanogr., 17, 33-42, 1970.
- Falkowski, P. G.: Evolution of the nitrogen cycle and its influence on the biological sequestration of CO2 in the ocean., Nature, 387, 272-275, 1997.
- Fawcett, S. E., Lomas, M. W., Casey, J. R., Ward, B. B., and Sigman, D. M.: Assimilation of upwelled nitrate by small eukaryotes in the Sargasso Sea, Nature Geoscience, 4, 717-722, 2011.
- Fong, A. A., Karl, D. M., Lukas, R., Letelier, R. M., Zehr, J. P., and Church, M. J.: Nitrogen fixation in an anticyclonic eddy in the oligotrophic North Pacific Ocean, ISME Journal, 2, 663-676, 2008.
- Foster, R. A., Kuypers, M. M. M., Vagner, T., Paerl, R. W., Musat, N., and Zehr, J. P.: Nitrogen fixation and transfer in open ocean diatom-cyanobacterial symbioses, ISME Journal, 5, 1484-1493, 2011.
- 39 Foster, R. A., Subramaniam, A., Mahaffey, C., Carpenter, E. J., Capone, D. G., and Zehr, J. P.: Influence
- 40 of the Amazon River plume on distributions of free-living and symbiotic cyanobacteria in the western
- 41 tropical north Atlantic Ocean, Limnology and Oceanography, 52, 517-532, 2007.

- 1 Foster, R. A., Sztejrenszus, S., and Kuypers, M. M. M.: Measuring carbon and N₂ fixation in field 2 populations of colonial and free living cyanobacteria using nanometer scale secondary ion mass 3
- spectrometry, Journal of Phycology, 49, 502-516, 2013.
- 4 Fu, F. X., Warner, M. E., Zhang, Y., Feng, Y., and Hutchins, D. A.: Effects of increased temperature and
- 5 CO2 on photosynthesis, growth, and elemental ratios in marine Synechococcus and Prochlorococcus
- 6 (Cyanobacteria), Journal of Phycology, 43, 485-496, 2007.
- 7 Fukuda, R., Ogawa, H., Nagata, T., and Koike, I.: Direct Determination of Carbon and Nitrogen Contents 8 of Natural Bacterial Assemblages in Marine Environments Nitrogen Contents of Natural Bacterial
- 9 Assemblages in Marine Environments, Applied and Environmental Microbiology, 64, 3352–3358, 1998.
- 10 Garcia, N., Raimbault, P., Gouze, E., and Sandroni, V.: Nitrogen fixation and primary production in 11 western Mediterranean, Comptes Rendus Biologies, 329, 742-750, 2006.
- 12 Garcia, N., Raimbault, P., and Sandroni, V.: Seasonal nitrogen fixation and primary production in the
- 13 Southwest Pacific: nanoplankton diazotrophy and transfer of nitrogen to picoplankton organisms, 14 Marine Ecology Progress Series, 343, 25-33, 2007.
- 15 Gasol, J. M., Zweifel, U. L., Peters, F., Fuhrman, J. A., and Hagstrom, A.: Significance of size and nucleic
- 16 acid content heterogeneity as measured by flow cytometry in natural planktonic bacteria, Applied and
- 17 Environmental Microbiology, 65, 4475-4483, 1999.
- 18 Gimenez, A., Baklouti, M., Bonnet, S., and Moutin, T.: Biogeochemical fluxes and fate of diazotroph 19 derived nitrogen in the food web after a phosphate enrichment: Modeling of the VAHINE mesocosms
- 20 experiment, Biogeosciences Discussions, doi: doi:10.5194/bg-2015-611, 2016. 2016.
- 21 Glibert, P. M. and Bronk, D.: Release of dissolved organic nitrogen by marine diazotrophic 22 cyanobacteria, Trichodesmium spp., Applied and Environmental Microbiology, 60, 3996-4000, 1994.
- 23 Goebel, N. L., Turk, K. A., Achilles, K. M., Paerl, R., Hewson, I., Morrison, A. E., Montoya, J. P., Edwards, 24 C. A., and Zehr, J. P.: Abundance and distribution of major groups of diazotrophic cyanobacteria and 25 their potential contribution to N_2 fixation in the tropical Atlantic Ocean, Environmental microbiology, 26 12, 3272-3789, 2010.
- 27 Grabowski, M. N. W., Church, M. J., and Karl, D. M.: Nitrogen fixation rates and controls at Stn ALOHA, 28 Aquatic Microbial Ecology, 52, 175–183, 2008.
- 29 Großkopf, T., Mohr, W., Baustian, T., Schunck, H., Gill, D., Kuypers, M. M. M., Lavik, G., Schmitz, R. A.,
- 30 Wallace, G. W. R., and LaRoche, J.: Doubling of marine dinitrogen-fixation rates based on direct 31 measurements, Nature, 488, 361-363, 2012.
- 32 Gruber, N.: The dynamics of the marine nitrogen cycle and its influence on atmospheric CO₂. In: The 33 ocean carbon cycle and climate., Follows, M. and Oguz, T. (Eds.), Kluwer Academic, Dordrecht, 2004.
- 34 Hasle, G. R.: The inverted microscope. In: Phytoplankton Manual, Sournia, A. (Ed.), UNESCO 35 Monographs on oceanographic methodology, Paris, 1978.
- 36 Höckelmann, C., Becher, P. G., von Reuss, S. H., and Jüttner, F.: Sesquiterpenes of the geosmin-
- 37 producing cyanobacterium Calothrix PCC 7507 and their toxicity to invertebrates, Z Naturforsch, 64, 38 49-55, 2009.
- 39 Hunt, B. P. V., Bonnet, S., Berthelot, H., Conroy, B. J., Foster, R., and Pagano, M.: Contribution and
- 40 pathways of diazotroph derived nitrogen to zooplankton during the VAHINE mesocosm experiment in
- 41 the oligotrophic New Caledonia lagoon, Biogeosciences Discussions, doi: doi:10.5194/bg-2015-614,
- 42 2016. 2016.

- 1 Kana, T. M., Darkangelo, C., Hunt, M. D., Oldham, J. B., Bennett, G. E., and Cornwell, J. C.: A membrane
- 2 inlet mass spectrometer for rapid high precision determination of N_2 , O_2 , and Ar in environmental
- 3 water samples, Analytical Chemistry, 66, 4166–4170, 1994.

4 Karl, D. M., Church, M. J., Dore, J. E., Letelier, R., and Mahaffey, C.: Predictable and efficient carbon

- sequestration in the North Pacific Ocean supported by symbiotic nitrogen fixation, Proceedings of the
 National Academy of Sciences, 109, 1842–1849, 2012.
- Karl, D. M., Letelier, R. M., Tupas, R., Dore, J., Christian, J., and Hebel, D. V.: The role of nitrogen
 fixation in biogeochemical cycling in the subtropical North Pacific Ocean, Nature, 388, 533-538, 1997.
- 9 Knapp, A. N., Dekaezemacker, J., Bonnet, S., Sohm, J. A., and Capone, D. G.: Sensitivity of
 10 *Trichodesmium erythraeum* and *Crocosphaera watsonii* abundance and N₂ fixation rates to varying
 11 NO₃⁻ and PO₄³⁻ concentrations in batch cultures, Aquatic Microbial Ecology, 66, 223-236, 2012.
- 12 Knapp, A. N., Fawcett, S. E., Martinez-Garcia, A., Leblond, N., Moutin, T., and Bonnet, S.: Nitrogen 13 isotopic evidence for a shift from nitrate- to diazotroph-fueled export production in VAHINE 14 mesocosm experiments, Biogeosciences Discussions, 12, 19901-19939, 2015.
- 15 Knapp, A. N., Sigman, D. M., and Lipschultz, F.: N isotopic composition of dissolved organic nitrogen
- and nitrate at the Bermuda Atlantic Time-series Study site, Global Biogeochemical Cycles, 19, 1-15, 2005.
- 18 Leblanc, K., Arístegui, J., Armand, L., Assmy, P., Beker, B., Bode, A., Breton, E., Cornet, V., Gibson, J.,
- 19 Gosselin, M.-P., Kopczynska, E., Marshall, H., Peloquin, J., Piontkovski, S., Poulton, A. J., Quéguiner, B.,
- 20 Schiebel, R., Shipe, R., Stefels, J., van Leeuwe, M. A., Varela, M., Widdicombe, C., and Yallop, M.: A
- 21 global diatom database abundance, biovolume and biomass in the world ocean, Earth System
- 22 Science Data, 4, 149–165, 2012.
- Leblanc, K., Cornet-Barthaux, V., Caffin, M., Rodier, M., Desnues, A., Berthelot, H., Turk-Kubo, K., and
 Héliou, J.: Phytoplankton community structure in the VAHINE MESOCOSM experiment, Biogeosciences
 Discussions, doi: doi:10.5194/bg-2015-605, 2016. 2016.
- Levitan, O., Rosenberg, G., Šetlík, I., Šetlíkova, E., Gtigel, J., Klepetar, J., Prášil, O., and Berman-Frank,
 I.: Elevated CO2 enhances nitrogen fixation and growth in the marine cyanobacterium Trichodesmium,
- 28 Global Change Biology, 13, 1-8, 2007.
- Luo, Y. W., Doney, S. C., Anderson, L. A., Benavides, M., Bode, A., Bonnet, S., Boström, K. H.,
- 30 Böttjer, D., Capone, D. G., Carpenter, E. J., Chen, Y. L., Church, M. J., Dore, J. E., Falcón, L. I.,
- 31 Fernández, A., Foster, R. A., Furuya, K., Gómez, F., Gundersen, K., Hynes, A. M., Karl, D. M.,
- 32 Kitajima, S., Langlois, R. J., LaRoche, J., Letelier, R. M., Maranón, E., McGillicuddy Jr, D. J.,
- 33 Moisander, P. H., Moore, C. M., Mourino-Carballido, B., Mulholland, M. R., Needoba, J. A.,
- Orcutt, K. M., Poulton, A. J., Raimbault, P., Rees, A. P., Riemann, L., Shiozaki, T., Subramaniam,
 A., Tyrrell, T., Turk-Kubo, K. A., Varela, M., Villareal, T. A., Webb, E. A., White, A. E., Wu, J., and
- 36 Zehr, J. P.: Database of diazotrophs in global ocean: abundances, biomass and nitrogen
- 37 fixation rates, Earth System Science Data 5, 47-106, 2012.
- Marie, D., Partensky, F., Vaulot, D., and Brussaard, C.: Enumeration of phytoplankton, bacteria and viruses in marine samples. In: Current Protocols in Cytometry, Robinson, J. P. (Ed.), John Wiley & Sons,
- 40 Inc., New York, 1999.
- 41 Meador, T. B., Aluwihare, L. I., and Mahaffey, C.: Isotopic heterogeneity and cycling of organic nitrogen
- 42 in the oligotrophic ocean, Limnology and Oceanography, 52, 934-947, 2007.

- 1 Mills, M. M., Ridame, C., Davey, M., La Roche, J., and Geider, J. G.: Iron and phosphorus co-limit 2 nitrogen fixation in the eastern tropical North Atlantic, Nature, 429, 292-294, 2004.
- Mohr, W., Grosskopf, T., Wallace, D. R. W., and LaRoche, J.: Methodological underestimation of oceanic nitrogen fixation rates, PloS one, 9, 1-7, 2010.
- 5 Moisander, A. M., Beinart, A., Voss, M., and Zehr, J. P.: Diversity and abundance of diazotrophs in the 6 South China Sea during intermonsoon, The ISME journal, 2, 954-967, 2008.
- 7 Moisander, P. H., Beinart, R. A., Hewson, I., White, A. E., Johnson, K. S., Carlson, C. A., Montoya, J. P.,
- 8 and Zehr, J. P.: Unicellular Cyanobacterial Distributions Broaden the Oceanic N₂ Fixation Domain,
- 9 Science, 327, 1512-1514, 2010.
- 10 Moisander, P. H., Rantajarvi, E., Huttunen, M., and Kononen, K.: Phytoplankton community in relation 11 to salinity fronts at the entrance to the Gulf of Finland, Baltic Sea, Ophelia, 46, 187-203, 1997.
- Montoya, J. P., Voss, M., Kahler, P., and Capone, D. G.: A simple, high-precision, high-sensitivity tracer
 assay for N₂ fixation, Applied and Environmental Microbiology, 62, 986-993, 1996.
- 14 Moore, C. M., Mills, M. M. M., Arrigo, K. R., Berman-Frank, I., Bopp, L., Boyd, P. W., Galbraith, E. D.,
- 15 Geider, R. J., Guieu, C., Jaccard, S. L., Jickells, T. D., La Roche, J., Lenton, T. M., Mahowald, N. M.,
- 16 Maranon, E., Marinov, I., Moore, J. K., Nakatsuka, T., Oschlies, A., Saito, M. A., Thingstad, T. F., A., T.,
- 17 and O., U.: Processes and patterns of oceanic nutrient limitation, Nature Geoscience, 6, 701–710, 2013.
- 18 Moutin, T., Thingstad, T. F., Van Wambeke, F., Marie, D., Slawyk, G., Raimbault, P., and Claustre, H.:
- 19 Does competition for nanomolar phosphate supply explain the predominance of the cyanobacterium
- 20 Synechococcus?, Limnology and Oceanography, 47, 1562-1567, 2002.
- 21 Moutin, T., Van Den Broeck, N., Beker, B., Dupouy, C., Rimmelin, P., and LeBouteiller, A.: Phosphate
- 22 availability controls Trichodesmium spp. biomass in the SW Pacific ocean, Marine Ecology-Progress
- 23 Series, 297, 15-21, 2005.
- Mulholland, M., Bronk, D. A., and Capone, D. G.: Dinitrogen fixation and release of ammonium and dissolved organic nitrogen by *Trichodesmium* IMS101, Aquatic Microbial Ecology, **37**, **85**-94, 2004.
- Mulholland, M. R., Bernhardt, P. W., Heil, C. A., Bronk, D. A., and O'Neil, J. M.: Nitrogen fixation and regeneration in the Gulf of Mexico, Limnology and Oceanography, 51, 176-177, 2006.
- Nelson, D. M., Treguer, P., Brezezinski, M. A., Leynaert, A., and Queguiner, B.: Production and dissolution of biogenic silica in the ocean: Revised global estimates, comparison with regional data and
- 30 relationship to biogenic sedimentation, Global Biogeochemical Cycles, 9, 359-372, 1995.
- 31 O'Neil, J. and Roman, M. R.: Grazers and Associated Organisms of *Trichodesmium*. In: Marine Pelagic
- 32 Cyanobacteria: *Trichodesmium* and other Diazotrophs, Carpenter, E. J., Capone, D.G., and Rueter, J.G.
- 33 (Ed.), NATO ASI Series, Springer Netherlands, 1992.
- Orcutt, K. M., Lipschultz, F., Gundersen, K., Arimoto, R., Michaels, A. F., Knap, A. H., and Gallon, J. R.: A
 seasonal study of the significance of N2 fixation by Trichodesmium spp. at the Bermuda Atlantic Time series Study (BATS) site, Deep Sea Research Part I, 48, 1583–1608, 2001.
- Polerecky, L., Adam, B., Milucka, J., Musat, N., Vagner, T., and Kuypers, M. M. M.: Look@NanoSIMS a
- 38 tool for the analysis of nanoSIMS data in environmental microbiology., Environmantal Microbiology, 4,
- 39 **1009-1023, 2012**.

- 1 Pujo-Pay, M. and Raimbault, P.: Improvement of the wet-oxidation procedure for simultaneous 2 determination of particulate organic nitrogen and phosphorus collected on filters, Marine and
- 3 Ecological Progress Series, 105, 203-207, 1994.

4 Qi, H., Coplen, T. B., Geilmann, H., Brand, W. A., and Böhlke, J. K.: Two new organic reference

- 5 materials for d13C and d15N measurements and a new value for the d13C of NBS 22 oil., Rapid 6 Communications in Mass Spectrometry, 17, 2483–2487, 2003.
- Rahav, E., Herut, B., Levi, A., Mulholland, M. R., and Berman-Frank, I.: Springtime contribution of
 dinitrogen fixation to primary production across the Mediterranean Sea, Ocean Sci, 9, 489-498, 2013.
- 9 Raveh, O., David, N., Rilov, G., and Rahav, E.: The temporal dynamics of coastal phytoplankton and 10 bacterioplankton in the eastern Mediterranean Sea, , PloS one, In press, 2015.
- 11 Redfield, A. C.: On the proportions of organic derivations in sea water and their relation to the 12 composition of plankton. In: James Johnstone Memorial Volume, R.J., D. (Ed.), University Press of 13 Liverpool, 1934.
- Rodier, M. and Le Borgne, R.: Population and trophic dynamics of Trichodesmium thiebautii in the SE
 lagoon of New Caledonia. Comparison with T. erythraeum in the SW lagoon, Marine Pollution Bulletin,
 61, 349-359, 2010.
- 17 Rodier, M. and Le Borgne, R.: Population dynamics and environmental conditions affecting 18 Trichodesmium spp. (filamentous cyanobacteria) blooms in the south-west lagoon of New Caledonia, 10 Iournal of Experimental Maxima Biology and Factory 258, 20, 22, 2008
- **Journal of Experimental Marine Biology and Ecology, 358, 20-32, 2008.**
- Scharek, R., Latasa, M., Karl, D. M., and Bidigare, R. R.: Temporal variations in diatom abundance and
 downward vertical fux in the oligotrophic North Pacific gyre, Deep Sea Research Part I, 46, 1051-1075,
 1999a.
- Sharek, R. M., Tupas, L. M., and Karl, D. M.: Diatom fluxes to the deep sea in the oligotrophic North
 Pacific gyre at Station ALOHA, Marine and Ecological Progress Series, 82, 55-67, 1999b.
- 25 Shiozaki, T., Nagata, T., Ijichi, M., and Furuya, K.: Nitrogen fixation and the diazotroph community in 26 the temperate coastal region of the northwestern North Pacific, Biogeosciences, 12, 4751–4764, 2015.
- Sigman, D. M., Casciotti, K. L., Andreani, M., Barford, C., Galanter, M., and Böhlke, J. K.: A bacterial
 method for the nitrogen isotopic analysis of nitrate in seawater and freshwater, Analytical Chemistry,
 73, 4145–4153, 2001.
- Smayda, T. J.: What to count. In Phytoplankton Manual, ed. . In: Monographs on oceanographic methodology 6, Sournia, A. (Ed.), UNESCO, Paris, 1978.
- Sohm, J. A. and Capone, D. G.: Phosphorus dynamics of the tropical and subtropical north Atlantic:
 Trichodesmium spp. versus bulk plankton, Marine and Ecological Progress Series, 317, 21-28, 2006.
- Staal, M., Meysman, F. J., and Stal, L. J.: Temperature excludes N2-fixing heterocystous cyanobacteria in the tropical oceans., Nature, 425, 504-507, 2003.
- Subramaniam, A., Yager, P. L., Carpenter, E. J., and al., e.: Amazon River enhances diazotrophy and carbon sequestration in the tropical North Atlantic Ocean., Proceedings of the National Academy of Sciences, 105, 10460–10465, 2008.
- Sun, J. and Liu, D.: Geometric models for calculating cell biovolume and surface area for phytoplankton, Journal of Plankton Research, 25, 1331-1346, 2003.

- 1 Thompson, A., Carter, B. J., Turk-Kubo, K., Malfatti, F., Azam, F., and Zehr, J. P.: Genetic diversity of the
- 2 unicellular nitrogen-fixing cyanobacteria UCYN-A and its prymnesiophyte host, Environmental
- 3 microbiology, 16, 3238–3249, 2014.
- Tripp, H. J., Bench, S. R., Turk, K. A., Foster, R. A., Desany, B. A., Niazi, F., Affourtit, J. P., and Zehr, J. P.:
 Metabolic streamlining in an open-ocean nitrogen-fixing cyanobacterium, Nature, 464, 90-94, 2010.
- Turk-Kubo, K. A., Frank, I. E., Hogan, M. E., Desnues, A., Bonnet, S., and Zehr, J. P.: Diazotroph
 community succession during the VAHINE mesocosms experiment (New Caledonia Lagoon),
 Biogeosciences, 12, 7435-7452, 2015.
- Van Wambeke, F., Pfreundt, U., Barani, A., Berthelot, H., Moutin, T., Rodier, M., Hess, W., and Bonnet,
 S.: Heterotrophic bacterial production and metabolic balance during the VAHINE mesocosm
 experiment in the New Caledonia lagoon Biogeosciences Discussions, 12, 19861-19900, 2015.
- Verity, P. G., Robertson, C. Y., Tronzo, C. R., Andrews, M. G., Nelson, J. R., and Sieracki, M. E.:
 Relationships between cell volume and the carbon and nitrogen content of marine photosynthetic
 nanoplankton, Limnology and Oceanography, 37, 1434–1446, 1992.
- Vincent, D., Slawyk, G., L'Helguen, S., Sarthou, G., Gallinari, M., Seuront, L., Sautour, B., and Ragueneau, O.: Net and gross incorporation of nitrogen by marine copepods fed on 15N-labelled diatoms: Methodology and trophic studies, Journal of experimental marine biology and ecology, 352, 295-305, 2007.
- 19 Walsby, A. E.: The gas vesicles and buoyancy of *Trichodesmium*, Marine Pelagic Cyanobacteria: 20 Trichodesmium and other Diazotrophs, 1992. 141-161, 1992.
- 21 White, A. E., Foster, R. A., Benitez-Nelson, C. R., Masqué, P., Verdeny, E., Popp, B. N., Arthur, K. E., and
- 22 Prahl, F. G.: Nitrogen fixation in the Gulf of California and the Eastern Tropical North Pacific, Progess in

23 Oceanography, 109, 1-17, 2012.

- Wilson, S. T., Böttjer, D., Church, M. J., and Karl, D. M.: Comparative assessment of nitrogen fixation
 methodologies conducted in the oligotrophic Noth Pacific Ocean, Applied and Environmental
 Microbiology, 78, 6516–6523, 2012.
- Yentsch, C. S. and Phinney, D. A.: Spectral fluorescence: An ataxonomic tool for studying the structure
 of phytoplankton populations, Journal of Plankton Research, 7, 617–632, 1985.
- Zehr, J. P., Bench, S. R., Carter, B. J., Hewson, I., Niazi, F., Shi, T., Tripp, H. J., and Affourtit, J. P.:
 Globally Distributed Uncultivated Oceanic N₂-Fixing Cyanobacteria Lack Oxygenic Photosystem II,
 Science, 322, 1110-1112, 2008.
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Table 1. N₂ fixation rates (nmol N L⁻¹ d⁻¹) measured in the mesocosms and in lagoon waters.
Table shows the range, median, mean, contribution of the <10 μm fraction to total rates (%), and
the number of samples analysed (n).

	Range	Median	Mean	% <10 µm	n
M1	0.5-69.7	15.9	19.7	38	61
M2	3.0-67.7	15.1	18.1	43	57
M3	2.9-60.4	14.2	17.7	29	59
Average mesocosms	2.1-65.9	15	18.5	37	177
Lagoon waters	1.9-29.3	8.7	9.2	n.a	61



Table 2. Spearman correlation matrix of N₂ fixation rates and hydrological parameters, biogeochemical stocks and fluxes, and planktonic communities (n=66). The significant correlations (p<0.05) are indicated in bold. n.a- not available.

	Parameter	M1	M2	M3	Lagoon waters
Hydrological	Temperature	0.394	0.319	0.347	0.228
parameters	Salinity	0.211	0.213	0.266	-0.122
Biogeochemical stocks and fluxes	NO ₃	-0.539	-0.302	-0.341	0.145
	$\mathrm{NH_4}^+$	0.152	0.103	0.006	0.197
	DIP	-0.613	-0.569	-0.482	-0.116
	DON	-0.329	-0.413	-0.235	-0.180
	DOP	-0.563	-0.157	-0.316	-0.243
	PON	0.575	0.293	0.494	0.077
	POP	0.514	0.001	0.439	0.036
	POC	0.399	0.352	0.356	-0.061
	Chl a	0.660	0.656	0.656	0.220
	Primary production	0.443	0.498	0.445	0.268
	Bacterial production	0.708	0.408	0.471	0.189
	T-DIP	-0.670	-0.603	-0.564	-0.190
	APA	0.575	0.568	0.273	-0.062
Planktonic communities	HNA	0.317	-0.043	0.458	n.a
	LNA	0.262	-0.021	0.000	n.a
	Prochlorococcus	0.429	-0.122	0.138	n.a
	Synechococcus	0.699	0.434	0.499	n.a
	Pico-eukaryotes	0.614	0.563	0.414	n.a
	Nanoeukaryotes	0.477	0.002	0.442	n.a
	Diatoms	-0.099	0.456	-0.200	n.a
	Dinoflagellates	0.242	-0.392	-0.321	n.a
	UCYN-A1	0.545	-0.521	-0.503	0.200
	UCYN-A2	0.127	-0.631	0.248	0.333
	UCYN-B	0.083	0.696	0.467	0.101
	UCYN-C	0.373	0.621	0.515	-0.167
	Trichodesmium	-0.145	0.147	0.285	-0.117
	DDAs	-0.036	-0.264	-0.527	0.262
	γ-24774A11	0.327	0.497	-0.750	0.733

Table 3. Average NO_3^- , DIP, DON, and DOP concentrations (µmol L⁻¹) measured over the P0, P1 and P2 periods. NO_3^- and DIP concentrations were determined using a segmented flow analyzer according to (Aminot and Kerouel, 2007). The detection limit was 0.01 and 0.005 µmol L⁻¹ for NO_3^- and DIP, respectively. DON and DOP concentrations were determined according to the wet oxidation procedure described in Pujo-Pay and Raimbault (1994) and Berthelot et al. (2015b).

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	Average P0	Average P1	Average P2
NO_3^-	0.04 ± 0.02	0.03±0.01	0.02±0.01
DIP	0.03±0.01	0.48±0.20	0.08 ± 0.05
DON	5.19±0.37	5.22±0.54	4.73±0.49
DOP	0.14±0.01	0.16±0.03	0.12±0.02

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