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Growth and production of the copepod community in the southern area of the Humboldt

Current System

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19 **Abstract**

20

21 Zooplankton production is a critical issue for understanding marine ecosystem ~~dynamics, structure~~
22 ~~and dynamics, however, its time-space variations are mostly unknown in most systems.~~ In this
23 study, estimates of copepod growth and production (CP) in the coastal upwelling and coastal
24 transition zones off central-southern Chile (~35-37° S) were obtained ~~from annual cycles~~ during a
25 3-year time series (2004, 2005, and 2006) at a fixed shelf station, and from spring-summer surveys
26 during the same years. To estimate copepod production (CP), we used published and unpublished
27 C-specific growth rates (g) of dominant copepod species from previous studies in the same
28 upwelling zone, and copepod biomasses assessed in our study. A preliminary analysis of compiled
29 data of copepod growth (g) showed that varied extensively among species and under variable
30 environmental conditions; however, g -values were not correlated to either near-surface
31 temperature or copepod size. Therefore, we used mean g 's of species and their corresponding
32 biomasses to assess species production and added CP. Copepod biomass (CB) and CP were higher
33 within the coastal upwelling zone (<50 km) and both decreased substantially from 2004 to 2006.
34 Integrated aAnnual CP ranged between 24 and 52 g C m⁻² year⁻¹ with a mean annual P/B ratio of
35 7.32-7. We estimated that CP could consume up to 60% of the annual primary production (PP) in
36 the upwelling zone but most of the time was around 8%. Interannual changes in CB and CP
37 values were associated with changes in the copepod community structure. Dominance of large-
38 sized forms shifted to small-sized species, the dominance of large-sized forms replaced by small-
39 sized species from 2004 to 2006. This change was also accompanied by more persistent ~~and time~~
40 ~~extended~~-upwelling during the same season of each yearal period. Extended upwelling may have
41 caused large losses of CB from the upwelling zone due to ~~an~~-increased ~~in~~-offshore advection of
42 coastal plankton. On a larger scale, these results suggest that climate-related impacts of

43 | incrementedasing wind-driven upwelling in ~~coastal~~-upwelling systems may causegenerate a
44 | negative trend in zooplankton biomass.

45

46 | **Key Words:** Copepods, Growth Rate, Humboldt Current, Biomass, Secondary Production

47 |

48 1. INTRODUCTION

49

50 Variability in biological production of lower trophic levels is a critical issue for understanding the
51 dynamics of marine ecosystems of the world ocean (Mann and Lazier, 1991). In this context,
52 zooplankton is a key component considering their ecological role in capturing, retaining and
53 transferring freshly produced phytoplankton-carbon toward higher levels (Poulet et al., 1995;
54 Kimmerer et al., 2007). Despite this wide recognition, there are not many studies targeting
55 zooplankton secondary production and its time-space variability in the ocean, making difficult to
56 assess the actual role of zooplankton in controlling or limiting biological production of high trophic
57 levels, including fish, mammals and seabird populations (Aebischer et al., 1990; Beaugrand et al.,
58 2003, Castonguay et al., 2008). Zooplankton secondary production is the total biomass produced
59 by a population or community per unit of area or volume over a unit of time (Kimmerer et al.,
60 2007), regardless the fate of such biomass (Winberg, 1971). There is, however, not a single and
61 simple method to estimate zooplankton production and growth rates, but several approaches
62 have been applied and different results obtained. For instance, daily estimates of production can
63 vary substantially when comparing mathematical models with enzymatic approaches (Avila et al.,
64 2012), or significant deviations can occur when applying artificial cohorts methods, as compared
65 with metabolic approaches; (Lin et al., 2013); even more, most of the traditionally applied
66 methods are logistically difficult to apply as to characterize time-space variations in these rates
67 (Sastri et al., 2013; ~~Mitra et al., 2014~~).

68 In the case of copepods, the dominant components of zooplankton biomass in the oceans, there
69 have been several attempts to develop theoretical and empirical relationships between
70 zooplankton production and the factors known to affect their growth. For instance, temperature
71 has been widely reported as a fundamental factor influencing copepod growth (Huntley and

72 | [Leópez, 1992; McLaren, 1995; Escribano et al., 2014](#)), while body size should also be considered as
73 | a fundamental driver on the basis that growth, as any other physiological rate, must be modulated
74 | by allometric effects (West et al., 1997). In fact, both variables have motivated the development of
75 | the metabolic theory of ecology (Brown et al., 2004) which proposes that animal growth is
76 | predictable from body size and environmental temperature. Meantime, other studies have
77 | provided evidence that food resources can often limit zooplankton growth (Hirst and Lampitt,
78 | 1998; Vargas et al., 2010). Which one of these factors is more important, or upon which
79 | conditions any of these effects are noticeable and can thus be used as a growth predictor are still
80 | open questions in pelagic ecology.

81 | ~~In order to obtain realistic estimates of zooplankton growth and production, specific approaches~~
82 | ~~should be adopted for any particular condition or community but~~ One of the critical problems for
83 | the calculation of secondary production is having reliable estimates of in situ growth rates of the
84 | species comprising the bulk of the zooplankton biomass in a given region or area. As mentioned
85 | above, weight or C-specific growth rate (g) has been related to temperature, food conditions, and
86 | body size, but in most cases direct estimates of g show no relation or very weak relationships with
87 | these factors (e.g. Lonsdale and Levinton, 1985; Chisholm and Roff, 1990; Hutchings et al., 1995).
88 | These relationships may also depend on the taxonomic group being considered (Hirst and Bunker,
89 | 2003).

90 | Zooplankton production has received very little attention in the eastern south Pacific. In this
91 | region, tThe Humboldt Current System (HCS) is one of the Eastern Boundary Currents (EBC's)
92 | known by its high biological productivity (Mann and Lazier, 1991), attributed usually to the high
93 | levels of primary production in the coastal zone ($>10 \text{ g C m}^{-2}\text{d}^{-1}$) sustained by wind-driven
94 | upwelling (Daneri et al., 2000; Montero et al., 2007). Copepods and euphausiids dominate the
95 | zooplankton biomass in the HCS off Chile (Escribano et al., 2007; Riquelme-Bugueño et al., 2012),

96 however, very few studies on zooplankton production are available. Escribano and McLaren (1999)
97 estimated secondary production for the dominant copepod *Calanus chilensis* in the upwelling
98 region off northern Chile, and Vargas et al. (2010) estimated growth and production of three
99 copepod species in the upwelling region off central-southern Chile during an annual cycle.
100 Riquelme-Bugueño et al. (2013) described the population dynamics and biomass production of the
101 Humboldt Current "Krill", *Euphausia mucronata*, for the same region. Although euphausiids may
102 occasionally become very abundant in this region, the bulk of zooplankton biomass in the coastal
103 upwelling zone is dominated by copepods, and, more specifically, by small-sized (<2 mm)
104 copepods (Escribano et al., 2007). Hence, the latter may well reflect the dynamics of the whole
105 zooplankton biomass and production in the southern area of the Humboldt Current (Hugget et al.,
106 2009). A group of about 10 copepod species comprises >90% of the total numerical abundance
107 (Escribano et al., 2007; Escribano et al., 2012), including the small calanoid *Paracalanus cf.*
108 *indicus*, which exhibits continuous reproduction throughout the year in the regions off northern
109 and central-southern Chile, apparently with >20 generations a year (Escribano et al., 2014). Also
110 included is the small calanoid *Acartia tonsa*, with multiple generations per year (Vargas et al.,
111 2010). The cyclopoids *Oithona similis* and *O. nana*, and the poecilostomadoids *Triconia conifera*,
112 *T. media*, and *Corycaeus typicus* are also abundant (Hidalgo et al., 2010). Larger-sized (>2 mm)
113 copepods are mainly represented by *Calanus chilensis* in the northern region and *Calanoides*
114 *patagoniensis* in the central-southern region (Hidalgo et al., 2010), and, occasionally, by
115 *Rhyncalanus nasutus* and *Eucalanaus* spp, including *E. inermis* and *E. glacialis* (Castro et al., 1993;
116 Hidalgo et al., 2010).

117 In this work, we first compiled published and unpublished (previous studies) data on assessed
118 growth rates of the dominant copepod species found in the coastal upwelling zone off central-
119 southern Chile, and assessed their relationships with temperature, under which their growth was

120 ~~measured, and with copepod body size. during the spring-summer period and under time-space~~
121 ~~variations in environmental conditions, including temperature and food resources, which allowed~~
122 ~~us to test the influence of copepod size and temperature on the C-specific growth rate (g).~~
123 Secondly, we used ~~these~~ species-dependent ~~g values~~ and copepod biomass to calculate copepod
124 ~~biomass~~ production and ~~its its~~ time-space variability in the domain of the coastal upwelling and
125 coastal transition zones, ~~thereby~~ contributing to provide the first estimates of copepod community
126 production in the Humboldt Current, ~~and to understanding the factors causing time-space~~
127 variability in copepod growth and production in this upwelling region.

128

129 2. METHODS

130

131 2.1. Field studies

132 Copepod abundance estimates were monthly obtained from a 3-years time series (January 2004
133 to 2005, and December 2006) at a fixed shelf station (Station 18, ~36.5° S, ~30 km from the coast)
134 off Concepción. During the spring-summer of the same years, spatial surveys were conducted to
135 sample including monthly samplings, and from spring-summer surveys during the same years
136 (Table 1), with the areastations located between the coastal and the coastal transition zones (up
137 to 180 km from the coast) off central-southern Chile (35°-39° S; Fig. 1). The spatial cruise in 2004
138 was carried out during 14-21 January completing a grid of 29 stations. In 2005 the cruise was
139 performed during 7-15 December with a grid of 17 stations. Finally, the 2006 cruise took place
140 during 10-25 November with a grid of 22 stations. Fig. 1 illustrates position of station 18 and
141 sampling stations for the spatial cruises.

142 At Station 18 and during the spatial surveys, zooplankton sampling was performed using a ←
143 1 m² Tucker Trawl net equipped with 200 µm mesh size nets and a calibrated flowmeter. At the

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144 time series, the net was trawled from 80 m deep to surface, providing integrated samples; details
145 on sampling procedures are described in Escribano et al. (2007). Samples were immediately split
146 onboard with a Motoda splitter and a fraction (usually ¼) was frozen at -20° C for zooplankton
147 biomass, and the rest of the sample was fixed with formalin 10%. Space variations in copepod
148 abundance were assessed through spring–summer, oceanographic cruises covering the upwelling
149 and the coastal transition zones (up to 180 km from the coast). The sampling grid varied slightly
150 from year to year (Fig. 1 and Table 1). During these spatial surveys, similar procedures as those of
151 the time series study were applied to obtain hydrographic data and zooplankton samples. For
152 zooplankton sampling, however, the Tucker Trawl equipment was deployed down to 200 m depth,
153 or near bottom in shallower stations. CTDO casts and bottle samples were obtained at each
154 station, together with chlorophyll–a (Chl-a) estimations in at least 7-9 depths in the upper 100 m
155 layerdepth at each station. Continuous profiles of temperature were used to assess depth of the
156 thermocline and derive mean temperature of the mixed layer, but thereafter temperature at 10 m
157 depth was used to relate with CB and CP, as suggested in previous studies (Escribano, 1998),
158 whereas Chla at 5 m depth was related to CB and CP, because Chla at this depth represents
159 approximately the phytoplankton maximum at Station 18 (Morales and Anabalón, 2012).
160 To assess upwelling conditions during the time series, ~~in addition,~~ wind data were obtained from a
161 meteorological station (shown in Fig. 1) since August 2004; speed and direction were measured
162 every 5 min. and vector averaged for every hour as to assess wind forcing and estimate an
163 upwelling index:

164

$$165 \tau = \rho cd V|V| \quad (1)$$

166

167 where τ = wind stress (kg s m^{-2}), ρ = air density assumed as 1.2 kg m^{-3} , c_d is an empirical constant=
168 0.0013, and V is the alongshore component of the wind in m s^{-1} .

169 ~~Station 18 represents an oceanographic time series study launched by COPAS Center (Escribano~~
170 ~~and Schneider 2007) and it includes rosette CTDO deployments and plankton monitoring.~~
171 ~~Zooplankton sampling was performed on a monthly basis using a 1 m^2 Tucker Trawl net equipped~~
172 ~~with $200 \mu\text{m}$ mesh size nets and a calibrated flow meter; deployments from 80 m depth to surface~~
173 ~~provided integrated samples; details on sampling procedures are provided in Escribano et al.~~
174 ~~(2007).~~

175 ~~Space variations in copepod abundance were assessed through spring-summer, oceanographic~~
176 ~~cruises covering the upwelling and the coastal transition zones (up to 180 km from the coast). The~~
177 ~~sampling grid varied slightly from year to year (Fig. 1 and Table 1). During these spatial surveys,~~
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179 ~~zooplankton samples. For zooplankton sampling, however, the Tucker Trawl equipment was~~
180 ~~deployed down to 200 m depth, or near bottom in shallower stations. CTDO casts and bottle~~
181 ~~samples were obtained at each station, together with chlorophyll a (Chla) estimations in at least~~
182 ~~7-9 depths in the upper 100 m depth at each station.~~

183

184 2.2. Copepod biomass and growth

185 Copepod biomass is needed to calculate secondary production. Biomass ~~estimates~~ for each
186 species ~~wase~~ obtained from mean weight estimates and length-weight relationships, available
187 from the literature, for copepodid stages of same species (REFS). For this purpose, we first
188 calculated the mean body size of all copepodid stages for each species (Table 12), and then we
189 applied a length-weight regression to estimate mean weight (as dry weight). Length-weight

190 regressions were obtained from literature (Chisholm and Roff, 1990; Hofcroft et al., 2002) and
191 species biomasses were estimated as:

192

$$193 \quad B_i = \sum_{i=1}^N (w_i n_i) 0.4 \quad (2)$$

194

195 where B_i is the species- i biomass ($\mu\text{g C m}^{-3}$), W_i and n_i are the mean dry weight (μg) and
196 abundance (number m^{-3}) of the i -species and 0.4 is the conversion factor to $\mu\text{g C}$ from dry weight
197 (Escribano et al., 2007).

198 Several studies carried out in the last few years in the upwelling zone off Chile have provided
199 estimates of in situ growth rates (g) of copepods for different copepodid stages and species (Table
200 ~~12~~). Most of these studies have applied the molting rate method, by using artificial cohorts
201 (review in Harris et al., 2000). We made use of this set of estimates to examine the influence of
202 temperature and copepod size on growth and, from that, we attempted to develop an empirical
203 equation to predict in situ g 's from these variables for each of the dominant species in the
204 samples.

205

206 2.3. Data analyses

207 Copepod production for each species was estimated from their biomass and g values as to obtain
208 total production, such that:

209

$$210 \quad CP = \sum_{i=1}^N (B_i g_i) \quad (3)$$

211

212 where CP = total copepod production ($\text{mg C m}^{-3} \text{d}^{-1}$), B_i = as defined above and g_i = C-specific growth
213 rate (d^{-1}) for each i -species.

214 CP was calculated for each sampling station during the spatial cruises and each of the monthly
215 samplings during the time series. CP integrated values in the water column (~~trapezoidal method~~)
216 were obtained by multiplying by the sampling depth and, in the case of the time series, an
217 integrated annual value was estimated by applying the trapezoidal method over the time axis. An
218 estimate of the annual production/biomass ratio (P/B) was also obtained. For this, annual
219 integrated CP and monthly mean biomass were used. Oceanographic data were all processed to
220 construct spatial contours (cruises) and temporal contours (time series) and the ranges of
221 temporal variation of temperature and Chla were derived for the mixed layer. Similar procedures
222 were applied to copepod abundance, biomass and production as to identify space-time patterns.
223 Relationships between g and temperature and body size were tested by linear and non-linear
224 regression methods, and goodness-of-fit was tested by correlation, the determination index and
225 ANOVA. Meantime, eventual associations among copepod variables and oceanographic factors
226 were assessed by General Linear Models (GLM) and Stepwise Multiple Regression (1-step) applied
227 on log-transformed data on copepod abundance, biomass and CP. GLM can test linear effects of
228 multiple continuous variables (e.g. environmental) on single dependent responses, such as CB or
229 CP. Influences of single or multiple variables are then tested with F-statistics. Meantime, Stepwise
230 Multiple Regression can selectively remove non-significant variables and find the best fitted model
231 to account for variation of the response variable (CB and CP in this case).

232

233 3. RESULTS

234

235 3.1. Oceanographic conditions

236 The three spatial surveys were carried out during the period of coastal wind-driven upwelling, as
237 evidenced by the surface distributions of temperature and salinity (Fig. 2). During the 2004 cruise,

238 recently upwelled waters ($<13^{\circ}\text{C}$ and $\text{salinity} >34\text{-psu}$) were found in the northern and central
239 areas in the coastal band; the offshore extension of these waters in the central area indicated
240 that there may have been one or more (sub) mesoscale eddies located in that area. Coastal
241 upwelling activity was also observed during the 2005 cruise but colder waters ($<12^{\circ}\text{C}$) with higher
242 salinities were restricted to the nearshore, except for lower salinity water (~ 33) off and within the
243 Arauco Gulf ($\sim 37^{\circ}\text{S}$). In the 2006 cruise, upwelling was concentrated in the northern area and
244 restricted to a narrow coastal band (<40 km from shore) so that, in general, waters were less
245 saline compared to 2004 and 2005. Upwelling conditions during the surveys were also evident
246 from the surface distributions of dissolved oxygen (DO) and Chl_a concentration (Fig. 3).
247 Remarkable differences in DO distribution among the cruises were found. In 2004, DO was within
248 saturated levels (>5 mL L⁻¹) over most of the region, except in the central area which showed lower
249 oxygenation (<4 mL L⁻¹) in a zonal band where higher salinities were detected, and an
250 oversaturated zone (>6 mL L⁻¹) in the northern part, coinciding also with higher salinities (Fig. 2),
251 together with high Chl_a levels (>810 mg m⁻³). During the 2005 cruise, oxygenation levels were
252 lower (<5 mL L⁻¹) in the entire region and more so in the coastal zone where higher salinities were
253 observed, except in area around the Arauco Gulf. Chl_a distribution was similar to that in 2004,
254 with high levels in the nearshore area in the northern and central areas. In 2006, highly
255 oxygenated conditions (>65 mL L⁻¹) prevailed in most of study region, except in the coastal band.
256 High levels of DO coincided with greater Chl_a concentrations in the entire region compared to
257 2004 and 2005 (Fig. 3, lower panel).

258 ~~During the time series study, daily integrated data of wind stress, and their monthly means, were~~
259 ~~estimated for the period from August 2004 to May 2007. Since data for the time series were~~
260 ~~available from January 2004 through December 2006, wind data were divided into two periods~~
261 ~~having the same number of days and same months as to assess whether upwelling conditions had~~

262 changed during these periods. There was a clear seasonality in upwelling favorable winds (T's with
263 positive values), such that potential for upwelling initiated in August and remained favorable until
264 April-May of each year, while downwelling conditions prevailed in the winter period between June
265 to early August (Fig. 4). In the first period of the time series (2004-2005), there were 205 days
266 favorable for upwelling (T>0) and downwelling was intense and persistent throughout early
267 winter (May 2005) to late winter (August 2005). In the second period (2006-2007), there were 255
268 days favorable for upwelling and downwelling conditions in winter were less intense and persisted
269 for a shorter time (~115 days) compared to the first period (Fig. 4).

270 At Station 18, the three annual cycles based on monthly sampling clearly revealed a seasonal
271 pattern, characterized by an upwelling period (September-March) and a non-upwelling period
272 (May-August) (Fig. 54). Mean temperature a 10 m depth was 12.6 °C for the entire time series
273 ranging between 10.6-14.5 °C. During the upwelling period, the ascent of cold waters (<12° C)
274 reacheds a shallow depth (below 20 m), together with surface warming in the upper layer during
275 the summer, whereas warmer waters (>13° C) prevailed below 30 m depth upon the non-
276 upwelling period in winter (Fig. 5A4A). Upwelling also brings to the surface layer high salinity
277 water (>34-psu) whereas a layer of freshwaters appears in winter such that higher stratification is
278 generated due to increased river runoff (Fig. 45B). During active upwelling-conditions, oxygen-
279 deficient conditions dominated in shallow subsurface waters (<20 m depth) due to the
280 shallowness of the oxygen minimum zone (OMZ) in the region. By contrast, the water column
281 becaemes oxygenated down to near bottom during non-upwelling (Fig. 5C4C). The annual bloom
282 of phytoplankton starteds in early spring (September-October), coinciding with the setup of
283 upwelling, and Chl-a remaineds high until the end of summer in most cases. Mean Chla in the
284 mixed layer (5 m depth) during the time series was 3.0 mg m⁻³ with a range of 0.01-46.3 mg m⁻³.

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285 Seasonality of upwelling conditions was also clearly reflected in the upwelling index, estimated as
286 wind stress (τ). This index was ranged between -54 and 34 ($\text{kg s}^{-1}\text{m}^{-2}$) and negative values
287 dominated during the autumn-winter, whereas in the spring-summer most of the daily integrated
288 values of τ resulted positive, indicating favorable winds for upwelling.

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290 3.2. Copepod growth rates

291 Several studies in the last few years have estimated g for dominant copepod species from the
292 Chilean upwelling zone; most of them are based on laboratory studies simulating-under a variety
293 of temperature and food conditions representing those in the field from where animals were
294 captured and, therefore, their results can be considered as~~represent~~ reliable estimates of copepod
295 growth in nature~~g~~. A summary of estimates of g for different developmental stages and species,
296 including body size of tested individuals, is provided in Table 12. Because of potential allometric
297 effects on growth rate, we attempted to develop a size-dependent model to predict g as a
298 function of body size (Fig. 56A). Although an apparent decrease in g with size is observed, no
299 significant correlation between these two variables was found ($P > 0.05$) after testing with different
300 lineal (GLM) and non-linear models.

301 Since temperature has been established as an important factor affecting growth rate of copepods
302 (Huntley and López, 1992; Gillooly et al., 2001), it has been suggested~~was thought~~ that this
303 variable could be a suitable predictor of g under variable environmental conditions (e.g. Huntley
304 and López, 1992; Lin et al., 2013). For all the available g estimates, we tested the influence of in
305 situ-simulated temperature under which g was estimated; ~~also~~, no significant effects were found
306 ($P > 0.05$) after lineal (GLM) and non-linear models (Fig. 6B5B). A combined model using both body
307 size and temperature, as that described by Lin et al. (2013), did not result in a significant
308 association~~differences~~ either. On this basis, mean values of g for some of the dominant species

309 were thus calculated and used to estimate their specific production rates (Fig. 76). In these
310 analyses we did not include some species, such as *Oithona* spp., *Oncaea* spp., *Corycaeus* spp.,
311 *Pleurommama* sp. and *Centropages* sp., for which no estimates of *g* were available. For these
312 species in which no estimates of *g* were available, a grand mean of copepod growth rate,
313 calculated from all available data (Table 1), rate was applied (mean \pm SD: $0.27 \pm 0.133 \text{ d}^{-1}$) for later
314 estimates of CP.

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316 3.3. Copepod biomass and production

317 Copepod abundance (N), biomass (CB) and production (CP) were estimated as annual means for
318 ~~both the spatial surveys and~~ the time series and mean values during the spatial surveys (Table 23).

319 In both cases, strong variability in N, CB and CP was observed (coefficient of variation:25-50%).

320 Spatial variability of N relates to a greater aggregation of copepods in the upwelling zone and
321 decreasing values towards the offshore (Fig. 78). The highest values of CB and CP were also
322 concentrated in the upwelling zone although there was a strong variation from year to year, with
323 lower values in 2004 (Fig. 98).

324 Copepod species in three size categories, in according to their total length: small (<1.5 mm),
325 medium (1.5-2.5 mm) and large (>2.5 mm), varied substantially from year to year (Table 34).

326 Small-sized species increased in abundance from 2004 to 2006, whereas large-size species tended
327 to slightly decrease in 2006 compared to 2005, although they were scarce in 2004 ~~the same years~~.

328 The distribution of these 3 size categories also varied from one year to another (Fig. 910). Medium
329 size species were absent in 2004 and large-sized species were more abundant in the upwelling
330 zone, while small-sized species became more abundant in 2005 and even more so in 2006 and
331 concentrated in the upwelling zone.

332 From the time series at Station 18, no seasonal pattern or trend in copepod abundance was
333 detected (Fig. 101A), as was the case for CB and CP (Fig. 101B); in both cases, lower values were
334 detected during 2006. Integrated annual CP at station 18 was $52.2 \text{ g C m}^{-2}\text{-y}^{-1}$, $32.8 \text{ g C m}^{-2}\text{-y}^{-1}$ and
335 $24.0 \text{ g C m}^{-2}\text{-y}^{-1}$ for 2004, 2005 and 2006, respectively. ~~From annual means of monthly integrated~~
336 ~~biomasses, the Annual P:B ratios obtained were $7.12-5$, $7.42-8$ and $7.52-9$ for 2004, 2005 and~~
337 2006, respectively. The daily P:B ratio was, on average, 0.24. The variance of CB for each year,
338 estimated from the coefficient of variation, was 16%, 27% and 24% for 2004, 2005 and 2006,
339 respectively.

340

341 3.4. Environmental effects on biomass and production

342 Using the data from the spatial surveys, a stepwise multiple regression was applied to test the
343 effect of year of sampling and oceanographic conditions on N, CB and CP. Copepod data were
344 previously log-transformed and a 1-step function was applied. Significant differences among years
345 in N and CP were found. Chl_a correlated positively with N, whereas Chl_a, DO and OMZ depth
346 correlated with CP (Table 45).

347 For the time series data, we used cross-correlations between copepod variables and
348 oceanographic conditions (including temperature, Chl_a, DO, and OMZ depth) to test for eventual
349 associations. Although all the oceanographic factors showed a seasonal pattern, characterized by
350 upwelling and downwelling periods (Fig. 54), copepod abundance, biomass and production did not
351 but their monthly fluctuations are rather random (Fig. 101). Therefore, it was not surprising that
352 no significant correlations ($P>0.05$) between N, CB and CP and derived oceanographic variables
353 were found.

354

355

356 4. DISCUSSION

357

358 The oceanographic conditions observed during this study are those expected from previous
359 studies in the upwelling zone (Strub et al., 1998; Hidalgo et al., 2012; Morales and Anabalón, 2012)
360 and the coastal transition zone (Letelier et al., 2009). The spatial surveys, conducted during spring-
361 summer conditions, show that upwelling conditions prevailed in a coastal band along the study are
362 of about 50 km width, coinciding with the isobath of 200 m (shown in Fig. 1). These conditions are
363 characterized by colder, more saline and less oxygenated water. This coastal band constitutes the
364 main habitat of a few dominant copepods species (Hidalgo et al., 2010), and as evidenced by their

365 aggregation over the shelf (Fig. 87). Dominant *Paracalanus Cf. indicus* and *Acartia tonsa* tend to
366 concentrate in the nearshore, or upwelling zone, whereas larger size calanoids, such as *Calanus*
367 *patagoniensis* and *Rhyncalanus nasutus* exhibit greater concentration offshore (Hidalgo et al.,
368 2010). However, at times during intense upwelling the whole community can be advected offshore
369 into mesoscale eddies (Morales et al., 2010). Therefore same species are, however, present in the
370 coastal transition zone although in lower abundances, probably as a result of their offshore
371 transport by mesoscale eddies (Morales et al., 2010), which are originated in the upwelling zone,
372 cannot move plankton mostly in a westward direction (Hormazabal et al., 2013). Our findings
373 however indicate that most of the copepod production (CP) takes place in the coastal upwelling
374 zone, where food resources (as represented by Chl-a) are also concentrated. Although, at times
375 during the year Chla becomes low (<1 mg m⁻³), even within the upwelling zone. In this respect,
376 Chla itself should not considered as a suitable indicator for food resources for omnivorous
377 copepods in this system, which are capable of efficiently use a heterotrophic diet when Chla is low
378 (Vargas et al., 2006).

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379 If most CP occurs in the upwelling zone, Station 18 is therefore a suitable location to assess its
380 temporal variability. Oceanographic variability there also clearly shows the upwelling signal
381 (Sobarzo et al., 2007a; Montero et al., 2007; Morales and Anabalón, 2012). At this location, the
382 copepod community has been well studied (Escribano et al., 2007; Hidalgo and Escribano, 2007),
383 and even though some seasonal signals in abundance and age-structure of some species have
384 been described (Castro et al., 1993; Hidalgo and Escribano, 2007), most populations can grow and
385 reproduce throughout the year (Vargas et al., 2010; Escribano et al., 2014). ~~suggesting that CP~~
386 ~~is a process taking place year-round.~~
387 ~~It has been suggested that food conditions (quantity and quality) may be an important factor~~
388 ~~determining copepod growth and production (Huntley and Boyd, 1984; Hirst and Lampitt, 1998;~~
389 ~~Lin et al., 2013). However, evidence from field studies has shown weak correlations between~~
390 ~~growth rate of copepods and food conditions. In this regard, Hutchings et al. (1995) found no~~
391 ~~correlation between copepod growth rate and Chl-a in the Benguelas upwelling system, while~~
392 ~~Peterson et al. (1991) suggested little effects of food availability on copepod production during~~
393 ~~summer in Skagerrak (north sea). In our study, we found a significant correlation between~~
394 ~~copepod abundance and Chl-a during the spatial surveys. This positive correlation was also found~~
395 ~~between CP and Chl-a.~~ It should be noted, however, that a significant correlation of copepod
396 abundance and CP with Chl-a does not necessarily mean that ~~copepod growth~~ is being
397 controlled by phytoplankton biomass. ~~Copepod abundance and production~~ also correlated
398 significantly with low oxygen and a shallow OMZ which, the same as for higher Chl-a levels,
399 coincide in the zone where greater CP occurs. ~~Thus,~~ the question on whether CP can be controlled
400 or limited by phytoplankton biomass cannot be answered from the spatial survey just because of
401 spatial correlation. On the other hand, no correlation between Chl-a and CP was found in the time
402 series data, and copepod abundance, CB and CP appeared uncoupled to the seasonal pattern of

403 Chl-a. Food conditions, in terms of quality, could also be considered as affecting copepod growth
404 (Peterson et al., 1991). During the same spatial surveys, Morales and Anabalón (2012) showed that
405 both nanoplankton and microplankton fractions as potential food resources for copepods were
406 abundant, diverse and widely distributed over the upwelling zone, so that food limitation for
407 copepodid growth seemed unlikely, at least over the upwelling zone. In the context of food effects
408 on CP, in this upwelling system dominant small-sized copepod can prey on both autotrophic and
409 heterotrophic components of the nano- and microplankton, being able to even shift their diet
410 depending on the available food offer during the year cycle (Vargas et al., 2006). In the upwelling
411 zone, the nanoplankton fraction remains high and diverse throughout seasons (Böttjer and
412 Morales, 2007; Anabalón et al., 2007). It has therefore been suggested that copepods in this
413 region can grow without food limitation year round (Escribano et al., 2014).

414 In our approach to estimate copepod production, the use of species-dependent growth rates may
415 be justified on the basis that g is a physiological rate controlled by two processes, development
416 rate (DR) and tissue accumulation. In fact, estimating DR is the most widely used approach to
417 assess growth rate of copepods. DR has been widely studied in copepods (McLaren and Leonard,
418 1995; Heinle, 1969) and it is considered a species-dependent attribute (Heinle, 1969; Atkinson,
419 1994). Nevertheless, within species g may strongly vary (Runge and Roff, 2000) and the use of a
420 constant g is certainly much more difficult to justify. In field and laboratory studies, g of copepods
421 has been found to vary as a function of temperature (Huntley and López, 1992; Hirst and
422 Bunker, 2003), food conditions (Checkley et al., 1980; Finlay and Roff, 2006) and body size (Hirst
423 and Shearer, 1997), although size effects may not be reflected at the intra-specific level but
424 among species (Banse, 1982; Peters, 1983). Temperature on the other hand may strongly affect g
425 by accelerating or retarding the development of copepods (McLaren, 1995). It is therefore
426 expected that g would correlate positively with temperature. From our oceanographic surveys,

427 | however, temperature seems to vary in a rather narrow range (~ ~~311.5~~ 13.5 °C) in the mixing
428 | layer of the upwelling zone, where dominant copepods aggregate and whose ~~diet~~ vertical
429 | distribution is restricted by a shallow (<50 m) OMZ during the upwelling period (Escribano et al.,
430 | 2009). Thus, because of its little variation within the upwelling zone of central/southern Chile,
431 | temperature may not be the key factor controlling *g*. Variation in temperature could also affect
432 | copepod size through its effect on development rate (Escribano et al., 2014), and thus affecting
433 | our biomass estimates. However, because of weak seasonal signal of temperature copepodid sizes
434 | do not follow a seasonal pattern in this region (Hidalgo and Escribano, 2007), and hence this
435 | potential effect remains uncertain.

436 | It is important to consider that copepod production represents only a fraction of total secondary
437 | production for the upwelling zone. Our estimate of CP does not consider molt and egg production
438 | of copepods either, but only somatic biomass production. Also, estimates of CB and CP are mostly
439 | based on late copepodid stages, since early stages (nauplii, C1-C3) were not assessed since they
440 | are difficult to identify and some of them could be undersampled by the 200 µm net. The
441 | assessment of all stages for a given species is a complex task, and previous studies have shown a
442 | numerical dominance of late stages through the year cycle, due to rapid development of young
443 | stages and their greater mortalities (Hidalgo and Escribano, 2007; Escribano et al., 2014). Thus,
444 | biomass and production of young stages have not been accounted in our estimates, although
445 | because of their small biomasses and possibly less abundance this missing fraction may not be
446 | substantial.

447 | Our estimates of copepod growth rates and production are comparable to values reported from
448 | other systems. Hutchings et al. (1995) found a CP in the range of 17 - 150 g C m⁻²y⁻¹ for the similar
449 | upwelling system of Benguelas, whereas Mackinnon and Duggan (2003) estimated a CP of 3.4 g C
450 | m⁻²y⁻¹ for subtropical water in north west Australia. Regarding the growth rate (*g*), Peterson et al.

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451 (1991) found a g of 0.1 d^{-1} for adults and about 0.27 d^{-1} for juveniles copepods, which is exactly the
452 same mean value we found for the whole copepod community and Lin et al. (2013) reported a g in
453 the wide range of $0.01 - 0.79 \text{ d}^{-1}$ for medium size copepods from the east China sea.

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454 With respect to the bulk of zooplankton, in the time series at Station 18 total zooplankton biomass
455 (TZB) was available for the same period, as published in Escribano et al. (2007). TZB shows more
456 variability than copepod biomass (CB) (Fig. 112A) and, on occasions, CB may account up to 96% of
457 TZB although on average our estimate of CB represents nearly 40% of TZB. Meantime, monthly
458 means of primary production ($\text{g C m}^{-2} \text{d}^{-1}$) in the same upwelling zone (Daneri et al., 2000) indicates
459 that copepod production could take up to 60% of the C being produced by phytoplankton (winter
460 2004) although the mean conversion of PP into CP was about 8% (Fig. 112B). This figure is in
461 according with global mean estimates of the effect of mesozooplankton on PP (as the percent PP
462 consumed per day: mode 6%, mean 23%) and decreases exponentially with increasing productivity
463 (Calbet, 2001).

464 Recently, Escribano et al. (2012) and Pino-Pinuer et al. (2014) have described a negative trend in
465 copepod biomass and abundance from the beginning of the time series at Station 18 (2002) to
466 more recent years (2010 - 2012). Both works related this decrease in copepods with a gradual
467 increment in upwelling intensity in the coastal zone off Chile upon increased southerly winds in
468 the last decade (Garreaud and Falvey, 2009). In this study, CB and CP significantly decreased from
469 2004 to 2006 at Station 18, although this trend was unclear in the spatial surveys. Copepods are
470 strongly subjected to offshore advection during upwelling (Peterson, 1998; Keister et al., 2009;
471 Morales et al., 2010). When examining the spatial patterns of oceanographic conditions, it appears
472 that in 2005 and 2006 the upwelling focuses were concentrated near the location of Station 18,
473 judging by low oxygen water in that area (Fig. 3), as compared to 2004 when the upwelling focus
474 was located farther from the nearshore. It is therefore likely that copepods populations were

475 | more subjected to offshore advection in 2005 and 2006. Most advective forces take place in the
476 | mixed layer (upper 50 m) and even though we obtained depth-integrated samples (0-200 m) most
477 | copepods, as indicated above, aggregate in this layer and thus being advected offshore.

478 | Another possibility for lower biomass and production in 2005 and 2006 at Station 18 could be
479 | explained in terms of food-limitation for copepod growth, as suggested for other systems (e.g.
480 | Hirst and Lampitt, 1998). Nevertheless, we found no significant differences in phytoplankton
481 | biomass (as food indicator) among the three years. Also, primary production (PP) estimated
482 | monthly at Station 18 during the same period, showed that the annual cycle of PP almost repeated
483 | every year (Montero et al., 2007). ~~Furthermore, off Central/southern Chile, copepods can sustain~~
484 | ~~their reproduction and growth throughout the year, despite the seasonal bloom of phytoplankton,~~
485 | ~~by switching their diet from autotrophic preys to an omnivorous diet (Vargas et al., 2006).~~
486 | ~~Therefore, it is unlikely that annual CP could be limited by food in this upwelling region.~~

487 |
488 | Since copepods mostly concentrate in the Ekman layer (<50 m) during upwelling, as constrained by
489 | a shallow OMZ (e.g. Escribano et al., 2009), more offshore advection, upon increased upwelling,
490 | can cause biomass loss from the coastal zone. In fact, from our analysis of wind data we
491 | ~~found~~ showed that favorable conditions for upwelling were more persistent (lasted longer) during
492 | the second part of the time series and hence promoting more export of CB to offshore areas. The
493 | upwelling index revealed that from year to year the number of days favorable for upwelling (-daily
494 | integrated $\tau > 0$) varied substantially during the upwelling season (September to March). The
495 | number of days having positive values of τ was greater in 2006 than in 2005 and 2004 (Fig. 12A).

496 | Active upwelling also promotes formation of mesoscale intra-thermocline eddies (Hormazabal et
497 | al., 2013) which can also enhance plankton export from the upwelling zone.

498 The effect of these physical processes would thus be reflected in lower CP at Station 18. In
499 according to Keister et al. (2009), offshore advection from the upwelling zone constitutes an
500 efficient mechanisms to transport [carbon](#) to oceanic regions. From the viewpoint of biological
501 production of the upwelling system, however, this should be considered as a C loss. Over an
502 annual basis, we estimated how much of the C produced by phytoplankton is converted into CP
503 and CB, and the annual deficit in CP that biomass loss can cause due to more advection driven by
504 increased upwelling. The impact of greatly incremented upwelling is shown in Fig. [12B3](#), which
505 illustrate how combined factors and processes, such as upwelling conditions, CP, CB and primary
506 production may have interacted during the time series as to cause a reduction in copepod
507 production in the upwelling zone. This schema illustrates observed copepod production (CP) and
508 biomass (CB) on an annual basis, as well as an estimate of expected CP (C_{Pe}) assuming a 8% of
509 conversion efficiency of PP into CP. C_{Pe} remains constant because PP was assumed as constant
510 also over an annual basis. Then the potential biomass loss due to incremented offshore advection
511 in 2005 and 2006 is illustrated (Fig. 12B). -It is important to stress that upwelling intensity may not
512 significantly change from one year to another in average, but the length and continuity of the
513 upwelling season can be the key process causing more biomass loss on an annual basis.

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695

696 **Table 1**

697 **Summary of cruises and the time series study in the coastal upwelling region of Central/southern**
698 **Chile to estimate copepod biomass and production in relation to upwelling conditions. Three**
699 **spatial cruises were conducted (FIP 2004, 2005 and 2006) and a monthly time series study at the**
700 **fixed Station 18.**

CRUISE	PERIOD	SEASONAL CONDITION	No. STATIONS
FIP 2004	14 – 21 November	spring/upwelling	29
FIP 2005	7 – 15 December	summer/upwelling	17
FIP 2006	10 – 25 November	spring/upwelling	22
Time Series	Jan 2004 to Dec 2006	all seasons	Single (fixed)

Table 12

C-specific growth rates and size ranges for different developmental stages of copepods from the coastal upwelling zone off Chile. Estimated growth rates (g) were obtained from the molting rate method applied under in situ simulated conditions of temperature. Satiating conditions of food were assumed. Developmental stages are copepodids (C1 to C5) and adult females (AD ♀). Estimates of g were obtained at the northern upwelling region of Chile (Nor) or at the southern region (Sou).

SPECIE	STAGES	SIZE RANGE	g	TEMP.	REFERENCE
		(mm)	(d ⁻¹)	(°C)	
<i>A. tonsa</i>	C1, C2, C3, C4,	0.53 – 1.33	0.248	12	Hidalgo P. (Unpublished data)
	C5, AD ♀				
<i>C. patagoniensis</i>	C1, C2, C3	0.85 – 2.15	0.294	12	Hidalgo P. (Unpublished data)
	C4, C5.				
<i>C. chilensis</i>	C1, C2, C3, C4,	0.94 – 2.45	0.214	12	Hidalgo P. (Unpublished data)
	C5, AD ♀				
	C1, C2, C3, C4,	ND – 2.33	ND	15	Escribano R. et al. (1998)
	C5, AD ♀				
	C1, C2, C3	0.61 – 1.96	ND	16.5	
C4, C5.				Escribano R. & I. McLaren (1999)	
AD ♀		2.55	ND		14.6
	AD ♀		2.35	ND	18.5

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<i>E. inermis</i>	AD ♀	4.77	0.193	16	Hidalgo et al. (2005)
<i>P. indicus</i>	C1, C2, C3, C4, C5, AD ♀	0.35 – 0.96	0.270	12	Yáñez et al. (2012)
<i>R. nasutus</i>	C1, C2, C3, C4, C5, AD ♀	1.45 – 5.88	ND	ND	Hidalgo P. (Unpublished data)

Specie	Stages	Size Range (mm)	$g \pm sSDd$ (d ⁻¹)	Temp. (°C)	Reference
<i>A. tonsa</i>	C1, C3, C5 (Nor)	0.53 – 0.87	0.22 ± 0.035	12	Hidalgo P. (Unpublished data)
	C1 to C5 (Nor) ND		0.20 ± 0.06	12	Hidalgo P. (Unpublished data)
	C1 to C5 (Sou) ND		0.25 ± 0.05	12	Hidalgo P. (Unpublished data)
	C1, C2, C4 (Sou)	0.53 – 0.73	0.093 ± 0.042	12	Vargas et al. (2010)
	C3 (Sou)	0.658	0.21 ± 0.01	13	Vargas et al. (2010)
<i>C. patagoniensis</i>	C1 to C5 (Sou)	0.85 – 2.15	0.29 ± 0.056	12	Hidalgo P. (Unpublished data)
<i>C. chilensis</i>	C1 to C5 (Nor)	0.61 – 1.96	0.27 ± 0.11	16.5	Escribano & McLaren (1999)
	C1 to C5 (Nor)	2.33	0.114	15	Escribano et al. (1998)
	C1 to C5 (Nor)	ND	0.21 ± 0.05	12	Hidalgo P. (Unpublished data)
	C1 to AD (Nor)	2.55	0.28	14.6	Ulloa et al. (2001)
	C1 to AD (Nor)	2.35	0.35	18.5	Ulloa et al. (2001)
<i>E. inermis</i>	C1 to AD (Nor)	4.77	0.193	16	Hidalgo et al. (2005)
<i>P. Cf indicus</i>	C1 to C5 (Nor)	ND	0.24 ± 0.07	12	Hidalgo P. (Unpublished data)
	C1 to C5 (Sou)	ND	0.27 ± 0.04	12	Hidalgo P. (Unpublished data)
	C2 (Sou)	0.427	0.13	14	Vargas et al. (2010)
	C2 (Sou)	0.427	0.25 ± 0.064	13	Vargas et al. (2010)
	C3 (Sou)	0.593	0.29 ± 0.072	13	Vargas et al. (2010)
	C3 (Sou)	0.593	0.35 ± 0.03	14	Vargas et al. (2010)
	C4 (Sou)	0.725	0.053 ± 0.006	12	Vargas et al. (2010)
	C1 to C5 (Nor)	0.35 – 0.84	0.23 ± 0.11	12	Yáñez et al. (2012)
	C1 to C5 (Sou)	0.35 – 0.84	0.3 ± 0.15	12	Yáñez et al. (2012)
<i>R. nasutus</i>	C1 to AD	1.45 – 5.88	ND	ND	Hidalgo P. (Unpublished data)

Con formato: Superíndice

Table 23

Estimated copepod abundance (N), copepod biomass (CB) and copepod production (CP) from the coastal upwelling zone of Central/southern Chile, based on spatial cruises (FP) and a time series study at Station 18. Mean \pm SD are shown. n represent the number of stations for each spatial/FP cruise and the number of samples months for each year, respectively.

YEAR	SPATIAL SURVEYS				TIMES SERIES			
	N (Ind/m ³)	CB (mgC m ⁻²)	CP (mgC m ⁻² d ⁻¹)	n	N (Ind/m ³)	CB (mgC m ⁻²)	CP (mgC m ⁻² d ⁻¹)	n
2004	50 \pm 76	449 \pm 65	12 \pm 17	22	435 \pm 337	711 \pm 577	164 \pm 128	9
2005	219 \pm 262	773 \pm 1412	193 \pm 351	17	199 \pm 306	281 \pm 451	65 \pm 99	10
2006	364 \pm 778	515 \pm 1138	130 \pm 288	22	215 \pm 178	279 \pm 276	67 \pm 65	12

Spatial Cruises

Year	N (Ind m ⁻³)	CB (mg C m ⁻²)	CP (mg C m ⁻² d ⁻¹)	n
2004	49.5 \pm 75.97	449 \pm 65.27	11.5 \pm 16.75	22
2005	219.0 \pm 261.50	772.6 \pm 1412.08	192.8 \pm 351.18	17
2006	364.0 \pm 778.29	515.0 \pm 1137.78	130.4 \pm 287.79	22

Time Series

<u>Year</u>	<u>N (Ind m⁻³)</u>	<u>CB (mg C m⁻²)</u>	<u>CP (mg C m⁻² d⁻¹)</u>	<u>n</u>
2004	435.1 ± 336.90	711.0 ± 577.45	164.1 ± 127.99	9
2005	199.3 ± 306.08	281.1 ± 450.60	65.2 ± 99.36	10
2006	215.3 ± 177.73	279.2 ± 276.39	67.2 ± 64.56	12

N: abundance, CB: Copepod biomass, C P: Copepod production, n: sample size

Table 34

Copepod species classified by size ranges as found during the spatial cruises in the coastal upwelling zone of Central/southern Chile. Species abundance is shown as mean \pm SD for each year.

Size	Specie	Abundance (n m ⁻³)		
		2004	2005	2006
Small [< 1.5 mm]	<i>Acartia tonsa</i>	-	72.2 \pm 5.90	60.0 \pm 190.07
	<i>Corycaeus</i> spp	-	19.0 \pm 33.36	5.3 \pm 8.85
	<i>Oithona similis</i>	34.6 \pm 49.62	23.8 \pm 53.26	48.1 \pm 60.20
	<i>Oncaea</i> spp.	-	8.6 \pm 7.93	58.5 \pm 190.62
	<i>Paracalanus</i> CF. <i>indicus</i>	33.5 \pm 37.0	65.2 \pm 60.99	180 \pm 440.31
Medium [1.5–2.5 mm]	<i>Centropages brachiatus</i>	-	4.2 \pm 4.97	3.0 \pm 3.20
	<i>Pleuromamma gracilis</i>	-	8.0 \pm 9.61	16.8 \pm 27.49
Large [> 2.5 mm]	<i>Calanoides patagoniensis</i>	-	62.2 \pm ND	62.3 \pm 71.23
	<i>Calanus chilensis</i>	1.1 \pm 2.41	90.2 \pm 191.2	29.9 \pm 76.04
	<i>Eucalanus</i> spp.	-	1.2 \pm 1.29	-
	<i>Metridia</i> spp.	-	-	5.5 \pm 13.54
	<i>Rhyncalanus nasutus</i>	2.6 \pm 4.22	15.2 \pm 32.56	5.9 \pm 10.10

Table 45

Results from a generalized linear model (GLM) to test the influence of oceanographic variability on copepod abundance (N) and copepod production (CP), estimated from spatial cruises carried out during upwelling conditions in the coastal upwelling zone off Central/southern Chile. Only significant ($P < 0.05$) or nearly significant ($0.05 < P < 0.10$) are shown. *=Significant ($P < 0.05$), **=Highly significant ($P < 0.01$).

DEPENDENT VARIABLE	SOURCE VARIATION	t-value	P
N	Year	3.077	0.003
	Chla	1.772	0.082
C-P	Year	2.845	0.006
	Chla	2.797	0.007
	DO	1.816	0.075
	OMZ	2.182	0.033

Dependent variable	Source variation	t-value	p-value
N	Year	3.077	0.003**
	Chl-a	1.772	0.082
CP	Year	2.845	0.006**
	Chl-a	2.797	0.007**
	DO	-1.816	0.075

Con formato: Fuente: (Predeterminado) +Cuerpo

OMZ -2.182 0.033*

Figure Captions

Figure 1: The coastal upwelling zone off Central/southern Chile in the eastern South Pacific showing the sampling stations for three oceanographic cruises carried out in 2004, 2005 and 2006 under upwelling conditions (spring-summer). The map also illustrate location of Station 18 were the monthly time series study was conducted during the same 3 years. Wind data to assess upwelling condition were obtained at the coastal Meteo Station.

Figure 2: Spatial distribution of nearly surface (5 m depth) temperature and salinity during the spring-summer cruises in 3 different years, illustrating upwelling conditions over the upwelling zone and the offshore area. The map shows the isotherms ($^{\circ}\text{C}$) for temperature distribution and the color scale for salinity.

Figure 3: Spatial distribution of nearly surface (5 m depth) dissolved oxygen (DO) and Chlorophyll-a concentration (Chla) during the spring-summer cruises in 3 different years, illustrating upwelling conditions over the upwelling zone and the offshore area. The map shows the isolines (mg m^{-3}) for Chla distribution and the color scale for DO.

Figure 4: Time series of wind stress at the coastal zone of Central/southern Chile as to represent variability in upwelling conditions. Data are plotted as daily integrated values and monthly means. Positive values of wind stress represent upwelling favorable conditions and negative ones downwelling conditions. The time series divided into two periods resulted in different number of days favorable for upwelling.

Figure 45: Time series of temperature (A), salinity (B), dissolved oxygen (C) and chlorophyll-a (D) at Station 18 of Central/southern Chile as observed from monthly data. Contours were constructed from CTD-O casts at 1 m resolution and Chl-a was measured at 9 depths from surface to 80 m. The salinity=34 isoline in B is noted to illustrate upwelling pulses. The red broken line in the DO contour illustrate the depth distribution of the upper limit of the oxygen minimum zone assumed as depth of $1 \text{ mL O}_2 \text{ L}^{-1}$.

Figure 56: Variability of growth rate (g) of copepods from the upwelling zone off Chile as a function of mean copepod size and temperature. g was estimated by the molting rate method under in situ temperature and copepod size represents the mean body size of all copepodid stage adult females for each species. No

significant relationship was found in either case (A or B). [Data are from published and unpublished works summarized in Table 1.](#)

Figure 67: Variability of C-specific growth of copepod species from the coastal upwelling zone off Chile estimated by the molting rate method. Mean values and standard deviation (vertical line) are shown. Data are from two upwelling regions, Northern and Central/south of Chile, [as described in Table 1-](#)

Figure 78: Spatial distribution of dominant copepods during the spring-summer cruises in 3 different years over the upwelling zone and the offshore area off Central/southern Chile. Total copepod abundance (N) was obtained from depth integrated zooplankton samples (0-200 m).

Figure 89: Spatial distribution of copepod biomass (CB) and copepod production (CP) during the spring-summer cruises in 3 different years over the upwelling zone and the offshore area off Central/southern Chile. Open circles in 4 categories ~~of size~~ show CB distribution, whereas isolines and the color scale show the distribution of CP.

Figure 910: Spatial distribution of copepods of 3 size categories: small (<1.5 mm), medium (1.5-2.5 mm) and large (>2.5 mm) during the spring-summer cruises in 3 different years over the upwelling zone and the offshore area off Central/southern Chile. Copepod abundance (N) was obtained from depth integrated zooplankton samples (0-200 m).

Figure 101: Copepod abundance (A) and copepod production ~~(CP)~~ and biomass (B) during the time series study at Station 18 from monthly zooplankton sampling (0-80 m). [Copepod production \(CP\)](#) (line-scatter) and biomass (vertical bars) were estimated from mean size and mean C-content of [copepodid stages of dominante adult](#) species, and their associated growth rates.

Figure 112: A) Time series of copepod biomass (CB) as related to total zooplankton biomass (TZB) at Station 18 off Central/southern Chile from monthly data. TZB is from Escribano et al. (2007). B) Time series of copepod production (CP) in relation to variability in primary production (From Daneri et al. 2000). Data are from monthly estimates of CP and PP.

Figure 123: [A\) Time series of wind stress at a coastal station near Station 18 off Concepción illustrating upwelling variability for near 3 years. Daily integrated wind stress is shown, including monthly means and the number of days favorable for upwelling \(wind stress>0\) for each annual upwelling period \(September to](#)

March) indicated by the shaded areas. B) Conceptual model to illustrate the interaction among factors and processes determining the variability of copepod biomass and copepod production in relation to upwelling conditions in the coastal upwelling zone of Central/southern Chile during a 3 year time series study.

CP=copepod production, CB=copepod biomass, PP=primary production, CPe= expected copepod production

assuming a daily P/B ratio=0.25. Biomass loss ~~(shaded area)~~ represents the CB being exported from the upwelling zone upon ~~a~~ more extended periods of upwelling in 2005 and 2006 as shown in B) ~~during the second part of the time series~~. This biomass loss due to increased offshore advection explains the reduction in CP as compared to CPe.

Fig. 1

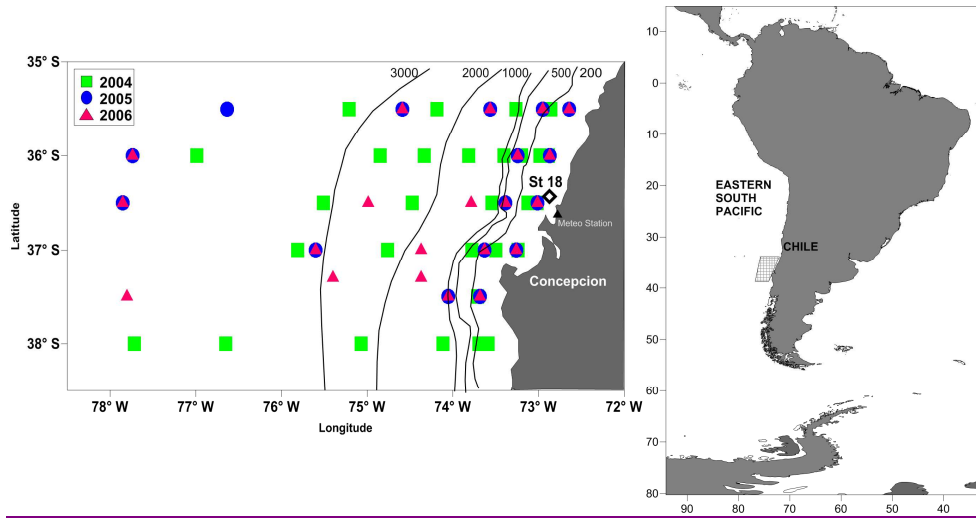


Fig. 2

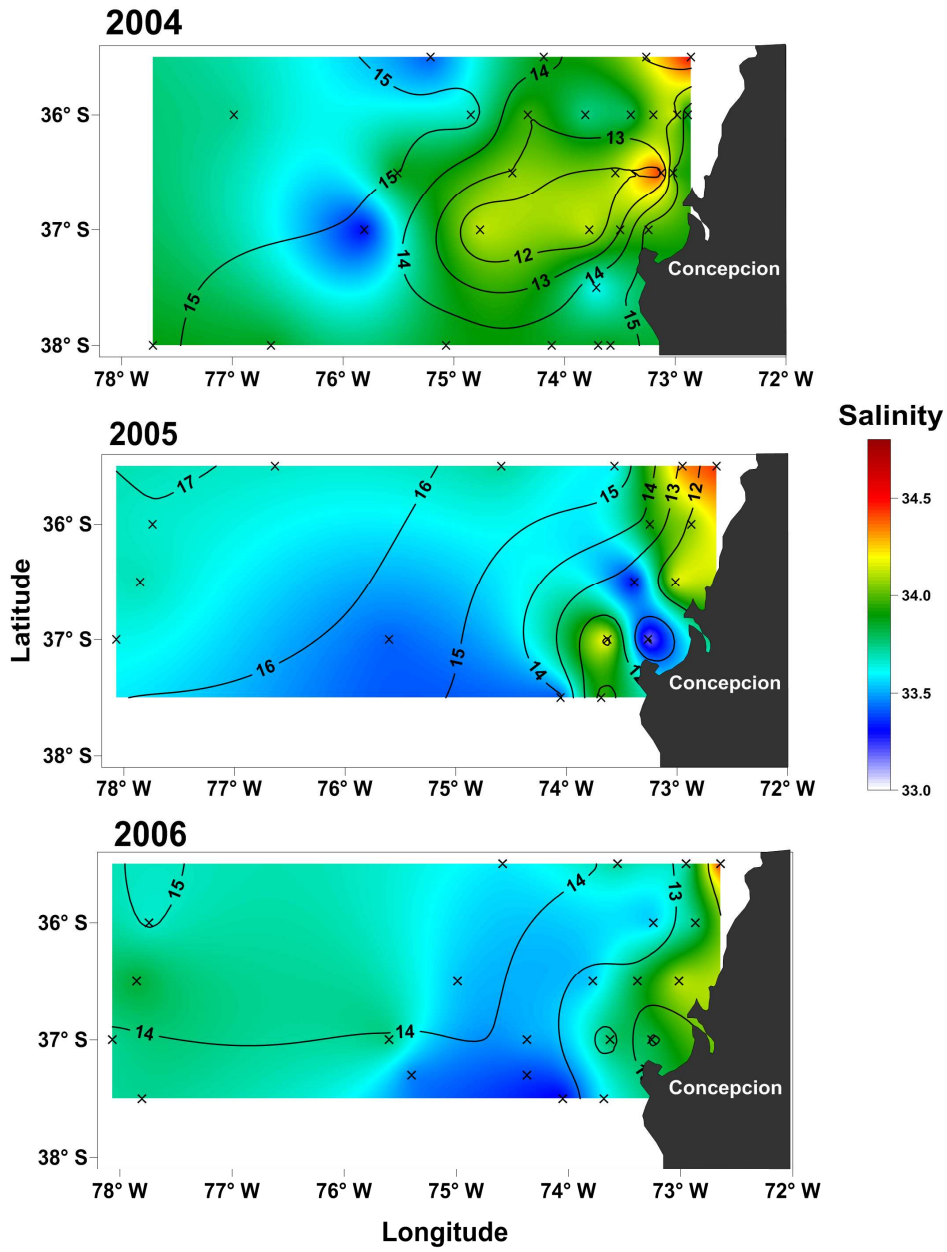


Fig. 3

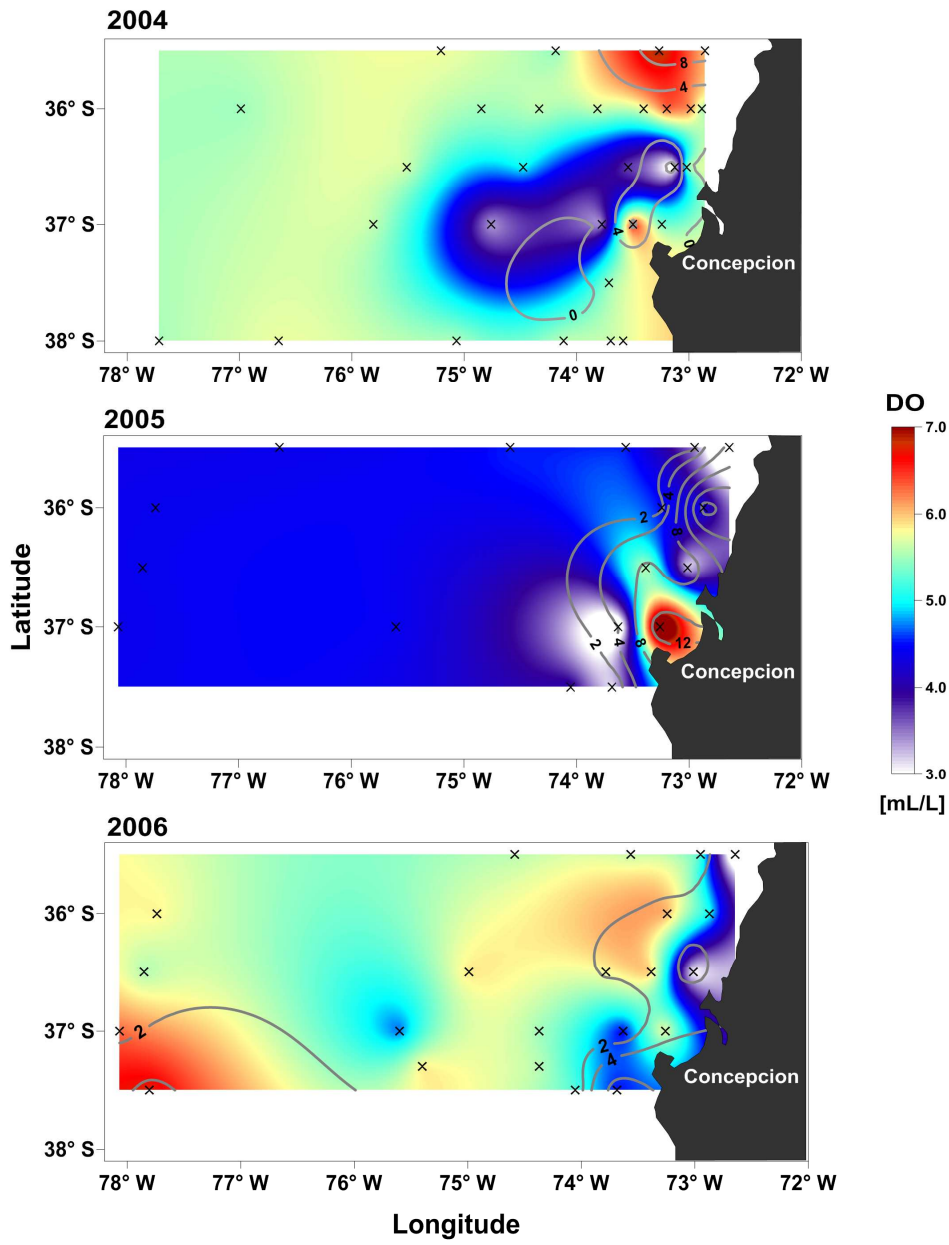


Fig. 4

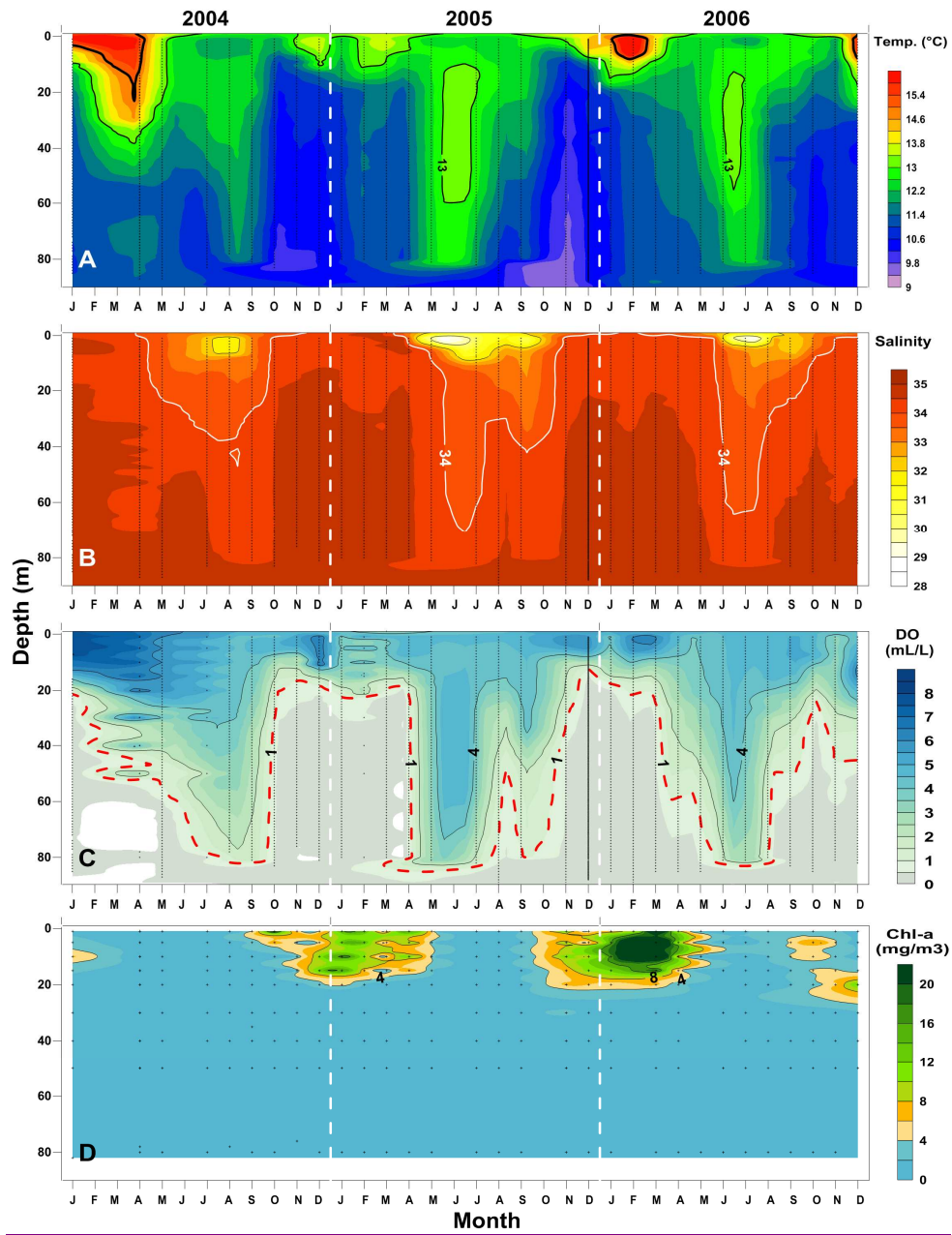


Fig. 5

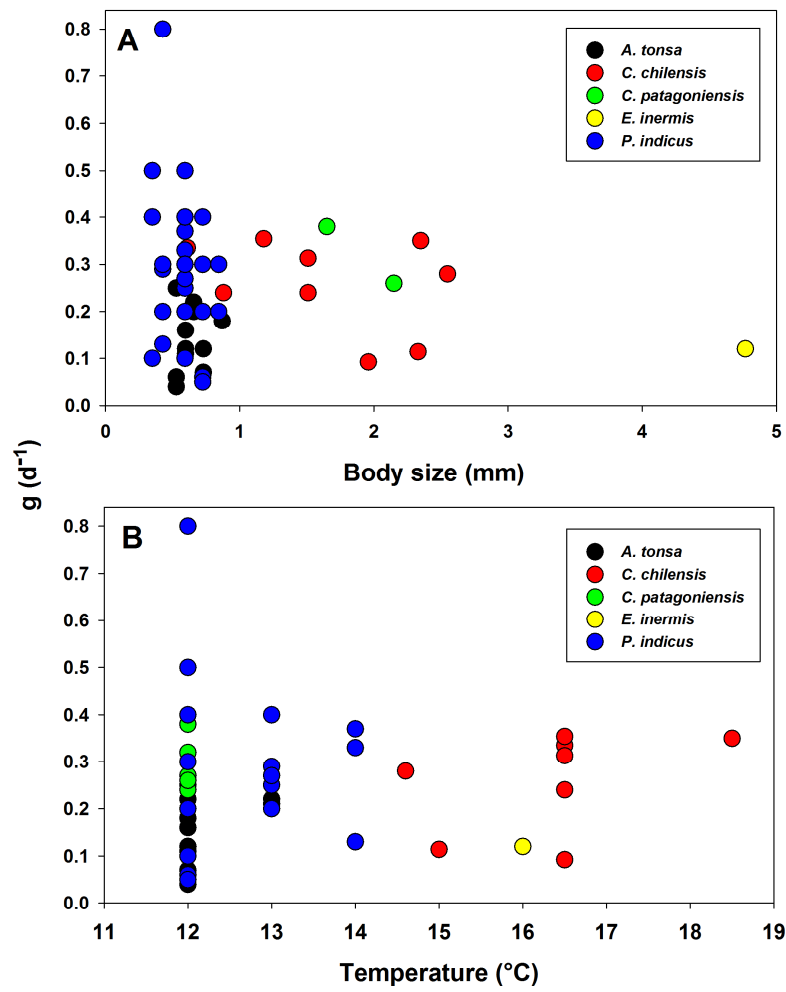


Fig. 6

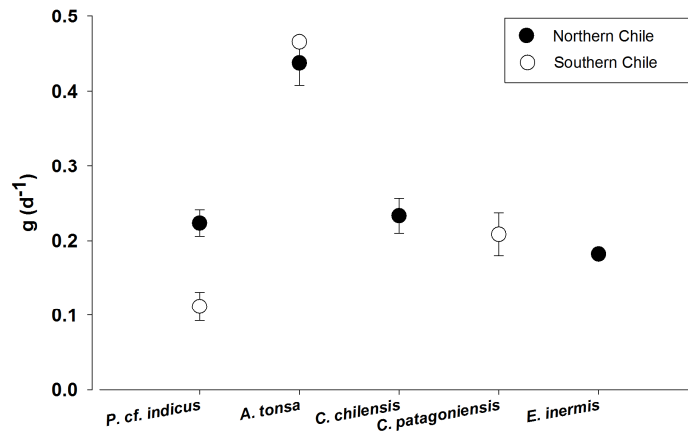


Fig. 7

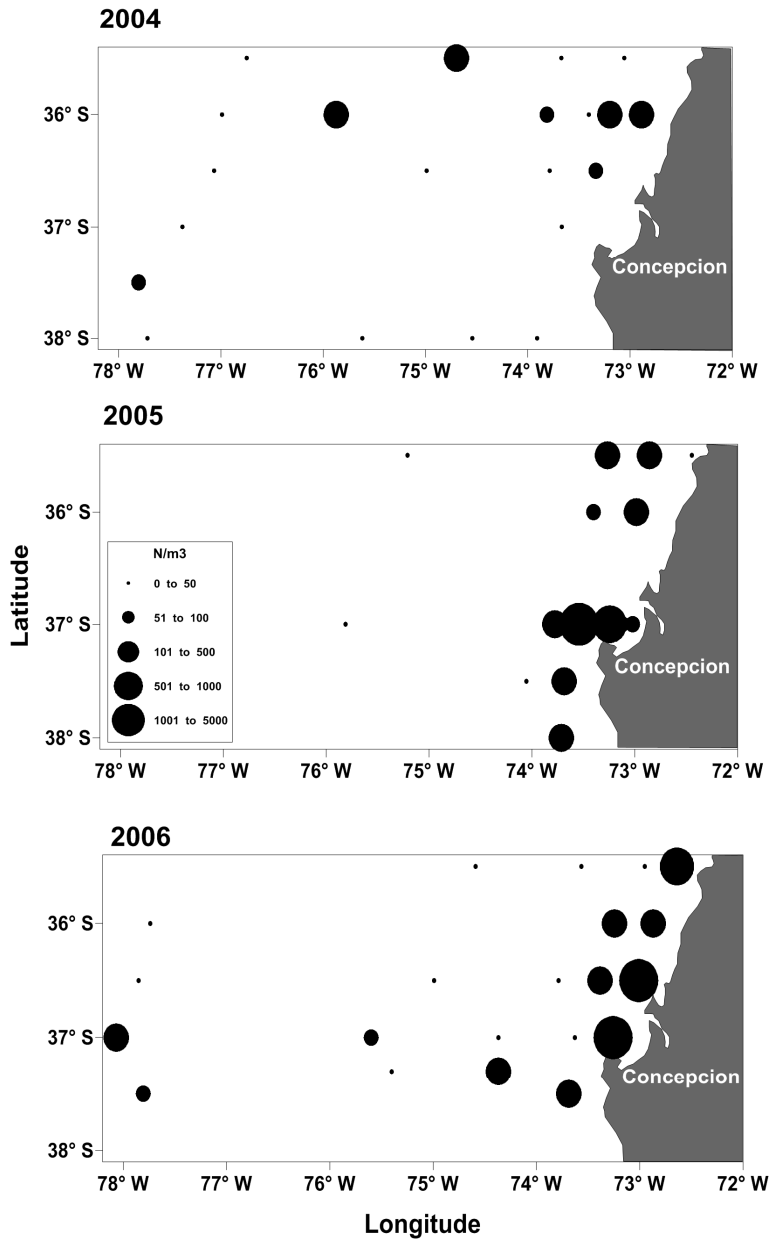


Fig. 8

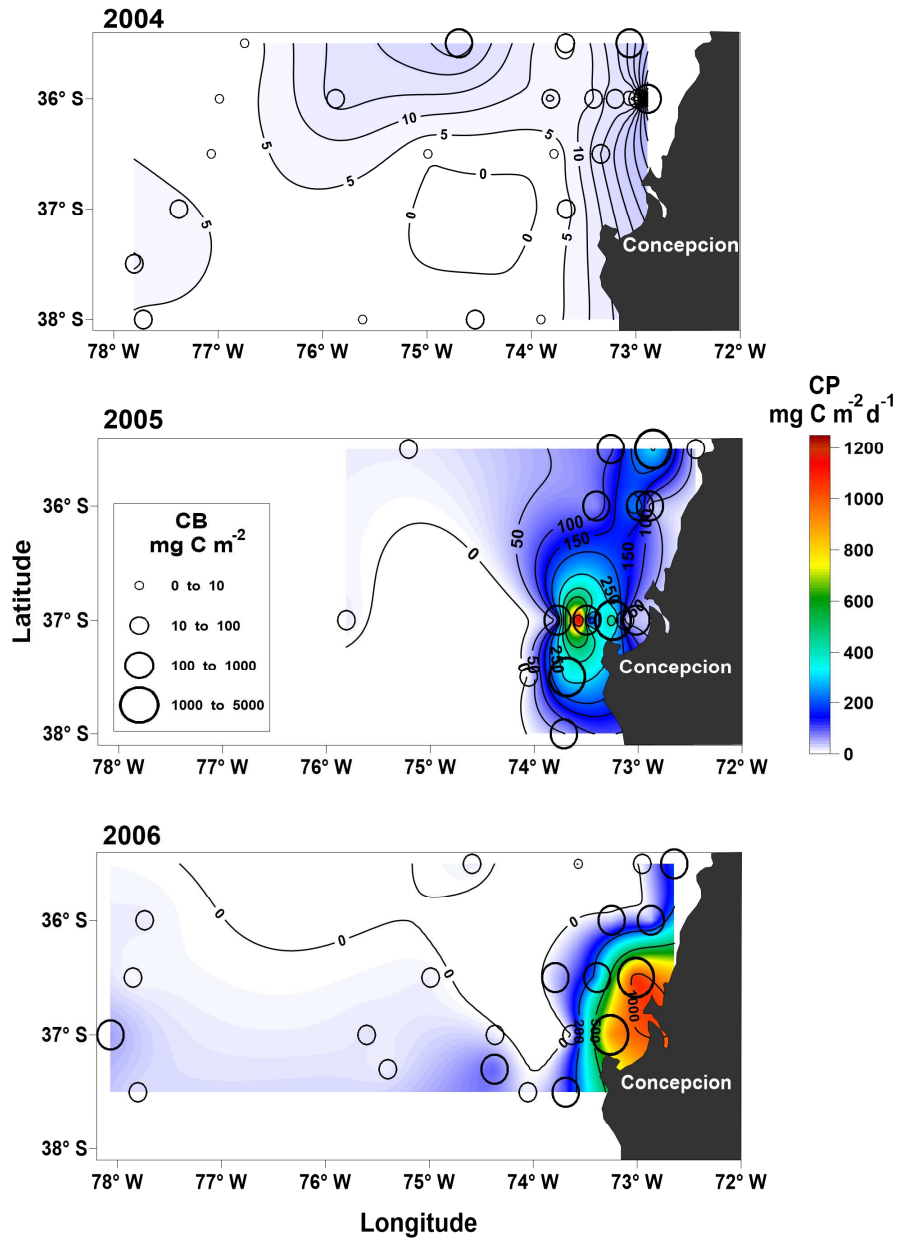


Fig. 9

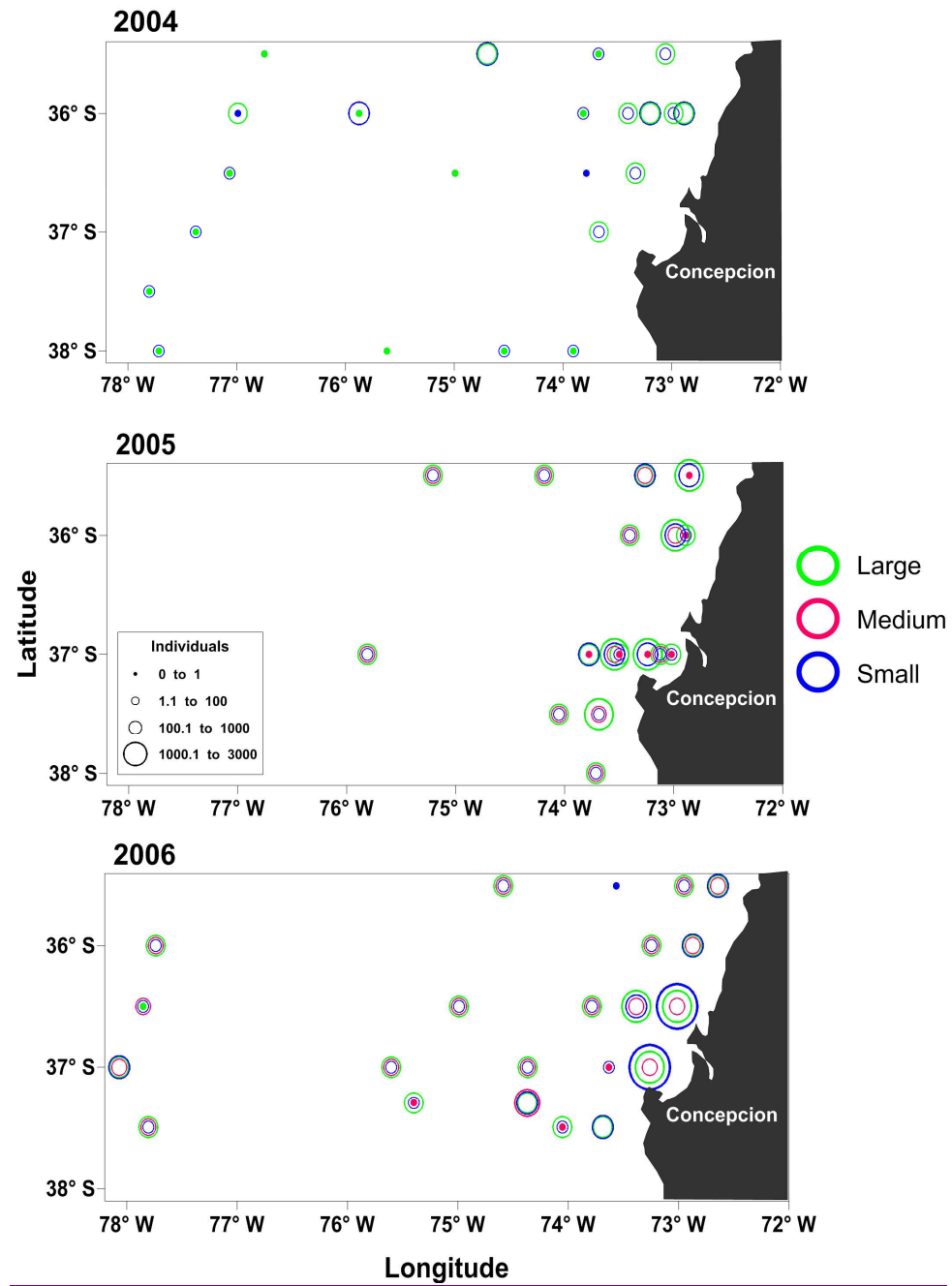


Fig. 10

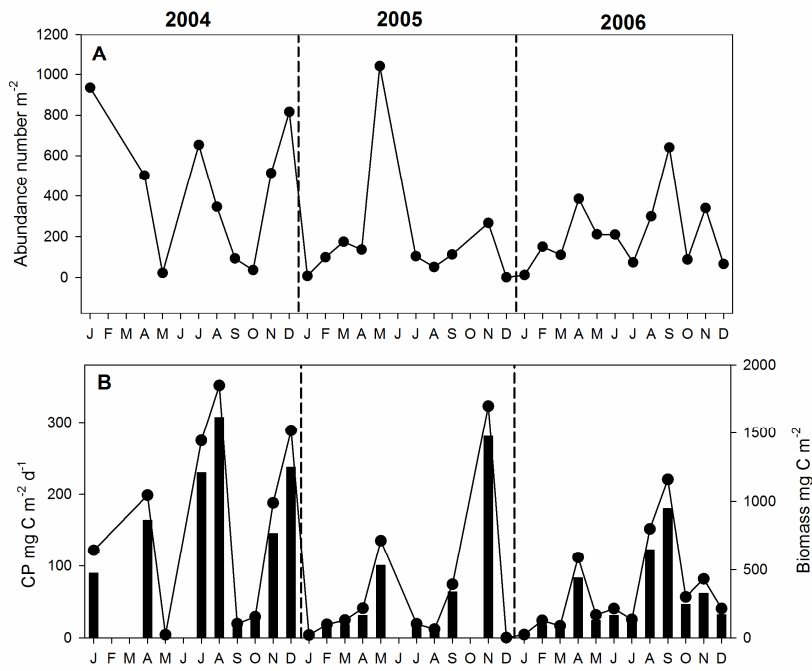


Fig. 11

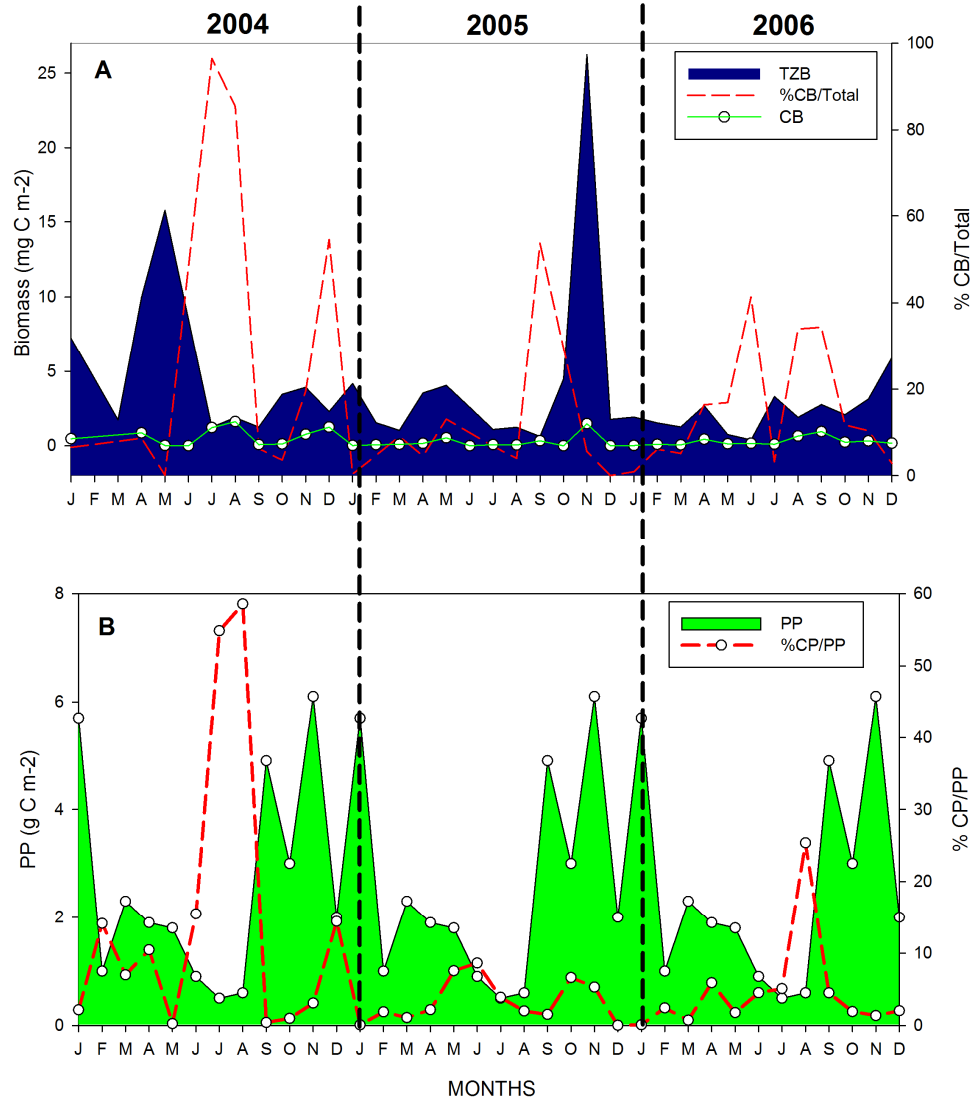


Fig. 12

