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## ***Interactive comment on “A parameterization of respiration in frozen soils based on substrate availability” by K. Schaefer and E. Jafarov***

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This manuscript presents an interesting application of a solute diffusion model within a soil carbon dynamics model to account for the effects of changing liquid water content as frozen soils thaw. I have two main concerns with this manuscript:

First, the same approach was recently published by Tucker (Soil Biology & Biochemistry 78 [2014] 90-96), but this work is not cited and Tucker is not given credit for having developed this idea. Indeed, Fig 2 of the present manuscript is nearly identical in form to Fig. 1 of the Tucker paper. Tucker used data from non-arctic areas, but the issue of freeze-thaw is still applicable. Tucker modified the Dual Arrhenius Michaelis-Menten (DAMM) model, which simulates diffusion of soluble C substrates in soil water films, and he showed how this diffusion is slowed drastically when the water is mostly as ice

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rather than in a liquid phase. He also included the effect of swelling ice occupying more pore space than liquid water, thus also limiting diffusion of O<sub>2</sub> into the soil. He demonstrated that the very large Q<sub>10</sub> values for soil respiration commonly observed across the small temperature increment between frozen and unfrozen soils is attributable to this diffusion effect rather to an actual high temperature dependence of the enzymatic activity. The present manuscript should cite the Tucker paper and the related DAMM papers as the source of this innovation.

Response: We incorporated references to Tucker [2014] (Line 94-6, 446, 453-5). We thank the reviewer for bringing this paper to our attention. We based our modeled liquid water fraction ( $\theta_{liq}$ ) on the Nicolsky et al. [2009] formulation while Tucker [2014] used the Romanovsky and Osterkamp [2000] formulation. Both are derived from the original power law formulation of Lovell [1957], so we would expect strong similarity between Figure 2 in our manuscript and Figure 1 in Tucker [2014]. We used a very different approach from Tucker [2014] and the DAMM model, since SiBCASA does not include Dissolved Organic carbon (DOC), solute diffusion, and oxygen diffusion (Line 115-7, 144). However, we give due credit to Tucker [2014] for linking liquid water fraction to simulated respiration (Line 94-6).

Second, I don't understand the discussion about the "original Q<sub>10f</sub>" formulation. The authors don't make it clear what their original formulation was. Is it simply a constant Q<sub>10</sub> across all temperatures? If so, what value of Q<sub>10</sub> was used? Or was the original formulation one in which a very high Q<sub>10</sub> was applied across the freeze-thaw temperature increment and more normal Q<sub>10</sub>s were applied above and below? I suspect that my lack of understanding of this might contribute to my sense that the authors' conclusion about long-term versus short-term effects is counter-intuitive. It would seem to me that it is the short-term respiration response that would not be adequately simulated by the conventional Q<sub>10</sub> model when soil temperature changes from -2C to +1C. For this short-term response across this small temperature range, the diffusional effect needs to be used to skillfully simulate the observed pulse in soil respiration. It also seems

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to me that the longer-term effect of a change of MAT from slightly below 0C to slightly above 0C could be simulated by the traditional Q10 approach. However, the authors have reached the opposite conclusion. I'm obviously missing something, but I believe that their explanation is inadequate.

Response: We clarified the text to state SiBCASA used a Q10 of 1.5 and a Q10f of 200 (Line 129-32). In SiBCASA, Q10 formulation applies to all temperatures and the Q10f formulation applied to  $T < 0\text{ }^{\circ}\text{C}$  (Line 129-32). We clarified the text to state that 'long term' in this context is 500-10,000 years for temperatures below  $-1\text{ }^{\circ}\text{C}$ , where the Q10f formulation depletes the frozen carbon (Line 89-91, 476). We already state that the Q10f formulation does a pretty good job for shorter time scales of 10 years or less (Line 472-3). We also clarified the text to state that there are two aspects of substrate availability: the amount of thawed organic matter and DOC diffusion in the thin water films (Line 54-5, 142-3). SiBCASA does not include a DOC pool and solute diffusion processes (Line 115-7, 144) so this parameterization focuses on the amount of thawed organic matter (Line 144-5, 156). We also clearly state that the new parameterization completely replaced the original Q10f formulation (Line 147-8).

Although this work is not entirely novel, because Tucker already applied this approach, the work is still worthy of publication because it is being implemented in a larger model that has broad applications to the fate of carbon in areas of permafrost. As long as Tucker is acknowledged (BTW, this reviewer is not Tucker) and the explanation of short-term versus long-term effects is better explained, I believe that this work could be suitable for publication.

Response: We referenced Tucker [2014] (Line 94-6, 446, 453-5) and we better explained what we mean by 'long-term' and 'short term' (Line 89-91, 476)

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