

1 **Estimation of nutrient contributions from the ocean**
2 **across a river basin using stable isotope analysis**

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11

12 **Abstract**

13 Total nitrogen (TN), consisting of total particulate nitrogen (TPN) and total dissolved
14 nitrogen (TDN) pools, may be transported in not only river channels but across entire
15 river basins such as via ground water and migratory animals. In general, the amount of
16 TPN exported from mountainous river basins to the ocean is expected to be larger than
17 TDN. Since marine derived nutrients (MDN) are hypothesized to largely be transported
18 in particulate form, it is necessary to investigate the contribution of particulate MDN in
19 forest ground surface soils to the total MDN at the river ecosystem scale. In this study
20 we investigated TN export from an entire river basin, and also estimated the
21 contribution of pink (*Oncorhynchus gorbuscha*) and chum salmon (*O. keta*) to total
22 oceanic nitrogen input across a river basin. The maximum potential contribution of TN
23 entering the river basin as salmon was 23.8 % relative to the total amount of TN
24 exported from the river basin. The contribution of MDN to particulate nitrogen in river
25 basin soils was estimated to be 22.9 % with SD of 3.6 % using stable isotope analysis
26 (SIA) of nitrogen ($\delta^{15}\text{N}$).

27

28 **1. Introduction**

29 Stable Isotope Analysis (SIA) is increasingly being used to examine connectivity in
30 coastal aquatic-terrestrial ecosystems, such as the input of marine derived nutrients
31 (MDN) from the open ocean to coastal and river ecosystems (Wyatt et al., 2010a; Wyatt
32 et al., 2010b; Wyatt et al., 2012, Havik et al. 2014, Adame et al. 2015). In the case of
33 river ecosystems, the transportation of nutrients, such as nitrogen and phosphorus, by
34 migrating fish results in enhancement of biofilms and planktonic productivity in river
35 systems (Juday et al., 1932; Cederholm and Peterson, 1985; Bilby et al., 1996; Gresh et
36 al., 2000; Chaloner et al., 2002; Moore and Schindler, 2004; Yanai and Kochi, 2005;
37 Levi and Tank, 2013, Marcarelli et al. 2014). In most of those cases, terrestrial
38 consumers like mammals, birds, fishes and insects have been shown to play a large role
39 in terms of providing MDN to watersheds (Donaldson, 1966; Ben-David et al., 1997a;
40 Hilderbrand et al., 1999; Gende et al., 2002; Naiman et al., 2002; Wilkinson et al., 2005;
41 Bartz and Naiman, 2005, Koshino et al. 2013). Moreover, MDN inputs have been
42 shown to be an important processes controlling the productivity of the ecosystem. For
43 example, Merz and Moyle (2006) found that the contribution of MDN to the foliar
44 nitrogen of wine grapes was about 18 to 25 %. Also, Hilderbrand et al. (1999)
45 demonstrated that trees and shrubs near spawning streams received 24 to 26 % of their
46 foliar nitrogen from MDN, while Helfield and Naiman (2002) suggested that 15.5 to

47 17.8 % of spruce foliage nitrogen may be provided by MDN. Thus, isotopic methods as
48 an intrinsic geospatial tracer can provide a means to quantify cross-ecosystem transfer
49 of nutrients. As a particularly important transfer mechanism, migrating fish such as
50 salmon have been shown to be necessary for sustainable nutrient-cycles due to their
51 important role as nutrient transporters (Ben-David et al., 1998; Wipfli et al., 1998; Yanai
52 and Kochi, 2005; Gende et al., 2007; Hocking and Reimchen, 2009; Hocking and
53 Reynolds, 2011). Additionally, MDN has been demonstrated to be important not only
54 for river ecosystems but also potentially for upstream lakes (Kline et al., 1990; Kline et
55 al., 1993; Schindler et al., 2003).

56

57 When we consider nutrient flux in a river flowing from its upstream end into the ocean,
58 the flux depends on nutrients supplied not only inside the river itself but also from the
59 entire river basin (Dutta and Nakayama, 2010; Alam and Dutta, 2012; Riggsbee et al.,
60 2008). Particulate nutrient flux from a basin, which is derived mainly from surface soils,
61 is generally larger than the flux from dissolved nutrients in mountainous regions
62 (Nakayama et al., 2011). Cederholm et al. (1989) demonstrated that mammals and birds
63 consume migrating fish, which may result in the secondary dispersion of MDN across
64 the river basin associated with the movement of these consumers. Other studies have
65 revealed that mammals incorporate MDN from salmon, which may subsequently lead to

66 re-export to the ocean through river flows (Bilby et al., 1996; Ben-David et al., 1997a;
67 Ben-David et al., 1997b; Hilderbrand et al., 1999; Szepanski et al., 1999; Reimchen,
68 2000, Holtgrieve et al. 2009). However, the contribution of MDN to surface soils, which
69 may be transported from a river basin to the ocean as suspended sediments, at the river
70 basin scale has not been adequately quantified in natural systems due to the difficulty in
71 quantifying complex food webs and making accurate biomass estimates.

72

73 In this study we present TN transport across an entire river basin to the ocean, the
74 potential contribution of TN from the ocean to the river basin by salmon, and the
75 contribution of MDN to surface soils in the river basin. Integrated stable isotope
76 analysis of geological, hydrological and biological compartments of the ecosystem
77 allowed us to estimate the nutrient budget for a natural river basin, suggesting it may be
78 important to conserve ocean-river connectivity in such systems.

79

80 **2. Geophysical setting**

81 Our target area, the Shiretoko Peninsula, was registered as a World Natural Heritage
82 area in July of 2005. Shiretoko is located at the southernmost extent of drift ice and its
83 ecological systems exhibit high biodiversity and high rates of nutrient circulation,
84 particularly due to runs of pink (*Oncorhynchus gorbuscha*) and chum (*O. keta*) salmon

85 from the Sea of Okhotsk. Potential runs of salmon along the coast of Hokkaido in the
86 Sea of Okhotsk have been estimated at about 29,900,000 individuals a year (Hokkaido
87 National Fisheries Research Institute, Fisheries Research Agency, 2009), equivalent to
88 2590 tons of total nitrogen. The size of the Okhotsk coastal region of Hokkaido is about
89 24,000 km², which corresponds to that mean total nitrogen input from the ocean of
90 about 108 kg km⁻² yr⁻¹ if we assume that all salmon run up rivers and the total nitrogen
91 is completely distributed into river basins. Shiretoko is located on the northeast coast of
92 Hokkaido, Japan (approximately 43°57' N to 44°21' N and 144°58' E to 145° 23'E),
93 and has a width, length and maximum altitude of about 15 km, 50 km and 1660 m,
94 respectively (Fig. 1). The Rausu River basin was selected as the main study area
95 because its watershed is the largest in the region and it is considered a representative
96 watershed in the Shiretoko Peninsula. The watershed area, river length, and the mean
97 river slope are 32.5 km², 7 km, 1/7, respectively. Because of the steep slope, nutrient
98 flux due to suspended sediments is larger than due to dissolved nutrients (Nakayama et
99 al., 2011). Field experiments were carried out over 5 years from 2008 to 2012. For
100 comparison with the Rausu River basin, stable isotope analyses were also carried out in
101 2014 in the Rusa River basin. Here, the watershed area, river length, and the mean river
102 slope are 9.2 km², 5.5 km, 1/7, respectively (Fig. 1).

103

104 **3. Methods**

105 **3.1 Nitrogen from a river basin to the ocean**

106 MDN supplied from the ocean to surface soils in a river basin generally includes feces
107 of mammals, droppings of birds, and the remains of salmon preyed upon by mammals,
108 birds and insects. These MDN are recycled within the terrestrial ecosystems and mainly
109 stored as soil organic matter (SOM). Thus, to focus on the influence of SOM on TPN
110 export, soil particles with diameter of less than 500 μm after rinsing in 1N-HCL
111 solution were used in this analysis. The analysis does not allow evaluation of TN
112 (TPN+TDN) export from the river basin to the ocean. However, TPN export from an
113 entire river basin has been revealed to be larger than TDN in the Rausu River basin due
114 to its steep slope (Nakayama et al., 2011).

115

116 We believe that the effect of denitrification on the $\delta^{15}\text{N}$ is negligible in our case. In
117 general, some proportion of the nitrogen is reduced due to denitrification, which results
118 in an increase in $\delta^{15}\text{N}$ of the soil (Yamada et al., 1996). However, Wada et al. (1984)
119 demonstrated that denitrification seems to have a small effect on the variation of $\delta^{15}\text{N}$ in
120 SOM under aerobic conditions close to the ground surface in a natural forest. Moreover,
121 Rennie et al. (1976) revealed that the isotope ratio of nitrogen in ground surface soils is
122 identical to that in organic nitrogen in the natural forest, which suggests that

123 denitrification does not involve any isotope fractionation. Mckinley et al. (2013) also
124 demonstrated that the $\delta^{15}\text{N}$ of surface soil is aerobic in forests when the water table is
125 not close to the ground surface. Since our sampling was carried out within the top 5 cm
126 of the soil and the surface soil is not saturated due to the steep slope, the SOM sampled
127 was considered to be under aerobic conditions.

128

129 We made an attempt to estimate the contribution of MDN to SOM resulting from the
130 accumulation of particulate organic matter, which should be directly related to the
131 potential riverine transport of MDN back to the ocean as suspended sediments, by
132 sampling surface soils across the Rausu River basin (Fig. 2). TN, TDN and TPN were
133 measured at St.0 around the river mouth from 2007 to 2009 in the Rausu River basin
134 (Fig. 2). The nitrogen concentration of filtered and non-filtered water samples were
135 analyzed by the cadmium reduction-colorimetric method. Annual TN and annual TDN
136 exports to the ocean were evaluated using the river discharge at St.0 with
137 TDN-discharge and TPN-discharge curves. The TDN-discharge and TPN-discharge
138 curves were produced using ten different peak discharge floods and base flow
139 discharges. As river discharge was not measured during the winter season from January
140 to March, a storage function method was applied to estimate river discharge from 2008
141 to 2012 (Michael, 1978; Michael et al., 1979). The validity of the storage function

142 method was confirmed through comparison with the observed river discharge from
143 April to December.

144

145 Surface soil samples were taken at 12, 20 and 21 stations in 2008, 2009 and 2012,
146 respectively. Three soil samples were collected at each sampling station in order to
147 account for small scale variability in SOM (Fig. 2 and Table 1). In 2008, fewer samples
148 were taken as we did not have permission to sample surface soils in special protection
149 zones. Surface soils were sampled from three different points at each station in a
150 volume of 15 cm × 15 cm × 5 cm (height × width × depth). Surface soil sampling
151 stations in 2012 are shown in Fig. 2. Since previous studies have revealed that surface
152 soil transport is related to the spatial distribution of surface soil type, land-use type and
153 vegetation (Ishida et al., 2010), the location of each sampling station was selected by
154 dividing the river basin into 21 domains (sub-basin areas) that vary in soil type and
155 vegetation (Figure 1). The spatial distribution of surface soil types was divided into 6
156 categories. Although the spatial pattern in vegetation is complicated, the vegetation can
157 generally be categorized in terms of altitude. Since Shiretoko is protected as a natural
158 World Heritage area, all areas studied are classified as forest and have high vegetation
159 cover.

160

161 **3.2 Salmon runs**

162 To evaluate the contribution of salmon to SOM, salmon runs were investigated in the
163 Rausu River. Salmon were caught at the river mouth for artificial incubation and release,
164 providing an estimate of the number of salmon caught by the apparatus (Hokkaido
165 National Fisheries Research Institute, Fisheries Research Agency, 2009). The apparatus
166 for catching salmon consisted of a lattice fence, which does not obstruct flood flows or
167 completely block the runs of salmon. Therefore, it was necessary to quantify the capture
168 rate of the apparatus in order to estimate the actual volume of salmon runs. Field
169 observations were conducted in the Tokorohoronai River, which is located in the same
170 region of Hokkaido but where it is customary to remove the catching apparatus before
171 and after the salmon run season, allowing us to monitor salmon escape rates from the
172 apparatus and the salmon run under open conditions at the same place. The capture rate
173 of the apparatus was calculated using the number of salmons passing the observation
174 point in a channel section of 3 m width and 0.2 m depth; the Rausu River width (about
175 15 m) is too wide for this type of observation. We used two infrared cameras
176 (SM-AVIR-602S, Hero Corp., Izumo, Japan) placed 2 m above the river surface and
177 recorded continuous videos to monitor the individual salmon passing this 3 m section.
178 Videos were taken from the 25th to 28th of November (before removal of the apparatus)
179 and from the 4th to 7th of December (after removal of the apparatus) in 2013. The

180 number of salmon was calculated as the net number of running upstream salmon by
181 identifying individual salmon at the observation point. No salmon were captured and
182 tagged for individual identification. There was no influence of rainfall during the
183 observation period.

184

185 **3.3 Stable isotope analysis**

186 Stable isotope ratios of carbon ($\delta^{13}\text{C}$) and nitrogen ($\delta^{15}\text{N}$) were measured using a Delta
187 Plus Advantage mass spectrometer (Thermo Electron) coupled with an elemental
188 analyzer (Flash EA 1112, Thermo Electron) at the Port and Airport Research Institute,
189 Japan (Table 1 for $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ of SOM in 2012). Stable isotope ratios are expressed
190 in δ notation as the deviation from standards in parts per thousand (‰) according to the
191 following equation:

192

$$193 \quad \delta^{13}\text{C}, \delta^{15}\text{N} = [\text{R}_{\text{sample}} / \text{R}_{\text{standard}} - 1] \quad (1)$$

194

195 where $\text{R} = {}^{13}\text{C}/{}^{12}\text{C}$ or ${}^{15}\text{N}/{}^{14}\text{N}$.

196

197 Vienna Pee Dee Belemnite and atmospheric nitrogen were used as the isotope standards
198 of carbon and nitrogen, respectively. The analytical precision in the mass spectrometer

199 system based on the standard deviation of the internal reference (L-histidine) replicates
200 was <0.15‰ for both $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$. The contribution of MDN to SOM was evaluated
201 by applying a two source mixing model based on stable isotope analysis (SIA) of
202 carbon ($\delta^{13}\text{C}$) and nitrogen ($\delta^{15}\text{N}$) (Kline et al., 1998; Moore and Semmens, 2008;
203 Hossler and Bauer, 2012). Salmon tissue isotopes were considered representative of the
204 isotope composition of ocean productivity. To isotopically characterize terrestrial
205 productivity, we considered one terrestrial end-member (source): Soil Samples
206 exhibiting the Lowest $\delta^{13}\text{C}$ at St.14 and $\delta^{15}\text{N}$ at St.18 in 2009 (hereafter SSL), and thus
207 assumed to have the highest terrestrial contribution to SOM. SSL was collected close to
208 the top of the mountain, where MDN is not expected to influence isotope values.
209 Representative soil samples collected in the same river basin were chosen because they
210 have isotopically similar characteristics to the target soil samples in this study.

211

212 The contribution of MDN to SOM was evaluated using a two sources mixing model
213 based on the measured $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$. The average contribution in the Rausu River
214 basin was computed using each sub-basin area obtained from the Thiessen method.

215

$$216 \quad f_{C_MDN} + f_{C_LDN} = 1 \quad (2)$$

$$217 \quad f_{C_MDN} \delta^{13}\text{C}_{salmon} + f_{C_LDN} \delta^{13}\text{C}_{SSL} = \delta^{13}\text{C}_{soil} \quad (3)$$

218 $f_{N_MDN} + f_{N_LDN} = 1$ (4)

219 $f_{N_MDN} \delta^{15}N_{salmon} + f_{N_LDN} \delta^{15}N_{SSL} = \delta^{15}N_{soil}$ (5)

220

221 where f_{C_MDN} and f_{C_LDN} are the contributions of MDN and land-derived nutrient (LDN)

222 by carbon, $\delta^{13}C_{salmon}$, $\delta^{13}C_{SSL}$ and $\delta^{13}C_{soil}$ are the stable isotope ratios of carbon for

223 salmon, SSL and soil samples, respectively, f_{N_MDN} and f_{N_LDN} are the contributions of

224 MDN and LDN by nitrogen, $\delta^{15}N_{salmon}$, $\delta^{15}N_{SSL}$ and $\delta^{15}N_{soil}$ are the stable isotope ratios

225 of nitrogen for salmon, SSL and soil samples, respectively.

226

227 As bamboo grass (*Sasa senanensis*) is the dominant species in the study area, bamboo

228 grass was collected at 13 soil sampling points (St.1, St.2, St.3, St.4, St.7, St.8, St.10,

229 St.11, St.12, St.13, St.14, St.17, and St.21). Furthermore, droppings of sea eagles

230 (*Haliaeetus* spp.) and feces of brown bear (*Ursus arctos*), which represent a typical

231 migratory bird and mammal in Shiretoko, were collected to investigate whether or not

232 they include MDN and thus contribute to SOM. Samples of feces and droppings for SIA

233 analysis offer a major advantage, i.e. little isotopic fractionation expected and thus ideal

234 to use the stable isotope values as a MDN tracer (Fry, 2006). Chum salmon tissues and

235 droppings of sea eagles were collected at the river mouth and feces of brown bear were

236 collected at St.14. The samples were pre-treated by rinsing with a chloroform-methanol

237 solution (2:1) prior to SIA to remove isotopically fractionated metabolites, such
238 metabolites in the samples were removed as urea and ammonium (Kuwae et al., 2008;
239 Kuwae et al., 2012).

240

241 **4. Results and Discussion**

242 **4.1 Estimation of nitrogen export to the ocean**

243 During 2007 to 2009 the concentration of TDN was observed to be constant, 0.090 mg
244 L⁻¹ (SD 0.022 mg L⁻¹), regardless of the discharge in the Rausu River. In contrast, TPN
245 was revealed to be a function of river discharge ($r^2=0.88$; Eq. 6) (Fig. 3). TPN showed a
246 strong correlation with suspended sediment (SS) concentrations, with SS concentration
247 increasing with increasing river discharge (Fig. 3). TPN was modeled using our field
248 observation results, discharge and TPN as in (6).

249

$$250 \quad \text{TPN} = 0.0032 \times Q^{1.771} \quad (6)$$

251

252 where Q is the river discharge (m³ s⁻¹).

253 The validity of the storage function method model was confirmed using the observed
254 river discharge from April to September of 2009, which resulted in a Coefficient of
255 Determination (CoD) of 0.61. The reliability of the model has been shown to be high

256 enough for the analysis of river discharge when the CoD is more than 0.6 (Dutta and
257 Nakayama, 2010). Annual mean exports of TDN, TPN and TN from 2008 to 2012 were
258 5210 kg yr⁻¹, 14750 kg yr⁻¹ and 19960 kg yr⁻¹, respectively. Since the size of the Rausu
259 River basin is 32.5 km², the annual mean exports of TDN, TPN and TN per unit
260 catchment area equate to 160 kg km⁻² yr⁻¹, 454 kg km⁻² yr⁻¹ and 614 kg km⁻² yr⁻¹,
261 respectively (Table 2). The average concentrations of TDN and TPN from 2008 to 2012
262 were 0.090 mg L⁻¹ and 0.216 mg L⁻¹, which agrees with a previous study at the site
263 (Nakayama et al., 2011).

264

265 **4.2 Contribution of salmon runs to nitrogen input from the ocean**

266 The average number of salmon passing the cameras in the Tokorohoronai River during
267 the 4 days while the apparatus for catching salmon was present was 0.49 hr⁻¹. The
268 average numbers for 4 days after the apparatus was removed from the river was 0.61
269 hr⁻¹, so the rate of capture of salmon by the apparatus (CS) was estimated as 20 %:
270 $(0.61-0.49) / 0.61 = 0.20$. Since the field observations were conducted at the end of
271 November and the beginning of December after the peak of salmon runs, floods may
272 damage the apparatus for catching salmon and the 20 % capture rate may be an
273 underestimate. Therefore, we attempted to apply two larger capture rates, 50 % and

274 80 %, in order to demonstrate the influence of this estimate on our calculations of the
275 possible nutrient re-export from the ocean due to salmon runs.

276

277 In the Rausu River, the annual average numbers of salmon caught by the apparatus at
278 the river mouth were 3075 and 10580 for chum and pink salmon, respectively, from
279 2001 to 2009. By assuming that all apparatuses have the same rate of capture, the
280 potential for chum and pink salmon runs can be estimated as 15375 and 52900 (CS
281 20 %), 6150 and 21160 (CS 50 %), and 3844 and 13225 (CS 80 %), respectively. The
282 average weight of chum and pink salmon in this region are 3.3 kg and 2.0 kg,
283 respectively (Makiguchi et al., 2007), which include a nitrogen content of about 0.100
284 kg and 0.0608 kg, respectively (Larkin and Slaney, 1997). Therefore, annual TN
285 potentially transported by chum and pink salmon is estimated to be 1542 kg yr⁻¹ and
286 3216 kg yr⁻¹ (CS 20 %), 617 kg yr⁻¹ and 1287 kg yr⁻¹ (CS 50 %), and 386 kg yr⁻¹ and
287 804 kg yr⁻¹ (CS 80 %), respectively. Finally, the annual TN transported by chum and
288 pink salmon per unit catchment area can be estimated as 146 kg km⁻² yr⁻¹ (CS 20 %), 59
289 kg km⁻² yr⁻¹ (CS 50 %), and 37 kg km⁻² yr⁻¹ (CS 80 %), (SD 19 kg km⁻² yr⁻¹), which
290 corresponds to the contribution of TN by salmon, 23.8 % (CS 20 %), 9.5 % (CS 50 %),
291 and 6.0 % (CS 80 %), relative to the annual outflow of TN per unit area (considered to
292 be 100 %) (Table 2).

293

294 **4.3 Contribution of MDN to SOM in the Rausu River basin**

295 Both $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ of SOM were lower than those of the salmon (Fig. 4). Interestingly,
296 SSL had almost the same value for mean $\delta^{15}\text{N}$ as bamboo grass, which confirms that
297 bamboo grass can be considered as a major source of LDN. The isotopic composition of
298 salmon as representative of oceanic $\delta^{15}\text{N}$ and $\delta^{13}\text{C}$ were 11.0 and -20.5, respectively.
299 The $\delta^{15}\text{N}$ and $\delta^{13}\text{C}$ of SSL were -3.2 and -29.5, respectively. Therefore, the three-year
300 average estimate of the contribution of MDN to SOM based on $\delta^{15}\text{N}$, depending on the
301 choice of terrestrial isotope values, was 22.9 % (SD 3.6 %) using a two sources mixing
302 model (Fig. 5). For reference, the three-year average estimate of the contribution of
303 MDN to SOM based on $\delta^{13}\text{C}$ was 17.7 % (SD 1.1 %) (Fig. 5). Since the annual export
304 of TPN per unit area from the Rausu River basin to the ocean was $454 \text{ kg km}^{-2} \text{ yr}^{-1}$,
305 annual re-export of TPN originally derived from the ocean is estimated to be 104 kg
306 $\text{km}^{-2} \text{ yr}^{-1}$ ($= 454 \text{ kg km}^{-2} \text{ yr}^{-1} * 22.9 \%$) (SD $16 \text{ kg km}^{-2} \text{ yr}^{-1} = 454 \text{ kg km}^{-2} \text{ yr}^{-1} * 3.6 \%$)
307 based on the contribution of MDN to SOM (Fig. 5 and Table 2).

308

309 We believe that the higher $\delta^{15}\text{N}$ of SOM in surface soils around the Rusa River is not
310 associated with human impacts but with higher contributions from MDN. Wada et al.
311 (1984) demonstrated that the $\delta^{15}\text{N}$ of SOM in forests may show significant variation in

312 the surface soil, for example by as much as about -3 ‰ to -2 ‰ at Jumonji in Chichibu
313 and at Mt. Shigayama, and about 1 ‰ to 5 ‰ at Memuro in the eastern Hokkaido. The
314 $\delta^{15}\text{N}$ variation at Memuro was obtained in the Hokkaido Agricultural Experimental
315 Station, which is located 10 km from the center of Obihiro city where 150,000 people
316 live. Therefore, the 1 ‰ to 5 ‰ variation there likely reflects the influence of
317 anthropogenic nitrogen emissions. The $\delta^{15}\text{N}$ variation of about -3 ‰ to -2 ‰ in surface
318 soils at Jumonji and at Mt. Shigayama may support our assumption that the larger $\delta^{15}\text{N}$
319 is, the higher the contribution of MDN. In order to confirm our assumption, we carried
320 out similar field observations in the Rusa River basin (Fig. 6). In the Rausu River, only
321 a part of the area is registered as a special protection zone of the Natural World Heritage
322 region, but the whole of the Rusa River basin is covered by a special protection zone.
323 The Rusa River basin is thus considered a more protected and natural area as defined by
324 the natural World Heritage conditions compared to the Rausu River. Therefore, the
325 contribution of MDN could be expected to be larger in the Rusa River basin compared
326 to the Rausu River basin (Fig.2). The spatial average of $\delta^{15}\text{N}$ in the Rusa River basin
327 was 1.1 ‰, which is 1.0 ‰ larger than in the Rausu River basin. It could thus be
328 suggested that the higher $\delta^{15}\text{N}$ of SOM in surface soils around the Rusa River is
329 associated with higher contributions from MDN. However, it should be noted that MDN
330 re-export as TN was estimated without the contribution of marine derived TDN, for

331 example, in dissolved form such as urine from bears, and thus should be considered the
332 minimum annual MDN re-export.

333

334 The stable isotope ratios in sea eagle droppings and brown bear feces were higher than
335 LDN, indicating that sea eagles and bears likely transport MDN to the SOM. In the case
336 of multiple food sources, feces and droppings are likely to be enriched in relatively
337 indigestible food sources, when compared with assimilated materials (Sponheimer et al.
338 2003; Kuwae et al.2008). Therefore, in the present study, feces and droppings are likely
339 to be enriched in LDN (e.g., plants) because LDN would be more indigestible than
340 MDN (e.g., fishes). However, such an enrichment does not affect the qualitative
341 investigation, i.e., whether or not feces and droppings include MDN and thus contribute
342 to SOM. Since brown bears are thought to be the major terrestrial consumer of
343 spawning salmon, they may impact re-export of nutrient from the ocean across the river
344 basin, such as through release of MDN-rich urine and feces (Hilderbrand et al. 1999).
345 Rennie et al. (1976) demonstrated that the $\delta^{15}\text{N}$ of surface organic matter is associated
346 with the total organic matter, which includes among other components leaf litter,
347 droppings from birds, and feces from animals. Wada et al. (1984) also revealed that
348 $\delta^{15}\text{N}$ of surface soils is almost identical within different organic nitrogen pools in
349 natural forests. Therefore, it is important to quantify the influence of sea eagles and

350 bears on the nutrient-cycle in these systems. However, based on Fig. 4, we cannot as yet
351 quantify the relative contribution of sea eagles and bears to total MDN transport.

352

353 **5. Conclusions**

354 In recent decades, field experiments and stable isotope analyses have been employed to
355 understand the contribution of salmon runs to river ecosystems. Runs of salmon are
356 thought to play a large role in the sustainability of nutrient circulation due to their
357 contribution to mammals that incorporate MDN and disperse it across the entire river
358 basin, with the MDN potentially re-exported to the ocean through river flows. The input
359 of TN from the ocean to river basin ecosystems has been actively investigated in
360 previous studies, since it can exert great control on ecosystems in which salmon run
361 upstream for spawning. However, the contribution of TN from the ocean across an
362 entire river basin has not previously been examined in detail. This is despite the fact that
363 waterfalls and other obstacles, which inhibit salmon runs, are known to reduce the
364 transport of MDN upstream. This study provides an important quantification of the role
365 of salmon in transporting MDN across an entire river basin of the Shiretoko World
366 Natural Heritage area using stable isotope analysis, and indicates that this is likely an
367 important nutrient pathway that should be preserved in these ecosystems.

368

369 **Author contribution**

370 K. Nakayama designed the field experiments and wrote most of the paper and
371 performed mixing model analysis. Y. Maruya produced the figures using the GIS
372 technical input and carried out runoff analysis. K. Komai helped with river discharge
373 and nitrogen concentration analysis. M. Komata, and K. Komai measured total nitrogen,
374 dissolved total nitrogen and particulate total nitrogen. K. Matsumoto carried out the
375 field experiments on salmon runs and conducted statistical analysis of stable isotopes. T.
376 Kuwae designed the field experiment regarding stable isotopes and carried out stable
377 isotope measurements. All authors read and commented on drafts of this paper.

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379

380

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388

389 **References**

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565

566 Fig. captions:

567

568 Fig. 1. (a) Coastline around the Shiretoko Peninsula and the Rausu River basin. (b)
569 Surface soil type. (c) Vegetation.

570

571 Fig. 2. (a) Elevation of the Rausu River basin. Green circles indicate surface soil
572 sampling stations in September of 2012. Red circles indicate field observation stations
573 for discharge, TDN (total dissolved nitrogen) and TPN (total particulate nitrogen). (b)
574 $\delta^{15}\text{N}$ and sampling stations in 2012. (c) $\delta^{13}\text{C}$ and sampling stations in 2012.

575

576 Fig. 3. River discharge, total particulate nitrogen and suspended sediment at the river
577 mouth of Rausu River. (a) River discharge and concentration of total particulate
578 nitrogen. (b) Concentration of suspended sediment and concentration of total particulate
579 nitrogen.

580

581 Fig. 4. $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$ of bamboo grass (*Sasa senamensis*), SSL (Soil Samples
582 exhibiting the Lowest values of $\delta^{13}\text{C}$ and $\delta^{15}\text{N}$), soil samples, bear feces (*Ursus arctos*),
583 salmon (*Oncorhynchus keta*), and sea eagles droppings (*Haliaeetus spp.*). The bars
584 indicate standard deviation.

585

586 Fig. 5. Contribution of MDN (marine derived nitrogen) from the ocean to the Rausu
587 River basin in 2008, 2009 and 2012 using the two sources mixing model. (a) Average
588 contributions of MDN based on SSL (Soil Samples exhibiting the Lowest values of $\delta^{13}\text{C}$
589 and $\delta^{15}\text{N}$) for $\delta^{15}\text{N}$ were 22.9 %. (b) Average contributions of MDN based on SSL for
590 $\delta^{13}\text{C}$ were 17.7 %.

591

592 Fig. 6. (a) Elevation of the Rusa River basin. Green circles indicate surface soil
593 sampling stations in September of 2012. (b) $\delta^{15}\text{N}$ and sampling stations in 2014. (c)
594 $\delta^{13}\text{C}$ and sampling stations in 2014.

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597 Table captions:

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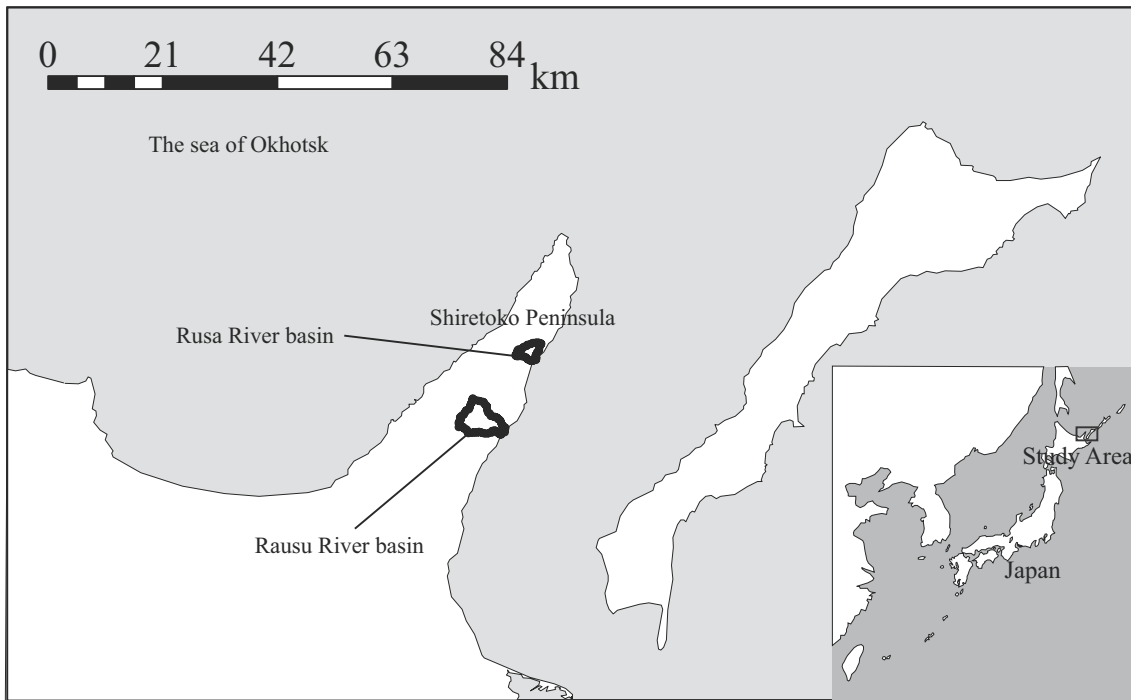
599 Table 1. $\delta^{15}\text{N}$ and $\delta^{13}\text{C}$ of SOM in 2012.

600

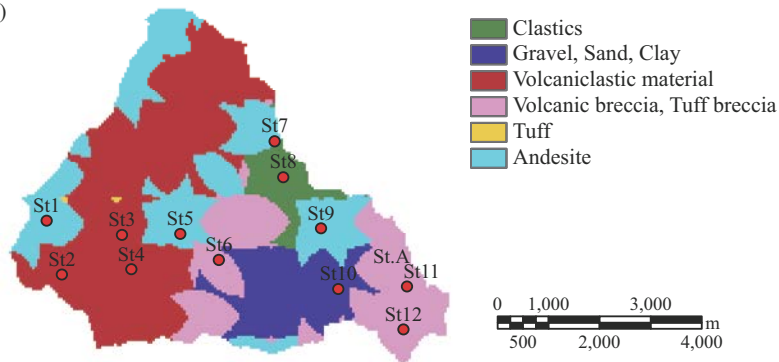
601 Table 2. Summary of annual export and re-export of nitrogen per unit area.

602

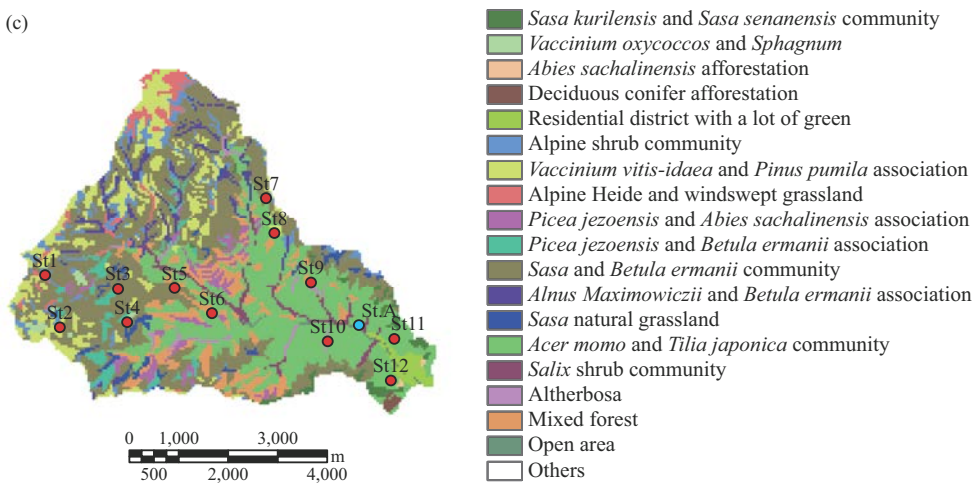
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(b)



(c)

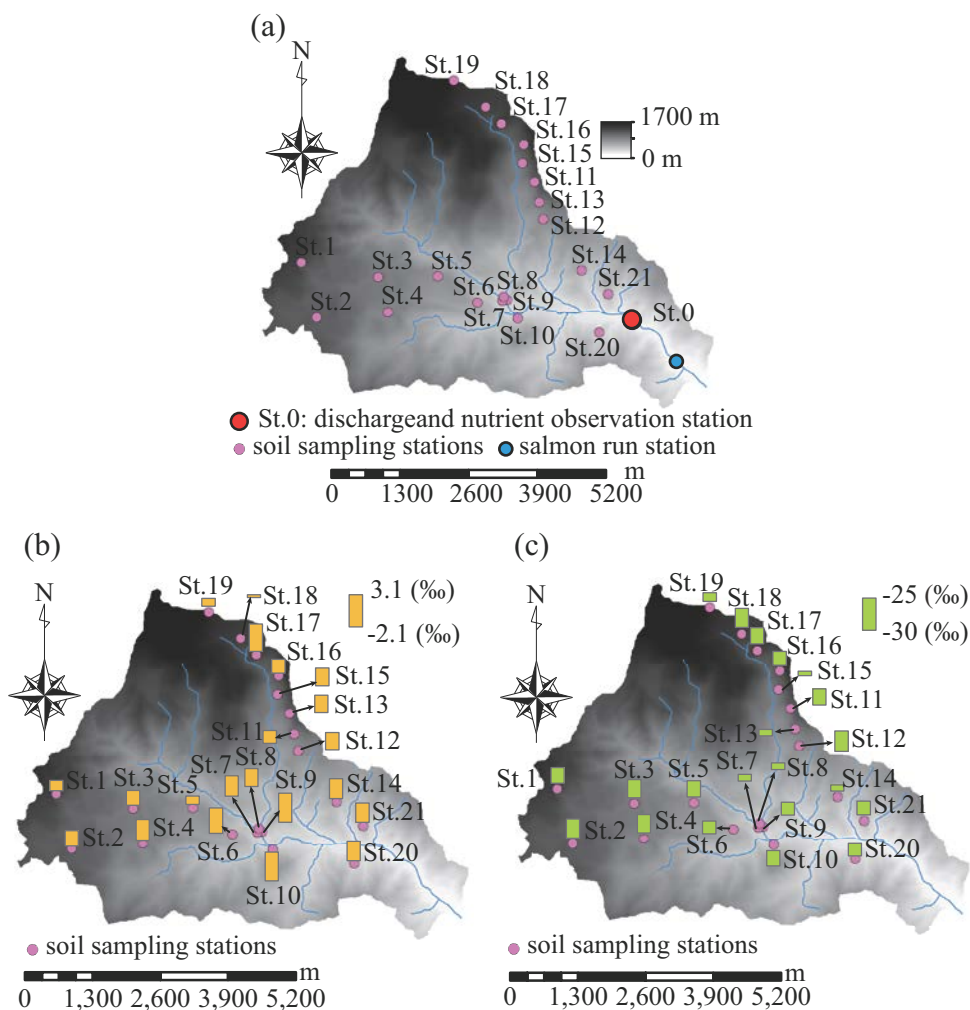


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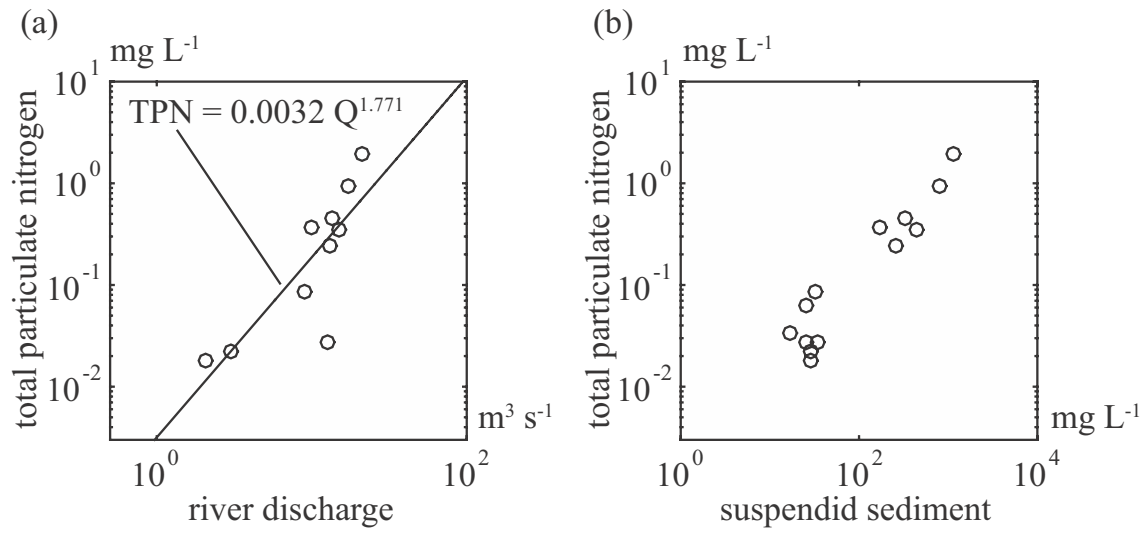
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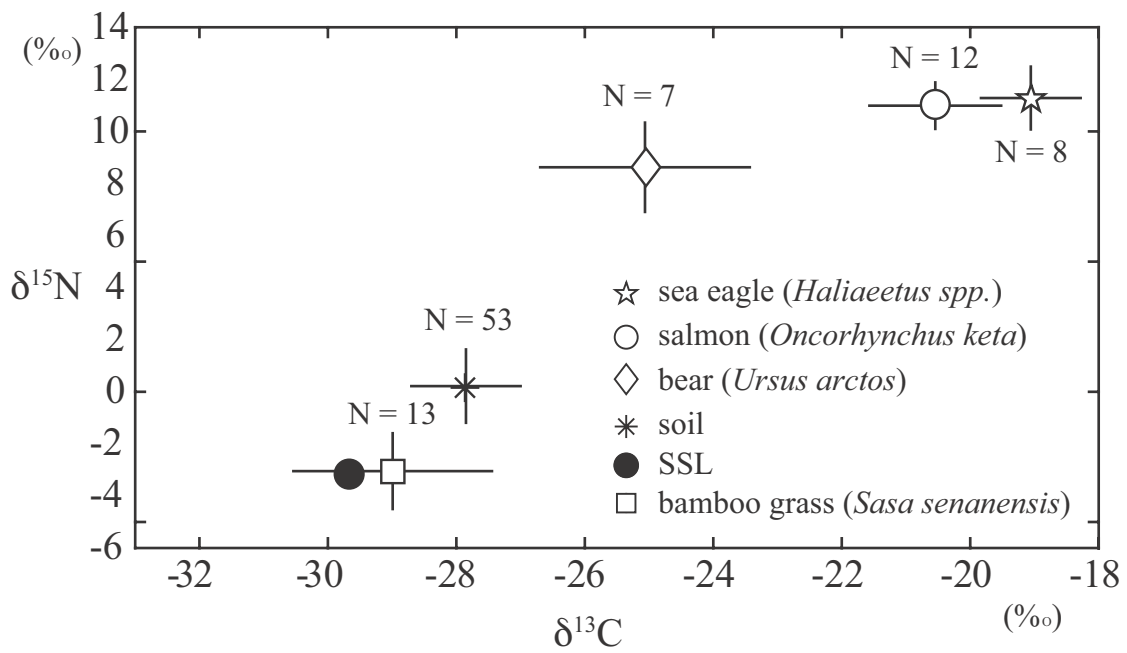
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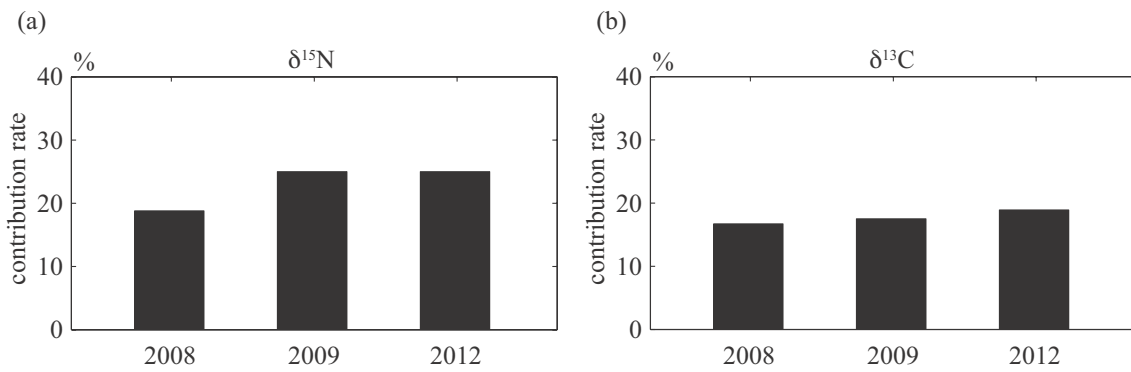
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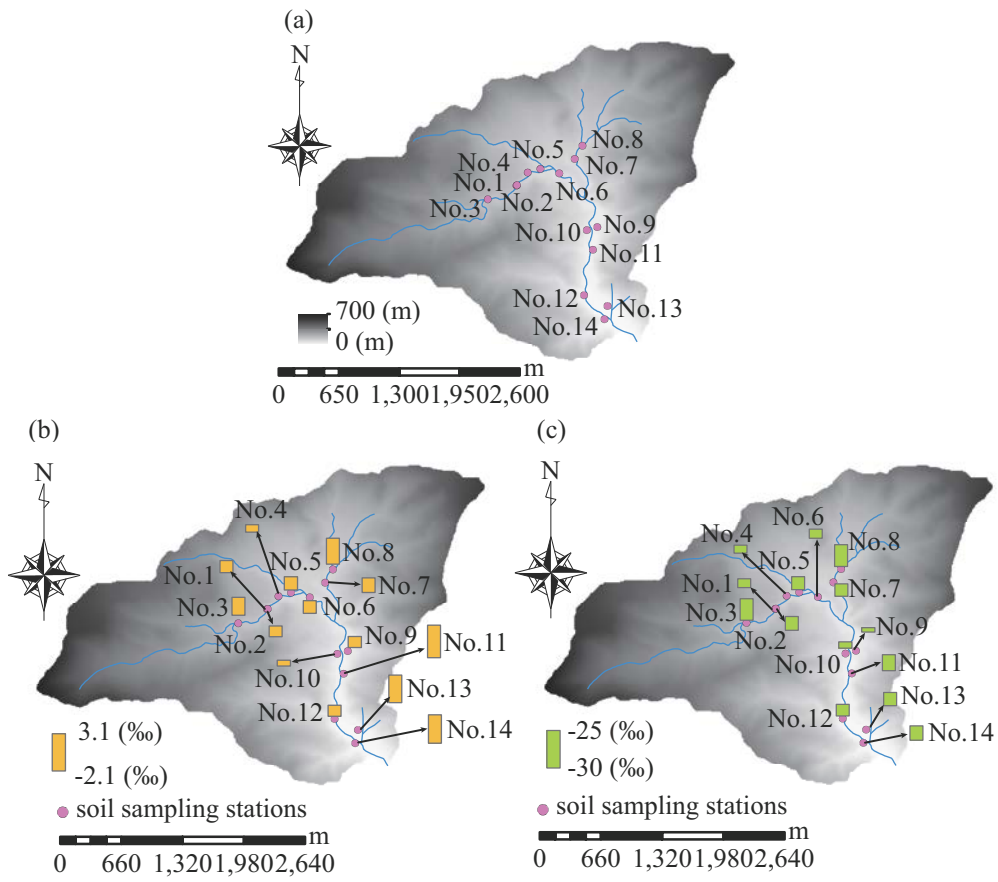
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645 $\delta^{13}\text{C}$ and sampling stations in 2014.

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Table 1. $\delta^{15}\text{N}$ and $\delta^{13}\text{C}$ of SOM in the Rausu River basin in 2012.

650

Station number	$\delta^{15}\text{N}$ (‰)	$\delta^{13}\text{C}$ (‰)
St.1	-0.8	-27.6
St.2	-0.1	-27.0
St.3	-0.1	-27.1
St.4	0.9	-27.1
St.5	-1.1	-27.5
St.6	1.7	-28.0
St.7	0.8	-29.0
St.8	0.4	-29.0
St.9	2.2	-28.0
St.10	2.2	-27.6
St.11	0.3	-27.5
St.12	0.3	-26.8
St.13	-0.4	-29.0
St.14	0.7	-29.0
St.15	0.4	-29.3
St.16	-0.3	-27.8
St.17	2.0	-27.5
St.18	-2.1	-27.1
St.19	-1.3	-28.7
St.20	0.6	-28.0
St.21	0.7	-27.8

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655 Table 2. Summary of annual export and re-export of nitrogen per unit area.

	N export		N re-export	
			Salmon run (%*)	MDN input (%)**
	N kg· y ⁻¹	N kg· km ⁻² · y ⁻¹	N kg· km ⁻² · y ⁻¹	N kg· km ⁻² · y ⁻¹
TDN	5210	160	-	-
TPN	14750	454	-	104 (22.9)
TN	19960	614	CS 20 %, 146 (23.8) CS 50 %, 59 (9.5) CS 80 %, 37 (6.0)	-

656 * = (Salmon run)/(N export)

657 ** = (N export) × (MDN contribution = 22.9)

658

659