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## Atmospheric concentrations of carbon dioxide and its isotopic composition in southern Poland: comparison of high-altitude mountain site and a near-by urban environment

L. Chmura, K. Rozanski, J. M. Necki, M. Zimnoch, T. Kuc, and A. Korus

AGH University of Science and Technology, Krakow, Poland

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Correspondence to: J. M. Necki (necki@agh.edu.pl)

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### Abstract

The results of regular observations of atmospheric CO<sub>2</sub> mixing ratios and its carbon isotope composition (δ<sup>13</sup>C, Δ<sup>14</sup>C), carried out at two continental sites located in central Europe are presented and discussed. The sites (Kasprowy Wierch, 49°14′ N, 19°59′ E, 1989 m a.s.l.; Krakow, 50°04′ N, 19°55′ E, 220 m a.s.l.), are located in two contrasting environments: (i) high-altitude mountaneous area, relatively free of anthropogenic influences, and (ii) typical urban environment with numerous local sources of carbon dioxide. Despite of relative proximity of those sites (ca. 100 km), substantial differences in both the recorded CO<sub>2</sub> levels and their isotopic composition were detected. The CO<sub>2</sub>
mixing ratios measured in the urban atmosphere revealed quasi-permanent excess concentration of this gas when compared with near-by background atmosphere. The annual mean CO<sub>2</sub> concentration recorded in Krakow in 2004 was almost 10% higher than that recorded at high-altitude mountain site (Kasprowy Wierch). Such effect is occuring probably in all urban centers. Carbon isotopic composition of atmospheric CO<sub>2</sub>

<sup>15</sup> proved to be efficient tool for identification the surface CO<sub>2</sub> fluxes into the atmosphere related to fossil fuel burning and their influence on the recorded levels of this gas in the local atmosphere. The available records of  $\Delta^{14}$ C for Krakow and Kasprowy Wierch suggest gradual reduction of <sup>14</sup>C-free CO<sub>2</sub> fluxes into the urban atmosphere of Krakow in the past several years.

#### 20 1. Introduction

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Abrupt increase of atmospheric  $CO_2$  concentration over the last 100 years, by far exceeding the natural variations, has been unequivocally attributed to human activities (Keeling, 1993; GLOBALVIEW, 2004). The anthropogenic impact on the global carbon cycle is mainly related to fossil fuel and biomass burning, land-use changes, as well as various industrial activities (Marland et al., 2002). Although the level of atmospheric  $CO_2$  is increasing, its growth rate is varying with time, responding to changes in the

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magnitude of the CO<sub>2</sub> fluxes entering and leaving the atmospheric carbon reservoir. These short-term changes have both anthropogenic and natural origin (Cias et al., 1995; Keeling et al., 1995). Identification and quantification of sources and sinks of carbon and their temporal and spatial variability on both global and regional scales,
 <sup>5</sup> is a prerequisite for a better understanding of the dynamics of carbon cycle and its

s is a prerequisite for a better understanding of the dynamics of carbon cycle and i response to ever increasing human impact (IPCC, 2001).

Development of reliable carbon cycle models requires adequate observational data. Up to now they are delivered mostly by the existing global and regional monitoring networks (GLOBALVIEW, 2004). Apart from the mixing ratios, also the isotopic composition of carbon dioxide is measured in number of locations. Isotope characteristics of carbon dioxide constitute a powerful tool for identification of sources and sinks of carbon and partitioning of carbon derived from burning of fossil fuels (Levin and Hes-

shaimer, 2000; Démeny and Haszpra, 2002; Kuc et al., 2003).

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The transition to market economy in east European countries is associated with <sup>15</sup> major changes of industrial technologies and resulting shifts in the structure of total energy consumption (gradual shift from coal to oil and natural gas). These changes have significant impact on strength and distribution of major sources of greenhouse gas emissions, in particular CO<sub>2</sub>, into the atmosphere in this part of Europe. Large urban and industrial centres are exposed to these emissions and the extent of their <sup>20</sup> changes over the past years need to be quantified.

This work presents data gathered in southern Poland, mostly during the time period 2000–2004. They illustrate the variability of atmospheric carbon dioxide, both with respect to its mixing ratios and isotopic composition, recorded at two ground-level continental sites located in contrasting environments: (i) high-altitude mountaneous area,

relatively free of anthropogenic influences, and (ii) typical urban environment with numerous local sources of carbon dioxide. The distance between both sites is approximately 100 km.

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### 2. Site description

#### 2.1. Kasprowy Wierch

The Kasprowy Wierch station is located in the south of Poland, within the High Tatra Mountains (Fig. 1a). The meteorological observatory which houses the monitoring
station is located at the top of a mountain peak called Kasprowy Wierch (49°14′ N, 19°59′ E, 1989 m a.s.l., 300 m above the tree line). Kasprowy Wierch is situated at the intersection of three main valleys, at the border between Poland and Slovak Republic. The nearest town, Zakopane, is located approximately 900 m below, 6 km north of Kasprowy Wierch. This is a small tourist town. During winter holiday period (January, February), relatively large amounts of wood and fossil fuels are combusted in the valley.

The Kasprowy Wierch observatory is equipped with electrical heating system and does not use any fossil fuel. During wintertime, diesel-operated snow cars are used in the near-by valleys to maintain proper conditions for skiing.

The climate of Kasprowy Wierch area is typical for a continental mountain location, with relatively large diurnal and seasonal variations of temperature, high precipitation rate, frequent changes of atmospheric pressure and strong winds. The winds are blowing predominantly along the N-S axis, with the average speed of around 7 m s<sup>-1</sup>. Winter season lasts typically for 8 months and ends up rapidly due to strong fohn circulation.

Regular observations of atmospheric CO<sub>2</sub> concentrations started at Kasprowy Wierch in 1994, as a joint project of the AGH University of Science and Technology, Krakow, Poland, and the University of Heidelberg, Germany. Continuous measurements using GC technique were initiated in 1996. The intake of outside air is located ca. 1 m above the roof of the observatory and ca. 6 m above the local ground. The data gathered at Kasprowy Wierch station till 2000 have been summarized by Necki et al. (2003).

Since Kasprowy Wierch is situated within the transition zone between free troposphere and the planetary boundary layer and is relatively free of local influences, this

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site can be considered as a regional reference station for trace gas measurements in the lower atmosphere.

#### 2.2. Krakow

Krakow (50°04' N, 19°55' E, 220 m a.s.l.), the second largest city of Poland, is located approximately 100 km north of Kasprowy Wierch. With more than 800 thousands inhabitants, rapidly growing car traffic and significant industrial activities, it represents a typical urban environment. Moreover, with the prevailing westerly circulation, the Krakow region is under substantial influence of a large coal mining and industrial centre (Upper Silesia) located approximately 60 km west of the city. Characteristic features of the local climate are generally weak winds (annual average around 2.7 m/s) and frequent inversions, extending sometimes over several days (particularly during winter time). These factors favour accumulation of surface gaseous emissions within the lower atmosphere above the city. The sampling site was situated on the University campus (Fig. 1b) located in the western part of the city, bordering recreation and sports grounds. The air intake was located on the roof of the Faculty building, approximately 20 m above the ground level.

#### 3. Analytical methods

The GC technique was employed at both sites to measure the mixing ratios of atmospheric CO<sub>2</sub>. Automated gas chromatograph (Hewlett Packard, Series 5890 with FID <sup>20</sup> detector and Ni catalyst for conversion of CO<sub>2</sub> to CH<sub>4</sub> and Porapak Q column) was used. Reference working standard was injected after each sample. Four determinations of CO<sub>2</sub> mixing ratios per hour were performed in automatic mode of operation. The air was cryogenically dried at  $-70^{\circ}$ C prior to analysis. Typical reproducibility of the mixing ratio measurements was around ±0.1 ppm. All CO<sub>2</sub> data are reported in WMO CO<sub>2</sub> mole fraction scale maintained via secondary standard gases periodically 2, 1849-1865, 2005

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calibrated against primary reference gases.

For isotope measurements the atmospheric  $CO_2$  was continuously sampled by sorption on molecular sieve, in bi-weekly intervals. After thermal desorption, the radiocarbon activity in the given  $CO_2$  sample was measured using benzene synthesis method and liquid scintillation spectrometry (Florkowski et al., 1975). The <sup>13</sup>C/<sup>12</sup>C ratio in the collected  $CO_2$  was determined by mass spectrometry.

Stable isotope composition of the analysed CO<sub>2</sub> samples is expressed in  $\delta$  notation defined as per mille deviation of the carbon isotope ratio measured in the sample (<sup>13</sup>C/<sup>12</sup>C) from that measured in the internationally accepted standard, VPDB (Coplen, 1996). The radiocarbon concentration of the measured samples was expressed as  $\Delta^{14}C=[(A_s/A_{st}-1]\times1000 \text{ where } A_s \text{ and } A_{st} \text{ stand for specific activity of the sample and the standard, respectively. A<sub>st</sub> is corrected for decay since 1950 and normalized to <math>\delta^{13}C=-19.0\%$ . A<sub>s</sub> is normalized to  $\delta^{13}C=-25\%$  (Stuiver and Pollach, 1977). Typical uncertainties of isotope analyses (±1 sigma) are in the order of ±0.1‰ for  $\delta^{13}C$  and  $\frac{15}{2} \pm 8\%$  for  $\Delta^{14}C$ .

#### 4. Results and discussion

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The CO<sub>2</sub> mixing ratios recorded at Kasprowy Wierch and Karkow stations are shown in Figs. 2a and b, respectively. The record for Kasprowy Wierch starts in June 1996 while in Krakow the measurements started only in 2003. Sampling at Kasprowy Wierch site
 was quasi continuous (except of breaks caused by technical problems), in contrary to Krakow where it had a form of campaigns, whenever the GC instrument was available for this type of measurements.

It is apparent from Fig. 2 that these two records of CO<sub>2</sub> mixing ratios differ considerably. The seasonal cycle, with reduced CO<sub>2</sub> mixing ratios during summer months, is well pronounced at Kasprowy Wierch. The presented record reveals typical behaviour of atmospheric CO<sub>2</sub> observed at mid-latitude continental sites of the Northern Hemisphere. The winter maximum ends in March, when the photosynthetic sink starts to 2, 1849–1865, 2005

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operate. From that time on, the  $CO_2$  mixing ratio gradually decreases, reaching a minimum in August/September. Afterwards, the CO<sub>2</sub> level rises, reaching maximum values in January-February. During winter months, the influence of intense fossil fuel combustion in the near-by town Zakopane and its surroundings can be visible in the recorded 5 concentrations of carbon dioxide. The peak-to-peak amplitude of the seasonal changes of the measured CO<sub>2</sub> mixing ratios varies significantly from year to year, the average being around 20 ppm. The measured mixing ratios reveal an increasing trend, from average value around 360 ppm in 1997 to around 380 ppm in 2004. The CO<sub>2</sub> levels recorded in Krakow are generally higher (around 440 ppm in 2004) with much higher maxima of CO<sub>2</sub> content, reaching more than 600 ppm.

The comparison of diurnal CO<sub>2</sub> cycles recorded at Kasprowy Wierch and in Krakow is presented in Fig. 3. Figure 3a shows typical example of diurnal variability of CO<sub>2</sub> mixing ratios, as recorded at both sites during summer time. Both the amplitude and the shape of diurnal CO<sub>2</sub> changes differ considerably. In the urban environment, the lowest CO<sub>2</sub> mixing ratios are recorded generally during mid-day, when the convective

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- activity of the lower atmosphere and resulting vertical mixing is at its maximum. In contrast, at the mountain site high CO<sub>2</sub> mixing ratios are generally recorded during mid-day and early afternoon. This stems from sun-driven convection within the planetary bounday layer over Kasprowy Wierch "sucking" the CO<sub>2</sub>-laden air from the valleys
- towards the top of the mountain. Large differences in the amplitude of diurnal CO<sub>2</sub> 20 changes recorded at both sites are striking (Fig. 3b). In urban environment they may reach even ten times higher values that those recorded at the near-by mountain site. This is connected with presence of numerous local CO<sub>2</sub> sources releasing relatively large amounts of CO<sub>2</sub> into the urban atmosphere, which further leads to elevated con-
- centrations and guasi-permanent excess of this gas above background concentrations, 25 which cannot be completely removed by mixing processes. This is illustrated by Fig. 4 which shows daily mean values of CO<sub>2</sub> mixing ratios recorded at both discussed sites during 2004. It is apparent from the data presented in Fig. 4 that the daily mean  $CO_2$ concentrations recorded at Krakow only seldom are comparable with those recorded

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at the same time at Kasprowy Wierch. At most cases they are significantly higher. The annual average  $CO_2$  concentration in 2004 recorded at Kasprowy Wierch was equal 378 ppm while in Krakow it reached 408 ppm.

- In addition to quasi-continuous observations of CO<sub>2</sub> mixing ratios, also carbon isotope composition of this gas was regularly measured at both monitoring sites. The measurements were performed on bi-weekly cumulative samples. This type of sampling is not adequate for studying short-term (diurnal) variability of the isotopic composition of atmospheric CO<sub>2</sub>.
- Measurements of carbon isotope composition of atmospheric CO<sub>2</sub> have been initiated in Krakow in 1983. The early record of  $\delta^{13}$ C and  $\Delta^{14}$ C of atmospheric CO<sub>2</sub> in Krakow was published by Kuc (1991) and later analysed by Kuc and Zimnoch (1998). A dedicated study focused daily variations of CO<sub>2</sub> mixing ratio and its stable isotope composition ( $\delta^{13}$ C,  $\delta^{18}$ O) within the urban atmosphere of Krakow was carried out by Zimnoch et al. (2004).
- Figure 5 shows variations of <sup>14</sup>C and <sup>13</sup>C content in bi-weekly cumulative samples of atmospheric CO<sub>2</sub> collected at Krakow and Kasprowy Wierch during the four-year period (2000–2003). The apparent offset of the  $\delta^{13}$ C record for Krakow with respect to Kasprowy Wierch data in the order of 1.2‰ (Fig. 5b), is attributed to local emissions of CO<sub>2</sub> originating mostly from burning of fossil fuels. The offset of Krakow  $\delta^{13}$ C record with respect to maritime reference station Mace Head is even higher (around
- 2% GLOBALVIEW, 2004). Upper envelope of the two records presented in Fig. 5b reveals a small increasing trend towards less negative  $\delta^{13}$ C values, pointing to diminishing burden of regional atmosphere with CO<sub>2</sub> of anthropogenic origin. The Kasprowy Wierch  $\delta^{13}$ C record approaches the values reported for "clean" reference sites (GLOB-
- <sup>25</sup> ALVIEW, 2004; Levin and Kromer, 1997). Seasonal fluctuations of  $\delta^{13}$ C visible at both discussed sites are shifted in phase. The Krakow record reveals lowest  $\delta^{13}$ C values during winter season, when local CO<sub>2</sub> emissions due to burning of fossil fuels in the city (heating plus car traffic) are most intense. At Kasprowy Wierch, less negative  $\delta^{13}$ C values are generally recorded during winter time, when advection of biospheric CO<sub>2</sub>

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from the valleys beneath the station is cut-off and free tropospheric air is "seen" by the station more frequently.

Substantial fluxes of CO<sub>2</sub> of fossil-fuel origin lead to generally lower <sup>14</sup>C content in atmospheric CO<sub>2</sub> recorded in urban environment, when compared to sites characterized by smaller anthropopression. This effect is clearly seen in Fig. 5a which shows records of atmospheric <sup>14</sup>CO<sub>2</sub> concentration available for Krakow and Kasprowy Wierch. The Krakow  $\Delta^{14}$ C record reveals distinct seasonality, with local minima of  $\Delta^{14}$ C occuring during winter months, when combustion of fossil fuels is most intense. This is not the case for Kasprowy Wierch where recorded fluctuations of  $\Delta^{14}$ C are much less pronounced and irregular. However, a distinct decreasing trend of <sup>14</sup>CO<sub>2</sub> concentration is visible at both discussed sites and the difference between mean  $\Delta^{14}$ C values recorded for Krakow and Kasprowy Wierch is smaller from year to year, pointing to diminishing fluxes of radiocarbon-free CO<sub>2</sub> in Krakow. This is confirmed by energy-usage statistics for Krakow which indicate gradual reduction of coal burning in the municipal heating

15 systems and reduction of low emission.

#### 5. Conclusions

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The presented work was focused on characterization of spatial and temporal variability of atmospheric concentrations of carbon dioxide and its isotopic composition, as recorded at two ground-level continental sites located in contrasting environments: (i) high-altitude mountaneous area, relatively free of anthropogenic influences, and (ii) typical urban environment with numerous local sources of carbon dioxide. Despite of relative proximity of those sites, substantial differences in both the recorded  $CO_2$  levels and their isotopic composition were detected. Human influence is clearly visible in the presented data, affecting both short- and long term changes of  $CO_2$  concentrations and its isotopic composition.

The CO<sub>2</sub> mixing ratios measured in the urban atmosphere revealed quasi-permanent excess concentration of this gas when compared with near-by background atmosphere.

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The magnitude of this surplus is significant; in the studied case the annual mean  $CO_2$  concentration recorded in Krakow in 2004 was almost 10% higher than that recorded at high-altitude mountain site (Kasprowy Wierch). Such effect is occuring probably in all urban centers. For large cities with several million inhabitants this permanent surpluss

- of CO<sub>2</sub> may be significantly higher than that observed in this study. The elevated concentrations of CO<sub>2</sub> are expected near the ground where people live and the biosphere is present. Moreover, one may argue that local climate in urban centers will be further modified by the presence of these elevated CO<sub>2</sub> concentrations enhancing the effect of urban heat islands.
- <sup>10</sup> Carbon isotopic composition of atmospheric  $CO_2$  proved to be an efficient tool for identification the surface  $CO_2$  fluxes into the atmosphere related to fossil fuel burning, their strength and temporal variability, as well as their influence on the recorded levels of this gas in the local atmosphere. The presented data clearly show that this influence can be monitored on long term basis. The available records of  $\Delta^{14}C$  for Krakow <sup>15</sup> and Kasprowy Wierch suggest gradual reduction of <sup>14</sup>C-free  $CO_2$  fluxes into the urban atmosphere of Krakow in the past several years.

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**Fig. 1.** Aerial view of the CO<sub>2</sub> monitoring sites discussed in the paper: **(a)** high-altitude mountain site Kasprowy Wierch, Tatra Mountains; **(b)** city of Krakow.

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**Fig. 4.** Daily mean values of atmospheric  $CO_2$  mixing ratio recorded during 2004 at Kasprowy Wierch and Krakow stations.



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**Fig. 5.** Changes of carbon isotopic composition of atmospheric CO<sub>2</sub> recorded at Kasprowy Wierch and Krakow stations during the time period 2000–2003. The data points represent bi-weekly cumulative samples. The lines are **(a)** changes of radiocarbon activity ( $\Delta^{14}$ C), **(b)** changes of <sup>13</sup>C content ( $\delta^{13}$ C). The smoothed curves were obtained using program recommended by NOAA/CMDL (CCGvu, ver. 4.40) (Thoning et al., 1989).