

Interactive comment on “Effects of multiple environmental factors on CO₂ emission and CH₄ uptake from old-growth forest soils” by H. Fang et al.

H. Fang et al.

huajunfang@yahoo.com.cn

Received and published: 12 October 2009

Thanks for the referee's comments concerning our manuscript entitled “Effects of multiple environmental factors on CO₂ emission and CH₄ uptake from old-growth forest soils”. We found the referee's comments most helpful and constructive, and have made some correction which we hope meet with his/her approval. The revised portions are underlined in red and added in the revised manuscript.

1. In the abstract, the authors mentioned that soil CO₂ flux in the old-growth forests were mainly driven by soil temperature (P7822, L12). I agreed with the opinion. But, I can't agree with the speculation that CO₂ fluxes will increase with increase in air tem-

C2471

perature (P7835, L5). This speculation is a just speculation, not scientific. This study was conducted in various climate conditions from boreal to tropical zone. When air temperature will be changed, the response of soil temperature will be different among zone. Not only scientist but also many people are concerned with an increase in carbon emission from soils due to an increase in air temperature. Therefore, the speculation should be deleted.

Our initial conclusion is not exact. Presently, both warming manipulative experiments and models suggest that there are conflicting conclusions on soil CO₂ emission to warming including acceleration and acclimation (Tian et al., 1998; Mellilo et al., 2002) depending on C substrate availability. In the warming experiment of Harvard forest, it took about ten years to observe the decrease of soil CO₂ emission, which was mainly attributed to limited size of the labile soil carbon pool (Mellilo et al., 2002). Contrast to warming experiment, the increase of ambient air temperature will not rapidly exhaust soil C substrates such as labile C pools within a short period regardless of temperate or tropical forests (Tang et al., 2006; Schwendenmann and Veldkamp, 2006).

Our results showed that soil CO₂ efflux increase with soil temperature. Moreover, the temperature sensitivity of soil CO₂ emission in the boreal and temperate forests is significantly higher than those in the subtropical and tropical forests. In the future, if boreal and tropical forests would be turned into temperate and subtropical forests due to climate warming, respectively, soil CO₂ emission will increase within a short period. Therefore, the original sentence was substituted by “Based on the gradient theory of exchange of time and space, increase in air temperature in the future would promote soil CO₂ emission in the old-growth forests within a short period, and northern forest soils were more sensitive than southern forest soils”.

Tian H Q, Melillo J M, Kicklighter D W, et al. 1998. Effect of interannual climate variability on carbon storage in Amazonian ecosystems. *Nature*, 396, 664-667
Mellion J M, Susan E T, Ronald A. 2002. Soil warming and carbon cycle feedback to the climate system. *Science*, 298: 2173-2176

C2472

Tang X L, Liu S G, Zhou G Y, Zhang D Q, Zhou C Y. 2006. Soil-atmospheric exchange of CO₂, CH₄, and N₂O in three subtropical forest ecosystems in southern China, *Global Change Biol.*, 12, 546-560.

Schwendenmann L, Veldkamp E. 2006. Long-term CO₂ production from deeply weathered soils of a tropical rain forest: Evidence for a potential positive feedback to climate warming. *Global Change Biology*, 12, 1878-1893.

2. I agreed with the result that there were relationships between soil C fluxes and mineral N (Fig. 6). However, I can't agree the authors' opinion that increasing in N deposition in eastern China would increase soil CO₂ emission but decrease soil CH₄ uptake in the old-growth forests (P7822, L21). In this study, CH₄ uptake was highest in subtropical forest (Table 2), where N deposition was highest (Table 1). How can you explain this result? In addition, it is necessary more explanation to assess the effect of mineral N on soil C flux in the discussion (4.3&4.4). The authors mentioned that NO₃ promote CO₂ emission and NH₄ inhibit CH₄ uptake (P7822, L16). In fact, figure 6 showed the clear relationship between soil C flux and mineral N. However, NH₄⁺ is consumed and NO₃ is produced via nitrification process, and NO₃ is consumed via denitrification process. Generally, nitrification rate is high under high temperature and mesic condition. These conditions would enhance a decomposition of soil organism and CO₂ emission from soil. The authors need to explain the considerable mechanism. As for CH₄, the authors mentioned CH₄ uptake was inhibited due to both the competitive and toxic inhibition (P7822, L19). But, in the subtropical forest soil, it is an acceptable explanation that high CH₄ uptake was observed due to that NH₄ oxidizers can oxidize CH₄. Anyway, the effect of mineral N on soil C flux should be discussed based on various possibilities.

We completely agree with the referee's opinions. Indeed, our results can only deduce the effect of soil N availability rather than nitrogen deposition on soil C effluxes. Warming, precipitation variation and N deposition can all result in change of soil mineral N storage along the environmental gradient. Because of covariation among these envi-

C2473

ronmental variables, we can not attribute variation of soil mineral N to one of them. Therefore, we revised the original sentence as follows. Soil NO₃-N increase and NH₄⁺-N decrease resulting from environmental change would increase soil CO₂ emission and CH₄ uptake.

Due to relative high N deposition rate (more than 30kgN ha⁻¹ yr⁻¹) and the climax of forest succession, the subtropical monsoon evergreen broadleaf forest at Dinghushan site has already reached N saturation (Mo et al., 2008), where soil mineral N is dominated by soil NO₃-N (Fig. 3) and soil N leaching and gaseous emission is very high (Fang et al., 2008, 2009; Zhang et al., 2008). However, the tropical seasonal rain forest in Xishuangbanna where economy is undeveloped and N deposition is low, soil mineral N is still dominated by soil NH₄⁺-N and is not unsaturated yet (Fig. 3). Because NH₄⁺-N rather than NO₃-N has the inhibitory effect on CH₄ production, the subtropical forest soil has the highest CH₄ uptake capacity. In addition, although the mean annual precipitation is the highest at subtropical forest site, the soil moisture content is the lowest in vegetative season from May to October (Table 1 and Table 2). Forest soil CH₄ uptake is generally driven by soil moisture or soil WFPS (water-filled pore space), which also result in subtropical forest soil has the highest capacity of CH₄ consumption.

If forest types were not taken into account, positive relationships between soil CO₂ efflux and soil NO₃-N as well as between soil CH₄ efflux and soil NH₄⁺-N were found (Fig.6 and Table 3). So we concluded that NO₃- can promote CO₂ emission and NH₄⁺ inhibit CH₄ uptake. Besides, there are significant difference for relationship between soil C effluxes and mineral N content among four forests. According the referee's advices, we added many explanations on these differences in discussion section.

(1) Nitrification rate and NO₃-N content are higher in the tropical and subtropical forests than in the boreal and temperate forest soils. NO₃-N accumulation will decrease C/N ratio of litter, debris and soil organic matter. Moreover, decomposition of soil organism and CO₂ emission from soil would be enhanced under high temperature and mesic conditions. On the other hand, soil CO₂ efflux increase with soil

C2474

NO₃-N, which could be partly explained by the difference in fine-root biomass. Forest with higher soil NO₃-N content has higher fine-root biomass (Table 1). Cleveland and Townsend (2006) also suggested that the increase in soil respiration in the N-fertilized plots may have been driven, at least in part, by changes in fine-root dynamics. However, inconsistent effects of N on soil CO₂ effluxes (decrease or no change) generally occur in the laboratory and field-fertilization experiments, which are mainly attributed to rapidly decline of soil microbial activities and decrease of root-biomass within a short period (three to five years)(Mo et al., 2008).

(2) Soil CH₄ uptake is controlled by soil temperature, moisture, soil mineral N, thickness of O-horizon etc. Contrast to the tropical and subtropical forests, soil NH₄⁺ in the boreal and temperate forests was mainly assimilated by plants or immobilized by soil clay mineral, the relatively lower CH₄ uptake was attribute to other reasons such as substrate availability or diffusion rather than NH₄⁺ inhibition. For example, low temperature and frozen layer in winter as well as the thicker O-horizon obstruct the diffusion of CH₄ in atmosphere to soil, which will indirectly decrease the uptake of soil CH₄ (Grosso et al. 2000). However, in the tropical and subtropical forest soils, the observed higher CH₄ uptake should be mainly attribute to the competitive and toxic inhibition of NH₄⁺.

Cleveland CC, Townsend AR (2006) Nutrient additions to a tropical rain forest drive substantial soil carbon dioxide losses to the atmosphere. PNAS, 103: 10316-10321

Fang YT, Gundersen P, Mo JM, Zhu WX (2008) Input and output of dissolved organic and inorganic nitrogen in subtropical forests of South China under high air pollution. Biogeosciences 5: 339-352

Fang YT, Zhu WX, Gundersen P, Mo JM, Zhou GY, Yoh M (2009b) Large Loss of Dissolved Organic Nitrogen from Nitrogen-Saturated Forests in Subtropical China. Ecosystems 12: 33-45

Grosso S J D, Parton W J, Mosier A R, et al. (2000) General CH₄ oxidation model and comparisons of CH₄ oxidation in natural and managed systems. Global Biogeochemi-

C2475

cal Cycles, 14:999~1019

Mo J, Zhang W, Zhu W, Gundersen P, Fang Y, Li D, Wang H (2008) Nitrogen addition reduces soil respiration in a mature tropical forest in southern China. Global Change Biology 14: 403-412

Zhang W, Mo JM, Yu GR, Fang YT, Li DJ, Lu XK, Wang H (2008a) Emissions of nitrous oxide from three tropical forests in Southern China in response to simulated nitrogen deposition. Plant and Soil 306: 221-236

Specific comments P7824, L17: The study sites were classified in boreal, temperate, subtropical, and tropical zone. However, I think the temperate site should be classified cool temperate zone, not temperate forest. Because mean annual temperature and precipitation (Table 1) are too low to be classified into temperate forest. In addition, soil was relatively wet due to the soil moisture data and vegetation (*Fraxinus mandshurica*), so that, this forest was not ordinal temperate forest. As same as temperate region, according to figure 1, the tropical region is too cool to regard as the tropical. Are the classification correct?

Based on Chinese climate and vegetation classification (mainly considering mean annual temperature, extreme temperature, annual accumulated temperature etc.), mean annual temperature of cool temperate, temperate, subtropical and tropical zone ranges from -2.2~-5.5oC, 2~8oC, 14~22 oC and 22~26 oC, respectively. Annual accumulated temperature ranges from 1100-1700oC, 1600-3200 oC, 4500-8000 oC and 8000-9000 oC, respectively. Therefore, the Daxinganling Mountains belong to cool temperate zone. Its vegetation is Chinese Larch (*Larix Gmelini*) ecosystem and is the southern margin of Taiga forest regions, which is also involved in boreal forest. The Changbai Mountain belongs to temperate zone. Its vegetation is a typical temperate needle-broadleaved mixed forest. Moreover, Manchurian ash (*Fraxinus mandshurica*) is a frequently constructive species in the temperate mixed forest. Dinghushan site belongs to southern subtropical climate zone, and its vegetation is monsoon evergreen broad leaved forest.

C2476

Xishuangbanna site is within the south of tropic of cancer, influencing by tropical monsoon climate. The vegetation is typical tropic seasonal rain forest. Due to higher altitude of Yunnan-Guizhou plateau, the mean annual temperature is slightly less than the lower bound of classification indexes. After comprehensive consideration of geomorphology, climate, vegetation, Xishuangbanna site customarily is ascribed to tropical forest.

P7825, L8: The authors mentioned soil classification, but the classification was very old. Soil classification was very important information. If the authors want to refer FAO/UNESCO taxonomy, soil classification should be conducted due to WRB 2006.

Comparing Chinese soil Taxonomy (Gong et al., 2007) and World reference base for soil resources (2006), we carefully compared and classified the four forest soils again. The soils are greyzems, luvisols, ferralsols and lixisols (IUSS Working Group WRB, 2006.) from north to south, respectively.

Zitong Gong, Ganlin Zhang, Zhicheng Cheng, et al. 2007. Pedogenesis and soil Taxonomy. Science press, Beijing.

IUSS Working Group WRB. 2006. World reference base for soil resources 2006. 2nd edition. World Soil Resources Reports No. 103. FAO, Rome.

P7826, L7: The measurement season was separated growing and non-growing season. But I think that the separation is not suitable for subtropical and tropical region. If the classification of subtropical and tropical region was incorrect, the authors can use growing and non-growing season, or summer and winter.

The boreal and temperate forest ecosystems are two northern forests with obvious growing and non-growing seasons, while subtropical and tropical forests are two evergreen forests, where there are no rigid division between growing season and non-growing season. However, there exists obvious cool-dry season from November to April and warm-humid season from May to October. According to season separation

C2477

by Yu et al. (2008), we consider the vegetative season from May to October and non-vegetative season from November to April for soil C effluxes comparison among the four forest ecosystems.

Yu GR, Song X, Wang QF, et al. Water use efficiency of forest ecosystems in eastern China and its relations to climatic variables. New Phytologist, 2008, doi: 10.1111/j.1469-8137.2007. 02316.x

Technical comments CO₂ and CH₄ flux: The authors use various words to express CO₂ and CH₄ flux. For example, CO₂ emission, CO₂ exchange, soil CO₂ flux, CO₂ efflux etc. In discussion, sometimes, it is correct to use CO₂ emission or CH₄ uptake, but the authors should take care of how to use the terms.

According the referee's suggestion, most of expressions about CO₂ and CH₄ fluxes were unified for soil CO₂ and CH₄ effluxes, except in discussion section where CO₂ emission and CH₄ uptake were used sometimes.

Table 1: Temperature mixed forest > temperate mixed forest. It is better to write forest vegetation name not Latin name, briefly. For example, larch in boreal forest. What is the measurement depth or range of soil texture, Total N, C/N and soil pH? Please write down as SOC density like as "0-20 cm".

"Temperature mixed forest" was replaced with "Temperate mixed forest". In each forest site, typical trees were expressed as English name and Latin name. For example, Chinese larch (*Larix gmelinii*) in boreal forest, Korean pine (*Pinus koraiensis*), basswood (*Tilia amurensis*), Manchurian ash (*Fraxinus mandshurica*) and oak (*Quercus mortgolicola*) in the temperate forest, guger-tree (*Schima superba*), rose apple (*Syzygium jambos*), henry chinkapin (*Castanopsis chinensis*) in the subtropical forest, and downy malugay (*Pometia tomentosa*), bayberry waxmyrtle-fruit (*Terminalia myriocarpa*), Yunan nutmeg (*Myristica yunnanensis*), South-Yunnan horsfieldia (*Horsfieldia tetratapa*), glabrous Homalium (*Homalium laoticum*) in the tropical forest.

C2478

All of soil properties showed in Table 1 are 0-20 cm soil layer, so the title of Table 1 is changed into "Stand characteristics and surface soil (0–20 cm) properties of the four old-growth forest sites". In addition, some soil properties in Table 1 were added standard error to express their range.

Table 2 & Fig. 2: Unit of soil moisture is not "%" but "m³ m⁻³". Please use SI unit. Reference: the authors cited "Hashimoto et al. 2004" (P7831, L13), but it is not showed in the reference list. Add in the list or delete from the text.

The unit of soil moisture has now been corrected in Table 2 and Fig. 2 replacing "%" with "m³ m⁻³". In addition, the lost reference was added in the list. Hashimoto, S., Tanaka, N., Suzuki, M., Inoue, A., Takizawa, H., Kosaka, I., Tanaka, K., Tantasirin C., and Tangtham, N.: Soil respiration and soil CO₂ concentration in a tropical forest, Thailand, Jap. Forest Res., 9, 75–79, 2004.

Please also note the Supplement to this comment.

Interactive comment on Biogeosciences Discuss., 6, 7821, 2009.