

**Seasonal and
latitudinal patterns of
pelagic community
metabolism**

S. Agustí

Seasonal and latitudinal patterns of pelagic community metabolism in surface waters of the Atlantic Ocean

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Abstract

Temporal and spatial patterns in the variability of the pelagic metabolism at the surface of the Atlantic Ocean were analyzed in a series of four oceanographic cruises (LATITUDE 1, 2, 3 and 4). The cruises crossed the oligotrophic waters of North and South subtropical gyres and this explained the low values of both gross primary production (GPP) and community respiration (*R*) found. Net community production (NCP), the balance between production and consumption, was strongly related to the variability in *R* rates ($R^2 = 0.72$, $P < 0.0001$). NCP was net heterotrophic in 83% of the data, but showed strong temporal and spatial patterns. At the inter-tropical zone, around 10°–12° N and 10°–12° S, a large variability was observed with values of NCP oscillating from net heterotrophic to net autotrophic seasonally. This variability implied NCP to be net autotrophic in boreal fall and austral spring, and net heterotrophic in boreal spring and austral fall, in the areas around the boundaries of the inter-tropical zone. The variability observed concur with the seasonal climatic and oceanographic regimes of the inter-tropical area, which documented seasonal changes of the North and South Atlantic equatorial currents system, the Guinea Dome, and the Benguela current. When considering the season of the data obtained, significant differences between spring and fall were found for the surface Atlantic, with water temperature and respiration increasing in autumn, showing a net heterotrophic metabolism, and with temperature and respiration decreasing in spring, where NCP were closer to the metabolic balance. In contrast, no seasonal differences were found for GPP and chlorophyll-*a* concentration. The results showed new spatial and temporal patterns in the pelagic metabolic balance of the surface Atlantic Ocean with consequences for the carbon flux.

1 Introduction

The importance of oceanic biota in the net absorption of atmospheric CO₂ depends on the net community production, i.e. on the balance between the carbon fixed by primary

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production and the consumption of organic carbon by the community by respiration. When production exceeds respiration, the net metabolism of the system is considered to be net autotrophic, contrary to the net heterotrophic metabolism that characterize a system where consumption exceeds production. The net metabolism of a system is an important descriptor of the role of the biological processes in the carbon flow (Smith and Hollibaugh, 1993; Duarte and Agustí, 1998; Williams, 1998). Besides its importance, our knowledge on the net metabolism of the open ocean is remarkably small especially when compared with the accumulation of data on other parameters, as primary production, a single component of the pelagic metabolism (e.g. Duarte and Regaudie de Gioux, 2009; Williams and del Giorgio, 2005; del Giorgio and Duarte, 2002). This paucity is even larger for the open oligotrophic ocean (del Giorgio and Duarte, 2002; Duarte and Regaudie de Gioux, 2009), for which we have in general less information.

Consumption should exceed respiration in unproductive areas of the ocean (Agustí and Duarte; 1998; Agustí et al., 2004; Duarte et al., 2001; Williams and Robinson 2005; Duarte and Regaudie de Gioux, 2009) but geographical and temporal variability of physical, chemical, and biological processes should generate different regional or seasonal variability in the net pelagic metabolism of the open ocean (e.g. Serret et al., 2002). Variability in the net metabolism of plankton generated by mesoscale processes has been identified in the oligotrophic North Atlantic (Gonzalez et al., 2001; Mouriño-Carballido and McGillicuddy 2006). For coastal waters there are reports describing patterns at the temporal scale (e.g. Duarte et al., 2004; Gonzalez et al., 2008); however, temporal variability of oceanic pelagic net metabolism is still poorly explored and we lack description of the patterns of its variability in the oceanic areas (e.g. Gonzalez et al., 2002).

The goal of this study is to analyze the spatial and temporal variability on the metabolic balance of the pelagic communities of the Atlantic Ocean by quantifying the net community production and the community respiration at the surface ocean in a series of oceanographic cruises (LATITUDE cruises) that crossed the Atlantic Ocean

from the North Atlantic (from the Canary Islands) to the South Atlantic (to the latitude of Buenos Aires or Montevideo). The measurements were performed on 4 cruises: two of them (LATITUDE-1 and LATITUDE-4) were conducted in March–April in years 1995 and 2000, respectively; and the other two (LATITUDE-2 and LATITUDE-3) were performed in October–November, in years 1995 and 1999, respectively. The sampling stations were located within the same meridional regions for the four cruises, thus helping the analysis of the temporal variability in the data by reducing inter-cruise spatial variability. The present study provides new data from both North and South Atlantic subtropical Gyres, contributing to increasing our understanding of the biological metabolism in the low productive areas of the ocean.

2 Methods

The study was performed during the cruises LATITUDE-1, 2, 3 and 4 across the Atlantic Ocean (Fig. 1). The LATITUDE-cruises comprised the sampling along transects from/to the South Atlantic, at latitudes of Uruguay-South Brazil, to/from the North Atlantic close to the Canary Archipelago area. LATITUDE-1 ran from March to April 1995, LATITUDE-2 from October to November 1995, and LATITUDE-3 and 4 were made during November 1999 and March–April 2000, respectively. LATITUDE-2 included the sampling across the NW African upwelling in the most north eastern area visited (Fig. 1). Sampling stations followed the 29° W Meridian from 20° S to 30° N during LATITUDE-1 (Fig. 1).

Seawater was sampled using Niskin bottles attached to a CTD-Rosette system. Subsamples of 250 ml were filtered through Whatman GF/F filters for fluorometric analysis of chlorophyll-*a* concentration as described in Agustí and Duarte (1999). Surface water (5 m) was gently siphoned into narrow mouth 125 ml replicated winkler bottles for the quantification of initial oxygen concentration, net community production (clear bottles) and community respiration (dark bottles), with 6–8 replicates each, depending on the station. Samples from LATITUDE-2, 3 and 4 were incubated for 24 h on deck, receiving screened (50 % reduction by a neutral net) natural light and submerged in

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a large incubator (i.e. 1000 L) refrigerated with running surface seawater. Samples from LATITUDE-I were placed in an incubator and maintained for 21–24 h at surface seawater temperature illuminated with artificial light ($500 \mu\text{mol photons m}^{-2} \text{s}^{-1}$).

Dissolved oxygen concentration was measured using high-precision Winkler titration after Carrit & Carpenter (1966), using a Mettler-DL21 Autotitrator for the potentiometric (redox electrode) end-point detection (Oudot et al., 1988). Respiration rates were calculated from the difference in oxygen concentration between initial and “dark” bottles, and net community production was calculated from the difference in oxygen concentration between “light” bottles and the initial oxygen concentration. Gross primary production was calculated as the sum of community respiration rates and net community production. In LATITUDE-1 samples were incubated under continuous light, and hourly net community production measured was converted to daily (24 h) estimates by considering the length of the day and night periods for each sampling event.

3 Results

Values of net community production (NCP), community respiration (R) and gross primary production (GPP) observed for the surface of the Atlantic Ocean were generally low, although the variability found was high (Table 1). Community respiration values ranged from 0.12 to $7.33 \mu\text{mol O}_2 \text{l}^{-1} \text{d}^{-1}$ and NCP varied from -6.87 to $3.65 \mu\text{mol O}_2 \text{l}^{-1} \text{d}^{-1}$. Values of GPP at the surface of the central Atlantic Ocean ranged from 0.03 to $3.91 \mu\text{mol O}_2 \text{l}^{-1} \text{d}^{-1}$ (Table 1). NCP values were mostly net heterotrophic along the study, as indicated by the high percentage of negative values observed (86%) indicative of a higher consumption over production in the pelagic community of the surface of the Atlantic Ocean. NCP was negatively and strongly related to community respiration ($\text{NCP} = -0.89 \times R + 0.71$, $R^2 = 0.72$, $P < 0.0001$, Fig. 2), but no relationship was found with other variables as chlorophyll- a concentration and GPP. GPP and R were not related ($R^2 = 0.059$). The ratio between GPP and R (P/R ratio) was lower than 1 in 83% of the samples, indicative of a dominance of net

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heterotrophic metabolism in the waters examined. P/R was positively correlated with GPP ($R^2 = 0.51$, $\log P/R = -0.188 + 1.14 \log \text{GPP}$, $P < 0.0001$, model II regression).

GPP varied with the latitude, showing increased values at the tropical zone between 10–20° in both hemispheres (Fig. 3), decreasing with increasing latitude at the temperate area and decreasing again at the equatorial zone (Fig. 3). This pattern was also observed in the distribution of data from individual cruises (Fig. 3). The variability of GPP with latitude differed from that observed for chlorophyll-*a* (Fig. 4), which showed lower values at the South Atlantic, with the minimum values around 10° S among the transects; and increasing at the North Atlantic, with maximum chlorophyll-*a* values around the area influenced by the equatorial upwelling (Fig. 4). Data of community respiration and NCP did not show a pattern with latitude when all data were analyzed; however, they followed a latitudinal variation when the data were analyzed by season. At spring, NCP at the North Atlantic showed negative values at the tropical area, with most negative values found around 10° N; NCP increased in the Southern Hemisphere in austral spring to reach positive values indicative of autotrophic net metabolism around 10° S (Fig. 5). This tendency was inverted in autumn, when negative values of NCP were observed at the South Atlantic, with the more negative values observed around 10° S, although NCP values increased around 10° N during the boreal autumn until reaching positive, net autotrophic values (Fig. 4). The seasonal changes observed resulted in highly significant differences between spring and fall (in the Northern and Southern Hemispheres, Table 2) for averaged values of NCP and R, showing more heterotrophic waters and significantly higher respiration rates and higher temperature during the autumn (Table 2). There were no significant relationships between values of net community metabolism and community respiration with water temperature. For the surface waters of the Atlantic Ocean studied, Chl-*a* and GPP showed not seasonal differences between spring and fall (Table 2).

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4 Discussion

20 The results presented here showed novel latitudinal and temporal patterns in the planktonic metabolism of the surface waters of the Atlantic Ocean. These patterns were identified despite the predominantly low values of productivity observed during the study. This low productivity was expected since the area included the sampling in the ultra oligotrophic waters of both North and South Atlantic subtropical gyres
25 (e.g. Duarte and Agustí, 1998; Williams, 1998; Duarte et al., 2001; Gonzalez et al., 2002; Serret et al., 2006). In 83% of the samples analyzed, community respiration exceeded gross primary production, indicating that the metabolism of the surface of the Atlantic Ocean was, predominantly, net heterotrophic. The gross production required to render the communities net autotrophic was $1,46 \mu\text{mol O}_2 \text{ l}^{-1} \text{ d}^{-1}$ (equivalent to $0,046 \text{ g O}_2 \text{ m}^{-3} \text{ d}^{-1}$), which was very close to the threshold value calculated
5 by Duarte and Agustí (1998) for aquatic ecosystems, and the value reported recently by Duarte and Regaudie de Gioux (2009), for a larger data set for marine waters.

Moreover, the low production characterizing the Atlantic Ocean and the dominance of a net heterotrophic metabolism at the surface waters indicated clear patterns in the metabolic balance, with communities showing changes from net heterotrophic to net autotrophic at the seasonal and latitudinal scales. The temporal patterns observed here at the surface waters has not been described before, since most studies analyzing the variability of the net metabolic balance in the Atlantic Ocean have focused on the variability across biogeographical regions and analyzed the integrated water column metabolism (e.g. Robinson et al., 2002; Gonzalez et al., 2002; Serret et al., 2002, 2006). Gonzalez et al. (2002) identified however the temporal variability in a year scale across the Atlantic, describing larger variability in the temperate areas of both S and N Atlantic, associated to its seasonality. Here, variability was also observed
10 in the temperate waters sampled, as described in other studies (e.g. Gonzalez et al., 2002), but the variability in the intertropical boundaries was identified here because the study included inter-annual data, and a more detailed sampling spatial resolution than
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previous studies (e.g. Robinson et al., 2002; Serret et al., 2002, 2006; Gonzalez et al., 2002), as the effort was done at a single layer.

The larger variability in the metabolic balance was observed at the margins of the inter-tropical zone, probably associated to the boundaries of the subtropical Gyres and its seasonal oscillation (e.g. Li and Philander, 1997; Chiang et al., 2002). At the North Atlantic, the major oscillation of the net metabolism was observed at around 10°–12° N, with the pelagic metabolism reaching net autotrophic values in boreal autumn, but showing strong net heterotrophic values during the boreal spring. This temporal variability should be related to the seasonal cycle of the equatorial Atlantic area that showed the highest and lowest temperatures in March and September (Li and Philander, 1997). The boundaries between the Gyres and the equatorial area are driven by the complex system of currents of the equatorial zone, whose dynamics also change seasonally. In the North Atlantic, the North Equatorial Current extends from approximately 16°–12° N with a uniform flow towards the south and west (e.g. Richardson and Walsh, 1986; Peterson and Stramma, 1991; Hooke et al., 2000). Southernly, the Guinea dome province should extend from 12°–8° N, and represents a highly seasonal mid-ocean upwelling, showing the main period of upwelling between July and September (e.g. Aiken et al., 2000). This seasonality described for the North Atlantic equatorial area was reflected in the net community metabolism measured here, which showed increased autotrophic metabolism during the boreal fall cruises, reflecting most probably the influence of the summer upwelling. A similar but opposite trend was observed at the South Atlantic where the net metabolism showed the highest negative values during the austral fall, but increased to reach a net autotrophic balance during the austral spring (i.e. at the boreal fall). At the South Atlantic, the equatorial province extends from 16° N–15° S and is also highly dynamic influenced by the South Equatorial Current (SEC), which is strongest during the austral winter (e.g. Krauss, 1986; Peterson and Stramma, 1991). The southern branch of the westward flowing SEC is fed by the Benguela Current, which flows through a strong, biologically productive upwelling region, and advects cool waters to the tropics with high speeds during the austral winter

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(Richardson and Walsh, 1986; Johns et al., 1998).

25 The seasonality and latitudinal variability found here for the net metabolism in the Atlantic surface was larger than that observed for chlorophyll-*a* concentration, that, despite displaying seasonal variability as well, always maintained higher values at the tropical North Atlantic as a result of being highly influenced by the guinea and equatorial upwelling systems (e.g. Aiken et al., 2002). Community respiration strongly explained the changes in NCP values in this study, indicating the importance of the heterotrophic metabolism in determining NCP. Heterotrophic organisms, such as zooplankton or bacteria, in the Atlantic Ocean are dependent on chlorophyll-*a* concentration but also on the variability of other processes. For example, both Chl-*a* and water temperature explained zooplankton egg production in the Atlantic Ocean, resulting in higher reproductive capacity at the equatorial area where both values increased (Calbet and Agustí, 1999). Low variability in Chl-*a* but high in zooplankton biomass was described by Finenko et al. (2003) for the Atlantic inter-equatorial zone, showing seasonal changes in the Chl-*a* to zooplankton biomass ratio. The lower Chl-*a* to zooplankton biomass ratio in this area was observed for winter–spring due to higher zooplankton biomass during this period (Finenko et al., 2003), which is in accordance with higher net-heterotrophic balance found here in the March–April cruises. Hoppe et al. (2002) identified the heterotrophic component of pelagic communities (i.e. heterotrophic bacteria) as key for the prediction of the metabolic balance of the Atlantic Ocean, and identified changes in temperature to strongly influence bacteria (Hoppe et al., 2002). Water temperature, considered to exert strong control of community respiration (López-Urrutia et al., 2006), varied significantly between spring and fall in the Atlantic surface waters examined. Regaudie-de-Gioux (2010), for a large data set of oceanic pelagic metabolism, calculated that P/R equals 1 at an average temperature of 23.5°C. Surface Atlantic waters in spring were closer to the metabolic balance than during the fall, and showed a lower averaged temperature of 24.13 ± 0.34 , close to the value calculated by Regaudie-de-Gioux (2010). This suggests that temperature is a proxy for the oceanographic and climatic changes influencing the metabolism of the

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surface Atlantic Ocean. The results support the idea that understanding and predicting the net metabolic balance of the oceans would require a better knowledge of the oceanographic and climatic processes influencing the heterotrophic component of the plankton communities (Hoppe et al., 2002; Serret et al., 2006).

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Table 1. Mean values (\pm SE) of Net Community Production (NCP), Community Respiration (R) and Gross Primary Production (GPP) observed in surface waters of the South and North Atlantic in the different LATITUDE cruises.

	N	NCP ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)	R ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)	GPP ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)
LATITUDE-1 (1995)	34	-1.78 ± 0.25	2.28 ± 0.25	0.50 ± 0.06
LATITUDE-2 (1995)	10	-0.89 ± 0.39	2.09 ± 0.31	1.61 ± 0.27
LATITUDE-3 (1999)	13	0.27 ± 0.40	1.32 ± 0.36	1.59 ± 0.28
LATITUDE-4 (2000)	10	-1.96 ± 0.68	3.39 ± 0.79	1.42 ± 0.39

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Table 2. Season averaged values (\pm SE) of chlorophyll-*a* concentration (Chl-*a*), Gross Primary Production (GPP), surface water temperature (T), Net Community Production (NCP) and Community Respiration (R), for the four LATITUDE cruises. * Significant One Way Anova probability.

	Spring (Boreal and Austral)	Autumn (Boreal and Austral)	Probability
Chl- <i>a</i> ($\mu\text{g l}^{-1}$)	0.116 ± 0.015	0.127 ± 0.018	0.64
GPP ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)	1.04 ± 0.14	0.91 ± 0.17	0.58
T ($^{\circ}\text{C}$)	24.13 ± 0.34	27.26 ± 0.36	$< 0.0001^*$
NCP ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)	-0.56 ± 0.26	-2.14 ± 0.30	$< 0.0002^*$
R ($\mu\text{mol O}_2\text{l}^{-1}\text{d}^{-1}$)	1.70 ± 0.24	3.19 ± 0.29	$< 0.0003^*$

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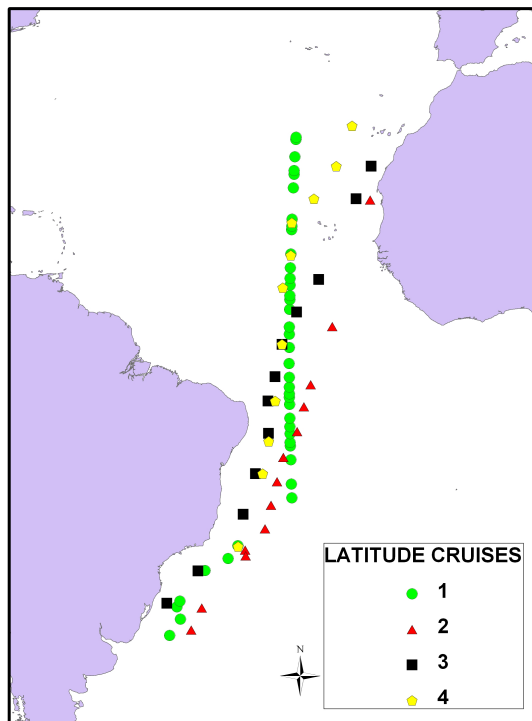


Fig. 1. Map showing the location of the sampling stations during the 4 LATITUDE cruises.

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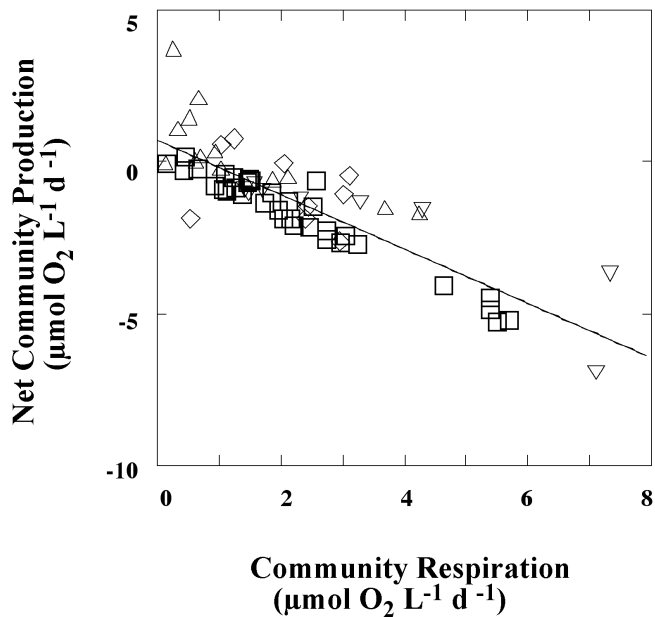


Fig. 2. The relationship between Net Community Production (NCP) and Community Respiration (R) for the surface pelagic communities sampled during the study. The solid line represents the fitted regression equation. The symbols corresponded to the 4 cruises: Squares: LATITUD-1, Diamonds: LATITUD-2, Triangles: LATITUD-3, and inverted Triangles: LATITUD-4.

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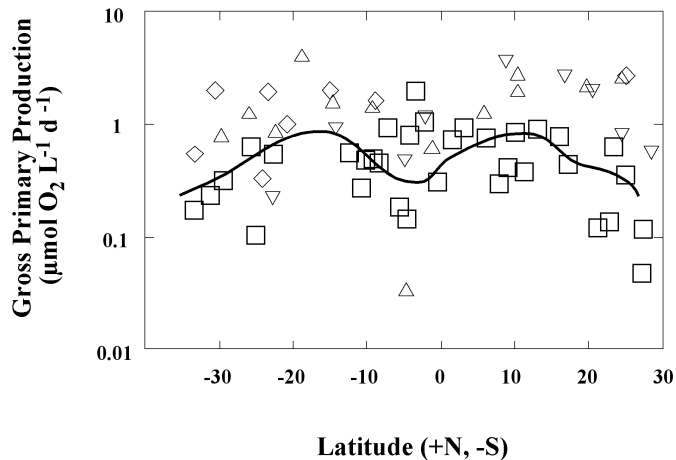


Fig. 3. Distribution of surface Gross Primary Production (GPP) with the latitude. The line represents the Lowest fit model. Symbols are as described in Fig. 2.

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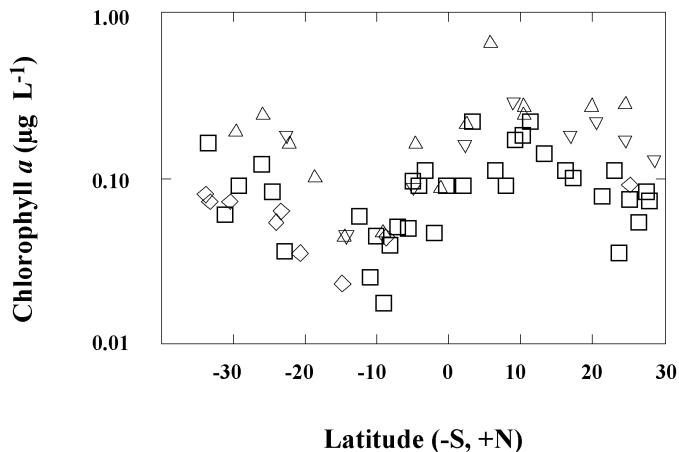


Fig. 4. Distribution of chlorophyll-*a* concentration with latitude, found at the surface of the Atlantic Ocean during the four LATITUDE cruises. Symbols are as described in Fig. 2.

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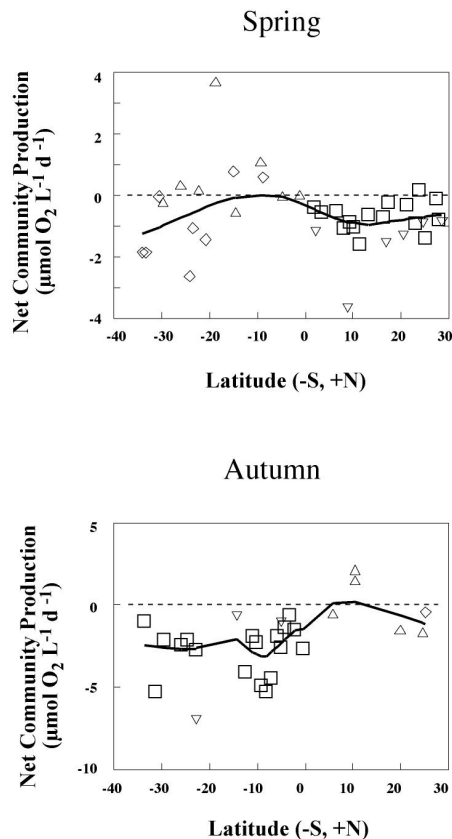


Fig. 5. Distribution with the latitude of surface Net Community Production (NCP) measured in the stations sampled during boreal and austral spring, and during boreal and austral autumn. The lines represent the lowest fitted models. Symbols are as described in Fig. 2.

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