Response to reviewers on 'Spring Blooms in the Baltic Sea have weakened but lengthened from 2000 to 2014' by P. M. M. Groetsch et al.

Anonymous Referee #2

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- The use of the term "variable-theshold" could still lead to confusion (i.e., annual median vs. climatological median). I suggest using "spatially variable-threshold" or similar.
- The sentence: "The observed negative tren in peak concentration introduces an artificial negative trend in bloom duration" can create confusion, because it actually does not (as it is explained afterwards). I suggest changing to "was expected to introduce".

Authors' response: We implemented both recommendation in the manuscript.

diff --git a/AlgalineSpringBloom.tex b/AlgalineSpringBloom.tex index e2e11b6..47a062b 100644 --- a/AlgalineSpringBloom.tex +++ b/AlgalineSpringBloom.tex @@ -114,7 +114,7 @@ Alg@line chla concentrations (see section \ref{ssec:algalin e}) were resampled to

1) Chla concentration exceeding a fixed-threshold of \SI{5}{\milli\gram\per\cubic\meter} w as defined as bloom by \citet{Fleming2006}, further referred to as \texttt{const5}. A 21-d ay centred-running-mean filter was used to keep results comparable to the other metrics co nsidered, whereas \citet{Fleming2006} used a 7-day centred-running-median filter.

2) \citet{Siegel2002} proposed a spatially variable-threshold metric based on the 5 \%-above-median concentration, but reported small quantitative differences for threshold s between 1 and 30 \%-above-median. Their threshold is based on the complete annual cycle, while here only the spring bloom period from day-of-year 31 to 160 is considered. We refe r to this metric as \texttt{median5}.

3) Distributions proposed to describe bloom phenology include shifted-Gaussian \citep{Plat t2009}, Gamma \citep{Vargas2009}, and Weibull distributions \citep{Rolinski2007}. The shif ted Gaussian is symmetric in shape, whereas Gamma distributions allow for different slopes of bloom rise and decline. In addition, Weibull functions recognize non-zero offsets befo re and after the bloom phase. The latter has proven essential to obtain a good fit for the transition phase between spring and summer bloom with the here analysed data set. A modif ied Weibull-function, as proposed by \cite{Rolinski2007}, was fitted non-linearly to the p reprocessed and scaled (to a range of zero to one) chla concentrations. The bloom initiati on and end are defined as the \$10^{th}\$ and \$90^{th}\$ percentiles before and after the blo om peak, respectively. This metric is further referred to as \texttt{weibull}.

@@ -137,7 +137,7 @@ Determination of the fluorescence yield was supported by an 'adaptive regression

Figure \ref{fig:diurnal_variability}A shows normalized fluorescence observations as a func tion of sampling time-of-day. Results are presented separately for summer (May to August), winter (November to February) and the transition periods (autumn, spring). Diurnal variab ility was most pronounced in summer, when the fluorescence signal varied on average 50 \% over the course of a day. In winter and during the transition periods (spring, autumn) a d iurnal variability of 35 and 38 \%, respectively, was contained in the fluorescence signal s. This seasonal effect is likely caused by variations in average irradiance intensity, wh ich are modulated primarily by sun elevation, but also by atmospheric conditions (e.g. clo ud cover, aerosol optical thickness) and optical properties of the water body (e.g. ice co ver, attenuation). Figure \ref{fig:diurnal_variability}B depicts normalized fluorescence a s a function of solar elevation. In this representation seasonal differences in diurnal va riability are essentially absent and the correspondence between solar elevation and averag e fluorescence response was approximately linear for daytime observations.

\subsection{Bloom Intensity and Timing}

Blooms generally developed first in the south and progressed towards the north (see Fig. \ ref{fig:phenology_geo_timing} and Table \ref{tab:bloomstats}). Bloom peak timing (not infl uenced by choice of metric) followed this pattern, as did metric-dependent bloom start and end dates. The fixed-threshold bloom metric \texttt{const5} suggested longer blooms in hi gh-biomass sea areas like the \textsc{gof}, compared to low-biomass areas such as the \tex tsc{sbs}. The spatially variable-threshold metric \texttt{median5} applies area-sp ecific bloom thresholds (\textsc{nbp}: \SI{3.52}{\milli\gram\per\cubic\meter}, \textsc{gof} }: \SI{4.95}{\milli\gram\per\cubic\meter}, \textsc{got}: \SI{2.51}{\milli\gram\per\cubic\m eter}, \textsc{sbs}: \SI{2.62}{\milli\gram\per\cubic\meter}, \textsc{bom}: \SI{4.02}{\milli i\gram\per\cubic\meter}) and resulted in approximately stable bloom duration in all sea ar eas. The \texttt{weibull} metric, which is not sensitive to absolute bloom intensity, also resulted in comparable bloom durations for all sea areas. The year-to-year variability of start, peak, and end days generally increased towards the south for all metrics.

Spring bloom intensity was described by three parameters: the metric-independent bloom pea

k concentration (\textsc{peakheight}), the chla concentration average during bloom conditi ons (\textsc{concavg}), and the sum of daily chla concentrations over the bloom period (\t extsc{bloomidx}). Similar patterns were observed for all these parameters and bloom metric s, as illustrated in Fig. \ref{fig:phenology_geo_intensity}. The highest bloom intensity w as found in the \textsc{gof} and \textsc{nbp}, followed by the \textsc{bom}. Low-intensity blooms were observed in the \textsc{sbp} and the \textsc{got}. Variability was generally proportional to bloom intensity, highest in the high-biomass and coastal \textsc{gof} and \textsc{bom}. Variability in \textsc{bloomidx} was comparable to that in \textsc{peakheigh t}, while \textsc{concavg} was considerably more stable. All calculated bloom phenology pa rameters can be found in the supplementary material.

@@ -183,7 +183,7 @@ Interannual variability in coastal systems exceeds long-ter
m trends by orders of

\cite{Helcom2014} reported stable or increasing chla concentrations for the period 2007-20
11 in several Baltic Sea areas despite signs of declining nutrient concentrations. More re
cently, eutrophication trend reversal and oligotrophication processes were reported by \ci
te{Andersen2015}, based on analysis of 112 years of consolidated Baltic Sea observations.
Both reports considered surface-layer chla concentration in summer as one of the direct in
dicators for eutrophication, but did not include spring bloom in their assessment. The tim
e series for 2000-2014 that we present here fills this gap: a negative trend in bloom inte
nsity was also found for spring bloom, providing further evidence for their hypothesis of
gradual nutrient load reduction.

Thresholds of \texttt{const5} and \texttt{median5} are fixed for the whole time series. Th e observed negative trend in peak concentration introduceswas expected to int roduce an artificial negative trend in bloom duration because an increasingly higher pe rcentile of the distribution is seen below the bloom threshold (Fig. \ref{fig:trend_scheme }). Contrary to this expected behaviour, however, \texttt{const5} and \texttt{median5} rev ealed no significant trends in bloom duration. This indicates that the anticipated negativ e trend in bloom duration was countered by a positive trend, e.g. in bloom intensity. The Weibull-metric is based on concentration distribution-ratios that are calculated individua lly for each bloom. Therefore, Weibull-metric results for bloom duration are not sensitive to long-term trends in peak concentration. Weibull-distribution metrics confirmed a highl y significant, positive trend in bloom duration. These two sets of results corroborate the conclusion that spring blooms in the Baltic Sea have become longer, while chla peak and a verage concentration levels declined.

This 'flattening' of the concentration distribution is supported by the absence of a trend in time-integrated biomass \textsc{bloomidx} and by shifts in nutrient concentration timi ng (earlier nutrient peak concentration, later 25 \%-of-peak-value day). These results ind icate that annually generated spring bloom biomass has not changed significantly over the study period, in contrast to bloom timing. \cite{Kahru2014} found a similar development fo r cyanobacterial summer surface bloom, and reported decadal oscillations, yet no long-term trend, of surface area covered by cyanobacteria in the period 1979-2013. In the same peri od, summer bloom initiation moved to earlier dates by \SI{-0.6}{\day\per\year}. These resu lts suggest that the gap has decreased between dinoflagelate- and diatom-dominated spring bloom and cyanobacterial summer bloom. Due to the shorter period covered here as compared to the time series presented by \cite{Kahru2014}, it cannot be ruled out that the spring b loom trends are caused by decadal oscillation. Moreover, Alg@line nutrient records often d id not commence sufficiently early in the season to record bloom onset. Trends in bloom st art and nutrient peak timing can therefore not be derived at the same accuracy and precisi on as the other phenological parameters. In future, additional data and longer time series may revise this analysis. To this end, nutrient metrics derived in this work are provided in the appendix.