

1 **The influence of tillage on N₂O fluxes from an intensively managed grazed** 2 **grassland in Scotland**

3 **Authors:** N.J. Cowan ^{a, b}, P.E. Levy ^a, D. Famulari ^a, M. Anderson ^a, J. Drewer ^a, M. Carozzi ^c, D.S. Reay ^b, U.M.
4 Skiba ^a

5 ^a Centre for Ecology and Hydrology, Penicuik, Edinburgh, UK, EH26 0QB

6 ^b School of Geosciences, Kings Buildings, University of Edinburgh, Edinburgh, UK, EH9 3JG

7 ^c INRA, INRA-AgroParisTech, UMR 1402 EcoSys, 78850 Thiverval-Grignon, France.

8

9 **Keywords:** plough, greenhouse gas, nitrous oxide, gap filling

10 *Correspondence to:* Nicholas Cowan (nicwan11@ceh.ac.uk)

11

12 **Abstract**

13 Intensively managed grass production in high-rainfall temperate climate zones is a globally important source of
14 N₂O. Many of these grasslands are occasionally tilled to rejuvenate the sward, and this can lead to increased N₂O
15 emissions. This was investigated by comparing N₂O fluxes from two adjacent intensively managed grazed
16 grasslands in Scotland, one of which was tilled. A combination of eddy covariance, high-resolution dynamic
17 chamber and static chamber methods was used.

18 N₂O emissions from the tilled field increased significantly for several days immediately after ploughing and
19 remained elevated for approximately two months after the tillage event contributing to an estimated increase in
20 N₂O fluxes of 0.85 ± 0.11 kg N₂O-N ha⁻¹. However, any influence on N₂O emissions after this period appears to
21 be minimal. The cumulative N₂O emissions associated with the tillage event and a fertiliser application of 70 kg-
22 N ammonia nitrate from one field were not significantly different from the adjacent un-tilled field, in which two
23 fertiliser applications of 70 kg-N ammonia nitrate occurred during the same period. Total cumulative fluxes
24 calculated for the tilled and un-tilled fields over the entire 175 day measurement period were 2.14 ± 0.18 and 1.65
25 ± 1.02 kg N₂O-N ha⁻¹, respectively.

26 1 Introduction

27 Modern agriculture and intensive land management practices are believed to contribute over 39 % of total global
28 anthropogenic emissions of the greenhouse gas (GHG) nitrous oxide (N₂O) (IPCC, 2014). N₂O is a naturally
29 occurring GHG released into the atmosphere by the microbial processes of nitrification and denitrification which
30 occur in soils and aquatic systems (Davidson et al., 2000; Seitzinger et al., 2000). Human activities which alter
31 environmental conditions can have a significant impact on natural microbial processes, which in turn can increase
32 N₂O emissions. Agricultural activities such as the use of nitrogen fertilisers, livestock production and land use
33 changes are all important sources of anthropogenic N₂O from agricultural soils (Fowler et al., 2013).

34 There is still large uncertainty associated with the quantification of N₂O emissions released from
35 agricultural soils on a national and global scale, due to the large spatial and temporal variability of N₂O fluxes
36 (Cowan et al., 2015; Jahangir et al., 2011; Mathieu et al., 2006). Many past experiments have focussed on the
37 release of N₂O from soils after the application of nitrogen fertilisers - which is the main cause of the rise of in
38 N₂O emissions since pre-industrial times (e.g. Bouwman et al., 2002; Dobbie et al., 1999). Other factors affecting
39 N₂O emissions from agricultural soils, such as tillage and compaction, are less well documented, thus preventing
40 effective assessment of their role in controlling N₂O fluxes from the agricultural sector.

41 The addition of organic nitrogen in the form of decaying plant matter (crop residues) is a recognised
42 potential source of N₂O following tillage, but the phenomenon is not well quantified (Baggs et al., 2003; Mutegi
43 et al., 2010). Currently the IPCC emission inventories estimate that 1 % of all organic nitrogen applied to soils as
44 crop residues will be emitted in the form of N₂O (IPCC, 2006). However, the degree to which tillage induces a
45 change in N₂O emissions may be determined by several factors: the prior use of nitrogen fertilisers (Abdalla et
46 al., 2010; Yamulki and Jarvis, 2002), soil compaction (Ball et al., 2008; Yamulki and Jarvis, 2002) and the method
47 of tillage (Sheehy et al., 2013). Changes in the bulk density, water filled pore space (WFPS) and oxygen
48 availability in soils which can lead to an increase or decrease in nitrification and denitrification rates depending
49 on environmental conditions (Elmi et al., 2003; Palma et al., 1997).

50 The large number of variables which may alter microbiological processes in tilled soils can lead to a wide
51 range of results between experiments carried out at different field sites, under different meteorological conditions.
52 Some experiments have reported large increases in annual N₂O emissions varying from 0.89 to 3.37 kg N ha⁻¹
53 dependent on application of fertiliser post-tillage (i.e. Chatskikh and Olesen, 2007; Merbold et al., 2014; Omonode

54 et al., 2011; Pinto et al., 2004; Yamulki and Jarvis, 2002), whereas others have shown a zero (i.e. Boeckx et al.,
55 2011; Choudhary et al., 2002) or potentially negative effect of tillage ($-0.88 \text{ kg N ha}^{-1}$, Tan et al., 2009). There is
56 little consensus among these studies on the relative effect of different drivers of N_2O production. However, it is
57 commonly reported that factors influencing the aeration of the soil (such as WFPS and bulk density) are cited as
58 influential in most tillage studies.

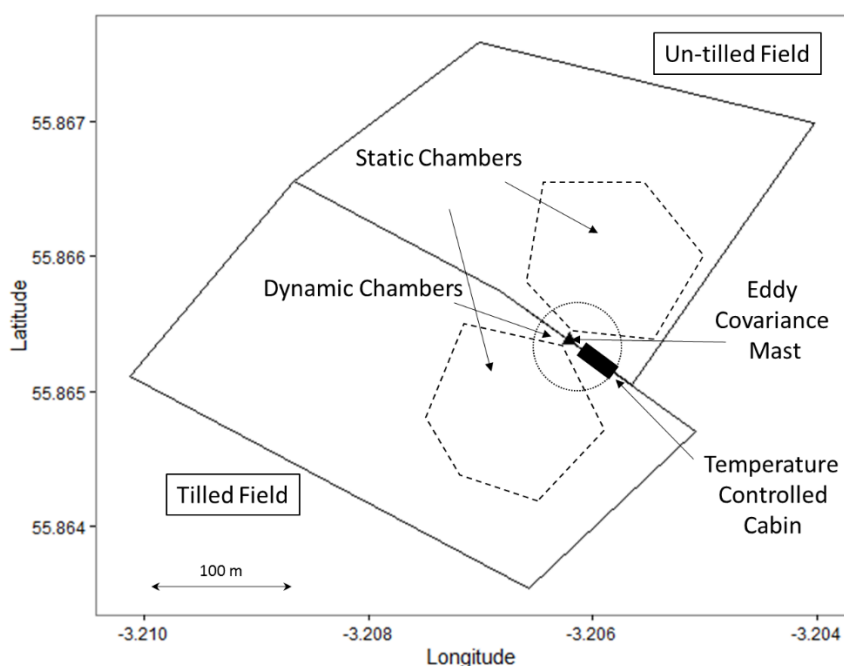
59 Improving our understanding of N_2O fluxes from tillage events is important, especially in countries such
60 as the UK, where agriculture accounts for approximately 70 % of the total land coverage (DEFRA, 2012) and
61 tillage is widely practiced. Improved grasslands alone account for 25 % of the total land coverage of the UK
62 (Morton et al., 2011). Tillage events occurs on rotational grasslands, for sward rejuvenation on permanent
63 grasslands, and in conversion to arable, and are a common enough occurrence that they could contribute
64 significantly to the total national inventory of anthropogenic N_2O emissions. However, few experiments have
65 been carried out on GHG emissions resulting from the tillage of grassland fields. The aim of this work was
66 therefore (i) to use multiple N_2O flux measurement methodologies to add to the understanding of the N_2O fluxes
67 from grasslands tilled for sward renewal, (ii) develop an improved statistical methodology which allows for
68 uncertainties in cumulative flux emissions to be calculated for these events, and (iii) compare our estimates with
69 those predicted using the current IPCC methodology.

70

71 2 **Materials and method**

72 2.1 **Field site**

73 Fluxes of N₂O were measured from an area of intensively managed, grazed grassland (Easter Bush, Scotland, 55°
74 51' 55.30"N, 3° 12' 22.17"W) before and after a tillage event on the 1st of May 2012, and were compared with
75 fluxes measured from an adjacent grassland which remained un-tilled (as described in Jones et al., 2011) (Figure
76 1). The climate is temperate maritime, with an average annual rainfall of 921 mm and average annual air
77 temperature of 9 °C (in the period 2001–2011). The two fields (each approximately 5.4 ha) have been managed
78 for intensive livestock production for at least twenty years, and since 2002 were predominately grazed by sheep.
79 The average stocking densities were 0.7 LSU ha⁻¹ (livestock units) and average N fertiliser application rates have
80 been approximately 200 kg N ha⁻¹ y⁻¹. Mainly NH₄NO₃ or NPK compound fertilisers were applied in three split
81 applications usually between March and July (Skiba et al., 2013).



82

83 **Figure 1** N₂O fluxes were measured from two adjacent grassland fields at the Easter Bush Farm (Penicuik,
84 Scotland). The north field remained un-tilled, while the south field was ploughed on the 1st of May 2012. An eddy
85 covariance mast was set up next to a permanent cabin positioned between the fields. Dynamic chamber
86 measurements were made within a 30 m radius of the cabin. Static chambers were located within the fetch of the
87 eddy covariance mast and moved periodically.

88 The soil in the fields is a clay loam with a sand/silt/clay texture of 52/20/28 and 57/19/24 for the top 30
 89 cm in the un-tilled and tilled fields, respectively with a pH of approximately 5.1 (in H₂O). They are classed as an
 90 imperfectly drained Macmerry soil of the Rowanhill association (eutric cambisol, FAO classification). A drainage
 91 system had been installed about 50 years ago, but is no longer functioning well, resulting in frequent occurrence
 92 of surface water during rainy periods. The fields had not been tilled for at least twenty years, and the farmer had
 93 reported reduced fertility and productivity. One field (also called the South Field in Jones et al., 2011) was
 94 therefore tilled in May 2012 (Table 1).

95 As standard practice, glyphosphate (1.5 l ha⁻¹) was applied to kill the grass three days prior ploughing
 96 on the 27th of April. The field was ploughed to a depth of 30 cm on the 1st of May 2012. Two days after ploughing,
 97 the field was harrowed, and then rolled and sown with ryegrass (*Lolium perenne L.*) on the third day after
 98 ploughing. The un-tilled field (also called the North field in Jones et al., 2011) was managed as usual and grazed
 99 by sheep (approximately 30 sheep ha⁻¹). Fertilisation events continued as normal on the un-tilled field which
 100 received two ammonium nitrate (Nitram) fertiliser applications of 70 kg-N ha⁻¹, one on the 28th of May and the
 101 second on the 9th of August. The tilled field only received a 70 kg-N ha⁻¹ Nitram application on the 9th of August,
 102 approximately four months after the tillage event.

103 **Table 1** Field management events for both the tilled and un-tilled fields in 2012.

Date	Tilled Field (South)	Un-Tilled Field (North)
16 th February 2012		Grazed by sheep (continuous)
27 th April 2012	Glyphosphate application (1.5 l ha ⁻¹)	
1 st May 2012	Ploughing at 30 cm depth	
3 rd May 2012	Harrowing, seeding & rolling	
28 th May 2012		70 kg-N ha ⁻¹ Nitram application
9 th August 2012	70 kg-N ha ⁻¹ Nitram application	70 kg-N ha ⁻¹ Nitram application
19 th September 2012	Grazed by sheep (continuous)	

104

105 Biomass samples were collected from the South Field prior to tillage in order tilled to estimate the grass
 106 biomass that would be tilled into the soil. Twenty soil cores (12 cm deep and 5.8 cm diameter) were extracted
 107 from the field. At these points, all above-ground biomass was harvested and dried in an oven at 80 °C to constant

108 weight. Once dry, the above-ground biomass was weighed. The soil cores were broken up by hand and dried at
109 100°C until constant weight. After drying, the root material was separated from the soil by hand and weighed. Sub
110 samples of the dried plant materials were prepared for elemental analysis of total carbon and nitrogen contents
111 (vario EL cube, Elemaentar, Hanau, Germany).

112 Total (above- and below-ground) biomass on the tilled field before tillage averaged of $369 \pm 310 \text{ g m}^{-2}$,
113 with a root to shoot ratio of ~ 1.5 . The nitrogen content was 2.5 %. Based on these measurements it is estimated
114 that the tillage event added a total of 93.6 kg ha^{-1} of nitrogen to the field in the form of crop residues.

115

116 2.2 Flux Measurements

117 N_2O fluxes were measured from both tilled and un-tilled fields over a seven month period using three measurement
118 methodologies; eddy covariance, static chamber and dynamic chamber techniques. The mixture of methods were
119 used to try to obtain as many measurements as practically possible, both temporally and spatially, during the
120 experiment. Eddy covariance was the primary measurement methodology used; however, due to unpredictable
121 changes in wind direction at the site it was necessary to deploy manual chamber methodology to ensure that both
122 fields were measured periodically during the experiment. The dynamic chamber measurements were used as a
123 cost effective way to provide many (> 30) high-resolution N_2O fluxes on the days immediately after tillage without
124 the need for time consuming GC lab analysis required by static chambers.

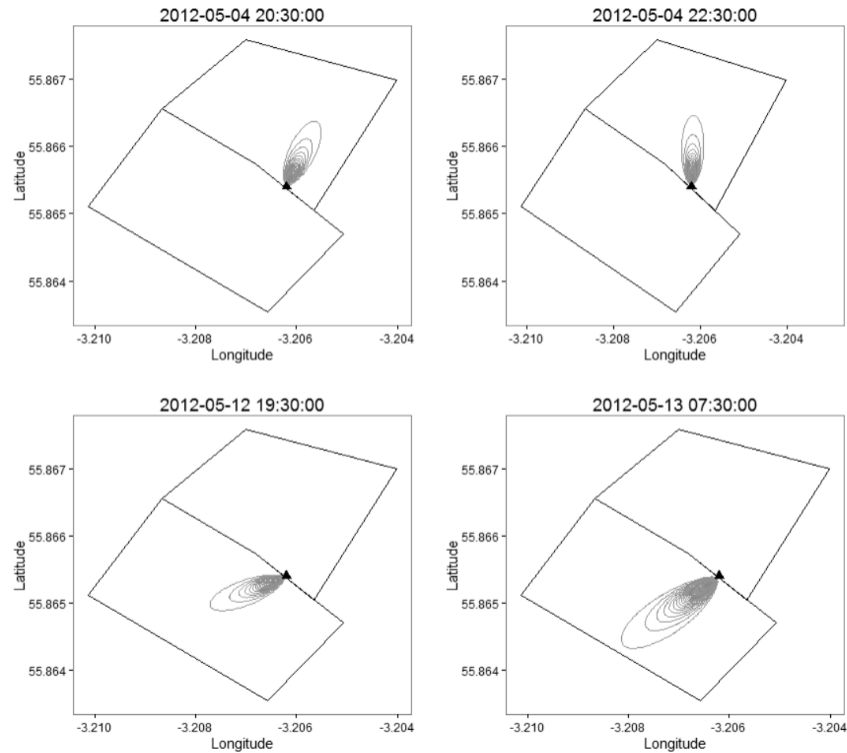
125 An eddy covariance system was installed on the 27th of March on the field boundary (Figure 1). An
126 ultra-sonic anemometer (WindMaster Pro 3-axis, Gill, Lymington, UK) mounted at 2.4 m was used to measure
127 fluctuations in 3-D wind components at a frequency of 10 Hz. Mixing ratios of N_2O , H_2O and CO_2 were measured
128 at 10 Hz by a quantum cascade laser (QCL) gas analyser (CW-QC-TILDAS-76-CS, Aerodyne Research Inc.,
129 Billerica, MA, USA), housed in a temperature controlled cabin. The inlet line to the QCL was a 13.5 m length of
130 Dekabon tubing (0.25 inch outer diameter), with a flow rate of approximately 13 l min^{-1} . Fluxes were calculated
131 at 30 min intervals using the EddyPro software (Version 5.2.1) (Li-Cor, Lincoln, NE, U.S.A.), based on the
132 covariance between the N_2O concentration (χ) and vertical wind speed (w):

$$133 \quad F_{\chi} = \overline{\chi'w'} \quad (\text{Eq. 1})$$

134 In the processing, we applied double coordinate rotation (vertical and crosswind), spike removal, block
135 averaging, and time lag removal by covariance maximisation. Correction for the frequency response of the system,
136 both high and low-frequency losses, were made using the method of Moncrieff et al., (1997). Corrections for
137 density fluctuations were applied on a half-hourly basis using the method of (Burba et al., 2012). The quality
138 control scheme of Mauder and Foken (2006), was used to remove poor quality flux measurements (their category
139 2). Initially, fluxes measured with a mean wind direction between 180 and 270 degrees from north were classed
140 as from the tilled field; those measured at greater than 330 and less than 100 degrees were classed as from the un-
141 tilled field. The remaining data were disregarded due to obstruction of the wind by the cabin and fence line.

142 Further footprint analysis was carried out in which we visually checked individual footprint plots of each
143 30 min flux (Figure 2). Any flux footprints that were very small or overlapped the two fields were removed from
144 the dataset. Standard meteorological variables (rainfall, air temperature and soil temperature) were recorded by a
145 tipping bucket, thermometers (2 m height & 10 cm depth) and TDR soil moisture probe at 10 cm depth. These
146 measurements were made adjacent to the flux tower at the site.

147



148

149 **Figure 2** Four example flux footprints, with contours showing the relative contribution to the measured eddy
 150 covariance flux, based on the model of Kormann and Meixner (2001). Half-hourly flux data were only included
 151 if 97.5 % of the measured flux was attributed to either the untilled (top) or tilled (bottom) fields.

152 N_2O fluxes were also measured from both fields using static chamber and dynamic chamber techniques.
 153 The static chambers consisted of a cylindrical polyvinyl chloride (PVC) plastic pipe of 38 cm inner diameter (ID)
 154 and 22 cm height. These chambers were inserted 5 cm into the soil, giving a headspace of approximately 20.4 l.
 155 Chambers were closed for 40 mins, during which time three 100 ml gas samples were collected via a syringe and
 156 a three-way tap fitted to the lid, at $t = 0, 20$ and 40 mins. After each measurement, chamber height was measured
 157 at five points to estimate the chamber volume. Gas samples were stored in 20 ml glass vials which were flushed
 158 with 100 ml of air in the syringe using a double needle. Samples were analysed using a Hewlett Packard 5890
 159 series II gas chromatograph (Agilent Technologies, Stockport, fitted with an electron capture detector) (Skiba et
 160 al., 2013).

161 Ten static chambers were positioned in each of the fields, within the estimated flux footprint of the eddy
 162 covariance system (10 to 200 m from the mast). Chambers in the fields were occasionally moved to prevent the
 163 effects of a micro-climate within the chambers that could bias measurements when compared to the surrounding

164 field area, and also to allow access to farm vehicles during the different stages of the tillage operation. Manual
165 chamber measurements were carried out between 9:00 and 15:00 on the measurement dates. Fluxes were
166 calculated as:

$$167 \quad F = \frac{dC}{dt} \cdot \frac{\rho V}{A} \quad (\text{Eq. 2})$$

168 where F is the gas flux from the soil ($\text{nmol m}^{-2} \text{s}^{-1}$), dC/dt is the rate of change in concentration with time in nmol
169 $\text{mol}^{-1} \text{s}^{-1}$ estimated by linear regression, ρ is the density of air in mol m^{-3} , V is the volume of the chamber in m^3
170 and A is the ground area enclosed by the chamber in m^2 . Static chamber measurements were made over a longer
171 period than shown in this paper and are discussed in relation to a second tillage event by Drewer et al. (2016).

172 Fluxes were also measured using the QCL in a closed, dynamic chamber system (Cowan et al., 2014a).
173 A chamber (39 cm inner diameter, 22 cm high) was placed onto a stainless steel collar inserted several cm into
174 the soil (on average 5 cm) at least 15 mins prior to measurement. Two 30 m lengths of 3/8 inch ID Tygon® tubing
175 connected the chamber to the inlet of the QCL and the outlet of a vacuum pump (SH-110, Varian Inc, CA, USA)
176 to form a closed system. This allowed a 30 m possible radius from the instrument cabin in which the chamber
177 could be placed (Figure 1). A flow rate of approximately 6 to 7 L min^{-1} was used, with a lag time of approximately
178 22 seconds between the chamber and analyser. Fluxes of N_2O were calculated with 1 Hz data over three minutes,
179 using both linear and non-linear asymptotic regression methods (Levy et al., 2011; Pedersen et al., 2010). Using
180 a mixture of goodness-of-fit statistics and visual inspection, the regression method that provided the best fit for
181 the time series of mixing ratios of N_2O was chosen for each individual measurement. The detection limit of
182 individual fluxes calculated by this method was approximately $0.04 \text{ nmol m}^{-2} \text{ s}^{-1}$ compared to $0.4 \text{ nmol m}^{-2} \text{ s}^{-1}$
183 when using the static chambers (Cowan et al., 2014a, 2014b).

184 In the first few days after the tillage event, the wind direction was north-easterly, meaning that the eddy
185 covariance system could not record fluxes from the tilled field (to the south-west). The dynamic chamber
186 measurements were primarily used to fill this gap in the eddy covariance time series with high precision chamber
187 measurements.

188

189 2.3 Gap filling

190 Because the eddy covariance system was placed on the field boundary, observations could only be made on a
191 single field at any given time. Furthermore, some data were missing because of instrument failure and some had
192 to be rejected according to the quality control criteria used. In order to estimate cumulative fluxes from both
193 fields, temporal interpolation of the missing data points was required. However, in the absence of a well-validated
194 process-based model for N₂O fluxes on which to base predictions, it is not obvious how this is best achieved. The
195 most common approach is to linearly interpolate in time between flux measurements. In this study, a general
196 additive model (GAM) was used as an alternative approach, which accounted for temporal patterns at a range of
197 time scales and nonlinear responses to environmental variables, implemented using the *mgcv* package in the R
198 software (Wood, 2006).

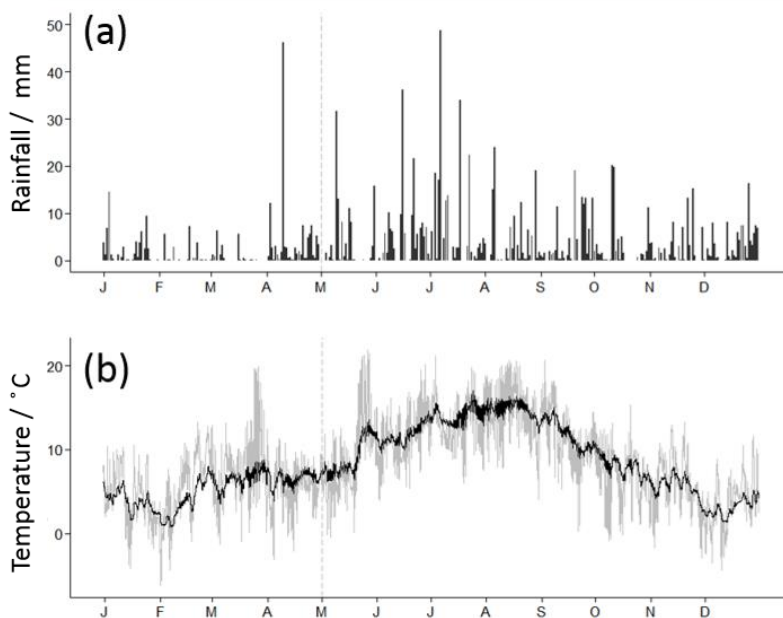
199 We fitted a GAM with the same environmental terms to the separate data sets from the tilled and un-
200 tilled fields. The terms included were air temperature, soil temperature, precipitation, and time. Additional terms
201 for temperature and precipitation aggregated over longer intervals (1, 6, 12, 24 and 48 hours preceding the flux
202 measurement) were examined and included where they improved the fit. The GAM allows for non-linearity by
203 fitting a smooth response with cubic splines. The degree of smoothing is optimised by the algorithm, but was also
204 adjusted subjectively, such that the model was not over-fitting to noise in the data. Observations from eddy
205 covariance and the two chamber methods were given equal weighting. Predictions from the GAM were used to
206 fill gaps when observations were not available. Uncertainty in predictions was estimated by simulating 2000
207 replicate time series from the GAM, using the uncertainty in the fitted parameters, to estimate the posterior
208 distribution. The quantiles of this posterior distribution provided the 95 % credibility interval at each predicted
209 30-min interval time step. To calculate cumulative fluxes, observed fluxes were used with their associated
210 uncertainties (Finkelstein and Sims, 2001) when available; otherwise the GAM predictions were used.

211 3 Results

212 3.1 Meteorological data

213 A total of 1191 mm of rain was recorded in 2012, higher than the average annual rainfall of 921 mm (2001 to
214 2011) for the Easter Bush area (Figure 3a). The annual variation in temperature was fairly typical of the field site
215 (Figure 3b). The prevailing wind direction at the field site is predominantly south-westerly (85 %). However,

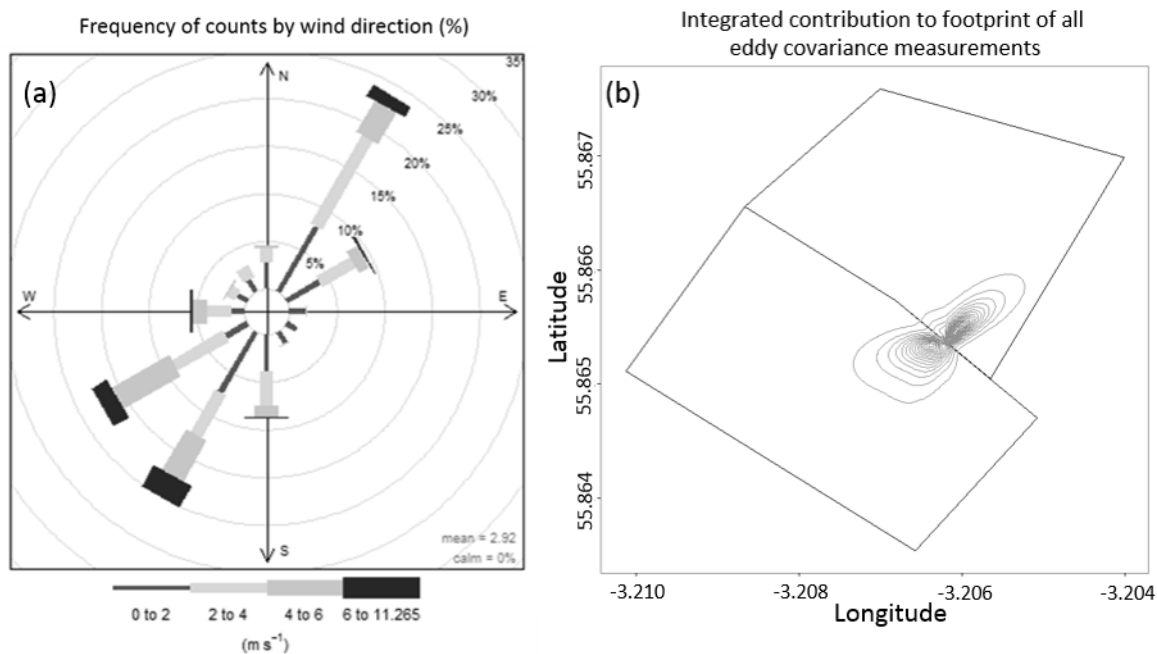
216 during the measurement campaign, the wind direction was split fairly evenly between the tilled and un-tilled fields
217 (Figure 4). This allowed a better basis for comparison of N₂O fluxes from the two fields, although data coverage
218 for each field was low, 34 % and 24 % for tilled and un-tilled respectively.



219

220 **Figure 3** (a) Accumulated daily rainfall at the Easter Bush Field site during the year 2012. (b) Air temperature at
221 height 3 m (grey) and soil temperature (black) recorded at the Easter Bush field site during the year 2012. Tillage
222 occurred on the 1st of May 2012 (grey dashed vertical line).

223



224

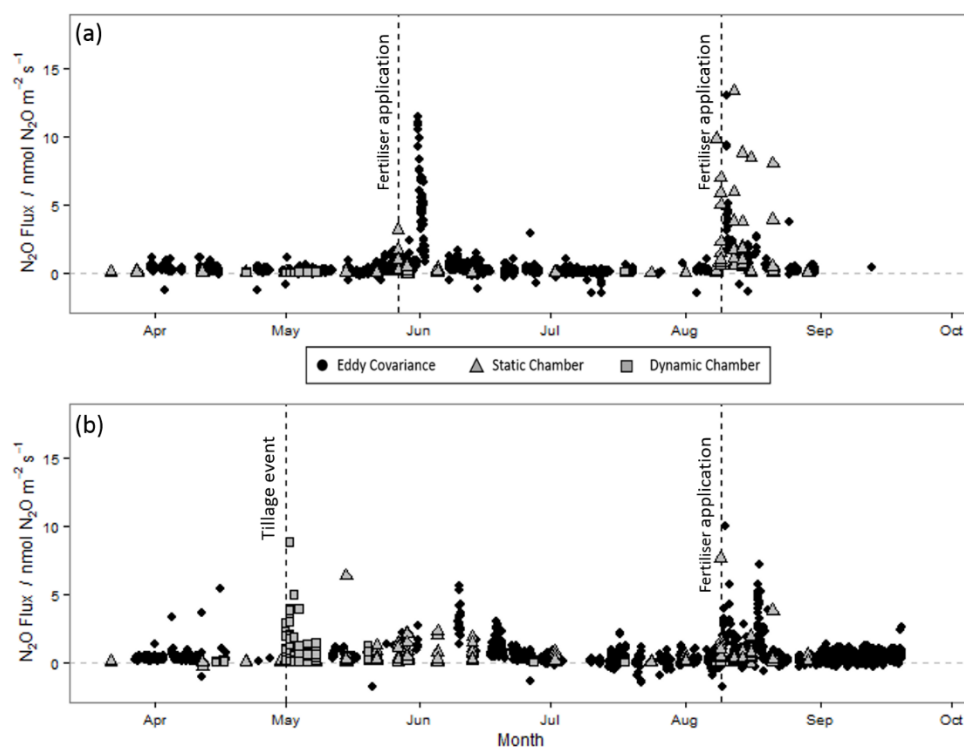
225 **Figure 4** (a) Wind rose plot for the Easter Bush field site during eddy covariance measurements (March – October
 226 2012). (b) Spatial distribution of the time-averaged flux footprint over the measurement period. The outer-most
 227 contour represents the area which, on average, contributed to 97.5 % of the measured half-hourly flux.

228 3.2 Comparison of N_2O fluxes measured from the un-tilled and tilled fields

229 Before the tillage event, N_2O fluxes were similar in the tilled and untilled fields. In both cases, around 90 % of
 230 measured fluxes were below $0.5 \text{ nmol m}^{-2} \text{ s}^{-1}$ (Figure 5). All three fertilisation events (the two fertiliser events in
 231 the untilled field and single fertiliser event in the tilled field) were characterised by an emission peak of 5-10 nmol
 232 $\text{m}^{-2} \text{ s}^{-1}$ lasting a few days, which declined over the following days and weeks, often with considerable variability
 233 and some apparent secondary peaks (Figure 5). Fluxes had returned to background levels ($<0.5 \text{ nmol m}^{-2} \text{ s}^{-1}$)
 234 within 28 days of each of the fertilisation events. Fluxes measured by all methods agreed reasonably well in
 235 magnitude, and there is no strong evidence for a systematic bias, given the differences in the spatial and temporal
 236 sampling (for a more specific insight see e.g. Cowan et al., 2014a).

237 The tillage event also produced an increase in emissions, and although the peak was less clearly defined,
 238 the effect was more prolonged. Fluxes generally ranged from ~ 0 to $1.0 \text{ nmol m}^{-2} \text{ s}^{-1}$ in the days before tillage and
 239 ~ 0 to $8.8 \text{ nmol m}^{-2} \text{ s}^{-1}$ in the week immediately after tillage (Figure 5b). Three exceptionally high individual
 240 chamber measurements measured in the days immediately after the second fertilisation event in the un-tilled field
 241 which are included in the data analysis (19.5 , 34.8 and $50 \text{ nmol m}^{-2} \text{ s}^{-1}$) are not included in Figures 5 or 7 in order

242 to keep the scale manageable. Fluxes from the tilled field from mid to late May were approximately $1 \text{ nmol m}^{-2} \text{ s}^{-1}$
 243 higher than from the untilled field (before the latter was fertilised). There followed an apparent increase in N_2O
 244 fluxes lasting approximately four weeks from the tilled field from late May to late June, peaking mid-June (Figure
 245 5b). Unfortunately, data coverage was rather low during this period due to changes in wind direction and a five
 246 day period in which the QCL was not operational. Because the tilled field had not been fertilised since the previous
 247 year, we infer that the increased fluxes were a result of the tillage event. Fluxes in the tilled field returned to pre-
 248 tillage magnitude during July. By July, a new sward of grass had grown in the tilled field, but sheep were not re-
 249 introduced into the field until September.

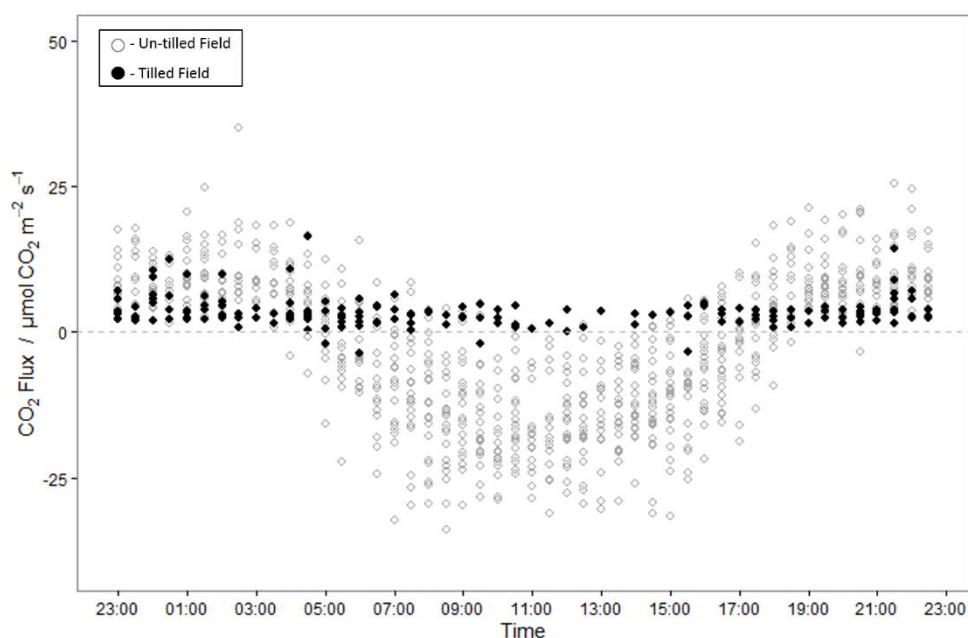


250

251 **Figure 5** Fluxes of N_2O from the (a) un-tilled and (b) tilled fields measured at the Easter Bush field site in 2012.
 252 Fertiliser was applied to the un-tilled field on the 28th of May and to both fields on 9th August (vertical dashed
 253 lines). Tillage began on 1st May. The Y-axis is limited to $15 \text{ nmol m}^{-2} \text{ s}^{-1}$ for better comparison between the
 254 fields. Only three static chamber measurements in the un-tilled field recorded fluxes above $15 \text{ nmol m}^{-2} \text{ s}^{-1}$ in the
 255 first few days after the August fertilisation.

256 The relatively high N_2O fluxes measured from the tilled field in the weeks after tillage (May to July)
 257 occur in a similar timeframe to the fertiliser event in the un-tilled field (Figure 5). Beyond the analytical footprint

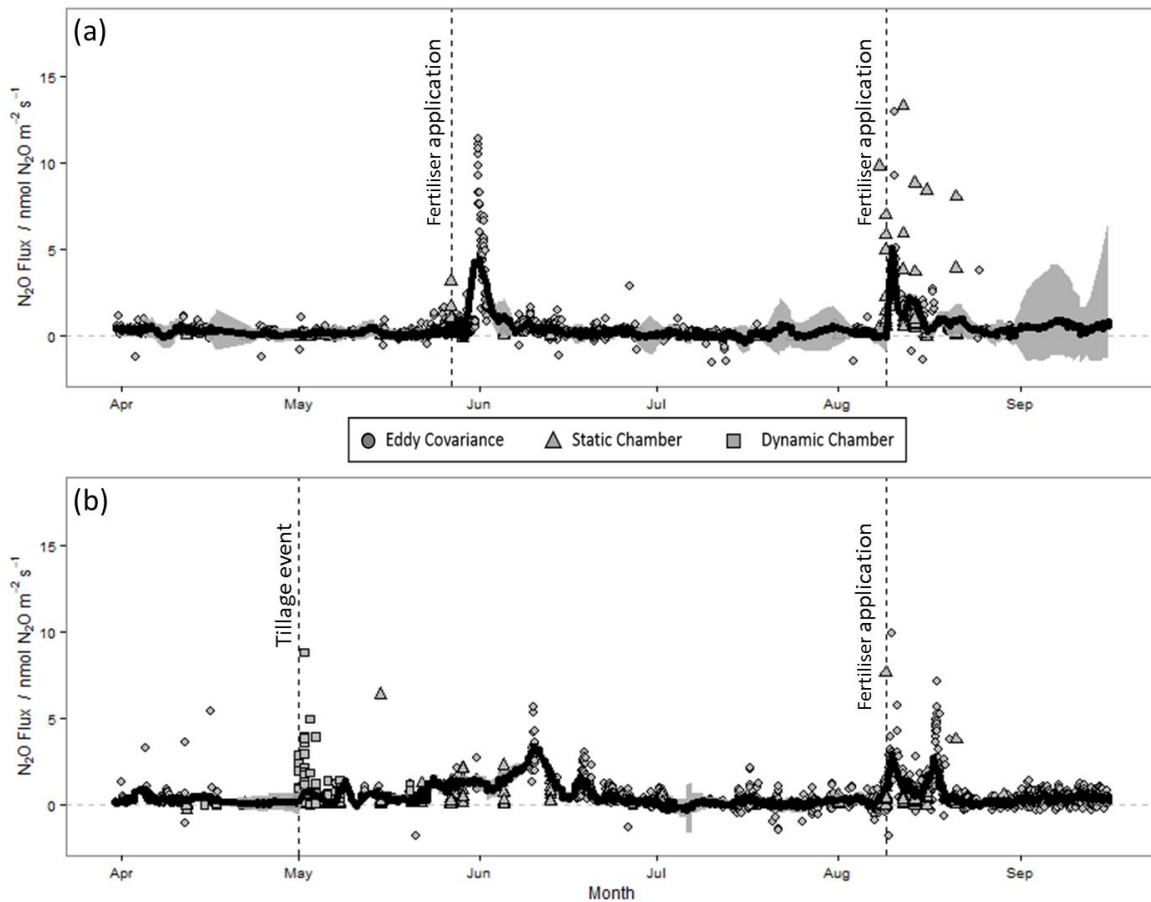
258 analysis, we wanted to check that the high N₂O fluxes, which we attribute to the tillage, actually do come from
259 the tilled field, and are not influenced by N₂O from fertilisation events on surrounding fields. The CO₂ fluxes
260 (measured by QCL instrument) provide a suitable tracer. We know that no significant photosynthesis took place
261 on the tilled field between 1st of May and 17th of June, as there was no green foliage visible until after this period.
262 Therefore, if the CO₂ fluxes showed no day-time uptake on the tilled field, we can be reasonably certain that the
263 measured N₂O fluxes were also coming from the tilled field. Figure 6 shows that this was the case: in fluxes
264 attributed to the tilled field, there was no day-time uptake of CO₂; in fluxes attributed to the un-tilled field, the
265 normal diurnal cycle in CO₂ flux is seen. By inference, we can attribute the high N₂O emissions after tillage to the
266 tilled field.



267
268 **Figure 6** CO₂ flux measurements made from the un-tilled (grey) and tilled (black) fields between the 1st of May
269 and the 17th of June. Uptake is denoted as a negative quantity. The results show a clear difference between the
270 fields, with no day-time uptake on the tilled field. This implies that the high N₂O fluxes measured after the
271 tillage event can also be attributed to the tilled field.

272 The GAM method was used to gap-fill flux data to calculate cumulative fluxes (Figure 7). Cumulative
273 N₂O fluxes calculated for the tilled and un-tilled fields for the 169 days from 1st of April to the 16th of September
274 were 2.14 ± 0.18 and 1.65 ± 1.02 kg N₂O-N ha⁻¹, respectively (Figure 8). Uncertainty in the GAM prediction is
275 particularly large when no measurements are available in which to fit the model. There are sustained periods in

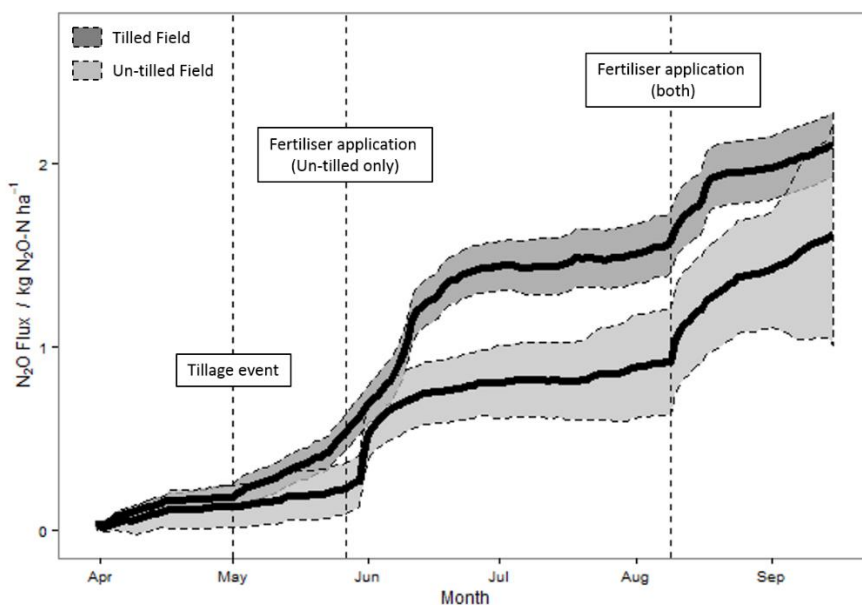
276 which very few eddy covariance measurements were recorded from the un-tilled field due to the wind direction
277 being predominantly south westerly (Figure 7a). The uncertainty in predicted flux becomes very large when
278 compared to periods when measurements data is available and these uncertainties propagate significantly in
279 cumulative flux estimates (Figure 8).



280

281 **Figure 7** The GAM method (black line) provides an estimated N₂O flux which can be used to gap-fill
282 measurements from both the (a) un-tilled and (b) tilled fields at 30 min intervals. The 95 % confidence interval in
283 the estimated flux reported by the GAM is included (grey). Tillage and fertiliser dates are indicated (vertical lines).

284



285

286 **Figure 8** Cumulative flux is calculated for the tilled (dark grey) and un-tilled fields (light grey) using the gap-
 287 filled flux data. The cumulative 95 % confidence intervals are shown (grey areas). Fertiliser was applied to the
 288 un-tilled field on the 28th of May and to both fields on the 9th of August and tillage occurred on the 1st of May
 289 (black dashed vertical lines).

290 4 Discussion

291 4.1 The influence of tillage on N₂O fluxes

292 The comparison of pre-tillage and post-tillage fluxes from the tilled field suggests that the tillage event was
 293 directly responsible for an immediate increase in N₂O fluxes (Figures 5 & 8). N₂O fluxes significantly larger than
 294 those measured pre-tillage were observed from the tilled field over two separate periods during which no changes
 295 in N₂O fluxes were observed in the adjacent un-tilled field. The initial increase in N₂O flux from the tilled field is
 296 a short lived peak which occurs directly after the disturbance of the soil caused by ploughing and harrowing. The
 297 second is a sustained increase which is observed throughout May and June. In the two month period in which
 298 fluxes from the tilled field were elevated, a total of 1.26 ± 0.12 kg N₂O-N ha⁻¹ was estimated to have been released.
 299 Assuming fluxes in the tilled field had remained at approximately pre-tillage magnitude had the tillage event not
 300 taken place (~ 0.27 nmol m⁻² s⁻¹, based on an average of fluxes measurements before the tillage event), it can be
 301 concluded that the tillage event contributed to an additional 0.85 ± 0.11 kg N₂O-N ha⁻¹ emitted from the field over
 302 a two month period.

303 Increases in N₂O flux lasting up to two months after grassland tillage events have been observed before
304 in other studies using both static chamber and eddy covariance measurements (Chatskikh and Olesen, 2007;
305 Merbold et al., 2014). Reported fluxes can be relatively high over a sustained period of time (several days or
306 weeks) and similar in magnitude to those recorded after fertilisation events. The mechanisms driving these large
307 sustained fluxes are believed to be partly due to the mineralisation of organic materials in the soils (decaying grass
308 materials from the previous sward in tilled grasslands) (Baggs et al., 2003; Hellebrand, 1998; Pimentel et al.,
309 2015). The large quantities of decaying organic matter ploughed into the soils would have provided a gradual
310 release of carbon and nitrogen into the soils, which provide substrate for the microbial processes of nitrification
311 and denitrification (Pimentel et al., 2015; Seastedt et al., 1992). According to IPCC estimates, 1 % of N added to
312 soils in the form of crop residues can be expected to be released as N₂O (IPCC, 2006). Based on our pre-tillage
313 biomass measurements made prior to tillage (93.6 kg N ha⁻¹), we would expect to see N₂O fluxes of approximately
314 0.94 kg N₂O-N ha⁻¹ from the field. This estimated value is within the range of uncertainty of our calculated
315 cumulative fluxes in this study (0.85 ± 0.11 kg N₂O-N ha⁻¹). High emissions from crop residues tilled into arable
316 crops have been recorded in similar wet soils with high clay content (Ball, 1999) which may indicate a similar
317 process is occurring under these conditions at other field sites in the area.

318 Large N₂O fluxes (> 0.5 nmol m⁻² s⁻¹) are observed from both fields after fertilisation events. Elevated
319 fluxes recorded from the fields after fertilisation typically last three to four weeks with an occasional large spike
320 lasting 24 to 48 hours before returning to pre-fertilisation levels. This month long period in which the majority of
321 large fluxes occur after fertilisation is also generally observed by other similar studies from the local area (Skiba
322 et al., 2013; Smith et al., 2012). Assuming the majority of N₂O emitted after a fertilisation event occurs within a
323 28 day period after the fertiliser application, the 28 day cumulative flux emissions associated with the fertilisation
324 events on the 28th of May and 9th of August on the un-tilled field were 0.55 ± 0.05 and 0.76 ± 0.24 kg N₂O-N ha⁻¹,
325 respectively. This equates to 0.79 and 1.09 % of the total nitrogen applied, respectively. The 28 day cumulative
326 flux emissions associated with the fertilisation event on the tilled field was 0.77 ± 0.34 kg N₂O-N ha⁻¹, or 1.10 %
327 of the total nitrogen applied. Assuming the 28 day periods account well for the emission factors of the fertiliser
328 events, these results are well within the range of uncertainty of the generic 1 (0.3 to 3.0) % value reported by the
329 IPCC for N fertiliser events (IPCC, 2014).

330

331 4.2 Gap filling of N₂O fluxes

332 Gap-filling N₂O flux measurements is challenging due to the lack of reliable process-based models on which to
333 base predictions. N₂O fluxes are believed to be driven primarily by the availability of nitrogen compounds in the
334 soils (ammonium and nitrate) (Davidson et al., 2000) as well as physical properties of the soil such as WFPS,
335 aerobic extent, soil type, temperature and compaction (Ball et al., 2008; Butterbach-Bahl et al., 2013; Choudhary
336 et al., 2002; Davidson et al., 2000; Turner et al., 2008). The collection of these data on a temporal/spatial scale
337 which would allow these models to be applied is not often logistically possible or affordable. The GAM method
338 used in this study incorporates readily-available meteorological data with the temporal pattern in the data, to
339 provide an empirical but practical means of temporal interpolation, which makes use of more information than
340 simple linear interpolation. Although the GAM method has proved useful, we would also emphasise the dangers
341 of extrapolating to conditions beyond those to which the model was fitted. For example, as we have not measured
342 fluxes during the cold months in winter, the GAM is unable to reliably predict fluxes in temperatures lower than
343 those measured during the study. The method deals appropriately with the large uncertainties where measurement
344 data are unavailable, contributing considerably to the total uncertainty in cumulative flux estimates.

345 In this study, spatial variability was not explicitly accounted for in the cumulative flux uncertainty and
346 this remains a potentially large error if extrapolating to areas larger than the measurement footprint. Eddy
347 covariance is able to integrate over a large area of the field (several 100 m²) (Eugster and Merbold, 2015) but
348 these measurements are still subject to an element of spatial variability which is difficult to fully account for given
349 the spatially heterogeneous nature of N₂O fluxes. Any study which plans to report cumulative flux estimates
350 should consider how to minimise the uncertainties which arise when interpolating and/or extrapolating
351 measurements to larger temporal and spatial scales (e.g. from occasional chamber measurements to annual field-
352 scale emissions). Further studies may require more complex statistical analysis, using methods such as Bayesian
353 statistics, to properly quantify the uncertainty in estimates of cumulative fluxes over large areas.

354

355 5 Conclusion

356 N₂O emissions from the grassland field after the tillage event were relatively large and sustained, similar in
357 magnitude to a nitrogen fertilisation event. The tillage event in this study is estimated to be responsible for a
358 period of high and sustained N₂O emissions lasting over a two month period after tillage (0.85 ± 0.11 kg N₂O-N

359 ha⁻¹), with a cumulative flux value akin to an 85 Kg-N fertiliser application according to IPCC emission factor
360 estimates. Relatively little difference in N₂O fluxes were observed between the tilled and un-tilled fields after a
361 subsequent identical application of nitrogen fertiliser in August 2012. Our results agree with several other similar
362 studies that tillage and the resultant addition of crop residues into soils can result in significant emissions of N₂O,
363 similar in magnitude to 1 % of the nitrogen available in those residues (0.9 % in this study). This study highlights
364 that the tillage of grassland fields can potentially result in a short term but significant increase in emissions of
365 N₂O, with the potential to affect regional or national greenhouse gas budgets.

366 **6 Acknowledgements**

367 We thank farm manager Wim Bosma, for the Easter Bush field site, who provided us with the opportunity to carry
368 out this experiment. We thank DEFRA and the UK Devolved Administrations for financial support through the
369 UK GHG Platform project AC0116 (The InveN2Ory project). We also thank the INGOS EU funded Integrating
370 Activity for support of the field infrastructure.

371

372 **7 References**

373 Abdalla, M., Jones, M., Ambus, P. and Williams, M.: Emissions of nitrous oxide from Irish arable soils: effects
374 of tillage and reduced N input, *Nutr. Cycl. Agroecosystems*, 86(1), 53–65, doi:10.1007/s10705-009-9273-8,
375 2010.

376 Baggs, E. M., Stevenson, M., Pihlatie, M., Regar, A., Cook, H. and Cadisch, G.: Nitrous oxide emissions
377 following application of residues and fertiliser under zero and conventional tillage, *Plant Soil*, 254(2), 361–370,
378 doi:10.1023/A:1025593121839, 2003.

379 Ball, B.: Soil and residue management effects on arable cropping conditions and nitrous oxide fluxes under
380 controlled traffic in Scotland 1. Soil and crop responses, *Soil Tillage Res.*, 52(3–4), 177–189,
381 doi:10.1016/S0167-1987(99)00080-X, 1999.

382 Ball, B., Crichton, I. and Horgan, G.: Dynamics of upward and downward N₂O and CO₂ fluxes in ploughed or
383 no-tilled soils in relation to water-filled pore space, compaction and crop presence, *Soil Tillage Res.*, 101(1–2),
384 20–30, doi:10.1016/j.still.2008.05.012, 2008.

385 Boeckx, P., Van Nieuland, K. and Van Cleemput, O.: Short-term effect of tillage intensity on N₂O and CO₂
386 emissions, *Agron. Sustain. Dev.*, 31(3), 453–461, doi:10.1007/s13593-011-0001-9, 2011.

387 Bouwman, A. F., Boumans, L. J. M. and Batjes, N. H.: Modeling global annual N₂O and NO emissions from
388 fertilized fields: N₂O AND NO EMISSIONS FROM FERTILIZERS, *Glob. Biogeochem. Cycles*, 16(4), 28-1-
389 28–9, doi:10.1029/2001GB001812, 2002.

390 Burba, G., Schmidt, A., Scott, R. L., Nakai, T., Kathilankal, J., Fratini, G., Hanson, C., Law, B., McDermitt, D.
391 K., Eckles, R., Furtaw, M. and Velgersdyk, M.: Calculating CO₂ and H₂O eddy covariance fluxes from an
392 enclosed gas analyzer using an instantaneous mixing ratio, *Glob. Change Biol.*, 18(1), 385–399,
393 doi:10.1111/j.1365-2486.2011.02536.x, 2012.

- 394 Butterbach-Bahl, K., Baggs, E. M., Dannenmann, M., Kiese, R. and Zechmeister-Boltenstern, S.: Nitrous oxide
 395 emissions from soils: how well do we understand the processes and their controls?, *Philos. Trans. R. Soc. B*
 396 *Biol. Sci.*, 368(1621), 20130122–20130122, doi:10.1098/rstb.2013.0122, 2013.
- 397 Chatskikh, D. and Olesen, J. E.: Soil tillage enhanced CO₂ and N₂O emissions from loamy sand soil under
 398 spring barley, *Soil Tillage Res.*, 97(1), 5–18, doi:10.1016/j.still.2007.08.004, 2007.
- 399 Choudhary, M. ., Akramkhanov, A. and Saggari, S.: Nitrous oxide emissions from a New Zealand cropped soil:
 400 tillage effects, spatial and seasonal variability, *Agric. Ecosyst. Environ.*, 93(1–3), 33–43, doi:10.1016/S0167-
 401 8809(02)00005-1, 2002.
- 402 Cowan, N. J., Famulari, D., Levy, P. E., Anderson, M., Bell, M. J., Rees, R. M., Reay, D. S. and Skiba, U. M.:
 403 An improved method for measuring soil N₂O fluxes using a quantum cascade laser with a dynamic chamber,
 404 *Eur. J. Soil Sci.*, 65(5), 643–652, doi:10.1111/ejss.12168, 2014a.
- 405 Cowan, N. J., Famulari, D., Levy, P. E., Anderson, M., Reay, D. S. and Skiba, U. M.: Investigating uptake of
 406 N₂O in agricultural soils using a high-precision dynamic chamber method, *Atmospheric Meas. Tech.*, 7(12),
 407 4455–4462, doi:10.5194/amt-7-4455-2014, 2014b.
- 408 Cowan, N. J., Norman, P., Famulari, D., Levy, P. E., Reay, D. S. and Skiba, U. M.: Spatial variability and
 409 hotspots of soil N₂O fluxes from intensively grazed grassland, *Biogeosciences*, 12(5), 1585–1596,
 410 doi:10.5194/bg-12-1585-2015, 2015.
- 411 Davidson, E. A., Keller, M., Erickson, H. E., Verchot, L. V. and Veldkamp, E.: Testing a Conceptual Model of
 412 Soil Emissions of Nitrous and Nitric Oxides: Using two functions based on soil nitrogen availability and soil
 413 water content, the hole-in-the-pipe model characterizes a large fraction of the observed variation of nitric oxide
 414 and nitrous oxide emissions from soils, *BioScience*, 50(8), 667–680, doi:10.1641/0006-
 415 3568(2000)050[0667:TACMOS]2.0.CO;2, 2000.
- 416 DEFRA, *Agriculture in the United Kingdom (Report)*. UK government, 2012.
- 417 Dobbie, K. E., McTaggart, I. P. and Smith, K. A.: Nitrous oxide emissions from intensive agricultural systems:
 418 Variations between crops and seasons, key driving variables, and mean emission factors, *J. Geophys. Res.*,
 419 104(D21), 26891, doi:10.1029/1999JD900378, 1999.
- 420 Drewer, J., Anderson, M., Levy, P.E., Scholtes, B. Helfter, C., Parker, J., Rees, R.M., Skiba, U.M.: The impact
 421 of ploughing intensively managed temperate grasslands on N₂O, CH₄ and CO₂ fluxes, *Plant & Soil*, submitted
 422 13/06/16.
- 423 Elmi, A., Madramootoo, C., Hamel, C. and Liu, A.: Denitrification and nitrous oxide to nitrous oxide plus
 424 dinitrogen ratios in the soil profile under three tillage systems, *Biol. Fertil. Soils*, 38(6), 340–348,
 425 doi:10.1007/s00374-003-0663-9, 2003.
- 426 Eugster, W. and Merbold, L.: Eddy covariance for quantifying trace gas fluxes from soils, *SOIL*, 1(1), 187–205,
 427 doi:10.5194/soil-1-187-2015, 2015.
- 428 Finkelstein, P. L. and Sims, P. F.: Sampling error in eddy correlation flux measurements, *J. Geophys. Res.*
 429 *Atmospheres*, 106(D4), 3503–3509, doi:10.1029/2000JD900731, 2001.
- 430 Fowler, D., Coyle, M., Skiba, U., Sutton, M. A., Cape, J. N., Reis, S., Sheppard, L. J., Jenkins, A., Grizzetti, B.,
 431 Galloway, J. N., Vitousek, P., Leach, A., Bouwman, A. F., Butterbach-Bahl, K., Dentener, F., Stevenson, D.,
 432 Amann, M. and Voss, M.: The global nitrogen cycle in the twenty-first century, *Philos. Trans. R. Soc. B Biol.*
 433 *Sci.*, 368(1621), 20130164–20130164, doi:10.1098/rstb.2013.0164, 2013.
- 434 Hellebrand, H. J.: Emission of Nitrous Oxide and other Trace Gases during Composting of Grass and Green
 435 Waste, *J. Agric. Eng. Res.*, 69(4), 365–375, doi:10.1006/jaer.1997.0257, 1998.
- 436 Intergovernmental Panel on Climate Change, Ed.: *Climate Change 2013 - The Physical Science Basis: Working*
 437 *Group I Contribution to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*,

- 438 Cambridge University Press, Cambridge. [online] Available from:
439 <http://ebooks.cambridge.org/ref/id/CBO9781107415324> (Accessed 24 November 2015), 2014.
- 440 (Intergovernmental Panel on Climate Change, I.: IPCC guidelines for national greenhouse gas inventories.
441 Chapter 11: N₂O Emissions from Managed Soils, and CO₂ Emissions from Lime and Urea Application., [online]
442 Available from: http://www.ipcc-nggip.iges.or.jp/public/2006gl/pdf/4_Volume4/V4_11_Ch11_N2O&CO2.pdf,
443 2006.
- 444 Jahangir, M., Roobroeck, D., Van Cleemput, O. and Boeckx, P.: Spatial variability and biophysicochemical
445 controls on N₂O emissions from differently tilled arable soils, *Biol. Fertil. Soils*, 47(7), 753–766,
446 doi:10.1007/s00374-011-0580-2, 2011.
- 447 Jones, S. K., Famulari, D., Di Marco, C. F., Nemitz, E., Skiba, U. M., Rees, R. M. and Sutton, M. A.: Nitrous
448 oxide emissions from managed grassland: a comparison of eddy covariance and static chamber measurements,
449 *Atmospheric Meas. Tech. Discuss.*, 4(1), 1079–1112, doi:10.5194/amtd-4-1079-2011, 2011.
- 450 Kormann, R. and Meixner, F. X.: An Analytical Footprint Model For Non-Neutral Stratification, *Bound.-Layer*
451 *Meteorol.*, 99(2), 207–224, doi:10.1023/A:1018991015119, 2001.
- 452 Levy, P. E., Gray, A., Leeson, S. R., Gaiawyn, J., Kelly, M. P. C., Cooper, M. D. A., Dinsmore, K. J., Jones, S.
453 K. and Sheppard, L. J.: Quantification of uncertainty in trace gas fluxes measured by the static chamber method,
454 *Eur. J. Soil Sci.*, 62(6), 811–821, doi:10.1111/j.1365-2389.2011.01403.x, 2011.
- 455 Mathieu, O., Leveque, J., Henault, C., Milloux, M., Bizouard, F. and Andreux, F.: Emissions and spatial
456 variability of N₂O, N₂ and nitrous oxide mole fraction at the field scale, revealed with 15N isotopic techniques,
457 *Soil Biol. Biochem.*, 38(5), 941–951, doi:10.1016/j.soilbio.2005.08.010, 2006.
- 458 Mauder, M. and Foken, T.: Impact of post-field data processing on eddy covariance flux estimates and energy
459 balance closure, *Meteorol. Z.*, 15(6), 597–609, doi:10.1127/0941-2948/2006/0167, 2006.
- 460 Merbold, L., Eugster, W., Stieger, J., Zahniser, M., Nelson, D. and Buchmann, N.: Greenhouse gas budget (CO₂
461 , CH₄ and N₂O) of intensively managed grassland following restoration, *Glob. Change Biol.*, 20(6), 1913–1928,
462 doi:10.1111/gcb.12518, 2014.
- 463 Moncrieff, J. B., Massheder, J. M., de Bruin, H., Elbers, J., Friborg, T., Heusinkveld, B., Kabat, P., Scott, S.,
464 Soegaard, H. and Verhoef, A.: A system to measure surface fluxes of momentum, sensible heat, water vapour
465 and carbon dioxide, *J. Hydrol.*, 188–189, 589–611, doi:10.1016/S0022-1694(96)03194-0, 1997.
- 466 Morton, D., Rowland, C., Wood, C., Meek, L., Marston, C., Smith, G., Wadsworth, R., Simpson, I.C.: Final
467 Report for LCM2007 - the new UK Land Cover Map. Countryside Survey Technical Report No. 11/07
468 NERC/Centre for Ecology and Hydrology 112pp. (CEH Project Number: C03259), 2011.
- 469 Omonode, R. A., Smith, D. R., Gál, A. and Vyn, T. J.: Soil Nitrous Oxide Emissions in Corn following Three
470 Decades of Tillage and Rotation Treatments, *Soil Sci. Soc. Am. J.*, 75(1), 152, doi:10.2136/sssaj2009.0147,
471 2011.
- 472 Palma, R. M., Rímolo, M., Saubidet, M. I. and Conti, M. E.: Influence of tillage system on denitrification in
473 maize-cropped soils, *Biol. Fertil. Soils*, 25(2), 142–146, doi:10.1007/s003740050294, 1997.
- 474 Pedersen, A. R., Petersen, S. O. and Schelde, K.: A comprehensive approach to soil-atmosphere trace-gas flux
475 estimation with static chambers, *Eur. J. Soil Sci.*, 61(6), 888–902, doi:10.1111/j.1365-2389.2010.01291.x, 2010.
- 476 Pimentel, L. G., Weiler, D. A., Pedroso, G. M. and Bayer, C.: Soil N₂O emissions following cover-crop
477 residues application under two soil moisture conditions, *J. Plant Nutr. Soil Sci.*, 178(4), 631–640,
478 doi:10.1002/jpln.201400392, 2015.
- 479 Pinto, M., Merino, P., del Prado, A., Estavillo, J. M., Yamulki, S., Gebauer, G., Piertzak, S., Lauf, J. and
480 Oenema, O.: Increased emissions of nitric oxide and nitrous oxide following tillage of a perennial pasture, *Nutr.*
481 *Cycl. Agroecosystems*, 70(1), 13–22, doi:10.1023/B:FRES.0000049357.79307.23, 2004.

- 482 Seastedt, T. R., Parton, W. J. and Ojima, D. S.: Mass loss and nitrogen dynamics of decaying litter of
483 grasslands: the apparent low nitrogen immobilization potential of root detritus, *Can. J. Bot.*, 70(2), 384–391,
484 doi:10.1139/b92-052, 1992.
- 485 Seitzinger, S. P., Kroeze, C. and Styles, R. V.: Global distribution of N₂O emissions from aquatic systems:
486 natural emissions and anthropogenic effects, *Chemosphere - Glob. Change Sci.*, 2(3–4), 267–279,
487 doi:10.1016/S1465-9972(00)00015-5, 2000.
- 488 Sheehy, J., Six, J., Alakukku, L. and Regina, K.: Fluxes of nitrous oxide in tilled and no-tilled boreal arable
489 soils, *Agric. Ecosyst. Environ.*, 164, 190–199, doi:10.1016/j.agee.2012.10.007, 2013.
- 490 Skiba, U., Jones, S. K., Drewer, J., Helfter, C., Anderson, M., Dinsmore, K., McKenzie, R., Nemitz, E. and
491 Sutton, M. A.: Comparison of soil greenhouse gas fluxes from extensive and intensive grazing in a temperate
492 maritime climate, *Biogeosciences*, 10(2), 1231–1241, doi:10.5194/bg-10-1231-2013, 2013.
- 493 Smith, K., Dobbie, K., Thorman, R., Watson, C., Chadwick, D., Yamulki, S. and Ball, B.: The effect of N
494 fertilizer forms on nitrous oxide emissions from UK arable land and grassland, *Nutr. Cycl. Agroecosystems*,
495 93(2), 127–149, doi:10.1007/s10705-012-9505-1, 2012.
- 496 Tan, I., Vanes, H., Duxbury, J., Melkonian, J., Schindelbeck, R., Geohring, L., Hively, W. and Moebius, B.:
497 Single-event nitrous oxide losses under maize production as affected by soil type, tillage, rotation, and
498 fertilization, *Soil Tillage Res.*, 102(1), 19–26, doi:10.1016/j.still.2008.06.005, 2009.
- 499 Turner, D. A., Chen, D., Galbally, I. E., Leuning, R., Edis, R. B., Li, Y., Kelly, K. and Phillips, F.: Spatial
500 variability of nitrous oxide emissions from an Australian irrigated dairy pasture, *Plant Soil*, 309(1–2), 77–88,
501 doi:10.1007/s11104-008-9639-8, 2008.
- 502 Wood, S. N.: *Generalized additive models: an introduction with R*, Chapman & Hall/CRC, Boca Raton, FL.,
503 2006.
- 504 Yamulki, S. and Jarvis, S.: Short-term effects of tillage and compaction on nitrous oxide, nitric oxide, nitrogen
505 dioxide, methane and carbon dioxide fluxes from grassland, *Biol. Fertil. Soils*, 36(3), 224–231,
506 doi:10.1007/s00374-002-0530-0, 2002.