

## Interactive comment on “Equatorward phytoplankton migration during a cold spell within the Late Cretaceous supergreenhouse” by Niels A.G.M. van Helmond et al.

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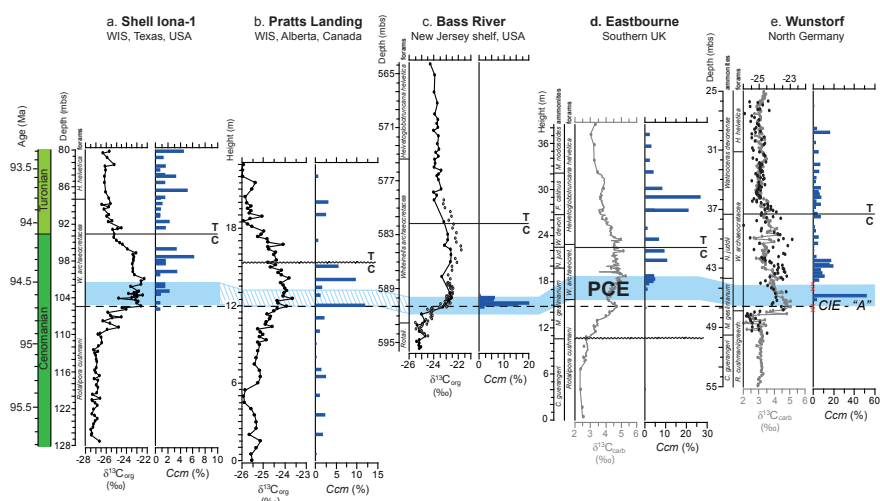
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In the original version of our manuscript the fonts of figure 6 are too small and therefore hard to read. We hereby upload an enlarged version of figure 6.

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**Figure 6.** Overview of  $\delta^{13}\text{C}_{\text{org}}$  and/or  $\delta^{13}\text{C}_{\text{carb}}$  abundances of *Cyclonephelium compactum-membraniphorum* morphological plexus (*Ccm*) and foraminiferal and/or ammonite zonation for the studied sections; (a) Shell Iona-1 core (Eldrett et al., 2014); (b) Pratts Landing (this study); (c) Bass River (van Helmond et al., 2014a) open symbols is  $\delta^{13}\text{C}_{\text{org}}$  derived from Bowman and Bralower (2005); (d) Eastbourne (Pearce et al., 2009), high-resolution  $\delta^{13}\text{C}_{\text{carb}}$  data derived from Paul et al. (1999); (e) Wunstorf – relative abundances of *Ccm* from van Helmond et al. (2015),  $\delta^{13}\text{C}_{\text{org}}$  from Du Vivier et al. (2014) and  $\delta^{13}\text{C}_{\text{carb}}$  from Voigt et al. (2008), a red cross marks a barren sample. Age is from the astronomically tuned age model for the Shell Iona-1 core (Eldrett et al., 2015). Dashed line represents the first maximum in the carbon isotope excursion, point “A” (cf. Voigt et al., 2008). Solid lines represent the Cenomanian–Turonian boundary. The blue shaded area represents the Plenus Cold Event according to its original definition (Gale and Christensen, 1996), the cooling in reconstructed sea surface temperatures at Bass River and Wunstorf (van Helmond et al., 2014a, 2015), and the (re)oxygenation of bottom waters in the Shell Iona-1 core (Eldrett et al., 2014). Note: the sections are plotted using different depth scales.

Fig. 1.

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