

1 **Impacts of extreme climate events and disturbances on**  
2 **carbon dynamics**

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15 **Abstract** The impacts of extreme climate events and disturbances (ECE&D) on the carbon  
16 cycle have received growing attention in recent years. This special issue showcases a  
17 collection of recent advances in understanding the impacts of ECE&D on carbon cycling.  
18 Notable advances include quantifying how harvesting activities impact forest structure,  
19 carbon pool dynamics, and recovery processes; observed drastic increases of the  
20 concentrations of dissolved organic carbon and dissolved methane in thermokarst lakes in  
21 western Siberia during a summer warming event; disentangling the roles of herbivores and  
22 fire on forest carbon dioxide flux; direct and indirect impacts of fire on the global carbon

23 balance; and improved atmospheric inversion of regional carbon sources and sinks by  
24 incorporating disturbances. Combined, studies herein indicate several major research needs.  
25 First, disturbances and extreme events can interact with one another, and it is important to  
26 understand their overall impacts and also disentangle their effects on the carbon cycle.  
27 Second, current ecosystem models are not skillful enough to correctly simulate the underlying  
28 processes and impacts of ECE&D (e.g., tree mortality and carbon consequences). Third,  
29 benchmark data characterizing the timing, location, type, and magnitude of disturbances must  
30 be systematically created to improve our ability to quantify carbon dynamics over large areas.  
31 Finally, improving the representation of ECE&D in regional climate/earth system models and  
32 accounting for the resulting feedbacks to climate are essential for understanding the  
33 interactions between climate and ecosystem dynamics.

34

## 35 **1 Introduction**

36 The biosphere plays an important role in regulating atmospheric carbon dioxide  
37 concentrations and thereby climate. Extreme climate events such as drought (Xiao et al.,  
38 2009; Zhao and Running, 2010) and disturbances such as fire (Bowman et al., 2009),  
39 hurricanes (Chambers et al., 2007; Dahal et al., 2014b; Xiao et al., 2011), wind storms  
40 (McCarthy et al., 2006), and insect outbreaks (Kurz et al., 2008a) can substantially alter  
41 ecosystem structure and function and influence terrestrial carbon dynamics. ECE&D are  
42 projected to increase in both frequency and severity during the remainder of the 21st century  
43 (IPCC, 2013), with important consequences for terrestrial carbon cycling. Projecting the  
44 impacts of these future events remains a challenge given the substantial uncertainty in  
45 forecasting these events and the insufficient representation of ECE&D in ecosystem and land  
46 surface models. A better understanding of the impacts of ECE&D on carbon dynamics across

47 different ecosystems is essential for projecting ecosystem responses to future climate change  
48 and feedbacks to the climate system.

49 Biospheric carbon fluxes often exhibit pronounced interannual variability (IAV) and  
50 ECE&D are believed to be primary sources of the IAV (Eimers et al., 2008; Reichstein et al.,  
51 2013; Xiao et al., 2014), which can be pronounced. For example, gross primary productivity  
52 (GPP) exhibited significant IAV over the period 2000-2014 on the global scale as identified  
53 by the MODIS GPP product (Zhao et al., 2005), with important regional differences (Fig. 1).  
54 The IAV is measured by the coefficient of variation (CV), defined as the standard deviation  
55 divided by the mean. Australia and southern Africa had the largest IAV; the U.S. Great  
56 Plains, the U.S. Southwest, Alaska, India, part of the Tibetan Plateau, eastern Mongolia,  
57 Kazakhstan, the Sahel region, and eastern Amazon had intermediate IAV; the remaining  
58 regions had relatively low IAV.

59 ECE&D can lead to exceptionally high or low annual carbon fluxes. We used the  
60 annual GPP data from the MODIS GPP product (Zhao et al., 2005) to identify extreme GPP  
61 values (outliers) that exceed the statistical normal range presumably caused by extreme  
62 climate events and/or disturbances (Fig. 2). For each grid cell, the outliers of annual GPP over  
63 the period 2000-2014 were identified using interquartile range (IQR) and quartiles (Q1: 25%  
64 quartile; Q3: 75% quartile). The outliers on the higher end were determined as values beyond  
65  $IQR + 1.5 \times Q3$ , and the outliers on the lower end were identified as values below  $IQR - 1.5$   
66  $\times Q1$ . Outliers on the lower end were observed in parts of Europe, Russia, North America, the  
67 Amazonia, and Africa (Fig. 2). These exceptionally low annual GPP were likely caused by  
68 drought, extreme low temperature, fire disturbance, or harvesting. Outliers on the higher end  
69 were observed in Alaska, the U.S. Southwest, Australia, and parts of the Amazonia and  
70 southern Africa (Fig. 2). These exceptionally high annual GPP were likely caused by

71 exceptionally moist conditions and/or warm temperatures. The U.S. Great Plains and  
72 Kazakhstan had large IAV and outliers on the lower end; part of Australia and southern  
73 Africa also exhibited large IAV but had outliers on the higher end; the large IAV of GPP did  
74 not correspond to outliers for other regions (Figs. 1 and 2). The IAV of carbon fluxes was  
75 likely driven by both outliers and moderate to strong anomalies in fluxes.

76 The impacts of ECE&D on carbon dynamics have received growing attention. We  
77 searched the number of journal articles on these topics using Web of Science (Fig. 3) and  
78 found a total of 497 and 1593 journal articles for extreme climate events and disturbances,  
79 respectively, over the period from 2000 to 2015. Notably, the annual number of publications  
80 on the impacts of these events on carbon dynamics has been growing at an average rate of 18  
81 articles per year from 2000 to 2015 and at an average rate of 25 articles per year over the past  
82 decade (2006-2015), emphasizing the growing scientific interest in these important topics.

83 Various approaches have been used to assess the impacts of ECE&D on ecosystem  
84 carbon dynamics. At the ecosystem scale, in-situ methods including field experiments  
85 (Barbeta et al., 2013), long-term observations (Turner et al., 2003), and the eddy covariance  
86 technique (Amiro et al., 2010; Schwalm et al., 2010) seek to understand the mechanisms  
87 underlying responses of ecosystem processes to ECE&D. Modeling approaches including  
88 process-based ecosystem models (Liu et al., 2011) or data-driven upscaling approaches (Jung  
89 et al., 2009; Xiao et al., 2008) have been used for regional to global assessments, which also  
90 rely heavily on satellite remote sensing (Xiao et al., 2014). Synthesizing these findings is an  
91 ongoing challenge, and multiple approaches are required to understand the consequences of  
92 different ECE&D for carbon cycling.

93 Spatially, the locations of the previous research activities have been largely aligned  
94 with the geography of the occurrence of ECE&D. For example, we have witnessed

95 pronounced impacts of insect outbreaks and fires in the northern Rocky Mountains (Hicke et  
96 al., 2012b; Kurz et al., 2008b; Law et al., 2004), the widespread deforestation in Amazon and  
97 other tropical regions (Achard et al., 2014; DeFries et al., 2002; Harris et al., 2012), peatland  
98 fires in Indonesia (Page et al., 2002; Turetsky et al., 2015), tropical cyclones in the United  
99 States (Dahal et al., 2014a), and drought and heat waves in Europe (Bréda et al., 2006; Ciais  
100 et al., 2005a; Reichstein et al., 2007) and the southwestern United States (Allen et al., 2010a;  
101 Carnicer et al., 2011; Zeppel et al., 2013). Temporally, most of the research has been on the  
102 impacts of individual ECE&D, with fewer studies involving long-term observations and  
103 monitoring records (Dahal et al., 2014a; Seidl et al., 2014). Abundant evidence has been  
104 collected globally in the past decades on increased tree mortality resulting from climate  
105 events such as prolonged mega droughts and heat waves (Allen et al., 2010a; McDowell,  
106 2011; Meddens et al., 2015; Meir et al., 2015). However, the mechanisms behind this  
107 increased mortality and the consequences on carbon dynamics still remain to be unveiled  
108 (Meddens et al., 2015; Meir et al., 2015).

109         The present special issue is the outcome of special sessions on the impacts of ECE&D  
110 on carbon dynamics at the American Geophysical Union Fall Meeting (2011-2013). This  
111 issue consists of 17 articles: 6 on extreme climate events and 11 on disturbances. This special  
112 issue, along with the special issue on climate extremes and biogeochemical cycles in  
113 *Biogeosciences* (Bahn et al., 2015), reflects recent advances in assessing how ECE&D  
114 influence terrestrial carbon cycling. We feel that the authors have provided a timely and  
115 valuable contribution to the research communities of carbon cycle and global change.

## 116 **2 Methods and Findings**

117 We highlight the findings in this special issue by grouping manuscripts that emphasize the  
118 impacts of drought and extreme precipitation events, herbivory (namely insect outbreaks),

119 fire, interactions between herbivory and fire, natural hazards (e.g. hurricanes and typhoons),  
120 and forest management.

121 *Drought and extreme precipitation events*

122 Piayda et al. (2014) quantified the impacts of the extreme drought event in 2012 on carbon  
123 and water cycling in a Mediterranean woodland. The drought reduced overstory GPP in 2012  
124 by 28% and carbon sink strength by 38% compared to 2011. Results indicated that successful  
125 simulation of drought effects on the montado ecosystem requires the incorporation of variable  
126 apparent maximum carboxylation rate, stomatal conductance, and vapor pressure deficit  
127 sensitivity into photosynthesis-stomatal conductance modeling.

128         The simulations of a process-based ecosystem model showed that drought from 2000  
129 to 2011 led to significant reduction in both GPP and net ecosystem productivity (NEP) of  
130 China's terrestrial ecosystems at regional to national scales (Liu et al., 2014). Relative to the  
131 long-term mean, the nationwide annual NEP in 2001, 2006, 2009, and 2011 decreased by *ca.*  
132 63, 88, 170, and 61 Tg C yr<sup>-1</sup>, respectively, due to droughts (Liu et al., 2014). These two  
133 studies were consistent with several previous synthesis and modeling studies indicating that  
134 severe droughts could reduce annual GPP and net ecosystem productivity (NEP), and the  
135 reduction in NEP was largely driven by the decrease in GPP due largely to reductions in GPP  
136 (Ciais et al., 2005b; Schwalm et al., 2010; Xiao et al., 2009).

137         The opposite of drought – extreme precipitation events - have received less attention  
138 in carbon cycle studies. Jiang et al. (2013) conducted a field experiment in three subtropical  
139 forests to study the responses of soil respiration to both drought and extreme high  
140 precipitation and found that altered precipitation strongly influenced soil respiration not only  
141 by controlling soil moisture but also by modifying moisture and temperature sensitivity of soil  
142 respiration. Their results indicate that soil respiration was more sensitive to soil moisture in

143 the presence of drought, and higher precipitation in the wet season could have a limited effect  
144 on the response of soil respiration to rising temperatures (Jiang et al., 2013).

145 Zeppel, Wilks, and Lewis (2014) reviewed studies of extreme precipitation and  
146 seasonal changes in precipitation on carbon metabolism in grassland and forested ecosystems.  
147 They found that extremely high precipitation is likely to increase aboveground net primary  
148 productivity (ANPP) of xeric biomes and reduce ANPP of mesic biomes. Changes in  
149 precipitation during the growing season are likely to have a greater impact on carbon cycle  
150 dynamics than precipitation during the non-growing season (Zeppel et al., 2014). These  
151 studies indicated that the direction and magnitude of the impacts of extreme precipitation  
152 events on carbon fluxes depend on the season (wet versus dry) and biome type (xeric versus  
153 mesic).

#### 154 *Extreme temperature events*

155 Extreme temperature events have been a feature of recent climate change, especially at high  
156 latitudes (IPCC, 2013). Previous studies showed that extreme temperature events often reduce  
157 GPP and NEP of terrestrial ecosystems (Ciais et al., 2005b; Qu et al., 2016). The effects of  
158 extreme temperature on the carbon dynamics of aquatic ecosystems, however, have received  
159 little attention. Pokrovsky et al. (2013) studied the impacts of the 5 – 15 °C summer warming  
160 event of 2012 on the carbon dynamics of thermokarst lakes in western Siberia. Dissolved  
161 organic carbon concentrations increased by a factor of two as a result of the warming event  
162 despite limited changes in conductivity and pH, and the concentration of dissolved methane  
163 increased by nearly fivefold (Pokrovsky et al., 2013). These results demonstrate a substantial  
164 increase in the methane emission capacity from lakes as a result of summertime warming in  
165 areas of permafrost thaw.

166 De Simon et al. (2013) examined the effects of manipulated warmer or cooler late  
167 winter/early spring conditions on the carbon budget and yield of soybean crops. Their results  
168 demonstrate that extreme temperature events in late winter did not result in significant  
169 changes in the net carbon balance (De Simon et al., 2013). These events may have larger  
170 impacts on natural ecosystems by advancing or delaying leaf-out dates.

171 Combined, these studies indicate that the effects of extreme temperature events on  
172 ecosystem carbon dynamics depend on the timing and magnitude of these events. Extreme  
173 temperature events occurring in the growing season could substantially alter carbon fluxes,  
174 while those events occurring during the remainder of the year had smaller effects than  
175 expected.

#### 176 *Insect outbreaks*

177 The coniferous forests of western North America have experienced an unprecedented  
178 herbivore outbreak over millions of hectares over the past decades (Hicke et al., 2012a; Raffa  
179 et al., 2008), part of the global tree die-off due to the combined effects of elevated  
180 temperatures, drought, and associated herbivory (Allen et al., 2010b). Measurements of the  
181 impacts of this disturbance at the site scale find minimal ecosystem carbon loss or even net  
182 uptake shortly after eruptive herbivory (Brown et al., 2010), which contrasts regional  
183 estimates of substantial carbon losses to the atmosphere (Ghimire et al., 2015; Kurz et al.,  
184 2008a). Mathys et al. (2013) in this issue used the eddy covariance technique to study carbon  
185 dioxide flux after a mountain pine beetle (*Dendroctonus ponderosae*, Hopkins) attack over a  
186 two-year period and compared these to an adjacent clearcut. They found that the mountain  
187 pine beetle-damaged forest was a carbon sink of *ca.* 50 g C m<sup>-2</sup> year<sup>-1</sup> two years after attack.  
188 This study also indicates that the residual forest and the understory vegetation contributed to  
189 carbon uptake and could enable the forest to return to carbon neutrality at a faster rate than  
190 clearcuts. The impacts of herbivore outbreak depend on the type of herbivore (e.g. foliivores



191 *versus* phloem-feeders) and the intensity of disturbance (Allen et al., 2010b; Brown et al.,  
192 2010; Ghimire et al., 2015; Hicke et al., 2012a; Kurz et al., 2008a; Mathys et al., 2013; Raffa  
193 et al., 2008).

#### 194 *Fire*

195 Fire causes direct and immediate carbon emissions into the atmosphere from biomass burning  
196 (the direct effect), and subsequent changes in NEP (the indirect effect) through changes in  
197 GPP and ecosystem respiration of the remaining live stand and the heterotrophic respiration  
198 of the damaged biomass. Li et al. (2014) in this special issue provided a quantitative  
199 assessment of the direct and indirect impacts of fire on the net carbon balance of global  
200 terrestrial ecosystems during the 20th century. Their results show that fire decreased the net  
201 carbon gain of global terrestrial ecosystems by  $1.0 \text{ Pg C yr}^{-1}$  averaged across the 20th century,  
202 as a result of the fire direct effect ( $1.9 \text{ Pg C yr}^{-1}$ ) partly offset by the indirect effect ( $-0.9 \text{ Pg C}$   
203  $\text{yr}^{-1}$ ). The effect of fire on the net carbon balance significantly declined until 1970 with a  
204 trend of  $8 \text{ Tg C yr}^{-1}$  due to an increasing indirect effect, and increased subsequently with a  
205 trend of  $18 \text{ Tg C yr}^{-1}$  due to an increasing direct effect (Li et al., 2014). These results help  
206 constrain the global-scale dynamics of fire and the terrestrial carbon cycle.

#### 207 *Insect outbreaks versus fire*

208 At the regional scale, Caldwell et al. (2013) simulated and evaluated the long-term impacts of  
209 the two characteristic disturbances in the Southern Rocky Mountains forests (i.e., the outbreak  
210 of mountain pine beetle and high-severity wildfire) on changes in species composition and  
211 carbon stocks. Wildfire caused larger changes in both patterns of succession and distribution  
212 of carbon among biomass pools than did mountain pine beetle disturbance; carbon in standing  
213 live biomass returned to pre-disturbance levels after 50 versus 40 years following wildfire and  
214 mountain pine beetle disturbances, respectively (Caldwell et al., 2013).

215 Clark et al. (2014) used the eddy covariance technique to study the impacts of fire and  
216 gypsy moth (*Lymantria dispar* L.) disturbance in oak-dominated, pine-dominated, and mixed  
217 forests in eastern North America. The net ecosystem exchange (NEE), GPP, and water use  
218 efficiency were of greater magnitude in the oak-dominated forest before disturbance during  
219 summer. Water use efficiency declined by 60% at the oak-dominated stand and by nearly  
220 50% at the mixed stand after gypsy moth disturbance, but prescribed fire had little impact on  
221 water use efficiency in the mixed or pine stands (Clark et al., 2014). These results  
222 demonstrate the importance of forest type, disturbance type, and time since disturbance on  
223 coupled carbon and water cycle functioning in temperate forests.

#### 224 *Hurricanes and typhoons*

225 Hurricane events in the U.S. have significant effects on regional carbon dynamics (Dahal et  
226 al., 2014b). Typhoons are natural disturbances to subtropical mangrove forests in Asia, and  
227 their effects on ecosystem carbon dynamics of mangroves are not well understood. Chen et al.  
228 (2014) examined the short-term effects of frequent strong typhoons on defoliation and the  
229 NEE of subtropical mangroves. The responses of daily NEE following typhoons were highly  
230 variable in mangrove ecosystems (Chen et al., 2014), demonstrating that the characteristics of  
231 the typhoon and antecedent ecosystem conditions are important for understanding hurricane  
232 impacts on carbon stocks and fluxes. Severe hurricanes and typhoons that destroy a large  
233 number of trees could have significant effects on regional carbon cycling, while those that  
234 lead merely to defoliation likely had transient effects on ecosystem carbon exchange.

#### 235 *Forest management*

236 Accurate quantification of the effects of partial cutting or clearcutting is essential for a better  
237 understanding of forest carbon dynamics and for informing forest management. Zhou et al.  
238 (2013a) conducted a meta-analysis on the impacts of partial cutting (i.e., cutting events with

239 aboveground biomass removal rate < 90%) on forest carbon stocks by collecting data on  
240 cutting intensity, forest structure, and carbon stock components. This is a global-scale meta-  
241 analysis, but the majority of the sites are distributed in the U.S. and Europe. The results  
242 showed that partial cutting reduced aboveground carbon by 43% and increased understory  
243 carbon storage by nearly 400% on average, but did not have significant effects on forest floor  
244 or mineral soil carbon stocks (Zhou et al., 2013a). This effort provides a new perspective on  
245 the impacts of forest harvesting as it covers the spectrum of harvest disturbances from partial  
246 cutting to clearcut and goes beyond previous reviews that mostly concentrated on the impacts  
247 of clearcutting (Johnson and Curtis, 2001; Nave et al., 2010). The impacts of partial cutting  
248 can be significant; for example, partial cutting accounted for about three quarters of the total  
249 C loss from timber harvesting in the eastern United States from 2002 to 2010 (Zhou et al.,  
250 2013b).

251 Wang et al. (2014) used a process-based forest ecosystem model, PnET-CN, to  
252 evaluate how clearcutting alters ecosystem carbon fluxes, biomass, and leaf area index in  
253 northern temperate forests. They found that harvest disturbance in northern temperate forests  
254 had significant effects on forest carbon fluxes and stocks, and increased harvesting intensity  
255 would delay the recovery of NEP. Evergreen needleleaf forests were slower to recover to full  
256 carbon assimilation capacity after stand-replacing harvests than deciduous broadleaf forests  
257 (Wang et al., 2014). Future modeling studies of disturbance effects should incorporate forest  
258 population dynamics (e.g., regeneration and mortality) and relationships between age-related  
259 model parameters and state variables (e.g., leaf area index).

#### 260 *Disturbance legacy*

261 The time since disturbance is an important controlling factor of carbon dynamics. Berryman  
262 et al. (2013) tested the impacts of experimental pinyon pine (*Pinus edulis* Englem.) mortality

263 on microbial respiration. They found that litter respiration responded to water availability at  
264 both treatment and control sites, and that soil respiration decreased at the site with  
265 experimental mortality. These results demonstrate ecosystem-level consequences of tree  
266 mortality that differs as a function of water availability (Berryman et al., 2013).

267 Yue et al. (2013) compared observations from post-fire vegetation trajectories in the  
268 boreal forest with simulations from the process-based ORCHIDEE vegetation model and  
269 supported the notion that the increase in atmospheric CO<sub>2</sub> concentrations and vegetation  
270 recovery were jointly responsible for current carbon sink conditions. It should be noted that  
271 nitrogen deposition – a global change factor enhancing ecosystem carbon uptake was not  
272 explicitly considered, although the effects of nitrogen deposition carbon sink strength have  
273 been controversial (Magnani et al., 2007; Nadelhoffer et al., 1999). Nevertheless, their results  
274 highlight the importance of understanding how global change and disturbance events interact  
275 to determine current – and likely future – carbon cycle dynamics (Yue et al., 2013). These  
276 two studies demonstrate that the legacy of disturbance and environmental factors jointly  
277 control the carbon dynamics following disturbance.

278 Modeling approaches have been widely used to simulate ecosystem carbon dynamics  
279 following disturbance. Wang et al. (2014) in this special issue simulated the dynamics of  
280 carbon fluxes and stocks following harvest. The simulated NEP and aboveground carbon  
281 stock after clearcuts generally followed the hypothesized trajectories (Chapin, 2011; Odum,  
282 1969) while the decline in NEP was due to relatively stable GPP and gradually increasing  
283 ecosystem respiration (ER). Evergreen needleleaf forests recovered more slowly from a net  
284 carbon source to a net sink, and lost more carbon than deciduous broadleaf forests.

285 Disturbance-induced tree mortality regulates the forest carbon balance, but tree  
286 mortality and its carbon consequences are not well represented in ecosystem models (Bond-

287 Lamberty et al., 2015). Bond-Lamberty et al. (2015) tested whether three ecosystem models –  
288 the classic big-leaf model Biome-BGC and the gap-oriented models ZELIG and ED - could  
289 reproduce the resilience of forest ecosystems to moderate disturbances. The models replicated  
290 observed declines in aboveground biomass well but could not fully capture observed post-  
291 disturbance carbon fluxes. This study indicates that ecosystem models are yet unable to  
292 correctly simulate the effects of disturbances.

293         Lack of critical geospatial data on disturbances and associated impacts on ecosystems  
294 has been identified as one of the main challenges in quantifying carbon dynamics over large  
295 areas (Liu et al., 2011). Recently, a continental-scale forest stand age map was developed for  
296 North America using forest inventory data, large fire data, and remotely sensed data,  
297 providing a new source of information that can benefit quantification of the carbon sources  
298 and sinks across the continent and contribute to studies of disturbance (Pan et al., 2011). Deng  
299 et al. (2013) in this special issue used these continental stand age maps as an additional  
300 constraint to atmospheric CO<sub>2</sub> inversions. They found that regions with recently disturbed or  
301 old forests are often nudged towards carbon sources while regions with middle-aged  
302 productive forests are shifted towards sinks, confirming stand age effects observed from many  
303 eddy covariance flux towers (Deng et al., 2013). These results were generally consistent with  
304 the synthesis results from eddy covariance flux data across North America (Amiro et al.,  
305 2010) but they were inconsistent with some other studies showing that old-growth forests  
306 were still carbon sinks (Desai et al., 2005; Luysaert et al., 2008). At the sub-continental  
307 level, their inverted carbon fluxes agreed well with continuous estimates of NEE upscaled  
308 from eddy covariance flux data (Xiao et al., 2008; 2011). Recent development in  
309 characterizing the timing, location, type, and magnitude of disturbances (Huang et al., 2010;  
310 Kennedy et al., 2010; Masek et al., 2013; Williams et al., 2014; Zhu and Woodcock, 2014)  
311 are helping to advance diagnosis and monitoring of carbon dynamics over large areas.

### 312 **3 Conclusions**

313 The contributions of this special issue reflect some of the most recent advances in the impacts  
314 of ECE&D on carbon dynamics. These studies address the impacts of different types of  
315 extreme events including forest management, hurricanes and typhoons, drought, extreme  
316 precipitation events, extreme temperature events, insect outbreaks, and fire as well as  
317 ecosystem recovery since disturbance. The direction and magnitude of the effects of these  
318 events on ecosystem carbon fluxes depend on the nature of the events (type, duration, and  
319 intensity), the timing of the events (e.g., wet versus dry season, summer versus winter), and  
320 the biome type (e.g., xeric versus mesic). These events typically have negative effects on net  
321 carbon uptake while some events such as extreme precipitation events may also have positive  
322 effects on net carbon uptake depending on antecedent conditions and the nature of the  
323 extreme events.

324         Importantly, studies in this special issue collectively indicate several major research  
325 needs. First, ECE&D can interact with one another, and it is important to disentangle their  
326 relative effects on the carbon cycle. Second, current ecosystem models in general are not  
327 skillful enough to correctly simulate the impacts of ECE&D such as disturbance-induced tree  
328 mortality and its carbon consequences, and therefore ecosystem models must be improved to  
329 correctly represent the underlying processes and impacts (Liu et al., 2011; Reichstein et al.,  
330 2013). For example, the processes of drought effects on ecosystem respiration are not well  
331 represented in models. Third, the lack of data on major disturbances is still one of the main  
332 challenges that hinder the improvement of quantifying carbon dynamics over large areas, and  
333 benchmark data characterizing the timing, location, type, and magnitude of disturbances must  
334 be created. With the ongoing continuous monitoring of earth surface conditions using a  
335 constellation of satellites and emerging data mining technologies, the characterization and

336 understanding of the impacts of ECE&D are expected to improve drastically over the next 5  
337 to 10 years. However, major challenges still remain on how to translate those conditional  
338 changes into carbon fluxes and understand the specific roles of ECE&D in particular. Finally,  
339 besides carbon fluxes and stocks, other biogeophysical properties such as albedo,  
340 evapotranspiration (ET), and surface energy exchange are also altered by ECE&D. Improving  
341 the representation of ECE&D in regional climate/earth system models and accounting for the  
342 resulting feedbacks to the climate are essential for understanding the interactions between  
343 climate and ecosystem dynamics. Ongoing research in these areas will continue to improve  
344 our emerging understanding of the impacts of ECE&D on carbon cycling and the feedbacks to  
345 the climate.

346

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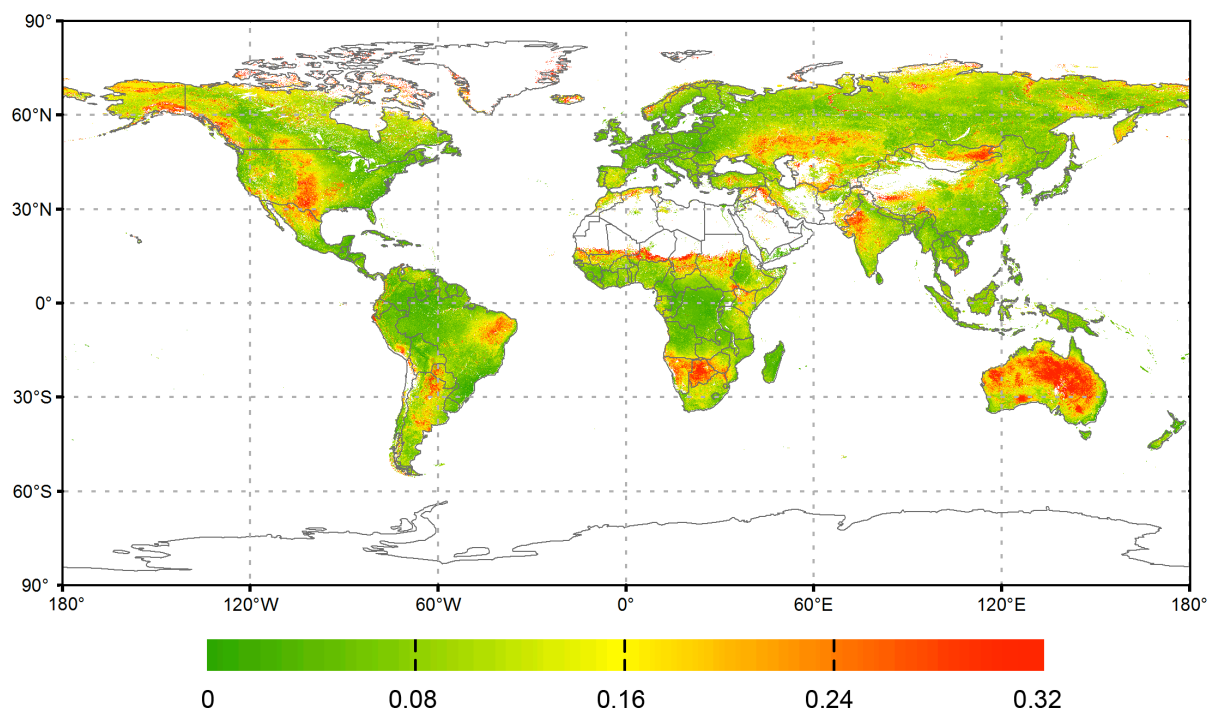
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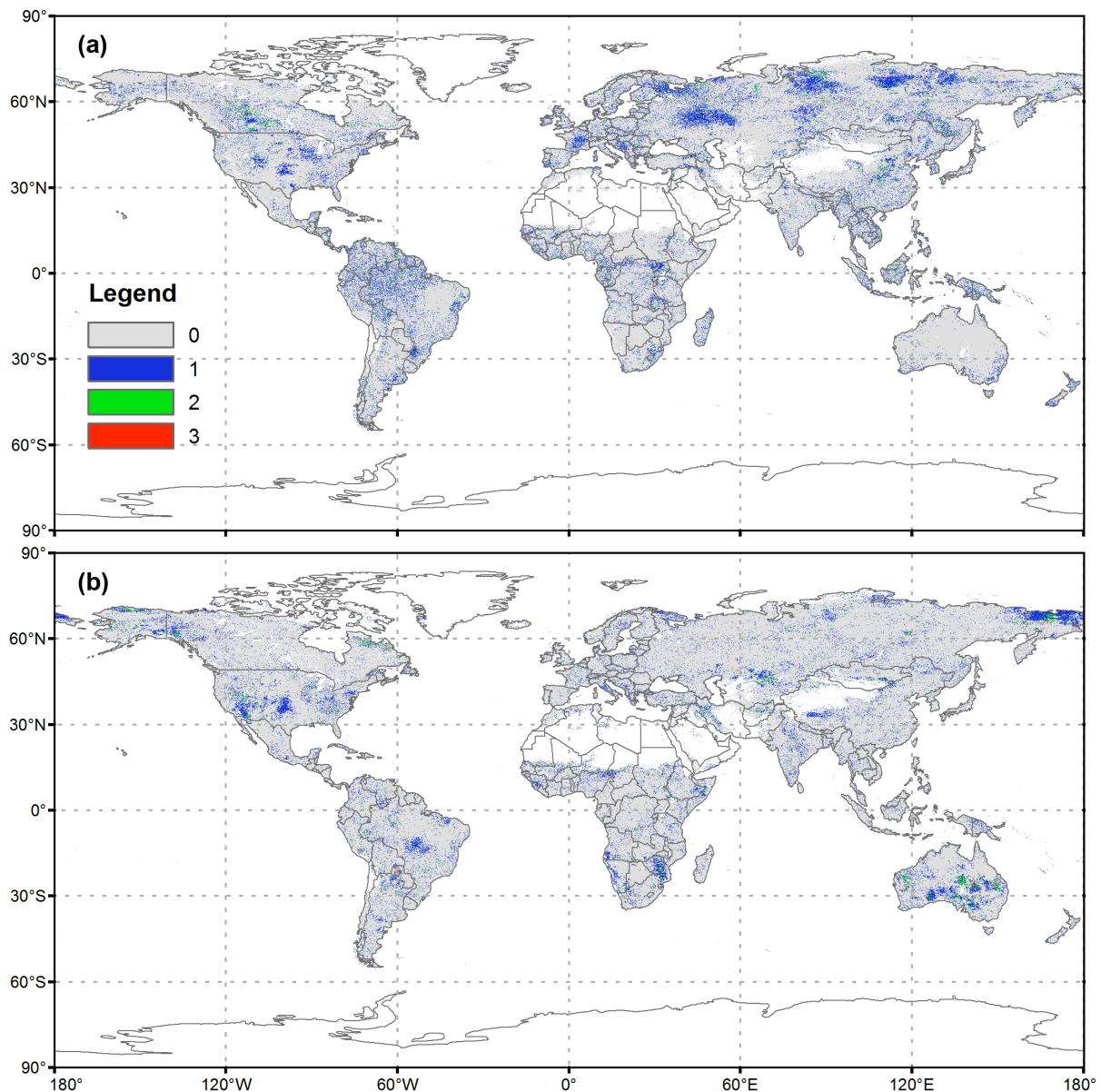
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635 **Fig. 1.** The interannual variability (i.e., the coefficient of variation or CV) of annual gross  
 636 primary productivity (GPP) over the period 2000-2014 from the MODIS GPP product  
 637 (MOD17A3). The CV is unitless.

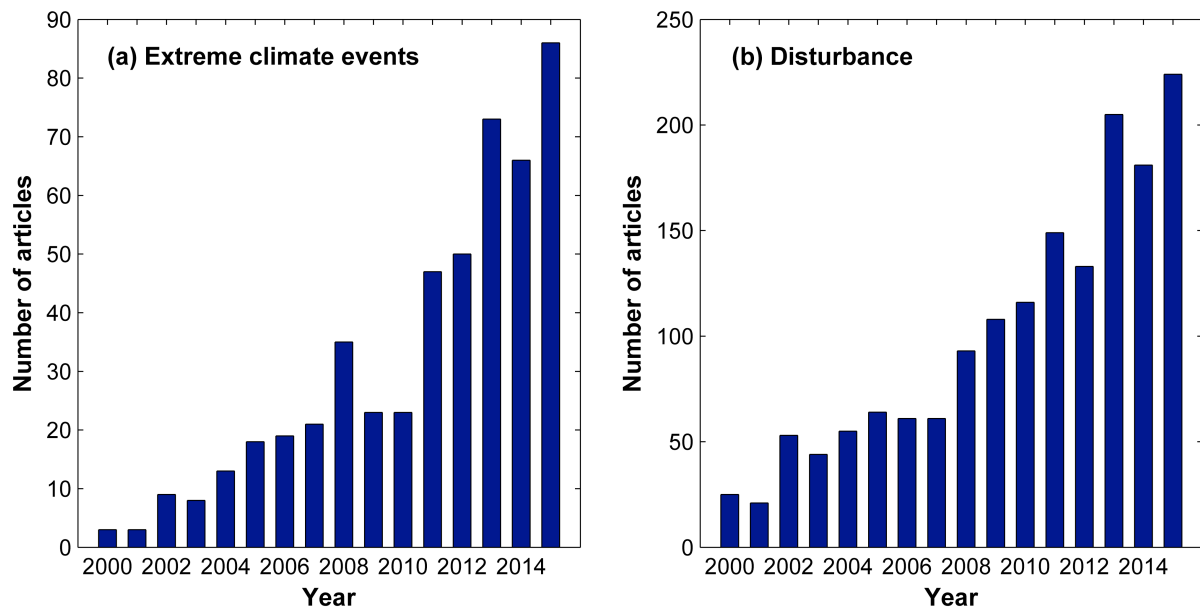
638



639

640 **Fig. 2.** The number of extreme years characterized by the number of outliers of the annual  
 641 gross primary productivity (GPP) distribution over the period 2000-2014: (a) outliers on the  
 642 lower end (i.e., exceptionally low annual GPP); (b) outliers on the higher end (i.e.,  
 643 exceptionally high annual GPP). The outliers were identified using interquartile range (IQR)  
 644 and quartiles (Q1: 25% quartile; Q3: 75% quartile) with the outliers on the higher end  
 645 determined as values beyond  $IQR + 1.5 \times Q3$  and the outliers on the lower end determined as  
 646 values below  $IQR - 1.5 \times Q1$ . The annual GPP values ( $g\ C\ m^{-2}\ year^{-1}$ ) were derived from the  
 647 MODIS GPP product (MOD17A3).  
 648





649

650 **Fig. 3.** The number of journal articles published over the period from 2000 to 2015 as  
 651 identified by Web of Science™ as of April 14, 2016 for the impacts of (a) extreme climate  
 652 events and (b) disturbance on carbon dynamics. The combination of key words that we used  
 653 to represent ‘extreme climate events’ is: TS=("extreme climate events" OR "climate  
 654 extremes" OR drought OR "extreme precipitation") AND TS=("carbon dynamics" OR  
 655 "carbon cycle" OR "carbon flux" OR "carbon stock" OR "carbon pool"), where TS stands for  
 656 Topic. The combination of key words used to represent ‘disturbance’ is: TS=(disturbance OR  
 657 fire OR harvesting OR logging OR hurricane or "insect outbreaks") AND TS=("carbon  
 658 dynamics" OR "carbon cycle" OR "carbon flux" OR "carbon stock" OR "carbon pool").