



1 **Preface: Impacts of extreme climate events and 2 disturbances on carbon dynamics**

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15 **Abstract** The impacts of disturbances and extreme climate on the carbon cycle have received
16 growing attention in recent years as evidenced by the increasing number of journal articles
17 published on these topics. This special issue showcases a collection of recent advancements in
18 understanding the impacts of disturbances and extreme events on the carbon cycle. Notable
19 advancements include, but are not limited to, quantifying how harvesting activities impact
20 forest structure, carbon pool dynamics, and recovery processes; observed drastic increases of
21 the concentrations of dissolved organic carbon and dissolved methane in thermokarst lakes in
22 western Siberia in a summer warming event; disentangling the roles of herbivores and fire on



23 forest carbon dioxide flux; direct and indirect impacts of fire on the global carbon balance;
24 and improved atmospheric inversion of regional carbon sources and sinks by incorporating
25 disturbances. Combined, studies herein indicate several major research needs. First,
26 disturbances and extreme events can interact with one another, and it is important to
27 understand their overall impacts and also disentangle their relative effects on the carbon cycle.
28 Second, benchmark data layers characterizing the timing, location, type, and magnitude of
29 disturbances must be systematically created to improve our ability to quantify carbon
30 dynamics over large areas. Third, current ecosystem models are not skillful enough to
31 correctly simulate the impacts of disturbances such as disturbance-induced tree mortality and
32 its carbon consequences, and therefore must be improved to correctly represent underlying
33 processes and impacts.

34

35 **1 Introduction**

36 The terrestrial biosphere plays an important role in regulating atmospheric carbon dioxide
37 concentrations and thereby climate. Terrestrial carbon fluxes often exhibit pronounced
38 interannual variability (IAV), and disturbances and extreme climate events are primary
39 sources of IAV (Eimers et al., 2008; Reichstein et al., 2013; Xiao et al., 2014). For example,
40 gross primary productivity (GPP) exhibited significant IAV over the period 2000-2014 on the
41 global scale as identified by MODIS, with important regional differences (Fig. 1). Some
42 tropical regions (e.g. Indonesia and parts of the Amazon) had the largest IAV with standard
43 deviation in annual GPP on the order of 200-250 g C m⁻² or greater, while the remaining
44 vegetated areas in the tropics also had relatively large IAV in annual GPP (with standard
45 deviations on the order of ~150 g C m⁻²) and vegetated temperate regions had intermediate
46 IAV (with standard deviation on the order of 80-120 g C m⁻²). Extreme climate events such as



47 drought (Xiao et al., 2009; Zhao and Running, 2010) and disturbances such as fire (Bowman
48 et al., 2009), hurricanes (Chambers et al., 2007; Dahal et al., 2014; Xiao et al., 2011), wind
49 storms (McCarthy et al., 2006), and insect outbreaks (Kurz et al., 2008) can substantially alter
50 ecosystem structure and function and influence terrestrial carbon dynamics. A better
51 understanding of the impacts of disturbances and extreme climate events on terrestrial carbon
52 dynamics across different ecosystems is essential for projecting ecosystem responses to future
53 climate change and feedbacks to the climate system.

54 Extreme climate events and disturbances are projected to increase in both frequency and
55 severity during the remainder of the 21st century (IPCC, 2013), with important consequences
56 for terrestrial carbon cycling. Projecting the impacts of these future events remains a
57 challenge given the substantial uncertainty in forecasting these events and the insufficient
58 representation of ecological disturbances and extreme climate events in ecosystem and land
59 surface models. We can only make progress in this grand challenge in Earth system science
60 by understanding how different ecosystems respond to different disturbances at different time
61 scales.

62 Various approaches have been used to assess the impacts of disturbances and extreme climate
63 events on ecosystem carbon dynamics. At the ecosystem scale, in-situ methods including field
64 experiments (Barbeta et al., 2013), long-term observations, and the eddy covariance technique
65 (Amiro et al., 2010; Dong et al., 2011) seek to understand the mechanistic responses of
66 ecosystem processes to disturbances and extreme climate events. Modeling approaches
67 including process-based ecosystem models (Liu et al., 2011) or data-driven upscaling
68 approaches (Jung et al., 2009; Xiao et al., 2008) have been used for regional to global
69 assessments, which also rely heavily on satellite remote sensing (e.g. Xiao et al., 2014).
70 Synthesizing these findings is an ongoing challenge, and multiple approaches are required to



71 understand consequences of different extreme climate events and disturbances for carbon
72 cycling.

73 The impacts of disturbances and extreme climate events on carbon dynamics have received
74 growing attention. We searched the number of journal articles on these topics using Web of
75 Science (Fig. 2) and found a total of 421 and 1495 journal articles for extreme climate events
76 and disturbances, respectively, over the period from 2000 to 2014. Notably, the total number
77 of publications on the impacts of these events on carbon dynamics has been growing at an
78 average rate of 20 articles per year over the past decade (2005-2014) (Fig. 2), emphasizing the
79 growing scientific interest in these important topics.

80 The present special issue is the outcome of special sessions on the impacts of extreme climate
81 events and disturbances on carbon dynamics at the American Geophysical Union Fall
82 Meeting. It consists of 17 articles: 6 on extreme climate events and 11 on disturbances. This
83 special issue, along with the special issue on climate extremes and biogeochemical cycles in
84 *Biogeosciences* (Bahn et al., 2015), reflects recent advances in assessing how disturbances
85 and extreme climate events influence terrestrial carbon cycling. We feel that the authors have
86 provided a timely and valuable contribution to the research communities of carbon cycle and
87 global change.

88 **2 Methods and Findings**

89 We separate the impacts of drought and extreme precipitation events, herbivory (namely
90 insect outbreaks), fire, interactions between herbivory and fire, natural hazards (e.g.
91 hurricanes and typhoons), and forest management when describing the findings of the
92 manuscripts in this special issue. These events interact with one another as noted, and it can
93 be difficult to disentangle their relative effects on the carbon cycle. That being said, it is
94 important to study and synthesize how different extreme events and disturbances impact



95 carbon cycling in our quest to understand every aspect of the carbon cycle (Baldocchi,
96 2008).

97 *Drought and extreme precipitation events*

98 Piayda et al. (2014) quantified the impacts of the extreme drought event in 2012 on carbon
99 and water cycling in a Mediterranean woodland. The drought reduced overstory GPP by 28%
100 and carbon sink strength by 38% compared to 2011. Results indicated that successful
101 simulation of drought effects on the montado ecosystem requires the incorporation of variable
102 apparent maximum carboxylation rate, stomatal conductance, and vapor pressure deficit
103 sensitivity into photosynthesis-stomatal conductance modeling.

104 The simulations of a process-based ecosystem model showed that drought from 2000 to 2011
105 led to significant reduction in both GPP and NEP of China's terrestrial ecosystems at regional
106 to national scales (Liu et al., 2014). Relative to the long-term mean, the nationwide annual
107 NEP in 2001, 2006, 2009, and 2011 decreased by *ca.* 63, 88, 170, and 61 Tg C yr⁻¹,
108 respectively, due to droughts (Liu et al., 2014) due largely to reductions in GPP (Ciais et al.,
109 2005; Schwalm et al., 2010; Xiao et al., 2009).

110 The opposite of drought – extreme precipitation events - have received less attention in
111 carbon cycle studies. Jiang et al. (2013) conducted a field experiment in three subtropical
112 forests to study the responses of soil respiration to drought and extreme precipitation and
113 found that altered precipitation strongly influenced soil respiration not only by controlling soil
114 moisture but also by modifying moisture and temperature sensitivity of soil respiration. Their
115 results indicate that soil respiration of these subtropical forests would decrease if soil moisture
116 continues to decrease in the future; higher precipitation in the wet season could have a limited
117 effect on the response of soil respiration to rising temperatures (Jiang et al., 2013).



118 Zeppel, Wilks, and Lewis (2014) reviewed studies of extreme precipitation and seasonal
119 changes in precipitation on carbon metabolism in grassland and forested ecosystems. They
120 found that on average xeric biomes are likely to respond positively to extreme precipitation,
121 but mesic biomes are likely to respond negatively, and that changes in precipitation during the
122 growing season are likely to have a greater impact on carbon cycle dynamics than
123 precipitation during the non-growing season (Zeppel et al., 2014). These studies indicated that
124 the direction and magnitude of the impacts of extreme precipitation events on carbon fluxes
125 depend on the season (wet versus dry) and biome type (xeric versus mesic).

126 *Extreme temperature events*

127 Extreme temperature events have been a feature of recent climate change, especially at high
128 latitudes (IPCC, 2013). Previous studies showed that extreme temperature events often reduce
129 GPP and NEP of terrestrial ecosystems (Ciais et al., 2005; Qu et al., 2016). The effects of
130 extreme temperature on the carbon dynamics of aquatic ecosystems, however, have received
131 little attention. Pokrovsky et al. (2013) studied the impacts of the 5 – 15 °C summer warming
132 event of 2012 on the carbon dynamics of thermokarst lakes in western Siberia. Dissolved
133 organic carbon concentrations increased by a factor of two as a result of the warming event
134 despite limited changes in conductivity and pH, and the concentration of dissolved methane
135 increased by nearly fivefold (Pokrovsky et al., 2013). These results demonstrate a substantial
136 increase in the methane emission capacity from lakes as a result of summertime warming in
137 areas of permafrost thaw.

138 De Simon et al. (2013) examined the effects of manipulated warmer or cooler late
139 winter/early spring conditions on the carbon budget and yield of soybean crops. Their results
140 demonstrate that extreme temperature events in late winter did not result in significant
141 changes in the net carbon balance, indicating that increasing heat and cold waves might have



142 smaller effects on the overall annual carbon balance of irrigated croplands than expected (De
143 Simon et al., 2013). Combined, these studies indicate that the effects of extreme temperature
144 events on ecosystem carbon dynamics depend on the timing and magnitude of these events.
145 Extreme temperature events occurred in the growing season could substantially alter carbon
146 fluxes, while those events occurred during the remainder of the year had smaller effects than
147 expected.

148 *Insect outbreaks*

149 The coniferous forests of western North America have experienced an unprecedented
150 herbivore outbreak over millions of hectares over the past decades (Hicke et al., 2012; Raffa
151 et al., 2008), part of the global tree die-off due to the combined effects of elevated
152 temperatures, drought, and associated herbivory (Allen et al., 2010). Measurements of the
153 impacts of this disturbance at the site scale find minimal ecosystem carbon loss or even net
154 uptake shortly after eruptive herbivory (Brown et al., 2010), which contrasts regional
155 estimates of substantial carbon losses to the atmosphere (Kurz et al., 2008). Mathys et al.
156 (2013) in this issue used the eddy covariance technique to study carbon dioxide flux after a
157 mountain pine beetle (*Dendroctonus ponderosae*, Hopkins) attack over a two-year period and
158 compared these to an adjacent clearcut. They found that the mountain pine beetle-damaged
159 forest was a carbon sink of *ca.* 50 g C m⁻² year⁻¹ two years after attack and suggested that
160 benefit of herbivory to undamaged trees needs to be accounted for when considering
161 ecosystem-scale carbon cycle consequences of herbivory (Mathys et al., 2013). These
162 observations suggest that the impacts of herbivore outbreak depend on the type of herbivore
163 (e.g. folivores *versus* phloem-feeders) and the intensity of disturbance.

164 *Fire*



165 Fire causes direct and immediate carbon emissions into the atmosphere from biomass burning
166 (the direct effect), and subsequent changes in NEP and carbon exchange between the land and
167 the atmosphere (the indirect effect). Li et al. (2014) in this special issue provided a
168 quantitative assessment of the direct and indirect impacts of fire on the net carbon balance of
169 global terrestrial ecosystems during the 20th century. Their results show that fire decreased
170 the net carbon gain of global terrestrial ecosystems by 1.0 Pg C yr^{-1} averaged across the 20th
171 century, as a result of the fire direct effect (1.9 Pg C yr^{-1}) partly offset by the indirect effect
172 ($-0.9 \text{ Pg C yr}^{-1}$). The effect of fire on the net carbon balance significantly declined until 1970
173 with a trend of 8 Tg C yr^{-1} due to an increasing indirect effect, and increased subsequently
174 with a trend of 18 Tg C yr^{-1} due to an increasing direct effect (Li et al., 2014). These results
175 help constrain the global-scale dynamics of fire and the terrestrial carbon cycle.

176 *Insect outbreaks versus fire*

177 At the regional scale, Caldwell et al. (2013) simulated and evaluated the long-term impacts of
178 the two characteristic disturbances in the Southern Rocky Mountains forests (i.e., the outbreak
179 of mountain pine beetle and high-severity wildfire) on potential changes in species
180 composition and carbon stocks. Wildfire caused larger changes in both patterns of succession
181 and distribution of carbon among biomass pools than did mountain pine beetle disturbance;
182 carbon in standing-live biomass returned to pre-disturbance levels after 50 vs. 40 years
183 following wildfire and mountain pine beetle disturbances, respectively (Caldwell et al., 2013).

184 Clark et al. (2014) used the eddy covariance technique to study the impacts of fire and gypsy
185 moth (*Lymantria dispar* L.) disturbance in oak-dominated, pine-dominated, and mixed forests
186 in eastern North America. The NEE, GPP, and water use efficiency were of greater magnitude
187 in the oak-dominated forest before disturbance during summer. Water use efficiency declined
188 by 60% at the oak-dominated stand and by nearly 50% at the mixed stand after gypsy moth



189 disturbance, but prescribed fire had little impact on water use efficiency in the mixed or pine
190 stands (Clark et al., 2014). These results demonstrate the importance of forest type,
191 disturbance type, and time since disturbance on coupled carbon and water cycle functioning in
192 temperate forests.

193 *Hurricanes and typhoons*

194 Hurricane events in the U.S. have significant effects on regional carbon dynamics (Dahal et
195 al., 2014). Typhoons are natural disturbances to subtropical mangrove forests in Asia, and
196 their effects on ecosystem carbon dynamics of mangroves are not well understood. Chen et al.
197 (2014) examined the short-term effects of frequent strong typhoons on defoliation and the
198 NEE of subtropical mangroves. The responses of daily NEE following typhoons were highly
199 variable in mangrove ecosystems (Chen et al., 2014), demonstrating that the characteristics of
200 the typhoon and antecedent ecosystem conditions are important for understanding hurricane
201 impacts on carbon stocks and fluxes. Severe hurricane and typhoons that destroy a large
202 number of trees could have significant effects on regional carbon cycling, while those that
203 lead merely to defoliation likely had transient effects on annual ecosystem carbon exchange.

204 *Forest management*

205 Accurate quantification of the effects of partial cutting or clearcutting is essential for a better
206 understanding of forest carbon dynamics and for informing forest management. Zhou et al.
207 (2013a) conducted a meta-analysis on the impacts of partial cutting on forest carbon stocks by
208 collecting data on cutting intensity, forest structure, and carbon stock components. The results
209 showed that partial cutting reduced aboveground carbon by 43% and increased understory
210 carbon storage by nearly 400% on average, but did not have significant effects on forest floor
211 or mineral soil carbon stocks (Zhou et al., 2013a). This effort provides a new perspective on
212 the impacts of forest harvesting as it covers the spectrum of harvest disturbances from partial



213 cutting to clearcut and goes beyond previous reviews that mostly concentrated on the impacts
214 of clearcutting (Johnson and Curtis, 2001; Nave et al., 2010). The impacts of partial cutting
215 can be significant; for example, partial cutting (i.e., cutting events with aboveground biomass
216 removal rate < 90%) accounted for about three quarters of the total C loss from timber
217 harvesting in the eastern United States from 2002 to 2010 (Zhou et al., 2013b).

218 Wang et al. (2014) used a process-based forest ecosystem model, PnET-CN, to evaluate how
219 clearcutting alters ecosystem carbon fluxes, biomass, and leaf area index in northern
220 temperate forests. They found that harvest disturbance in northern temperate forests had
221 significant effects on forest carbon fluxes and stocks, and evergreen needleleaf forests were
222 more vulnerable to stand-replacing harvests than deciduous broadleaf forests (Wang et al.,
223 2014).

224 *Time since disturbance*

225 The time since disturbance is an important controlling factor of carbon dynamics. Berryman
226 et al. (2013) tested the impacts of experimental pinyon pine (*Pinus edulis* Engelm.) mortality
227 on microbial respiration. They found that litter respiration responded to water availability at
228 both treatment and control sites, and that soil respiration decreased near the site with
229 experimental mortality. These results demonstrate ecosystem-level consequences of tree
230 mortality that differs as a function of water availability (Berryman et al., 2013).

231 Yue et al. (2013) compared observations from post-fire vegetation trajectories in the boreal
232 forest with simulations from the process-based ORCHIDEE vegetation model and found that
233 the increase in atmospheric CO₂ concentrations in addition to vegetation recovery were jointly
234 responsible for current carbon sink conditions. Their results highlight the importance of
235 understanding how global change and disturbance events interact to determine current – and
236 likely future – carbon cycle dynamics (Yue et al., 2013). These two studies demonstrate that



237 the legacy of disturbance and environmental factors jointly control the carbon dynamics
238 following disturbance.

239 Modeling approaches have been widely used to simulate ecosystem carbon dynamics
240 following disturbance. Wang et al. (2014) in this special issue simulated the dynamics of
241 carbon fluxes and stocks following harvest. The simulated NEP and aboveground carbon
242 stock after clear-cuts generally followed the hypothesized trajectories (Chapin, 2011; Odum,
243 1969) while the decline in NEP was due to relatively stable GPP and gradually increasing
244 ecosystem respiration (ER). Evergreen needleleaf forests recovered more slowly from a net
245 carbon source to a net sink, and lost more carbon than deciduous broadleaf forests.

246 Disturbance-induced tree mortality regulates the forest carbon balance, but tree mortality and
247 its carbon consequences are not well represented in ecosystem models (Bond-Lamberty et al.,
248 2015). Bond-Lamberty et al. (2015) tested whether three ecosystem models – the classic big-
249 leaf model Biome-BGC and the gap-oriented models ZELIG and ED - could reproduce the
250 resilience of forest ecosystems to moderate disturbances. The models replicated observed
251 declines in aboveground biomass well but could not fully capture observed post-disturbance
252 carbon fluxes. These two studies indicate that ecosystem models are yet unable to correctly
253 simulate the effects of disturbances, and future modeling studies of disturbance effects should
254 incorporate forest population dynamics (e.g., regeneration and mortality) and relationships
255 between age-related model parameters and state variables (e.g., leaf area index).

256 Lack of critical geospatial data layers on disturbances and associated impacts on ecosystems
257 has been identified as one of the main challenges in quantifying carbon dynamics over large
258 areas (Liu et al., 2011). Recently, a continental-scale forest stand age map was developed for
259 North America using forest inventory data, large fire polygons, and remotely sensed data,
260 providing a new source information that can benefit quantification of the carbon sources and



261 sinks across the continent and contribute to studies of disturbance (Pan et al., 2011). Deng et
262 al. (2013) in this special issue used these continental stand age maps as an additional
263 constraint to atmospheric CO₂ inversions. They found that regions with recently disturbed or
264 old forests are often nudged towards carbon sources while regions with middle-aged
265 productive forests are shifted towards sinks, conforming stand age effects observed from
266 many eddy covariance flux towers (Deng et al., 2013). These results were generally consistent
267 with the synthesis results from eddy covariance flux data across North America (Amiro et al.,
268 2010) but they were inconsistent with some other studies showing that old-growth forests
269 were still carbon sinks (Desai et al., 2005; Luyssaert et al., 2008). At the sub-continental
270 level, their inverted carbon fluxes agreed well with continuous estimates of net ecosystem
271 carbon exchange (NEE) upscaled from eddy covariance flux data (Xiao et al., 2008; 2011).
272 Recent development in characterizing the timing, location, type, and magnitude of
273 disturbances (Huang et al., 2010; Kennedy et al., 2010; Masek et al., 2013; Zhu and
274 Woodcock, 2014) will likely help advance diagnosis and monitoring of carbon dynamics over
275 large areas.

276 **3 Conclusions**

277 The contributions of this special issue reflect some of the most recent advances in the impacts
278 of disturbances and extreme climate events on carbon dynamics. These studies address the
279 impacts of different types of extreme events including forest management, hurricanes and
280 typhoons, drought, extreme precipitation events, extreme temperature events, insect
281 outbreaks, and fire as well as ecosystem recovery since disturbance. The direction and
282 magnitude of the effects of these events on ecosystem carbon fluxes depend on the nature of
283 the events (type, duration, and intensity), the timing of the events (e.g., wet versus dry season,
284 summer versus winter), and the biome type (e.g., xeric versus mesic). These events typically



285 have negative effects on net carbon uptake while some events such as extreme precipitation
286 events may also have positive effects on net carbon uptake depending on antecedent
287 conditions and the nature of the extreme event.

288 Importantly, studies in this special issue collectively indicate several major research needs.
289 First, disturbances and extreme events can interact with one another, and it is important to
290 disentangle their relative effects on the carbon cycle. Second, the lack of data layers on major
291 disturbances is still one of the main challenges that hinder the improvement of quantifying
292 carbon dynamics over large areas, and benchmark data layers characterizing the timing,
293 location, type, and magnitude of disturbances must be systematically created. Third, current
294 ecosystem models are not skillful enough to correctly simulate the impacts of disturbances
295 such as disturbance-induced tree mortality and its carbon consequences, and therefore
296 ecosystem models must be improved to correctly represent the underlying processes and
297 impacts. Ongoing research in these areas will continue to improve our emerging
298 understanding of the impacts of extreme events on carbon cycling.

299

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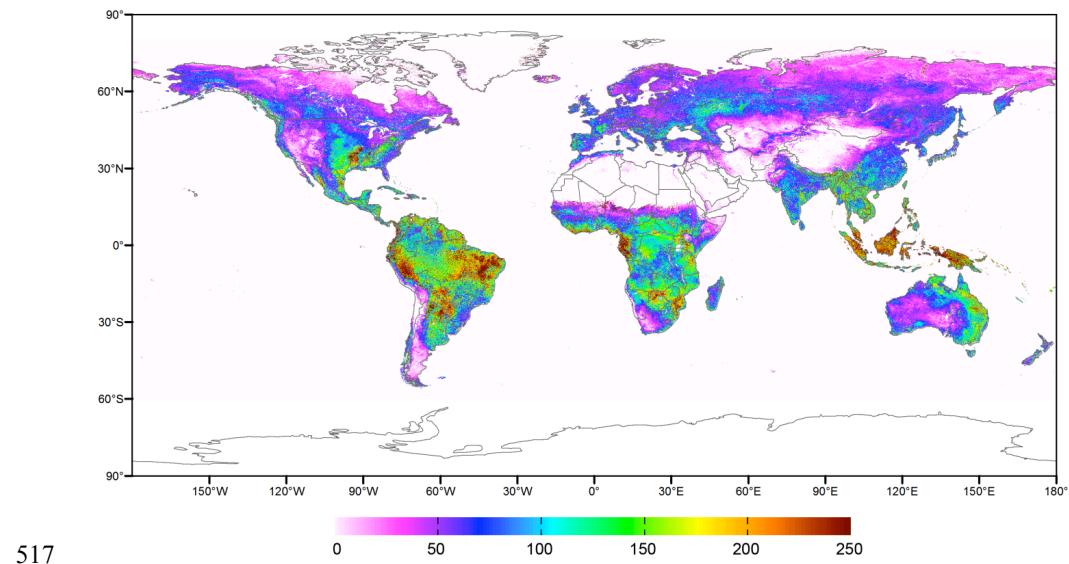
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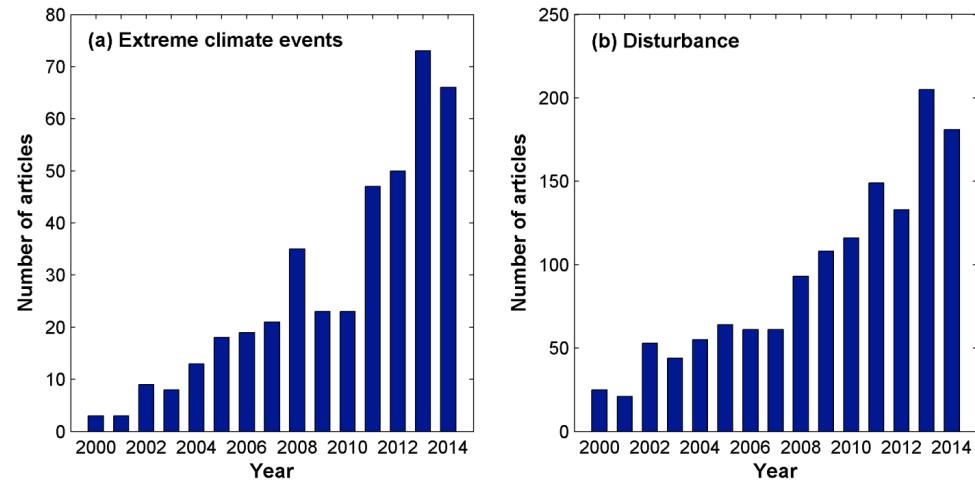
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518 **Fig. 1.** The standard deviation (as a metric of the interannual variability, IAV)
519 primary productivity (GPP) over the period 2000 to 2014 from the MODIS GPP product
520 (MOD17A3). Units are $\text{g C m}^{-2} \text{ year}^{-1}$.

521



522 523 **Fig. 2.** The number of journal articles published over the period from 2000 to 2014 as
524 identified by Web of Science™ for the impacts of (a) extreme climate events and (b)
525 disturbance on carbon dynamics. The combination of key words that we used to represent
526 'extreme climate events' is: TS=(“extreme climate events” OR “climate extremes” OR
527 drought OR “extreme precipitation”) AND TS=(“carbon dynamics” OR “carbon cycle” OR
528 “carbon flux” OR “carbon stock” OR “carbon pool”), where TS stands for Topic. The
529 combination of key words used to represent ‘disturbance’ is: TS=(disturbance OR fire OR
530 harvesting OR logging OR hurricane or “insect outbreaks”) AND TS=(“carbon dynamics” OR
531 “carbon cycle” OR “carbon flux” OR “carbon stock” OR “carbon pool”).