

Dear Editor,

We are pleased to resubmit the revised version of our manuscript after minor revisions. Here are our point-by-point responses to each of the comments of referee #5.

#### **Referee #5**

**The referee's comment: 1.1** Former comment #2: I am pleased to see that the methods section has been detailed further with help of an appendix. I do think that showing (maybe in the supplement) a revised Figure 1 (as given in your recent response to reviewers) would be of interest to the reader, i.e. how measured GPP relates to measured AOT40 and - much more importantly - F<sub>st</sub> and POD1 over time. But rather than showing growing season GPP means (they can indeed lead to a misinterpretation), it would be interesting to see for example daily values of GPP vs. POD1 up until that moment or the ozone dose of the preceding days; in other words, is there a visible effect of (acc.) ozone flux on measured daily GPP? I accept that this might be too cumbersome for this paper, but would like to mention it at least to show that your excellent set of measured NEE/GPP data might hold a lot of very useful additional information on the accumulated and/or instantaneous effect of ozone on forest growth, which potentially contradicts the main statement of this paper.

#### **The authors' response:**

We show here below the requested plots of measured daily GPP as function of the seasonal accumulated F<sub>st</sub> (Figure 1) and as function of the F<sub>st</sub> accumulated over the 5 preceding days (Figure 2). As can be seen on these graphs, GPP increases with accumulated F<sub>st</sub> in the low F<sub>st</sub> range - these are data from days in the early and late growing season - and then rather stagnates in the middle and high F<sub>st</sub> range. There is no GPP decrease with increasing F<sub>st</sub> detectable on these graphs and from this we can infer that there is no visible effect of accumulated ozone dose on measured daily GPP.

We prefer not to include these two figures in the paper, because we share the opinion of the referee that including them would make the paper too "cumbersome". We would however like to include in the supplement the figure that shows measured growing season GPP vs O<sub>3</sub> loads (see Figure 3 further below).

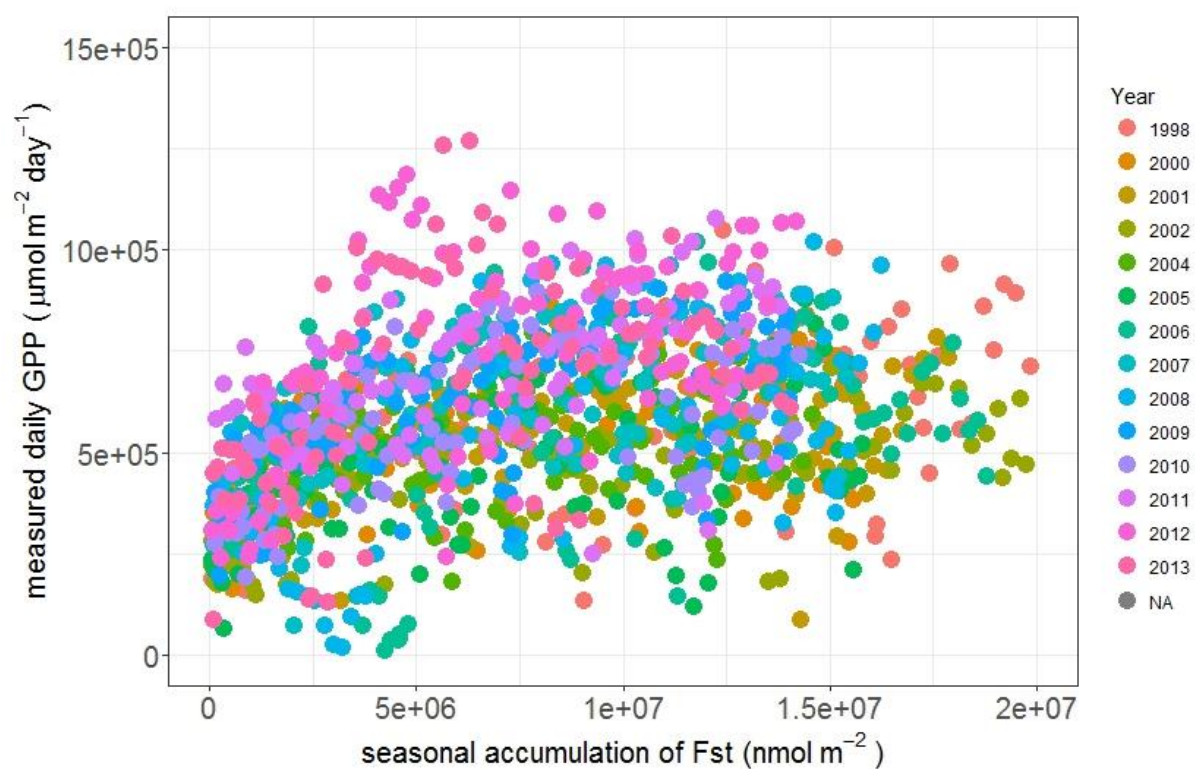


Figure 1. Measured daily GPP ( $\mu\text{mol m}^{-2} \text{ day}^{-1}$ ) as function of the stomatal  $O_3$  flux ( $\text{nmol m}^{-2}$ ) accumulated over all preceding days of the growing season.

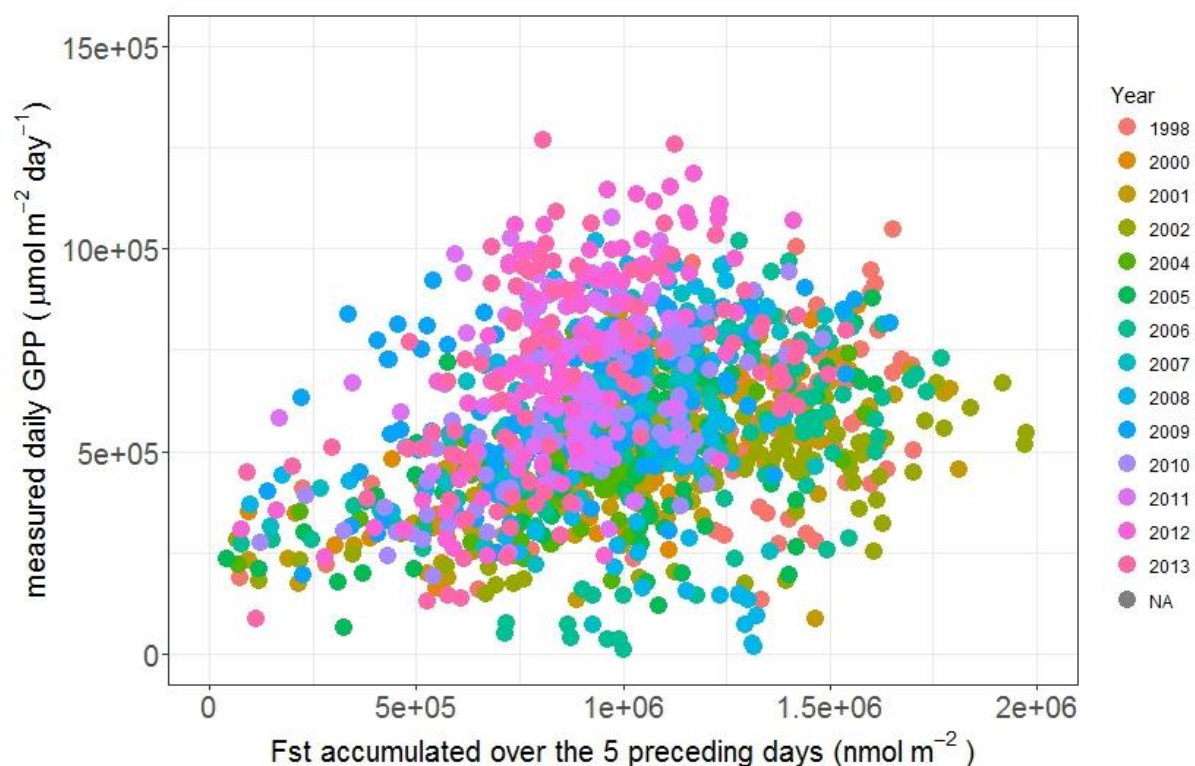


Figure 2. Measured daily GPP ( $\mu\text{mol m}^{-2} \text{ day}^{-1}$ ) as function of the stomatal  $O_3$  flux ( $\text{nmol m}^{-2}$ ) accumulated over the five preceding days.

**The referee's comment: 1.2** Furthermore, your comment "The increasing GPP involves forest recovery from acidification (Neiryck et al., 2008). The negative relation between GPP and  $O_3$  dose for this period might wrongly be interpreted as an  $O_3$  effect." is worrying, given the scope of your paper. You basically say that there is another parameter that has influenced the GPP development over time and has hence interacted with the ozone effect. Can you please describe how your methodological approach is able to disentangle the effect of a) recovery from acidification and b) ozone on GPP?

**The authors' response:**

Our methodological approach is able to disentangle the two effects because it involves the modelling of GPP from a set of input variables that excludes ozone but includes the main drivers of GPP - also accounting for the GPP increase as a recovery response from acidification. These drivers are, amongst others, the environmental drivers and LAI. Those together explain the bulk of daily and yearly variability in GPP.

**The referee's comment: 1.3** Also, line 457 of the paper "Overall, no significant  $O_3$  effects on daily and growing season GPP accumulated over the growing season were found" does seem to counteract Figure 1 (as given in your recent response to reviewers), and this figure includes even measured as opposed to modelled data. You might want to briefly comment on this.

**The authors' response:**

We agree with the referee that this statement seems to contradict the information contained in Figure 1 (= Figure 3 A-C below). To be more correct, we will rephrase the line to "Overall, no significant  $O_3$  effects on daily and growing season GPP accumulated over the growing season were found with our modelling approach." (Line 451).

We nevertheless strongly believe that Figure 1 with the plots of measured growing season GPP versus  $F_{st}$ , AOT40, and  $POD_1$  is misleading in the sense that it suggests a strong  $O_3$  effect on GPP, whereas based on our expert judgement we think the plots reflect more the fact that the period of increasing GPP by forest recovery from acidification happens to coincide with a period of decreasing  $O_3$  concentrations. It is actually only with a modelling approach like the one we applied here in this study that these "confounding" effects can be disentangled.

We furthermore believe that the negative correlation at low stomatal  $O_3$  fluxes is not causal, because the trends shown in Figure 1 (= Figure 3 A-C below) are not entirely what one would expect. For example, if we look at the plot of measured GPP versus  $F_{st}$  (Figure 3 A), we see GPP decreasing with increasing  $F_{st}$  in the low  $F_{st}$  range (50 – 70  $mmol\ O_3\ m^{-2}$ ) but we don't see a further GPP decline in the middle and high  $F_{st}$  range (70-150  $mmol\ O_3\ m^{-2}$ ). If  $O_3$  would negatively affect GPP, we would a priori expect a decrease in GPP at high  $F_{st}$  and much less at low  $F_{st}$ .

We added the figure below to the supplement of the manuscript (Figure S4) as suggested by the referee in comment 1.1 and added information about this figure at the appropriate text locations in the Results in paragraph 3.3 Ozone effects on GPP and in Discussion in paragraph 4.3 Ozone effects on GPP.

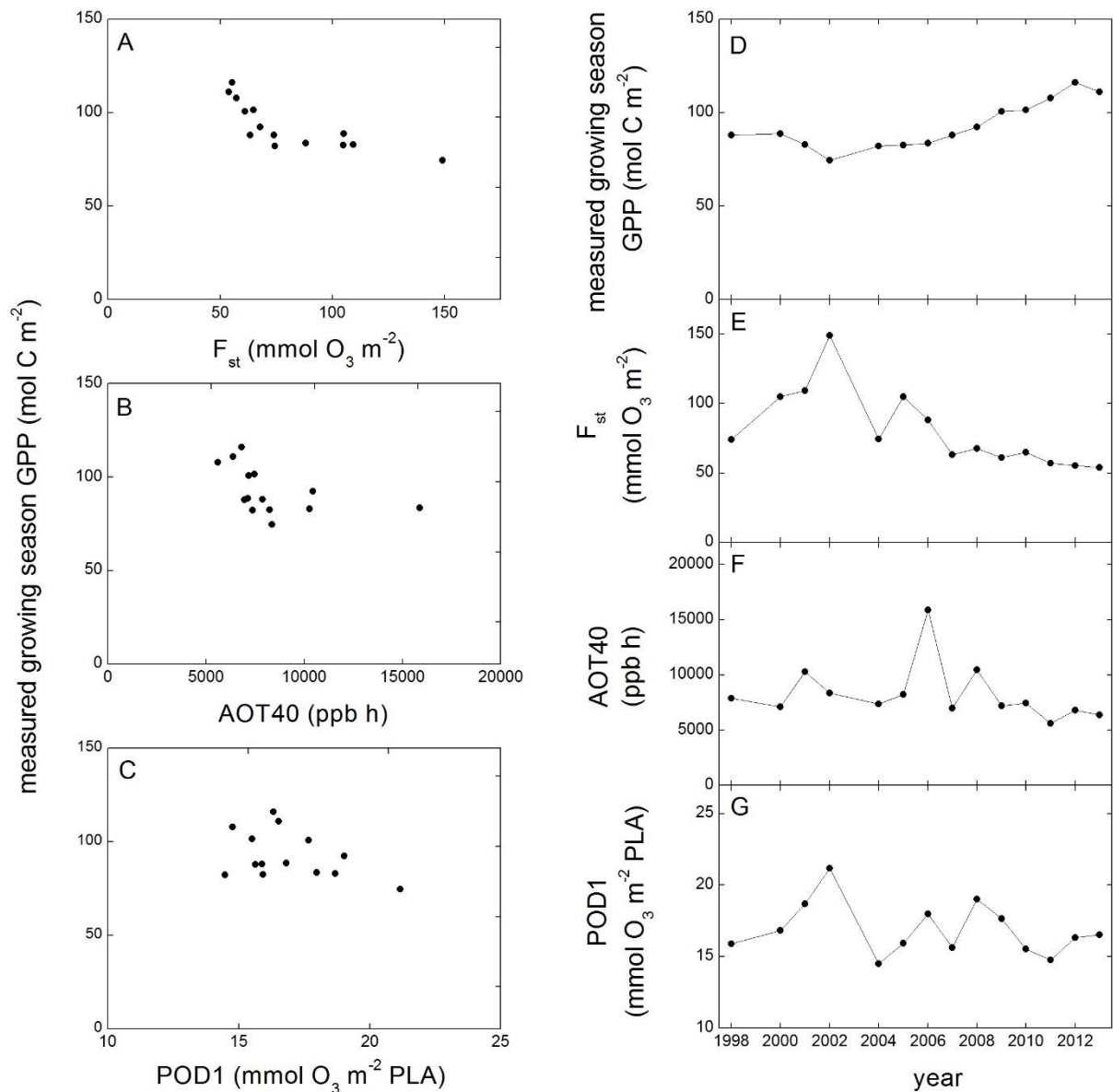


Figure 3. (A, B, C) The measured growing season GPP as function of total stomatal O<sub>3</sub> flux ( $F_{st}$ ), AOT40, and POD<sub>1</sub>. PLA = projected leaf area. (D, E, F, G) Time series of measured growing season GPP,  $F_{st}$ , AOT40, and POD<sub>1</sub>.

**The referee's comment:** 2. Former comment #6: I am fine with the response and don't have to see the results of the 1-day and 6-day delay period.

General comment: Line 794 and 819: Change "in function of" to "as a function of"

**The authors' response:**

We thank the referee for this comment and corrected this in the manuscript (Line 806 and 831).

We also thank the referee for his multitude of critical and constructive comments that helped us to improve our manuscript.

# No impact of tropospheric ozone on the gross primary productivity of a Belgian pine forest

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**Abstract** High stomatal ozone (O<sub>3</sub>) uptake has been shown to negatively affect crop yields and the growth of tree seedlings. However, little is known about the effect of O<sub>3</sub> on the carbon uptake by mature forest trees. This study investigated the effect of high O<sub>3</sub> events on gross primary productivity (GPP) for a Scots pine stand near Antwerp, Belgium over the period 1998-2013. Stomatal O<sub>3</sub> fluxes were modelled using in situ O<sub>3</sub> mixing ratio measurements and a multiplicative stomatal model, which was parameterised and validated for this Scots pine stand. Ozone-induced GPP reduction is most likely to occur during or shortly after days with high stomatal O<sub>3</sub> uptake. Therefore, a GPP model, an artificial neural network, parameterised for days with low stomatal O<sub>3</sub> uptake rates was used to simulate GPP during periods of high stomatal O<sub>3</sub> uptake. Possible negative effects of high stomatal O<sub>3</sub> uptake on GPP would then result in an overestimation of GPP by the model during or after high stomatal O<sub>3</sub> uptake events. The O<sub>3</sub> effects on GPP were linked to AOT40 and POD<sub>1</sub>. Although the critical levels for both indices were exceeded in every single year, no significant negative effects of O<sub>3</sub> on GPP were found and no correlations between GPP residuals and AOT40 and POD<sub>1</sub> were found. Overall, we conclude that no O<sub>3</sub> effects were detected on the carbon uptake by this Scots pine stand.

## 1 Introduction

Tropospheric ozone (O<sub>3</sub>) is a secondary air pollutant that has the potential to negatively affect vegetation, leading to reduced growth and carbon sequestration potential (ICP Vegetation, 2012; Subramanian et al., 2015). Background concentrations of tropospheric O<sub>3</sub> have increased with 30 % since pre-industrial times (Young et al., 2013) and are projected to further increase considerably until about 2050 (IPCC, 2007). Depending on the scenarios, background O<sub>3</sub> levels might either increase or decrease after 2050 (IPCC, 2007).

In recent years, many studies have been conducted to investigate the mechanisms underlying the O<sub>3</sub> impacts on vegetation. Ozone reduces plant growth by altering photosynthetic rates, carbohydrate production, carbon sequestration, carbon allocation, and carbon translocation (Beedlow et al., 2004; Ashmore, 2005; Wittig et al., 2009). Once O<sub>3</sub> enters the leaves through the stomata, it can affect plant growth by direct cellular damage (Mauzerall and Wang, 2001), leading to visible leaf injury and reduced leaf longevity (Li et al., 2016). In response to O<sub>3</sub>, respiratory processes increase, which will also affect the tree's carbon balance (Ainsworth et al., 2012). Skärby et al. (1987) proved that dark respiration of Scots pine shoots increased after long-term exposure to a low level of O<sub>3</sub>. Protective responses, such as compensation (e. g. repair of injured tissue), avoidance (e. g. stomatal

closure), and tolerance (e. g. alteration of metabolic pathways), all consume carbon and, hence, resistance to O<sub>3</sub> damage costs energy. The size of this cost affects the amount of carbon remaining to support growth (Skärby et al., 1998).

To assess the impact of O<sub>3</sub>, several indices have been created, e. g. AOT40 (ppb h), the cumulated O<sub>3</sub> mixing ratio in excess of a threshold of 40 ppb, and POD<sub>y</sub>, the accumulated O<sub>3</sub> flux above a flux threshold  $y$  (nmol m<sup>-2</sup> s<sup>-1</sup>). Critical levels, quantitative estimates of exposure to O<sub>3</sub> above which direct adverse effects may occur (CLRTAP, 2015), have been determined for these indices based on O<sub>3</sub> dose-response relationships from fumigation experiments with enhanced O<sub>3</sub> mixing ratios (Karlsson et al., 2004). The magnitude of the O<sub>3</sub> impact on plants depends on the intensity of O<sub>3</sub> exposure, environmental factors influencing both plant photosynthesis and the O<sub>3</sub> flux to plant surfaces, and plant species-specific defensive mechanisms (Musselman and Massman, 1999). Because of the variable plant responses to similar O<sub>3</sub> mixing ratios, the question arises whether widely applicable tolerable limits of O<sub>3</sub> mixing ratio exist (Skärby et al., 1998).

While high stomatal O<sub>3</sub> fluxes have been shown to affect the yield of crops and the growth of tree seedlings and saplings (e.g. (Büker et al., 2015)), little is known about the effect on mature forest trees. When scaling up the results from seedlings to mature trees the resulting data should be viewed with caution, due to differences in energy budgets, canopy:root balances and architecture and carbon allocation patterns (McLaughlin et al., 2007; Huttunen and Manninen, 2013). In addition to the uncertainties related with the up-scaling from seedlings to mature trees, data from controlled experiments should also be used with caution, because trees can react differently in field conditions (Skärby et al., 1998). The effect of O<sub>3</sub> uptake on carbon uptake under ambient O<sub>3</sub> mixing ratios by trees has hardly been studied in situ. Some studies showed reductions in plant growth due to stomatal O<sub>3</sub> uptake (Zapletal et al., 2011; Fares et al., 2013; Yue and Unger, 2013), while other studies did not show any effect (Samuelson, 1994; Zona et al., 2014). Whether or not an effect of stomatal O<sub>3</sub> uptake was found was species- and site- specific, and there is a clear need for more studies investigating the effect of O<sub>3</sub> on carbon uptake by mature trees in the field (Huttunen and Manninen, 2013).

In this study we investigated the effect of O<sub>3</sub> at ambient levels on the gross primary productivity (GPP) of a mature Scots pine stand in Flanders, Belgium over a period of 14 growing seasons between 1998 and 2013. The investigation of O<sub>3</sub> effects on GPP is relevant because GPP represents the first step in the process of C assimilation and quantifies the rate at which C substrate is provided for growth, wood production, et cetera. Critical levels of AOT40 and POD<sub>1</sub> are being exceeded for this stand (Neiryneck et al., 2012), indicating a potential effect of O<sub>3</sub> on tree productivity already at current ambient levels. To detect O<sub>3</sub> effects on GPP, we adopted a modelling approach that involved simulating GPP with a model with an O<sub>3</sub>-damage free parameterisation and evaluating model overestimations of GPP. We used an artificial neural network (ANN) to model GPP. ANN's are a power tool to process multidimensional data in which complex nonlinear interrelationships between the parameters can be expected. ANN's are successfully used in remote sensing, evolutionary ecology, et cetera, and have previously been used to model GPP (Lek and Guegan, 1999; Rochelle-Newall et al., 2007; Akhand et al., 2016; Liu et al., 2016). In this study we used ANN's since they don't employ predefined model conditions compared to conventional statistical models.

## 2 Materials and methods

### 2.1 Study area

The study area consisted of a 2-ha Scots pine stand in a 150-ha coniferous/deciduous forest named 'De Inslag', situated in Brasschaat (+51° 18' 33'' N, +04° 31' 14'' E), northeast of the Antwerp agglomeration and east-northeast of the Antwerp harbour (Neiryndck et al., 2008). The site has a temperate maritime climate with a mean annual temperature of 11 °C and a mean annual precipitation of 830 mm (Neiryndck et al., 2008). The soil has been classified as Albic Hypoluvic Arenosol (Gielen et al., 2011), a moderately wet sandy soil with a distinct humus and/or iron B-horizon (Janssens et al., 1999). The sandy layer overlays a clay layer which is situated at a depth of 0.7 - 2 m. As a result of the poor drainage groundwater depth is typically high, fluctuating between 0.5 and 2 m (Carrara et al., 2003).

The pine stand was planted in 1929 (Neiryndck et al., 2008). Until the autumn of 1999, when the forest was thinned, tree density amounted to 542 trees ha<sup>-1</sup>. The thinning decreased tree density to 376 trees ha<sup>-1</sup>. Average canopy height is 21.4 m (Op de Beeck et al., 2010). With a peak in leaf area index (LAI) of  $1.3 \pm 0.5 \text{ m}^2 \text{ m}^{-2}$  in 2007 (Op de Beeck et al., 2010) and an average LAI of  $1.2 \pm 0.5 \text{ m}^2 \text{ m}^{-2}$  in the period 1998-2007, the stand canopy is very sparse. Only two needle-age classes are present: current-year needles and one-year-old needles (Op de Beeck et al., 2010).

The stand is part of the ICP Forests level II and Fluxnet/CarboEurope-IP networks, and is equipped with a 41 m tall instrumentation tower. Measurements of ecosystem CO<sub>2</sub> exchange with the eddy covariance technique and meteorological measurements are being conducted at the site on a continuous basis since 1996 (Gielen et al., 2013).

### 2.2 Measurements

The period of study covered the period 1998 - 2013, with the years 1999 and 2003 excluded due to poor data quality or coverage.

#### 2.2.1 Meteorology

Air temperature ( $T_{\text{air}}$ ; °C) and humidity (RH; %) were measured with a PT100 and a HMP 230 dew point transmitter (both Vaisala, Finland) in aspirated radiation shields mounted on the tower at 2, 24 and 40 m height. Wind speed (WS, m s<sup>-1</sup>) was measured with a cup anemometer (LISA, Siggelkow GMBH, Germany) at 24, 32 and 40 m height. Ingoing and outgoing short-wave and long-wave radiation were measured at the top of the tower with a CNR1-radiometer and a CMP6-pyranometer (Kipp and Zonen, the Netherlands). Rainfall was registered by a tipping bucket rain gauge (NINA precipitation pulse transmitter, Siggelkow GMBH, Germany). Both  $T_{\text{air}}$  and RH were used to calculate vapour pressure deficit (VPD; kPa). Soil temperature ( $T_{\text{soil}}$ ; °C) was measured at 9 cm below the soil surface with temperature probes (Didcot DPS-404, UK). Soil water content (SWC; m<sup>3</sup> m<sup>-3</sup>) was measured at 25 cm below the soil surface with Time Domain Reflectometers (CS616, Campbell Scientific, UK). Instant SWC was read manually from the Reflectometers every three to 14 days and values were interpolated to obtain daily estimates, taking into account water inputs via precipitation (Gielen et al., 2010). Soil water potential (SWP; MPa) was derived from the SWC measurements with the model of van Genuchten (van Genuchten, 1980). All

meteorological variables (except SWC and rainfall) were measured every 10 seconds and half hourly means were calculated. Data gaps were filled with data from nearby weather stations.

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### 2.2.2 Ozone mixing ratio

The O<sub>3</sub> mixing ratio ([O<sub>3</sub>]; ppb) was measured at a 10 s resolution above the canopy at 24 m height with an UV Photometric Analyzer (model TEI 49I, Thermo Environmental Instruments) and converted to half hourly averages.

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Data gaps were filled with [O<sub>3</sub>] measurements done at 40 m height. If these were not available, gaps were filled with [O<sub>3</sub>] measurements from a nearby weather station from the Flemish Environmental Agency (VMM) at Luchtbal, which is less than 10 km from the site.

### 2.2.3 Leaf Area Index

A continuous time series with daily LAI values was reconstructed for the pine stand based on the historical data.

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The general approach was to keep the seasonal pattern measured in 2009 by Op de Beeck et al. (2010) fixed for each year and to scale it year per year to the seasonal maximum LAI (LAI<sub>max</sub>). LAI<sub>max</sub> had been measured with the LAI-2050 (LI-COR, Lincoln, Nebraska, USA) in 1997 and 2003 by Gond et al. ((1999) and Konôpka et al. (2005), respectively, and with digital hemispherical photography in 2007 by Op de Beeck et al. (2010). To assure consistency across the time series, measurements were corrected for clumping using a factor 0.83 (Jonckheere et al., 2005). The three measurements of LAI<sub>max</sub> were interpolated linearly to derive LAI<sub>max</sub> values for the missing years. The thinning event in 1999 was accounted for by subtracting the removed leaf biomass, determined with allometric relations from Yuste et al. (2005) and specific leaf area measurements from Op de Beeck et al. (2010).

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### 2.2.4 Gross Primary Productivity

Gross primary productivity ( $\mu\text{mol C m}^{-2} \text{ s}^{-1}$ ) was derived from net ecosystem exchange (NEE) measured with the eddy covariance technique and following the standard data quality procedures as explained in Appendix A. Half-hourly averaged values of GPP were derived for the 14 entire growing seasons of the study period, and integrated to daily and growing season totals.

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### 2.2.5 Stomatal conductance

Measurements of stomatal conductance to H<sub>2</sub>O ( $g_{\text{st, H}_2\text{O}}$ ) were done at needle level during the summers of 2007 (Op de Beeck et al., 2010) and 2013 to obtain data for parameterisation of the multiplicative stomatal model used in the calculation of stomatal O<sub>3</sub> fluxes (see sections 2.3 and 2.4). The two summers were marked by quite different environmental conditions: cold and wet in 2007 and warm and dry in 2013. Measurements were carried out with the LI-6400 Gas Exchange System (LI-COR, Lincoln, Nebraska, USA) and included diurnal stomatal courses as well as stomatal responses to PAR, T<sub>air</sub>, and VPD. Measurements were carried out on sets of three or four live fascicles, i.e. six to eight needles, which were enclosed in the LI-6400's leaf chamber while attached to the tree. Twenty-six needle sets were measured in total, equally divided between current-year and one-year-old needles. Each needle set was harvested after being measured and hemi-surface needle area was determined in order to express  $g_{\text{st, H}_2\text{O}}$  on the correct needle area basis. Needle area was derived from needle dimensions (length and width

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at top, middle, and base), assuming a hemi-circular cross-sectional needle area. Measurements of  $g_{st,H_2O}$  were converted to stomatal conductance to  $O_3$  ( $g_{st}$ ) by multiplying  $g_{st,H_2O}$  with the ratio of the molecular diffusivities of water vapour and  $O_3$  in the air (= 0.61).

### 2.3 Calculation of stomatal $O_3$ fluxes

Stomatal  $O_3$  fluxes were calculated at a half-hourly resolution from continuous series of half-hourly  $[O_3]$  and meteorology and daily LAI with an electric analog model built from three resistances in series:

$$R_{tot} = R_{aero} + R_{bl} + R_{can} \quad (1)$$

where  $R_{tot}$  is the total resistance to  $O_3$ ,  $R_{aero}$  is the aerodynamic resistance to  $O_3$ ,  $R_{bl}$  is the quasi-laminar boundary layer resistance to  $O_3$ , and  $R_{can}$  is the canopy resistance to  $O_3$  (all expressed in  $s\ m^{-1}$ ).

The aerodynamic resistance was calculated following (Grünhage, 2002) with:

$$R_{aero} = \frac{1}{\kappa u^*} \left[ \ln\left(\frac{z-d}{z_0}\right) - \Psi_h\left(\frac{z-d}{L}\right) + \Psi_h\left(\frac{z_0}{L}\right) \right] \quad (2)$$

where  $\kappa$  is the von Karman constant (0.43),  $u^*$  ( $m\ s^{-1}$ ) is the friction velocity,  $L$  is the Obukhov length,  $z$  is the  $[O_3]$  measurement height (24 m),  $d$  is the zero plane displacement (= 0.1 h),  $z_0$  is the momentum roughness parameter (= 0.65 h),  $h$  is the canopy height, and  $\Psi_h$  is the atmospheric stability function. This function is calculated using the set of coefficients published by Dyer (1974):

- for unstable atmospheric stratification ( $L < 0m$ )

$$\Psi_h = 2 * \ln\left[\frac{1}{\varphi_h(\zeta)} + 1\right] \quad (3)$$

$$\varphi_h = (1 - 16 * \zeta)^{-0.5} \quad (4)$$

$$\zeta = \frac{z-d}{L} \text{ with } z = z_2 = z_{ref,T} \text{ and } z = z_1 = d + z_0 \quad (5)$$

- for stable atmospheric stratification ( $L > 0m$ ):

$$\Psi_h = -5 * \zeta \quad (6)$$

$$\zeta = \frac{z-d}{L} \text{ with } z = z_2 = z_{ref,T} \text{ and } z = z_1 = d + z_0 \quad (7)$$

- for neutral atmospheric stratification ( $|L| \rightarrow \infty$ ):

$$\Psi_h = 0 \quad (8)$$

The quasi-laminar boundary layer resistance was calculated following (Baldocchi et al., 1987) with:

$$R_{bl} = \frac{2}{\kappa * u^*} \left( \frac{Sc}{Pr} \right)^{2/3} \quad (9)$$

where  $\kappa$  is the von Karman constant (0.43),  $u^*$  ( $m\ s^{-1}$ ) is the friction velocity, which is derived from the measured momentum fluxes,  $Sc$  is the Schmidt number (1.07 for  $O_3$ ), and  $Pr$  is the Prandtl number (0.72 for  $O_3$ ).

The canopy resistance was calculated from a stomatal resistance ( $R_{st}$ ) and a non-stomatal resistance ( $R_{nst}$ ), mounted in parallel:

$$R_{can} = \left( \frac{1}{R_{st}} + \frac{1}{R_{nst}} \right)^{-1} \quad (10)$$

175 The stomatal resistance  $R_{st}$  was calculated with an algorithm that divides the pine canopy into eight horizontal leaf layers, with LAI being divided equally between the layers, and that simulates the transfer of radiation through the layered canopy. The algorithm then calculates the stomatal resistance for the sunlit and shaded area fraction of each leaf layer with the multiplicative stomatal model described by Jarvis (1976) and reformulated by (Emberson et al., 2000). Resistance values are then integrated over all layers to obtain canopy level  $R_{st}$ . The algorithm is  
180 explained in more detail in Op de Beeck et al. (2010). The version of the multiplicative stomatal model used in this study is described in detail in Appendix B. This model was given a site-specific parameterisation as explained in section 2.4.

The non-stomatal resistance  $R_{nst}$  was assumed to be constant in time and set to  $279 \text{ s m}^{-1}$ . This value was derived from long-term  $O_3$  flux measurements in Brasschaat (Neiryneck et al., 2012).

185 Total and stomatal  $O_3$  fluxes ( $F_{tot}$  and  $F_{st}$ ;  $\text{nmol m}^{-2} \text{ s}^{-1}$ ) were calculated on a halfhourly basis with:

$$F_{tot} = 44.64 \frac{[O_3]}{R_{tot}} \quad (11)$$

$$F_{st} = F_{tot} \frac{R_{can}}{R_{st}} \quad (12)$$

where 44.64 is the molar density of air in  $\text{mol m}^{-3}$  at an air pressure of 101.3 kPa and an air temperature of  $0^\circ\text{C}$ , used here to convert flux units from  $\text{m s}^{-1}$  to  $\text{mol m}^{-2} \text{ s}^{-1}$ . Half-hourly fluxes were aggregated to daily and yearly  
190 values.

## 2.4 Parameterisation and validation of the multiplicative stomatal model

The multiplicative stomatal model was parameterised and validated against the data set of  $g_{st}$  measurements collected at the site. This data set included besides measured  $g_{st}$  also PAR,  $T_{air}$ , VPD, and SWP, and was split into a parameterisation set and a validation set by grouping the odd and even rows of data after being ranked by PAR.

195 Parameterisation was done by optimising model parameters with the function ‘lsqcurvefit’ in Matlab (Matlab and Statistics Toolbox Release 2013a), which finds the best parameter values starting from an initial values and which can be used to fit nonlinear functions with more than two independent variables. The parameters of the boundary functions  $f_{PAR}$ ,  $f_{Tair}$ ,  $f_{VPD}$ , and  $f_{SWP}$  were optimised separately, starting from initial values that were estimated visually from plots of  $g_{st}$  versus each of the input variables (PAR,  $T_{air}$ , VPD, and SWP). The phenology function  
200  $f_{phen}$  was set to 1 for parameterisation of  $f_{PAR}$ ,  $f_{Tair}$ ,  $f_{VPD}$ , and  $f_{SWP}$  since  $g_{st}$  had been measured on mature needles only. We included  $f_{phen}$  in the final model to estimate the stomatal  $O_3$  fluxes over the growing season (Appendix B).

The parameterised model was then tested against the validation data set. Model performance was evaluated with the linear regression  $y = ax + b$  fitted to the plot of measured versus modelled  $g_{st}$ , and with the following set of performance statistics: the coefficient of determination ( $R^2$ ), mean bias (MB), relative mean error (RME), Willmott's index of agreement (d), model efficiency (ME), and root mean squared error (RMSE) and its systematic (RMSE<sub>s</sub>) and unsystematic component (RMSE<sub>u</sub>). These statistics are explained briefly in Appendix C. To evaluate visually the goodness-of-fit of each boundary function, modelled  $g_{st}$  was plotted versus each of the input variables and the corresponding boundary function added to the scatter plot.

## 2.5 Detecting O<sub>3</sub> effects on GPP

We adopted a modelling approach to detect possible O<sub>3</sub> effects on GPP. Under the assumption that O<sub>3</sub>-induced GPP reduction is most likely to occur during and shortly after days of high stomatal O<sub>3</sub> fluxes, we parameterised a GPP model against a data set from which such days were removed and then simulated daily and growing season GPP with this supposedly O<sub>3</sub>-damage free model. A reduction of GPP due to O<sub>3</sub> would become apparent as a model overestimation of daily GPP for the days on which an O<sub>3</sub> effect was assumed, and possibly also as an overestimation of growing season GPP. The physiological mechanism beyond the assumption made hereabove is that at high stomatal O<sub>3</sub> fluxes the trees' defensive mechanisms cannot detoxify all O<sub>3</sub> entering the needles and damage is caused to the photosynthetic apparatus (Dizengremel, 2001; Matyssek and Sandermann, 2003). This leads to decreased gross photosynthetic rates and GPP. The damage is repaired afterwards when the stomatal O<sub>3</sub> load decreases.

We used as GPP model a feed-forward back propagation Artificial Neural Network (ANN) in Matlab (Matlab and Statistics Toolbox Release 2013a). The ANN contained 10 nodes organised in 1 layer, which came out as the best performing network after comparing networks containing different number of nodes and/or layers (data not shown). The default settings of the Matlab Neural Network Toolbox were used. A normalisation process was applied for training and testing the data: data were scaled to [-1 1] based on the lowest and highest value in the dataset. We used the Levenberg-Marquardt algorithm to train the ANN for 1000 iterations (Marquardt, 1963). Progress of training procedure was monitored using the mean squared error (MSE) of the network. The daily GPP data were used as dependent target variable in the ANN. The input variables were year, day of year,  $T_{min}$ ,  $T_{max}$ ,  $T_{mean}$ , average VPD, SWC,  $R_g$ , average  $T_{soil}$ , and average WS. Daily totals of the variables were used, with the exception of VPD,  $T_{soil}$ , and WS for which daily averaged values were used. The individual weights of these parameters on our model were estimated by replacing each input variable with a random permutation of its values. This was done for the GPP model as described above, and a GPP model containing O<sub>3</sub> as input variable to test if O<sub>3</sub> had any explanatory power on GPP.

To obtain an O<sub>3</sub>-damage free GPP model, the days for which an O<sub>3</sub> effect on GPP was expected were removed from the dataset. We assumed that if an O<sub>3</sub> effect occurs, it would occur at the days with the highest stomatal O<sub>3</sub> fluxes. Because the defensive capacity of the pine trees was not quantified and, hence, the O<sub>3</sub> load above which O<sub>3</sub> would affect GPP not known, we repeated the analysis three times by removing the days with the 2 %, 5 % and 10 % highest stomatal O<sub>3</sub> fluxes. Because the results for a 2 % and 10 % cut-off were equal to those for a 5 % cut-off,

we report only results for a 5 % cut-off. The model was trained with 2/3 of the remaining dataset, while the other 1/3 was used to test the model. This O<sub>3</sub>-damage free model was then run with the full dataset.

Model overestimation of daily GPP was evaluated (1) from the linear regression on the data of measured versus modelled GPP for the days on which an O<sub>3</sub> effect was assumed, testing whether the regression slope and intercept were different from 1 and 0, and (2) by comparing measured and modelled daily GPP for these days by means of a paired-samples t-test or a Wilcoxon signed-rank test if differences were not normally distributed (Shapiro-Wilk test). A significant outcome of this test in combination with a regression slope significantly lower than 1 (and an intercept not different from 0) would together point to a significant overestimation of GPP. Furthermore, (3) the regression slope and intercept were compared with the slope and intercept of the regression fitted to the dataset used to train and test the GPP model. This was done to evaluate whether GPP estimations for the days on which we assumed an O<sub>3</sub> effect were, in relative terms, significantly higher than GPP estimations for the days used for model training and testing. This would become apparent as a significantly lower slope (with an intercept not different from 0). Model overestimation of growing season GPP was evaluated with the first two tests above on the growing season data. Additionally, the residuals of growing season GPP (model - measurement) were plotted against AOT40, POD<sub>1</sub>, and total growing season stomatal O<sub>3</sub> uptake, and linear regression lines fitted. It was tested whether regression slope and intercept were significantly different from 0 to assess the presence of a statistically significant O<sub>3</sub> dose response relationship.

Since it may take some time to repair damage to the photosynthetic apparatus induced by O<sub>3</sub>, O<sub>3</sub> effects might last several days after a peak of O<sub>3</sub> exposure. They might thus not be detected with the model parameterised as explained above. To account for such a sustained O<sub>3</sub> effect, the modelling was repeated, now not only excluding the days with the highest stomatal O<sub>3</sub> fluxes from the dataset for model training but also the following days. The modelling was repeated with three different such delay periods, being the first, the first two, and the first six days following each flux peak. The results were evaluated with the same statistical tests as mentioned above. Because the results were similar for the three delay periods, only the results for the two-day period are shown.

High O<sub>3</sub> events are often coupled with specific meteorological conditions, i.e. high radiation and air temperatures. Since the dataset for model training had been compiled by removing the days with the highest stomatal O<sub>3</sub> fluxes, it was not unlikely that these conditions were underrepresented in the training dataset. If so, this could induce a bias in the model response to radiation and temperature and possibly result in overestimations of GPP for the days on which an O<sub>3</sub> effect was expected, which we then might wrongly attribute to O<sub>3</sub>. To evaluate the risk for such model bias, we compared the frequency distribution and range of radiation, T<sub>min</sub>, T<sub>max</sub>, T<sub>mean</sub>, and also VPD between the training dataset and the dataset with the days on which we expected an O<sub>3</sub> effect.

One of the assumptions in our approach is that O<sub>3</sub> effects on GPP only last on the short term, i.e. just a few days, and are hence not carried over. The presence of a carry-over effect would compromise the validity of our approach. We can rule out a carry-over effect by testing whether trees exposed to low stomatal O<sub>3</sub> fluxes late in the growing season behave in the same way as when exposed to similar low O<sub>3</sub> fluxes early in the growing season. To test this, we compiled a dataset that contained per growing season only the days after the first major peak of stomatal O<sub>3</sub> flux in the growing season. From this period, we further selected only the days with low stomatal O<sub>3</sub> fluxes for

which moreover no short-term  $O_3$  effect was expected. In other words, we excluded the days with a peak of stomatal  $O_3$  flux plus the six following days. We trained the GPP model with these data and then predicted GPP for the days before the first major  $O_3$  peak in each growing season. If a carry-over effect would be present, at least an effect induced during the first major  $O_3$  flux peak, it would be somehow included in the trained model. This would then underestimate GPP for the days before each first major  $O_3$  peak, where a carry-over effect has assumptively not yet occurred. Model underestimation of GPP was evaluated from a linear regression on the data of measured versus modelled GPP, testing whether the regression slope and intercept were different from 1 and 0. This slope and intercept were also compared with the slope and intercept of the regression line fitted to the training data. Also, measured and modelled GPP were compared with a paired-samples t-test or a Wilcoxon signed-rank test if differences were not normally distributed (Shapiro-Wilk test).

All statistics were performed with R 3.2.3 (R Core Team, 2015) at a significance level of  $p = 0.05$ .

### 3 Results

#### 3.1 Measurements: meteorology, GPP, and LAI

Figure 1 shows a fingerprint of the multi-annual average diel and seasonal patterns of the main meteorological variables, being  $T_{air}$ , incoming global radiation ( $R_g$ ) and VPD, and measured GPP. This figure gives a good overview of how meteorology and GPP typically changed over time in this forest; interannual anomalies from the average patterns can be found in Fig. S1. Distinct daily and seasonal patterns can be observed for, reaching highest values in summer, in the afternoon. Similar patterns can also be observed in GPP, which basically follows the pattern of  $R_g$ . As seen in Fig. 1, the photosynthetic period extends, on average, from day of year 115 (end of April) till day of year 300 (end of October). The time series of precipitation and SWP are provided in Fig. 2, while the seasonal LAI courses are shown for each year in Fig. 3. The yearly maximum LAI ranged from 1.4 to 1.9  $m^2 m^{-2}$ . The thinning of the forest in 1999 can clearly be observed in the LAI pattern. After the thinning, the canopy never fully closed.

#### 3.2 Multiplicative stomatal model and simulated $O_3$ fluxes

The optimized parameter values of the model are presented in Table 1. The different statistics to evaluate the model performance are presented in Table 2 and this for both the parameterisation and validation dataset. For the parameterisation dataset, the measured data were plotted against modelled  $g_{st}$  and plotted in Fig. 4A. The slope of the linear fit was not significantly different from 1 ( $p = 0.87$ ) and the intercept was not significantly different from 0 ( $p = 0.81$ ). Model evaluation for the validation dataset was equally good as for the parameterisation dataset (Table 2). Also in the linear fit for the validation set (Fig. 4, B), the slope was not significantly different from 1 ( $p = 0.98$ ) and the intercept was not significantly different from 0 ( $p = 0.70$ ).

Figure 5 shows the scatter plots of measured  $g_{st}$  versus each of the model input variables: PAR,  $T_{air}$ , VPD, and SWP, and for each plot the fitted boundary function.

The average daily  $O_3$  fluxes for the different years are presented in Fig. S2. Daily  $F_{st}$  ranges from 1.12 to 1.52  $nmol\ O_3\ m^{-2}\ day^{-1}$ . In 2011 the daily  $F_{st}$  was the lowest, while the highest values were observed in 2002. The annual average ratio  $F_{st}/F_{tot}$  varied between 24-28 % (Fig. S2). We observed the lowest ratios in the beginning and at the end of the growing season. Above-average ratios were observed at the peak of the growing season.

### 3.3 Ozone effects on GPP

Figure 6 shows the frequency distributions of  $R_g$ ,  $T_{min}$ ,  $T_{max}$ ,  $T_{mean}$ , and VPD for the training data set and the dataset with days on which we assumed an  $O_3$  effect. Days in the latter data set are generally more concentrated in the upper half of each variable's range. The training data set includes more days in the lower half, but conditions of high radiation, temperature or VPD do not seem to be underrepresented as the data set also included a substantial number of days in the higher part. For all variables, the variable range of the data set with days for we assumed an  $O_3$  effect is fully contained range of the training data set.

All parameters in the GPP model were ranked according to their contribution to GPP prediction (Table 3). Global radiation is the most important parameter in defining GPP with a mean squared error (MSE) of  $37500.81\ mol\ m^{-2}\ s^{-1}$ , followed by day (30240.61  $mol\ m^{-2}\ s^{-1}$ ) and year (27486.63  $mol\ m^{-2}\ s^{-1}$ ). The maximum air temperature and VPD contribute equally to the model with a MSE of about  $15300\ mol\ m^{-2}\ s^{-1}$ . Wind velocity,  $T_{min}$  and SWC contribute the least to GPP. Ozone as input variable had a MSE of  $11885.73\ mol\ m^{-2}\ s^{-1}$  (Table 3, B) and contributed the least with a MSE similar to the overall model ( $10019.30\ mol\ m^{-2}\ s^{-1}$ ).

To test for carry-over  $O_3$  effects, we evaluated and compared the linear regressions of measured versus modelled GPP of a dataset with low  $O_3$  fluxes after the first major  $O_3$  flux peak in the growing season and a dataset before this peak (Fig. 7). For both regressions, intercept and slope were not significantly different from 0 and 1 respectively (training:  $p_{slope} = 1$ ,  $p_{intercept} = 1$ , testing:  $p_{slope} = 0.83$ ,  $p_{intercept} = 0.44$ ). The slopes were also not significantly different from each other ( $p = 0.86$ ) and neither were the intercepts ( $p = 0.53$ ).

Figure 8 shows measured versus modelled daily GPP for the model trained without the days with the highest stomatal  $O_3$  fluxes (GPP model 1) and the model trained to test also for lag effects (GPP model 2). Both models reproduced daily GPP well for the dataset against which they were trained and tested, as indicated by the high  $R^2$  values and the fitted regression lines falling on the 1:1 line (Fig. 8 A, B). For both models, the regression slope for the data set with the days on which we assumed an  $O_3$  effect was significantly lower than 1 and the intercept significantly higher than 0 (Fig. 8 C, D). For GPP model 1, the regression slopes were not significantly different between the two data sets ( $p = 0.46$ ), but the intercepts were ( $p < 0.05$ ). For GPP model 2, both the regression slopes and intercepts differed significantly ( $p < 0.001$ ) and  $p < 0.001$ ). However, a Wilcoxon signed-rank test showed for both models that modelled daily GPP was not significantly higher than measured daily GPP for the days on which an  $O_3$  effect was assumed ( $p = 0.83$  and  $p = 0.64$ , respectively). Also, a paired samples t-test showed for both models that modelled growing season GPP was not significantly higher than measured growing season GPP ( $p = 0.93$  and  $p = 0.55$ , respectively). The slope and intercept of the linear regression line were not significantly different from 1 and 0 (Fig. 8 E, F).

No statistically significant correlations were found between the model residuals of growing season GPP and total stomatal  $O_3$  uptake ( $F_{st}$ ), AOT40, and  $POD_1$  (Fig. 9). Figure S4 shows the relation between the measured growing season GPP,  $F_{st}$ , AOT40 and  $POD_1$ , as well as time series of these parameters. No statistically significant correlations were found between the measured growing season GPP and AOT40 and  $POD_1$ . A significantly negative correlation between measured growing season GPP and  $F_{st}$  ( $p = 0.006$ ), due to a decline in GPP at low  $F_{st}$  ( $F_{st} < 70 \text{ mmol } O_3 \text{ m}^{-2}$ ).

## 4 Discussion

### 4.1 Multiplicative stomatal model

All statistics shown in Table 2 clearly indicated that the fitted multiplicative stomatal model performed well. For both parameterisation and validation datasets, the model explained 72 % of the variance in  $g_{st}$ . For both datasets, slope and intercept of the linear regression lines of measured versus modelled  $g_{st}$  were not significantly different from 1 and 0, respectively (Fig. 4). Moreover, the model efficiency (ME in Table 2) of 0.72 and the Wilmott's index (d) close to 1 both indicate that the modelled values matched the measured values well. A good model provides low root-mean-square error (RMSE), while the systematic component (RMSE<sub>s</sub>) should approach zero and the unsystematic component (RMSE<sub>u</sub>) should approach RMSE (Willmott et al., 1985), which was the case for this model. Low mean bias (MB) and low mean relative error (MRE) further indicated very good performance. The good performance of the model can also be observed in Fig. 5, in which the boundary lines represented the response of  $g_{st}$  to the independent variables when other variables were not limiting. The boundary lines fitted close to the data points, which is an indication of a good model, because the multiplicative stomatal model is based on the assumption that the variables act more or less multiplicatively and independently from each other (Grüters et al., 1995).

As explained in the mapping manual of the Convention on Long-Range Transboundary Air Pollution (CLRTAP), Scots pine is the representative species to assess the risk of  $O_3$  damage to coniferous forests in Atlantic Central Europe (CLRTAP, 2015). This risk is assessed on the basis of  $O_3$  doses calculated with the DO<sub>3</sub>SE algorithm, which employs a Jarvis type stomatal model that has been parameterised for Scots pine based on a compilation of primary and secondary data (Emberson et al., 2007; Büker et al., 2015; CLRTAP, 2015). The parameterisation for our Scots pine stand differs in some numbers from the one used in the DO<sub>3</sub>SE algorithm. The most remarkable difference is that  $g_{max}$  of the Scots pines in Brasschaat is much lower (0.14 vs 0.18 mol  $O_3 \text{ m}^{-2} \text{ s}^{-1}$ ). This low  $g_{max}$  may imply that during episodes of high  $O_3$  mixing ratio, the Brasschaat site is unlikely to take up very high amounts of  $O_3$  (Altimir et al., 2004; Emberson et al., 2007). This may have contributed to the absence of a clear  $O_3$  response at our site. A second difference is that the stomata of the pine trees remain opened at night ( $g_{min} = 0.02 \text{ mol } O_3 \text{ m}^{-2} \text{ s}^{-1}$ ), while the DO<sub>3</sub>SE model simulates full stomatal closure. Furthermore, the response to temperature is for our Scots pine stand shifted to a slightly higher temperature ( $T_{opt} = 25$  vs  $20 \text{ }^\circ\text{C}$ ) and the response to soil drought is much stronger ( $SWC_{max} = -0.19$  vs  $-0.7 \text{ MPa}$  and  $SWC_{min} = -1.18$  vs  $-1.5 \text{ MPa}$ ). From these differences it can be inferred that stomatal  $O_3$  uptake rates at the Brasschaat site are considerably lower than would be simulated with the DO<sub>3</sub>SE model for generic Scots pine. This highlights the importance of a site-specific parameterisation when aiming to assess stomatal  $O_3$  loads at site level.

## 4.2 Stomatal O<sub>3</sub> fluxes

The stomatal O<sub>3</sub> flux contributed on average for 26 % to the total O<sub>3</sub> flux over the study period (Fig. S2). This fraction is similar to the 21 % stomatal O<sub>3</sub> flux in a Danish Norway spruce stand (Mikkelsen et al., 2004) and the 30 % stomatal O<sub>3</sub> flux in *Quercus ilex* in Italy (Vitale et al., 2005; Gerosa et al., 2005). Cieslik (2004) showed that in Southern Europe stomatal O<sub>3</sub> flux of different vegetation types, such as pine forest and Mediterranean shrubs, is typically less than 50 % of the total O<sub>3</sub> flux. A five-year study on a Mediterranean *Pinus ponderosa* stand showed a stomatal O<sub>3</sub> flux contribution of 57 % (Fares et al., 2010). Clearly, species- and site-specific differences such as tree age or micro-climate are introducing large variability in stomatal O<sub>3</sub> uptake (Neiryneck et al., 2012).

The low relative stomatal O<sub>3</sub> flux in the Scots pine stand in Brasschaat could be the result of the sparse canopy with low LAI. Although no relation between stomatal O<sub>3</sub> flux and LAI was found in a previous site study on this site (Neiryneck et al., 2012), interannual and seasonal variation in LAI is very small, rendering such a correlation analysis very difficult.

## 4.3 Ozone effects on GPP

A comparison of the frequency distributions of radiation, temperature, and VPD between the training dataset and the dataset with the days on which we expected an O<sub>3</sub> effect showed that the meteorological conditions in the latter data set were fully represented in the training dataset. From the full overlap we can rather safely assume that the GPP model did not include a biased response to these variables that could result in a GPP overestimation that we might wrongly interpret as an effect of O<sub>3</sub>. Also, O<sub>3</sub> as input variable in the ANN did not have any explanatory power on GPP as it had the lowest MSE value close to the overall model MSE. Furthermore, a GPP model parameterised to include a carry-over effect of O<sub>3</sub> on GPP did not overestimate GPP at a statistically detectable level for days on which such an effect was not assumed to occur. From these results, we infer that carry-over effects of O<sub>3</sub> were unlikely to have occurred and that the assumption on the absence of (detectable) carry-over effects was valid.

The statistical tests ran on the data sets of measured and modelled GPP did not reveal a statistically significant model overestimation of daily GPP for the days on which we assumed an O<sub>3</sub> effect, nor an overestimation of growing season GPP. Also no significant correlations between growing season GPP residuals and stomatal O<sub>3</sub> flux, AOT40, and POD<sub>1</sub> were found, even though critical levels for AOT40 and POD<sub>1</sub> were exceeded in every single year of our study period. From these results and within the limits of the modelling approach applied in this study, we can infer that no significant effect of O<sub>3</sub> on GPP occurred.

Some earlier studies have investigated the effect of O<sub>3</sub> on forest carbon uptake. Cumulative stomatal uptake of 27 mmol m<sup>-2</sup> over the growing season did not result in any visible damage or a reduction in NEE of a poplar plantation in Belgium (Zona et al., 2014). Zapletal et al. (2011), on the other hand, reported that CO<sub>2</sub> uptake of a Norway spruce forest in the Czech Republic increased with increasing stomatal O<sub>3</sub> flux, followed by a sudden decrease in CO<sub>2</sub> uptake, suggesting that an O<sub>3</sub> flux threshold exists. Fares et al. (2013) showed a negative correlation between GPP and O<sub>3</sub> uptake at two Mediterranean ecosystems (a forest dominated by *Pinus ponderosa* in California, USA and an orchard site of *Citrus sinensis* cultivated in California, USA). A GPP reduction of 1-16 % in response to



O<sub>3</sub> uptake under ambient O<sub>3</sub> mixing ratio of 30-50 ppb was determined across vegetation types and environmental conditions in the United States by Yue and Unger (2013). The magnitude of reduction depended on the sensitivity to O<sub>3</sub> of the species and on the biome types.

AOT40 is, at present, the European standard for forest protection (EEA, 2014), with a critical level of 5000 ppb h, equivalent to a growth reduction of 5 % (Mills et al., 2011). In this study on Scots pine in Brasschaat, this value was far exceeded in all years (Fig. 9), yet no negative effect on GPP was observed in years with higher AOT40 values.

POD<sub>1</sub> is considered a more appropriate index for potential O<sub>3</sub> damage because it considers O<sub>3</sub> flux. The critical level of POD<sub>1</sub> is species-specific; a critical level of 8 mmol m<sup>-2</sup> with 2 % growth reduction is used for Norway spruce and a critical level of 4 mmol m<sup>-2</sup> with 4 % growth reduction is used for birch and beech (Mills et al., 2011). A critical level for Scots pine has not yet been determined and therefore the value of 8 mmol m<sup>-2</sup> for Norway spruce is often adopted as critical level for Scots pine. During this study, this critical level was exceeded every single year, and again no significantly negative correlation between total GPP residuals and POD<sub>1</sub> was observed. In comparison to the AOT40 level, 2006 was not the year with the highest POD<sub>1</sub>. This difference between AOT40 and POD<sub>1</sub> during 2006 was due to stomatal closure; during high O<sub>3</sub> mixing ratio events, g<sub>st</sub> was rather low (Fig. S3). POD<sub>1</sub> was highest in the year 2002, when O<sub>3</sub> mixing ratios were relatively low, but g<sub>st</sub> was high. The low O<sub>3</sub> mixing ratios explain the lower AOT40 for 2002.

Notwithstanding the absence of a statistically significant positive correlation between GPP residuals and both AOT40 and POD<sub>1</sub>, critical levels for both AOT40 and POD<sub>1</sub> were exceeded every single year. AOT40 is based on O<sub>3</sub> mixing ratio and these concentration-based indices have been shown to be weaker indicators for O<sub>3</sub> damage than flux-based indices (Karlsson et al., 2007; Simpson et al., 2007). The critical level of POD<sub>1</sub> for Scots pine was adopted from the critical level for Norway spruce (Mills et al., 2011). Possibly this critical level is too low for Scots pine. As shown by Reich (1987), pines are less sensitive to O<sub>3</sub> compared to hardwoods and crops. This supports the idea of a too low critical level.

Figure S4 shows a negative relationship between measured growing season GPP and O<sub>3</sub> dose, most notably and only significant between GPP and F<sub>st</sub> (Fig. S4, A). These trends suggest a strong effect of O<sub>3</sub> on GPP, which contradicts the outcome of our modelling analysis. The relationships are negative because the steady GPP increase that can be observed from the year 2005 until the end of the study period coincides with a steady decrease in O<sub>3</sub> loads (Fig. S4, D, E, F). To our judgement, this GPP increase is more likely to be the result of forest regrowth in response to decreased acidification at the site (Neiryck et al., 2008) than to be a response to decreased O<sub>3</sub> loads. This forest regrowth is accounted for in our modelling analyses by means of the LAI input in the ANN, allowing us to disentangle the effect of both factors on GPP. We furthermore believe that the observed trends do not reflect a causal relationship between GPP reduction and O<sub>3</sub> loads, because the GPP decrease is strongest at low O<sub>3</sub> loads but virtually absent at high O<sub>3</sub> loads (Fig. S4, A, B). This is not what would be expected in case of an O<sub>3</sub> effect.

Overall, no significant O<sub>3</sub> effects on daily and growing season GPP were found with our modelling approach. It can thus be concluded that O<sub>3</sub> did not affect GPP of the pine forest, at least if the assumptions we made in our

approach to detect O<sub>3</sub> effects are valid. The most crucial assumption involves the distinction between days at which a GPP effect did and did not occur. It was not possible to identify these days with great precision, due to lack of knowledge on the defensive capacity of the trees and their ability to repair O<sub>3</sub> damage. To overcome this, we repeated our analysis with three different peak thresholds for daily stomatal O<sub>3</sub> uptake rates above which an effect would occur and with three different delay periods over which an induced O<sub>3</sub> effect would last. The fact that all nine analyses produced the same outcome provides validity to our conclusions, despite the uncertainty involved in the identification of days with O<sub>3</sub> effects.

The lack of a detected O<sub>3</sub> effects on GPP does not mean that O<sub>3</sub> didn't negatively affect this Scots pine stand in Brasschaat. Stomatal O<sub>3</sub> uptake has here been linked to reductions in GPP only. As already stated in the introduction, protective responses such as compensation and enhanced tolerance occur in trees (Skärby et al., 1998). It is likely that trees at our study site were able to fully detoxify the O<sub>3</sub> taken up. The respiratory cost involved might have come at the expense of biomass production and growth, while gross C uptake remained unaffected. Future analyses, such as tree ring analysis, may provide an answer to whether this is the case.

## 5 Summary

We parameterised a multiplicative stomatal model for a Scots pine stand in Brasschaat. This species- and site-specific parameterised model performed very well. With this model embedded in a resistance scheme, stomatal O<sub>3</sub> fluxes were calculated and used to test for O<sub>3</sub> effects on GPP. Only very small reductions in growing season GPP were calculated. Although critical levels for AOT40 and POD<sub>1</sub> were exceeded in every single year, no significant correlations between total GPP residuals and stomatal O<sub>3</sub> flux, AOT40, and POD<sub>1</sub> were found. Within the limitations of the approach used in this study, we can thus conclude that O<sub>3</sub> did not affect the gross carbon uptake by the Scots pine stand in Brasschaat.

## 475    **Appendix A Gross Primary Productivity measurements**

*This study investigates O<sub>3</sub> effects on GPP. Below is briefly explained how GPP was measured.*

Gross primary productivity ( $\mu\text{mol C m}^{-2} \text{ s}^{-1}$ ) was derived from net ecosystem exchange (NEE) measured with the eddy covariance technique (Baldocchi and Meyers, 1998). The eddy covariance system was set up in august 1996. It consists of a sonic anemometer (Model Solent 1012R2, Gill Instruments, Lymington, UK) to measure turbulence  
480 and an infrared gas analyser (IRGA) (Model LI-6262, LI-COR Inc., Lincoln, NE, USA) to measure the CO<sub>2</sub> concentration. The measurements were conducted at the top of the tower at a height of 41 m, about 19 m above the canopy. Half-hourly NEE fluxes were calculated following the guidelines of the standard EUROFLUX methodology (Aubinet et al., 1999) as described in detail by Carrara et al. (2003; 2004). All half-hourly fluxes originating from outside the footprint were removed according to the criteria described by Nagy et al. (2006). A  
485 detailed description of the composition of the footprint can be found in the same paper. After filtering for non-forest fluxes, the remaining data have been filtered for not optimal turbulence conditions using the  $u^*$  approach (Aubinet et al., 1999); the method described in Reichstein et al. (2005) has been used as basis, including the bootstrapping to estimate 100 thresholds per year. After all the filtering on average about 55 % of the half hourly fluxes were discarded. The remaining data were used to gapfill the missing data following the non-linear  
490 regressions method (NLR; (Falge et al., 2001a)) and the Marginal Distribution Sampling method (MDS; (Reichstein et al., 2005)). Gross primary productivity was derived from NEE by adding the modelled total ecosystem respiration (autotrophic plus heterotrophic) to NEE. The ecosystem respiration was modelled with standardised algorithms as presented in Falge et al. (2001b).

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## Appendix B The multiplicative stomatal model

*In this work the multiplicative stomatal model described by Jarvis (1976) is modified specifically for the Scots pine stand in Brasschaat. The basic model is explained below.*

Stomatal conductance to O<sub>3</sub> at needle level ( $g_{st}$ ) was modelled with the multiplicative stomatal model first described by Jarvis (1976) and later reformulated by (Emberson et al., 2000). In this study we used a modified version of the model (Eq. 1).

$$g_{st} = g_{max} * f_{phen} * (f_{min} + (1 - f_{min}) * (f_{PAR} * f_T * f_{VPD} * f_{SWP})) \quad (A1)$$

Here  $g_{st}$  is the stomatal conductance to O<sub>3</sub> and  $g_{max}$  is the maximal stomatal conductance to O<sub>3</sub>. The functions  $f_{PHEN}$ ,  $f_{PAR}$ ,  $f_T$ ,  $f_{VPD}$ , and  $f_{SWP}$  represent the modification of  $g_{max}$  by, respectively, phenology, PAR, T<sub>air</sub>, VPD, and SWP. The function  $f_{min}$  is the ratio of  $g_{min}$  and  $g_{max}$  where  $g_{min}$  is the minimal stomatal conductance to O<sub>3</sub>. Impaired stomatal aperture mechanisms (stomatal sluggishness) due to O<sub>3</sub> exposure (Paoletti and Grulke, 2010) were not included in this model. In this modified version PAR, T<sub>air</sub>, VPD, and SWP influence the range between  $g_{max}$  and  $g_{min}$  instead of  $g_{max}$  and zero. This modification was needed to allow for a constant  $g_{st}$  during night time (=  $g_{min}$ ) that increases as soon as PAR > 0  $\mu\text{mol m}^{-2} \text{s}^{-1}$ , in accordance with our observations (Op de Beeck et al., 2010).

Phenology modifies  $g_{max}$  because of the variation in  $g_{st}$  due to differences in needle age. The function  $f_{PHEN}$  is modelled as follows:

$$\text{if } SGS \leq doy \leq (SGS + c), \text{ then } f_{PHEN} = f_{min} + (1 - f_{min}) * (1 - b) * \left( \frac{doy - SGS}{c} \right) + b$$

$$\text{if } SGS + c \leq doy \leq EGS - d, \text{ then } f_{PHEN} = f_{min} + (1 - f_{min}) * 1$$

$$\text{if } EGS - d \leq doy \leq EGS, \text{ then } f_{PHEN} = f_{min} + (1 - f_{min}) * (1 - b) * \left( \frac{EGS - doy}{d} \right) + b$$

where SGS is the start of the growing season (doy = 115), EGS is the end of the growing season (doy = 300), and b (= 0.8), c (= 20), and d (= 20) are species-specific parameters representing the minimum of  $f_{PHEN}$ , the number of days for  $f_{PHEN}$  to reach its maximum and the number of days during the decline of  $f_{PHEN}$  for the minimum to reach again, assuming linear increase and decrease at the start and end of the growing season.

The stomatal response to PAR is described by a rectangular hyperbola, where  $a_{PAR}$  is a species-specific parameter determining the shape of the hyperbola (Emberson et al., 2000);

$$f_{PAR} = 1 - \exp(-a_{PAR} * PAR) \quad (A3)$$

The stomatal response to T<sub>air</sub> is given by a parabolic function, where T<sub>min</sub> is the minimum temperature at which stomatal opening occurs, and T<sub>opt</sub> is the optimum temperature of stomatal opening (Emberson et al., 2000);

$$f_T = \max(0; 1 - \frac{(T - T_{opt})^2}{(T_{opt} - T_{min})^2}) \quad (A4)$$

525 The stomatal response to VPD is described by the following relationship, where  $VPD_{min}$  is a threshold for minimal stomatal opening, and  $VPD_{max}$  is a threshold for full stomatal opening (Emberson et al., 2000);

$$f_{VPD} = \min(1; \max(0; \frac{VPD_{min} - VPD}{VPD_{min} - VPD_{max}})) \quad (A5)$$

The stomatal response to SWP is described by the following relationship, where  $SWP_{min}$  is a threshold for minimal stomatal opening, and  $SWP_{max}$  is a threshold for full stomatal opening (Emberson et al., 2000);

530  $f_{SWP} = \min(1; \max(0; \frac{SWP_{min} - SWP}{SWP_{min} - SWP_{max}})) \quad (A6)$

## Appendix C Statistics of model performance

*In order to test how well the modified stomatal model performed, several model statistics were calculated. These model statistics are explained below.*

- 535 The mean bias (MB) is the mean difference between the simulations ( $S_i$ ) and the observations ( $O_i$ ), with  $n$  being the number of data points (Stone, 1993);

$$MB = n^{-1} \sum_{i=1}^n (S_i - O_i) \quad (B1)$$

The mean relative error (MRE) is the mean relative difference between the simulations and the observations (Peierls, 1935);

540 
$$MRE = n^{-1} \sum_{i=1}^n \frac{|S_i - O_i|}{O_i} \quad (B2)$$

Willmott's index of agreement ( $d$ ) is a dimensionless goodness-of-fit coefficient, with  $\bar{O}$  being the mean observation (Willmott, 1981); The index can vary between 0 and 1, with  $d$  equals 1 for a perfect agreement between simulations and observations.

$$d = 1 - \frac{\sum_{i=1}^n (S_i - O_i)^2}{\sum_{i=1}^n (|S_i - \bar{O}| + |O_i - \bar{O}|)} \quad (B3)$$

- 545 The model efficiency (ME) gives an indication of how well the observations match the simulations (Nash and Sutcliffe, 1970); Model efficiency can range from  $-\infty$  to 1 and is 1 when simulations and observations match perfectly. An efficiency of 0 indicates that the simulations are as accurate as the mean observation and an efficiency of less than zero indicates that the mean observation is a better predictor than the model.

$$ME = 1 - \frac{\sum_{i=1}^n (S_i - O_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (B4)$$

- 550 The root-mean-squared error (RMSE) is a measure of the mean absolute difference between the simulations and the observations, weighting large differences heavily (Willmott et al., 1985); The systematic component (RMSE<sub>s</sub>) estimates the model's linear or systematic error, hence, the better the regression between simulations and observations, the smaller the systematic component (Willmott et al., 1985). The unsystematic component is a measure of how much of the discrepancy between simulations and observations is due to random processes
- 555 (Willmott et al., 1985). A good model will provide low values of RMSE, with RMSE<sub>s</sub> close to zero and RMSE<sub>u</sub> close to RMSE (Willmott et al., 1985).

$$RMSE = \sqrt{n^{-1} \sum_{i=1}^n (S_i - O_i)^2} \quad (B5)$$

$$RMSE_s = \sqrt{n^{-1} \sum_{i=1}^n (S'_i - O_i)^2} \quad (B6)$$

$$RMSE_u = \sqrt{n^{-1} \sum_{i=1}^n (S_i - S'_i)^2} \quad (B7)$$

560  $S'_i = a * O_i + b$  , where 'a' and 'b' are slope and intercept, respectively, of the linear regression of the simulations versus the observations.

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## References

Ainsworth, E. A., Yendrek, C. R., Sitch, S., Collins, W. J., and Emberson, L. D.: The Effects of Tropospheric Ozone on Net Primary Productivity and Implications for Climate Change, *Annu. Rev. Plant Biol.*, 63, 637-661, 10.1146/annurev-arplant-042110-103829, 2012.

Akhand, K., Nizamuddin, M., Roytman, L., and Kogan, F.: Using remote sensing satellite data and artificial neural network for prediction of potato yield in Bangladesh, 2016, 997508-997508-997515,

Altimir, N., Tuovinen, J.-P., Vesala, T., Kulmala, M., and Hari, P.: Measurements of ozone removal by Scots pine shoots: calibration of a stomatal uptake model including the non-stomatal component, *Atmos. Environ.*, 38, 2387-2398, 10.1016/j.atmosenv.2003.09.077, 2004.

Ashmore, M. R.: Assessing the future global impacts of ozone on vegetation, *Plant Cell Environ.*, 28, 949-964, 10.1111/j.1365-3040.2005.01341.x, 2005.

Aubinet, M., Grelle, A., Ibrom, A., Rannik, Ü., Moncrieff, J., Foken, T., Kowalski, A. S., Martin, P. H., Berbigier, P., Bernhofer, C., Clement, R., Elbers, J., Granier, A., Grünwald, T., Morgenstern, K., Pilegaard, K., Rebmann, C., Snijders, W., Valentini, R., and Vesala, T.: Estimates of the Annual Net Carbon and Water Exchange of Forests: The EUROFLUX Methodology, in: *Adv. Ecol. Res.*, edited by: Fitter, A. H., and Raffaelli, D. G., Academic Press, 113-175, 1999.

Baldocchi, D., and Meyers, T.: On using eco-physiological, micrometeorological and biogeochemical theory to evaluate carbon dioxide, water vapor and trace gas fluxes over vegetation: a perspective, *Agr. Forest Meteorol.*, 90, 1-25, [http://dx.doi.org/10.1016/S0168-1923\(97\)00072-5](http://dx.doi.org/10.1016/S0168-1923(97)00072-5), 1998.

Baldocchi, D. D., Hicks, B. B., and Camara, P.: A canopy stomatal resistance model for gaseous deposition to vegetated surfaces, *Atmos. Environ.*, 21, 91-101, [http://dx.doi.org/10.1016/0004-6981\(87\)90274-5](http://dx.doi.org/10.1016/0004-6981(87)90274-5), 1987.

Beedlow, P. A., Tingey, D. T., Phillips, D. L., Hogsett, W. E., and Olszyk, D. M.: Rising atmospheric CO<sub>2</sub> and carbon sequestration in forests, *Front. Ecol. Environ.*, 2, 315-322, 2004.

Büker, P., Feng, Z., Uddling, J., Briolat, A., Alonso, R., Braun, S., Elvira, S., Gerosa, G., Karlsson, P. E., Le Thiec, D., Marzuoli, R., Mills, G., Oksanen, E., Wieser, G., Wilkinson, M., and Emberson, L. D.: New flux based dose-response relationships for ozone for European forest tree species, *Environ. Pollut.*, 206, 163-174, 10.1016/j.envpol.2015.06.033, 2015.

Carrara, A., Kowalski, A. S., Neiryck, J., Janssens, I. A., Yuste, J. C., and Ceulemans, R.: Net ecosystem CO<sub>2</sub> exchange of mixed forest in Belgium over 5 years, *Agr. Forest Meteorol.*, 119, 209-227, 10.1016/s0168-1923(03)00120-5, 2003.



600 Carrara, A., Janssens, I. A., Curiel Yuste, J., and Ceulemans, R.: Seasonal changes in photosynthesis, respiration and NEE of a mixed temperate forest, *Agr. Forest Meteorol.*, 126, 15-31, 10.1016/j.agrformet.2004.05.002, 2004.

Cieslik, S. A.: Ozone uptake by various surface types: a comparison between dose and exposure, *Atmos. Environ.*, 38, 2409-2420, 10.1016/j.atmosenv.2003.10.063, 2004.

CLRTAP: Mapping Critical Levels for Vegetation, Chapter III of Manual on methodologies and criteria for  
605 modelling and mapping critical loads and levels and air pollution effects, risks and trends., in: UNECE Convention on Long-range Transboundary Air Pollution, 2015.

Dizengremel, P.: Effects of ozone on the carbon metabolism of forest trees, *Plant Physiol. Biochem.*, 39, 729-742, [http://dx.doi.org/10.1016/S0981-9428\(01\)01291-8](http://dx.doi.org/10.1016/S0981-9428(01)01291-8), 2001.

Dyer, A. J.: A review of flux-profile relationships, *Boundary-Lay. Meteorol.*, 7, 363-372, 10.1007/bf00240838,  
610 1974.

European Environment Agency: <http://www.eea.europa.eu/>, access: 1 May, 2014.

Emberson, L. D., Ashmore, M. R., Cambridge, H. M., Simpson, D., and Tuovinen, J. P.: Modelling stomatal ozone flux across Europe, *Environ. Pollut.*, 109, 403-413, 2000.

Emberson, L. D., Büker, P., and Ashmore, M. R.: Assessing the risk caused by ground level ozone to European  
615 forest trees: a case study in pine, beech and oak across different climate regions, *Environ. Pollut.*, 147, 454-466, 10.1016/j.envpol.2006.10.026, 2007.

Falge, E., Baldocchi, D., Olson, R., Anthoni, P., Aubinet, M., Bernhofer, C., Burba, G., Ceulemans, R., Clement, R., Dolman, H., Granier, A., Gross, P., Grünwald, T., Hollinger, D., Jensen, N.-O., Katul, G., Keronen, P., Kowalski, A., Lai, C. T., Law, B. E., Meyers, T., Moncrieff, J., Moors, E., Munger, J. W., Pilegaard, K., Rannik, Ü., Rebmann, C., Suyker, A., Tenhunen, J., Tu, K., Verma, S., Vesala, T., Wilson, K., and Wofsy, S.: Gap filling strategies for defensible annual sums of net ecosystem exchange, *Agr. Forest Meteorol.*, 107, 43-69, [http://dx.doi.org/10.1016/S0168-1923\(00\)00225-2](http://dx.doi.org/10.1016/S0168-1923(00)00225-2), 2001a.

Falge, E., Baldocchi, D., Olson, R., Anthoni, P., Aubinet, M., Bernhofer, C., Burba, G., Ceulemans, R., Clement, R., Dolman, H., Granier, A., Gross, P., Grünwald, T., Hollinger, D., Jensen, N.-O., Katul, G., Keronen, P.,  
625 Kowalski, A., Ta Lai, C., Law, B. E., Meyers, T., Moncrieff, J., Moors, E., Munger, J. W., Pilegaard, K., Rannick, Ü., Rebmann, C., Suyker, A., Tenhunen, J., Tu, K., Verma, S., Vesala, T., Wilson, K., and Wofsy, S.: Gap filling strategies for long term energy flux data sets, *Agr. Forest Meteorol.*, 107, 71-77, 2001b.

Fares, S., McKay, M., Holzinger, R., and Goldstein, A. H.: Ozone fluxes in a *Pinus ponderosa* ecosystem are dominated by non-stomatal processes: Evidence from long-term continuous measurements, *Agr. Forest Meteorol.*,  
630 150, 420-431, <http://dx.doi.org/10.1016/j.agrformet.2010.01.007>, 2010.

Fares, S., Vargas, R., Detto, M., Goldstein, A. H., Karlik, J., Paoletti, E., and Vitale, M.: Tropospheric ozone reduces carbon assimilation in trees: estimates from analysis of continuous flux measurements, *Glob. Change Biol.*, 19, 2427-2443, 10.1111/gcb.12222, 2013.

Gerosa, G., Vitale, M., Finco, A., Manes, F., Denti, A., and Cieslik, S.: Ozone uptake by an evergreen  
635 Mediterranean Forest () in Italy. Part I: Micrometeorological flux measurements and flux partitioning, *Atmos. Environ.*, 39, 3255-3266, 10.1016/j.atmosenv.2005.01.056, 2005.

Gielen, B., Verbeeck, H., Neirynck, J., Sampson, D. A., Vermeiren, F., and Janssens, I. A.: Decadal water balance of a temperate Scots pine forest (*Pinus sylvestris* L.) based on measurements and modelling, *Biogeosciences*, 7, 1247-1261, 2010.

640 Gielen, B., Neiryck, J., Luyssaert, S., and Janssens, I. A.: The importance of dissolved organic carbon fluxes for the carbon balance of a temperate Scots pine forest, *Agr. Forest Meteorol.*, 151, 270-278, 10.1016/j.agrformet.2010.10.012, 2011.

Gielen, B., De Vos, B., Campioli, M., Neiryck, J., Papale, D., Verstraeten, A., Ceulemans, R., and Janssens, I. A.: Biometric and eddy covariance-based assessment of decadal carbon sequestration of a temperate Scots pine  
645 forest, *Agr. Forest Meteorol.*, 174-175, 135-143, 10.1016/j.agrformet.2013.02.008, 2013.

Gond, V., De Pury, D. G. G., Veroustraete, F., and Ceulemans, R.: Seasonal variations in leaf area index, leaf chlorophyll, and water content; scaling-up to estimate fAPAR and carbon balance in a multilayer, multispecies temperate forest, *Tree Physiol.*, 19, 673-679, 1999.

Grünhage, L.: An O<sub>3</sub> flux-based risk assessment for spring wheat., Joint ICP Vegetation/EMEP. Ad-hoc Expert  
650 Panel Meeting on Modelling and Mapping of Ozone Flux and Deposition to Vegetation to be held under the UN/ECE Convention on Long-range Transboundary Air Pollution, 2002.

Grüters, U., Fangmeier, A., and Jäger, H.-J.: Modelling stomatal responses of spring wheat (*Triticum aestivum* L. cv. Turbo) to ozone and different levels of water supply, *Environ. Pollut.*, 87, 141-149, 1995.

Huttunen, S., and Manninen, S.: A review of ozone responses in Scots pine (*Pinus sylvestris*), *Environ. Exp. Bot.*,  
655 90, 17-31, <http://dx.doi.org/10.1016/j.envexpbot.2012.07.001>, 2013.

ICP Vegetation: Yield response and ozone injury on *Phaseolus vulgaris*. Experimental Protocol., 2012.

IPCC: Climate Change 2007: Synthesis Report. Contribution to Working Group I, II and III to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change., Geneva, Switzerland, 104, 2007.

Janssens, I. A., Sampson, D. A., Cermak, J., Meiresonne, L., Riguzzi, F., Overloop, S., and Ceulemans, R.:  
660 Above- and belowground pyhtomass and carbon storage in a Belgian Scots pine stand, *Ann. For. Sci.*, 56, 81-90, 1999.

Jarvis, P. G.: The interpretation of the variations in leaf water potential and stomatal conductance found in canopies in the field, *Philos. T. Roy. Soc. B.*, 273, 593-610, 1976.

Jonckheere, I., Muys, B., and Coppin, P.: Allometry and evaluation of in situ optical LAI determination in Scots  
665 pine: a case study in Belgium, *Tree Physiol.*, 25, 2005.

Karlsson, P. E., Uddling, J., Braun, S., Broadmeadow, M., Elvira, S., Gimeno, B. S., Le Thiec, D., Oksanen, E., Vandermeiren, K., Wilkinson, M., and Emberson, L.: New critical levels for ozone effects on young trees based on AOT40 and simulated cumulative leaf uptake of ozone, *Atmos. Environ.*, 38, 2283-2294, 10.1016/j.atmosenv.2004.01.027, 2004.

670 Karlsson, P. E., Braun, S., Broadmeadow, M., Elvira, S., Emberson, L., Gimeno, B. S., Le Thiec, D., Novak, K., Oksanen, E., Schaub, M., Uddling, J., and Wilkinson, M.: Risk assessments for forest trees: the performance of the ozone flux versus the AOT concepts, *Environ. Pollut.*, 146, 608-616, 10.1016/j.envpol.2006.06.012, 2007.

Konôpka, B., Yuste, J. C., Janssens, I. A., and Ceulemans, R.: Comparison of Fine Root Dynamics in Scots Pine and Pedunculate Oak in Sandy Soil, *Plant Soil*, 276, 33-45, 10.1007/s11104-004-2976-3, 2005.

675 Lek, S., and Guegan, J. F.: Artificial neural networks as a tool in ecological modelling, an introduction, *Ecol. Model.*, 120, 65-73, Doi 10.1016/S0304-3800(99)00092-7, 1999.

Li, P., Calatayud, V., Gao, F., Uddling, J., and Feng, Z.: Differences in ozone sensitivity among woody species are related to leaf morphology and antioxidant levels, *Tree Physiol.*, 36, 1105-1116, 10.1093/treephys/tpw042, 2016.

680 Liu, S., Zhuang, Q., He, Y., Noormets, A., Chen, J., and Gu, L.: Evaluating atmospheric CO<sub>2</sub> effects on gross primary productivity and net ecosystem exchanges of terrestrial ecosystems in the conterminous United States using the AmeriFlux data and an artificial neural network approach, *Agr. Forest Meteorol.*, 220, 38-49, <http://dx.doi.org/10.1016/j.agrformet.2016.01.007>, 2016.

Marquardt, D.: An Algorithm for Least-Squares Estimation of Nonlinear Parameters, *J. Soc. Ind. Appl. Math.*, 11, 431-441, 10.1137/0111030, 1963.

685 Matyssek, R., and Sandermann, H.: Impact of Ozone on Trees: an Ecophysiological Perspective, in: *Progress in Botany: Genetics Physiology Systematics Ecology*, edited by: Esser, K., Lüttge, U., Beyschlag, W., and Hellwig, F., Springer Berlin Heidelberg, Berlin, Heidelberg, 349-404, 2003.

Mauzerall, D. L., and Wang, X.: Protecting agricultural crops from the effects of tropospheric ozone exposure: Reconciling Science and Standard Setting in the United States, Europe, and Asia, *Annu. Rev. Energy Env.*, 26, 237-268, doi:10.1146/annurev.energy.26.1.237, 2001.

690 McLaughlin, S. B., Nosal, M., Wullschlegel, S. D., and Sun, G.: Interactive effects of ozone and climate on tree growth and water use in a southern Appalachian forest in the USA, *New Phytol.*, 174, 109-124, 10.1111/j.1469-8137.2007.02018.x, 2007.

695 Mikkelsen, T. N., Ro-Poulsen, H., Hovmand, M. F., Jensen, N. O., Pilegaard, K., and Egeløv, A. H.: Five-year measurements of ozone fluxes to a Danish Norway spruce canopy, *Atmos. Environ.*, 38, 2361-2371, <http://dx.doi.org/10.1016/j.atmosenv.2003.12.036>, 2004.

Mills, G., Pleijel, H., Büker, P., Braun, S., Emberson, L., Harmens, H., Hayes, F., Simpson, D., Grünhage, L., Karlsson, P.-E., Danielsson, H., Bermejo, V., and Fernández, I. G.: Mapping critical levels for vegetation, in: 700 *Manual on methodologies and criteria for Modelling and Mapping Critical Loads & Levels and Air Pollution Effects, Risks and Trends*, III.1-III.114, 2011.

Musselman, R. C., and Massman, W. J.: Ozone flux to vegetation and its relationship to plant response and ambient air quality standards, *Atmos. Environ.*, 33, 65-73, 1999.

Nagy, M. T., Janssens, I. A., Curiel Yuste, J., Carrara, A., and Ceulemans, R.: Footprint-adjusted net ecosystem 705 CO<sub>2</sub> exchange and carbon balance components of a temperate forest, *Agr. Forest Meteorol.*, 139, 344-360, <http://dx.doi.org/10.1016/j.agrformet.2006.08.012>, 2006.

Nash, J. E., and Sutcliffe, J. V.: River flow forecasting through conceptual models, *J. Hydrol.*, 10, 282-290, 1970.

Neiryneck, J., Janssens, I. A., Roskams, P., Quataert, P., Verschelde, P., and Ceulemans, R.: Nitrogen biogeochemistry of a mature Scots pine forest subjected to high nitrogen loads, *Biogeochemistry*, 91, 201-222, 10.1007/s10533-008-9280-x, 2008.

710 Neiryneck, J., Gielen, B., Janssens, I. A., and Ceulemans, R.: Insights into ozone deposition patterns from decade-long ozone flux measurements over a mixed temperate forest, *J. Environ. Monit.*, 14, 2012.

Op de Beeck, M., Gielen, B., Jonckheere, I., Samson, R., Janssens, I. A., and Ceulemans, R.: Needle age-related and seasonal photosynthetic capacity variation is negligible for modelling yearly gas exchange of a sparse 715 temperate Scots pine forest, *Biogeosciences*, 7, 199-215, 2010.

Paoletti, E., and Grulke, N. E.: Ozone exposure and stomatal sluggishness in different plant physiognomic classes, *Environ. Pollut.*, 158, 2664-2671, 10.1016/j.envpol.2010.04.024, 2010.

Peierls, R.: Statistical Error in Counting Experiments, *P. Roy. Soc. A – Math. Phys.*, 149, 467-486, 10.1098/rspa.1935.0076, 1935.

- 720 Reich, P. B.: Quantifying plant response to ozone - a unifying theory, *Tree Physiol.*, 3, 63-91, 1987.
- Reichstein, M., Falge, E., Baldocchi, D., Papale, D., Aubinet, M., Berbigier, P., Bernhofer, C., Buchmann, N., Gilmanov, T., Granier, A., Grünwald, T., Havránková, K., Ilvesniemi, H., Janous, D., Knohl, A., Laurila, T., Lohila, A., Loustau, D., Matteucci, G., Meyers, T., Miglietta, F., Ourcival, J.-M., Pumpanen, J., Rambal, S., Rotenberg, E., Sanz, M., Tenhunen, J., Seufert, G., Vaccari, F., Vesala, T., Yakir, D., and Valentini, R.: On the  
725 separation of net ecosystem exchange into assimilation and ecosystem respiration: review and improved algorithm, *Glob. Change Biol.*, 11, 1424-1439, 10.1111/j.1365-2486.2005.001002.x, 2005.
- Rochelle-Newall, E. J., Winter, C., Barrón, C., Borges, A. V., Duarte, C. M., Elliott, M., Frankignoulle, M., Gazeau, F., Middelburg, J. J., Pizay, M.-D., and Gattuso, J.-P.: Artificial neural network analysis of factors controlling ecosystem metabolism in coastal systems, *Ecol. Appl.*, 17, S185-S196, 2007.
- 730 Samuelson, L. J.: Ozone-exposure responses of Black-Cherry and Red Maple seedlings, *Environ. Exp. Bot.*, 34, 355-362, 10.1016/0098-8472(94)90017-5, 1994.
- Simpson, D., Ashmore, M. R., Emberson, L., and Tuovinen, J. P.: A comparison of two different approaches for mapping potential ozone damage to vegetation. A model study, *Environ. Pollut.*, 146, 715-725, 10.1016/j.envpol.2006.04.013, 2007.
- 735 Skärby, L., Troeng, E., and Boström, C. A.: Ozone uptake and effects on transpiration, net photosynthesis and dark respiration in Scots pine, *For. Sci.*, 33, 801-808, 1987.
- Skärby, L., Ro-Poulsen, H., Wellburn, F. A. M., and Sheppard, L. J.: Impacts of ozone on forests: a European perspective, *New Phytol.*, 139, 109-122, 1998.
- Stone, R. J.: Improved statistical procedure for the evaluation of solar radiation estimation models, *Solar Energy*,  
740 51, 289-291, 1993.
- Subramanian, N., Karlsson, P. E., Bergh, J., and Nilsson, U.: Impact of Ozone on Sequestration of Carbon by Swedish Forests under a Changing Climate: A Modeling Study, *For. Sci.*, 61, 445-457, 10.5849/forsci.14-026, 2015.
- van Genuchten, M. T.: A Closed-form Equation for Predicting the Hydraulic Conductivity of Unsaturated Soils,  
745 44, 892-898, 10.2136/sssaj1980.03615995004400050002x, 1980.
- Vitale, M., Gerosa, G., Ballarindenti, A., and Manes, F.: Ozone uptake by an evergreen mediterranean forest ( L.) in Italy—Part II: flux modelling. Upscaling leaf to canopy ozone uptake by a process-based model, *Atmos. Environ.*, 39, 3267-3278, 10.1016/j.atmosenv.2005.01.057, 2005.
- Willmott, C. J.: On the validation of models, *Phys. Geogr.*, 2, 184-194, 1981.
- 750 Willmott, C. J., Ackleson, S. G., Davis, R. E., Feddema, J. J., Klink, K. M., Legates, D. R., O'Donnell, J., and Rowe, C. M.: Statistics for the evaluation and comparison of models, *J. Geophys. Res.*, 90, 8995-9005, 1985.
- Wittig, V. E., Ainsworth, E. A., Naidu, S. L., Karnosky, D. F., and Long, S. P.: Quantifying the impact of current and future tropospheric ozone on tree biomass, growth, physiology and biochemistry: a quantitative meta-analysis, *Glob. Change Biol.*, 15, 396-424, 10.1111/j.1365-2486.2008.01774.x, 2009.
- 755 Young, P. J., Archibald, A. T., Bowman, K. W., Lamarque, J. F., Naik, V., Stevenson, D. S., Tilmes, S., Voulgarakis, A., Wild, O., Bergmann, D., Cameron-Smith, P., Cionni, I., Collins, W. J., Dalsoren, S. B., Doherty, R. M., Eyring, V., Faluvegi, G., Horowitz, L. W., Josse, B., Lee, Y. H., MacKenzie, I. A., Nagashima, T., Plummer, D. A., Righi, M., Rumbold, S. T., Skeie, R. B., Shindell, D. T., Strode, S. A., Sudo, K., Szopa, S., and Zeng, G.: Pre-industrial to end 21st century projections of tropospheric ozone from the Atmospheric Chemistry and Climate

- 760 Model Intercomparison Project (ACCMIP), *Atmos. Chem. Phys.*, 13, 2063-2090, 10.5194/acp-13-2063-2013, 2013.
- Yue, X., and Unger, N.: Ozone vegetation damage effects on gross primary productivity in the United States, *Atmos. Chem. Phys. Discussions*, 13, 31563-31605, 10.5194/acpd-13-31563-2013, 2013.
- Yuste, J. C., Konôpka, B., Janssens, I. A., Coenen, K., Xiao, C. W., and Ceulemans, R.: Contrasting net primary  
765 productivity and carbon distribution between neighboring stands of *Quercus robur* and *Pinus sylvestris*, *Tree Physiol.*, 25, 701-712, 2005.
- Zapletal, M., Cudlin, P., Chroust, P., Urban, O., Pokorny, R., Edwards-Jonasova, M., Czerny, R., Janous, D., Taufarova, K., Vecera, Z., Mikuska, P., and Paoletti, E.: Ozone flux over a Norway spruce forest and correlation with net ecosystem production, *Environ. Pollut.*, 159, 1024-1034, 10.1016/j.envpol.2010.11.037, 2011.
- 770 Zona, D., Gioli, B., Fares, S., De Groote, T., Pilegaard, K., Ibrom, A., and Ceulemans, R.: Environmental controls on ozone fluxes in a poplar plantation in Western Europe, *Environ. Pollut.*, 184, 201-210, 10.1016/j.envpol.2013.08.032, 2014.

775 Table 1. Optimised parameter values of the multiplicative stomatal model.

$g_{\max}$ (mol O <sub>3</sub> m <sup>-2</sup> s <sup>-1</sup> )	0.14
$g_{\min}$ (mol O <sub>3</sub> m <sup>-2</sup> s <sup>-1</sup> )	0.02
$a_{\text{PAR}}$	0.0057
$T_{\text{opt}}$ (°C)	25.61
$T_{\min}$ (°C)	5.47
$\text{VPD}_{\min}$ (kPa)	3.16
$\text{VPD}_{\max}$ (kPa)	0.51
$\text{SWP}_{\min}$ (MPa)	-1.18
$\text{SWP}_{\max}$ (MPa)	-0.19

Table 2. Performance statistics for the multiplicative stomatal model: mean bias (MB), relative mean error (RME), systematic and unsystematic root mean squared error (RMSE<sub>s/u</sub>), Willmott's index of agreement (d), model efficiency (ME), coefficient of determination (R<sup>2</sup>).

Statistics	Parameterisation	Validation
MB	0.002	0.002
RME	0.34	0.33
RMSE	0.019	0.019
RMSE <sub>s</sub>	0.006	0.006
RMSE <sub>u</sub>	0.017	0.017
d	0.99	0.99
ME	0.72	0.72
R <sup>2</sup>	0.72	0.72

785

Table 3: Ranking of the parameters defining GPP in the ANN by replacing each input variable with a random permutation of its values. (A) The parameters with their mean squared error (MSE, mol m<sup>-2</sup> day<sup>-1</sup>) for the model without O<sub>3</sub> (B) The parameters with their MSE for the model with O<sub>3</sub>. The overall model MSE without any random permutation is also shown.

Ranking Nr.	A	B
1	R <sub>g</sub> – 37500.81	R <sub>g</sub> – 41358.93
2	doy – 30240.61	year – 33978.09
3	year – 27486.63	doy – 31127.90
4	VPD – 15380.68	T <sub>soil</sub> – 24893.78
5	T <sub>max</sub> – 15323.22	T <sub>max</sub> – 23567.45
6	T <sub>soil</sub> – 15076.75	T <sub>mean</sub> – 21354.76
7	T <sub>mean</sub> – 13858.91	VPD – 16395.14
8	WV – 13369.01	T <sub>min</sub> – 15418.16
9	T <sub>min</sub> – 12732.96	WV – 14685.97
10	SWC – 12402.04	SWC – 12831.19
11		O <sub>3</sub> – 11885.73
Overall model MSE	11360.85	10019.30



Fig. 1. Fingerprint of air temperature ( $T_{\text{air}}$ ), incoming global radiation ( $R_g$ ), vapour pressure deficit (VPD), and measured gross primary productivity (GPP), averaged over the period 1998-2013. Day of year is plotted on the y-axis and hour of day on the x-axis.

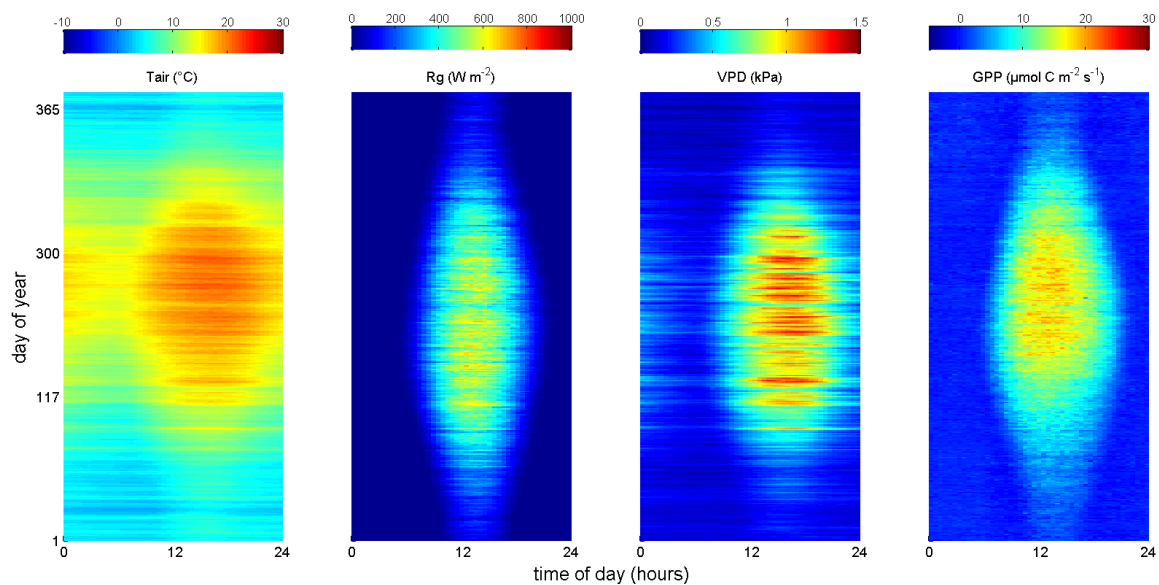
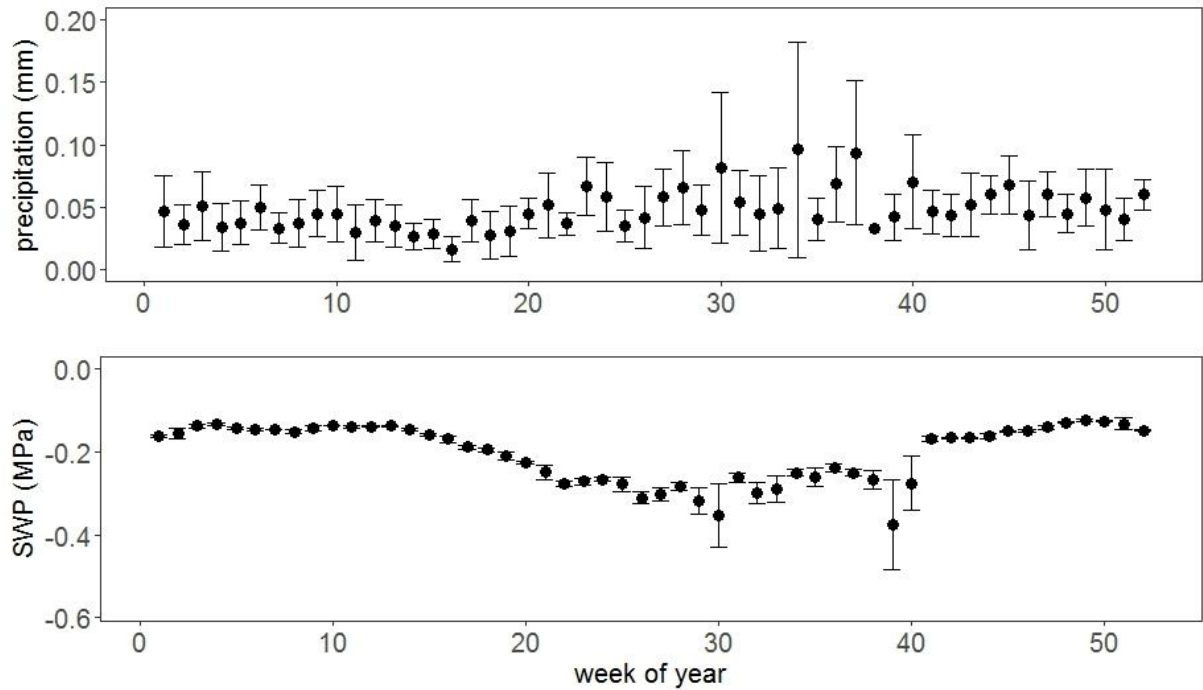
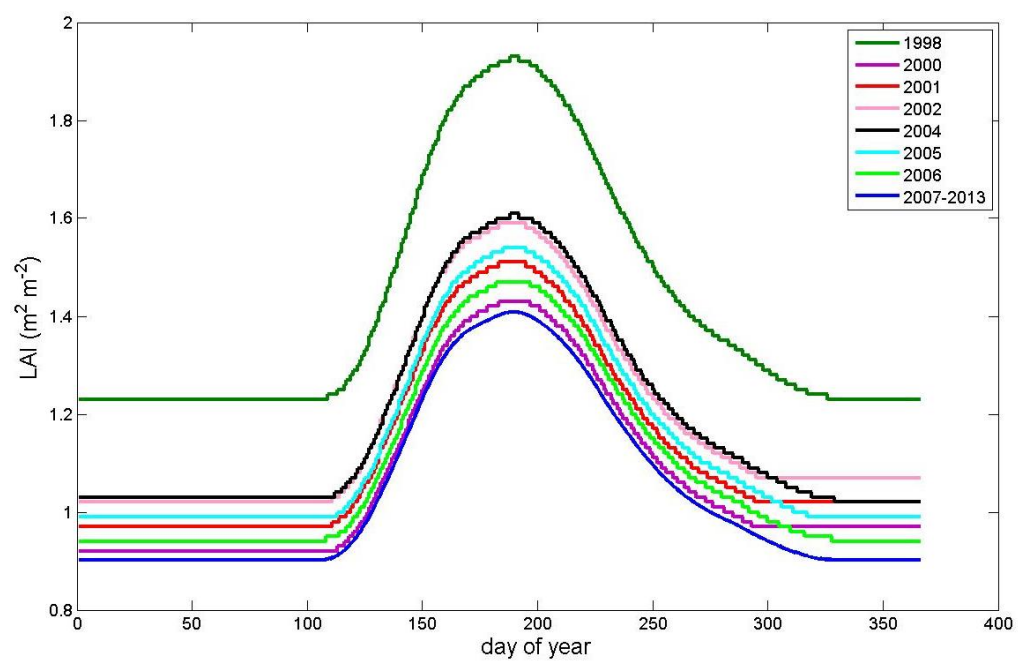


Fig. 2. Time series of the weekly total precipitation and mean soil water potential (SWP). The precipitation and SWP data are averaged over the period 1998-2013. Error bars represent the 95 % confidence intervals.

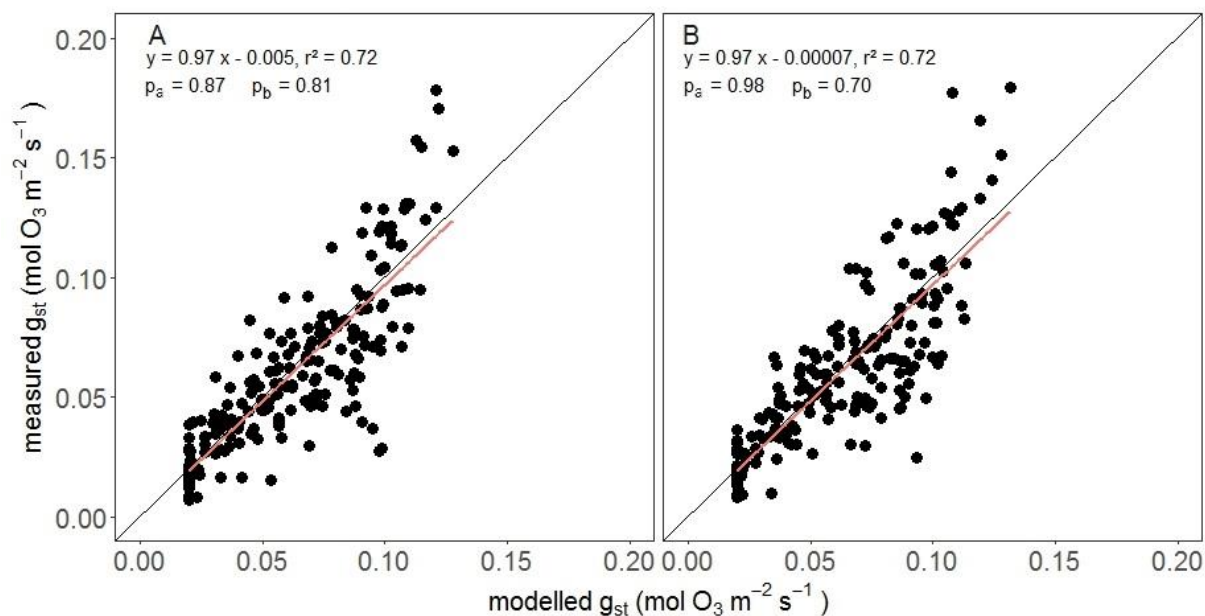


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Fig. 3. Seasonal course of LAI for each of the 14 growing seasons used in this study.



800 Fig. 4. Measured versus modelled stomatal conductance ( $g_{st}$ ) for the parameterisation dataset (A) ( $n = 205$ ) and  
the validation dataset (B) ( $n = 205$ ). The black line is the 1:1 line. The red line is the linear fit for which the equation  
is given in the figure. Also shown are the p-values of test for the slope being different from 1 ( $p_a$ ) and the intercept  
different from 0 ( $p_b$ ).



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Fig. 5. Measured stomatal conductance ( $g_{st}$ ) as a function of the different variables used in the multiplicative model: photosynthetically active radiation (PAR), air temperature ( $T_{air}$ ), vapour pressure deficit (VPD), and soil water potential (SWP). The red line represents the boundary line for which the functions are given in Appendix B (A3-A6). (n = 205)

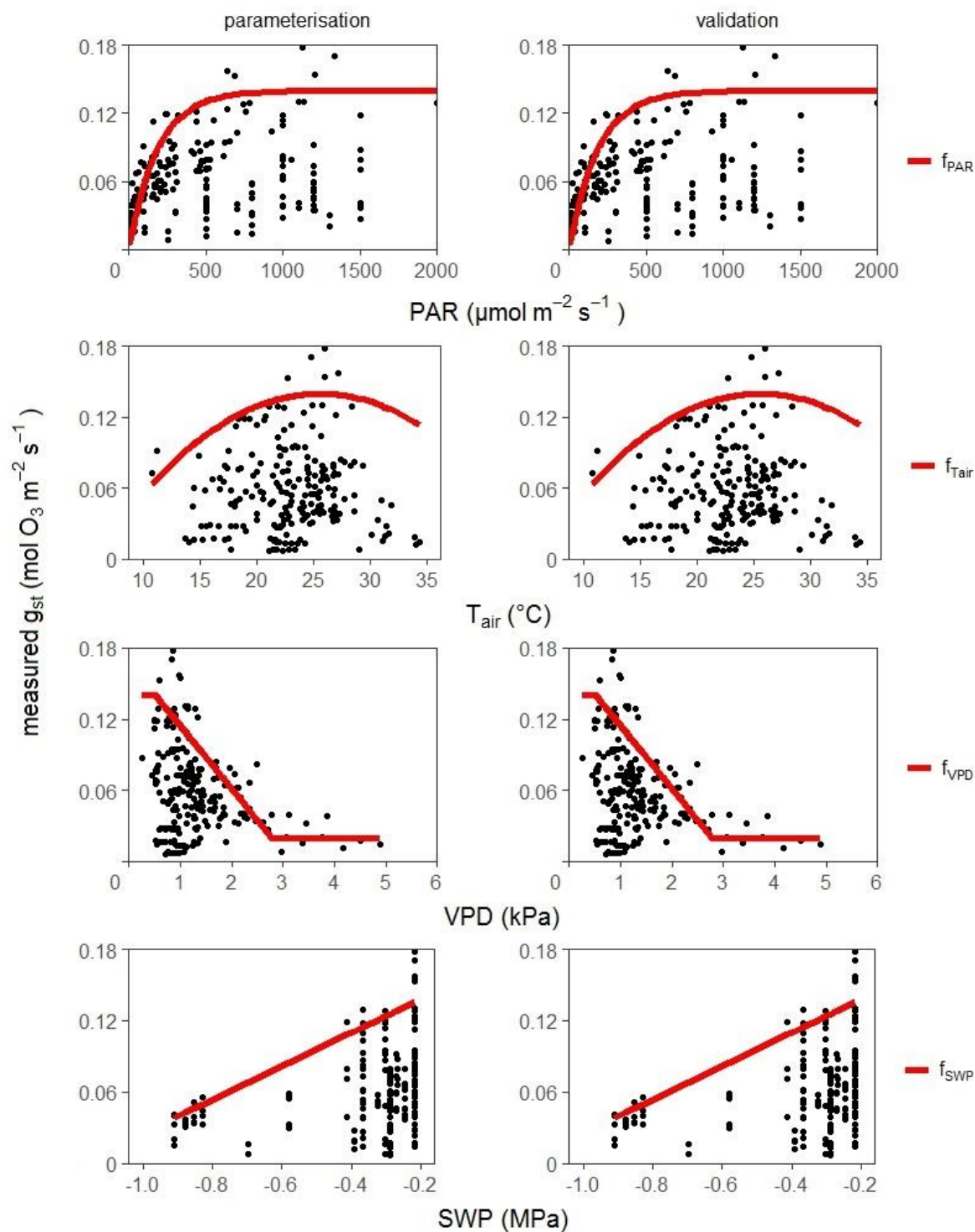


Fig. 6. Histograms of meteorological variabls for the training dataset (red) and the high O<sub>3</sub> uptake dataset (blue). The subplots represent global radiation  $R_g$  (A), minimum temperature  $T_{min}$  (B), maximum temperature  $T_{max}$  (C), mean temperature  $T_{mean}$  (D) and vapour pressure deficit VPD (E).

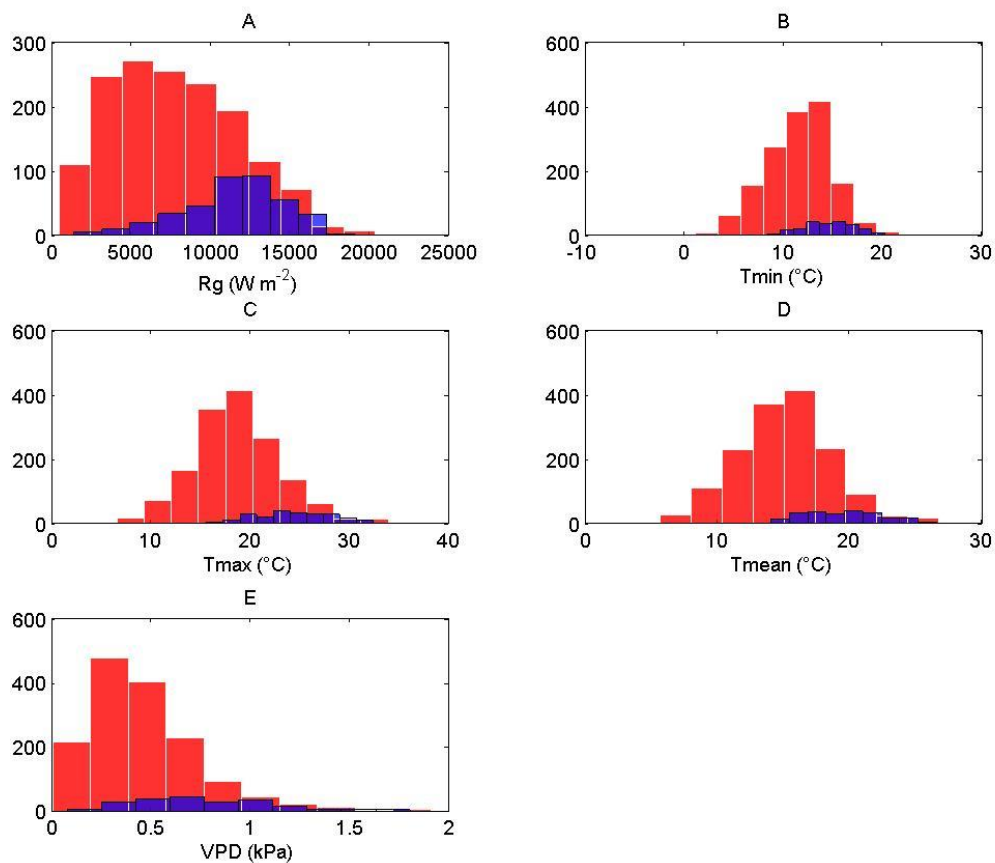
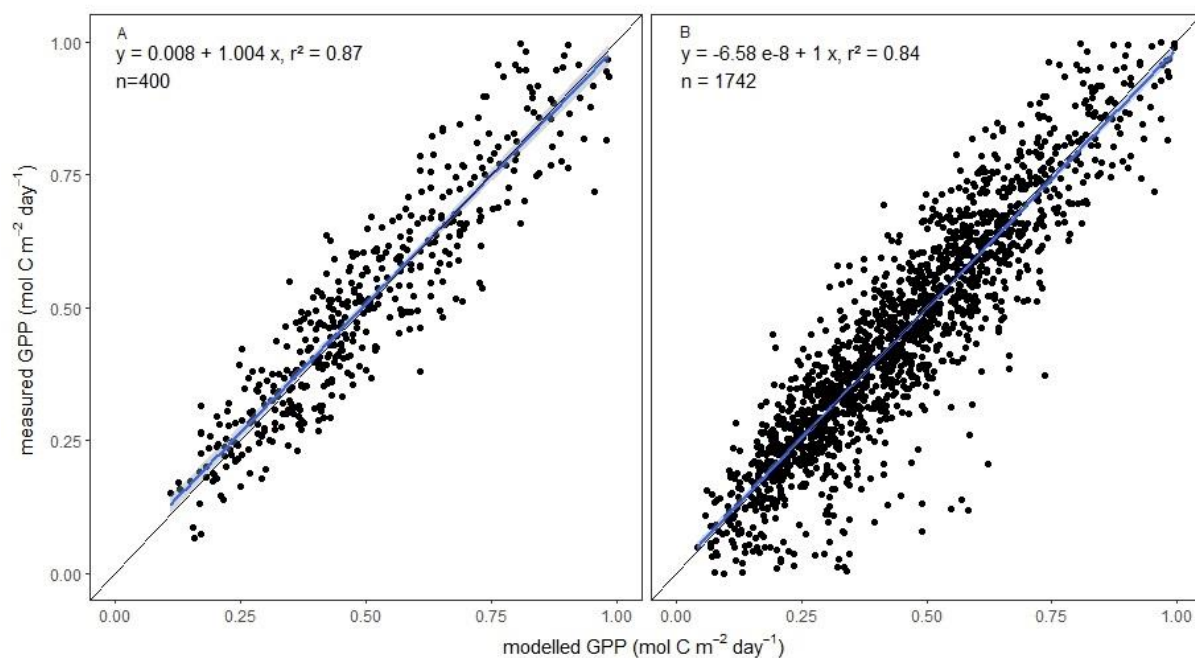


Fig. 7. Measured GPP is plotted as function of modelled GPP for two different datasets: (a) only the days before the first major O<sub>3</sub> peak in every year, (b) the training dataset with the days after the first major O<sub>3</sub> peak in every year, excluding those with high O<sub>3</sub> fluxes + six following days to train the network. The black line is the 1:1 line. The blue line is the regression fit including 95 % confidence intervals (in grey).



820

Fig. 8. Measured versus modelled gross primary productivity (GPP) for days used for model training and testing (A, B), for days on which an O<sub>3</sub> effect was assumed (C, D), and for the entire growing season (E, F). GPP model 1 was trained without days with the highest stomatal O<sub>3</sub> uptake, whereas GPP model 2 was trained to test for possible lag effects of O<sub>3</sub> on GPP. Black lines are fitted linear regression lines and grey lines mark the 95 % confidence bands. Also shown are p-values for the tests of slope and intercept from the regression  $y = a x + b$  being different from 1 and 0, respectively. Black lines are the fitted linear regression lines and grey lines are the 95 % confidence bands.

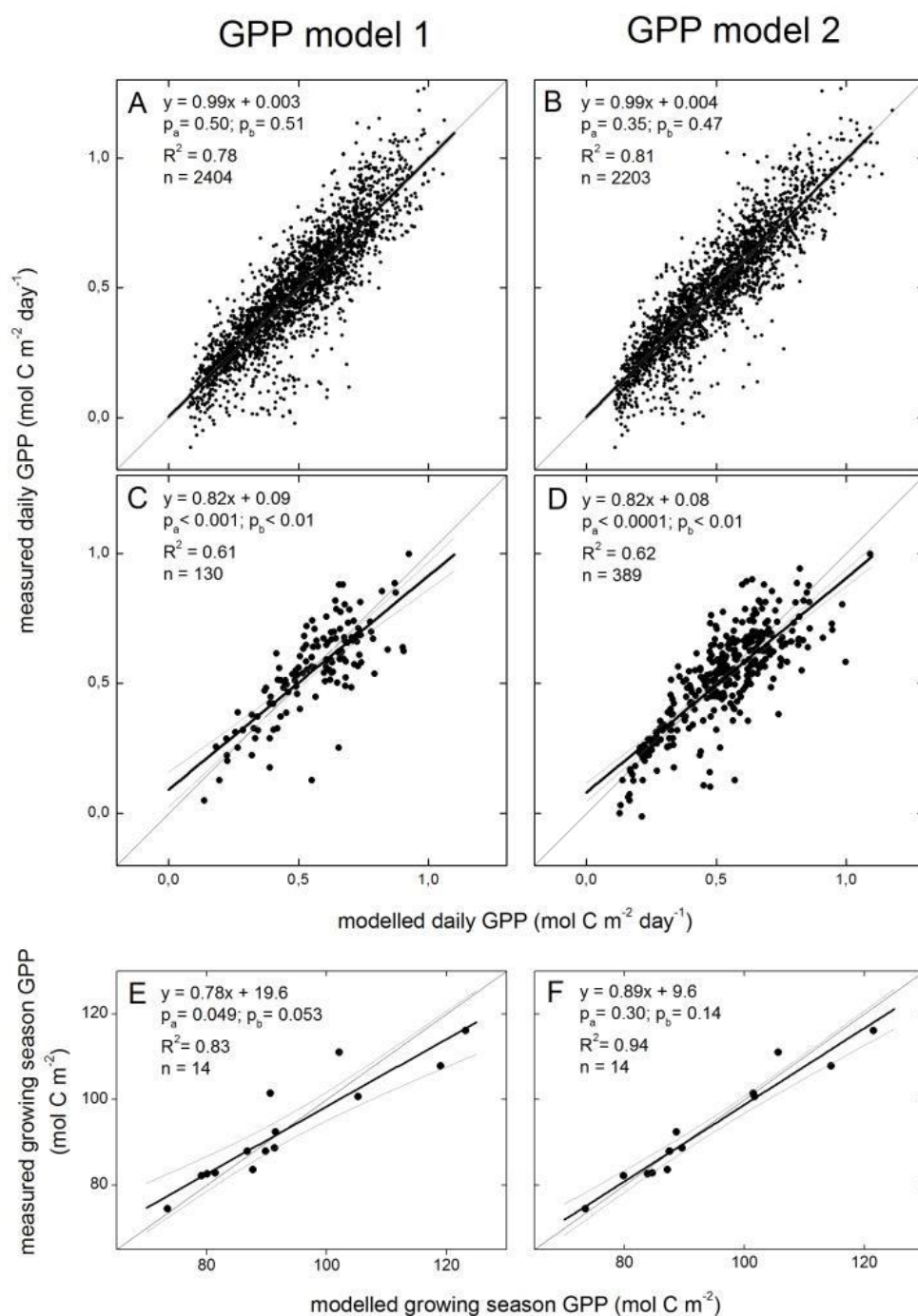
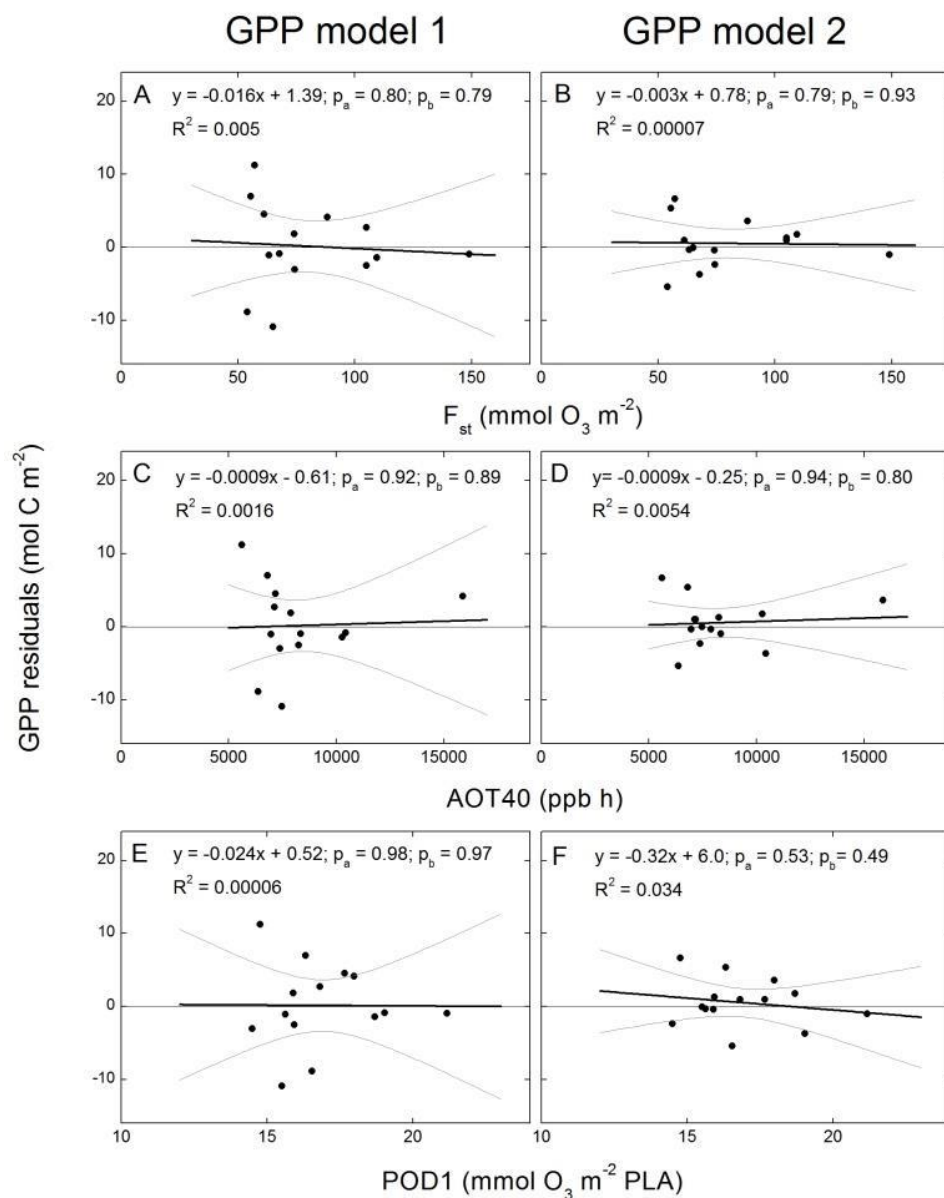




Fig. 9. Residuals of growing season gross primary productivity (GPP) ~~in-as a~~ function of (A, B) total stomatal O<sub>3</sub> flux over the growing season (F<sub>st</sub>), (C, D) AOT40, and (E, F) POD<sub>1</sub>. PLA = projected leaf area. Negative residuals indicate model overestimation of GPP. GPP model 1 was trained without days with the highest stomatal O<sub>3</sub> uptake, whereas GPP model 2 was trained to test for possible lag effects of O<sub>3</sub> on GPP. Black lines are fitted linear regression lines and grey lines mark the 95 % confidence bands. Also shown are p-values for the test of the slope and intercept from the regression  $y = a x + b$  being different from 0. (n = 14).



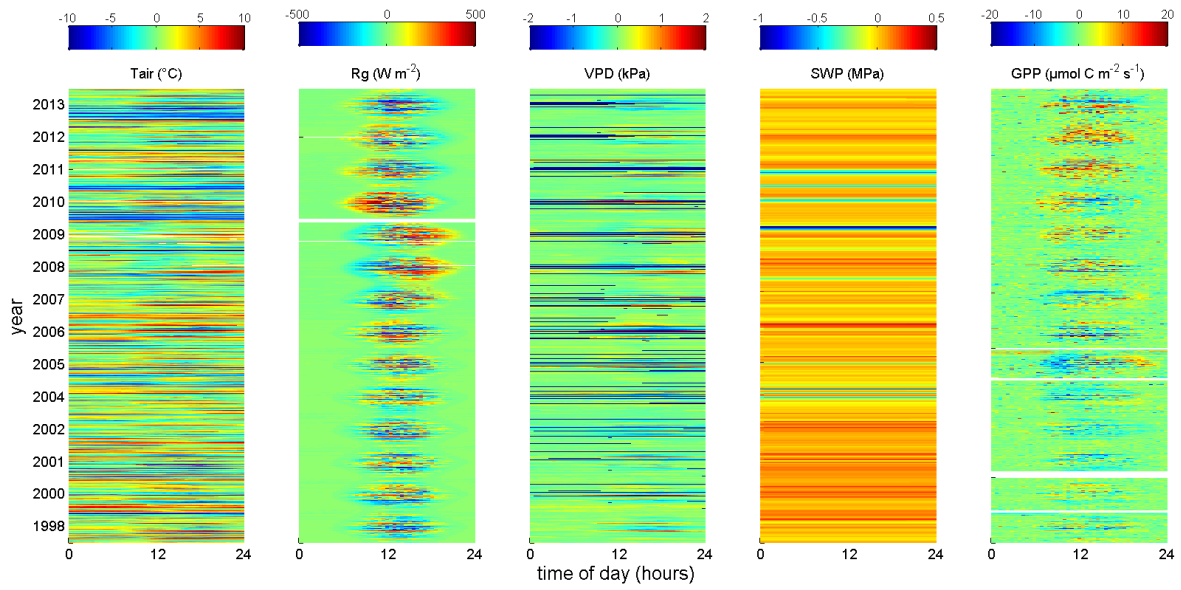


Fig. S1. Fingerprint of annual anomalies of the meteorological, soil water, and carbon exchange measurements compared to the averaged values shown in Fig. 1 and 2 of the main text. For each day of the year air temperature ( $T_{\text{air}}$ ), incoming global radiation ( $R_g$ ), vapour pressure deficit (VPD), soil water potential (SWP), and gross primary productivity (GPP) are plotted in function of time of the day.

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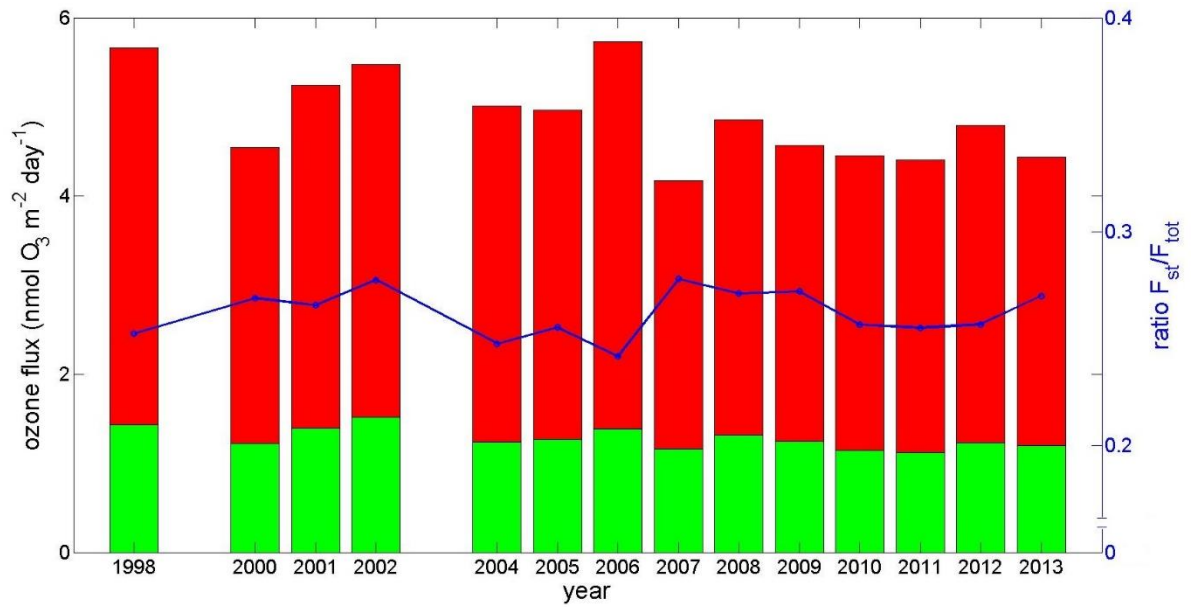
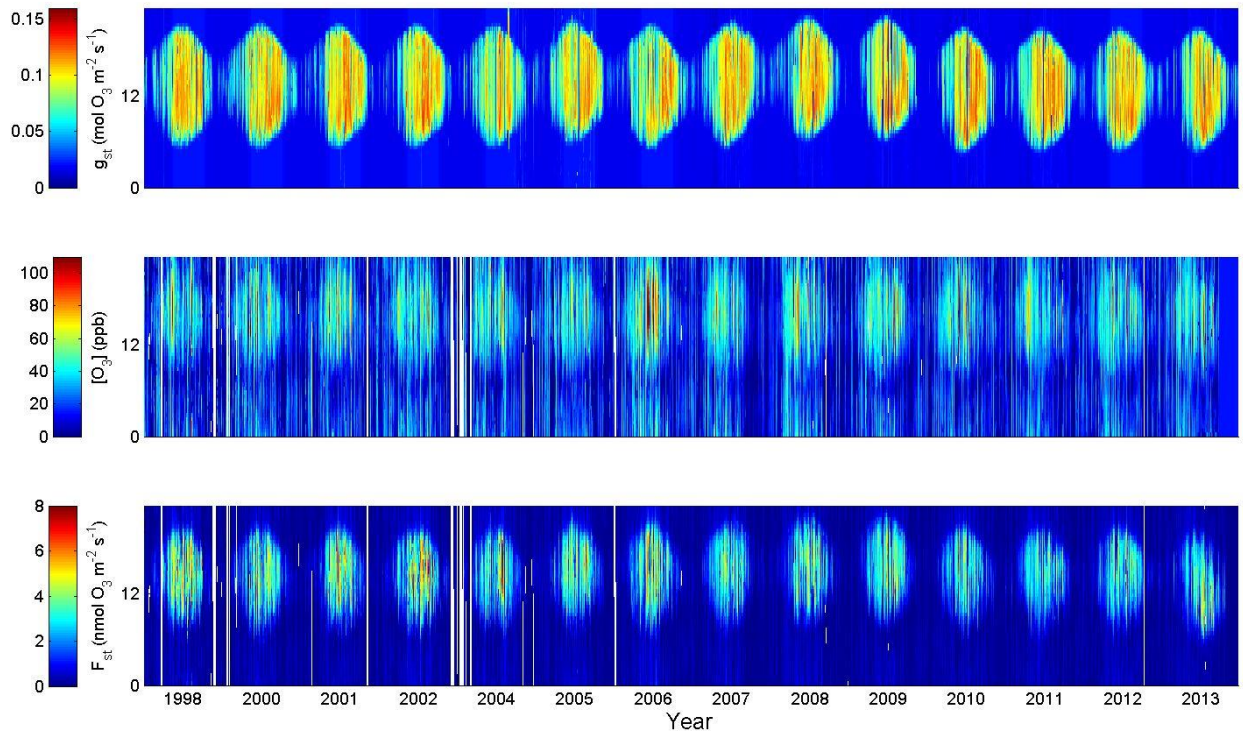


Fig. S2. Average daily  $O_3$  flux during the growing season and ratio of stomatal versus total ozone flux ( $F_{\text{st}}/F_{\text{tot}}$ ) for each year. The height of the bars represents the average total  $O_3$  flux. The green part of the bar is the stomatal  $O_3$  flux, whereas the red part is the non-stomatal  $O_3$  flux. The ratio  $F_{\text{st}}/F_{\text{tot}}$  is represented by the blue line.



10

Fig. S3. Fingerprints of stomatal conductance ( $g_{st}$ ), ozone mixing ratio ( $[\text{O}_3]$ ), and stomatal  $\text{O}_3$  uptake ( $F_{st}$ ). Day of year is plotted on the x-axis and hour of day on the y-axis.

Fig. S4. (A, B, C) The measured growing season GPP as function of total stomatal O<sub>3</sub> flux (F<sub>st</sub>), AOT40, and POD<sub>1</sub>. PLA = projected leaf area. (D, E, F, G) Time series of measured growing season GPP, F<sub>st</sub>, AOT40, and POD<sub>1</sub>.

