C/N ratio, stable isotope (δ¹³C, δ¹⁵N), and *n*-alkane patterns of brown mosses along hydrological gradients of low-centred polygons of the Siberian Arctic Romy Zibulski^{1,2}, Felix Wesener⁴, Heinz Wilkes^{3,5}, Birgit Plessen³, Luidmila A. Pestryakova⁶, Ulrike Herzschuh^{1,2,7}

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21 Abstract

Mosses are a major component of the arctic vegetation, particularly in wetlands. We present C/N 22 ratiosmolar, δ^{13} C and δ^{15} N data of 400 brown moss samples belonging to 10 species that were 23 24 collected along hydrological gradients within polygonal mires located on the southern Taymyr Peninsula and the Lena River delta in northern Siberia. Additionally, *n*-alkane patterns of six of these 25 species (16 samples) were investigated. The aim of the study is to see whether the inter- and 26 27 intraspecific differences in C/N, isotopic compositions and *n*-alkanes are indicative of habitat with particular respect to water-level. Overall, we find high variability in all investigated parameters for 28 two different moisture-related groups of moss species. The C/N_(m) ratios range between 11 and 53 29 (median: 32) and show large variations at the intraspecific level. However, species preferring a dry 30

habitat (xero-mesophilic mosses) show higher C/N(m) ratios than those preferring a wet habitat (meso-31 hygrophilic mosses). The δ^{13} C values range between -37.0 and -22.5‰ (median = -27.8‰). The δ^{15} N 32 values range between -6.6 and +1.7‰ (median = -2.2‰).We find differences in $\delta^{13}C$ and $\delta^{15}N$ 33 34 compositions between both habitat types. For some species of the meso-hygrophilic group, we suggest 35 that a relationship between the individual habitat water-level and isotopic composition can be inferred 36 as a function of microbial symbiosis. The *n*-alkane distribution also shows differences primarily 37 between xero-mesophilic and meso-hygrophilic mosses, i.e. having a dominance of *n*-alkanes with long $(n-C_{29}, n-C_{31})$ and intermediate $(n-C_{25})$ chain lengths, respectively. Overall, our results reveal that 38 39 C/N_(m) ratios, isotopic signals and *n*-alkanes of studied brown moss taxa from polygonal wetlands are 40 characteristic of their habitat.

41 **1. Introduction**

42 Specific physiological and morphological traits enable mosses to attain extensive surface coverage in 43 low-temperature ecosystems such as arctic tundra (Turetsky et al., 2012; Wasley et al., 2006). Mosses 44 are a major component of the biomass in arctic wetlands and contribute strongly to the biodiversity. In particular, the vegetation of the widely distributed polygonal tundra is rich in moss taxa, which partly 45 46 originates from the strong compositional turnover along a small-scale hydrologic gradient (Zibulski et 47 al. 2016). As an intermediate layer between air and the permafrost soil, mosses control the water cycle, 48 greenhouse gas and energy exchange (Blok et al., 2011; McFadden et al., 2003), and the structure of 49 the habitats of vascular plant communities in arctic landscapes (Beringer et al., 2001; Gornall et al., 50 2011; Malmer et al., 1994). Because of their low decomposition rates (Aerts et al., 1999; Turetsky, 51 2003), they contribute strongly to the peat and permafrost carbon pool (Tarnocai et al., 2009).

52 Despite the significance of mosses in high-latitude biodiversity and matter cycles only little is 53 known about their C/N ratio, stable isotope and *n*-alkane characteristics in comparison to vascular 54 plants. Such information is not only necessary to improve our understanding of the physiological 55 adaptation or plasticity of mosses to certain environmental characteristics, but can also be used when similar measurements of fossil plant material are interpreted as proxies of former ecological or
 environmental conditions (e.g. Birks, 1982).

Carbon-to-nitrogen ratios (C/N), stable carbon and nitrogen isotope values (δ^{13} C, δ^{15} N) and the 58 59 *n*-alkane fingerprints of bulk organic matter are among the most commonly measured parameters of 60 plant matter (Chambers and Charman, 2004). The C/N ratios of mosses are generally in the range of 61 those of higher terrestrial plants (Prahl et al., 1980) but are known to depend on the available nitrogen, 62 which originates in pristine regions from decomposition of organic matter, microbial activity or 63 atmosphere deposits (e.g. Chanway et al., 2014; Lee et al., 2009). We provide C/N ratios by weight of 64 arctic brown mosses, in anticipation that they will be useful for comparative palaeo-environmental 65 reconstructions (Andersson et al. 2011,) and in the evaluation of organic matter sources in Russian 66 permafrost soils with regard to species and habitat-specific patterns. Furthermore, C/N ratios may be related to growth form, i.e. higher ratios can be expected for mosses which compete with vascular 67 68 plants for light and thus need to invest in a high stem-stability (Sveinbjörnsson and Oechel, 1992).

As well as the C/N ratio, the δ^{13} C composition is barely altered by decomposition processes in the 69 soil and also suitable for palaeo-environmental reconstructions. The $\delta^{13}C$ ratio of an individual plant 70 are a mixed signal of the physiological traits of the species and the individual's direct environment. 71 72 Mosses use the C₃ pathway for carbon assimilation (O'Leary, 1988, Farquhar et al., 1989) and despite a lack of stomata in the photosynthetic active parts, they have a similar range in their δ^{13} C values of 73 74 between -24 and -32 ‰ (Ménot and Burns, 2001; Rundel et al., 1979; Smith and Epstein, 1971) as vascular C₃ plants. Differences in δ^{13} C values among several species can be explained by individual 75 plant physiology and biochemistry (Galimov, 2000). Differences within a single species have been 76 77 linked with environmental conditions such as temperature (Skrzypek et al., 2007; Waite and Sack, 78 2011), growing depth below water-level (Raghoebarsing et al., 2005), position within a cushion or 79 hummock (Price et al., 1997), lipid content (Rundel et al., 1979) or the influence of microbial 80 symbioses (Liebner et al., 2011; Vile et al., 2014). Furthermore, even differences among branches and 81 stems of single individuals have been reported (Loader et al., 2007). However, almost all of these 82 studies were made on Sphagnaceae, which are not representative of all mosses because of their

specific morphology (i.e. the occurrence of a photosynthetic active cell type and a dead cell type, which is responsible for water storage and often an additional coating of the photosynthetic active cell) and their specific habitat preferences (i.e. preferring acidic wetlands). Hence, this information cannot simply be transferred to brown mosses – which form the major component in northern Siberian lowlands – due to their different morphology.

88 Most studies on δ^{15} N values of moss material have focused on the anthropogenic impact on the 89 nitrogen cycle (Harmens et al., 2011; Liu et al., 2008; Poikolainen et al., 2009), whereas reports on the 89 δ^{15} N compositions of mosses from relatively pristine ecosystems such as the Arctic are rare or have 91 been investigated in relation to the study of bird colonies (e.g. Lee et al., 2009). Potentially, such 92 information can indicate pathways and sources of the nutrient supply in these N-limited ecosystems 93 (Kielland, 1997; Michelsen et al., 1996, 1998).

94 Compared with vascular plants and Sphagnum species, relatively few investigations of n-alkane 95 patterns of brown mosses are available. Palaeoenvironmental reconstructions use the potential of 96 *n*-alkanes to distinguish between different plant groups (Ficken et al., 1998, moisture conditions 97 (Pancost et al., 2000, Nichols et al 2006, Zhou et al. 2010), whether organic material is to decomposed 98 and changes in temperature (Feakins et al., 2016, Sachse et al., 2006) along distinct gradients. 99 Sphagnum species, for example, show a dominance of $n-C_{23}$ and $n-C_{25}$ homologues (Baas et al., 2000) 100 that are comparable to the pattern of vascular submerged plants (Ficken et al., 2000), and ratios are 101 used as a proxy for wet moisture conditions. Other studies discuss the suitability of *n*-alkane patterns 102 in moss species for chemotaxonomical studies on recent and fossil material (Bush and McInerney, 103 2013; Nott et al., 2000; Schellekens and Buurman, 2011). A greater protection potential of waxes with 104 a higher content of long-chain *n*-alkanes against solar irradiation or alternatively an enhanced loss of 105 short-chain *n*-alkanes by evaporation have been suggested as possible mechanisms to cause different 106 *n*-alkane patterns in leaf waxes of trees (Sachse et al., 2006). However, with respect to mosses the 107 pattern and mechanisms are even less understood.

108 This study presents C/N_(m), isotopic (δ^{13} C, δ^{15} N) and *n*-alkane characteristics of mosses from low-109 centred polygons in northern Siberia. Low-centred polygons are geomorphological forms in arctic 110 landscapes originating from frost-heave processes in the soil. They are characterized by elevated dry 111 rims and a water-saturated or water-filled centre. This centimetre-scale hydrological gradient is well 112 reflected by a strong turnover in the vascular plant and moss composition (Zibulski et al., 2016). We 113 investigate the relationship between the $C/N_{(m)}$, isotopic and *n*-alkane compositions and the 114 hydrological conditions within low-centred polygons. In particular, we aim to reveal whether 115 fingerprints are habitat-specific or rather species-specific.

116 **2. Material and methods**

117 **2.1 Sites**

118 The plant material was collected during the vegetation season (July-August) from eight low-centred polygons located along a zonal vegetation gradient ranging from open forest via the forest-tundra 119 120 intersection to subarctic tundra (Matveev, 1989) to obtain a representative sample set of northern Siberian lowlands (Fig. 1). Six of the polygons (06/P, 17/P, P3/I, P3/II, P3/III, 12/P), sampled in 2011, 121 are located in the Khatanga River region (70-72° N, 97-102° E, southern Taymyr Peninsula) and a 122 further two polygons (LP1 and LP2), sampled in 2012, are located on Samoylov Island in the Lena 123 River Delta (72.4° N, 126.5° E). The climate is cold-humid (Khatanga climate station annual mean 124 temperature and precipitation: -13.2°C and 272 mm, Rivas-Martinez and Rivas-Saenz, 2009; 125 126 Samoylov climate station annual mean temperature and precipitation: -12.5°C and 232.7 mm, Boike et al., 2013)). 127

128 **2.2 Sampling and studied moss species**

A representative continuous transect of adjacent one square-metre plots (from rim to rim) was selected for sampling in each polygon. The surface height in relation to water-level (measured at the centre of each plot) and vegetation (abundance information) were recorded. Individual characteristics of each low-centred polygon are presented in Table A1. Moss plants were hand-picked, dried in the field and transported to the Alfred Wegener Institute in Potsdam. Taxa identification is based on the relevant literature (Frahm and Frey, 2004; Lobin et al., 1995; Smith, 1978) and the Bryophyte Flora of North America internet source (http://www.efloras.org). In total, eight species that were observed to have different water-regime preferences were included in the analyses. To approximate water-regime preferences of each species, we calculated the mean plantposition in relation to water-level (h) from the recorded plot surface height. *Warnstorfia exannulata* was observed to grow as part of a swinging mat at water-level, it was assigned a preference for waterlevel. Using this information the species were classified as either as of xero-mesophilic mosses (mean plant-position between 13 and 16 cm in relation to water-level) or to the group of meso-hygrophilic mosses (mean plant-position between 3 and -30 cm in relation to water-level) to ease the presentation

143 of results.

144**Table 1:** Calculated mean plant-position relative to the water-level and the classification of the sampled brown moss species.145(* Samples of Warnstorfia examulata were growing on a swinging mat at water-level and mean plant position was thus set to

146 *zero.*)

species	abbreviation	h (mean plant-position in relation to water-level in cm)	classification		
Hylocomium splendens	Hyl_spl	16	xero-mesophilic		
Tomentypnum nitens	Tom nit	13	xero-mesophilic		
Aulacomnium turgidum	Aul tur	13	xero-mesophilic		
Aulacomnium palustre	Aul pal	13	xero-mesophilic		
Hamatocaulis lapponi	Ham_lap	3	meso-hygrophilic		
Warnstorfia exannulata	War exa	0 *	meso-hygrophilic		
Meesia triquetra	Mee tri	-1	meso-hygrophilic		
Drepanocladus revolvens	Dre rev	-5	meso-hygrophilic		
Scorpidium scorpioides	Sco_sco	-13	meso-hygrophilic		
Calliergon giganteum	Cal gig	-30	meso-hygrophilic		

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148 2.4 Measurements of C/N_(m) ratios, stable carbon and nitrogen isotope values and 149 *n*-alkane distributions

Selected plant material (i.e. apical parts of a specimen) was rinsed with de-ionised water and 150 151 mechanically cleaned from organic particulate material. The content of carbon and nitrogen and the 152 ratio of stable isotopes were measured with a DELTAplusXL isotope ratio mass spectrometer (Thermo 153 Fischer Scientific) coupled to an elemental analyser (NC2500 Carlo Erba) via a CONFLOW III Interface. Due to the relatively wide range of C/N ratios of mosses, we used about 1.5 mg for each 154 155 carbon stable isotope measurement (n = 400) and a replicate of 3 mg for each nitrogen stable isotope 156 measurement (n = 326) and the analysis of elemental composition. The high weight needed for the nitrogen sample replicates prevented the measurement of $\delta^{15}N$ and thus the C/N_(m) for some samples. 157 The calibration for carbon was performed using an urea standard and a δ^{13} C isotopic standard (IAEA 158 159 CH-7). The nitrogen contents were calibrated against an acetanilide standard and the nitrogen isotopic 160 composition with ammonium sulphate standard (IAEA N-1). The reliability of the method was 161 checked with the NIST plant standard SRM 1547. The isotopic ratios are given in delta notation 162 relative to VPDB for δ^{13} C values and relative to air for δ^{15} N values, respectively. The reproducibility 163 for replicate analyses is 0.2% for carbon and nitrogen and 0.2‰ for δ^{13} C and δ^{15} N values.

n-Alkane analyses were performed on a subset of 16 samples. We took material from the polygon 164 complex P3 from the Khatanga region (72.149° N, 102.693° E), which comprises three successive 165 166 low-centred polygons (Table A1) to exclude effects of environmental conditions at different locations. 167 The moss samples were washed, identified and air-dried. They were weighed (150-1000 mg dry 168 weight) and samples extracted with an accelerated solvent extractor (ASE) (Dionex, Sunnyvale; USA) 169 using Cl₂Me:MeOH (10:1) at 5 bar and 75°C. The extract was separated in to aliphatic hydrocarbon, 170 aromatic hydrocarbon and nitrogen, sulphur and oxygen compound fractions using medium pressure 171 liquid chromatography following Radke et al. (1980). Five μg of the quantification standard (5 α -172 androstane, 1-ethylpyrene, 5 α -androstan-17-one and erucic acid) were added. Gas chromatography (GC) of aliphatic hydrocarbon fractions was performed using a GC Agilent 6890 equipped with an 173 174 Ultra-1 fused silica capillary column (Model Agilent 19091A-105, length 50 m, inner diameter 200 µm, film thickness 0.33 µm). Helium was used as a carrier gas at a constant flow rate of 1 ml 175 176 min⁻¹. The GC oven was heated from 40 °C (2 min hold time) to 300 °C (65 min hold time) at a rate of 5°C per minute. The samples were injected by means of splitless injection. Quantitative evaluation of 177 178 data was done with ChemStation software. Additionally, we calculated the average chain length (ACL₂₁₋₃₃) assumed to represent a proxy for moisture (Andersson et al, 2011) and temperature (Bush 179 and McInerney, 2015) with a comprehensible extended range from $n-C_{21}$ to $n-C_{33}$ and the proxy ratio 180 P_{aq}, which was developed as a proxy ratio to distinguish submerged or floating aquatic macrophytes 181 182 from emergent and terrestrial plants (Ficken et al., 2000):

$$ACL_n = \frac{\sum (n \times C_n)}{\sum C_n}$$
, $n = 21 - 33$

$$P_{aq} = \frac{(C_{23} + C_{25})}{(C_{23} + C_{25} + C_{29} + C_{31})}$$

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184 **2.5 Statistical Tests**

We calculated the range, .25, .5, and .75 quantiles of δ^{13} C and δ^{15} N values and the C/N_(m) ratios for all species (Table A2). Significant differences in the C/N_(m) ratios and δ^{13} C and δ^{15} N values among different groups were assessed with a t-test. In addition, we performed linear regression between the obtained values for each individual species and surface height. All analyses were implemented in R version 3.2.0 (R Core Team, 2015). Furthermore, we performed a PCA with the percentage of *n*-alkane homologues (square-root transformed) in R using the vegan package version 2.0-10 (Oksanen et al., 2013).

192 **3. Results**

The overall ranges in C/N_(m) ratios of both groups have a broad overlap (xero-mesophilic: 22.5–67.9; meso-hygrophilic: 15.4–70.4). However, the medians of the C/N_(m) values of the xero-mesophilic species ranging from 47.6 to 52.9 (Fig. 2) are significantly higher than those of the meso-hygrophilic group, which range from 37.1 to 46.5 (W = 18280, $p \ll 0.001$). The C/N_(m) ratios show no intraspecific relations among individual species and water-level (Fig. 3a), except for *Tomentypnum nitens* (r² = 0.11, p < 0.05).

The δ^{13} C values of the meso-hygrophilic group (-34.9 to -22.5‰) cover the range of the xero-199 mesophilic group (-32.1 to -24.2%), which have a noticeably lower variance in δ^{13} C values (Fig. 2). 200 The medians of the individual species in the xero-mesophilic group (range: -29.4 to -27.1‰) are 201 significantly different (W = 16232, p = 0.008) from those of the meso-hygrophilic group (range: -29.2 202 to -24.8%). δ^{13} C values of Meesia triquetra, Drepanocladus revolvens, and Scorpidium scorpioides 203 (all belonging to the meso-hydrophilic group) are significantly positively related with the position of 204 205 the water-level (Fig. 3b), while no member of the xero-mesophilic group revealed such relationship 206 (Fig. 3a).

207 The ranges of δ^{15} N values of both groups are rather similar (Fig. 2). However, individual species 208 medians of both groups are significantly different (t = -6.96, *p* << 0.001; xero-mesophilic group; -3.2 209 to -2.7‰, meso-hygrophilic group: -2.9 to -0.1). *Drepanocladus revolvens, Scorpidium scorpioides,* and *Calliergon giganteum*, all belonging to the meso-hygrophilic group, exhibit a positive relationship between the δ^{15} N values and position relative to water-level (Fig. 3b.

212 *n*-Alkane distributions show the expected predominance of the odd chain length (Table 2). The absolute *n*-alkane concentrations (*n*-C₁₉ to *n*-C₃₃) range from 34 to 238 μ g g⁻¹ of dry weight. The 213 214 individual taxa show significant differences with respect to carbon number of the most abundant 215 *n*-alkane (*n*-C_{max}) forming unimodal distribution patterns. It is *n*-C₃₁ for *Tomentypnum nitens*, which is 216 also the only species containing n-C₃₃ in detectable amounts, n-C₂₉ for Aulacomnium turgidum, n-C₂₇ 217 for Aulacomnium palustre, $n-C_{27}$ and $n-C_{25}$ for the Drepanocladus taxa, and $n-C_{max} = 25$ for Scorpidium scorpioides. These differences in the n-alkane composition largely relate with the species-218 219 preferred mean plant-position relative to water-level (Fig. 4).

Evaluations of the *n*-alkane biomarker proxies, ACL_{21-33} and P_{aq} , also show a clear division between the xero-mesophilic and the meso-hygrophilic species groups (Table 2), whereas intraspecific variations are rather small (with the exception of *Drepanocladus*). The xero-mesophilic group is notable for high averages of ACL_{21-33} (28.41) and P_{aq} (0.17) compared to low averages of ACL_{21-33} (25.61) and P_{aq} (0.87) for the meso-hygrophilic group.

The observed trend is also assumed in the biplot of the first two PCA axes, even though their explained variance is relatively low (16.9%) in the dataset (Fig. 5). The first axis separates xeromesophilic from meso-hygrophilic taxa. *Aulacomnium* taxa are located in the upper range and *Tomentypnum nitens* in the lower range of the second PCA axis, while no trend is observed within the meso-hygrophilic group along the second axes.

Table 2: The concentration ($\mu g/g dry$ weight) and ACL_{21-33} and P_{aq} (after Ficken et al., 2000) of n-alkanes in brown moss samples. (Numbers in brackets relates to the sample number.)

species	<i>n</i> -C ₁₉	<i>n</i> -C ₂₁	<i>n</i> -C ₂₃	<i>n</i> -C ₂₅	<i>n</i> -C ₂₇	<i>n</i> -C ₂₉	<i>n</i> -C ₃₁	<i>n</i> -C ₃₃	ACL21-33	Paq
Xero-mesor	hilic hab	itat group								4
Tom_nit (1)	0	2.0689	2.6374	6.0816	10.7454	17.7273	34.0742	0	28.86	0.14
Tom_nit (2)	0.5571	1.041	1.4493	2.6964	9.1008	8.4867	20.2818	7.2312	29.47	0.13
Tom_nit (3)	0.9936	1.456	2.175	2.8712	9.6118	7.2721	21.2309	4.9717	29.14	0.15
Aul_tur (1)	0	0.968	1.126	2.2001	8.9548	23.5267	3.5595	0	28.15	0.11
Aul_tur (2)	1.028	1.3468	1.2794	4.8797	18.7427	50.9291	5.5645	0	28.22	0.10
Aul_pal (1)	0	1.9298	2.5459	4.7646	20.5085	7.5826	7.979	0	27.35	0.32
Aul_pal (2)	0.7341	1.0767	1.5183	2.9728	12.3293	11.0909	4.9109	0	27.69	0.22

Meso-hygrophilic habitat group										
Dre_rev (1)	0.7868	1.5426	2.3659	43.833	19.0836	5.1551	3.5637	0	25.92	0.84
Dre_rev(2)	0.981	1.5227	2.7605	22.5638	14.5103	6.5356	7.5623	0	26.6	0.640
Dre_sp.	1.4696	1.9968	5.0986	29.6729	30.4582	9.9108	5.1317	0	26.38	0.70
Sco_Sco(1)	0	3.7612	11.7002	133.4207	29.0024	6.2023	1.9425	0	25.3	0.95
Sco_Sco (2)	0	3.8911	10.4693	93.7009	21.4601	5.8531	1.7706	0	25.29	0.93
Sco_Sco (3)	0.8856	2.7949	11.8988	134.378	24.2348	5.4035	2.4969	0	25.28	0.95
Sco_Sco (4)	1.6217	5.5813	11.705	101.7602	21.3126	6.9161	2.4522	0	25.26	0.92
Sco_Sco (5)	1.4083	3.8857	11.8966	121.5701	23.3167	4.315	1.7261	0	25.21	0.96
Sco_Sco (6)	1.345	2.6672	19.4796	170.3015	34.3255	6.8599	3.1732	0	25.28	0.95

4.

232

Discussion

233 4.1 C/N_(m) ratios

234 The C/N ratios of mosses from polygonal tundra in Northern Siberia are relatively low compared with 235 those obtained for mosses from Antarctic bogs that range between 80 and 100 (Björck et al., 1991) or 236 from western Canada that range between 55 and 76 (Kuhry and Vitt, 1996). However, neither the taxa 237 xeric and mesic growing conditions were sampled in Antarctica and Canada. All investigated species 238 are considered as ectohydric mosses, which receive nitrogen mostly from precipitation deposits (Ayres 239 et al., 2006). Our results reveal that averaged C/N_(m) ratios for the xero-mesophilic moss group are 240 higher than for the meso-hygrophilic group, probably reflecting the known difference between 241 terrestrial and aquatic plants (Meyers and Ishiwatari, 1993). There are two possible impacts, which can 242 influence the C/N ratio of these groups: (1) competition with vascular plants and (2) accessibility of 243 nitrogen pools (2)g. (1) If moss plants invest in a high stem-to-leaf biomass ratio, which results in a 244 high C/N_(m) ratio, they will increase their height and stability, and thus their competitive ability 245 against vascular plants for light (Sveinbjörnsson and Oechel, 1992). Furthermore, the low N input by 246 precipitation and a low N content of moss litter slows down the fungal and bacterial N mineralisation 247 which increases the thickness of moss litter mats (Gornall et al., 2007; Turetsky, 2003). This in turn 248 will increase the isolating function of moss mats, thus negatively affecting seed germination of 249 vascular plants (Gornall et al., 2007). (2) Lower C/N_(m) ratios of meso-hygrophilic mosses may 250 originate from higher amounts of dissolved nitrogen in polygon waters as a result of high net primary productivity, the presence of N₂-fixers such as cyanobacteria, and the exudations of zooplankton. 251 252 Frahm (2001) assumes that loose epiphytic and endophytic symbiotic relationships between mosses

253 and cyanobacteria are probably restricted to wetland taxa. Lindo et al. (2013) report such associations 254 between brown mosses and cyanobacteria. Thus, the N supply is better for brown mosses preferring 255 meso-hygrophilic than xero-mesophilic habitats, and the respective taxa accordingly have lower 256 C/N_(m) ratios indicating habitat-specific variation in C/N_(m) ratios. We expected to also find intraspecific variations between C/N_(m) ratios and water-level. The large variability in the C/N data 257 may be a result of atmospheric conditions and organic matter degradation being the principal sources 258 259 at xeric sites, whereas in mesic and wet sites microbial symbionts play an important role in the C/N ratio. However, the signal-to-noise ratio is probably too low to give a meaningful result because only 260 261 the average water level of each plot but not of each individual plant was recorded.

262 **4.2** δ^{13} C values

263 With respect to bryophytes, most isotopic studies have hitherto been performed on Sphagnum (Markel et al., 2010, Ménot and Burns, 2001) while our study focuses on brown mosses - a major component 264 265 in Siberian wetlands. The intraspecific variability for some meso-hygrophilic species (i.e. Meesia triquetra, Drepanocladus revolvens, Scorpidium scorpioides) show that the δ^{13} C signals are related to 266 267 the hydrological conditions at the growing site of each individual, i.e. individuals growing at dry sites showed higher medial δ^{13} C values than those growing at wet sites. A difference among the two habitat 268 269 groups is observed; they partly contradict the intraspecific findings in that some of the xero-mesophilic 270 species known to prefer dry rims such as Hylocomium splendens and Tomentypnum nitens have particularly low δ^{13} C medians. 271

The detected differences in moss δ^{13} C values, particularly of the meso-hygrophilic group, either 272 273 reflect a source signal depending on water-level or a physiological reaction of the plant related to water-level (Bramley-Alves et al., 2014; Proctor et al., 1992). Mosses are typical C₃ plants (Farquhar 274 275 et al., 1989, Rundel et al., 1979) characterized by a high CO₂ compensation point (Bain and Proctor, 276 1980; Dilks and Proctor, 1975; Salvucci and Bowes, 1981). The high availability of atmospheric CO₂ and elevated diffusion rates of CO₂ in air compared to water (O'Leary) result in typical terrestrial C₃ 277 land plant δ^{13} C characteristics, because of a decreasing cell water pressure in dry habitats which entails 278 a strong discrimination rate against ¹³CO₂ induced by RuBisCO (Rice and Giles, 1996). With respect 279 to the xero-mesophilic group, we observe an increase in discrimination against ¹³C from taxa 280

281 preferring a low position relative to the water-level (e.g. Aulacomnium taxa) than those preferring high positions (e.g. Hylocomium splendens). In contrast, if plant tissue is coated by a water film, the cell 282 water pressure should reach an optimum, which is expected to results in a weaker discrimination rate 283 against ¹³C by RuBisCO (Rice and Giles, 1996), because of a source restriction by the slower diffusion 284 rate of CO₂ in water (Lloyd and Farquhar, 1994). A lower carbon isotope discrimination related to 285 286 water saturation is observed for only three species out of six meso-hygrophilic mosses. However, this basic signal may be masked by variations in δ^{13} C values of different carbon sources, which are 287 expected to be more influential for meso-hygrophilic mosses in water-saturated conditions. Ménot and 288 Burns (2001) studied intraspecific variations for three Sphagnum species, which prefer three different 289 habitat types (dry, meso, wet) along an elevational gradient, which was positively correlated with 290 precipitation. They find a decline in discrimination against ¹³C with increasing wetness, and similar to 291 292 our results no relationship for species with a strong wet -preference. This is attributed to the variation in δ^{13} C from highly varying dissolved inorganic carbon (Proctor et al., 1992). Mosses potentially 293 access ¹³C-depleted CO₂ that originates from oxidation of typically strongly ¹³C-depleted biogenic 294 295 methane by methanotrophic microorganisms (Kip et al., 2010; Liebner et al., 2011; Raghoebarsing et 296 al., 2005). Studies by Nichols et al. (2009) show that a higher water level at the peat surface is crucial 297 for high methane-derived CO₂ release. Furthermore, symbiosis with methanotrophs enhances the moisture-related effect on the δ^{13} C signal of bryophytes. Endophytic microorganisms in hyalocytes of 298 299 submerged Sphagnum (Raghoebarsing et al., 2005) or epiphytic microorganisms on submerged brown mosses (Liebner et al., 2011) are presumed to provide ¹³C-depleted CO₂ directly to the lamina cells of 300 301 mosses. The studies of Ruttner (1947) and Bain and Proctor (1980) show that, in general, moss taxa 302 are incapable of bicarbonate uptake. Hence, bicarbonate, known to be a carbon source for submerged 303 vascular plants and algae (Herzschuh et al., 2010; Merz, 1992), can most probably be excluded as a carbon source for moss and thus as a factor influencing the δ^{13} C value. Moreover, the bicarbonate 304 content in pond waters in northern Siberian landscapes is very low (Wetterich et al., 2008). Other 305 sources of ¹³C-depleted CO₂ are surface run-off during spring flooding, rain events and decomposition 306 307 processes in the pond (Leng and Marshall, 2004; Maberly et al., 2013). Yet we cannot fully eliminate 308 the possibility that the measured bulk material was contaminated in parts with epiphytic or endophytic

309 microorganisms. The overall isotopic composition would, however, likely be unaffected, as Ménot and Burns (2001) have shown that the δ^{13} C values of bulk organic material and alpha-cellulose of 310 311 Sphagnum are very similar. Thus, the large ranges within several species of meso-hgryophilic habitats in arctic regions suggest that the existence of open water leads to more depleted $\delta^{13}C$ values and 312 measurements of the isotopic composition of methane when present and microbial groups in the water 313 314 and terrestrial litter should be possible. Finally, considering the relationship of selected brown mosses 315 to mean plant position, the complex origin of plant-available carbon makes it difficult to interpret the δ^{13} C record, especially for meso-hygrophilic brown mosses as well as *Sphagnum* (Prince et al., 1997, 316 Ménot and Burns, 2001). 317

318 **4.3** δ^{15} N isotopes

Like δ^{13} C, the interpretation of stable nitrogen isotope compositions of mosses is challenging because, again, source signals need to be separated from those originating from physiological isotopic discrimination processes. Our results yield relatively ¹⁵N-depleted δ^{15} N values for xero-mesophilic mosses growing preferentially on rims compared to meso-hygrophilic mosses.

The terrestrial arctic systems are generally thought to be nitrogen limited (Gordon et al., 2001; 323 Kielland, 1997). On the rim sites, atmospheric deposition can be considered to be the most important 324 325 source for nitrogen (Jonasson and Shaver, 1999) originating from fog, dew, precipitation and surface 326 run-off (Sveinbjörnsson and Oechel, 1992). However, most of the nitrogen available to rim mosses originates from recycling of already ¹⁵N-depleted higher plant and moss litter (Turetsky, 2003). The 327 328 ectohydric morphology enables an efficient nutrient uptake across the entire moss plant surface via 329 trapped water. In fact, the meso-hygrophilic group has a higher N content than the xero-mesophilic group (see section on C/N ratio). Inorganic nitrogen, but especially the high amounts of organic 330 nitrogen provided by N-mineralization in tundra soils (Kielland, 1995) are important for mosses 331 332 growing on the rather dry sites such as the polygonal rims (Atkin, 1996).

333 Three of the investigated submerged or floating moss species show a significant positive 334 relationship between water-level and δ^{15} N values. These results are similar to those of Asada et al. 335 (2005) who tested a relationship between δ^{15} N values of different *Sphagnum* species and their position relative to the groundwater level, which they assumed to originate from different nitrogen sources and different internal fractionating processes. We assume that the often heavier nitrogen isotope composition of meso-hygrophilic brown moss individuals originates from the high degree of symbiotic associations with aquatic atmospheric nitrogen-fixing autotrophic microorganisms such as *Nostoc* or *Anabena* (Lindo et al., 2013) or methanotrophs (Vile et al., 2014). The high spatial degree with endoor ectosymbiotic N₂-fixing microorganisms enables the direct uptake of their nitrogen products, which is similar to that of N₂ in air.

343 **4.4** *n*-alkane patterns

344 Compared to vascular plants that are characterized by a thick leaf-wax layer, mosses produce only a 345 small amount of *n*-alkanes (Baas et al., 2000; Ficken et al., 1998). Like previous studies on vascular 346 plants (Aichner et al., 2010; Ficken et al., 2000; Meyers and Ishiwatari, 1993), our results generally 347 reveal a differentiation between terrestrial taxa (i.e. xero-mesophilic group) characterized mainly by 348 *n*-alkanes maximizing at n-C₂₉ and n-C₃₁ and submerged living taxa (i.e. the meso-hygrophilic group) 349 maximizing at n-C₂₅ and n-C₂₇. Earlier investigations of Nott et al. (2000), Baas et al. (2000) and 350 Bingham et al. (2010), who compared the n-alkane fingerprints of Sphagnum taxa growing along a 351 hydrological gradient, agree with our results.

352 Huang et al. (2012a) and Ficken et al. (1998) used proxy ratios (ACL, P_{ag}) to divide moss taxa roughly 353 by their moisture preferences. They calculate the ACL₂₃₋₃₃ for samples of lichens and Racomitrum 354 lanuginosum, which have rather similar hydrological requirements to our xero-mesophilic mosses. 355 Despite the slightly narrower ACL₂₃₋₃₃ range, their results show similarities to our xero-mesophilic 356 group. A comparison between ACL₂₁₋₃₃, ratios of Sphagnum (plant position nearly at water-level) of 357 Huang et al. (2012b) and our brown mosses shows that the ACL₂₁₋₃₃ ratios of Sphagnum species are 358 rather lower. The intraspecific conclusion of Huang et al. (2012b) (wetter moisture conditions entail 359 lower ACL₂₁₋₃₃) for *Sphagnum* is reflected by our measurements.

$$dry (ACL_{21-33} = 29.1 - 27.5) < moisture condition > (ACL_{21-33} = 26.4 - 25.2) wet$$

$$Tom_nit < Aul_tur < Aul_pal < Dre_rev < Sco_sco$$

$$360$$

As we observed a clear difference in the ACL_{21-33} between the xero-mesophilic and the mesohygrophilic group, we suggest that the inclusion of mid-chain *n*-alkanes (*n*-C₂₁ to *n*-C₂₅) in the equation of ACL improves its value as a proxy for moisture conditions. Andersson et al. (2011) inferred ACL_{27-31} values of 29 for brown-moss peat from western Russian during wet phases, which is however, poorly comparable to our results because they investigated total peat organic matter instead of pure moss material.

Ficken et al. (2000) proposed P_{aq} as a semi-quantitative proxy ratio for the differentiation of terrestrial and aquatic plants (<0.1 terrestrial plants, 0.1–0.4 emergent macrophytes, 0.4-1 submerged/floating macrophytes). Our inferred P_{aq} results for the individual species agree with these assumptions. If we consider that the proxy ratio levels were created by vascular plants from a limited dataset of lakes in Kenya and as we focus on non-vascular plants of the arctic, we chose other level terms.

terms by Ficken et al. (2000)		emergent macrophytes					submerged/floating macrophytes			
adapted terms for mosses	2	xero-	mesophilic r	nosse	es	1	meso-hygroj	philic	e mosses	
species sorted by P_{aq}	Aul_tur	<	Tom_nit	<	Aul_pal	<	Dre_rev	<	Sco_sco	

Overall, our results do not support the inference of Nichols et al. (2006) that a hydrological classification is possible between *Sphagnum* and non-*Sphagnum* formed peat as the latter show wide variations between different habitats. The inferred broad P_{aq} range of *Drepanocladus* and between both *Aulacomnium* probably indicates that intraspecific variation is related to the individual's growing condition, which could provide the basis to develop P_{aq} as a proxy for water-level when measured on taxonomically identified fossil plant material.

373

As with P_{aq} , *n*-alkanes seem to be species-specific given stable environmental parameters and are related to the species-specific moisture requirements, which are adapted to changing environmental conditions. Thus, our results confirm the conclusions of Bingham et al. (2010), Bush and McInerney (2015) and Nott et al. (2000) that the pattern of *n*-alkanes has the potential to become a valuable proxy for chemotaxonomic identification and moisture conditions. *Scorpidium scorpioides*, a species with a rather narrow preference range (i.e. it is limited to open water conditions), shows low intraspecific variations. This matches the results for *Sphagnum* compiled by Bingham et al. (2010), which also show minor intraspecific variations. *Aulacomnium* in contrast, which grows in a rather wide range of moisture conditions, shows strong variations in its *n*-alkane spectra: whether this is a function of the individual's growing conditions, however, needs to be investigated in a more extensive study.

390 5. Conclusions

The habitat and intraspecific isotopic and chemical patterns of 10 brown-moss species detected along
 small-scale hydrological gradients in Siberian polygonal tundra were studied.

The observed higher $C/N_{(m)}$ ratios of xero-mesophilic mosses compared to those of the mesohygrophilic mosses originate from the different environmental requirements when living emergent (i.e. investment in a higher stability resulting in high C/N ratios) as opposed to submerged. Furthermore, the latter group may also gain a better nitrogen supply through microbial symbioses.

With respect to the isotopic source pools, the meso-hygrophilic species have greater access than xero-mesophilic species, which is seen in their large ranges. The approximate habitat-specific division of δ^{13} C values as a result of discrimination by RuBisCO under different hydrological regimes is overturned by the influence of different sources and cannot provide a clear distinction from a single measurement of either habitat type. For species, growing near the water level, no intraspecific relationship with water level was observed probably as a result of the parallel impact of processes causing opposing δ^{13} C trends.

Our analyses reveal that, compared with xero-mesophilic mosses, meso-hygrophilic mosses are characterized by enriched δ^{15} N values probably originating from microbial symbioses. Both carbon and nitrogen isotopic ratios seem to be valuable proxies to differentiate between taxa preferring the polygon rim or pond. Moreover, with respect to meso-hygrophilic mosses, the detected positive relations between intraspecific variations and the individuals' relative growing position could allow even more semi-quantitative information about water-level changes to be inferred. The *n*-alkane patterns of brown mosses (limited 16 individuals belonging to five species) indicate
that they are species-specific and have thus the potential to be developed as a chemotaxonomic proxy.
The applicability of proxy ratios (ACL and Paq) could be attested for arctic mosses after adjustments
of the levels.

414 Overall, our study indicates that C/N, isotopic and *n*-alkane analyses of brown moss material has 415 a high environmental indicator potential, particularly if species-specific material instead of bulk 416 material is analysed.

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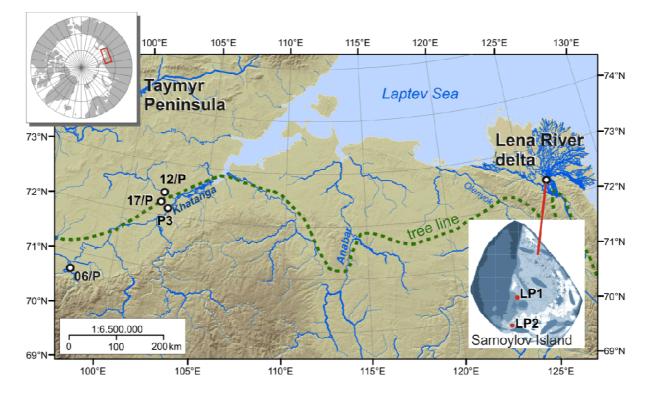
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658 Figures



659

660 Fig. 1: The Khatanga study sites are located in the south-east of the Taymyr Peninsula and both the sites on

661 Samoylov Island are in the southern Lena River delta (Map by Th. Böhmer).

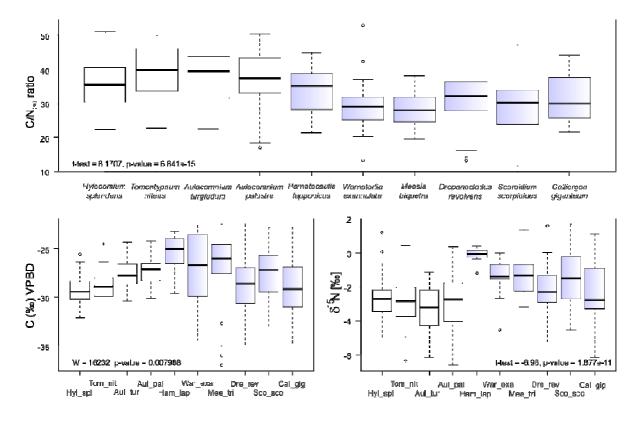
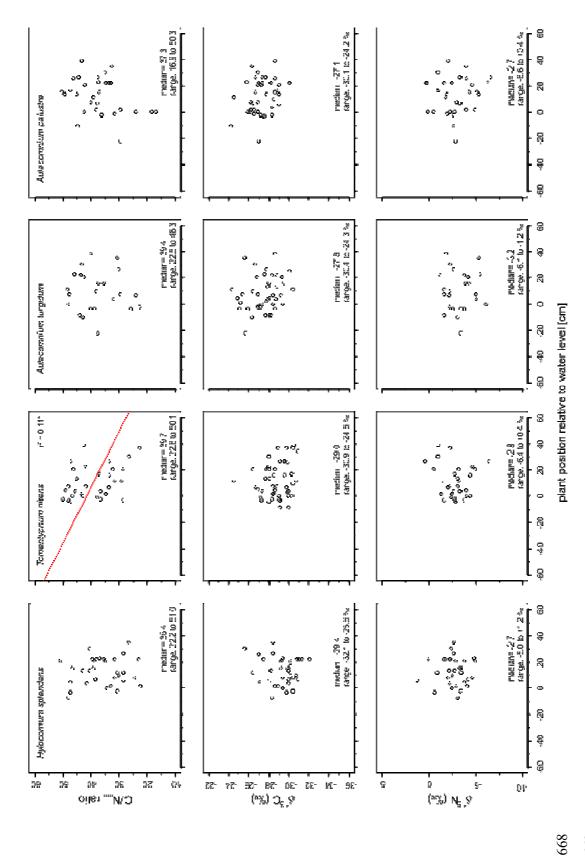
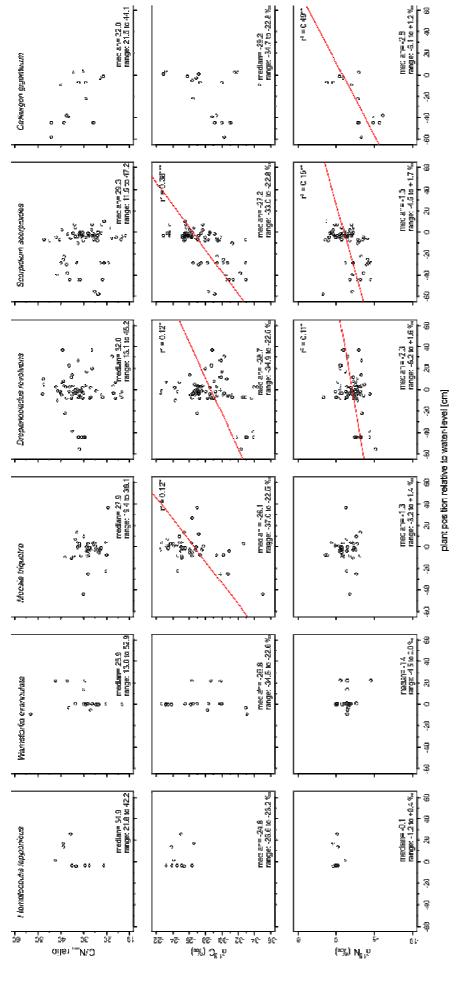


Fig. 2: C/N(m) ratios, $\delta^{l_3}C$ and $\delta^{l_5}N$ values of the 10 moss taxa studied. White boxplots are the xero-mesophilic

665 group and boxplots shaded in blue are the meso-hygrophilic group. A t-test was done to distinguish the signals 666 between the two habitat groups.









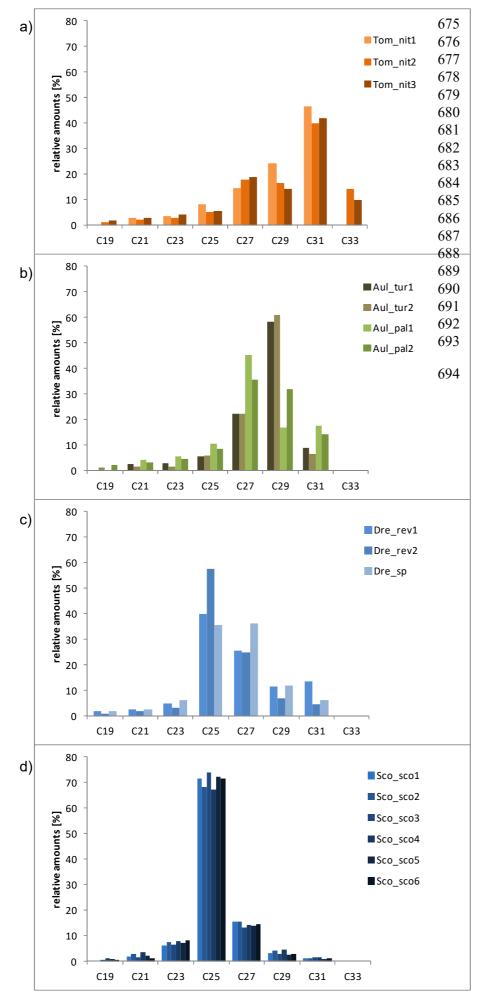


Fig. 4: The relative amounts [%] of n-alkanes with an odd number of carbon atoms in selected brown mosses grouped by species and ordered by their preferences with respect to their mean plant-position relative to water-level for xeromesophilic (a, b) and mesohygrophilic (c, d) mosses. Tom-nit: Tomentypnum nitens; Aul tur: Aulacomnium turgidum; Aul_pal: Aulacomnium palustre; Dre rev: Drepanocladus revolvens; Dre sp: Drepanocladus sp.; Sco sco: Scorpidium scorpioides

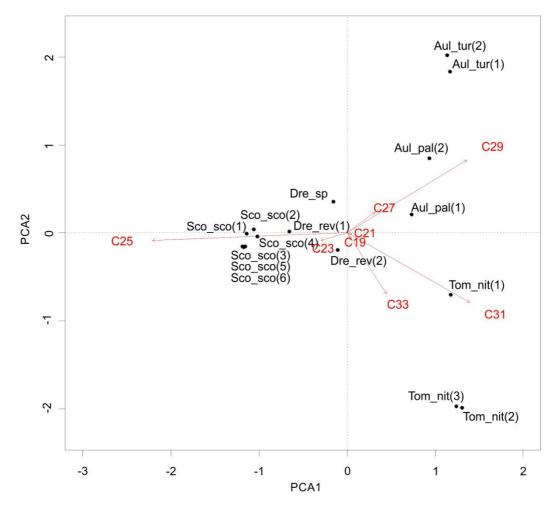


Fig 5: PCA of n-alkanes separates the meso-hygrophilic group on the left side of axis 1 from the xero-mesophilic
 group on the right side. Along the second axis the Aulacomniaceae are distinguished from Tomentypnum nitens

698 according to the distribution of long-chain n-alkanes with more or less than 30 carbon atoms.

699

700 Appendix

polygon cross section, (length of transect = polygon size) moss layer surface height permafrost table water level		coordinates	short description vegetation type; additiona information
G G V 0 2 4 6 8 10 12 14 16 18 20 22 24 26	06/P	70.666° N, 97.708° E	open forest; swinging bog (mat of mosses)
9 	17/P	72.243° N, 102.233° E	forest-tundra intersection; shallow and sparsely vegetated
9 0 2 4 6 8 10 12 14 16 18 20 22	P3/I		forest-tundra intersection; a complex of three individual polygons:
9 0 2 4 6 8 10 12 14 16 18 20 22	P3/II	72.149° N, 102.693° E	P3/I - deep, open water body P3/II - shallow, open wate body
2 7 0 2 4 6 8 10 12 14 16 18	P3/III		P3/III - shallow completel vegetated
0 2 4 6 8 10 12 14 16 18 20 22 24 26 28	12/P	72.431° N, 102.373° E	tundra; shallow and vegetated
2 0 2 4 6 8 10 12 14 16 18	LP1	72.375° N, 126.483° E	tundra; deep polygon without thaw depth below the water body
Q 	LP2	72.370° N, 126.481° E	tundra; shallow and vegetated

Table A1: Brief description of studied polygons. (For further information see Zibulski et al., 2016)

703 **Table A2:** $C/N_{(m)}$ ratio, $\delta^{3}C$ and $\delta^{5}N$ data of individual mosses depending on their position relative to the water-level (cm).

6	*	δ ¹³ C	[%6]	<u> </u>				
Species	median	range	n median		range	n	median	range
H. splendens (Hyl_spl)	47.1	29.5 to 67.9	34	-29.4	-32.1 to -25.6	31	-2.7	-5.0 to +1.2
T. nitens (Tom_nit)	52.9	30.1 to 66.6*	46	-29.0	-30.9 to -24.5	30	-2.8	-6.4 to +0.4
A. turgidum (Aul_tur)	52.4	29.9 to 64.3	41	-27.8	-30.4 to 24.3	24	-3.2	-6.1 to -1.2
A. palustre (Aul_pal)	49.6	22.5 to 66.9	40	-27.1	-30.1 to -24.2	30	-2.7	-6.6 to +0.4
H. lapponicus (Ham_lap)	46.5	28.3 to 56.1	10	-24.8	-26.6 to -23.2	9	-0.1	-1.2 to +0.4
W. exannulata (War_exa)	38.4	17.3 to 70.4	20	-26.8	-34.5 to -22.6	19	-1.4	-4.5 to 0.0
<i>M. triquetra</i> (Mee_tri)	37.1	25.8 to 50.7	45	-26.1	-37.0 to -22.5*	34	-1.3	-3.2 to +1.4
D. revolvens (Dre_rev)	42.6	17.5 to 64.1	72	-28.7	-34.9 to -22.5*	67	-2.3	-5.2 to +1.6**
S. scorpioides (Sco_sco)	38.9	15.4 to 62.8	69	-27.2	-33.0 to -22.8***	65	-1.5	-4.5 to +1.7**
C. giganteum (Cal_gig)	42.6	28.6 to 58.7	23	-29.2	-34.7 to -22.8	17	-2.9	-6.1 to +1.2**

Stars designate significant linear regressions between parameter and the plant position relative to water-level (* $p \le 0.01$, ** $p \le 0.05$, *** $p \le 0.001$).