



# Distinctly different bacterial communities in surface and oxygen minimum layers in the Arabian Sea

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# 12 Abstract

13 Contributions of microbial communities to biogeochemical processes in oxygen minimum oceanic zones are being realized through the applications of molecular techniques. To understand 14 seasonal and depth-wise variations in bacterial community structure (BCS) in the Arabian Sea 15 oxygen minimum region, extensive sampling and molecular analyses were carried out. 16S 16 rRNA gene sequencing was done to profile the BCS from five depths, surface (5m), deep 17 18 chorophyll maximum (43-50m, DCM), 250m, 500m and 1000m during Spring intermonsoon (SIM), Fall intermonsoon (FIM), and Northeast monsoon (NEM) seasons. Sequencing of 19 743 chimera-free clones revealed a clear vertical partitioning of BCS between the surface 20 (surface + DCM) and OMZ (250 + 500 + 1000m) layers. There was no distinct seasonal 21 difference in the BCS. Most 16S rRNA gene sequences were affiliated to Gammaproteobacteria 22 (39.31%), Alphaproteobacteria (23.56%) and Cyanobacteria (20.2%). Higher diversity and 23 OTUs in OMZ predominantly consisting of Alteromonodales, Sphinogomonadales, 24 Rhodobacterales, Burkholderales, and Acidimicrobiales we observed might be due to their 25 microaerophilic metabolism, ability to degrade recalcitrant substrates and assimilate sinking 26 particulate matter. Further hitherto undescribed diversity both in surface and OMZ layers was 27 28 evidenced. Implicit role of extant bacterial community in denitrification and anammox and in sulphur oxidation is highlighted. 29

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31 Keywords: Oxygen Minimum Zone, Bacterial diversity, 16S rRNA gene, Gammaproteobacteria,

32 Cyanobacteria, Arabian Sea





# 33 **1 Introduction**

34 Poor ventilation of intermediate layers, higher microbial respiration and organic matter oxidation within the water column (Wyrtki 1962) lead to oxygen minimum zones (OMZ) with 35 dissolved oxygen (DO) concentrations often below  $< 20\mu M L^{-1}$  (Lisa 2003). Major OMZs are 36 found in the intermediate depths of eastern tropical North Pacific (ETNP, Wyrtki 1966), eastern 37 tropical South Pacific (ETSP, Wyrtki 1966), eastern Arabian Sea (Wyrtki 1973; Madhupartap et 38 al. 1996; Naqvi and Jayakumar 2000) and eastern South Atlantic (Karstensen et al. 2008). Within 39 the OMZ, intense anaerobic and related processing of nitrogenous compounds (Stramma 2008) 40 lead to loss of fixed nitrogen to the atmosphere. Denitrification (Naqvi 1994) and anaerobic 41 oxidation of ammonia (anammox, Dalsgaard et al. 2012) are the major pathways for this nitrogen 42 loss. Hypoxic conditions often select resilient microbes and restrict their vertical distribution 43 (Wishner et al. 1995). 44

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The OMZs, earlier believed to occur mostly in nutrient rich upwelling regions, are 46 47 currently expanding and/or intensifying due to anthropogenic impacts (Diaz and Rosenberg 2008). The expansion of OMZs is affecting benthic ecosystems and marine fisheries due to 48 49 habitat alterations and/or changes in nutrient cycling (Stramma et al. 2008). Further, the OMZ expansion is ascribed to increased production of climate active trace gases including carbon 50 dioxide  $(CO_2)$ , methane  $(CH_4)$  and nitrous oxide  $(N_2O)$ . Microbes are involved in all such 51 processes but little is known about their community structure and metabolism within OMZs, in 52 53 particular from the Arabian Sea-OMZ (AS-OMZ).

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Coinciding with an active denitrification zone (Naqvi 1994; Naqvi et al. 1998) the AS-55 OMZ is among the largest anoxic regions in the oceans with DO levels  $< 20\mu$ M, sometimes 56 dipping as low as 0.1 µM (Naqvi 2006; Paulmier and Ruiz-Pino 2009). This alone accounts for 57 20% of the oceanic denitrification (Codispoti et al. 2001) and contributes to 40% of the global 58 pelagic dinitrogen (N<sub>2</sub>) production (Naqvi et al. 2008). Previous studies mostly focused on 59 delineating the role of denitrifying and anammox bacteria in the loss of fixed nitrogen in the AS-60 OMZ (Jayakumar et al. 2009; Bulow et al. 2010; Pitcher et al. 2011; Newell et al. 2011; Bouskill 61 et al. 2012). However, detailed attempts to understand bacterial community structure and/or 62 bacterial diversity of the AS-OMZ and its overlying surface waters are a few and far in between 63





64 (Fuchs et al. 2005; Jain et al. 2014). The present study thus aimed to delineate phylogenetic 65 diversity of the overall bacterial communities in the AS-OMZ, as well as the diversity of bacterial phylotypes contributing to temporally stable bacterial community structure (BCS) in the 66 OMZ (vide Jain et al. 2014). For this, we analyzed small subunit ribosomal RNA gene (16S 67 rRNA or SSU rRNA) clone libraries prepared from water samples collected from five depths 68 from the Arabian Sea Time Series station (ASTS; 17°0.126'N, 67°59.772'E), during three 69 different seasons. ASTS is akin to well known HOTS (Hawaii Ocean Time-Series) in the Pacific 70 Ocean and BATS (Bermuda Atlantic Time-Series) in the Altantic Ocean. 71

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# 73 **2 Materials and methods**

# 74 2.1 Sampling

Water samples were collected under the SIBER program in May 2012 (Spring 75 76 intermonsoon [SIM]), September 2012 (Fall intermonosson [FIM]), and February 2013 (Northeast monsoon [NEM]), as described previously by Jain et al (2014). In brief, samples for 77 DNA extractions were collected using pre-cleaned Niskin bottles on a CTD rosette. Samples 78 from surface (0.5m), deep chlorophyll maxima (DCM) (~35-50m), intense denitrification zone 79 80 (250m, 500m) and deep denitrification zone (1000m) at the ASTS station (17°0.126' N, 67°59.772'E) were strained through 200µm pore sized bolting silk, immediately after collection, 81 2.5 L of seawater sample from each depth was filtered peristatically through a Sterivex cartridge 82 fitted with 0.22 µm pore size membrane filter (Millipore, USA). The Sterivex cartridge was then 83 84 filled with 1.8 ml of lysis buffer (50 mM 131 Tris pH 8.3, 40 mM EDTA and 0.75 M sucrose), sealed, and stored frozen at -80°C until nucleic acid extraction was performed in the laboratory. 85 86

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# 87 2.2 DNA extraction and PCR amplification of 16s rRNA genes

<sup>88</sup> DNA was extracted from Sterivex filters using using the modified method of Ferrari & <sup>89</sup> Hollibaugh (1999). The precipitated DNA was hydrated in 50  $\mu$ l sterile deionised water. All <sup>90</sup> DNA samples were subjected to PCR amplification using universal 16S rRNA primers 27F and <sup>91</sup> 1492R (to confirm that the DNA was of PCR quality. The 16S rRNA gene was amplified <sup>92</sup> following conditions mentioned by Sambrook (1989). PCR amplification was performed in a <sup>93</sup> final volume of 50  $\mu$ l in a thermocycler (Applied Biosystems, USA) and correct amplification <sup>94</sup> was ensured by checking for the amplicons electrophoretically.





#### 95 2.3 Clone library construction and DNA sequencing

96 PCR amplified 16S rRNA gene products were purified using Axyprep-96 PCR Clean up kit (Axygen, Biosciences), cloned into pCR4-TOPO vector using a TOPO-TA cloning kit for 97 sequencing (Invitrogen, USA) and transformed by chemical transformation into TOP-10 cells as 98 per manufacturer's instructions. As the transformation efficiency was low to moderate, at least 99 thrice cloning trials were repeated to collect a minimum of 65 clones for further analyses. All 100 positive clones/transformants from each sample were picked out, grown overnight at 37°C on LB 101 plates and subjected to the colony-PCR with primers sets pucM13F/pucM13R using temperature 102 conditions as per manufacturer's instructions. PCR products were purified with the Axyprep-96 103 PCR Clean up kit (Axygen, Biosciences) and then sequenced using an ABI 3130XL genetic 104 analyzer (Applied Biosystems, USA). Clone libraries were constructed from particular depths 105 keeping DO profile in mind. However, for the purpose of comparison, clone libraries from 106 surface and DCM depth were considered as surface group (SIM, FIM and NEM). Similarly, 107 108 clone libraries from 250, 500 and 1000m were considered as OMZ group for each season

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#### 110 **2.4 Phylogenetic analysis of clone sequences**

The sequences were assembled into contigs using DNA Baser sequence assembly 111 software version 2 (DNA baser, USA). Vector contamination was removed from the sequences 112 using the VecScreen tool (http:// www. ncbi. nlm. nih. gov /tools/vecscreen/). Only consensus 113 sequences without vector and primer residues and with a quality score of 20 (which translates 114 115 into more than 99.5% correct bases, Allex 1999), were used for further analyses. Chimeric 116 sequences were excluded using а chimera detection program (http://decipher.cee.wisc.edu/FindChimeras.html). It is to be noted that close to 15% of the total 117 sequences were chimeric. Phylogenetic affiliation of proper sequences was determined using the 118 119 naïve Bayesian classifier of the RDP sequence classifier tool (Wang et al. 2007) using the 16S rRNA Training Set 26 2015; 120 14 (updated released on May https://rdp.cme.msu.edu/classifier/classifier.jsp). The classification of sequences was done using 121 1,000 pseudo-bootstrap replications at a bootstrap value of 80%, which results in a standard error 122 123 of only 1.3%. The sequences were also compared with other databases including SILVA, NCBI and Greengenes, using MOTHUR. The sequences were assigned to a phylum if their identity was 124 > 90% in the different databases searched. The sequences obtained and described in this study 125





were submitted to the NCBI GenBank database and are available under accession numbers
KJ589647 to KJ590044, KR269603 to KR269693, KR919859 to KR920002, KR919859 to
KR920002 and KR673365 to KR819266.

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# 130 2.5 Determination of OTUs

131 16S rRNA gene sequences were aligned using ClustalX and the alignment was curated 132 using Gblocks (Castresana 2000) to remove poorly aligned positions and divergent regions. 133 Distance matrices were created from the curated alignments with Phylip 3.66. The sequences 134 were then assigned into phylotypes (operational taxonomic units, OTUs) using MOTHUR by 135 applying the average neighbor rule (Schloss and Westcott 2011). 97% cut-off for sequence 136 similarity was used to delimit an OTU.

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## 138 **2.6 Estimation of shared and unique OTUs**

In order to elucidate seasonal and depth wise differences in bacterial phylotypes, the fraction of shared and unique OTUs was estimated using MOTHUR. Further, the numbers of shared and unique OTUs between all three seasons surface samples (SIM-surface, FIM-surface, NEM-surface) and all three seasons OMZ samples (SIM-OMZ, FIM-OMZ, NEM-OMZ) were also estimated.

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#### 145 **2.7 Construction of Phylogenetic tree**

A phylogenetic tree of representative shared OTUs from the clone libraries was constructed to visualize their relationship and affiliations with the closest relative sequences from the database (Kemble et al. 2011). The tree was constructed as per MEGA-6 using maximum composite likelihood as substitution model and bootstrap values were calculated using neighborjoining method with a resampling size of n = 500.

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# 152 **2.8 Statistical analysis of clone libraries**

The clone libraries from the ASTS were analysed statistically using RDP Library Compare (http://rdp.cme.msu.edu/wiki/index.php/Lib\_Compare) and ∫ LIBSHUFF analyses. The Ribosomal Database Project (RDP) II library compare tool uses the naive Bayesian rRNA Classifier Version 1.0 at an 80% confidence threshold (Wang et al. 2007; Cole et al. 2007) to





157 calculate the difference between two libraries. In brief, the RDP LibCompare provides P values 158 for determining statistical significance of abundance differences for individual taxa instead of estimating overall difference between samples. This tool uses the RDP Classifier to assign 159 sequences to taxa. Depending on the abundance of sequences assigned to each taxon, one of two 160 statistical tests is used to compute a P value to determine if a taxon is differentially represented 161 in the libraries. The f LIBSHUFF program was applied to compare bacterial community structure 162 between libraries in a phylogenetic context. It measures differences between communities based 163 on differences between sequences using Monte Carlo test (Schloss et al. 2004; Schloss 2008). ∫ 164 LIBSHUFF implements the Cramer-von Mises statistic to test the generic hypothesis that two 165 communities are the same (Singleton et al. 2001; Schloss et al. 2004). All statistical analyses 166 were performed using MOTHUR software, except RDP. 167

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# 169 **2.9 Diversity and richness estimation**

The indices for diversity (Simpson's and Shannon's indices), richness estimates (Jackknife, Chao 1 and ACE), rarefaction and collectors curve were performed using MOTHUR (Schloss et al. 2009). Total richness of the clone libraries was extrapolated from the observed number of OTUs using the three nonparametric richness estimators.

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# 175 **3 Results**

# 176 **3.1 Physico-chemical characteristics of the sampling site**

177 Physico-chemical characteristics (i.e. temperature, salinity, dissolved oxygen, pH, nitrite, nitrate, ammonium, silicate, phosphate and total organic carbon) at the ASTS are described 178 earlier in Jain et al. (2014) and vertical distribution of dissolved oxygen (DO), nitrite (NO<sub>2</sub>) and 179 180 nitrate (NO<sub>3</sub>) during three seasons is shown in Figure 1. In general, surface and DCM depths (43-50 m) are well oxygenated followed by a steep oxycline between DCM and 250 m. The 181 average DO concentration ranged from 185.24  $\pm$  31.1  $\mu$ M L<sup>-1</sup>at DCM to 5.56  $\pm$  5.5  $\mu$ M L<sup>-1</sup> at 182 250m. The DO concentrations were slightly more at 1000 m. The nitrite concentration was 0.29 183 in the upper thermocline region (50m) with higher oxygen concentrations (during FIM) and 2.5 184  $\mu M L^{-1}$  in the intermediate depths (250m) with low oxygen (during NEM). Surface waters 185 186 during SIM and FIM were devoid of nitrate and it was quite low during the NEM.





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# 189 **3.2 Bacterial Community Structure**

Seasonal distribution pattern of clone sequences from surface and OMZ (Figure 2) 190 indicate that maximum numbers of clone sequences were affiliated to three major phylogenetic 191 groups. These are Gammaproteobacteria (39.31%), Alphaproteobacteria (23.56%) and 192 Cyanobacteria (20.2%). These groups represented up to 82% of the usable sequences. The 193 relative proportions of sequences in these groups varied temporally, as well as spatially, with 194 depth. The percentages of Gamma- and Alphaproteobacteria were much higher in the OMZ than 195 in the surface layers. During NEM the percentages of Cyanobacteria and Gammaproteobacteria 196 were the highest at the surface and in the DCM. Notably, the percentage of 197 Gammaproteobacteria within OMZ (250, 500 and 1000m) did not vary much between seasons. 198 The distribution of Alphaproteobacteria in OMZ was quite similar during FIM and NEM. The 199 highest proportion of unclassifiable bacterial sequences in the surface layers, as well as OMZ 200 201 depths, was observed during SIM, followed by FIM and NEM.

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# 203 3.3 Shared and Unique OTUs

The shared and unique OTUs between all three seasons in the surface and OMZ samples 204 were divided into four categories: (1) OTUs common to all seasons, (2) OTUs common to two 205 seasons, (3) season-specific OTUs having >1 sequence, and (4) season-specific OTUs having 206 only one sequence (singleton). Three OTUs common to all season's within surface (Common to 207 208 all season's surface-OTUs, CTASS-OTU) was represented by a minor fraction (11-17 %) of the sequences in individual seasons (Figure 3). Further, 20 OTUs (Common to 2 seasons surface, 209 CT2SS-OTU) were shared within surface during two of the three seasons and represented by 19 210 to 38 % of the sequences in individual seasons. The proportion of season-specific OTUs 211 212 (containing only one sequence) in the surface was the largest (43-73%) and represented by 19-57% of sequences in individual season. 213

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A total of eight OTUs were common within OMZ (Common to all seasons OMZ, CTASO-OTU) during three seasons and represented by 14-37% of individual season sequences. Further, 20 OTUs (Common to 2 seasons OMZ, CT2SO-OTU) shared within OMZ during two of the three seasons represented 15-20 % of the individual season sequences. The proportion of





season-specific OTUs in the OMZ was the largest (50-63%) and represented by 28-46% of the
individual sequences from any of the seasons. More than half (63%) of season-specific OTUs in
the OMZ are represented by single sequences.

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# 223 **3.4 Phylogenetic affiliation and phylogenetic tree of shared OTUs**

The relative abundance and phylogenetic affiliation of the OTUs shared by all seasons 224 and any of the two seasons in surface and OMZ are listed in **Table 1(a)** and **Table 1(b)**. To study 225 the evolutionary differences between shared OTUs, representative sequences were subjected to 226 phylogenetic analysis and the relationship with the database sequences was depicted using a 227 phylogenetic tree for surface (Figure 4) and OMZ (Figure 5). All CTASS-OTUs (CTASS-OTU-228 1, 2 and 3) represented 13% of all surface sequences and were related to the Cyanobacteria, 229 more precisely to the order Synechococcales. On the phylogenetic tree CTASS-OTU-1 and 230 CTASS-OTU-3 appears on the same branch while CTASS-OTU-2 branches independently from 231 232 them and their closest relatives are sequences from Red Sea. The OTUs shared between two of the three seasons (CT2SS-OTUs) were mainly affiliated with the Cyanobacteria, Proteobacteria 233 and Acidobacteria. SIM-Surface and FIM-Surface shared OTUs were related to 234 Gammaproteobacteria (CT2SS-OTU-4, 9, 11 and 13), Betaproteobacteria (CT2SS-OTU-10), 235 Alphaproteobacteria (CT2SS -OTU-14 and 15), Cyanobacteria (CT2SS-OTU-8 and 16) and 236 Acidobacteria (CT2SS-OTU-20). FIM-Surface and NEM-Surface shared ten OTUs, of which six 237 were affiliated to Cyanobacteria (CT2SS-OTU-3, 5, 12, 17 and 18), two to 238 239 Gammaproteobacteria (CT2SS-OTU-2 and 21) and one each to Alphaproteobacteria (SS-OTU-1) and Betaproteobacteria (CT2SS-OTU-19). SIM-Surface and NEM-Surface shared only one 240 OTU affiliated to Gammaproteobacteria (CT2SS-OTU-6). 241

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Of the eight CTASO-OTUs, the two most abundant ones, namely CTASO-OTU-1 and CTASO-OTU-2, represented 16% of all OMZ sequences and were related to *Gammaproteobacteria*, in the Order *Alteromonadales*. CTASO-OTU-1 and -2 branched independently from each other and their closest relatives are sequences from Arctic Ocean, Ishigaki Jima Island (Japan), South China Sea and Red Sea. The next most abundant CTASO-OTU-3 and CTASO-OTU-4 represented 5% of all OMZ sequences and belong to the Order *Sphingomonadales* of Class *Alphaproteobacteria*. Both these aligned on the same branch in the





tree and are related to sequences from South Atlantic Ocean and South Pacific gyre. The CTASO-OTU-5, CTASO-OTU-6, CTASO-OTU-7 and CTASO-OTU-8, represented 1% of the total OMZ sequences. They are affiliated with *Burkholderilaes*, *Acidimicrobiales*, uncultured *Gammaproteobacteria*, and *Rhodobacterales*, respectively. The closest relatives of CTASO-OTU-5, 6, 7, and 8 are reported from Sri Lanka thermal spring, Indian Ocean, Red Sea, and Submarine basalt loihi Seamount, respectively.

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The OTUs shared in the OMZ between two seasons were mainly from three different 257 Phyla, including Proteobacteria, Actinobacteria and marine group A (MGA). SIM-OMZ and 258 FIM-OMZ were found to share OTUs related to Gammaproteobacteria (CT2SO-OTU-2, 3, 10, 259 11, 12, and 15), MGA (CT2SO-OTU-12 and 15) and Alphaproteobacteria (SO-OTU-11). The 260 FIM-OMZ and NEM-OMZ shared four OTUs. Of those, three were affiliated to 261 Alphaproteobacteria (CT2SO-OTU-1, 14, and 16) and one to Actinobacteria (CT2SO-OTU-20). 262 The SIM-OMZ and NEM-OMZ shared 10 OTUs. Four of them were affiliated with to 263 Alphaproteobacteria (CT2SO-OTU-6, 7, 8, and 18), another four to Gammaproteobacteria 264 (CT2SO-OTU-9, 13, 17, and 20), and one each to MGA (CT2SO-OTU-4) and Actinobacteria 265 (CT2SO-OTU-5). 266

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# 268 **3.5 Comparison of clone libraries**

Using the RDP classifier, it was seen that at Class level, the surface samples during all 269 270 three seasons had more representatives of Cyanobacteria than in the OMZ samples. This was entirely different in the OMZ where Alphaproteobacteria, Gammaproteobacteria and 271 Marinimicrobia were the dominant groups (Table 2). Further, from the RDP libcompare it was 272 clear that the abundance of Gammaproteobacteria, (Altermonodales), Alphaproteobacteria 273 274 (Spingomondales), and SAR11 differed significantly between seasons in the surface clone libraries. In contrast, the abundance of Cyanobacteria and Alphaproteobacteria was significantly 275 lower and different in the OMZ (Table 3). The LIBSHUFF (p < 0.0016) analysis, used to note 276 that bacterial community, was significantly different between surface and OMZ during all three 277 278 seasons. The bacterial community in surface layers differed significantly between seasons (LIBSHUFF, p < 0.0001), whereas no such significant seasonal difference was found in the 279 bacterial community extant in the OMZ (Table 4). 280





#### 281 **3.6 Analysis of clone diversity and richness**

282 Shannon and Simpson indices, as well as rarefaction curves, clearly indicate that at an evolutionary distance of 3% the bacterial diversity in both surface and OMZ was the highest 283 during SIM followed by NEM and FIM (Table 5). However, during each season, higher 284 diversity indices were evident in the OMZ than in the surface layers. The nonparametric 285 Jackknife, Chao 1 and ACE estimators also revealed that the estimated OTUs are much higher in 286 the OMZ than the observed number of OTUs during all the seasons. Rarefaction (Figure 6a) and 287 collector curves (Figure 6b) implied that no saturation was reached either at sequence or OTU 288 level at an evolutionary distance of 1% and 3%, respectively. However, at Class or Family level, 289 at distances of 10%, the rarefaction implied some saturation. 290

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# 292 **4 Discussion:**

Traditionally the OMZs were seen as regions dominated by heterotrophic denitrification 293 fueled by sinking of organic matter produced via photosynthesis in the sunlit surface ocean. They 294 were also considered to process a fundamentally different microbial community and operate on a 295 different biogeochemistry. The discovery Anammox, and active but cryptic sulfur cycle in 296 anoxic OMZs have significantly shifted the old paradigms. For almost a decade now the OMZ 297 gene surveys have focused extensively on microbes performing denitrification and anaerobic 298 ammonia oxidation (anammox) with less emphasis on the overall microbial community. 299 Microbial communities within OMZs play central roles in ocean, yet we still lack a fundamental 300 301 understanding of how microbial biodiversity is distributed across the OMZs. In this regard, our 302 efforts are useful in providing some novel insights. Arabian Sea is modulated seasonally by upwelling, winter cooling (Prasanna Kumar et al. 2001) and semi-annual reversal of monsoonal 303 winds (Madhupratap et al. 1996). This greatly influences primary production, organic carbon 304 concentration and flux (Hansell and Peltzer 1998) and bacterial abundance (Ramaiah et al. 1996, 305 306 2000; Jain et al. 2014). In spite of the global biogeochemical and climatic importance of AS-OMZ, spatio-temporal variation (Riemann et al. 1999; Fuchs et al. 2005; Jain et al. 2014) and 307 phylogenetic diversity of the microorganisms inhabiting therein are only sparsely addressed 308 (Riemann et al. 1999; Jayakumar et al. 2009; Divya et al. 2010, 2011). Oxygen deficient waters 309 in the intermediate depths is a perennial feature of the Northeastern Arabian Sea which gets 310 intensified during winter or NEM season due to poor water circulation and high surface 311





productivity (Naqvi et al. 1990; Prasanna Kumar and Prasad 1996). Lower DO level and higher nitrite concentration at the intermediate depth during NEM season signifies intense denitrification processes and the presence of nitrite in the thermocline region signifies the nitrification process (Sen Gupta et al. 1976). During intermonsoon (SIM and FIM) periods, the intense solar heating and weak winds stratify the Northeastern Arabian Sea surface layer, leading to depletion of nitrate in the upper euphotic zone (Muraleedharan and Prasanna Kumar 1996).

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The SIM is a transition period from winter to summer and is generally known to be a 319 period of low primary productivity caused by water column stratification and oligotrophic 320 conditions (Madhupratap et al. 1996). During this period, the bacterial community is mainly 321 sustained by the slow-to-degrade dissolved organic carbon (DOC) from earlier phytoplankton 322 blooms of the NEM (Ramaiah et al. 2000). Despite these environmental conditions we observed 323 higher bacterial diversity during SIM (both in surface as well as OMZ). Most microbial habitats 324 are spatially heterogeneous (Kassen and Rainey 2004) and can contain a large number of 325 potential niches. Previous studies reported a positive correlation between habitat heterogeneity 326 ("patchiness") and the phylogenetic diversity of bacteria (Korona et al. 1994; Rainey et al. 2000; 327 Kassen et al. 2000; Zhou et al. 2002). We find that the bacterial diversity at the ASTS varies 328 seasonally in the surface layers and in OMZ and is at its highest during SIM. High rates of 329 denitrification might be among the many possible causes for this higher diversity. Further, the 330 high proportion of unclassified bacteria may also contribute to the increased diversity to some 331 332 extent as observed during SIM.

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Though over 60 clones were finally available from each depth for sequencing, they do 334 not capture the full bacterial diversity in the AS-OMZ. This notwithstanding, our efforts are 335 336 useful to suggest that overall bacterial diversity at Phylum and OTU level in the ASTS is higher in OMZ than in the surface layers. Interestingly, the depth-dependent variation of OMZ bacterial 337 diversity is not consistent among different studies (Ganesh et al. 2014). Further, Bryant et al 338 (2012) reported a consistent decline in the bacterial diversity within the ETSP-OMZ. Similarly, 339 Zaikova et al (2010) observed low diversity associated with the seasonal OMZ off British 340 Columbia. In contrast, Stevens and Ulloa (2008), based on 16S rRNA clone libraries identified 341 higher OTU diversity at the ETSP-OMZ, a pattern consistent with the observation in our study. 342





343 Similarly, Brown et al (2009) and Kemble et al (2011) showed elevated OTU richness to 344 coincide with the zone of minimum oxygen at the HOTS. Bryant et al (2012) and Jain et al (2014) attributed lower diversity in the OMZ to competition for limited resources, environmental 345 filtering, and lower redox potential and less readily available organic matter. On the other hand, 346 higher diversity in the OMZ has been linked to the use of a wider range of terminal oxidants 347 348 compared to the oxic depths where oxygen is the dominant electron acceptor (Stevens and Ulloa 2008). Most clades of *Gammaproteobacteria* are known to denitrify using nitrate as electron 349 acceptor (Miller et al. 2010). Among them, the predominant Alteromonodales in the AS-OMZ 350 appears to be a proficient denitrifier group. 351

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353 Bacterial community structure at the ASTS is dominated by phylotypes affiliated with three major phylogenetic groups viz., Gammaproteobacteria, Alphaproteobacteria and 354 Cyanobacteria. The dominance of Alphaproteobacteria and Gammaproteobacteria, has been 355 reported earlier from a variety of pelagic marine environments (Giovannoni and Rappé 2000), 356 including the OMZ of the ETSP (Stevens and Ulloa 2008; Ganesh et al. 2014) and the Southern 357 Arabian Sea (Fuchs et al. 2005). Our results suggest seasonal variation in the abundance of 358 certain major bacterial groups mainly in the surface layers than in the OMZ depths. As Hansell 359 & Peltzer (1998) suggested, the organic carbon concentration and changes in primary 360 productivity patterns in the surface layers may be responsible for variation in the abundance of 361 Gammaproteobacteria (Altermonodales) and Alphaproteobacteria (Spingomondales). Higher 362 363 abundance of Gammaproteobacteria (Altermonodales) and Alphaproteobacteria (Spingomondales) and SAR11 in the OMZ libraries might be supported by microaerophilic 364 metabolism, particle associated lifestyle and/or their ability to use abundantly available nitrate as 365 their terminal electron acceptor for energy generation. 366

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The total number of common OTUs shared within surface clone libraries was significantly lower than that in the OMZ. All common surface OTUs constitute only a minor fraction (13 %) of the total surface sequences and were mostly affiliated with *Cyanobacteria*, more precisely to the genus *Synechococcus* and *Prochlorococcus*. Interestingly, majority of surface OTUs (shared between two of the three clone libraries) were also affiliated with *Cyanobacteria*, specifically to *Prochlorococcus*. *Cyanobacteria* of the genera *Prochlorococcus* 





374 and *Synechococcus* have been found inhabiting the upper lit part of OMZs (Johnson et al. 1999; 375 Goericke et al. 2000; Ulloa et al. 2006; Galán et al. 2009). However, phylotypes of Synechococcus are often restricted to the upper part of the euphotic zone and do not show a clear 376 vertical partitioning with depth as Prochlorococcus does. Moreover Prochlorococcus is 377 dominant in the oligotrophic waters (Olson et al. 1990; Li 1995; Liu et al. 1997a; Grob et al. 378 2007). The dominance of *Prochlorococcus* over *Synechococcus* in the OMZ is presumably a 379 result of specific physiological adaptation to the prevailing light and nutrient conditions within 380 the region. Even though few Prochlorococcus ecotypes are able to utilize oxidized forms of 381 nitrogen, particularly nitrite (Moore et al. 2002), the biological processes of nitrate and nitrite 382 reduction carried out by the local OMZ microbial community (Farías et al. 2007) could support 383 their dominant presence in the AS-OMZ upper stratum. 384

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It is apparent that the higher number of common OTUs affiliated with Alteromonodales 386 387 (16% of the total OMZ sequences) and Sphinogomonadales (5% of the total OMZ sequences) might contribute to the temporally stable BCS in the OMZ. However, it should be noted that 388 proportions of phylotypes affiliated with Alteromonodales and Sphinogomonadales also varied 389 seasonally in the surface layers. Alteromonas macleodii clustered into two major genotypic 390 groups or ecotypes, one found in the upper water column and another in the deep water column. 391 Martinez et al (2008) reported that the deep ecotype is better suited to microaerophilic conditions 392 and for the degradation of recalcitrant compounds and colonizes relatively large particles that 393 394 sink rapidly to meso and bathypelagic depths. On the other hand, the surface ecotype, a typical rstrategist (investing most energy in multiplying fast) has more potential for regulation and 395 degradation of sugars and amino acids, and specializes in colonizing smaller particulate organic 396 matter with much slower sinking rates (Perez et al. 2012). Therefore, it is possible that the 397 398 persistence of Alteromonodales throughout the year in the AS-OMZ could be due to its microaerophilic metabolism and due to its ability to degrade recalcitrant compounds as well as to 399 be able to degrade large sinking particulate matter, for eg. transparent exopolymeric particles 400 (Dileep kumar et al. 1998 and Ramaiah et al. 2005). Alteromonadales, particularly in the genera 401 402 Idiomarina and Alteromonas are known to metabolize both labile and semi-labile high molecular weight dissolved organic matter (McCarren et al. 2010). Presence of Sphingomonadales in the 403 upper OMZ has been reported earlier by Riemann et al (1999), and although their ecological 404





significance in OMZ is not clear, *Sphingomonadales* has been documented to have wide
metabolic capabilities (Miller et al. 2010) and can degrade aromatic compounds (Fredrickson et
al. 1995). Thus, their abundance, especially of *Novosphingobium* and *Sphingobium*, in the ASOMZ may be driven by high concentrations of phosphate and ammonia as postulated by Liu et al
(2015).

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The SAR11 group is known to limit the flux of sulfur from the ocean by demethylating 411 dimethylsulfoniopropionate (DMSP), which is usually degraded to volatile dimethyl sulfide 412 (Moran et al. 2003; Howard et al. 2006). Demethylated DMSP can serve as a substrate in sulfur 413 oxidation in the deeper layer. This group is also known to assimilate labile amino acids, thereby 414 playing a significant role in C, N, and S cycling (Malmstrom et al. 2004). The other common 415 OMZ-OTUs were those Burkholderilaes, Acidimicrobiales, uncultured Gammaproteobacteria, 416 and *Rhodobacterales* seems to contribute to the stable BCS in the AS-OMZ. It is interesting to 417 note that only a few of the common OMZ OTUs affiliated to known denitrifiers, suggesting that 418 denitrification is spatially heterogenous (niche segregation). Many of the other shared OMZ 419 OTUs (between any two seasons) and season specific OMZ OTUs were affiliated with 420 *Nitratireducter*.sp, a known denitrifying bacterium. 421

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One of the most significant findings in OMZ biogeochemistry and microbiology has been 423 the recognition of a cryptic pelagic sulphur cycle. Traditionally, it was believed that sulphur 424 425 reduction would not initiate until the oxygen and nitrate/nitrite is fully consumed (Froelich et al. 1979). Surprisingly, recent metagenomic studies have discovered an abundant and diverse sulfur-426 oxidizing microbial community in the OMZ layers. This community is particularly enriched in 427 Deltaproteobacteria related to sulfur-oxidizing symbionts of deep sea bivalves (Canfield et al. 428 429 2010). Radiolabeled sulfate tracer experiments from the OMZ off the Chilean coast revealed significant sulfate reduction in the upper reaches of the OMZ water (Canfield et al. 2010). 430 Sulphur oxidising bacterial clades, such as Thiotrichales and Desulfobacteriales, were present in 431 the libraries made from 250m and 500m, although represented by only a few sequences. These 432 433 are the very first indications of the existence of sulphur metabolizing bacterial community in the AS-OMZ. 434





436 Based on diversity and richness indicators, the estimated total numbers of OTUs for all 437 samples from the AS-OMZ were much higher than the observed number of OTUs, indicating that additional sampling would have revealed greater diversity. The same trend was observed 438 when rarefaction and collector curves were calculated as no stationarity was reached at levels 439 below Phylum approximated by an evolutionary distance of 10%. Together, the richness 440 estimators ACE and Chao 1, and the three different diversity indices indicate that there is a high 441 degree of undescribed diversity in both surface and OMZ, a large fraction of which likely belong 442 to clusters with sequences whose taxonomic resolution is finer than that defined for species. 443

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Results of this study demonstrate that the OMZ bacterial community is diverse yet 445 distinctively different communities predominate in the surface and OMZ layers. Many of the 446 groups that have adapted to low oxygen are likely to have important roles in C, N and S cycling 447 within the OMZs. Organisms contributing to high species-richness may act as regulators of 448 449 fluxes of C, N and S into and out of the OMZ. Their rates of cycling and metabolism in the OMZ seem to be stable. While seasonal variation in the OMZ bacterial community is minimal, the 450 diversity in the OMZ depths is greater than that in the surface layers, with deep-branching groups 451 and many novel and as yet uncultivated clades. Also, the sulphur oxidation appears to be 452 operated by certain microbes in OMZs exemplied by sparse yet obvious presence of 453 Thiotrichales and Desulfobacteriales in the AS-OMZ. Prevalence of denitrification, anammox 454 and sulphur oxidation brings forth the role of microbes in the AS-OMZ ecosystem functioning. 455

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# 462 **Conflict of interest.** None declared

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468	Re	ferences:
469	1	
470	1.	Allex, C.F.: Computational methods for fast and accurate DNA fragment assembly,
471		Department of Computer Sciences, University of wisconsin-Madison, UW Technical
472		Report CS-1R, 1999.
473	2	Bouskill N Eveillard D Chien D Javakumar A and Ward B B : Environmental factors
475	2.	determining ammonia oxidizing organism distribution and diversity in marine environments.
476		Environ Microbiol, 14, 714–729, doi: 10.1111/j.1462-2920.2011.02623.x, 2012.
477		, , , , , , , , , , , , , , , , , , ,
478	3.	Bulow, S., Rich, J., Naik, H., Pratihary, A.K. and Ward B.B.: Denitrification exceeds
479		anammox as a nitrogen loss pathway in the Arabian Sea oxygen minimum zone, Deep- Sea
480		Res, 57, 384–393, doi:10.1016/j.dsr.2009.10.014, 2010.
481		
482	4.	Brown, M., Philip, G., Bunge, J., Smith, M.C., Bissett, A., Lauro, F.M., Fuhrman, J.A. and
483		Donachie, S.P.: Microbial community structure in the North Pacific Ocean, ISME J, 3,
484		1374–1386, doi:10.1038/ismej.2009.86, 2009.
485		
486	5.	Bryant, J., Stewart, F., Eppley, J. and Delong E.: Microbial community phylogenetic and
487		trait diversity declines with depth in a marine oxygen minimum zone, Ecology, 93,
488		1659–1673, doi: 10.1890/11-1204.1, 2012.
489		
490	6.	Canfield, D., Stewart, F., Thamdrup, B., De Brabandere, L., Dalsgaard, T., Delong,
491		E.F., Revsbech, N.P. and Ulloa, O.: A cryptic sulfur cycle in oxygen-minimum-zone waters
492		off the Chilean coast, Science, 330, 1375–1378, 2010.
493	-	
494	7.	Castresana, J.: Selection of conserved blocks from multiple alignments for their use in
495		phylogenetic analysis, Molecular Biology and Evolution, 17, 540-552, 2000.
496	0	
497	8.	Codispoti, L., Brandes, J., Christensen, J., Devol, A.H., Naqvi, S.W.A., Paeri, H.W. and Veshinori. T. The eccentric fixed nitrogen and nitrous oxide hydratic maying tenants as we
498		onter the anthronocene? Scientia Marine 65, 85, 105, 2001
499 500		enter the anthropocene? Scientia Marina, 05, 85–105, 2001.
501	0	Cole I Chai B Farrie P Wang O Kulam Sved Mohidean A S McCarrell D.M.
502	9.	Bandela A M Cardenas E Garrity G M and Tiedie I M : The ribosomal database
502		project (RDP-II): introducing my RDP space and quality controlled public data Nucleic
503 504		Acids Research 35 169–172 doi: 10.1093/nar/gkl889.2007
505		10105 10500101, 55, 16, 172, 001. 10110,5/100/GR100, 2007.
506	10	Dalsgaard, T., Thamdrup, B., Farias, L. and Revsbech, N.: Anammox and denitrification in
507		the oxygen minimum zone of the eastern South Pacific, Limnol. Oceanogr, 57, 1331-1346,
508		doi: 10.4319/lo.2012.57.5.1331,2012.
509		
510	11	Divya, B., Soumya, K. and Nair, S.: 16S rRNA and enzymatic diversity of culturable
511		bacteria from the sediments of oxygen minimum zone in the Arabian Sea, Antonie van
512		Leeuwenhoek, 98, 9-18, doi: 10.1007/s10482-010-9423-7, 2010.
513		





514

515

516 517

13. Diaz, R.J. and Rosenberg, R.: Spreading dead zones and consequences for marine 518 ecosystems, Science, 321, 926–929, doi: 10.1126/science.1156401, 2008. 519 520 14. Farías, L., Paulmier, A. and Gallegos, M.: Nitrous oxide and N-nutrient cycling in the 521 oxygen minimum zone off northern Chile, Deep Sea Res Part I, 54, 164-180, 522 doi:10.1016/j.dsr.2006.11.003, 2007. 523 524 15. Ferrari, V.C. and Hollibaugh, J.T.: Distribution of microbial assemblages in the Central 525 Arctic Ocean Basin studied by PCR/DGGE: analysis of a large data set, Hydrobiologia, 401, 526 55-68, 1999. 527 528 16. Fuchs, B.M., Woebken, D., Zubkov, M.V., Burkill, P. and Amann, R.: Molecular 529 identification of picoplankton populations in contrasting waters of the Arabian Sea, Aquat 530 Microb Ecol, 39, 145–157, 2005. 531 532 17. Fredrickson, J.K., Balkwill, D.L., Drake, G.R., Romine, M.F., Ringelberg, D.B. and White, 533 534 D.C.: Aromatic-degrading Sphingomonas isolates from the deep subsurface, Appl Environ. Microbiol, 61, 1917–1922, 1995. 535 536 18. Froelich, P.N., Klinkhammer, G.P., Bender, M.L., Luedtke, N.A., Heath G.R., Cullen D., 537 Hammond P.D.D. and Hartman, B.: Early oxidation of organic matter in pelagic sediments 538 of the eastern equatorial Atlantic: suboxic diagenesis, Geochimica et Cosmochimica Acta, 539 43, 1075-1090, doi:10.1016/0016-7037(79)90095-4, 1979. 540 541 19. Ganesh, S., Parris, D.J., DeLong, E.F. and Stewart, F.J.: Metagenomic analysis of size 542 fractionated picoplankton in a marine oxygen minimum zone, The ISME Journal, 8, 187-543 211, doi:10.1038/ismej.2013.144, 2014. 544 545 20. Galán, A., Molina, V., Thamdrup, B., Woebken D., Lavik, G., Kuypers, M.M.M. and Ulloa, 546 547 O.: Anammox bacteria and anaerobic oxidation of ammonium in the oxygen minimum zone off northern Chile, Deep Sea Res Part II, 56, 1021-1031, doi:10.1016/j.dsr2.2008.09.016, 548 2009. 549 550 21. Giovannoni, S.J. and Rappe, M.: Evolution, diversity and molecular ecology of marine 551 prokaryotes, Microbial ecology of the oceans, 47-48, 2000. 552 553 22. Goericke, R., Olson, R.J. and Shalapyonok, A.: A novel niche for Prochlorococcus sp in 554 low-light suboxic environments in the Arabian Sea and the Eastern Tropical North Pacific, 555 Deep Sea Res Part I, 47, 1183-1205, doi:10.1016/S0967-0637(99)00108-9, 2000. 556 557 17

12. Divya, B., Parvathi, A., Loka Bharathi, P.A. and Nair, S.: 16S rRNA-based bacterial

Arabian Sea, World J Microbiol Biotechnol, 27, 2821–2833, 2011.

diversity in the organic-rich sediments underlying oxygen deficient waters of eastern





561

565

569

573

- Grob, C., Ulloa, O., Li, W.K.W., Alarcón, G., Fukasawa, M. and Watanabe, S.: Picoplankton
  abundance and biomass across the eastern South Pacific Ocean along latitude 32.5 degrees
  S. Mar Ecol Prog Ser, 332, 53–62, 2007.
- 562 24. Hansell, D.A. and Peltzer, E.T.: Spatial and temporal variations of total organic carbon in
  563 the Arabian Sea, Deep- Sea Res II, 45, 2171–2193, doi:10.1016/S0967-0645(98)00067-8,
  564 1998.
- 566 25. Howard, E.C., Henriksen, J.R., Buchan, A., Reisch, C.R., Bürgmann, H., Welsh, R., Ye,
  567 W., González, J.M., Mace, K., Joye, S.B., Kiene, R.P., Whitman, W.B. and Moran, M.A.:
  568 Bacterial taxa that limit sulfur flux from the ocean, Science, 314, 649–652, 2006.
- 570 26. Jayakumar, A., O'Mullan, G.D., Naqvi, S.W.A. and Ward, B.B.: Denitrifying bacterial
  571 community composition changes associated with stages of denitrification in oxygen
  572 minimum zones, Microb Ecol, 58, 350–362, doi: 10.1007/s00248-009-9487-y, 2009.
- Jain A., Bandekar, M., Gomes, J., Shenoy, D., Meena, R.M., Naik, H., Khandeparkar, R.
  and Ramaiah, N.: Temporally invariable bacterial community structure in the Arabian Sea oxygen minimum zone, Aquatic Microbial Ecology, 73, 51–67, doi:10.3354/ame01704, 2014.
- 578
- 28. Johnson, Z., Landry, M.L., Bidigare, R.R., Brown, S.L., Campbell, L., Gunderson, J., Marra,
  J. and Trees, C.: Energetics and growth kinetics of a deep *Prochlorococcus spp* Population
  in the Arabian Sea, Deep Sea Res Part II, 46, 1719–1743, doi:10.1016/S0967-0645(99)00041-7, 1999.
- 583
- Karstensen, J., Stramma, L. and Visbeck, M.: Oxygen minimum zones in the eastern
  tropical Atlantic and Pacific oceans, Progress in Oceanography, 77, 331–350,
  doi:10.1016/j.pocean.2007.05.009, 2008.
- 30. Kassen, R. and Rainey, P.: The ecology and genetics of microbial diversity, Annual Rev
   Microbiol, 58, 207-231, doi: 10.1146/annurev.micro.58.030603.123654, 2004.
- 590

593

- 591 31. Kassen, R., Buckling, A., Bell, G. and Rainey, P. B.: Diversity peaks at intermediate
   592 productivity in a laboratory microcosm, Nature, 406, 508-512, 2000.
- 32. Korona, R., Nakatsu, C., Forney, L. and Lenski, R.: Evidence for multiple adaptive peaks
  from populations of bacteria evolving in a structured habitat. Proc. Natl. Acad. Sci., 91,
  9037-9041, 1994.
- 597
- 33. Kembel, S.W., Eisen, J.A., Pollard, K.S. and Green, J.L.: The Phylogenetic Diversity of
   Metagenomes, PLoS ONE, 6, 23-214, doi:10.1371/journal.pone.0023214, 2011.
- 600





- 34. Kumar, M.D., Sarma, V.V.S.S., Ramaiah, N., Gauns, M., and de Sousa, S.N.:
  Biogeochemical significance of transport exopolymer particles in the Indian Ocean,
  Geophys. Res. Lett, 25, 81-84, doi: 10.1029/97GL03481, 1998.
- 35. Lisa, A.L.: Oxygen minimum zone benthos: adaptation and community response to hypoxia,
   Oceanography and Marine Biology, 41, 1–4, 2003.
- 36. Li, W.K.W.: Composition of ultra phytoplankton in the Central North-Atlantic, Mar Ecol
   Prog Ser, 122, 1–8, 1995.
- 610

604

607

- 37. Liu, H.B., Nolla, H.A. and Campbell, L.: *Prochlorococcus* growth rate and contribution to
  primary production in the equatorial and subtropical North Pacific Ocean, Aquat Microb
  Ecol, 12, 39–47, 1997a.
- 614
- 38. Liu, J., Fu, B., Yang, H., Zhao, M., He, B. and Zhang, X.H.: Phylogenetic shifts of
  bacterioplankton community composition along the Pearl Estuary: the potential impact of
  hypoxia and nutrients, Front Microbiol, 6,64-75, doi: 10.3389/fmicb.2015.00064, 2015.
- 618
- 39. Madhupratap, M., Prasanna, K.S., Bhattathiri, P.M.A., Kumar, D.M., Raghukumar, S., Nair,
  K. K. C. and Ramaiah, N.: Mechanism of the biological response to winter cooling in the
  north eastern Arabian Sea, Nature, 384, 549–552, doi:10.1038/384549a0, 1996.
- 622
- 40. Malmstrom, R.R., Kiene, R.P., Cottrell, M.T. and Kirchman, D.L.: Contribution of SAR11
  bacteria to dissolved dimethylsulfoniopropionate and amino acid uptake in the North
  Atlantic Ocean, Appl. Environ. Microbiol, 70, 4129–4135, 2004a.
- 626
- 41. Martinez, E.I., Martin, A.B., D'Auria, G., Mira, A., Ferriera, S., Johnson, J., Friedman, R.
  and Rodriguez, V.F.: Comparative genomics of two ecotypes of the marine planktonic
  copiotroph *Alteromonas macleodii* suggests alternative lifestyles associated with different
  kinds of particulate organic matter, The ISME Journal, 2, 1194–1212, doi:
  10.1038/ismej.2008.74, 2008.
- 42. McCarren, J., Becker, J.W., Repeta, D.J., Shi Y., Young, C.R., Malmstrom, R.R., Chisholm,
  S.W. and DeLong E.F.: Microbial community transcriptomes reveal microbes and metabolic
  pathways associated with dissolved organic matter turnover in the sea, Proc Natl Acad Sci,
  107, 16420–16427, doi: 10.1073/pnas.1010732107, 2010.
- 637

- 43. Miller, T.R., Delcher, A.L., Salzberg, S.L., Saunders, E., Detter, J.C. and Halden, R.U.:
  Genome sequence of the dioxin-mineralizing bacterium *Sphingomonas wittichii* RW1, J.
  Bacteriol, 192, 6101–6102, doi: 10.1128/JB.01030-10, 2010.
- 641 642
- 44. Moran, M.A., Gonzalez, J.M. and Kiene, R.P.: Linking a bacterial taxon to sulfur cycling in
  the sea: studies of the marine *Roseobacter* group, Geomicrobiol J, 20, 375–388, doi:
  10.1080/01490450303901, 2003.





646 647 648	45. Moore, L.R., Post, A.F., Rocap, G. and Chisholm, S.W.: Utilization of different nitrogen sources by the marine cyanobacteria <i>Prochlorococcus</i> and <i>Synechococcus</i> , Limnol Oceanogr 47, 989–996 doi: 10.4319/lo.2002.474.0989.2002
640	Geomogi, 17, 707 770, doi: 10.1517/10.2002.171.10707, 2002.
650 651	46. Muraleedharan, P.M. and Prasanna, K.S.: Arabian Sea upwelling—a comparison between coastal and open ocean regions, Curr Sci, 71, 842–846, 1996.
652	
653	47. Naqvi, S.W.A., Noronha, R., Somasundar, K. and SenGupta, R.: Seasonal changes in the depitrification regime of the Arabian Sea, Deep Sea Res L 37, 593–611, doi:10.1016/0108
655 656	0149(90)90092-A, 1990.
657 658 659	48. Naqvi, S.W.A.: Denitrification processes in the Arabian Sea, Proceedings of Indian Academy of Sciences, 103, 279-300, 1994.
660 661 662 663	49. Naqvi, S.W.A., Yoshinari, T., Jayakumar, D.A., Altabet, M.A., Narvekar, P.V., Devol, A.H., Brandes, J.A. and Codispoti L.A.: Budgetary and biogeochemical implication of N <sub>2</sub> O isotope signatures in the Arabian Sea, Nature, 394, 462–464, doi: 10.1038/28828, 1998.
664 665 666	50. Naqvi, S.W.A. and Jayakumar, D.A.: Ocean biogeochemistry and atmospheric composition: significance of the Arabian Sea, Curr Sci, 78, 289–299, 2000.
667 668 669 670	51. Naqvi, S.W.A., Naik, H., Pratihary, A.K., D'Souza, W., Narvekar, P. V., Jayakumar, D. A., Devol, A. H., Yoshinari, T. and Saino, T.: Coastal versus open-ocen denitrification in the Arabian Sea, Biogeosciences, 33, 621-633, doi:10.5194/bg-3-621-2006, 2006.
671 672 673	52. Naqvi, S.W.A., Voss, M. and Montoya, J.P.: Recent advances in the biogeochemistry of nitrogen in the ocean, Biogeosciences, 54, 1033-1041, 2008.
674 675 676 677	53. Newell, S.E., Babbin, A.R., Jayakumar, A. and Ward, B.B.: Ammonia oxidation rates and nitrification in the Arabian Sea, Global Biogeochem Cycles, 25, 621-629, doi: 10.1029/2010gb003940, 2011
678 679 680 681	54. Olson, R.J., Chisholm, S.W., Zettler, E.R., Altabet, M.A. and Dusenberry J.A.: Spatial and temporal distributions of prochlorophyte picoplankton in the North-Atlantic Ocean, Deep Sea Res Part I, 37, 1033–1051, doi:10.1016/0198-0149(90)90109-9, 1990.
682 683 684 685	55. Pitcher, A., Villanueva, L., Hopmans, E.C., Schouten, S., Reichart, G.J. and Sinninghe D.J.S.: Niche segregation of ammoniaoxidizing archaea and anammox bacteria in the Arabian Sea oxygen minimum zone. ISME J, 5, 1896–1904, doi: 10.1038/ismej.2011.60, 2011.
687 688 689	56. Paulmier, A. and Ruiz-Pino, D.: Oxygen minimum zones OMZs in the modern ocean, Progress in oceanography, 80, 113–128, doi:10.1016/j.pocean.2008.08.001, 2009.





690 691 692 693 694	57.	Pe'rez, M.L., Gonzaga, A., Martin, C.A.B., Onyshchenko, O., Ghavidel, A., Ghai, R. and Rodriguez-Valera, F.: Genomes of surface isolates of Alteromonas macleodii: the life of a widespread marine opportunistic copiotroph, Scientific Reports Nature, 2, 696-707, doi:10.1038/srep00696, 2012.
695 696 697	58.	Prasanna, K.S. and Prasad, T.G.: Winter cooling in the northern Arabian Sea, Curr Sci, 71, 834–841, 1996;.
698 699 700 701 702	59.	Prasanna, K.S., Ramaiah, N., Gauns, M., Sarma, V.V.S.S., Muraleedharan, P.M., Raghukumar, S., Kumar D.M. and Madhupratap, M.: Physical forcing of biological productivity in the Northern Arabian Sea during the northeast monsoon, Deep-Sea Res II, 48, 1115–1126, doi:10.1016/S0967-0645(00)00133-8, 2001.
703 704 705 706	60.	Ramaiah, N., Sarma, V.V.S.S., Gauns, M., Kumar, D.M. and Madhupratap, M.: Abundance and relationship of bacteria with transparent exopolymer particles during the 1996 summer monsoon in the Arabian Sea, Proc. Indian Acad. Sci. Earth Planet. Sci, 109, 443-451, 2000.
707 708 709	61.	Ramaiah, N., Raghukumar, S. and Gauns, M.: Bacterial abundance and production in the central and eastern Arabian Sea, Curr Sci, 71, 878-882, 1996.
710 711 712 713	62.	Ramaiah, N., Raghukumar, S., Mangesh, G. and Madhupratap, M.: Seasonal variations in carbon biomass of bacteria, thraustochytrids and microzooplankton in the Northern Arabian Sea, Deep-Sea Res II, 52, 1910-1921, doi:10.1016/j.dsr2.2005.05.004, 2005.
714 715 716 717	63.	Rainey, P., Buckling, A., Kassen, R. and Travisano, M.: The emergence and maintenance of diversity: insights from experimental bacterial populations, Trends Ecol. Evol, 15, 243-247, 2000.
718 719 720 721 722	64.	Riemann, L., Steward, G.F., Fandino, L.B., Campbell, L., Landry, M.R. and Azam, F.: Bacterial community composition during two consecutive NE Monsoon periods in the Arabian Sea studied by denaturing gradient gel electrophoresis (DGGE of rRNA genes), Deep-Sea Res II, 46, 1791–1811, doi:10.1016/S0967-0645(99)00044-2, 1999.
723 724 725	65.	Sambrook, J., Fritsch, F.F. and Maniatis, T.: Molecular cloning: a laboratory manual (2 <sup>nd</sup> edn., Cold Spring Harbor, N.Y., Cold Spring Harbor laboratory), 1989.
726 727 728	66.	Schloss, P.D., Larget, B.R. and Handelsman, J.: Integration of mi-crobial ecology and statistics: a test to compare gene libraries, Appl. Environ. Microbiol, 70, 5485–5492, 2004.
729 730 731	67.	Schloss, P.D.: Evaluating different approaches that test whether microbial communities have the same structure, ISME J, 2, 265–275, doi: 10.1038/ismej.2008.5, 2008.
732 733 734	68.	Schloss, P.D., Westcott, S.L., Ryabin, T., Hall J.R., Hartmann, M., Hollister, E.B., Lesniewski, R.A., Oakley, B.B., Parks, D.H., Robinson, C.J., Sahl, J.W., Stres, B., Thallinger, G.G., Horn, D,V. and Weber C.F.: Introducing mothur: open-source, platform-





735 736 737	independent, community- supported software for describing and comparing microbial communities, Appl Environ Microbiol, 75, 7537–7541, doi:10.1128/AEM.01541-09, 2009.
738 739 740 741	69. Schloss, P.D. and Westcott, S.L.: Assessing and Improving Methods Used in Operational Taxonomic Unit-Based Approaches for 16S rRNA Gene Sequence Analysis, Applied and Environmental microbiology, <b>77</b> , 3219-3226, doi: 10.1128/AEM.02810-10, 2011.
742 743 744 745	70. Singleton, D.R., Furlong, M.A., Rathburn, S.L. and Whitman, W.B.: Quantitative comparisons of 16SrRNA gene sequence libraries from environmental samples, Appl Environ Microbiol, 67, 4374-4376, doi: 10.1128/AEM.67.9.4374-4376.2001, 2001.
746 747 748 749	71. Stevens, H. and Ulloa, O.: Bacterial diversity in the oxygen minimum zone of the eastern tropical South Pacific, Environ Microbiol, 10, 1244–1259, doi: 10.1111/j.1462-2920.2007.01539.x, 2008.
750 751 752	72. Stramma, L., Johnson, G.C., Sprintall, J. and Mohrholz, V.: Expanding Oxygen-Minimum Zones in the Tropical Oceans, Science, 320, 655-658, doi: 10.1126/science.1153847, 2008.
753 754 755 756 757 758	73. Ulloa, O., Belmar, L., Farías, L., Castro-González, M., Galán, A., Lavín, P., Molina, V., Ramírez, S., Santibáñez, F. and Stevens, H.: Microbial communities and their biogeochemical role in the water column of the oxygen minimum zone in the eastern South Pacific, Gayana (Concepción), 70, 83–86, doi.org/10.4067/S0717-65382006000300018, 2006.
759 760 761 762	74. Wang, Q., Garrity, G.M., Tiedje, J.M. and Cole. Naïve Bayesian Classifier for Rapid Assignment of rRNA Sequences into the New Bacterial Taxonomy, Appl Environ Microbiol, 73, 5261–5267, doi:10.1128/AEM.00062-07, 2007.
763 764 765 766 767	75. Wishner, K.F., Ashjian, C.J., Gelfman, C., Gowing, M.M., Kann, L., Levin L.A., Mullineaux L.S. and Saltzman J.: Pelagic and benthic ecology of the lower interface of the eastern tropical Pacific oxygen minimum zone, Deep Sea Res Part I Oceanogr Res Pap, 42, 93–115, doi:10.1016/0967-0637(94)00021-J, 1995.
768 769 770	76. Wyrtki, K.: The oxygen minima in relation to ocean circulation, Deep-Sea Res, 9, 11–23, doi:10.1016/0011-7471(62)90243-7, 1962.
771 772 773	77. Wyrtki, K.: Oceanography of the eastern Pacific Ocean, Oceanogr Mar Biol Annu Rev, 4, 33–68, 1966.
774 775 776	<ol> <li>Wyrtki, K.: Physical oceanography of the Indian Ocean, The biology of the Indian Ocean, 8, 18–36, 1973.</li> </ol>
777 778 779	79. Zaikova, E., Walsh, D.A., Stilwell, C.P., Mohn, W.W., Tortell, P.D. and Hallam, S.J.: Microbial community dynamics in a seasonally anoxic fjord: Saanich Inlet British Columbia, Environ Microbiol, 12, 172–191, doi: 10.1111/j.1462-2920.2009.02058.x, 2010.





780	
781	80. Zhou, J., Xia, B., Treves, D.S., Wu L.Y., Marsh T.L. O'Neill, R.V. Palumbo A.V. and
782	Tiedje J.M.: Spatial and resource factors influencing high microbial diversity in soil, Appl. Environ Microbial 68, 326 334, doi: 10.1128/AEM.68.1.326.334.2002, 2002
783 784	Environ. Microbiol, 08, 520-554, doi: 10.1128/AEM.08.1.520-554.2002, 2002.
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Table 1 (a). Percentage of OTUs from Arabian Sea Time Series (ASTS) location being common to all three seasons and in two







Table 1 (b). Percentage of OTUs from Arabian Sea Time Series (ASTS) location being common to all three seasons and in two

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815 seasons in OM	IZ dur	ing dif	ferent s	eason and tl	he phylogenetic affiliation of the repre	esentative sequence	e from each OTU.	
OMZ								
Sampling season	SIM	FIM	NEM	Total seqs	Phylogenetic affiliation	Accession number	Environment	% similarity
OTUS from OMZ common to :	all seas	suo						
CTASO-OTU-01 (KJ589883)	7.8	15.4	12.2	46	Uncultured Alteromonas sp.	AB262378.1	Ishigaki jima Island, Japan	%66
CTASO-OTU-02 (KR673368)	0.7	<i>L.T</i>	5.3	17	Alteromonas sp.	JX533666.1	South china sea	100%
CTASO-OTU-03 (KJ589777)	0.7	2.6	3.8	6	Erythrobacter sp.	AM990645.1	South Pacific Gyre	96%
CTASO-OTU-04 ( KR269635)	1.3	3.4	2.3	6	Uncultured bacterium clone	KM110216.1	South Atlantic Ocean	%66
CTASO-OTU-05 (KR919998)	0.7	2.6	0.8	5	Uncultured Burkholderia sp.	KF305507.1	Sri Lanka thermal spring	100%
CTASO-OTU-06 (KJ589826)	0.7	0.9	1.5	4	Uncultured bacterium clone	JX441387.1	Indian Ocean	100%
CTASO-OTU-07 (KJ589739)	1.3	0.9	0.8	4	Uncultured Proteobacterium clone	KM018890.1	Red Sea	%66
CTASO-OTU-08 (KR269655)	0.7	0.9	0.8	3	Methylarcula sp.	DQ412076.1	Submarine basalt loihi Seamount	%66
OTUs from OMZ found in any	v two se	casons						
CT2SO-OTU-01 (KR919865)	0.0	0.9	3.1	5	Sphingobium sp.	KF672731.1	Haihe estuary	100%
CT2SO-OTU-02 (KJ589884)	1.3	2.6	0.0	5	Uncultured gamma proteobacterium clone	JN232986.1	Continental slope Cape Lookout	98%
CT2SO-OTU-03 (KJ589779)	2.6	0.9	0.0	5	Uncultured marine microorganism clone	JN166297.1	Subtropical North Pacific	96%
CT2SO-OTU-04 (KJ589746)	2.0	0.0	0.8	4	Uncultured bacterium clone, Marine Group A,	НQ674579.1	Northeast subarctic pacific Ocean	98%
CT2SO-OTU-05 (KJ589789)	1.3	0.0	0.8	3	Euzebya sp.	KP735966.1	Pacific Ocean	96%
CT2SO-OTU-06 (KR819220)	1.3	0.0	0.8	3	Nitratireductor sp.	EU440986.1	Southwest Indian Ridge	97%
CT2SO-OTU-07 (KJ589735)	1.3	0.0	0.8	3	Uncultured marine microorganism clone	JN166353.1	Subtropical North Pacific	98%
CT2SO-OTU-08 (KJ589873)	1.3	0.0	0.8	3	Uncultured bacterium clone	JN165748.1	Subtropical North Pacific	
CT2SO-OTU-09 (KR673384)	0.7	0.0	0.8	2	Uncultured bacterium clone	JX391520.1	Marine sediments	98%
CT2SO-OTU-10 (KR919866)	0.7	0.9	0.0	2	Oleibacter marinus strain 201,	NR_112787.1	Tropical ocean	%66
CT2SO-OTU-11 (KR269632)	0.7	0.9	0.0	2	Uncultured alpha proteobacterium clone	GQ337156.1	Deep Arctic Ocean	%66
CT2SO-OTU-12 (KR269629)	0.7	0.9	0.0	2	Uncultured bacterium clone	HQ674365.1	Pacific Ocean	%66
CT2SO-OTU-13 (KJ589749)	0.7	0.0	0.8	2	Uncultured bacterium clone	JX441447.1	Indian Ocean	95%
CT2SO-OTU-14 (KR919982)	0.0	0.9	0.8	2	Alterierythrobacter sp.	EU440971.1	Southwest Indian Ridge	98%





CT2SO-OTU-15 (KJ589784)	0.7	0.9	0.0	5	Uncultured bacterium clone, Marine Group A	HQ673196.1	Northeast subarctic pacifi Ocean	%66
CT2SO-OTU-16 (KR819230)	0.0	0.9	0.8	2	Uncultured bacterium clone	KF146540.1	Yellow Sea	98%
CT2SO-OTU-17 (KR673387)	0.7	0.0	0.8	2	Uncultured gamma proteobacterium clone	KP076569.1	Coastal Arabian Sea	97%
CT2SO-OTU-18 (KJ589880)	0.7	0.0	0.8	2	Uncultured bacterium clone	M110216.1	South Altantic Ocean	98%
CT2SO-OTU-19 (KJ589799)	0.7	0.0	0.8	2	Uncultured bacterium clone	AY375071.1	Deep sea sediments pacifi ocean	97%
CT2SO-OTU-20 (KR919989)	0.0	0.9	0.8	5	Uncultured actinobacterium	FJ615153.1	Saanich Inlet OMZ	%66
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Table 2. RDP library comparison results for the groups of significantly differing (p <0.01) sequences between surface and OMZ clone

libraries and their phylogenetic affiliations according to RDP library compare (Cole et al 2007). The more abundant group is shown in 843

bold. 844

Spring Inter-monsoon (SIM)	Surface	OMZ	p-value
Phylum Cyanobacteria	41	6	6.00E-14
class Cyanobacteria	34	8	1.05E-11
Phylum <i>Proteobacteria</i>	39	120	4.23E-06
class Gammaproteobacteria	16	72	4.13E-05
Fall Inter-monsoon (FIM)			
Phylum Cyanobacteria	99	0	1.50E-16
class Cyanobacteria	56	0	3.82E-14
Phylum <i>Proteobacteria</i>	87	105	1.06E-09
class Alphaproteobacteria	28	33	3.85E-02
order Sphingomonadales	0	6	3.89E-04
class Gammaproteobacteria	40	62	2.60E-06
order Alteromonadales	32	40	9.32E-03
Northeast Monsoon (NEM)			
Phylum Cyanobacteria	31	2	1.10E-10
class Cyanobacteria	29	1	6.10E-11
Phylum Proteobacteria	55	109	3.40E-04
class Alphaproteobacteria	15	46	2.98E-03









855	Table 3. RDP library comparison results for the groups of significantly differing (p $<\!0.01)$
856	sequences within surface and OMZ clone libraries obtained during different seasons and their
857	phylogenetic affiliation according to RDP library compare (Cole et al 2007). The more abundant

858 group is shown in bold.

RDP classifier taxonomic rank for sequences			p-value
	SIM-surface	<b>FIM-surface</b>	
Phylum Proteobacteria	39	87	1.71E-1
class Gammaproteobacteria	16	40	2.50E-1
order Alteromonadales	5	32	4.77E-3
	<b>FIM-surface</b>	NEM-surface	
Phylum Proteobacteria	87	55	3.03E-1
class Alphaproteobacteria	28	15	8.65E-3
order Sphingomonadales	0	5	4.38E-3
family SAR11	16	0	1.01E-3
class Gammaproteobacteria	40	39	2.78E-3
-	NEM-surface	SIM-surface	
Phylum Proteobacteria	55	39	3.49E-2
class Gammaproteobacteria	39	16	4.00E-4
order Alteromonadales	29	5	3.97E-5
	SIM-OMZ	FIM-OMZ	
Phylum Cyanobacteria	9	0	8.05E-3
•	<b>FIM-OMZ</b>	NEM-OMZ	
NA	NA	NA	NA
	NEM-OMZ	SIM-OMZ	
Class Alphaproteobacteria	46	33	6.34E-3

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- **Table 4.** J-LIBSHUFF comparisons of clone libraries constructed from the surface and OMZ
- 871 depths during spring intermonsoon (SIM), fall intermonsoon (FIM), and northeast monsoon
- 872 (NEM).

		Surface			OMZ		
		SIM-S	FIM-S	NEM-S	SIM-O	FIM-O	NEM-O
	SIM-s	0	0.2783	0.0525	0.0902	<0.0001	0.0001
Surface	FIM-s	<0.0001	0	<0.0001	<0.0001	<0.0001	<0.0001
	NEM-s	<0.0001	<0.0001	0	0.0006	<0.0001	<0.0001
	SIM-0	<0.0001	<0.0001	<0.0001	0	0.0036	0.0022
OMZ	FIM-o	<0.0001	<0.0001	<0.0001	0.0262	0	0.0352
	NEM-o	<0.0001	<0.0001	<0.0001	0.0615	0.0071	0

With an experiment wise error rate of 0.05 and taking into account a Bonferroni's correction for multiple comparisons, the libraries were considered significantly different when either of the two P values generated for an individual pairwise comparison was lower than 0.0016 (significant values are marked in bold).

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- Table 5. Comparison of bacterial diversity in clone libraries constructed from the surface layers
  and OMZ during spring intermonsoon (SIM), fall intermonsoon (FIM), and northeast monsoon
- 890 (NEM) at the Arabian Sea Time Series (ASTS) location.

Surface SIM 84 FIM 158 NEM 93 OMZ	uences	<b>OTUs</b> 60 59 58	Shannon           3.86           3.57           3.82	<b>Simpson</b> 0.02 0.04 0.02	<b>Chao</b> 405.50 106.75	ACE 1350.94	Jackknife	
Surface SIM 84 FIM 158 NEM 93 OMZ	3	60 59 58	3.86 3.57 3.82	0.02 0.04 0.02	405.50 106.75	1350.94	322.88	
SIM         84           FIM         158           NEM         93           OMZ	3 ) 7	60 59 58	3.86 3.57 3.82	0.02 0.04 0.02	405.50 106.75	1350.94	322.88	
FIM 158 NEM 93 OMZ	3 ) 7	59 58 119	3.57 3.82	0.04	106.75			
NEM 93 OMZ	) 7	58 119	3.82	0.02		150.83	114.32	
OMZ	) 7	110		0.02	158.33	310.95	152.12	
	) 7	110						
SIM 160	7	11)	4.63	0.01	378.67	1052.11	374.61	
FIM 117		68	3.79	0.04	162.23	171.30	169.81	
NEM 131		92	4.21	0.02	425.67	1004.50	958.52	







Figure 1 Vertical distribution of dissolved oxygen (DO), nitrate (NO<sub>3</sub>) and nitrite (NO<sub>2</sub>) concentrations during spring intermonsoon (SIM), fall intermonsoon (FIM), and northeast monsoon (NEM) at the Arabian Sea Time Series (ASTS) location.

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Figure 2 Vertical distribution of major bacterial groups at the Arabian Sea time series (ASTS)
location during spring intermonsoon (SIM), fall intermonsoon (FIM), and northeast monsoon
(NEM).

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Figure 3 Percentage of shared and unique OTUs in the surface and OMZ clone libraries obtained
during spring intermonsoon (SIM), fall intermonsoon (FIM), and northeast monsoon (NEM) at
the Arabian Sea Time Series (ASTS) location.









Figure 4 Phylogenetic tree of representative sequences of the OTUs common between surface
clone libraries obtained during spring intermonsoon (SIM), fall intermonsoon (FIM), and
northeast monsoon (NEM).

- 966 **CTASS** Common to all season surfaces, **CT2SS** Common to two season surface
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Figure 5 Phylogenetic tree of representative sequences of the OTUs common between OMZ
clone libraries obtained during spring intermonsoon (SIM), fall intermonsoon (FIM), and
northeast monsoon (NEM).

- 974 CTASO- Common to all season OMZ, CT2SO- Common to two season OMZ
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Figure 6 Rarefaction (a) and Collectors (b) curves for the surface and OMZ clone libraries at
90%, 97%, and 99% similarity level obtained during spring intermonsoon (SIM), fall
intermonsoon (FIM), and northeast monsoon (NEM) at the Arabian Sea Time Series (ASTS)
location