



Interactions between nocturnal turbulent flux, storage and advection at an ‘ideal’ eucalypt woodland site

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Abstract.

While the eddy covariance technique has become an important technique for estimating long-term ecosystem carbon balance, under certain conditions the measured turbulent flux of carbon at a given height above an ecosystem does not represent the true surface flux. Profile systems have been deployed to measure periodic storage of carbon below the measurement height, but have not been widely adopted. This is most likely due to the additional expense and complexity, and possibly also the perception – given that net storage over intervals exceeding 24 hours is generally negligible – that these measurements are not particularly important. In this study, we used a three year record of net ecosystem exchange of carbon and simultaneous measurements of carbon storage to ascertain the relative contributions of turbulent carbon flux, storage and advection (calculated as a residual quantity) to the nocturnal carbon balance, and to quantify the effect of neglecting storage. The conditions at the site are in relative terms highly favourable for eddy covariance measurements, yet we found a substantial contribution (~40%) of advection to nocturnal turbulent flux underestimation. The most likely mechanism for advection is cooling-induced drainage flows, the effects of which were observed in the storage measurements. The remaining ~60% of flux underestimation was due to storage of carbon. We also showed that substantial underestimation of carbon uptake (approximately 80 gC m⁻² a⁻¹, or 25% of annual carbon uptake) arose when standard methods of nocturnal flux correction were implemented in the absence of storage estimates. These biases were much larger than quantifiable uncertainties in the data. Neglect of storage also distorted the relationships between the carbon exchange processes (respiration and photosynthesis) and their key controls (light and temperature, respectively). We conclude that addition of storage measurements to eddy covariance sites with all but the lowest measurement heights should be a high priority for the flux measurement community.



1 Introduction

Over the past 2 decades, eddy covariance measurements have been widely adopted as a tool for aggregate flux measurement (Baldocchi, 2003), and there are now over 650 operational monitoring sites registered with the international flux network (Fluxnet: fluxnet.ornl.gov). Within the Australian regional network (OzFlux: www.ozflux.org.au), there are 29 active sites (Beringer et al., 2016, this issue). The use of the eddy covariance technique allows continuous automated monitoring of mass and energy fluxes, and long-term multi-site datasets have yielded valuable ecological insights in recent years (Baldocchi, 2008).

It has long been documented that eddy covariance measurements are prone to underestimation of the true surface efflux of carbon at night recognised (e.g. Goulden et al., 1996b). The key processes associated with this underestimation are storage and advection (Aubinet et al., 2012). In the former process, carbon may be stored below the measurement height under calm conditions. However, since the CO₂ mole fraction must be approximately preserved over longer time scales (e.g. 24 hours +), carbon respired and accumulated nocturnally is generally released with the initiation of buoyancy-generated turbulence following sunrise, such that the net storage of CO₂ is zero.

Advection involves mean transfer of carbon due to the development of horizontal and vertical gradients in scalar fields; the primary process by which this occurs is the initiation of gravity-driven drainage currents on sloping terrain (Aubinet, 2008). In contrast to storage, this mechanism generally results in a net loss of carbon from the observing system. Since substantial respiration occurs nocturnally, this loss causes a selective systematic error towards carbon uptake for cumulative carbon budgets at time scales > 24 hours. These drainage flows have been observed to occur on slopes of < 1° (Aubinet et al., 2003; Staebler and Fitzjarrald, 2004). Most non-agricultural ecosystem measurement sites are on sloped terrain because historically, flat, arable land has been cleared for agriculture. Drainage-induced advection is therefore thought to occur to varying extent at most sites.

The storage term can be calculated using relatively simple instrumentation that measures CO₂ concentrations along a vertical profile between the eddy covariance instrumentation and the ground (Yang et al., 2007). In contrast, attempts to measure advection have involved deployment of complex instrumentation and delivered highly uncertain results. Thus indirect approaches to data correction have been devised, the most common of which is the identification of a threshold below which the nocturnal turbulent carbon flux declines with turbulent activity (as expressed by friction velocity [u^*]) (Goulden et al., 1996b). Since there should be no relationship between u^* and ecosystem respiration, this decline is interpreted as an increase in the storage and advection terms at the expense of the turbulent flux. Data below this threshold (herein u^*_{th}) are discarded and replaced using functional relationships between known physical respiratory drivers (primarily temperature) and nocturnal carbon fluxes.



The u_* correction has been criticised on theoretical and practical grounds (Aubinet et al., 2012; Aubinet and Feigenwinter, 2010; Van Gorsel et al., 2007), but remains the most widely adopted approach to nocturnal data correction. It should be applied to the sum of the measured turbulent flux and storage terms (Papale, 2006), but is quite commonly applied to measurements of turbulent flux in isolation, rather than the sum of the turbulent flux and storage terms. This is because only 10-30% of sites have deployed profile systems to measure storage (Papale, pers. comm., 23/11/2015). In Australia, only four of the 29 active sites have profile measurements, whereas ≥ 15 sites have canopies of sufficient height to warrant them (using measurement height > 3 m as a threshold for requirement).

While the net storage over time is approximately zero, it is generally positive nocturnally as carbon accumulates below the measurement height. In the morning, the sign reverses due to two processes: 1) the turbulent transfer of the accumulated carbon upwards through the measurement plane, and; 2) photosynthetic carbon uptake by the canopy. Thus neglect of storage means that nocturnal respiratory carbon release is underestimated, but this is balanced by underestimation of morning ecosystem photosynthetic carbon uptake.

However, when the nocturnal u_* correction is applied, it implicitly accounts for both storage and advection. Since there is no corresponding correction for the reversal of the storage term after sunrise (i.e. a negative change in storage balancing the vertical turbulent transfer of the nocturnally accumulated carbon), the requirement for the storage term to be approximately zero is violated, and the nocturnally respired carbon is effectively counted twice (Aubinet et al., 2002). This unavoidably biases measurements towards net carbon efflux, and also affects the apparent relationship between ecosystem carbon fluxes and climatic controls.

Given the number of sites that do not have profile systems, it is thus important to quantify the effects of failing to measure storage. In this study, we use a three year-record of carbon exchange (including storage measurements) for an Australian eucalypt woodland to investigate the interaction between nocturnal turbulent flux, storage and advection. We devise a simple method to infer the magnitude of advection and in turn quantify the apportionment of the nocturnal ecosystem carbon source between turbulent flux, storage and advection. We quantify the biases in annual carbon exchange that arise from neglecting the storage term and discuss its effects on interpretation of carbon fluxes in the context of climatic drivers. Given the significant additional investment and complexity associated with the construction and deployment of profile systems alongside eddy covariance systems, it might be argued that the incurred bias of neglecting storage could be ignored if it is small relative to other measurement uncertainties. We therefore also propagate the errors associated with determination of u_{*th} , random measurement error and imputation (gap-filling) error to annual estimates of net carbon exchange, and assess their magnitude relative to biases due to neglect of storage.



2 Methods

2.1 Site description

The site became operational in December 2011 (see Table 1 for site characteristics). It was established as part of a project investigating the concurrent effects of catchment reforestation on biodiversity, carbon sequestration and stream water yields. From an eddy covariance perspective, the site is considered 'ideal' in that it is relatively flat and homogeneous within the footprint area, and the canopy is very open (leaf area index ≈ 1). While the tower is situated on a slope of approximate south-easterly aspect (Figure 1), the slope is generally less than 1° . Vegetation is relatively homogeneous, consisting of a sparse eucalypt overstorey (dominant species is *E. microcarpa*, stand leaf area index [LAI] ≈ 1 ; see Table 1) and a sparse shrub understorey. The open canopy reduces the potential for strong decoupling of the sub-canopy space from the overlying air, and underestimation of respiration by above-canopy eddy covariance systems relative to direct chamber measurements correspondingly decreases with declining leaf area (Speckman et al., 2014). Mean annual temperature for 2012-2014 was 15.9°C (minimum and maximum temperatures of -3.1°C and 45.0°C). Average annual rainfall for the nearest long-term rainfall measurement site (Mangalore Airport; Bureau of Meteorology station ID 088109) is 560 mm (1971-2000 average).

2.2 Instrumentation

The eddy covariance (herein EC) method was used to measure CO_2 fluxes at 36 m (mean $[\pm\text{SD}]$ canopy height was only 15.3 $[\pm 6.4]$ m, but emergent individual trees were up to 30 m). The EC method requires fast-response instrumentation to measure simultaneous variations in scalar (here carbon) and vector (3D wind velocities) quantities. A Campbell Scientific (Logan, USA) CSAT3 sonic anemometer (Table 2) was used to measure wind velocities and a Licor (Lincoln, USA) LI7500 infra-red gas analyser (herein IRGA) to measure CO_2 and H_2O vapour mole fractions. EC data were logged at 10 Hz (post-processing is described below), and 30-minute averages for radiant and subsurface energy fluxes and standard meteorology (temperature, humidity, wind speed and direction, rainfall, barometric pressure) were also logged. All data were transferred telemetrically to a central server at Monash University, Clayton, Australia.

A custom-built profile system using a Licor LI840 IRGA measured changes in CO_2 storage below the EC measurement height. The system consisted of an array of 6 gas intakes (configured logarithmically in the vertical - 0.5, 2, 4, 8, 16, 36 m) connected to a sampling system (Figure 2) via tubing (of equal length - 18m - for all heights; enclosure mounted at 18 m). A KNF Neuberger (Freiberg, Germany) NMP 850.1.2 vacuum pump drew air from all levels through a common manifold with sample and exhaust chambers. A bank of 12V SMC (Tokyo, Japan) VO307 solenoid valves switched each of the gas lines sequentially to a sampling loop (flow rate = 0.5 lpm) consisting of a gas analyser (Licor LI840) and Alicat Scientific (Tucson, USA) mass flow meter, while the remaining lines bypassed the loop and were exhausted to the atmosphere. Dwell time for



each level was 20 s (first 15 s for flushing of the manifold, average of last 5 s logged), translating to a measurement cycle of 2 minutes.

2.3 Data Processing and Analysis

- 5 Post-processing of EC data (including quality assurance and quality control) was undertaken using OzFluxQC, a software package developed by the OzFlux community (primarily P. Isaac, the data manager of OzFlux) in the Python programming language (Isaac et al., 2016, this issue). The Python programming language was used for all subsequent data analysis. Thirty-minute fluxes were calculated from the 10 Hz data using block averaging (Moncrieff et al., 2004). Corrections applied to the raw data included: 2-D coordinate rotation (Lee et al., 2004), in which the coordinate frame is rotated to force first the
 10 mean cross-wind, then second the mean vertical wind, to zero over the measurement period; frequency attenuation corrections, which account for loss of covariance associated with high frequency cut-off, sensor separation and path averaging of then instrumentation (Massman and Clement, 2004), and; density corrections associated with the effects of surface sensible and latent heat fluxes (Webb et al., 1980). Data were processed to level 4. In OzFluxQC, this represents the point at which all QA / QC except u^* correction has been applied to flux data, and meteorological data have undergone QA /
 15 QC and gap-filling (Isaac et al., 2016).

- We identified the u^* threshold (u^*_{th}) using change-point detection (CPD) following Barr et al. (2013). The time series was divided into multiple temperature-stratified samples, and a two-phase linear regression model was fitted to all possible change points (change in carbon exchange as a function of u^*) for each sample. If the change point that minimises the sum of
 20 squares error shows statistically significant improvement over a null model (no change point), the change point (i.e. u^*_{th}) is retained (subject to additional quality control criteria, as described by Barr et al., 2013). A bootstrapping procedure (in which the data for each year were randomly sampled with replacement 1000 times; see Papale, 2006) was used to generate a probability distribution for u^*_{th} , the mean and 95% confidence interval of which provides a best estimate and uncertainty interval for u^*_{th} .

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The rate of change of carbon storage was calculated from the difference between quasi-instantaneous (2-minute) vertical concentration profiles at the tower at the beginning and end of the flux averaging period (Finnigan, 2006). We adopted the approach for storage calculation of Yang et al. (2007):

$$30 \quad \left(\frac{\Delta C}{\Delta t}\right)_{k=1} \times z_{k=1} + \sum_{k=2}^n \left\{ \left[\left(\frac{\Delta C}{\Delta t}\right)_k + \left(\frac{\Delta C}{\Delta t}\right)_{k-1}\right] \times \frac{z_k - z_{k-1}}{2} \right\} \quad (1)$$



Here, $\Delta C / \Delta t$ is the time rate of change of carbon molar density ($\mu\text{mol C m}^{-3}$), k is the profile level, z is height above the surface (m) and n is the number of profile levels. Mole fraction reported by the IRGA was converted to carbon molar density using the ideal gas law (temperature measurements were drawn from instruments co-located with the air inlet for each profile level, whereas pressure measurements were drawn from ground level). The 2-minute period preceding the beginning and the end of the 30-minute period (e.g. 1128-1130 and 1158-1200 for the 1130-1200 period) were used to calculate ΔC . Given that the pump draw was simultaneously divided across 6 lines, there was a lag > 1 minute, so that the sampling was approximately temporally centred on the half hour. The average of the time derivative for two levels (k and $k-1$) is the best estimate of the time derivative for the *layer* that has k and $k-1$ as its upper and lower boundaries, except for the lower layer, for which it is assumed that $\Delta C / \Delta t$ for $k = 1$ is representative for the layer. Layers were scaled according to the layer thickness and the storage term represented the sum over all layers.

2.4 Analytical framework

We assess the carbon balance in the familiar context of a notional control volume with an orthogonal coordinate system (Finnigan et al., 2003), the mass balance of which (neglecting horizontal turbulent flux divergence) is:

$$NEE = \underbrace{\overline{w'c'}(h_m)}_I + \underbrace{\int_0^{h_m} \frac{\partial \bar{c}(z)}{\partial t} dz}_{II} + \underbrace{\int_0^{h_m} \left(\bar{u}(z) \frac{\partial \bar{c}(z)}{\partial x} + \bar{v}(z) \frac{\partial \bar{c}(z)}{\partial y} \right) dz}_{III} + \underbrace{\int_0^{h_m} \left(\bar{w}(z) \frac{\partial \bar{c}(z)}{\partial z} \right) dz}_{IV} \quad (2)$$

Here, NEE (net ecosystem exchange of carbon) is the true source term, term *I* is the turbulent flux across the upper horizontal plane of the control volume at instrument height h_m (\mathbf{w} is the vertical velocity, and overbar and prime denote mean and quasi-instantaneous fluctuation from mean, respectively), term *II* is the storage term integrated over finite time period (t) and control volume depth (z), term *III* is the sum of the advection components in the horizontal dimensions (x and y , with corresponding vectors \mathbf{u} and \mathbf{v}) and term *IV* the vertical advection. We adopted the standard micrometeorological convention suggested by Chapin et al. (2006) in which NEE is positive (negative) when the net transfer of carbon is from ecosystem to atmosphere (atmosphere to ecosystem). In the following text, equation 2 is simplified to:

$$NEE = F_c + S_c + Av_c + Ah_c \quad (3)$$

Here, F_c and S_c are the turbulent flux and storage terms, and Av_c and Av_h are the vertical and horizontal advection terms, respectively. During the day, when turbulence is well-developed, the turbulent flux (F_c) is generally the dominant term, but at night, the other terms may become dominant under weak mixing. Following the identification of u_{*th} , nocturnal data were



rejected where $u_* < u_{*th}$. While nocturnal advection was not measured, it was inferred as the residual of the terms in equation 3. Nocturnally, NEE is equivalent to ecosystem respiration (herein ER). While ER is unknown when $u_* < u_{*th}$, it can be estimated (\widehat{ER}) using an empirical model, the parameters of which are optimised for periods in which the sum of turbulent flux and storage approximates ER (*i.e.* when $u_* > u_{*th}$). Equation 3 thus becomes:

$$\widehat{ER} - F_c - S_c = Av_c + Ah_c \quad (4)$$

2.5 Imputation

Carbon flux and temperature data were used to optimise the parameters of an empirical temperature response function (optimisation used the Levenberg Marquardt algorithm implemented in the Python Scipy package) that was then used to estimate ER for $u_* < u_{*th}$ (and for subsequent gap-filling). The model was an Arrhenius-style function proposed by Lloyd and Taylor (1994):

$$ER = rb e^{E_o \left(\frac{1}{T_{ref} - T_0} - \frac{1}{T - T_0} \right)} \quad (5)$$

Here, rb is the reference respiration at a reference temperature (T_{ref}), E_o is an activation energy parameter that controls temperature sensitivity, and T_0 is the temperature at which metabolic activity approaches zero. T_0 and T_{ref} are fixed at -46.02 °C and 10 °C, respectively (Lloyd and Taylor 1994); the unconstrained version of the function is overparameterised (Reichstein et al., 2005; Richardson and Hollinger, 2005). We fitted E_o annually and then derived rb for successive 15-day windows (5-day step with linear interpolation to generate a continuous time series) (Reichstein et al., 2005). This avoids confounding of diurnal and seasonal temperature responses, and allows the model to capture low-frequency variation in ER associated with variables not explicitly represented in the model (*e.g.* soil moisture and substrate availability). We used temperature measured at the EC height (36 m) as the respiratory driver because it was found to have the lowest RMSE (relative to soil – which had the highest – and 0.5, 2, 4, 8 and 16 m air temperatures; data not shown).

Gap-filling was also required for daytime to assess the effects of nocturnal data treatment on annual NEE. We used a Michaelis Menten-type rectangular hyperbolic model (Ruimy et al., 1995) of modified form (Falge et al., 2001) to estimate NEE, where ER was calculated from equation 5 using daytime temperatures in conjunction with nocturnally-derived parameter estimates,;



$$NEE = \frac{\alpha Q}{1 - Q/2000 + \alpha Q/\beta} + ER. \quad (6)$$

Here, α is the initial slope of the photosynthetic light response, Q is photosynthetic photon flux density, and β is photosynthetic capacity at 2000 $\mu\text{mol photons m}^{-2} \text{ s}^{-1}$. The same window size, window step and interpolation procedure was used as for the nocturnal fitting of the respiration model. We adopted the additional light-response model criterion in which A_{opt} is modified to include a non-linear scaling factor to account for the effects of vapour pressure deficit (VPD) on stomatal conductance (Lasslop et al., 2010):

$$\beta = \begin{cases} \beta_0 e^{(-k(VPD - VPD_0))}, & VPD > VPD_0 \\ \beta_0, & VPD < VPD_0 \end{cases} \quad (7)$$

Here, VPD_0 is a threshold value above which stomatal conductance becomes sensitive to VPD, and k is a fitted parameter defining the β response to VPD.

2.6 Uncertainty estimation

We quantified sources of uncertainty in the data arising from random measurement and model error. We calculated random error from a daily differencing procedure (Hollinger and Richardson, 2005). When differences in critical drivers are sufficiently small ($<35 \text{ W m}^{-2}$ for insolation, $<3 \text{ }^\circ\text{C}$ for air temperature, and $<1 \text{ m s}^{-1}$ for wind speed), differences between NEE data pairs separated by 24 hours were considered to represent random error. Since random error in EC data is heteroschedastic and follows the Laplacian distribution, we calculated the standard deviation of the error ($\sigma[\delta]$) for j samples in i u_* bins as:

$$\sigma(\partial)_i = \sqrt{2} \frac{1}{n_i} \sum_{j=1}^{n_i} |\partial_{i,j} - \bar{\partial}_i|. \quad (8)$$

$\sigma[\delta]$ was regressed on the flux magnitude to derive a linear relationship that was in turn used to estimate $\sigma[\delta]$ for each datum. Monte Carlo simulations were used to translate these estimates to annual uncertainty, whereby for each of 10^4 simulations, estimates of random error for all observational data were aggregated over one year. This yielded a normal distribution of uncertainties, the 2σ bounds of which were taken as the annual uncertainty.



With respect to model error, we followed Keith *et al.* (2009) in which, for day and night conditions, a sub-sample of 10^3 records with observational data was randomly selected from the annual dataset. A proportion of the observational data in the subsample equal to the observed proportion of data missing in the annual time series was then replaced with model estimates, and the summed difference between the complete observational and gap-filled subsamples was calculated and expressed as a proportion of the observational sum. This was repeated 10^4 times, and the 2σ bounds of the proportional error was calculated, and then applied to the annual sum to produce an absolute annual model error.

To combine uncertainties, we assumed independence of the random and model estimates and sum in quadrature:

$$\varepsilon_{tot} = \sqrt{\varepsilon_r^2 + \varepsilon_m^2}, \quad (9)$$

Here, ε_{tot} , ε_r and ε_m are combined total, random and model uncertainty, respectively.

Other unquantified sources of error may also be present. Barr *et al.* (2013) argued that one of the key sources of uncertainty in annual NEE is the estimation of u_{*th} . We used their bootstrapping approach to derive robust confidence intervals for u_{*th} , and we included estimates of the effect on annual NEE of setting u_{*th} to the upper and lower bounds of the 95% confidence interval for u_{*th} .

3 Results and Discussion

3.1 Contribution of mass balance components to nocturnal carbon dynamics

A nocturnal relationship was observed between friction velocity (u_*) and F_c (Figure 3), with F_c declining quasi-linearly below $u_{*th} = 0.42 \text{ m s}^{-1}$ and approaching zero at zero turbulence. There was little seasonal variation in u_{*th} (data not shown), and the estimates and uncertainty bounds for all years were similar (Table 3); we conservatively used the highest estimate across all years. Given that the primary abiotic respiration controls, temperature and soil moisture, showed no relationship with u_* except at $u_* < 0.08 \text{ m s}^{-1}$ which is linked to declines in temperature, we interpreted this as an increase in the non-turbulent terms of the mass balance.

The decline in F_c with declining u_* was accompanied by a corresponding increase in S_c because as turbulence is progressively suppressed, carbon is expected (excluding advective losses) to be increasingly stored below the measurement height. The strong sensitivity of CO_2 accumulation to u_{*th} below the measurement height was observed in raw time series data under varying u_* (Figure 4); the CO_2 mole fraction responded very sensitively on variations in u_* , and the effect of u_* crossing u_{*th} is evident even in the raw data.



However, S_c was not the only important additional term in the nocturnal mass balance. A u_* dependent decline in $F_c + S_c$ was also evident (Figure 5) below a threshold of $u_* = 0.32 \text{ m s}^{-1}$, although this was inherently more uncertain due to the much higher random error in storage relative to the EC measurements (Table 3). There is no plausible explanation for u_* -dependent declines in the respiratory carbon source (again, excluding changes in relevant controls), and so we inferred that this represents carbon losses associated with the remaining mass balance terms (*i.e.* advection). This can be quantitatively estimated as a residual following equation 4. Note that parameter optimisation of the temperature response function used $F_c + S_c$ as the target variable because S_c is observed to be non-zero when $u_* \gg u_{*th}$ (see Figure 3, and subsequent discussion).

The terms in equation 4 are plotted as a function of u_* in Figure 5. The inferred advection estimate increased rapidly below u_* threshold = 0.32 m s^{-1} , and was comparable to S_c at the lowest u_* values. This indicates that under the calmest conditions, F_c approached zero, and approximately half of the carbon respired by the ecosystem was stored below the measurement height while the remainder was advected away. Integrated over the interval $0 < u_* < u_{*th}$, S_c accounted for 61% of the difference between F_c and \widehat{ER} , with the other 39% attributed to the advection components ($Av_c + Ah_c$). This indicates that even on very flat terrain, the nocturnal advection term is significant.

3.2 Inferred advection mechanisms

Given that $Av_c + Ah_c$ is inferred, it is not possible to assess the relative contributions of the two components to the mass balance. However, gravity-induced drainage flow is a common mechanism for horizontal advective carbon losses. These flows are expected to be most readily identifiable from changes in S_c because drainage flows are initiated under stable conditions when turbulence is suppressed. F_c is expected to be small and S_c correspondingly large (prior to the initiation of drainage flows) under these conditions.

Drainage flows are also expected to have a vertical spatial fingerprint. Specifically, their onset may result in changes in the magnitude of S_c primarily in the lower layers of the control volume. At the study site, mean canopy height was $15.3 \pm 6.4 \text{ m}$ (SD; Table 1). Given that drainage flows generally confined to the trunk space below the canopy (Aubinet et al., 2003), and conservatively assuming that the canopy comprises the upper 30% of tree height, drainage flows may be confined to depths of $<10 \text{ m}$, which are comparable to commonly reported values (Goulden et al., 2006; Mahrt et al., 2001). The impact on the carbon balance may nonetheless be large because as much as 70% of carbon is sourced from the soil in temperate forest ecosystems (Goulden et al., 1996a; Janssens et al., 2001; Law et al., 1999).



S_c for the individual layers is presented in Figure 6. There was a clear decline in S_c for all layers below 8 m when u_* was less than approximately 0.25 m s^{-1} . In contrast, storage in the higher 8-16 m and 16-36 m layers continued to increase near linearly. We hypothesise that this indicates the onset of drainage flows at low levels under stable conditions, causing horizontal advective (i.e. Ah_c) losses of carbon from the lower layers of the control volume. The ongoing increases in storage in the 8-16 and 16-36 m layers may indicate that the carbon source in these layers originated primarily from the vegetation than from upward transfer from lower layers. In the interval between the u_* thresholds for F_c alone and $F_c + S_c$ (i.e. $0.32 \leq u_* \leq 0.42 \text{ m s}^{-1}$), $Av_c + Ah_c$ was not significantly different to zero (see Figure 5). The linear relationship between each of the lower layers (0-0.5, 0.5-2, 2-4 and 4-8 m) and the mean 8-36 m layer in this interval $0.32 \leq u_* \leq 0.42 \text{ m s}^{-1}$ can be used to extrapolate the expected rate of change for those layers in the absence of advection when $u_* < 0.31 \text{ m s}^{-1}$.

Extrapolation of this linear relationship to conditions where $u_* < 0.32$ provides an estimate of the expected magnitude of S_c in the absence of advection. If drainage flows are the primary advective mechanism, then the sum of F_c and the linearly adjusted storage term should approximate \overline{ER} . The correction to the storage in the 0-8 m layers when $u_* < 0.31 \text{ m s}^{-1}$ increased the 0-36 m storage term such that for the interval $0 < u_* < 0.31$, the mass balance was approximately closed because $\overline{ER} - F_c \approx S_c$ to within the uncertainty (95% confidence interval) of the bin means over this interval (Figure 7). This indicates that the decline in S_c at lower layers was of approximately the same magnitude as the inferred advection, consistent with the presence of low-level drainage flows removing carbon from the control volume.

Our approach is subject to substantial uncertainty because the assumption that the linear relationship between levels holds for declining u_* may not be correct. Direct observation of the proposed advective mechanism (terrain-induced drainage flows) would provide clarification of the mechanisms driving the nocturnal C dynamics at the site. While wind speed and direction were measured for all levels, the instrumentation lacked the resolution to detect the weak winds that generally characterise drainage flows in moderate terrain (typically less than 0.5 m/s ; Aubinet et al., 2003; Goulden et al., 2006; Mahrt et al., 2001) (typically less than 0.5 m s^{-1} ; Aubinet et al., 2003; Goulden et al., 2006; Mahrt et al., 2001). Nonetheless, the available data are consistent with a primary advective mechanism of terrain drainage flows.

3.3 Effects of correction methods on diurnal and annual carbon balances

The importance of the contribution of S_c to the diurnal carbon balance is presented in Figure 8a. Given that the contribution averages approximately zero over the diurnal cycle, annual NEE sums for both F_c and $F_c + S_c$ were comparable: approximately -450 , -400 and $-560 \text{ gC m}^{-2} \text{ a}^{-1}$ for 2012, 2013 and 2014, respectively (Table 4; small differences [$< 20 \text{ gC m}^{-2} \text{ a}^{-1}$] were observed for F_c versus $F_c + S_c$, reflecting small differences in parameters of functions used for gap filling). However, the diurnal dynamics were substantially changed with the addition of S_c , with increased amplitude and phase shift of peak carbon uptake (from midday – synchronous with the solar radiative peak – on average, to about 1100).



Nocturnal u_* dependency of $F_c + S_c$ indicates the presence of advection under weak turbulence. Following u_* correction for this dependency (Figure 8b), estimated annual carbon uptake was reduced by 50–75 $\text{gC m}^{-2} \text{a}^{-1}$.

Given that the majority of sites both internationally and in Australia do not have profile measurement systems, we discuss the effects of neglect of S_c because this is the *de facto* approach taken for sites without profile systems. As a secondary option, a single point storage term (herein S_{c_pt}) can be derived from the EC gas analyser. This will underestimate storage for taller towers, where much carbon accumulates within the control volume, and is subject to substantial error (Gu et al., 2012; Yang et al., 2007) but may nonetheless potentially reduce bias.

Summing the turbulent flux and single point estimate (i.e. $F_c + S_{c_pt}$) and applying the u_* correction, the diurnal cycle was relatively well-approximated on average (Figure 9a). S_{c_pt} sums to zero over each year, but substantially underestimates S_c over the diurnal cycle (Figure 10). We expect to see a bias towards efflux because although the u_* correction rectifies the nocturnal underestimation of efflux, the ventilation of accumulated carbon through the turbulent flux is not balanced by a decline of corresponding magnitude in storage. However, the results differed among years (Table 4). In 2012, u_* -corrected NEE from $F_c + S_{c_pt}$ was comparable to the best estimate for $F_c + S_c$ (–375 and –382 $\text{gC m}^{-2} \text{a}^{-1}$ for $F_c + S_{c_pt}$ and $F_c + S_c$, respectively). In 2013 uptake was actually higher for $F_c + S_{c_pt}$ (–362 vs –338 $\text{gC m}^{-2} \text{a}^{-1}$ for $F_c + S_c$). In 2014, the value was substantially lower (–431 versus –480 $\text{gC m}^{-2} \text{a}^{-1}$ for $F_c + S_c$).

The inconsistency of the annual NEE estimate between years arose due to the effect of error on model parameterisation and subsequent imputation. Random error estimates (derived from equation 8) are shown in Figure 11; at low flux magnitudes the error associated with $F_c + S_{c_pt}$ was large relative to that for F_c alone or $F_c + S_c$. This may be due to a mismatch in the timing of peak S_{c_pt} relative to S_c ; small day-to-day variations in this timing may – in conjunction with noise in S_{c_pt} – result in large variations in NEE estimates for a given set of environmental conditions (as essentially demonstrated in Figure 11). Large errors in point-based storage estimates were also reported by Gu et al. (2012) for a forest ecosystem. This in turn affects the parameter estimates for rb in the respiration function (equation 5) and α in the light response function (equation 6) in particular; even if this translates to small relative errors in GPP and ER, the effect on NEE – as the small difference between two large exchanges – may in relative terms be substantial.

Correction of F_c yielded comparable nocturnal NEE to that estimated from u_* -corrected $F_c + S_c$ because it accounted for the nocturnal effects of storage and advection (Figure 9b). However, the morning increase in turbulent flux mixes nocturnally respired carbon stored in the control volume up through the measurement height. In the absence of a corresponding change in storage to account for this, the respiratory flux is counted twice, and annual NEE calculated using this method was underestimated by approximately 80 $\text{gC m}^{-2} \text{a}^{-1}$ relative to the best estimate (Table 4). This is larger than the correction for $F_c + S_c$ associated with advection (50–75 $\text{gC m}^{-2} \text{a}^{-1}$), which is consistent with the previously established fact that S_c



contributes more to F_c underestimation than advection. In the absence of S_c estimates, lower annual bias would be obtained for this study if the u_* correction were *not* applied.

The double-counting problem might be mitigated at sites without storage measurements if, after nocturnal data correction, the morning 'flush' period, when the accumulated carbon in the control volume is vented following the re-establishment of thermally generated turbulence after sunrise, is identified and removed. This flush is expected to be manifested as an upward spike in F_c following sunrise (e.g. Aubinet et al., 2012, Figure 5.2 therein). However, it may be difficult to identify objectively the flushing of accumulated CO_2 from the control volume from the behaviour of F_c alone.

- 10 We observed no spike in F_c indicating morning venting of accumulated carbon (Figure 8a). The magnitude of the expected F_c spike depends on: (i) the quantity of carbon stored, which depends on prior respiration and advection, and; (ii) the relative timing and magnitude of increase of source / sink activity and turbulent activity. On average, u_* reached u_{*th} for F_c approximately two hours after sunrise, whereas the storage term began to decline immediately when insolation increased (Figure 10). This indicates that carbon stored within the control volume began to be consumed by photosynthesis before efficient ventilation of the control volume was underway. The CO_2 mole fractions of the lowest levels measured by the profile system (0.5 and 2 m) also began to decline first (Figure 12), suggesting that shrubs reach the light-compensation point earlier after sunrise than do trees probably because lower near-surface temperatures are expected to suppress respiration. The effect on storage was small since the layers represent a small proportion of the control volume.
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- 20 Early-morning photosynthesis may be substantial in eucalypt-dominated ecosystems in which the characteristically pendulous (in some species up to 75% of mature leaves typically hang at angles > 80 degrees from horizontal; Pereira et al., 1987), amphistomatous leaves evolved to maximise incident radiation at low sun angles, which shifts photosynthetic activity towards periods with lower vapour pressure deficit (James and Bell, 1996). Mutual shading would partially counteract this effect at low sun angles, but this may have less effect in systems with sparse canopies such as the woodland in this study.
- 25 Similarly, an Australian temperate eucalypt forest site with long-term turbulent flux and CO_2 profile measurements showed no morning spike in CO_2 efflux (Van Gorsel et al., 2007, Figure 4 therein).

- Alternatively, the flush period may be defined as occurring between sunrise and the time at which the quasi-instantaneous, vertically integrated control volume mean of the CO_2 mole fraction approaches its long-term temporal mean. The long term temporal mean in this study was 395 ppm CO_2 , within 2 ppm of the mean southern-hemisphere ambient CO_2 mole fraction during 2012-2014. As long as the CO_2 mole fraction in the control volume exceeds that of the overlying atmosphere, there is potential for flushing of previously respired carbon from the control volume. If it is assumed that, over a sufficient period, mean CO_2 mole fraction in the control volume is in equilibrium with the overlying atmosphere, then the concentration gradient between the control volume and the overlying atmosphere will continue to facilitate upward turbulent transfer of
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CO₂ until the control volume CO₂ mole fraction declines to its mean. In this study, this occurred on average at approximately 0900 local standard time compared with austral / autumnal vernal equinox sunrise times of 0622 and 0611, respectively. Ensuring that no data were included during the period where flushing of nocturnally respired carbon could potentially have occurred would require removal of data between 0600-0900.

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Essentially the reverse process to that described above occurred during daytime and early evening. Negative daytime S_c after 0900 reflects ongoing photosynthetic CO₂ drawdown within the control volume. S_c approached zero at approximately 1600 on average, because insolation, which drives photosynthesis, is out of phase with surface heating, which drives u^* and respiration and so photosynthesis declines earlier than respiration and turbulent mixing. Therefore, NEE estimates continued to be biased throughout the day in the absence of S_c . However, this has little effect on long-term carbon balance because the daytime bias in F_c due to neglect of S_c is offset by nocturnal bias of opposite sign in the early evening. Given that the CO₂ mole fraction dropped below the mean at 0900 and remained below it until 2100, S_c summed over this period must be approximately zero.

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The period between sunset (austral / autumnal vernal equinox sunset times: 1830 and 1813) and 2100 is therefore the inverse of the flush period between 0600-0900, with the CO₂ mole fraction in the control volume *lower* than ambient which, in the absence of respiratory carbon production, would facilitate downward turbulent transfer of carbon. The increase in the storage term after sunset (Figure 10) partly reflects the 'refilling' of the control volume reservoir following daytime drawdown, which may contribute to the observation that nocturnal S_c reached its highest value when $u^* > u_{*th}$. While the effect on annual NEE may be minor, it violates the assumption in the u^* -correction approach that only the turbulent flux term is significant when $u^* > u_{*th}$. The implications of this for sites without storage measurements are discussed below.

3.4 Effects of neglecting carbon storage on physiological interpretation of data

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From the perspective of deriving annual NEE sums, it is daytime rather than nocturnal measurements that are more critical; applying the u^* correction to either F_c or $F_c + S_c$ resulted in similar estimates of nocturnal NEE on average. But the effects of neglecting S_c depend on time of night. F_c underestimated NEE following sunset, even where $u^* > u_{*th}$ for F_c .

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However, a secondary nocturnal problem is recognised in the literature: when u^* increases following extended calm periods, stored CO₂ is vented from the control volume, which artificially inflates F_c relative to the true source term, the extent of which effect will depend on the importance of advection (Aubinet et al., 2012). When a u^* threshold is imposed, such periods are likely to be included in the retained data. This has the opposite effect to the early-evening effect, and is more likely to be problematic later in the evening when stable stratification and substantial storage of respired carbon is more likely. Both effects were observed in the nocturnal progression of S_c for periods when $u^* > u_{*th}$ (Figure 13): S_c was on



average > 0 for the first 4–6 hours after sunset and < 0 afterwards. On balance, the effect in this study was to increase slightly the estimation of ER. This explains why S_c was slightly positive when $u_* > u_{*th}$ in Figure 3, and why F_c was slightly lower than $F_c + S_c$ at night (primarily in the early evening) in Figure 9b.

- 5 However, given that temperature decreases over the evening, this suggests that the slope of temperature response functions will be slightly increased for $F_c + S_c$ versus F_c alone. Given that the optimisation procedure minimises the prediction error, this may not have a large quantitative effect averaged over the evening, but interpretation of system response to temperature is distorted. Moreover, extrapolation beyond the parameterisation domain (e.g. estimation of daytime ER) may result in
- 10 error (because the function optimised using F_c will underestimate NEE at high temperatures). Any bias in estimated daytime ER will then necessarily propagate to estimation of GPP (commonly calculated as $NEE - ER$). Because these errors are offsetting, this is not likely to have a large effect on annual NEE estimates.

- Similar distortion of response to insolation occurs during the day. The addition of S_c substantially affects diurnal NEE
- 15 dynamics, particularly during the morning, which affects the interpretation of the controls on NEE. For example, Figure 14 shows the difference in radiation use efficiency (RUE – here simply defined as the ratio of mean NEE to mean insolation) during daylight hours for F_c alone versus $F_c + S_c$. RUE was higher in the early morning, and declined more sharply, when $NEE = F_c + S_c$. Such declines in RUE are often associated with stomatal response to increasing VPD, and so the importance of this driver may be missed or minimised when S_c is not measured.

- 20 Application of light-response function analysis to daytime data to extract either photosynthetic or respiratory parameters is problematic in the absence of storage measurements because $F_c \neq NEE$ during most of the day. The estimation of ER (and quantum efficiency) derived from light response function analysis (e.g. Gilmanov et al., 2003; Lasslop et al., 2010) is strongly dependent on the magnitude of observed NEE when insolation is low (sunrise and sunset), and thus the effect of
- 25 neglecting the storage term may be particularly distorting to these parameters (Aubinet et al., 2012).

3.5 Sources of uncertainty

- One of the largest sources of uncertainty in annual NEE estimates is expected to derive from uncertainty in u_{*th} (Barr et al., 2013; Papale, 2006). We propagated this uncertainty to annual NEE (for both F_c and $F_c + S_c$) by filtering and gap-filling the
- 30 data using the lower and upper bounds of the 95% confidence interval (CI) of the normally distributed population ($N = 10^3$) of u_{*th} derived from CPD (Table 3). Much larger effects were evident for the lower uncertainty bound ($\mu - 2\sigma$, where μ is the best estimate for u_{*th} and σ is the standard deviation), which is to be expected because systematic errors in nocturnal flux measurement occur at low u_* . However, there should be no systematic variation in NEE when $u_* > u_{*th}$. The direct effect of



the upper uncertainty bound ($\mu + 2\sigma$) is expected to be minimal. While the reduction of nocturnal data availability for higher u_{*th} is expected to increase parameter imputation uncertainty, the effect here was minor, with annual NEE for $u_{*th} = \mu$ and $u_{*th} = \mu + 2\sigma$ differing by $<10 \text{ gC m}^{-2} \text{ a}^{-1}$ in all years.

- 5 The uncertainty in u_{*th} was greater for $F_c + S_c$ than for F_c alone due to the additional random error inherent in S_c (see Finnigan, 2006 for further discussion), which feeds into the change-point detection process. This propagated to larger uncertainty in the lower bound for annual NEE ($50\text{--}75 \text{ gC m}^{-2} \text{ a}^{-1}$ compared to $20\text{--}40 \text{ gC m}^{-2} \text{ a}^{-1}$ for F_c), despite the effect of u_* correction for F_c was almost double that for $F_c + S_c$. This is because for $F_c + S_c$, the lower bound of the u_{*th} uncertainty is below the 1st percentile of the nocturnal data, such that the full effect of advection was propagated to annual NEE but for
10 F_c alone, it was closer to the 40th percentile, such that only a small proportion of storage and advection occurred in the interval between $u_{*th} = \mu$ and $u_{*th} = \mu - 2\sigma$.

While the lower bound uncertainty due to u_* threshold determination error was therefore of a magnitude only slightly smaller than the bias introduced by applying the u_* correction in the absence of storage measurements, the uncertainty resulted in an
15 increase in the potential uptake of carbon, whereas the bias was of opposite sign. Thus the uncertainty ranges for F_c and for $F_c + S_c$ do not overlap. In this sense, the uncertainty range for F_c alone is meaningful only in a very narrow sense (i.e. in terms of what has been measured as opposed to true source / sink uncertainty), and should not be reported as determining the uncertainty in annual NEE. We have determined that the annual NEE estimate derived from $F_c + S_c$ is not contained within the u_{*th} uncertainty interval for annual NEE derived from F_c alone.

- 20 It should also be noted that the large lower-bound uncertainty in annual NEE derived from $F_c + S_c$ is very likely overestimated. If the lack of correlation between u_* and F_c above u_{*th} indicates that additional terms in the mass balance are negligible, then S_c should approach zero at u_{*th} . Given that this is what we observed (Figure 3), and the measurement system (and associated measurement errors) for S_c is independent of that for F_c , this is an independent validation of u_{*th} . It is not
25 clear how such information might be used in the context of frequentist statistical analysis, but it strongly suggests that the uncertainty bounds for NEE that include the effects of u_{*th} uncertainty presented here are unrealistically large.

The NEE uncertainty contribution of combined random and model error was small ($<25 \text{ gC m}^{-2} \text{ a}^{-1}$) by comparison (Table 3), $<10\%$ of annual NEE for each year. Model error was greater than random uncertainty for both day and night, although
30 the difference was larger at night. This is because the majority of nocturnal observational data was removed by the u_* correction (thereby reducing random error) and replaced with model data. Given that the signal:noise is lower at night, model error is expected to be relatively large. The method used calculates and compounds observation – model data differences, but the observational data already contain random error so this necessarily inflates the propagated model uncertainty above errors associated with missing driver information or systematic measurement error. This problem is partially propagated to



the daytime because random error contributes to uncertainty in the nocturnally derived parameters of equation 5, which are then used to calculate the ER component of daytime NEE (equation 6). Moreover, the light-response function estimates are affected by daytime random error, which explains why model error was relatively large comparable to nocturnal values during the day despite 70-90% of the data being retained.

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Annual NEE uncertainty due to random and model error was approximately 20% greater for $F_c + S_c$ than for F_c alone, due to additional random error to that for F_c arising from the addition of S_c (Figure 11). This is largely nocturnally determined because the storage term is smaller and less variable during the daytime when fluxes are largest. The increased annual uncertainty of $F_c + S_c$ is largely due to higher model uncertainty, which most likely reflects the propagation of random error to model uncertainty through effects on non-linear parameter estimation.

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The interdependence of model and random error technically renders invalid the assumption of independence in equation 9. However, the effect is to increase rather than to decrease uncertainty, which is small. While there are methods to separate model and random components (see Dragoni et al., 2007), this generally requires co-location of two instrument arrays.

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However, the daily differencing procedure we used is known to overestimate error by up to a factor of 2 (Billesbach, 2011; Dragoni et al., 2007) due both to potential wind-dependent temporal variations in source / sink strength (which may materially affect annual NEE estimates; Griebel et al., 2016) and because some signal is included in the differencing procedure.

20

It should be emphasised that there are numerous sources of uncertainty that have not been quantified here. Perhaps most important of these is systematic errors in the measurements themselves, which may be an extremely important source of true uncertainty (Lasslop et al., 2008). Thus the uncertainties reported here for $F_c + S_c$ also should not be formally interpreted as total uncertainty in the true source / sink term, but as the uncertainty contributed by a subset of quantifiable errors.

4 Conclusions

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We used a simple method to infer advection from measurements combined with a simple and widely used empirical respiration model. Even at our very flat site, approximately 40% of flux underestimation was attributable to advection. Observation of reductions in storage at lower levels (within 8m of the surface) in response to declining u_* indicate that the most likely advective mechanism is terrain drainage flows. High resolution measurements of wind directions within the control volume would be invaluable for directly detecting the presence of terrain-aligned flows, and are planned for this site.

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Given that this site is relatively ideal, this follows earlier findings that drainage-induced advection is likely to affect most sites to some degree. Where the NEE time series consists of turbulent flux and storage, nocturnal correction for these effects



should be made under all circumstances. The level of bias incurred in the absence of storage measurements is contingent on the contribution of advection to the nocturnal mass balance. Nocturnal u_* correction is not advisable if the dominant term in nocturnal flux underestimation is expected to be storage (as found in this study), because the biasing effect of correcting the nocturnal component without the counterbalancing daytime component is larger than correcting for the smaller advection term. At sites with severe advection problems, neglect of storage is relatively less important, because if CO_2 is drained away nocturnally, the early-morning venting that causes bias following nocturnal correction is small.

But this contingency underscores the intractable nature of the problem: the relative contributions of storage and advection to the nocturnal mass balance cannot be quantitatively assessed in the absence of profile measurements. Moreover, even at sites where drainage flows are known to regularly occur at night, it is likely that shear-induced turbulence penetrates below canopy only under strong winds; the rarity of such conditions may result in the rejection of an unacceptably large number of data. Where this is the case, profile measurements are required to increase the proportion of available nocturnal data because storage increases prior to the onset of drainage flows, which only occur once the cooling air mass adjacent to the surface achieves sufficient density to overcome friction and begin to flow (Van Gorsel et al., 2007).

Storage measurements nonetheless introduce some complications for data interpretation. The additional random error in nocturnal storage measurements increases uncertainty in u_* threshold and, correspondingly, annual NEE (uncertainties due to direct random observation error and imputation error were small by comparison). But as we have argued, the lower bound uncertainty for u_* threshold is unrealistic, since the storage term on average approaches zero at the u_* threshold. This behaviour is expected if the central u_* threshold estimate from change point detection is approximately correct. Even if these uncertainties are considered accurate, when propagated to annual NEE, the resulting uncertainty intervals for the sum of turbulent flux and storage versus turbulent flux alone do not overlap. This indicates that biases are not subsumed within (quantified) uncertainties; effectively, profile measurements reduce precision and increase accuracy of annual NEE estimates.

We therefore believe that for both OzFlux and Fluxnet, the installation of profile systems for sites with trees (woodlands, forests, savannas) is extremely important to ensure that both determination of annual carbon exchange and interpretation of ecosystem processes are accurate. At the very least, the issues explored here need to be taken into consideration during data analysis, and alternative methods of estimating uncertainties at sites without profile systems need to be developed. For sites under the auspices of the Integrated Carbon Observation System (ICOS: www.icos-ri.eu), profile systems are mandatory; while this is not yet the case for OzFlux and Fluxnet, for accurate estimates of annual NEE, profile systems are vital.



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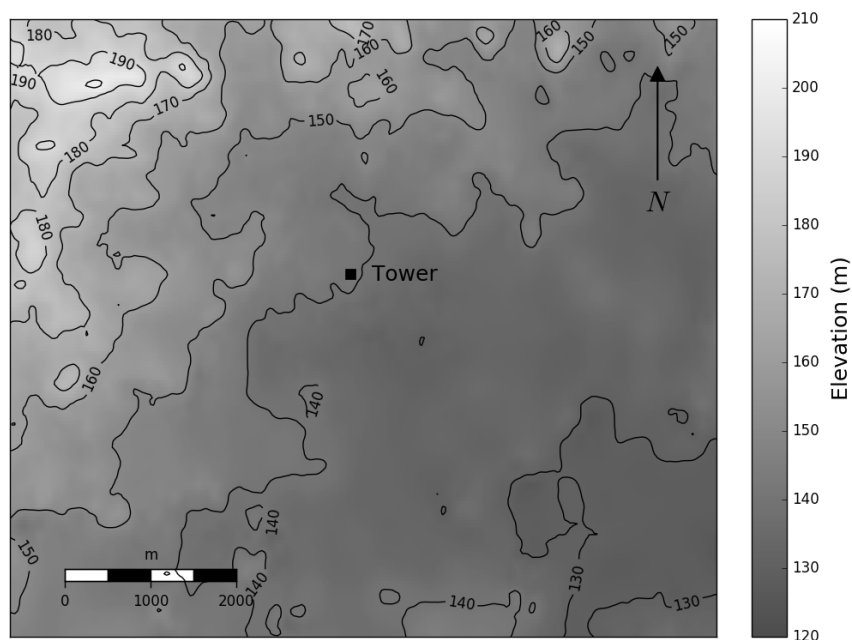


Figure 1: topography of terrain surrounding tower.

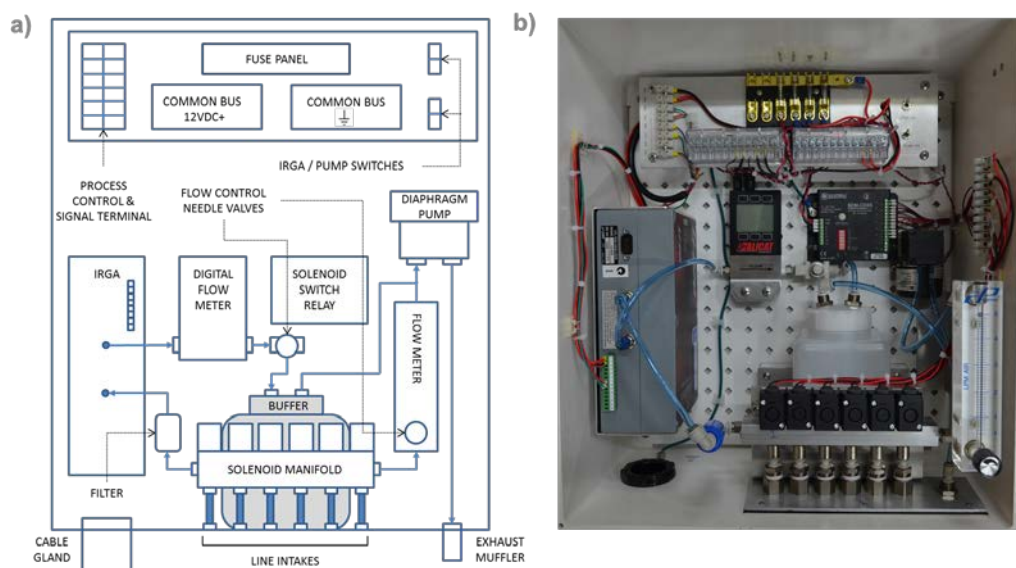


Figure 2: a) schematic, and; b) photographic layout of profile gas analysis system (note that a Licor LI820 is pictured here; an LI840 was used for measurement).

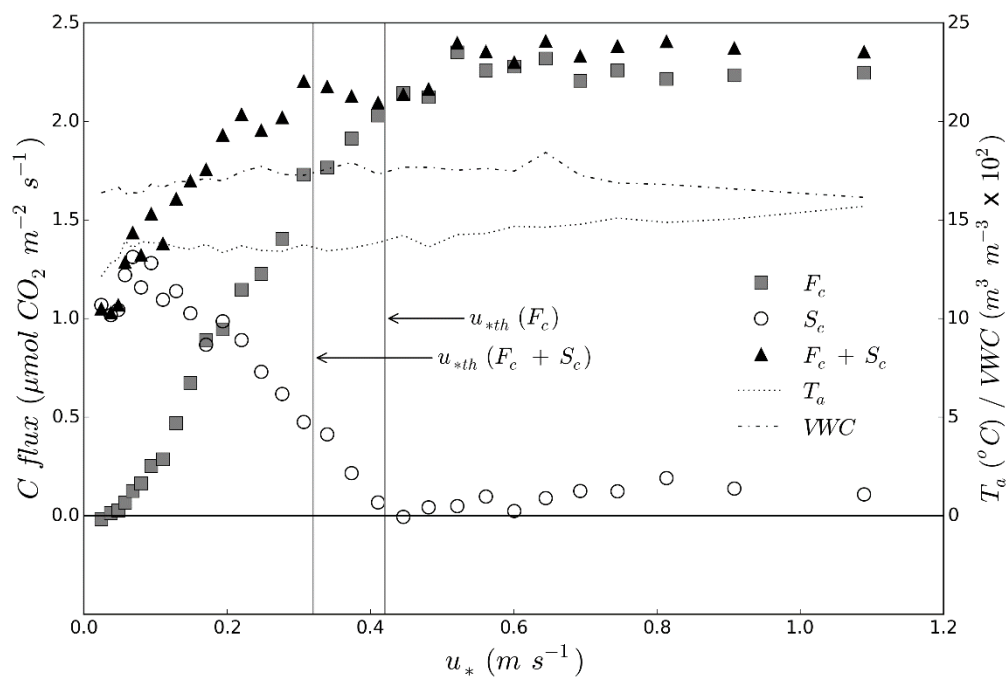


Figure 3: LH axis) dependence of mean measured nocturnal carbon mass balance components (turbulent flux [F_c], storage [S_c] and $F_c + S_c$) on friction velocity (u_*); RH axis) air temperature at EC instrumentation height (36m) and volumetric soil moisture content at 10cm. Vertical lines denote u_{*th} for both F_c and $F_c + S_c$, as labelled.

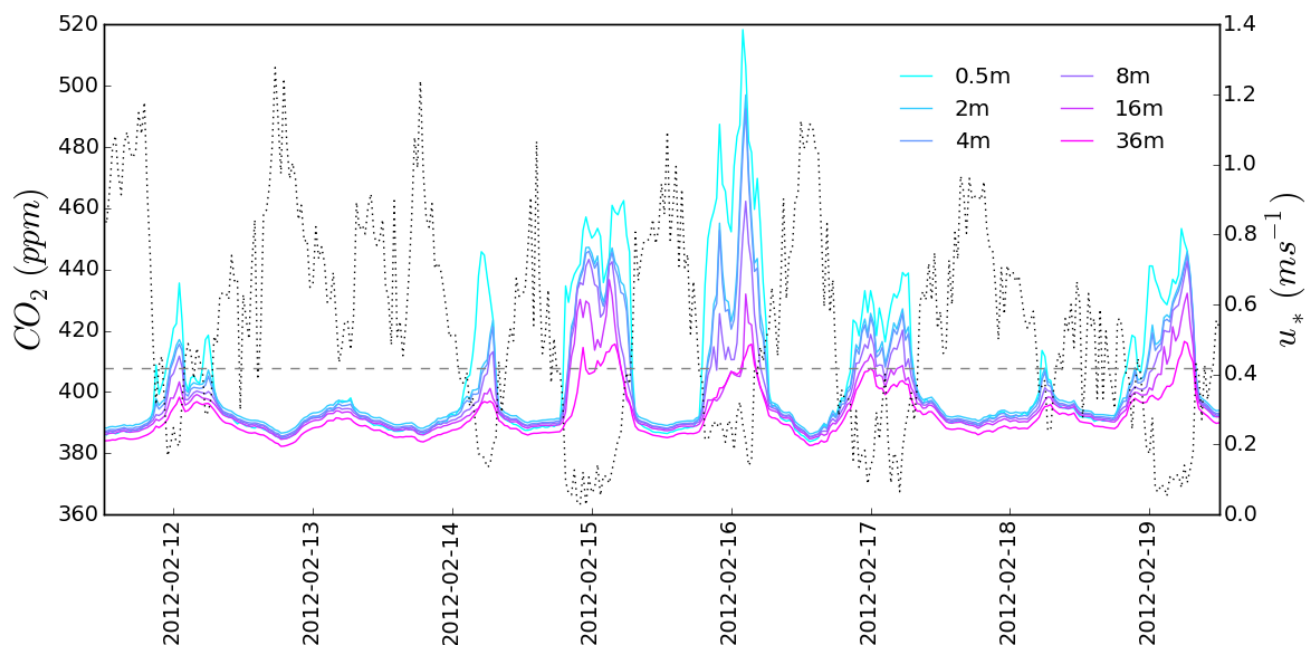


Figure 4: LH axis) profile system time series (date labels correspond to midnight) of CO_2 mole fraction (coloured solid lines); RH axis) time series of u_* measured at 36m (grey dotted line). Note horizontal dashed line marks u_{*th} for F_c .

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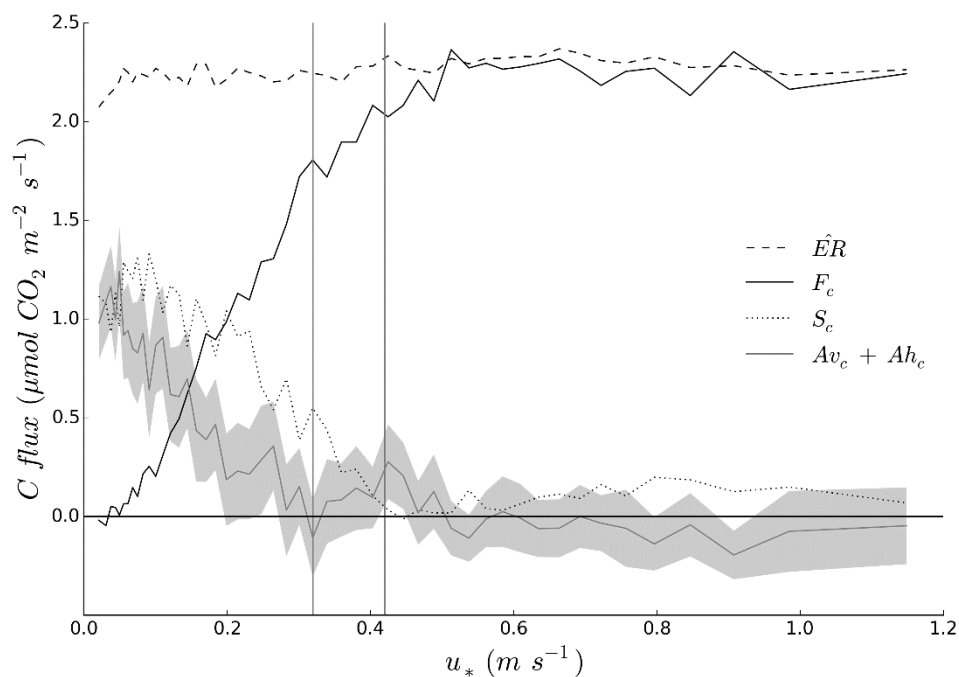




Figure 5: dependence of measured (F_c , S_c), model-estimated (\widehat{ER}) and inferred (Av_c , Ah_c) mass balance components on friction velocity (u_* ; the grey shaded area represents the 95% confidence interval for the sample bin mean of the inferred advection components). Vertical lines denote u_{*th} for both F_c and $F_c + S_c$, as per Figure 3.

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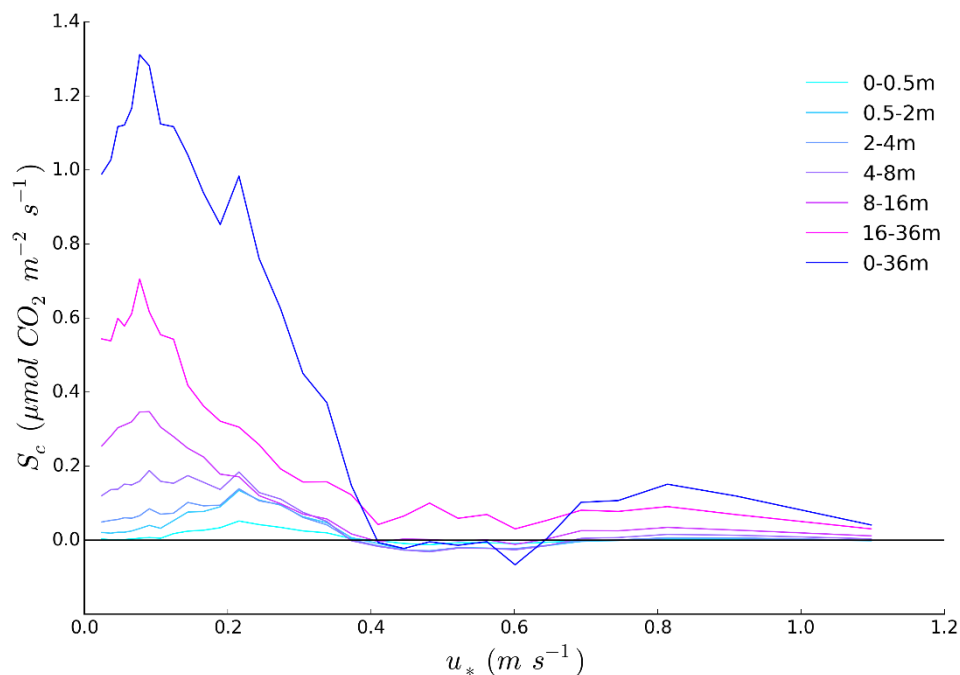
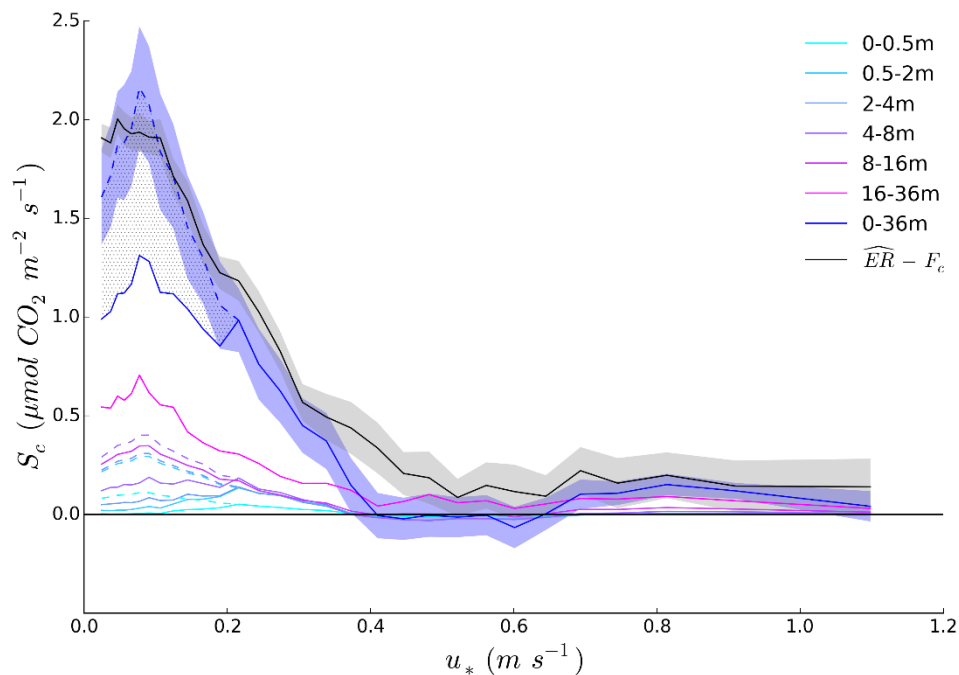


Figure 6: dependence of measured storage components on friction velocity for individual layers.

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5 **Figure 7: dependence of corrected storage components and $\widehat{R}_e - F_c$ on friction velocity (dashed lines represent corrected storage estimates; stippled region represents difference between measured and corrected 0-36m storage; shaded regions represent 95%CI for the u_* bin mean).**

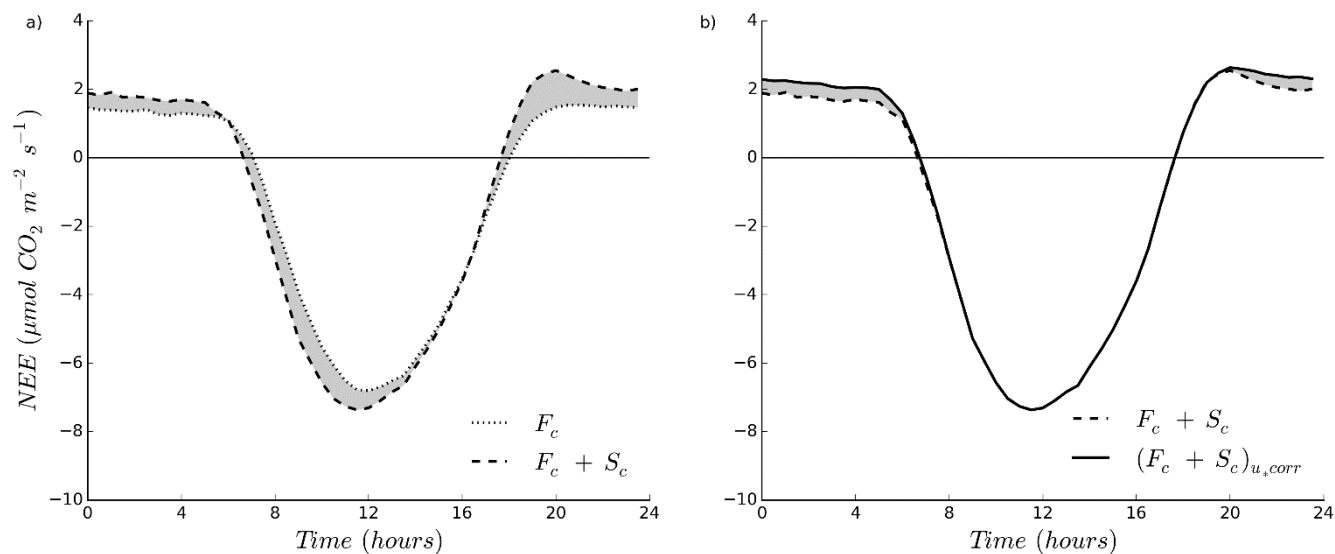


Figure 8: effects of (a) storage addition and (b) u^* correction on diurnal mean NEE dynamics.

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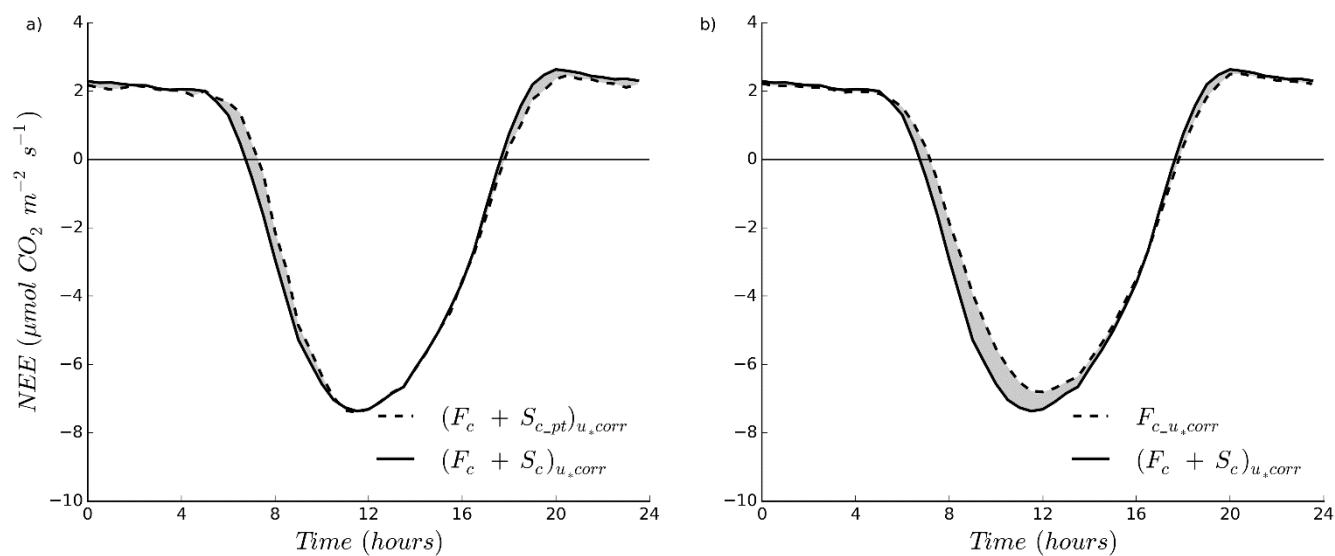
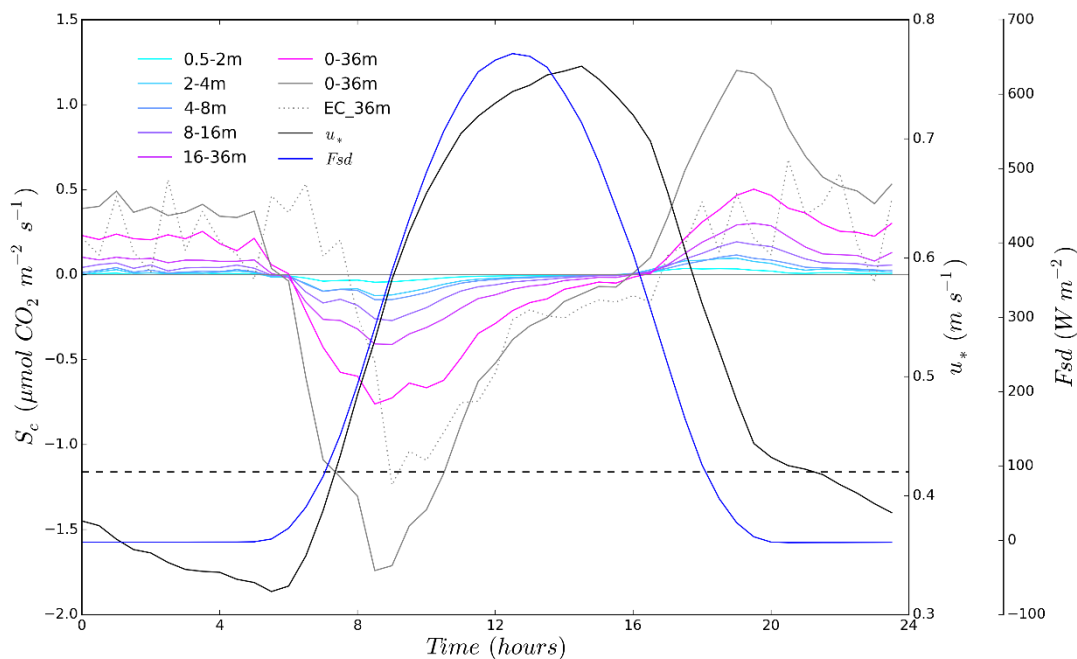


Figure 9: effects of (a) using storage term derived from eddy covariance IRGA, and (b) neglecting storage, on diurnal mean NEE dynamics.

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5 **Figure 10: mean diurnal cycle of aggregate and component S_c (LH axis; horizontal solid line represents mean S_c) and u_* (RH axis; horizontal dashed line represents change point-derived nocturnal u_* threshold); note that day length as indicated by insolation is slightly greater than 12 hours due to missing data for several months during winter / spring 2013, thereby slightly biasing the data towards longer photoperiod.**

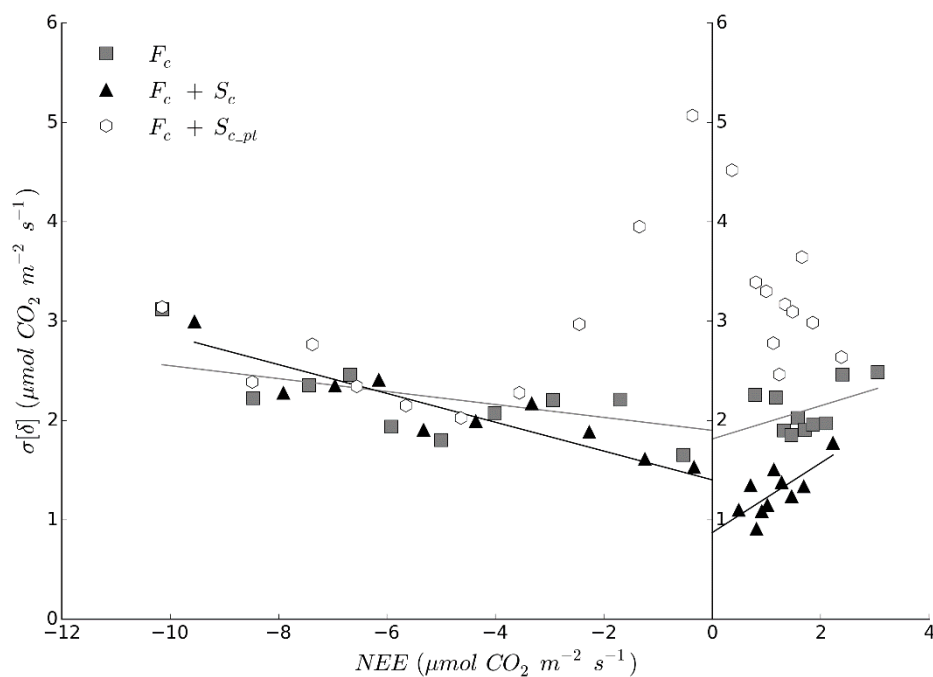


Figure 11: standard deviation of estimated random error ($\sigma[\delta]$) as a function of flux magnitude for turbulent flux (F_c), turbulent flux plus profile-based storage estimate ($F_c + S_c$) and turbulent flux plus point-based storage estimate ($F_c + S_{c_pt}$).

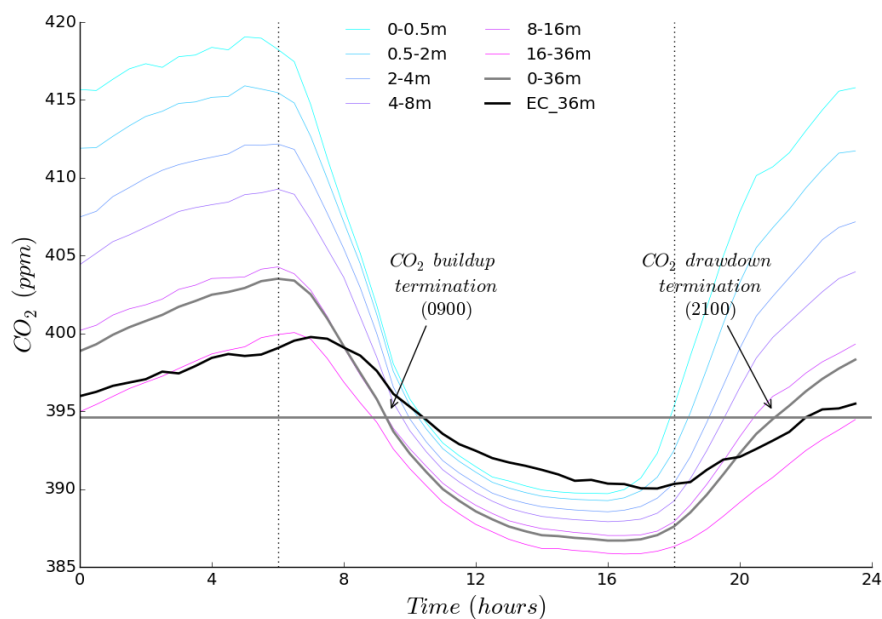


Figure 12: mean diurnal cycle of CO_2 mole fraction measured by profile system for given layers (coloured lines; grey line represents 0-36m height-integrated control volume mean) and by eddy covariance infra-red gas analyser at 36m (black line). Note that due to minor discrepancies between instruments, mole fractions for the eddy covariance IRGA are here baselined to the height-integrated profile IRGA multi-annual mean (394.6ppm).

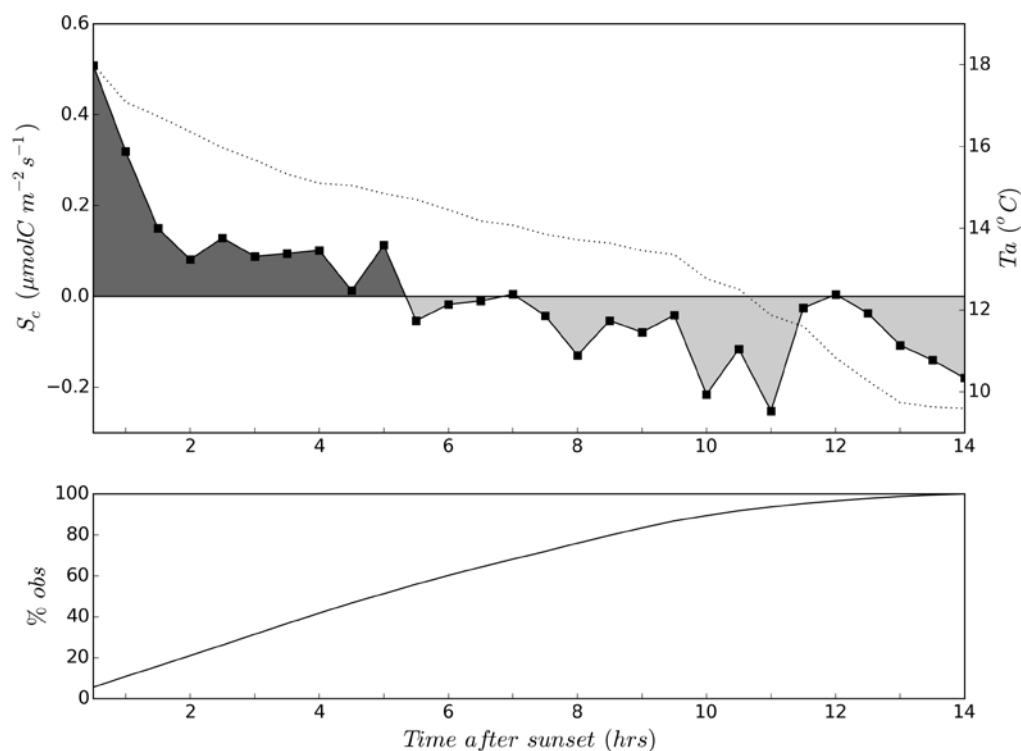


Figure 13: dependence of S_c (including only data where $u^* > u^*_{th}$) on time after sunset (upper panel; dotted line is air temperature); cumulative percentage of total nocturnal S_c observations (lower panel).

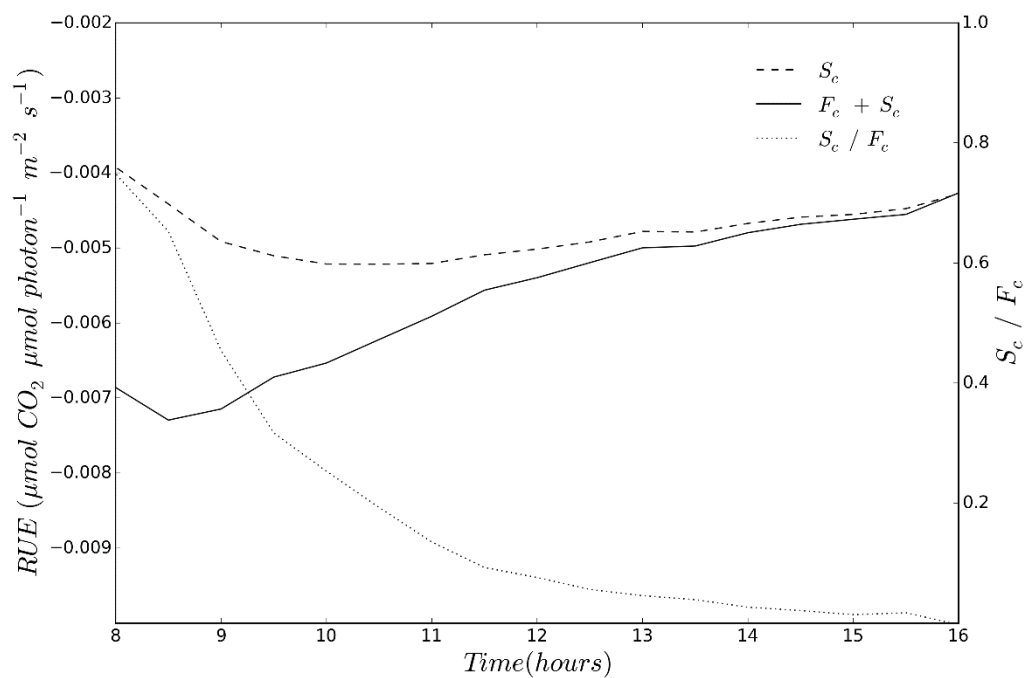


Figure 14: LH axis) effects of addition of F_c to S_c on radiation use efficiency (RUE); RH axis) S_c as a proportion of F_c .



Table 1: site characteristics

Latitude, longitude (° dec.)	-36.673215, 145.029247
Slope (°)	<1
Aspect	N/A
Dominant overstorey (>90%*) species	<i>Eucalyptus microcarpa</i>
Dominant understorey (>90%*) species	<i>Cassinia arculeata</i>
Mean canopy height ±SD (m)	15.3±6.4
Leaf area index (m ² m ⁻²)	~1.1
Mean annual temperature (°C)	15.9
Mean annual precipitation (mm)**: long term (1971-2000)	560

* By biomass (although also by number of individuals in the case of the overstorey)

** From nearest long-term rainfall measurement site (Mangalore Airport; Bureau of Meteorology station ID 088109)

Table 2: site instrumentation

Measurement	Instrument	Manufacturer
Wind vectors / virtual temperature	CSAT3	Campbell Scientific Instruments
Radiation components	CNR4	Kipp and Zonen
CO ₂ mole fraction (eddy covariance)	LI7500	Licor Biosciences
CO ₂ mole fraction (profile)	LI840	Licor Biosciences
Temperature / humidity	HMP45C	Vaisala
Wind speed / direction (profile)	Wind Sentry Set	RM Young
Barometric pressure	PTB110	Vaisala
Volumetric soil water content	CS616	Campbell Scientific Instruments
Soil heat flux	HFP01	Hukseflux
Soil temperature	TCAV	Campbell Scientific Instruments
Data logging	CR3000	Campbell Scientific Instruments



Table 3: lower 95%CI bound ($\mu - 2\sigma$), mean (μ), and upper 95%CI bound ($\mu + 2\sigma$) of Gaussian PDF of u_{*th} (derived from change point detection of bootstrapped samples - see Methods), data percentile (i.e. percentage data excluded for each u_{*th}) and resulting imputed annual estimate of NEE. Note: i) $\mu - 2\sigma$ set to zero if < 0 (e.g. $F_c + S_c$ in 2013); ii) respiration and light response function analysis could not find a solution for $F_c + S_c$ in 2013 when $u_* = 0.73$ (insufficient data for robust statistical fit).

Year	Condition	F_c			$F_c + S_c$		
		u_* ($m\ s^{-1}$)	Data percentile	Annual NEE ($gC\ m^{-2}\ a^{-1}$)	u_* ($\pm 95\% CI$)	Data percentile	Annual NEE ($gC\ m^{-2}\ a^{-1}$)
2012	$\mu - 2\sigma$	0.26	47	-323.1	0.01	<1	-450.6
	μ	0.39	60	-299.0	0.30	50	-381.5
	$\mu + 2\sigma$	0.52	72	-304.7	0.59	77	-387.1
2013	$\mu - 2\sigma$	0.19	38	-285.2	0	0	-364.1
	μ	0.40	61	-250.4	0.32	53	-313.3
	$\mu + 2\sigma$	0.61	79	-251.4	0.73	87	-
2014	$\mu - 2\sigma$	0.23	43	-433.4	0.02	<1	-552.7
	μ	0.42	63	-395.2	0.32	53	-483.2
	$\mu + 2\sigma$	0.61	79	-398.8	0.62	80	-481.9

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Table 4: gap-filled annual NEE ($gC\ m^{-2}\ a^{-1}$) for 2012-2014 obtained following different data treatment (F_c : turbulent flux only; $F_{c_{u_*corr}}$: turbulent flux with low u_* conditions removed; $F_c + S_{c_{pt}}$: summed turbulent flux and point-based storage estimate; $(F_c + S_{c_{pt}})_{u_*corr}$: summed turbulent flux and point-based storage estimate with low u_* conditions removed; $F_c + S_c$: summed turbulent flux and profile-based storage estimate; $(F_c + S_c)_{u_*corr}$: summed turbulent flux and profile-based storage estimate with low u_* conditions removed).

Year	F_c	$F_{c_{u_*corr}}$	$F_c + S_{c_{pt}}$	$(F_c + S_{c_{pt}})_{u_*corr}$	$F_c + S_c$	$(F_c + S_c)_{u_*corr}$
2012	-463	-301	-489	-375	-447	-382
2013	-403	-267	-492	-362	-388	-338
2014	-572	-393	-586	-431	-551	-480

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